



1 Traveltime calculations for qP-, qSV- and qSH- waves 2 in two-dimensional tilted transversely isotropic media

3 Guangnan Huang^{1,2}, Songting Luo², Juzhi Deng¹, Václav Vavryčuk³

4 ¹Department of Geophysics, East China University of Technology, Nanchang 330013, China.

5 ²Department of Mathematics, Iowa State University, Ames, IA 50011, USA.

6 ³Institute of Geophysics, The Czech Academy of Science, 14100 Praha, Czech Republic.

7 **Key Points:**

- 8 • A fast sweeping method is developed for computing the first-arrival traveltimes
- 9 by solving the anisotropic eikonal equation in 2D TTI media
- 10 • The method is consistent and monotone

Corresponding author: Guangnan Huang, HappyHuangSir@outlook.com

This article has been accepted for publication and undergone full peer review but has not been through the copyediting, typesetting, pagination and proofreading process which may lead to differences between this version and the Version of Record. Please cite this article as doi: 10.1029/2019JB018868

Abstract

This paper presents a fast sweeping method (FSM) to calculate the first-arrival travel-times of the qP, qSV, and qSH waves in two-dimensional (2D) transversely isotropic media, whose symmetry axis may have an arbitrary orientation (tilted transverse isotropy, TTI). The method discretizes the anisotropic eikonal equation with finite difference approximations on a rectangular mesh, and solves the discretized system iteratively with the Gauss-Seidel iterations along alternating sweeping orderings. At each mesh point, a highly nonlinear equation needs to be solved to update the numerical solution until its convergence. For solving the nonlinear equation, an interval that contains the solutions is first determined and partitioned into a few subintervals such that each subinterval contains one solution; then the false position method is applied on these subintervals to compute the solutions; after that, among all possible solutions for the discretized equation, a causality condition is imposed and the minimum solution satisfying the causality condition is chosen to update the solution. For problems with a point-source condition, the fast sweeping method is extended for solving the anisotropic eikonal equation after a factorization technique is applied to resolve the source singularities, which yields clean first-order accuracy. When dealing with the triplication of the qSV wave, solutions corresponding to the minimal group velocity are chosen such that continuous solutions are computed. The accuracy, efficiency, and capability of the proposed method are demonstrated with numerical experiments.

1 Introduction

The traveltimes of seismic waves are often used to study the interior structure of the Earth. Numerical methods for calculating the traveltimes of seismic waves play an important role in many seismic techniques, such as raypath backtracking, quality factor inversion, formation stress inversion, and Kirchhoff prestack depth migration. Accurate traveltimes can be calculated by ray-tracing methods and finite difference eikonal solvers. The ray-tracing method computes the traveltimes by solving an appropriate initial or boundary value problem for a system of ordinary differential equations. It can provide high-order accuracy for the traveltime tables (Cerveny, 1972; Shearer & Chapman, 1988). However, (1) raypaths often diverge from each other, and large spatial gaps often exist between two adjacent rays, especially in complicated heterogeneous velocity models (Vidale, 1990); (2) traveltimes are only calculated for shot-receiver pairs such that they must be interpolated onto a large number of grid nodes when they are used for seismic migration and tomography (Vinje et al., 1993; Gray & May, 1994; J. Huang & Bellefleur, 2012); (3) the two-point ray-tracing problem can be highly nonlinear such that it is difficult to solve efficiently; and (4) it can be difficult or time-consuming to distinguish whether the solution is a first or later arrival where triplication occurs. On the contrary, the eikonal solvers such as the finite difference traveltime calculation methods have no such disadvantages. For the past four decades, many eikonal solvers have been developed (Vidale, 1988; Podvin & Lecomte, 1991; Cao & Greenhalgh, 1993; Hole & Zelt, 1995; Sethian & Popovici, 1999; Kim, 2002; Zhao, 2005; Fomel et al., 2009; Stovas & Alkhalifah, 2012; Lan et al., 2014). Among these eikonal solvers, the fast marching method (FMM) and the fast sweeping method (FSM) are the two most popular ones. It is worth noting that finite difference eikonal solvers in general can only compute first arrivals, they may require well-designed numerical procedures for solving a complicated nonlinear system, and the number of iterations is problem-dependent if an iterative scheme is used.

The kinematic and dynamic features of seismic waves have great differences when they propagate in isotropic and anisotropic media. In isotropic media, only compressional and shear waves exist. While in anisotropic media, there may have three wavemodes: one quasi-compressional wave (qP) and two quasi-shear waves (qS1 and qS2). Each wave-mode propagates with its own wavespeed and polarization. The phase and group velocities of each wavemode are not only functions of elastic moduli parameters, but also func-

tions of the seismic wave direction. Many ~~of the~~ finite difference eikonal solvers have already been extended to calculate the traveltimes of seismic waves in anisotropic media (Dellinger & Symes, 1997; Eaton, 1993; Lecomte, 1993; Kim & Cook, 1999; Perez & Bancroft, 2001; Qian & Symes, 2002a). But most of them have been developed to deal with the tilted elliptically anisotropic (TEA) case only, for example the fast marching method for the TEA eikonal equation (Sethian & Vladimirovsky, 2003; Cristiani, 2009; Lou, 2006) and the fast sweeping method for the TEA eikonal equation (Tsai et al., 2003; Qian et al., 2007; Luo & Qian, 2012b). Han et al. (2017) developed a fast sweeping method which uses a quartic solver to tackle the quartic equation of the slowness surface with limited value range for the possible solutions, and obtained the traveltime of the qP wave. Bouteiller et al. (2018) developed a time-dependent discontinuous Galerkin method for computing the traveltime of the qP wave in 2D cases by transferring the static quartic anisotropic eikonal equation into a time-dependent equation. Waheed et al. (2015) and Waheed and Alkhalifah (2017) proposed an iterative fast sweeping method, like the fixed-point iteration method, to compute the traveltime of the qP wave by solving a sequence of TEA eikonal equations, where the slowness field of the TEA eikonal equation is updated iteratively whenever the numerical solution is updated until it converges to the solution of the original quartic anisotropic eikonal equation.

Finite difference eikonal solvers are efficient. However, they all suffer from the source singularities due to the non-differentiability of the solution at the point source (Qian & Symes, 2002b). The source singularities induce large errors near the source which will further spread to the whole computational domain and make the traveltime inaccurate. Without any treatments of the source singularities, such methods, even high-order methods, have only $O(h \log h)$ convergence order with h the mesh size. Moreover, this poses a problem to calculate some quantities involving derivatives of the traveltime, such as take-off angles and geometric spreading factors (Noble et al., 2014). This also poses a problem for iterative eikonal solvers involving derivatives of the traveltime as in Waheed et al. (2015). In order to overcome the difficulty caused by the source singularities, a few different methods have been proposed. The first method wraps a small region around the source, assumes the medium in the region is homogeneous such that the analytical solution can be obtained, and carries out the computation only outside of this region (Sethian & Popovici, 1999). This method is feasible only when the medium around the source is homogeneous. The second method refines the grid around the source in order to compensate the truncation error, but it involves *ad hoc* parameters without a clear selection criterion (Kim & Cook, 1999; Rawlinson et al., 2008). The third method uses an adaptive grid refinement near the source to control the error, but it incurs an additional heavy computational burden (Qian & Symes, 2002b). And the fourth method makes finite difference approximations for the eikonal equation on spherical grids centred on the source point in order to reduce inaccuracy (Alkhalifah & Fomel, 2001). However, the final result has to be interpolated to traveltime tables in Cartesian coordinates, which increases the computational cost.

In order to resolve the source singularities effectively without involving *ad hoc* parameters, a factorization approach has been proposed in Pica (1997), Zhang et al. (2005), Fomel et al. (2009), Luo and Qian (2012b), and Waheed and Alkhalifah (2017) for the isotropic eikonal equation and anisotropic eikonal equation with weak anisotropy. The traveltime is factored into two factors. One factor is a known function that captures the singularities around the source, and the other factor is smooth near the source. The smooth factor satisfies a modified/factored equation that can be solved efficiently with high accuracy. Hence, the original traveltime is recovered with high accuracy. Luo and Qian (2011) and Luo et al. (2012a) extended this factorization method to higher-order schemes to calculate first-arrival traveltimes and amplitudes. Treister and Haber (2016) and Treister and Haber (2017) used the first- and second-order finite difference schemes in the fast marching method to solve the factored eikonal equation. Luo and Qian (2012b) gave a systematic procedure to obtain analytical approximations for the known factor that cap-

tures the source singularities, and extended the factorization approach to eikonal equation in the TEA media. Following this approach, Waheed et al. (2014), Tavakoli et al. (2015), and Waheed and Alkhalifah (2017) proposed an iterative factored eikonal solver for computing the first-arrival traveltimes of the qP wave in TTI media with a simplified formulation of the anisotropic eikonal equation. Bouteiller et al. (2018) extended the factorization approach to a high-order method in the framework of discontinuous Galerkin method by transforming the simplified anisotropic eikonal equation into a time-dependent equation.

In review of the traveltimes calculation methods in anisotropic media with the finite difference schemes, one finds that most of them are the eikonal solvers for the qP wave governed by a weak or simplified anisotropic eikonal equation, because the qSV wave involves a triplication phenomenon when several qSV waves can propagate along the same raypath (Vavryuk, 2003a, 2006). This phenomenon is mostly associated with strong anisotropy or with directions close to point singularities in anisotropy (Vavryuk, 2003b). In this work, we propose an anisotropic eikonal solver for the qP, qSV and qSH waves, in the framework of the fast sweeping method, in 2D TTI media with arbitrary anisotropic strength. The method has the following important features: (1) the anisotropic eikonal equation is discretized on a mesh covering the computational domain; (2) the coupled system of the discretized equations among all grid points is solved iteratively by combining the Gauss-Seidel iterations with alternating sweeping orderings; (3) at each grid point, the subintervals that contain the solutions of the discretized equation are predetermined such that each subinterval contains exactly one solution and the false position method can be applied to compute the solutions efficiently; (4) among all possible solutions at a grid point, a causality condition is imposed to pick the one that corresponds to the first-arrival traveltimes; (5) the scheme is monotone and the numerical solution will converge to the viscosity solution as the mesh size approaches zero; and (6) the medium can have arbitrary anisotropic strength, and the symmetric axis of the anisotropic medium can have arbitrary orientation. These features make the eikonal solver more applicable to general situations and can obtain first-arrival traveltimes for the three waves. In order to resolve the source singularities for the anisotropic eikonal equation with point-source conditions, the factorization approach is applied such that a factored anisotropic eikonal equation is derived. The proposed anisotropic eikonal solver is further extended to solve the factored anisotropic eikonal equation following the similar procedures, which results in the fast sweeping methods for the factored anisotropic eikonal equation. The proposed methods enjoy all the appealing features of the usual fast sweeping method. The number of iterations is independent of the mesh size, and the numerical solution will converge to the desired weak solution as the mesh size approaches zero.

The rest of the paper is organized as follows: the general anisotropic eikonal equation for TTI media, and the factored anisotropic eikonal equations by use of the multiplicative and additive factorization techniques are introduced in Section 2. The numerical schemes in the framework of the fast sweeping method for solving the general and factored anisotropic eikonal equations are presented in Section 3. A few anisotropic models are used in the numerical experiments to verify the accuracy and efficiency of the proposed methods, which is discussed in Section 4. Conclusive remarks are given at the end.

2 Anisotropic eikonal equation

The determination of the traveltimes of seismic waves in general anisotropic media involves solving a sixth-order partial differential equation, i.e., the Christoffel equation (Cerveny, 2001),

$$\det |a_{ijkl}n_jn_l - v^2\delta_{ik}| = 0, \quad (1)$$

where a_{ijkl} is a rank-4 density normalized stiffness tensor, $\mathbf{n}$ is the normal vector of the wavefront, v is the phase velocity, and δ_{ik} is the Kronecker delta function. From the Christoffel

167 fel equation, and by introducing the slowness vector $\mathbf{p} = \frac{\mathbf{n}}{v}$, one can derive the anisotropic
 168 eikonal equation as

$$v|\nabla T| = 1, \quad (2)$$

169 where T is the traveltime and $\mathbf{p} = \nabla T$. In 2D cases, equation (2) can be rewritten as

$$H(P, Q) \equiv v_m \sqrt{P^2 + Q^2} - 1 = 0, \quad (m = 1, 2, 3). \quad (3)$$

170 where v_m ($m = 1, 2, 3$) is the phase velocity for the qP-, qSV- and qSH-wave, respec-
 171 tively, and $(P, Q) = (T_x, T_y)$. A general 2D TTI medium can be defined by five elas-
 172 tic moduli $\{a_{11}, a_{13}, a_{33}, a_{44}, a_{66}\}$ and the angle of the symmetry axis θ_0 (Thomsen, 1986).
 173 The expressions for v_m corresponding to the three wavemodes are given as (Daley & Hron,
 174 1977; Zhou & Greenhalgh, 2004),

$$\begin{aligned} v_{1,2} &= \sqrt{M \pm \sqrt{M^2 - N}}, \\ v_3 &= \sqrt{a_{44} + (a_{66} - a_{44}) \sin^2 \vartheta}, \end{aligned} \quad (4)$$

175 where M and N are defined as,

$$\begin{aligned} M &= 0.5(K_1 + K_2), \\ N &= K_1 K_2 - K_3, \end{aligned} \quad (5)$$

176 and

$$\begin{aligned} K_1 &= a_{44} \cos^2 \vartheta + a_{11} \sin^2 \vartheta, \\ K_2 &= a_{33} \cos^2 \vartheta + a_{44} \sin^2 \vartheta, \\ K_3 &= 0.25(a_{13} + a_{44})^2 \sin^2 2\vartheta. \end{aligned} \quad (6)$$

177 Here the angle ϑ is formed by the phase slowness direction and the direction of the sym-
 178 metry axis of the medium, i.e., $\vartheta = \theta - \theta_0$ with θ the phase slowness angle. The rela-
 179 tionship of these three angles is illustrated in Figure 1. According to Thomsen (1986),
 180 the phase slowness angle θ is formed by the wavefront normal and the vertical axis of
 181 the medium, and it can be computed by

$$\theta = \arccos \left(\frac{Q}{\sqrt{P^2 + Q^2}} \right). \quad (7)$$

182 With equations (4) to (7), the phase velocity v_m can be computed for an anisotropic
 183 TTI media, and v_m depends on the phase slowness angle θ .

184 2.1 Multiplicatively factored anisotropic eikonal equation

185 The multiplicative factorization method decomposes the solution of equation (2)
 186 as a product of two factors: the first factor is calculated analytically or numerically to
 187 capture the source singularities, and the second factor is a smooth correction near the
 188 source. Let us consider a multiplicatively factored decomposition,

$$T = T_0 \tau, \quad (8)$$

189 where T_0 is the pre-determined factor to capture the source singularities, and τ is the
 190 unknown factor that is smooth near the source.

191 Substituting equation (8) into equation (2) yields the following multiplicatively fac-
 192 tored anisotropic eikonal equation for τ ,

$$v_m |(P_1, Q_1)| = 1, \quad (m = 1, 2, 3), \quad (9)$$

193 with P_1 and Q_1 defined as,

$$\begin{aligned} P_1 &= T_{0x}\tau + T_0\tau_x, \\ Q_1 &= T_{0y}\tau + T_0\tau_y. \end{aligned} \quad (10)$$

194 Then, equation (9) can be rewritten as

$$H(P_1, Q_1) \equiv v_m \sqrt{P_1^2 + Q_1^2} - 1 = 0, \quad (m = 1, 2, 3). \quad (11)$$

195 2.2 Additively factored anisotropic eikonal equation

196 For the additively factored method, the traveltime T is decomposed as

$$T = T_0 + \tau, \quad (12)$$

197 where T_0 and τ are defined similarly.

198 Substituting equation (12) into equation (2) yields the additively factored anisotropic
199 eikonal equation for τ ,

$$v_m |(P_2, Q_2)| = 1, \quad (m = 1, 2, 3), \quad (13)$$

200 with P_2 and Q_2 defined as,

$$\begin{aligned} P_2 &= T_{0x} + \tau_x, \\ Q_2 &= T_{0y} + \tau_y. \end{aligned} \quad (14)$$

201 Then, equation (13) can be rewritten as

$$H(P_2, Q_2) \equiv v_m \sqrt{P_2^2 + Q_2^2} - 1 = 0, \quad (m = 1, 2, 3). \quad (15)$$

202 3 Fast sweeping method

203 To compute the first-arrival traveltimes for the three wavemodes, we will solve the
204 anisotropic eikonal equation (3) numerically in the sense of viscosity solutions, for which
205 the fast sweeping method is presented. For simplicity, we illustrate the scheme on a uni-
206 form mesh ($n_x \times n_y$) covering the rectangular computational domain, with mesh size
207 (h_x, h_y). We take $h_x = h_y = h$ for notational simplicity.

208 3.1 General eikonal equation

209 Figure 2 shows an interior grid point C with four neighboring grid points W, E, N, S .
210 The anisotropic eikonal equation can be discretized on the four triangles associated with
211 point C : ΔCEN , ΔCNW , ΔCWS and ΔCSE . Taking ΔCWS as an example, the dis-
212 cretized eikonal equation can be written as

$$v_m \left| \left(\frac{T_C - T_W}{h}, \frac{T_C - T_S}{h} \right) \right| - 1 = 0, \quad (m = 1, 2, 3), \quad (16)$$

213 where T_W and T_S are traveltimes at grid points W and S , respectively.

214 Given T_W and T_S , equation (16) must be solved to find solutions for T_C at C . Sim-
215 ilarly, the anisotropic eikonal equation (3) is discretized on the remaining three trian-
216 gles, and is solved for solutions T_C at C with given neighbor values. For each possible
217 solution for T_C , it is required to satisfy a causality condition such that it becomes a can-
218 didate for updating T_C at C . The causality condition is related to the characteristic di-
219 rection,

$$\begin{aligned} \frac{\partial H}{\partial P} &= \frac{P}{\sqrt{P^2 + Q^2}} v_m(\theta) - \frac{Q}{\sqrt{P^2 + Q^2}} \frac{\partial v_m}{\partial \theta}, \\ \frac{\partial H}{\partial Q} &= \frac{Q}{\sqrt{P^2 + Q^2}} v_m(\theta) + \frac{P}{\sqrt{P^2 + Q^2}} \frac{\partial v_m}{\partial \theta}. \end{aligned} \quad (17)$$

220 In triangle ΔCWS , it requires $\frac{\partial H}{\partial P} \geq 0$ and $\frac{\partial H}{\partial Q} \geq 0$. In general, the causality condition
 221 requires that (H_P, H_Q) passes through C and lies in the triangle used in the discretiza-
 222 tion. Then for all possible candidates for T_C from all the four triangles, we pick the min-
 223 imum one that corresponds to the first-arrival traveltime. If there are no candidates, T_C
 224 will be updated along the edges, for example on triangle ΔCWS ,

$$T_C = \min \left(T_W + \frac{h}{U_m^{WC}}, T_C + \frac{h}{U_m^{SC}} \right), \quad (m = 1, 2, 3), \quad (18)$$

225 where U_m^{WC} and U_m^{SC} are group velocities along edges $\vec{WC}$ and $\vec{SC}$, respectively.

226 The discretized equation (16) at all grid points are coupled together to form a sys-
 227 tem of nonlinear equations that can be solved using the Gauss-Seidel iteration with al-
 228 ternating sweeping orderings, which is the fast sweeping method.

229 **Algorithm Sketch: Fast Sweeping Method for Anisotropic Eikonal Equa-** 230 **tion**

- 231 1. Initialization: assigning exact/approximate values at grid points according to given
 232 boundary conditions, which will be fixed during the iterations, and assigning large
 233 positive values at all other grid points.
- 234 2. Gauss-Seidel iteration: sweeping the computational domain with four alternating
 235 orderings iteratively:

$$\begin{aligned} (a) \quad & i = 1 : n_x, j = 1 : n_y, & (b) \quad & i = 1 : n_x, j = n_y : 1, \\ (c) \quad & i = n_x : 1, j = 1 : n_y, & (d) \quad & i = n_x : 1, j = n_y : 1, \end{aligned}$$

236 and at each grid point C , updating T_C according to the above numerical proce-
 237 dure.

- 238 3. Termination: terminating the iteration if the L_1 -norm difference of the solutions
 239 between two successive iterations is smaller than the specified accuracy require-
 240 ment.

241 During the Gauss-Seidel iteration of the fast sweeping method, the discretized equa-
 242 tion (16) must be solved efficiently, and the group velocity along edges needs to be com-
 243 puted.

244 **3.2 Multiplicatively factored eikonal equation**

245 Taking ΔCWS as an example, the discretized equation of the multiplicatively fac-
 246 tored eikonal equation can be written as

$$v_m |(P_1, Q_1)| - 1 = 0, \quad (m = 1, 2, 3), \quad (19)$$

247 where P_1 and Q_1 are defined as,

$$\begin{aligned} P_1 &= T_{0x}\tau_C + T_0 \frac{\tau_C - \tau_W}{h}, \\ Q_1 &= T_{0y}\tau_C + T_0 \frac{\tau_C - \tau_S}{h}. \end{aligned} \quad (20)$$

248 Given τ_W and τ_S , this equation can be solved to find solutions for τ_C at C . Sim-
 249 ilarly, each possible solution for τ_C should satisfy a causality condition such that it be-
 250 comes a candidate for updating τ_C at C . The causality condition is similar as above with
 251 the characteristic direction given by,

$$\frac{\partial H}{\partial P_1} = \frac{P_1}{\sqrt{P_1^2 + Q_1^2}} v_m(\theta) - \frac{Q_1}{\sqrt{P_1^2 + Q_1^2}} \frac{\partial v_m}{\partial \theta},$$

$$\frac{\partial H}{\partial Q_1} = \frac{Q_1}{\sqrt{P_1^2 + Q_1^2}} v_m(\theta) + \frac{P_1}{\sqrt{P_1^2 + Q_1^2}} \frac{\partial v_m}{\partial \theta}. \quad (21)$$

252 In triangle ΔCWS , it requires $\frac{\partial H}{\partial P_1} \geq 0$ and $\frac{\partial H}{\partial Q_1} \geq 0$. Similarly, the factored anisotropic
 253 eikonal equation must be discretized and solved on the remaining three triangles. And
 254 from all possible candidates for τ_C that satisfy the causality condition, we pick the min-
 255 imum one corresponding to the first-arrival traveltimes. If there are no candidates, we will
 256 update τ_C along the edges in the following way as in Fomel et al. (2009) and Luo and
 257 Qian (2012b).

258 The characteristic equations of the multiplicatively factored eikonal equation are
 259 given as

$$\begin{aligned} \left(\frac{dx}{dt}, \frac{dy}{dt} \right) &= \left(\frac{\partial H}{\partial p}, \frac{\partial H}{\partial q} \right) = T_0 \left(\frac{\partial H}{\partial P_1}, \frac{\partial H}{\partial Q_1} \right), \\ \frac{d\tau}{dt} &= (p, q) \left(\frac{\partial H}{\partial p}, \frac{\partial H}{\partial q} \right)^T = 1 - \left(T_{0x} \frac{\partial H}{\partial P_1} + T_{0y} \frac{\partial H}{\partial Q_1} \right) \tau, \end{aligned} \quad (22)$$

260 where $(p, q) = (\tau_x, \tau_y)$ are derivatives of τ with respect to x and y , respectively. Ac-
 261 cording to the first equation, we have

$$\left(\frac{dx}{dt} \right)^2 + \left(\frac{dy}{dt} \right)^2 = T_0^2 U_m^2, \quad (m = 1, 2, 3), \quad (23)$$

262 where U_m is the group velocity that will be defined in equation (34). Using the method
 263 of characteristics, we can approximate τ_C at C along the edge $\overrightarrow{WC}$ (or $\overrightarrow{SC}$) by impos-
 264 ing that the ray falls on $\overrightarrow{WC}$ (or $\overrightarrow{SC}$). Let us take the edge $\overrightarrow{WC} = (\delta x, \delta y)$ as an exam-
 265 ple. According to equation (23), we have

$$\delta t = \frac{\sqrt{\delta x^2 + \delta y^2}}{T_0 U_m}, \quad (m = 1, 2, 3). \quad (24)$$

266 Then from the second equation of the characteristic equations, the approximation for
 267 τ_C , denoted as τ_{WC} , can be computed by,

$$\tau_{WC} = \frac{\tau_W + \delta t}{1 + T_{0x} \frac{\delta x}{T_0} + T_{0y} \frac{\delta y}{T_0}}. \quad (25)$$

268 Similarly, τ_C can also be calculated along $\overrightarrow{SC}$, denoted as τ_{SC} . And we will pick
 269 the minimal one by $\min \{ \tau_{WC}, \tau_{SC} \}$ to update τ_C at C .

270 The discretized equation (19) at all grid points are coupled together to form a sys-
 271 tem of nonlinear equations. This set of nonlinear equations can be solved similarly us-
 272 ing the fast sweeping method. The algorithmic sketch of the fast sweeping method for
 273 the multiplicatively factored eikonal equation is similar to that of the general eikonal equa-
 274 tion. However, the latter one involves three extra parameters T_0 , T_{0x} and T_{0y} .

275 3.3 Additively factored eikonal equation

276 Similarly, taking ΔCWS as an example, the discretized equation of the additively
 277 factored eikonal equation can be written as

$$v_m |(P_2, Q_2)| - 1 = 0, \quad (m = 1, 2, 3), \quad (26)$$

278 where P_2 and Q_2 are defined as,

$$\begin{aligned} P_2 &= T_{0x} + \frac{\tau_C - \tau_W}{h}, \\ Q_2 &= T_{0y} + \frac{\tau_C - \tau_S}{h}. \end{aligned} \quad (27)$$

Given τ_W and τ_S , this equation can also be solved to find solutions for τ_C at C . A similar causality condition is imposed on the solution such that it becomes a candidate for updating τ_C at C . The characteristic direction for the additively factored eikonal equation is given as,

$$\begin{aligned}\frac{\partial H}{\partial P_2} &= \frac{P_2}{\sqrt{P_2^2 + Q_2^2}} v_m(\theta) - \frac{Q_2}{\sqrt{P_2^2 + Q_2^2}} \frac{\partial v_m}{\partial \theta}, \\ \frac{\partial H}{\partial Q_2} &= \frac{Q_2}{\sqrt{P_2^2 + Q_2^2}} v_m(\theta) + \frac{P_2}{\sqrt{P_2^2 + Q_2^2}} \frac{\partial v_m}{\partial \theta}.\end{aligned}\quad (28)$$

In triangle ΔCWS , it requires $\frac{\partial H}{\partial P_2} \geq 0$ and $\frac{\partial H}{\partial Q_2} \geq 0$. Similarly, from all possible candidates that satisfy the causality condition, the minimum one is chosen to update τ_C at C . If there are no candidates, τ_C will be calculated along the two edges $\overrightarrow{WC}$ and $\overrightarrow{SC}$, respectively.

The characteristic equations of the additively factored eikonal equation are given as

$$\begin{aligned}\left(\frac{dx}{dt}, \frac{dy}{dt}\right) &= \left(\frac{\partial H}{\partial p}, \frac{\partial H}{\partial q}\right) = \left(\frac{\partial H}{\partial P_2}, \frac{\partial H}{\partial Q_2}\right), \\ \frac{d\tau}{dt} &= (p, q) \left(\frac{\partial H}{\partial p}, \frac{\partial H}{\partial q}\right)^T = 1 - \left(T_{0x} \frac{\partial H}{\partial P_2} + T_{0y} \frac{\partial H}{\partial Q_2}\right).\end{aligned}\quad (29)$$

According to the first equation, we have

$$\left(\frac{dx}{dt}\right)^2 + \left(\frac{dy}{dt}\right)^2 = U_m^2, \quad (m = 1, 2, 3). \quad (30)$$

Using the method of characteristics, we can compute τ_C at C along the edge $\overrightarrow{WC}$ (or $\overrightarrow{SC}$) by imposing that the ray falls on $\overrightarrow{WC}$ (or $\overrightarrow{SC}$). Let us take the edge $\overrightarrow{WC} = (\delta x, \delta y)$ as an example. According to equation (30), we have

$$\delta t = \frac{\sqrt{\delta x^2 + \delta y^2}}{U_m}, \quad (m = 1, 2, 3). \quad (31)$$

According to the second equation of the characteristic equations, the approximation for τ_C , denoted as τ_{WC} , can be computed as,

$$\tau_{WC} = \tau_W + \delta t - (T_{0x} \delta x + T_{0y} \delta y). \quad (32)$$

Similarly, τ_C can also be calculated along $\overrightarrow{SC}$, denoted as τ_{SC} . And the minimum one, $\min\{\tau_{WC}, \tau_{SC}\}$, is chosen to update τ_C at C .

The algorithmic sketch of the fast sweeping method for the additively factored eikonal equation is also similar to that of the general eikonal equation. And it also involves three extra parameters T_0 , T_{0x} and T_{0y} .

3.4 Calculation of traveltime T_0

For the two factorization techniques, T_0 should be computed for a homogeneous anisotropic model, where the medium parameters are assigned as those of the original model at the source point. T_0 can be computed as

$$T_0^m(\mathbf{x}) = \frac{|\mathbf{x} - \mathbf{x}_0|}{U_m(\theta)}, \quad (m = 1, 2, 3), \quad (33)$$

where $\mathbf{x}_0$ is the source position, and $\mathbf{x}$ is a position in the model domain. $U_m(\theta)$ is the group velocity along the ray direction $\mathbf{x} - \mathbf{x}_0$.

The expression of the group velocity can be found in previous work (Berryman, 1979),

$$U_m^2 = v_m^2 + \left(\frac{\partial v_m}{\partial \vartheta} \right)^2, \quad (m = 1, 2, 3), \quad (34)$$

where

$$\begin{aligned} \frac{\partial v_{1,2}}{\partial \vartheta} &= \frac{1}{2v_{1,2}} \left[\frac{\partial M}{\partial \vartheta} \pm \frac{M \frac{\partial M}{\partial \vartheta} - 0.5 \frac{\partial N}{\partial \vartheta}}{\sqrt{M^2 - N}} \right], \\ \frac{\partial v_3}{\partial \vartheta} &= \frac{(a_{66} - a_{44})}{2v_3} \sin 2\vartheta, \end{aligned} \quad (35)$$

and

$$\begin{aligned} \frac{\partial M}{\partial \vartheta} &= 0.5(a_{11} - a_{33}) \sin 2\vartheta, \\ \frac{\partial N}{\partial \vartheta} &= [K_1(a_{44} - a_{33}) + K_2(a_{11} - a_{44})] \sin 2\vartheta - 0.5(a_{44} + a_{13})^2 \sin 4\vartheta. \end{aligned} \quad (36)$$

The phase slowness angle θ , as well as the angle ϑ , are defined implicitly in that they depend on the solution T . We will present a numerical procedure to compute the group velocity and phase velocity along a ray direction in Section 3.8.

3.5 Solving the discretized equation (16)

At grid point C , the discretized equation (16) must be solved among all the four triangles. The equation is highly nonlinear in T_C , and it may have multiple solutions for T_C . Therefore, solving equation (16) for T_C is challenging. We present our numerical procedures for solving the equation: (1) determine an interval that contains all possible solutions, (2) partition the interval into subintervals such that each subinterval contains exactly one solution, and (3) apply false position method to find the solution in each subinterval (Press et al., 1992). We elaborate the numerical procedures by taking ΔCWS as an example.

According to the Fermat's principle, the interval that contains all possible values for T_C is

$$I_{WS} \equiv \left[\min(T_W, T_S), \min \left(T_W + \frac{h}{U_m^{WC}}, T_S + \frac{h}{U_m^{SC}} \right) \right], \quad (m = 1, 2, 3). \quad (37)$$

In order to determine subintervals that contain exactly one solution, we need to locate the extrema of H as a function of T_C , and use the extreme points to partition the interval into a few subintervals (see Figure 6). We can determine the extreme points by calculating the critical points through

$$\frac{\partial H}{\partial T_C} = \frac{\partial H}{\partial P} \frac{\partial P}{\partial T_C} + \frac{\partial H}{\partial Q} \frac{\partial Q}{\partial T_C} = 0, \quad (38)$$

which is an equation of the angle θ after algebraic manipulation, i.e.,

$$F(\theta) \equiv \left(\sin \theta v_m(\theta) - \cos \theta \frac{\partial v_m}{\partial \theta} + \cos \theta v_m(\theta) + \sin \theta \frac{\partial v_m}{\partial \theta} \right) \frac{1}{h} = 0. \quad (39)$$

Therefore, we can solve equation (39) for all possible solutions θ_i , and then find corresponding extreme points for H as a function of T_C through

$$T_C^i = \frac{\sin \theta_i T_S - \cos \theta_i T_W}{\sin \theta_i - \cos \theta_i}, \quad (i = 1, 2, 3, \dots). \quad (40)$$

We note that the solutions for $F(\theta)$ can be precomputed and saved. Then during the Gauss-Seidel iteration, T_C^i are computed through formula (40) to partition the interval I_{WS} into subintervals. Once the subintervals are determined, we can simply apply the false position method to find the solution in each subinterval, hence we can find all solutions in the interval I_{WS} .

3.6 Solving the discretized equation (19)

Similarly, at grid point C , the discretized equation (19) needs to be solved in all the four triangles. The equation is also highly nonlinear in τ_C , and it may have multiple solutions for τ_C . In order to introduce the procedures conveniently, we also take ΔCWS as an example.

According to the Fermat's principle, the interval that contains all possible values for τ_C is

$$I_{WS} \equiv \left[\min \left(\frac{\tau_W T_{0W}}{T_{0C}}, \frac{\tau_S T_{0S}}{T_{0C}} \right), \min(\tau_{WC}, \tau_{SC}) \right], \quad (41)$$

where τ_{WC} and τ_{SC} can be calculated by the method of characteristics along two edges $\overrightarrow{WC}$ and $\overrightarrow{SC}$, respectively.

In order to partition the interval I_{WS} into a few subintervals such that each subinterval contains exactly one solution, we can locate the extreme points of H as a function of τ_C , and use the extreme points as the partitioning points. We can determine the extreme points by calculating the critical points through

$$\frac{\partial H}{\partial \tau_C} = \frac{\partial H}{\partial P_1} \frac{\partial P_1}{\partial \tau_C} + \frac{\partial H}{\partial Q_1} \frac{\partial Q_1}{\partial \tau_C} = 0, \quad (42)$$

which is an equation of the angle θ after algebraic manipulation, i.e.,

$$F(\theta) \equiv \left(\sin \theta v_m(\theta) - \cos \theta \frac{\partial v_m}{\partial \theta} \right) L_1 + \left(\cos \theta v_m(\theta) + \sin \theta \frac{\partial v_m}{\partial \theta} \right) L_2 = 0, \quad (43)$$

with

$$\begin{aligned} L_1 &= T_{0x} + \frac{T_0}{h}, \\ L_2 &= T_{0y} + \frac{T_0}{h}. \end{aligned} \quad (44)$$

From the solutions of equation (43), denoted as θ_i , ($i = 1, 2, 3, \dots$), the extreme points of H as a function of τ_C can be computed by,

$$\tau_C^i = \frac{\sin \theta_i T_0 \tau_S - \cos \theta_i T_0 \tau_W}{\sin \theta_i T_{0y} h + \sin \theta_i T_0 - \cos \theta_i T_{0x} h - \cos \theta_i T_0}, \quad (i = 1, 2, 3, \dots). \quad (45)$$

The solutions of $F(\theta)$ can be pre-computed and saved for repeated use during the Gauss-Seidel iterations. In the local solver, τ_C^i can be calculated using formula (45) to partition the interval I_{WS} into a few subintervals. And then the false position method is applied to find the solution in each subinterval, hence all solutions in the interval can be found.

3.7 Solving the discretized equation (26)

Similarly, at grid point C , the discretized equation (26) needs to be solved for τ_C in all the four triangles. We also take ΔCWS as an example to demonstrate the procedures.

According to the Fermat's principle, the interval that contains all possible values for τ_C is

$$I_{WS} \equiv [\min(T_{0W} - T_{0C} + \tau_W, T_{0S} - T_{0C} + \tau_S), \min(\tau_{WC}, \tau_{SC})]. \quad (46)$$

The interval I_{WS} is also partitioned into a few subintervals such that each subinterval contains exactly one solution. The partitioning points are also the extreme points of H as a function of τ_C , and they can be determined by calculating the critical points through,

$$\frac{\partial H}{\partial \tau_C} = \frac{\partial H}{\partial P_2} \frac{\partial P_2}{\partial \tau_C} + \frac{\partial H}{\partial Q_2} \frac{\partial Q_2}{\partial \tau_C} = 0, \quad (47)$$

which is also an equation of the angle θ after algebraic manipulation, i.e.,

$$F(\theta) \equiv \left(\sin \theta v_m(\theta) - \cos \theta \frac{\partial v_m}{\partial \theta} + \cos \theta v_m(\theta) + \sin \theta \frac{\partial v_m}{\partial \theta} \right) \frac{1}{h} = 0, \quad (48)$$

Similarly, from the solutions of equation (48), denoted as θ_i , ($i = 1, 2, 3, \dots$), we can compute the extreme points of H as a function of τ_C by,

$$\tau_C^i = \frac{\cos \theta_i T_{0x} h - \sin \theta_i T_{0y} h - \cos \theta_i \tau_W + \sin \theta_i \tau_S}{\sin \theta_i - \cos \theta_i}, \quad (i = 1, 2, 3, \dots). \quad (49)$$

The solutions of $F(\theta)$ can be pre-computed and saved for repeated use during the Gauss-Seidel iterations. In the local solver, these critical points, associated with each grid point, can be used to partition the interval I_{WS} into a few subintervals. After that, we can use the false position method to find the solution in each subinterval, hence all solutions in the interval can be found.

3.8 Group velocity U_m along ray direction

When calculating T_0 , τ_{WC} and τ_{SC} , the group velocity U_m along a given ray direction must be determined. For example, U_m along edges is used in the local solver for a given grid point C . However, the group velocity is a function of the phase slowness direction, but not a function of the ray direction. If the phase slowness direction for a given ray direction can be determined, then the group velocity along the ray direction can be computed. Previous work (Vavrycuk, 2006, 2008; Zhang & Zhou, 2018) has investigated how to calculate the slowness vector for a given ray direction.

If the phase slowness direction $\mathbf{n} = (\sin \theta, \cos \theta)$ and the phase velocity v_m are given, the slowness vector $\mathbf{p}_m$ can be written as

$$\mathbf{p}_m = \frac{\mathbf{n}}{v_m}, \quad (m = 1, 2, 3). \quad (50)$$

According to Cerveny (2001), the phase slowness vector $\mathbf{p}_m$ and the group velocity vector $\mathbf{U}_m$ should satisfy the following equation,

$$\mathbf{p}_m \cdot \mathbf{U}_m = 1. \quad (51)$$

The phase slowness direction $\mathbf{n}$ and the ray direction $\mathbf{N}$ are given as

$$\mathbf{n} = \frac{\mathbf{p}_m}{|\mathbf{p}_m|}, \quad \mathbf{N} = \frac{\mathbf{U}_m}{|\mathbf{U}_m|}. \quad (52)$$

By dividing equation (51) with $|\mathbf{p}_m| |\mathbf{U}_m|$, one can derive the following equation,

$$\mathbf{n} \cdot \mathbf{N} - \frac{v_m}{U_m} = 0. \quad (53)$$

For a given ray direction, denoted as $\mathbf{N} = (N_1, N_2)$, equation (53) provides a way to calculate the phase slowness angle θ , as well as the phase velocity v_m and the group velocity U_m , i.e., by solving the following equation,

$$G(\theta) \equiv N_1 \sin \theta + N_2 \cos \theta - \frac{v_m(\theta)}{U_m(\theta)} = 0. \quad (54)$$

From equation (54), we can see that the only unknown is θ . If the phase slowness angle θ is computed, then group velocity U_m and the phase velocity v_m along the ray direction $\mathbf{N}$ can be calculated.

Equation (54) can be pre-solved for θ , as well as for v_m and U_m , along a set of ray directions. For example, on $\overrightarrow{\Delta CWS}$, this equation can be pre-solved for the two ray directions along two edges $\overrightarrow{WC}$ and $\overrightarrow{SC}$, respectively, and hence the group velocity along these two directions can be saved for repeated use during the Gauss-Seidel iterations.

3.9 Discussion of the methods

The proposed method is developed in the framework of the fast sweeping method. Therefore, it has all the desired properties of the fast sweeping method, such as consistency, monotonicity and convergence (Zhao, 2005; Qian et al., 2007; Luo & Zhao, 2016). The scheme is consistent with the first-order finite difference approximations, i.e., the discretized equation will converge to the original equation as the mesh size approaches zero. The causality condition implies that the scheme is monotone, i.e., at each grid point C , the numerical Hamiltonian H is non-decreasing with respect to the solution at C , and non-increasing with respect to the solutions at neighbor points. The consistency and monotonicity assure the stability of the scheme such that the numerical solution will converge to the viscosity solution (Barles & Souganidis, 1991; Zhao, 2005; Qian et al., 2007; Luo & Zhao, 2016) which corresponds to the first-arrival traveltime (Lions, 1982).

Similarly as in the usual fast sweeping method, the number of iterations depends on the problems and the desired accuracy requirement. However, for a given problem with a prescribed accuracy requirement, it is independent of the mesh size as the mesh size approaches zero (Zhao, 2005; Qian et al., 2007; Luo & Zhao, 2016).

In the local solver for solving the highly nonlinear equations to compute all possible updates at each grid point, necessary ingredients can be pre-determined prior to the Gauss-Seidel iterations. That is, equations (39), (43), (48) and (54) can be pre-solved with any appropriate root-finding methods, and their solutions can be saved for repeated use during the Gauss-Seidel iterations. Moreover, their solutions can be computed efficiently in parallel.

4 Numerical examples

We present several numerical experiments to demonstrate the efficiency and accuracy of the developed methods. In the numerical implementations, we denote one iteration as four sweeps over all grid points. Numerical errors at all grid points in L_1 -norm are recorded. The stopping criterion is 10^{-9} . Wherever applicable, the solutions computed by the shortest path method (SPM) on densely sampled meshes are used as the reference solutions (Zhou & Greenhalgh, 2006; G. Huang et al., 2014).

4.1 Homogeneous anisotropic model

We first use a homogeneous anisotropic model to test the effectiveness and feasibility of the proposed methods, along with demonstration of the necessary ingredients in the methods. The moduli parameters are $a_{11} = 5.2$, $a_{13} = 0.93$, $a_{33} = 4.0$, $a_{44} =$

1.0, $a_{66} = 1.0$, and the inclination angle θ_0 is set to 0° (VTI), 45° (TTI), or 90° (HTI). The computational domain is a 5×2.5 km rectangular domain, with a point source located at $x = 2.5$ km, $y = 0$ km.

For computing the group velocity along a given ray direction, equation (54), i.e., $G(\theta) = 0$, needs to be solved. Figure 3 shows an example of the function $G(\theta)$ with $\theta_0 = 0^\circ$ for the three wavemodes. From Figure 3, one can see that at least one root of $G(\theta)$ exists for each of the three wavemodes. If more than one root exist, for example in the case of triplication for the qSV wave (Vavrycuk, 2003a, 2003b, 2006), the one corresponding to the minimal group velocity is chosen. Figure 4 shows an example of the triplication for the qSV wave.

For solving the discretized equation on a triangle at a given grid point, the roots of $F(\theta)$ are used to partition the solution interval into subintervals (see Figure 5). Figure 6 shows an example of the function $F(\theta)$ that has about 2-6 roots for the three wavemodes. These roots correspond to the extreme points of H in the solution interval.

Traveltime tables of the qP, qSV and qSH-wave computed by the proposed methods are compared with the reference solutions in Figures 7-9. The number of iterations, L_1 -norm errors, CPU times and convergence orders are listed in Table 1. We observe the expected order of convergence $O(h \log(h))$ that is normal for the fast sweeping method. For the two factored methods, the machine error is dominant. For example with $\theta_0 = 0^\circ$: for the qP wave, the maximal relative error of the original method is 0.14, and the maximal relative errors of the two factored methods are close to 2.75×10^{-5} ; for the qSV wave, the maximal relative error of the original method is 0.33, and the maximal relative errors of the two factored methods are close to 8.5×10^{-6} ; for the qSH wave, the maximal relative error of the original method is 0.21, and the maximal relative errors of the two factored methods are close to 1.7×10^{-5} .

4.2 Overthrust TTI model

We further test the proposed methods on the overthrust TTI model, with the model parameters shown in Figure 10. The computational domain is a 6×4 km rectangular domain with a point source located at $x = 3$ km, $y = 0$ km. The reference solution is computed by the irregular grid shortest path method (SPM) (Zhou & Greenhalgh, 2006; G. Huang et al., 2014) on a 601×401 mesh, with 5 secondary nodes added to each cell boundary in the computation. The number of iterations, L_1 -norm errors, convergence orders and CPU times are listed in Tables 2-4, where we observe a clean first-order convergence for the two factored methods. The numerical plots are presented in Figures 11-13, where we can see that the solutions match very well, and the solutions computed by the two factored methods have better resolutions than those computed by the original FSM.

5 Conclusions

We present an efficient fast sweeping method (FSM) for calculating first-arrival traveltimes of the three wavemodes (qP, qSV, and qSH) in 2D heterogeneous, transversely isotropic media with arbitrary dipping symmetry axes. No weak anisotropy is assumed, and no simplification is made to the phase and group velocities. The proposed methods enjoy all the appealing features as in the fast sweeping method for the isotropic eikonal equation (Zhao, 2005), i.e., consistency, monotonicity and convergence.

For problems with a point-source condition, a factorization approach is applied to resolve the source singularities such that clear first-order convergence is obtained. Numerical experiments, including a homogeneous model, a three-layered model (Supple-

mentary Section) and the Overthrust TTI model, verify the effectiveness of the proposed methods.

Extension of the proposed methods to 3D anisotropic eikonal equation in TTI media is underway. The formulations of the methods are similar as those in 2D cases. The main difference is in equations (39), (43) and (48) for pre-computing the partitioning points and equation (54) for computing the group velocity along a given ray direction. For 3D cases, these equations will depend on two unknown angles (just like the spherical coordinate system in 3D versus the polar coordinate system in 2D), which is more challenging than 2D cases where these equations depend on one unknown angle. Solving a nonlinear equation of two unknowns is non-trivial. An extra condition/equation is required, which can be provided through the relation among the slowness vector, the ray vector and the symmetric axis. The resulting system of two equations will be solved to determine the two unknown angles, and the solutions can be saved and repeated used to compute the partitioning points in the Gauss-Seidel iterations, similarly as in 2D cases. We will report the results once the work is completed.

Acknowledgments

This research is jointly supported by the National Natural Science Foundation of China (41504095 and 41674077) and the China Scholarship Council. Luo was partially supported by NSF DMS (1418908 and 1719907). Model files in numerical examples can be downloaded from https://zenodo.org/record/3678494#.Xk_rj2gzY2x.

References

- Alkhalifah, T., & Fomel, S. (2001). Implementing the fast marching eikonal solver: Spherical versus cartesian coordinates. *Geophysical Prospecting*, 49(2), 165-178.
- Barles, G., & Souganidis, P. E. (1991). Convergence of approximation schemes for fully nonlinear second order equations. *Asymptotic Analysis*, 4(3), 271-283.
- Berryman, J. G. (1979). Long-wave elastic anisotropy in transversely isotropic media. *Geophysics*, 44(5), 896-917.
- Bouteiller, P. L., Benjema, M., Métivier, L., & Virieux, J. (2018). An accurate discontinuous galerkin method for solving point-source eikonal equation in 2D heterogeneous anisotropic media. *Geophysical Journal International*, 212(3), 1498-1522.
- Cao, S., & Greenhalgh, S. (1993). Calculation of the seismic first-break time field and its ray path distribution using a minimum traveltimes tree algorithm. *Geophysical Journal International*, 114(3), 593-600.
- Cerveny, V. (1972). Seismic rays and ray intensities in inhomogeneous anisotropic media. *Geophysical Journal International*, 29(1), 1-13.
- Cerveny, V. (2001). Seismic ray theory. *Cambridge University Press*, 21-26.
- Cristiani, E. (2009). A fast marching method for hamilton-jacobi equations modeling monotone front propagations. *Journal of Scientific Computing*, 39(4), 189-205.
- Daley, P. F., & Hron, F. (1977). Reflection and transmission coefficients for transversely isotropic media. *Bulletin of the Seismological Society of America*, 67(3), 661-675.
- Dellinger, J., & Symes, W. (1997). Anisotropic finite-difference traveltimes using a hamilton-jacobi solver. *SEG Technical Program Expanded Abstracts*, 1786-1789.
- Eaton, D. W. S. (1993). Finite difference traveltimes calculation for anisotropic media. *Geophysical Journal of the Royal Astronomical Society*, 114(2), 273-280.
- Fomel, S., Luo, S., & Zhao, H. (2009). Fast sweeping method for the factored eikonal equation. *Journal of Computational Physics*, 228(17), 6440-6455.

- Gray, S. H., & May, W. P. (1994). Kirchhoff migration using eikonal equation travel-times. *Geophysics*, 59(5), 810-817.
- Han, S., Zhang, W., & Zhang, J. (2017). Calculating qP-wave traveltimes in 2-D TTI media by high-order fast sweeping methods with a numerical quartic equation solver. *Geophysical Journal International*, 210(3), 1560-1569.
- Hole, J. A., & Zelt, B. C. (1995). 3-D finite-difference reflection traveltimes. *Geophysical Journal International*, 121(2), 427-434.
- Huang, G., Zhou, B., Li, H., Zhang, H., & Li, Z. (2014). 2D seismic reflection tomography in strongly anisotropic media. *Journal of Geophysics and Engineering*, 11(6), 1-8.
- Huang, J., & Bellefleur, G. (2012). Joint transmission and reflection traveltime tomography using the fast sweeping method and the adjoint-state technique. *Geophysical Journal International*, 188(2), 570-582.
- Kim, S. (2002). 3D eikonal solvers, Part I: First-arrival traveltimes. *Geophysics*, 67(4), 1225-1231.
- Kim, S., & Cook, R. (1999). 3-D traveltime computation using second-order ENO scheme. *Geophysics*, 64(6), 1867-1876.
- Lan, H., Chen, J., & Zhang, Z. (2014). A fast sweeping scheme for calculating P wave first-arrival travel times in transversely isotropic media with an irregular surface. *Pure and Applied Geophysics*, 171(9), 2199-2208.
- Lecomte, I. (1993). Finite difference calculation of first traveltimes in anisotropic media. *Geophysical Journal International*, 113(2), 318-342.
- Lions, P. L. (1982). *Generalized solutions of hamilton-jacobi equations*. Boston: Pitman Advanced Pub. Program.
- Lou, M. (2006). Traveltime calculation in 3D TTI media by the fast-marching method. *Expanded Abstracts in EAGE workshop*, 34-39.
- Luo, S., & Qian, J. (2011). Factored singularities and high-order lax-friedrichs sweeping schemes for point-source traveltimes and amplitudes. *Journal of Computational Physics*, 230(12), 4742-4755.
- Luo, S., & Qian, J. (2012b). Fast sweeping methods for factored anisotropic eikonal equations: Multiplicative and additive factors. *Journal of Scientific Computing*, 52(2), 360-382.
- Luo, S., Qian, J., & Zhao, H. (2012a). Higher-order schemes for 3D first-arrival traveltimes and amplitudes. *Geophysics*, 77(2), T47-T56.
- Luo, S., & Zhao, H. (2016). Convergence analysis of the fast sweeping method for static convex hamilton-jacobi equations. *Research in the Mathematical Sciences*, 3(1), 35-62.
- Noble, M., Gesret, A., & Belayouni, N. (2014). Accurate 3-D finite difference computation of traveltimes in strongly heterogeneous media. *Geophysical Journal International*, 199(3), 1572-1585.
- Perez, M., & Bancroft, J. C. (2001). Finite-difference anisotropic traveltimes and raypaths. *CREWES Research Report*, 13, 445-461.
- Pica, A. (1997). Fast and accurate finite-difference solutions of the 3D eikonal equation parametrized in celerity. *SEG Technical Program Expanded Abstracts*, 1774-1777.
- Podvin, P., & Lecomte, I. (1991). Finite difference computation of traveltimes in very contrasted velocity models: a massively parallel approach and its associated tools. *Geophysical Journal of the Royal Astronomical Society*, 105(1), 271-284.
- Press, W. H., Teukolsky, S. A., Vetterling, W. T., & Flannery, B. P. (1992). Numerical Recipes in C. *Cambridge University Press*, 353-359.
- Qian, J., & Symes, W. W. (2002a). Finite-difference quasi-P traveltimes for anisotropic media. *Geophysics*, 67(1), 147-155.
- Qian, J., & Symes, W. W. (2002b). An adaptive finite-difference method for traveltimes and amplitudes. *Geophysics*, 67(1), 167-176.

- Qian, J., Zhang, Y. T., & Zhao, H. K. (2007). A fast sweeping method for static convex hamilton-jacobi equations. *Journal of Scientific Computing*, 31(1-2), 237-271.
- Rawlinson, N., Hauser, J., & Sambridge, M. (2008). Seismic ray tracing and wavefront tracking in laterally heterogeneous media. *Advances in Geophysics*, 49, 203-273.
- Sethian, J. A., & Popovici, M. A. (1999). 3-D traveltimes computation using the fast-marching method. *Geophysics*, 64(2), 516-523.
- Sethian, J. A., & Vladimirov, A. (2003). Ordered upwind methods for static hamilton-jacobi equations: Theory and algorithms. *SIAM Journal on Numerical Analysis*, 41(1), 325-363.
- Shearer, P. M., & Chapman, C. H. (1988). Ray tracing in anisotropic media with a linear gradient. *Geophysical Journal International*, 94(3), 575-580.
- Stovas, A., & Alkhalifah, T. (2012). A new traveltimes approximation for TI media. *Geophysics*, 77(4), C37-C42.
- Tavakoli, B., Ribodetti, A., Virieux, J., & Operto, S. (2015). An iterative factored eikonal solver for TTI media. *SEG Technical Program Expanded Abstracts*, 3576-3581.
- Thomsen, L. (1986). Weak elastic anisotropy. *Geophysics*, 51(10), 1954-1966.
- Treister, E., & Haber, E. (2016). A fast marching algorithm for the factored eikonal equation. *Journal of Computational Physics*, 324(1), 210-225.
- Treister, E., & Haber, E. (2017). A multigrid solver to the helmholtz equation with a point source based on travel time and amplitude. *Numerical Linear Algebra with Applications*, 26(1), 1-20.
- Tsai, Y.-H. R., Cheng, L.-T., Osher, S., & Zhao, H.-K. (2003). Fast sweeping algorithms for a class of hamilton-jacobi equations. *SIAM Journal on Numerical Analysis*, 41(2), 673-694.
- Vavrycuk, V. (2003a). Generation of triplications in transversely isotropic media. *Physical Review B*, 68(5), 054107.
- Vavrycuk, V. (2003b). Parabolic lines and caustics in homogeneous weakly anisotropic solids. *Geophysical Journal International*, 152(2), 318-334.
- Vavrycuk, V. (2006). Calculation of the slowness vector from the ray vector in anisotropic media. *Proceedings of the Royal Society A Mathematical Physical and Engineering Sciences*, 462(2067), 883-896.
- Vavrycuk, V. (2008). Real ray tracing in anisotropic viscoelastic media. *Geophysical Journal International*, 175(2), 617-626.
- Vidale, J. (1988). Finite-difference calculation of travel times. *Bulletin of the Seismological Society of America*, 78(6), 2062-2076.
- Vidale, J. (1990). Finite-difference calculation of traveltimes in three dimensions. *Geophysics*, 55(5), 521-526.
- Vinje, V., Iversen, E., & Gjøystdal, H. (1993). Traveltimes and amplitude estimation using wavefront construction. *Geophysics*, 58(8), 1157-1166.
- Waheed, U. B., & Alkhalifah, T. (2017). A fast sweeping algorithm for accurate solution of the tilted transversely isotropic eikonal equation using factorization. *Geophysics*, 82(6), WB1-WB8.
- Waheed, U. B., Yarman, C. E., & Flagg, G. (2014). An iterative fast sweeping based eikonal solver for tilted orthorhombic media. *SEG Technical Program Expanded Abstracts*, 480-485.
- Waheed, U. B., Yarman, C. E., & Flagg, G. (2015). An iterative, fast-sweeping based eikonal solver for 3D tilted anisotropic media. *Geophysics*, 80(3), C49-C58.
- Zhang, L., Rector, J. W., & Hoversten, G. (2005). Eikonal solver in the celerity domain. *Geophysical Journal International*, 162(1), 1-8.
- Zhang, L., & Zhou, B. (2018). Calculation of slowness vector from ray directions for qP- qSV- and qSH -wave in tilted transversely isotropic media. *Geophysics*,

- 640 83(4), C153-C160.
- 641 Zhao, H. (2005). A fast sweeping method for eikonal equations. *Mathematics of*
 642 *Computation*, 74(250), 603-628.
- 643 Zhou, B., & Greenhalgh, S. (2004). On the computation of elastic wave group veloc-
 644 ities for a general anisotropic medium. *Journal of Geophysics and Engineering*,
 645 1(3), 205-215.
- 646 Zhou, B., & Greenhalgh, S. (2006). Raypath and travelttime computations for 2D
 647 transversely isotropic media with dipping symmetry axes. *Exploration Geo-*
 648 *physics*, 37(3), 150-159.

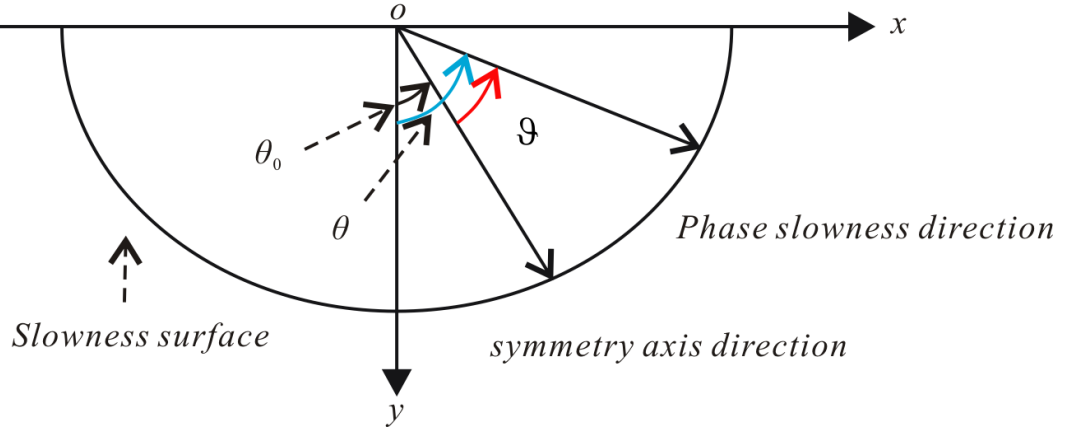


Figure 1. Definition of three different angles: (a) ϑ is the angle between the phase slowness direction and the symmetry axis direction. (b) θ is the angle between the phase slowness direction and the vertical y -axis. (c) θ_0 is the angle between the symmetry axis direction and the vertical y -axis.

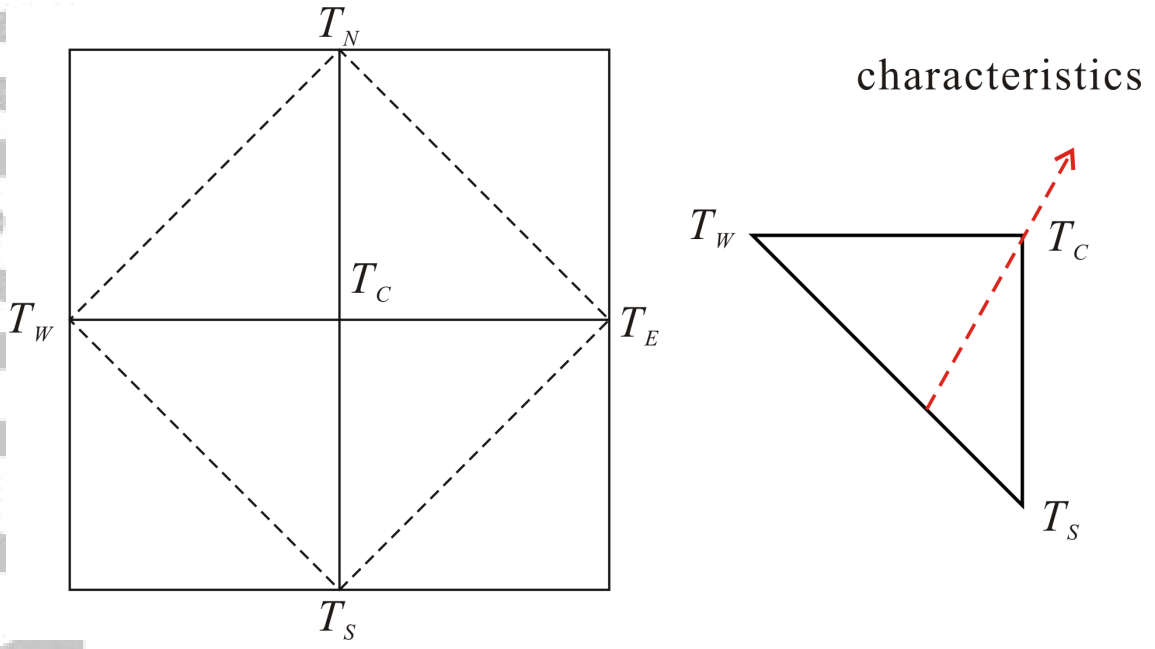


Figure 2. Rectangular mesh in the 2D case. Four triangles (ΔCEN , ΔCNW , ΔCWS and ΔCSE) are used to calculate traveltime candidates for the center grid point C .

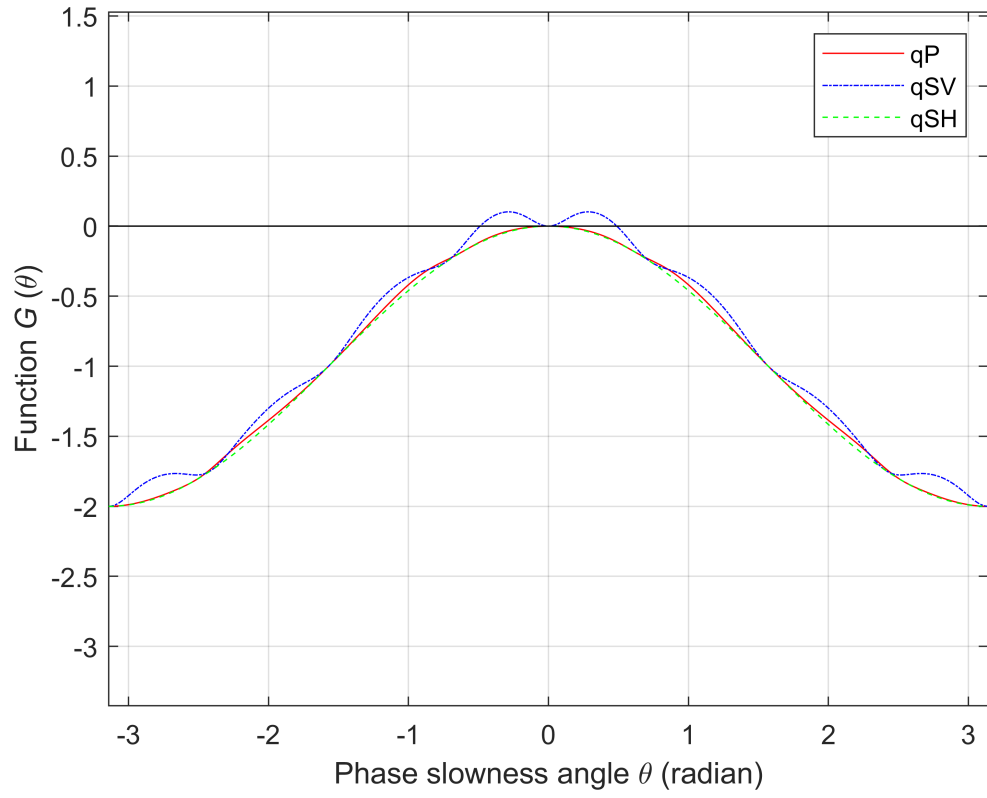


Figure 3. Function $G(\theta)$ for three wavemodes (qP-, qSV- and qSH-wave) in the homogeneous anisotropic model with the inclination angle $\theta_0 = 0^\circ$.

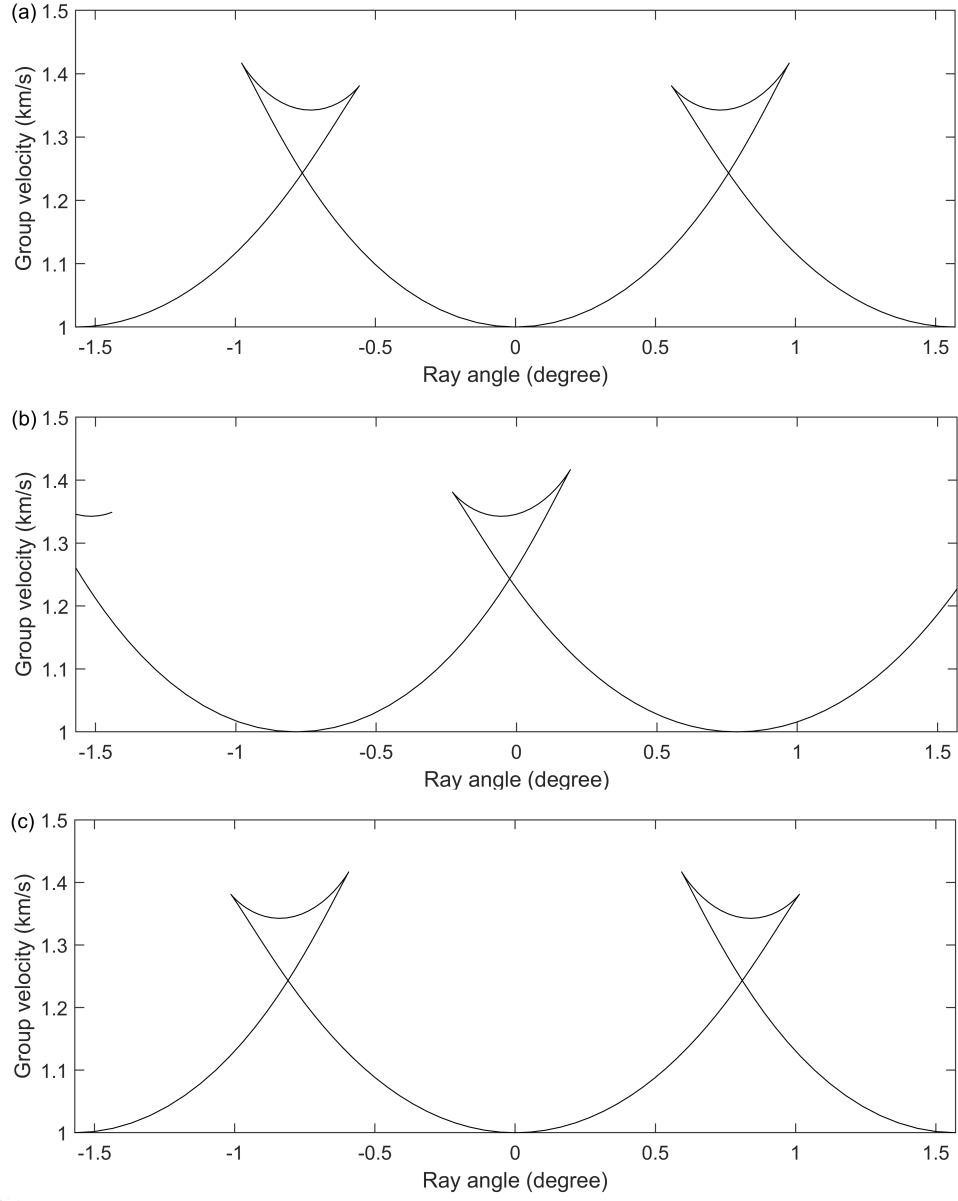


Figure 4. Group velocities of the qSV wavemode for three different orientations of the symmetry axis ($\theta_0 = 0^\circ, 45^\circ$ and 90°). The group velocity is plotted against the ray angle from $-\pi/2$ to $\pi/2$.

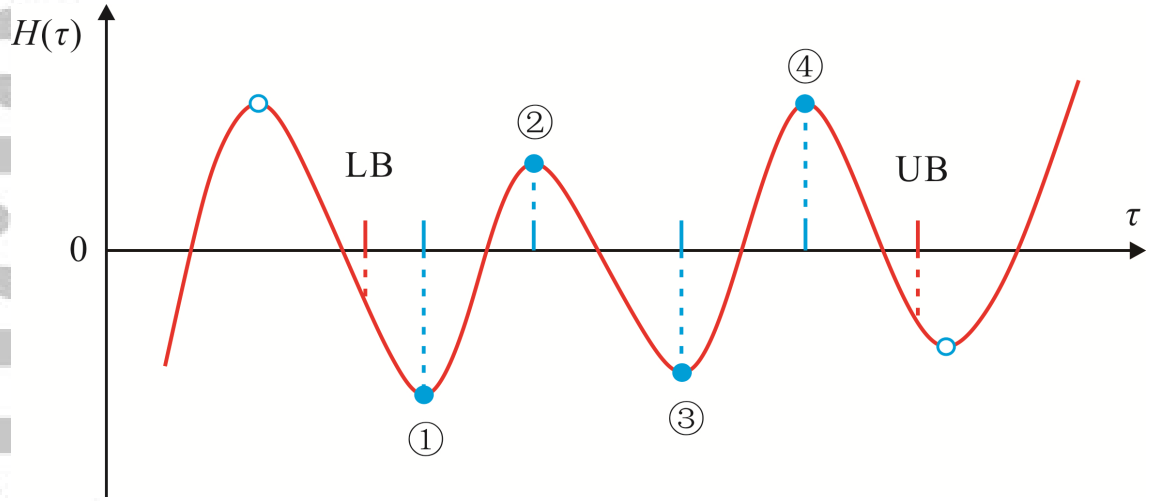


Figure 5. Extreme points of the eikonal equation partition the solution interval I_{WS} into subintervals. LB and UB are the lower and upper bounds of the solution interval, respectively. The blue solid circles represent the extreme points in the solution interval, while the blue hollow circles represent the extreme points outside of the solution interval.

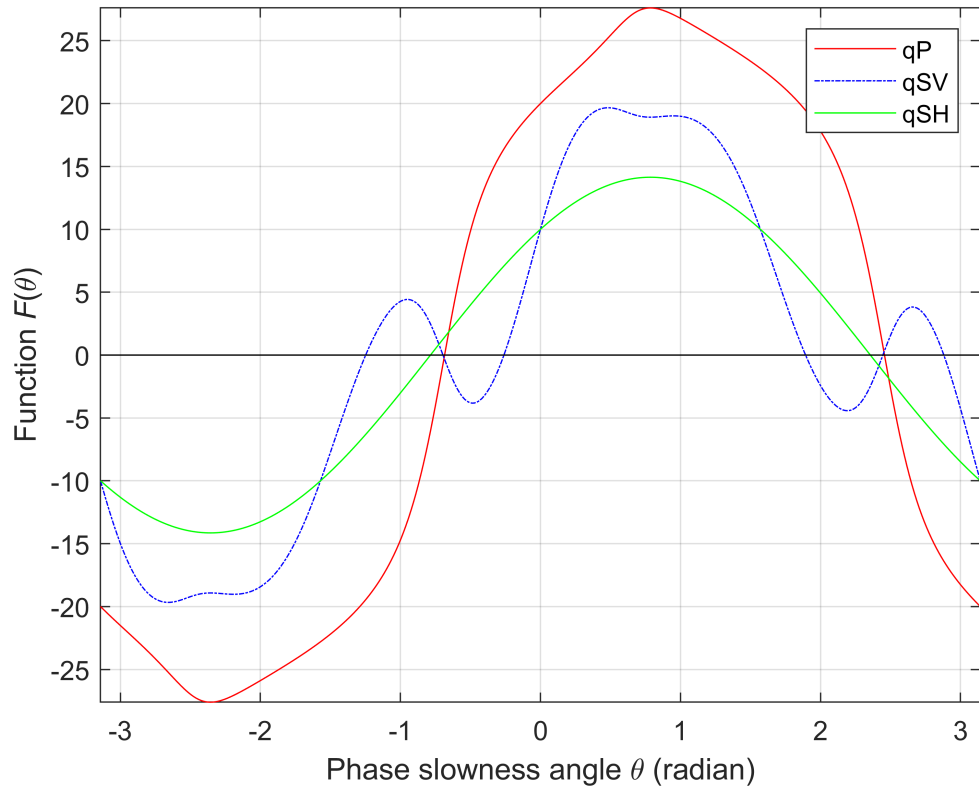


Figure 6. Curves of the critical point function $F(\theta)$ for the three wavemodes in homogeneous anisotropic model.

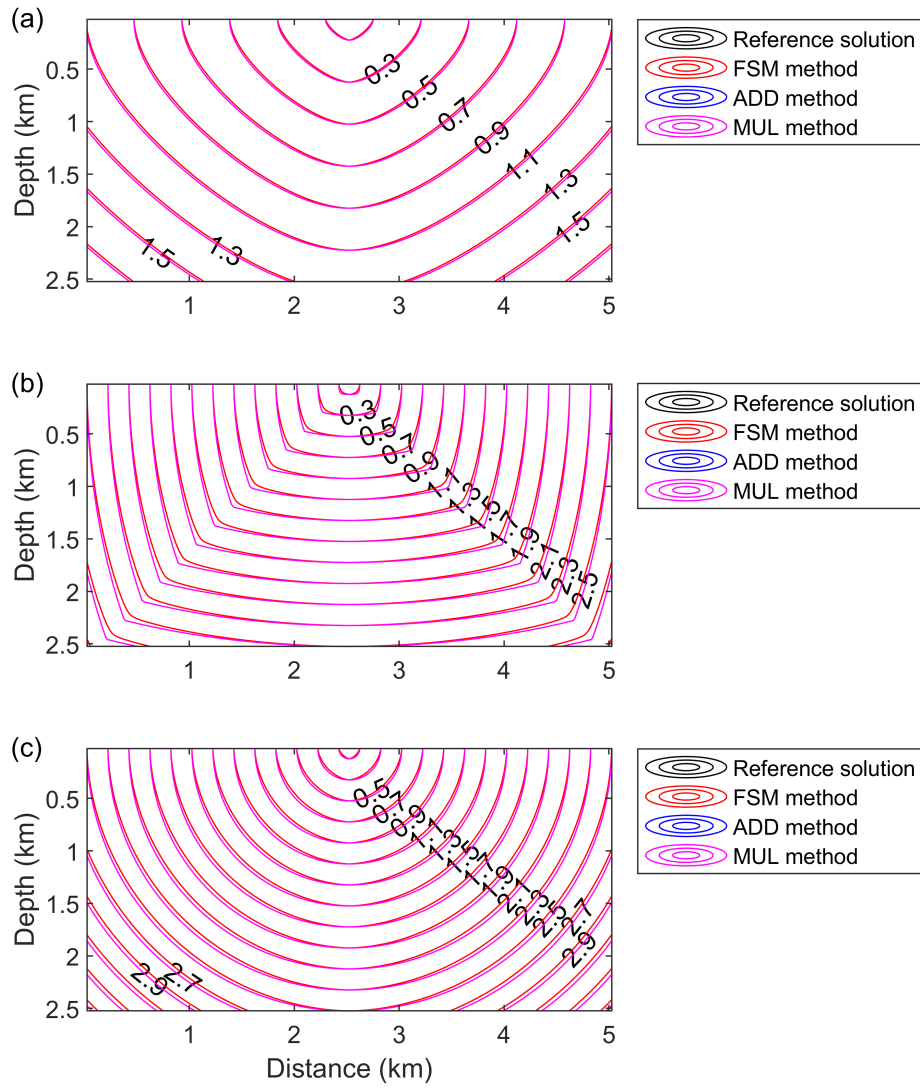


Figure 7. Traveltime comparison for the three wavemodes between the reference and numerical solutions in the homogeneous anisotropic model with $\theta_0 = 0^\circ$. Black contour line stands for the reference solution; Red contour line stands for the numerical solution calculated by the FSM method; Blue and magenta contour lines represent the numerical solutions calculated by the additively and multiplicatively factored FSM methods respectively.

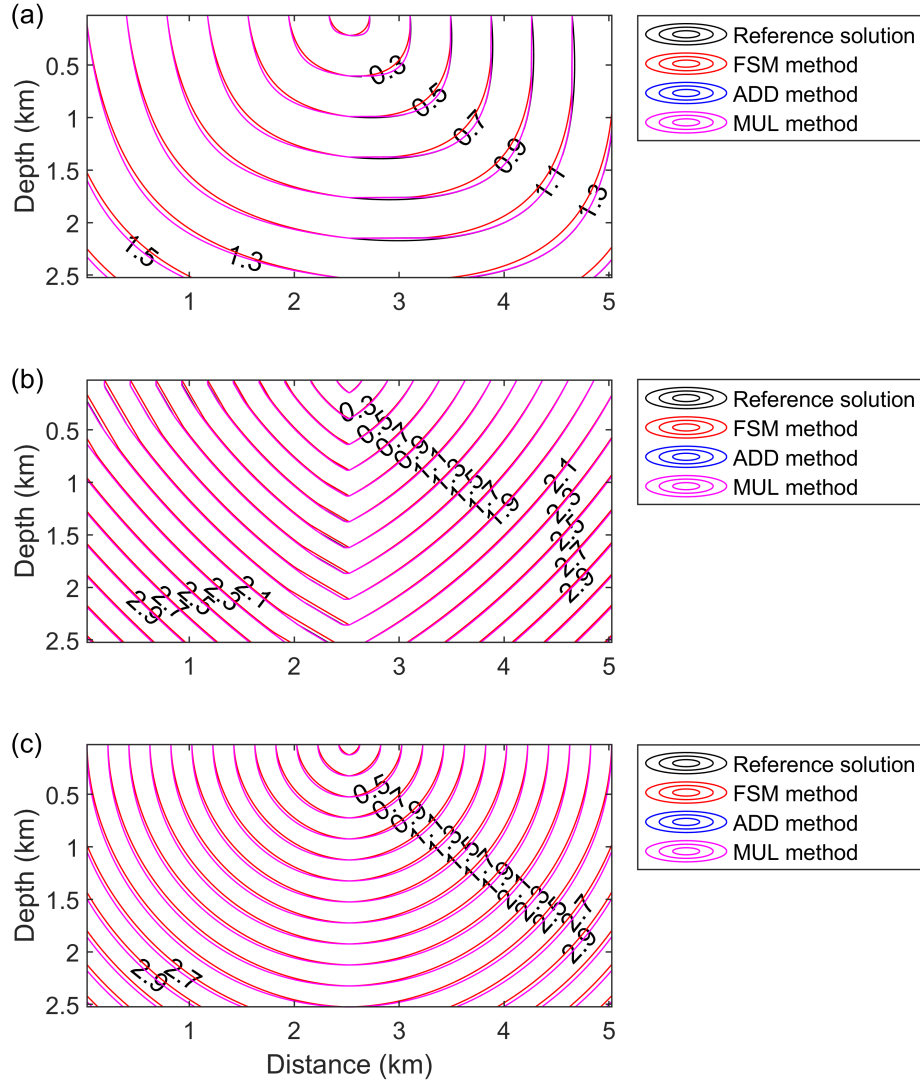


Figure 8. Traveltime comparison for the three wavemodes between the reference and numerical solutions in the homogeneous anisotropic model with $\theta_0 = 45^\circ$. Black contour line stands for the reference solution; Red contour line stands for the numerical solution calculated by the FSM method; Blue and magenta contour lines represent the numerical solutions calculated by the additively and multiplicatively factored FSM methods respectively.



Acc

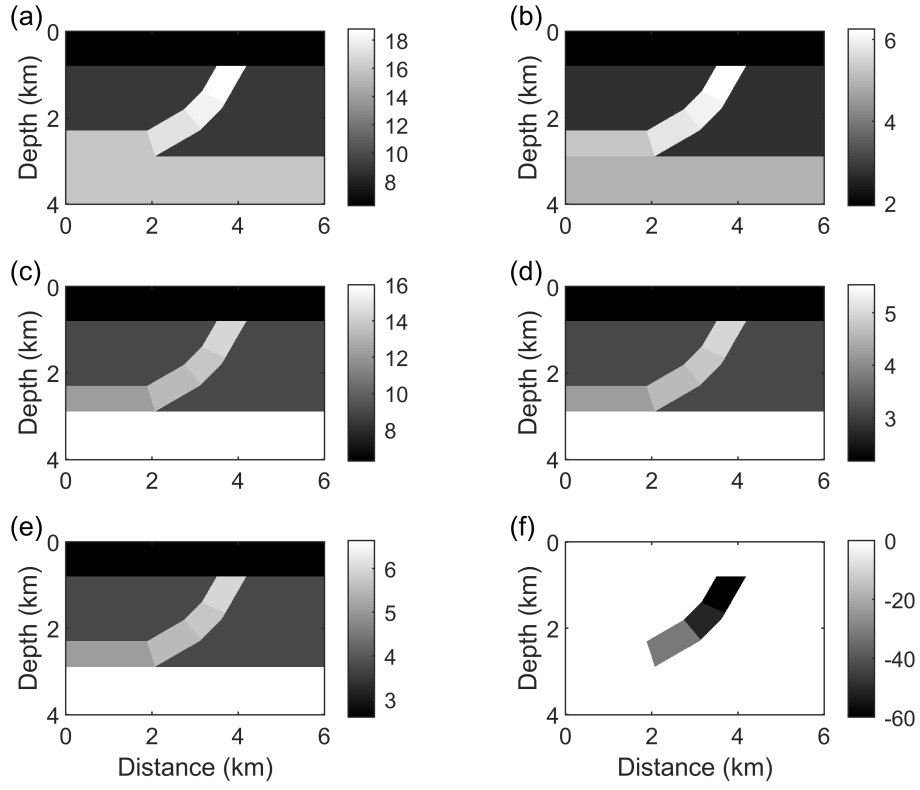


Figure 10. Anisotropic parameters of the overthrust TTI model. (a) a_{11} model, (b) a_{13} model, (c) a_{33} model, (d) a_{44} model, (e) a_{66} model, (f) θ_0 model.

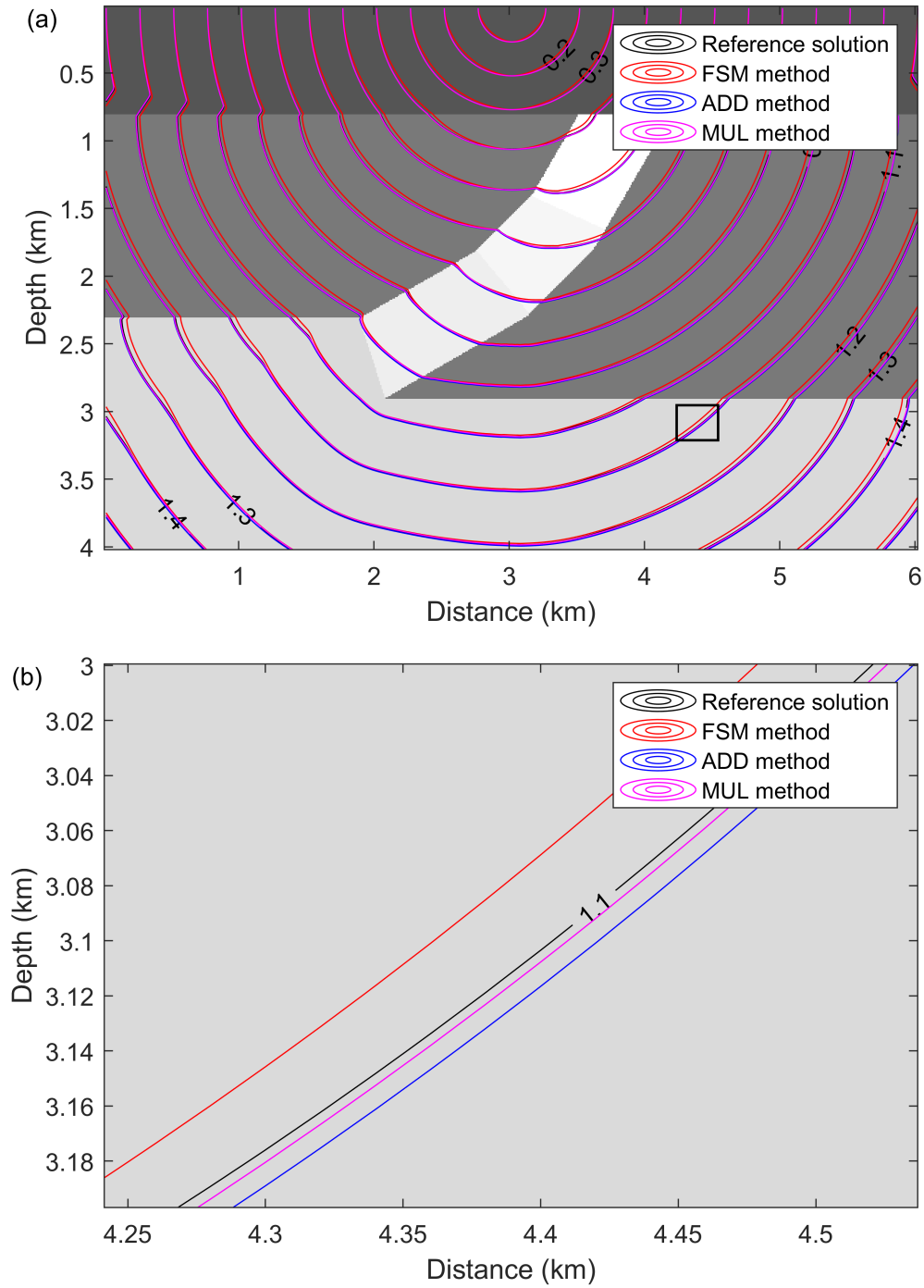


Figure 11. Traveltime comparison for the qP wave between the reference and numerical solutions in the overthrust TTI model. (a) Black contour line stands for the reference solution, red contour line stands for the numerical solution calculated by the FSM method, blue and magenta contour lines represent the numerical solutions generated by the additively and multiplicatively factored FSM methods respectively. (b) Zoom-in map of the square area as shown in (a), from which we can see traveltimes solved by the factored methods have better accuracy.

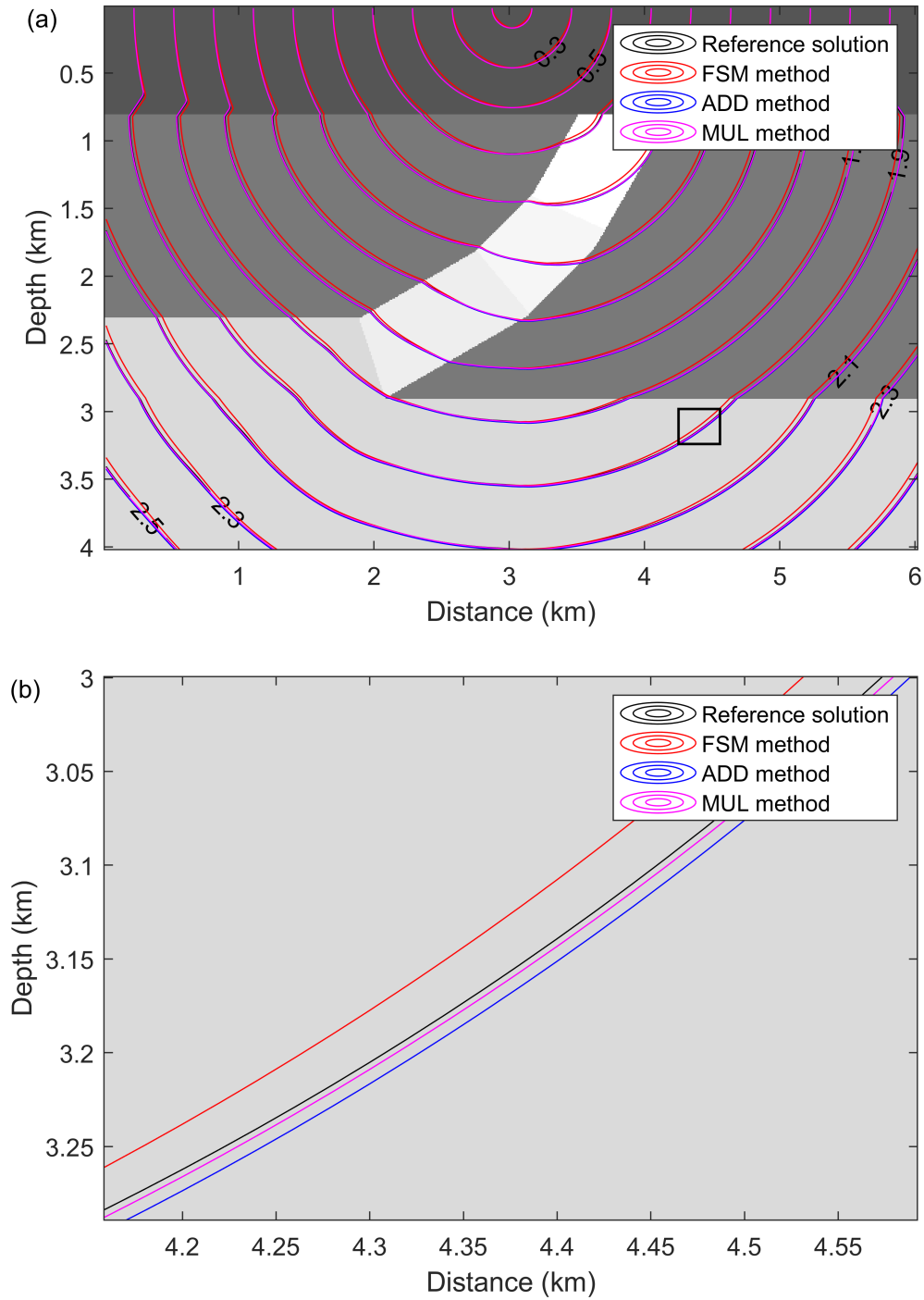


Figure 12. Traveltime comparison for the qSV wave between the reference and numerical solutions in the overthrust TTI model. (a) Black contour line stands for the reference solution, red contour line stands for the numerical solution calculated by the FSM method, blue and magenta contour lines represent the numerical solutions generated by the additively and multiplicatively factored FSM methods respectively. (b) Zoom-in map of the square area as shown in (a), from which we can see traveltimes solved by the factored methods have better accuracy.

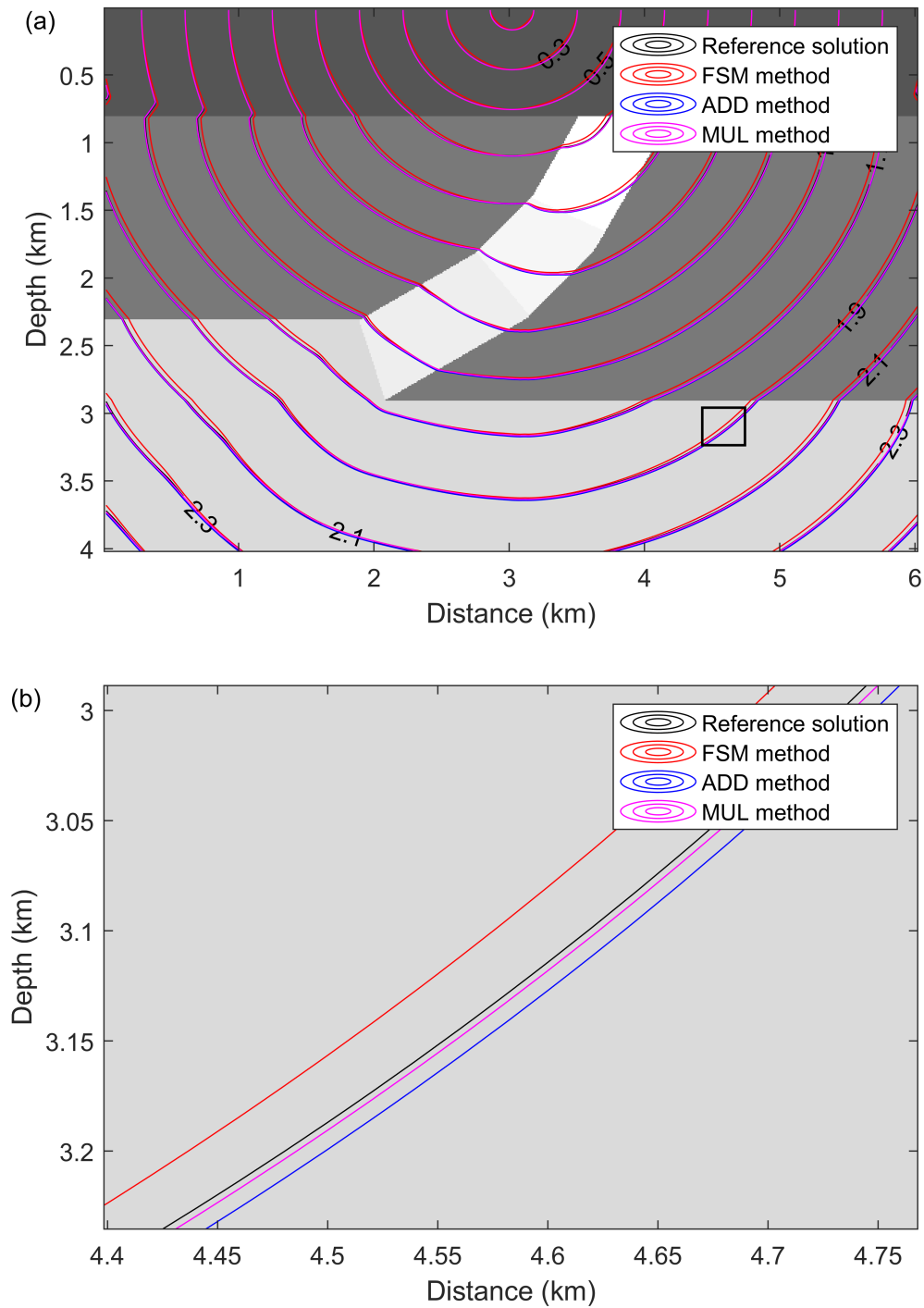


Figure 13. Traveltime comparison for the qSH wave between the reference and numerical solutions in the overthrust TTI model. (a) Black contour line stands for the reference solution, red contour line stands for the numerical solution calculated by the FSM method, blue and magenta contour lines represent the numerical solutions generated by the additively and multiplicatively factored FSM methods respectively. (b) Zoom-in map of the square area as shown in (a), from which we can see traveltimes solved by the factored methods have better accuracy.

Table 1. Accuracy of the first-order FSM method in the homogeneous anisotropic model.

qP wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
101×51	1	0.0138	-	0.7
201×101	1	0.0082	0.7510	1.0
401×201	1	0.0047	0.8030	4.0
801×401	1	0.0027	0.8090	20.0
qSV wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
101×51	1	0.0272	-	1.0
201×101	1	0.0153	0.8301	2.0
401×201	1	0.0084	0.8651	8.0
801×401	1	0.0046	0.8789	29.0
qSH wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
101×51	1	0.0369	-	1.0
201×101	1	0.0221	0.7396	1.0
401×201	1	0.0129	0.7767	6.0
801×401	1	0.0074	0.8111	25.0

Table 2. Accuracy of the first-order FSM method in the overthrust TTI model.

qP wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0181	-	1.0
151×101	2	0.0101	0.8416	2.0
301×201	2	0.0061	0.7275	10.0
qSV wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0298	-	0.8
151×101	2	0.0173	0.7845	2.0
301×201	2	0.0106	0.7067	9.0
qSH wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0283	-	1.0
151×101	2	0.0161	0.8137	2.0
301×201	2	0.0099	0.7016	10.0

Table 3. Accuracy of the first-order additively factored FSM method in the overthrust TTI model.

qP wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0058	-	1.0
151×101	1	0.0030	0.9511	2.0
301×201	2	0.0015	1.0000	11.0
qSV wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0103	-	2.0
151×101	2	0.0053	0.9586	5.0
301×201	2	0.0026	1.0275	25.0
qSH wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0093	-	2.0
151×101	2	0.0051	0.8667	8.0
301×201	2	0.0025	1.0286	23.0

Table 4. Accuracy of the first-order multiplicatively factored FSM method in the overthrust TTI model.

qP wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0034	-	0.8
151×101	1	0.0014	1.2801	3.0
301×201	2	5.6251e-04	1.3155	11.0
qSV wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0045	-	1.0
151×101	2	0.0020	1.1699	3.0
301×201	2	7.5947e-04	1.3969	13.0
qSH wavemode				
Mesh	Iteration	L_1 error	Convergence order	Time cost (seconds)
76×51	1	0.0054	-	1.0
151×101	2	0.0022	1.2955	4.0
301×201	2	9.6072e-04	1.1953	15.0