



Site U1600¹

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Core descriptions

Supplementary material

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S. Kutterolf, T.H. Druitt, T.A. Ronge, S. Beethe, A. Bernard, C. Berthod, H. Chen, S. Chiyonobu, A. Clark, S. DeBari, T.I. Fernandez Perez, R. Gertisser, C. Hübscher, R.M. Johnston, C. Jones, K.B. Joshi, G. Kletetschka, O. Koukousioura, X. Li, M. Manga, M. McCanta, I. McIntosh, A. Morris, P. Nomikou, K. Pank, A. Peccia, P.N. Polymenakou, J. Preine, M. Tominaga, A. Woodhouse, and Y. Yamamoto²

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² **Expedition 398 Scientists' affiliations.**

1. Background and objectives

Site U1600 is located 10 km south of Anhydros Island within a small graben atop the Anhydros Horst (Figure F1). The Anhydros Horst separates the Anhydros Basin to the west from the Anafi Basin to the east (Preine et al., 2022a, 2022b). The water depth is 326 meters below sea level (mbsl). Permission to drill in this location was requested as Site CSK-24A and granted by the International Ocean Discovery Program (IODP) Environmental Protection and Safety Panel during the expedition. Three holes (U1600A–U1600C) were drilled for a total recovery depth of 184.2 meters below seafloor (mbsf) (all depths below seafloor are given using the core depth below seafloor, Method A [CSF-A] scale, except in **Operations**, where the drilling depth below seafloor [DSF] scale is used), with average recoveries ranging 32%–75%.

The site was chosen because of its situation on the Anhydros Horst immediately east of the Kolumbo chain of volcanoes and for the well-stratified nature of the graben fill on seismic profiles

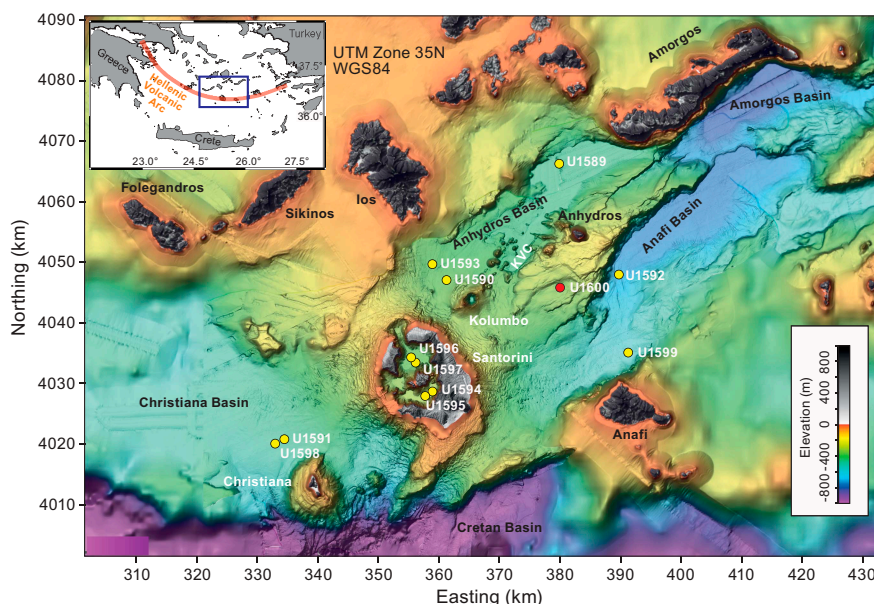


Figure F1. Site map. Red = Site U1600, yellow = other sites. Inset = location map. See Figure F1 in the Site U1589 chapter (Druitt et al., 2024a) for the swath data on which this map is based. KVC = Kolumbo volcanic chain.

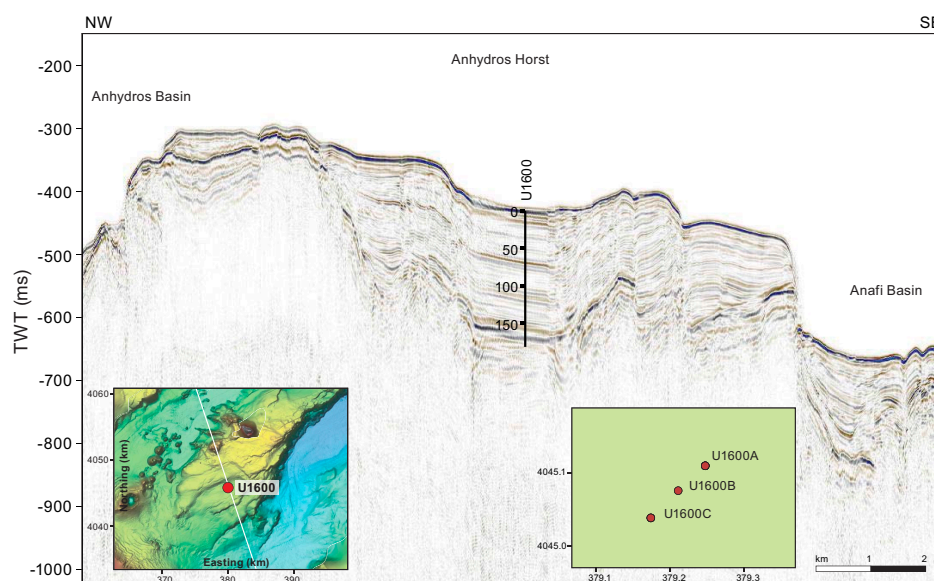


Figure F2. Profile across the Anhydros Horst along Seismic Line HH06-17. Insets: location of Site U1600 and Holes U1600A–U1600C. Depths in meters. TWT = two-way traveltime.

(Figure F2). It seemed to be a likely site at which to drill a condensed sequence of muds and tephra for chronology, sheltered from the large-scale mass wasting of the main basins. Site U1600 is located within the area of the 2015 PROTEUS seismic tomography experiment, during which sub-bottom profiling, gravity, and magnetic data were also recorded (Hooft et al., 2017).

Drilling at Site U1600 provided the possibility of reconstructing a near-complete volcanic stratigraphy consistent with both onshore and offshore constraints and pinned by chronological markers from biostratigraphy, magnetostratigraphy, and sapropel records. Benthic foraminifera from fine-grained sediments provided estimates of paleowater depths and, via integration with seismic profiles and chronologic data, of time-integrated basin subsidence rates. Drilling on the Anhydros Horst addressed scientific Objectives 1–4 and 6 of the Expedition 398 Scientific Prospectus (Druitt et al., 2022).

2. Operations

The vessel arrived on location at Site U1600 on 29 January 2023 at 0430 h. The vessel was switched to full dynamic positioning mode and was on auto control at 0450 h, beginning Site U1600.

2.1. Hole U1600A

With the all clear given, the rig crew started assembling the advanced piston corer/extended core barrel (APC/XCB) bottom hole assembly (BHA) with the same bit used previously. Hole U1600A (36°32.6277'N, 25°39.0553'E) was spudded with Core 1H on 29 January 2023 at 0710 h. The recovery of 4.0 m gave a calculated seafloor of 326.2 mbsl (Table T1). APC coring continued with Cores 2H–4H. Prior to Core 4H, two misfires (pins did not shear) were experienced with the APC system, and an XCB barrel was run as a deplugger. Core 4H was fired on the third attempt with good recovery but with 45,000 lb overpull and a completely fractured liner.

The switch was made to the half-length APC (HLAPC) system with Cores 5F–10F at 56 mbsf. Core 10F required overpull of 50,000 lb. Coring continued with Cores 11X–13X to 84.4 mbsf, the final depth for Hole U1600A. The recovery for the XCB coring was extremely poor, and the decision was made to repeat the piston section in a new hole. The drill string was pulled up with the top drive, clearing the seafloor at 1905 h and ending Hole U1600A.

Table T1. Core summary, Site U1600. mbsf = meters below seafloor. NA = not applicable. DSF = drilling depth below seafloor. H = APC, F = HLAPC, X = XCB, numeric core type = drilled interval, R = RCB. (Continued on next page.) [Download table in CSV format.](#)

Hole U1600A							
Latitude: 36°32.6277'N							
Longitude: 25°39.0553'E							
Water depth (m): 326.18							
Date started (UTC): 0245 h; 29 January 2023							
Date finished (UTC): 1700 h; 29 January 2023							
Time on hole (days): 0.59							
Penetration (mbsf): 84.4							
Cored interval (m): 84.4							
Recovered length (m): 51.71							
Recovery (%): 61.27							
Drilled interval (m): NA							
Drilled interval (no.): 0							
Total cores (no.): 13							
APC cores (no.): 4							
HLAPC cores (no.): 6							
XCB cored (no.): 3							
Hole U1600B							
Latitude: 36°32.6092'N							
Longitude: 25°39.0311'E							
Water depth (m): 326.34							
Date started (UTC): 1700 h; 29 January 2023							
Date finished (UTC): 1000 h; 30 January 2023							
Time on hole (days): 0.6							
Penetration (mbsf): 91.4							
Cored interval (m): 91.4							
Recovered length (m): 68.67							
Recovery (%): 75.13							
Drilled interval (m): NA							
Drilled interval (no.): 0							
Total cores (no.): 19							
APC cores (no.): 1							
HLAPC cores (no.): 18							
Hole U1600C							
Latitude: 36°32.5890'N							
Longitude: 25°39.0066'E							
Water depth (m): 326.34							
Date started (UTC): 0200 h; 4 February 2023							
Date finished (UTC): 0400 h; 5 February 2023							
Time on hole (days): 1.08							
Penetration (mbsf): 188.5							
Cored interval (m): 113.5							
Recovered length (m): 36.78							
Recovery (%): 32.41							
Drilled interval (m): 75							
Drilled interval (no.): 1							
Total cores (no.): 13							
RCB cores (no.): 13							
Core	Top depth drilled DSF (m)	Bottom depth drilled DSF (m)	Interval advanced (m)	Recovered length (m)	Core recovery (%)	Core on deck date (2023)	Core on deck time UTC (h)
398-U1600A-							
1H	0.0	4.0	4.0	4.08	102	29 Jan	0515
2H	4.0	13.5	9.5	9.64	101	29 Jan	0600
3H	13.5	23.0	9.5	6.58	69	29 Jan	0630
4H	23.0	32.5	9.5	8.74	92	29 Jan	0815
5F	32.5	37.2	4.7	4.52	96	29 Jan	0940
6F	37.2	41.9	4.7	2.74	58	29 Jan	1000
7F	41.9	46.6	4.7	1.56	33	29 Jan	1035
8F	46.6	51.3	4.7	4.43	94	29 Jan	1105
9F	51.3	56.0	4.7	4.70	100	29 Jan	1200
10F	56.0	60.7	4.7	4.66	99	29 Jan	1225
11X	60.7	68.7	8.0	0.00	0	29 Jan	1345
12X	68.7	74.7	6.0	0.00	0	29 Jan	1445
13X	74.7	84.4	9.7	0.06	1	29 Jan	1545
398-U1600B-							
1H	0.0	6.8	6.8	6.83	100	29 Jan	1820
2F	6.8	11.5	4.7	4.66	99	29 Jan	1920
3F	11.5	16.2	4.7	4.10	87	29 Jan	1950
4F	16.2	20.9	4.7	3.66	78	29 Jan	2015
5F	20.9	25.6	4.7	4.71	100	29 Jan	2050
6F	25.6	30.3	4.7	4.28	91	29 Jan	2125
7F	30.3	35.0	4.7	2.26	48	29 Jan	2155
8F	35.0	39.7	4.7	2.69	57	29 Jan	2230
9F	39.7	44.4	4.7	0.00	0	29 Jan	2300
10F	44.4	49.1	4.7	0.58	12	29 Jan	2330
11F	49.1	53.8	4.7	4.84	103	30 Jan	0005
12F	53.8	58.5	4.7	4.72	100	30 Jan	0035
13F	58.5	63.2	4.7	4.44	94	30 Jan	0120
14F	63.2	67.9	4.7	4.70	100	30 Jan	0150
15F	67.9	72.6	4.7	4.47	95	30 Jan	0220
16F	72.6	77.3	4.7	1.86	40	30 Jan	0310
17F	77.3	82.0	4.7	4.76	101	30 Jan	0340
18F	82.0	86.7	4.7	0.16	3	30 Jan	0400
19F	86.7	91.4	4.7	4.95	105	30 Jan	0425

Table T1 (continued).

Core	Top depth drilled DSF (m)	Bottom depth drilled DSF (m)	Interval advanced (m)	Recovered length (m)	Core recovery (%)	Core on deck date (2023)	Core on deck time UTC (h)
398-U1600C-							
11	0.0	75.0	75.0	****Drilled interval****		4 Feb	0805
2R	75.0	84.7	9.7	0.00	0	4 Feb	0930
3R	84.7	94.4	9.7	0.17	2	4 Feb	1035
4R	94.4	104.1	9.7	0.93	10	4 Feb	1130
5R	104.1	113.8	9.7	1.43	15	4 Feb	1250
6R	113.8	123.5	9.7	3.39	35	4 Feb	1400
7R	123.5	133.2	9.7	4.20	43	4 Feb	1500
8R	133.2	142.9	9.7	1.74	18	4 Feb	1620
9R	142.9	152.6	9.7	9.03	93	4 Feb	1730
10R	152.6	162.3	9.7	9.37	97	4 Feb	1825
11R	162.3	172.0	9.7	3.77	39	4 Feb	1920
12R	172.0	179.7	7.7	1.15	15	4 Feb	2130
13R	179.7	183.7	4.0	1.16	29	4 Feb	2255
14R	183.7	188.5	4.8	0.44	9	5 Feb	0020
Totals:			364.3	157.16			

2.2. Hole U1600B

The ship was positioned 50 m southwest of Hole U1600A. Hole U1600B (36°32.6092'N, 25°39.0311'E) was spudded with Core 1H on 29 January 2023 at 2005 h. Recovery of 6.8 m established the water depth at 326.3 mbsl. Coring was switched to the HLAPC system from Core 2F and continued on 30 January through Core 19F at 91.4 mbsf, the final depth for Hole U1600B.

The drill string was tripped up with the top drive, clearing the seafloor at 0735 h. The bit cleared the rotary table at 0900 h, and the drill floor was secured for transit.

The crew started raising the thrusters at 0954 h, and the vessel was under bridge control at 0956 h. All thrusters were up and secured, and the sea passage started at 1000 h, ending Hole U1600B.

2.3. Hole U1600C

The vessel returned to Site U1600 on 4 February 2023. Once on site, the rig crew assembled the rotary core barrel (RCB) BHA with a used bit. Hole U1600C (36°32.5890'N, 25°39.0067'E) was spudded at 0655 h. A drill-ahead was done to 75.0 mbsf. RCB coring was initiated with Core 2R from 75.0 mbsf through Core 14R at 188.5 mbsf on 5 February, the final depth for Hole U1600C. After confirming that the basement was reached, the string was tripped up with the top drive, and the bit cleared the seafloor at 0410 h and the rig floor at 0525 h. The drill floor was secured at 0548 h, the vessel was switched to bridge control, and the thrusters were raised starting at 0550 h. All thrusters were up and secure, starting the sea passage at 0600 h and ending Site U1600.

Hole U1600C took 26.00 h (1.1 days).

3. Lithostratigraphy

Cores from Site U1600 recovered a relatively coherent sedimentary stratigraphy from 0 to 184.24 mbsf (Sections 398-U1600A-1H-1 through 13X-CC [0–74.73 mbsf]; 398-U1600B-1H-1 through 19F-CC [0–91.65 mbsf]; and 398-U1600C-3R-CC through 14R-1) (Figure F3).

The recovered material is sedimentary and unlithified in Holes U1600A and U1600B, with a gradual transition to stiffer and more consolidated material toward the bottom of Hole U1600B. Hole U1600C consists of more consolidated sediments, but paradoxically this is not supported by higher *P*-wave velocity (see [Physical properties](#)). However, a change from sediment to sedimentary rock was clearly identified at ~166 mbsf; hence, the lithology name changes from sediment (e.g., sand) to sedimentary rock (e.g., sandstone).

X-ray diffraction (XRD) data were obtained from three interstitial water (IW) squeeze cake sediment residues from Hole U1600A and five each from Holes U1600B and U1600C.

The sedimentary succession in the uppermost 92 m is dominated by volcanic and tuffaceous lithologies with minor intermittent intervals of nonvolcanic muds and sands (Lithostratigraphic Unit I). All cores in Unit I are rich in bioclasts (shells and shell fragments), even the volcanic intervals. In cases where ash or lapilli intervals have >25% bioclasts, they are denoted by the name bioclastic ash, and where bioclasts are present in abundances <25% that name becomes “with shells.” There are few ooze intervals in Unit I; most tuffaceous and nonvolcanic sediments are muds and sands, and the criteria for the “bioclastic” prefix and the “with shells” suffix is the same as described above. Smear slides for microscopic analyses were prepared to confirm macroscopic descriptions of distinct lithology changes at the section level, such as identification of vitric ash particles in tuffaceous lithologies or crystals in ash layers.

Unit II (77.30–172 mbsf), observed in the lowermost 14 m of Hole U1600B and in the uppermost ~87 m of Hole U1600C, is a continuation of the calcareous muds and sands present in Unit I but with three important distinctions: (1) these sediments are devoid of any volcanic and tuffaceous intervals, whereas Unit I in Holes U1600A and U1600B contains abundant volcanic and tuffaceous intervals; (2) the sediments in Unit II are almost entirely organic rich; and (3) XRD data show significant dolomite (see [X-ray diffraction](#)) that was not observable macroscopically. Given these differences, all of Hole U1600C is designated as Unit II down to a sharp discontinuity with ultramafic basement rock at 172.1 mbsf. These basement rocks are designated as Unit III.

Figure [F3](#) summarizes the lithostratigraphy of Site U1600, displaying core recovery and lithostratigraphic units and subunits in Holes U1600A–U1600C. Table [T2](#) provides the upper and lower boundaries, thicknesses, biostratigraphic ages (see [Biostratigraphy](#)), and lithologic summary of the units and subunits. Figure [F4](#) graphically presents the relative proportions of volcanic, tuffaceous, and nonvolcanic lithologies with lithostratigraphic units and subunits. Figure [F5](#) presents the grain size distribution of the sediments in all units and subunits, in particular those within the volcanic- and tuffaceous-dominated subunits of Unit I, to graphically show the distribution of ash, lapilli-ash, and lapilli and the variation between mud and sand in the lowermost part of Unit I and in Unit II. Figure [F6](#) displays different types of core disturbance observed at Site U1600.

The following sections describe (1) the effects of core disturbance; (2) the lithostratigraphic units and subunits; (3) the correlations between Holes U1600A, U1600B, and U1600C; and (4) XRD results from Site U1600.

3.1. Core disturbance

Several types of core disturbance disrupt the lithostratigraphy at Site U1600 (Figure [F6](#)):

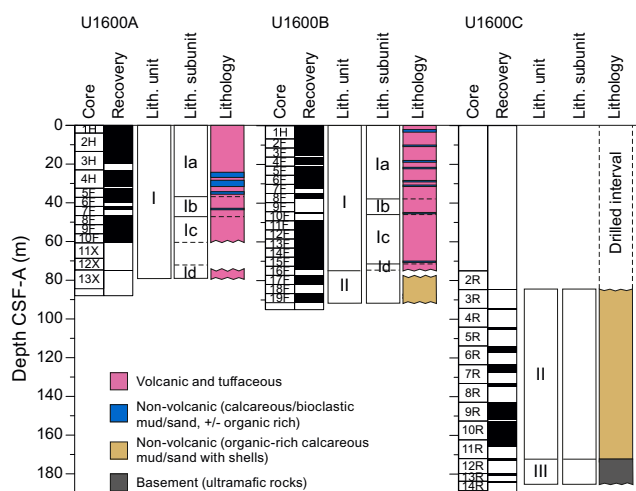
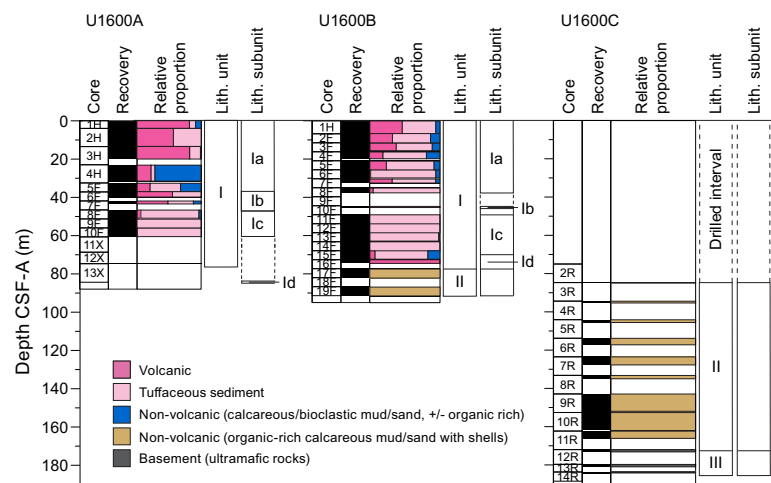
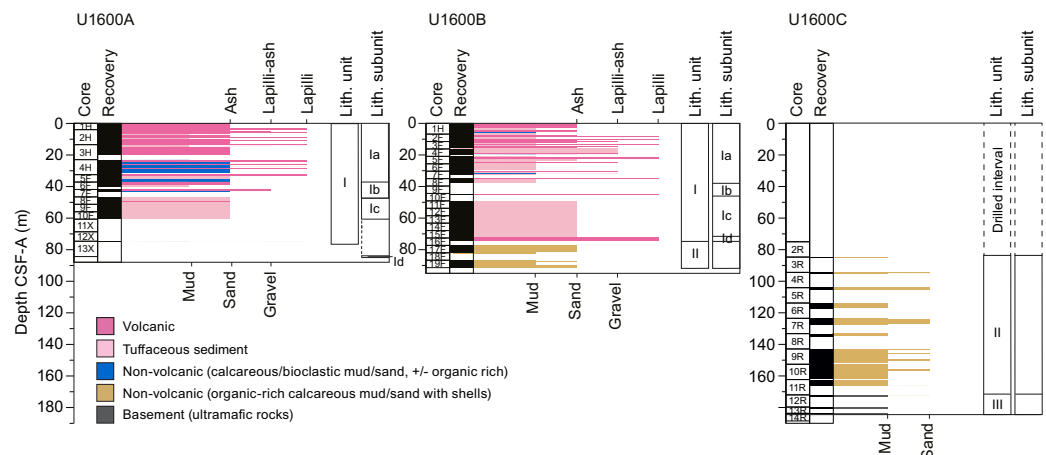


Figure F3. Lithostratigraphy, Site U1600. Unit color = dominant lithology.

Table T2. Lithostratigraphic units and subunits, Holes U1600A–U1600C. * = bottom not recovered or end of hole. [Download table in CSV format.](#)

Lith. unit	Subunit	Top depth (mbsf)	Top hole, core, section, interval (cm)	Bottom depth (mbsf)	Bottom hole, core, section, interval (cm)	Thickness (m)	Lithologic summary
I	1a	398-0.00	U1600A-1H-1, 0	398-37.00	U1600A-5F-CC, 8	37.00	Tuffaceous mud, ash, ash with lapilli, lapilli-ash, and calcareous tuffaceous mud
		0.00	U1600B-1H-1, 0	37.67*	U1600B-8F-2, 117	37.67	
	1b	37.20	U1600A-6F-1, 0	46.74	U1600A-8F-1, 14	9.54	Bioclastic ash and lapilli
		44.70	U1600B-10F-1, 0	45.28*	U1600B-10F-CC, 18	0.58	
	1c	46.74	U1600A-8F-1, 14	60.61	U1600A-10F-CC, 10	13.87	Calcareous tuffaceous sand with shells
		49.10	U1600B-11F-1, 0	72.03	U1600B-15F-3, 113	22.93	
	1d	74.70	U1600A-13X-CC, 0	74.73*	U1600A-13X-CC, 3	0.03	Lapilli
		72.03	U1600B-15F-3, 113	74.46*	U1600B-16F-CC, 9	2.43	
	II	—	—	—	—	—	Organic-rich calcareous mud/sand with shells, calcareous mud/sand, sandstone, and rare ash pods
		77.30	U1600B-17F-1, 0	91.63*	U1600B-19F-CC, 16	14.33	
III		84.70	U1600C-3R-CC, 0	172.10	U1600C-12R-1, 10	87.40	Ultramafic rocks (basement)
		172.10	U1600C-12R-1, 10	184.24*	U1600C-14R-1, 54	12.14	

**Figure F4.** Relative percentages of volcanic, tuffaceous, and nonvolcanic lithologies, Site U1600. Unit I is volcanic and tuffaceous, and Units II and III are nonvolcanic dominated.**Figure F5.** Grain size distributions of volcanic, tuffaceous, and nonvolcanic sediments. Length of colored bars = relative grain size (ash = <2 mm; lapilli = 2–64 mm; mud = <63 µm; sand = 0.063–2 mm), with separate scales shown for volcanic grain size (top) and nonvolcanic grain size (bottom; used for tuffaceous and nonvolcanic sediments). Mixed lithologies such as lapilli-ash (dark pink) that have relative grain sizes between two categories are plotted between ticks.

- Uparching: slight to moderate coring-induced shear between the sediment and core liner results in bedding uniformly dragged downward along the core margins (Figure F6A). In these intervals, original lithofacies and sedimentary structures are usually slightly to severely disturbed but can still be recognized visually.
- Fall-in: granular material from the top of the hole may fall in and accumulate at the bottom until the next core is recovered (Figure F6B). We exclude these intervals from lithostratigraphy interpretations.
- Mixed sediment: in water-saturated, granular core sections, grains and clasts flow and mix, producing moderately to severely disturbed, partly overprinted, sedimentary structures and stratigraphy (Figure F6C).
- Sediment flowage: high shearing rates between cored sediments and the core liner leave a smear or thin trail of displaced sediment along the inside of the core liner (Figure F6D).
- Special variety of midcore flow-in: deformation (shearing and truncation) of mud caused by underpressure stress during retrieval of the core from the host sediments is aggravated by midcore flow-in of allochthonous, dark sandy sediment injected between the segmented mud units (Figure F6E; Jutzeler et al., 2014).
- Cracks: slight drilling-related brecciation produces cracks in the original lithologies (Figure F6F).
- Brecciated: more intense drilling-related brittle rock failure disturbs the original lithofacies and sedimentary structures more severely, although they usually remain recognizable (Figure F6G).
- Biscuiting: fractured disc-shaped pieces, ranging in size from a few to more than 10 cm, often packed with sheared and remolded core material mixed with drill slurry, filling gaps between brittle biscuits (Figure F6H). The degree of biscuiting ranges from slight to severe, depending on the size of the biscuits and the proportion of biscuits to infill material.

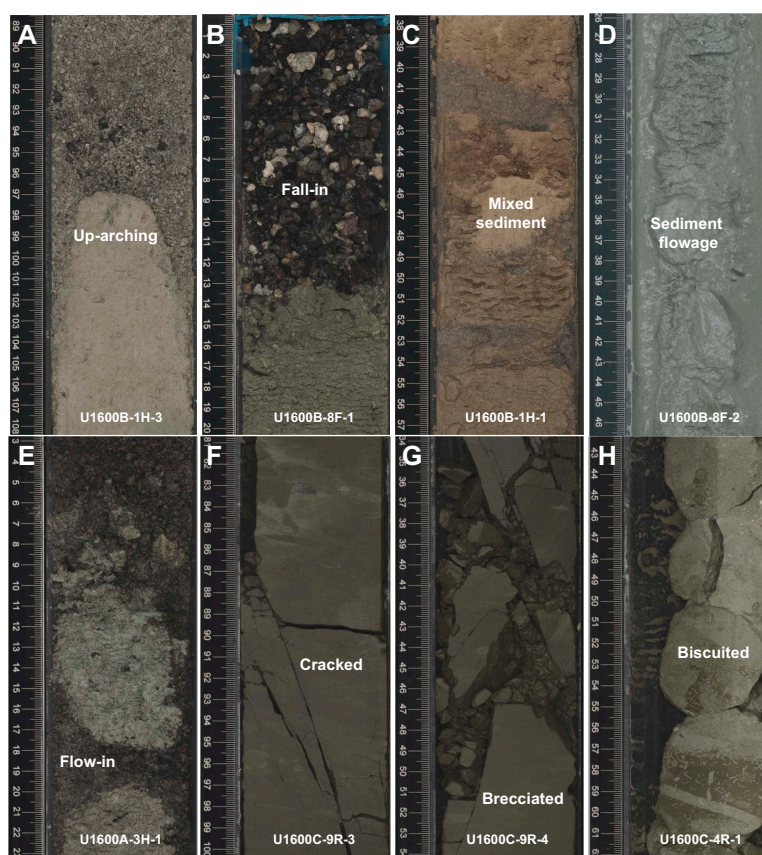


Figure F6. Core disturbances, Site U1600. A. Uparching. B. Fall-in. C. Mixed sediments. D. Sediment flowage. E. Midcore flow-in. F. Cracks. G. Brecciation. H. Biscuiting.

- Artificial size and density segregation: likely to occur during drilling, with postrecovery core handling processes on board (e.g., inclining, shaking, and plunging cores on the catwalk to compact sediments), or while the core is lying flat on deck (Jutzeler et al., 2014). Such core disturbance is observed most often in volcanic sediments because increased porosity allows sucking in of seawater during hydraulic piston coring. The resulting soupy texture allows material to flow within the core liner. Secondary normal or reverse grading or density separation of clasts may occur as a result of this disturbance and obscure primary sedimentary features.

3.2. Description of units and subunits

3.2.1. Unit I

Intervals: 398-U1600A-1H-1, 0 cm, to 13X-CC (bottom of the hole); 398-U1600B-1H-1, 0 cm, to 16F-CC

Depths: Hole U1600A = 0–74.73 mbsf; Hole U1600B = 0–74.46 mbsf

Thicknesses: Hole U1600A = 74.73 m; Hole U1600B = 74.46 m

Age: Holocene to Early Pleistocene

Lithology: volcanic sediment (ash, lapilli-ash, and lapilli) and tuffaceous sediment (calcareous tuffaceous mud and sand) with lesser nonvolcanic sediment (calcareous mud, bioclastic sand, and gravel), with organic-rich varieties

The sediment recovered from Hole U1600A and most of Hole U1600B is summarized in one main volcanic- and tuffaceous-dominated unit (Unit I) divided into four subunits (1a–1d) (Figures F3, F4, F5; Table T2). Figure F7 displays images of the main lithologies of Unit I.

Unit I extends from 0 to 74.73 mbsf in Hole U1600A and from 0 to 74.46 mbsf in Hole U1600B (Table T2). Unit I consists predominantly of calcareous tuffaceous sediment (tuffaceous oozes, muds, and sands) and volcanic lithologies (ash, lapilli-ash, and lapilli) with lesser amounts of non-volcanic sediments (ooze, calcareous mud and sand, and bioclastic sand and their organic-rich counterparts). Grain sizes in Hole U1600B transition from mud in the uppermost 33 m to sand in the remainder of the hole, but this trend is not apparent in sediments from Hole U1600A.

Volcanic lithologies (>75% volcanic particles; glass shards, pumice, lithics, and crystals) comprise ash, lapilli-ash, and lapilli. Description of ash, lapilli-ash, and lapilli in volcanic intervals was based on the relative abundance of ash-sized (<2 mm) and lapilli-sized (2–64 mm) particles, as described in **Lithostratigraphy** in the Expedition 398 methods chapter (Kutterolf et al., 2024), with ash and lapilli used when the proportion of one size was >75% and lapilli-ash described when both sizes were present at <75% abundance (Fisher and Schmincke, 1984).

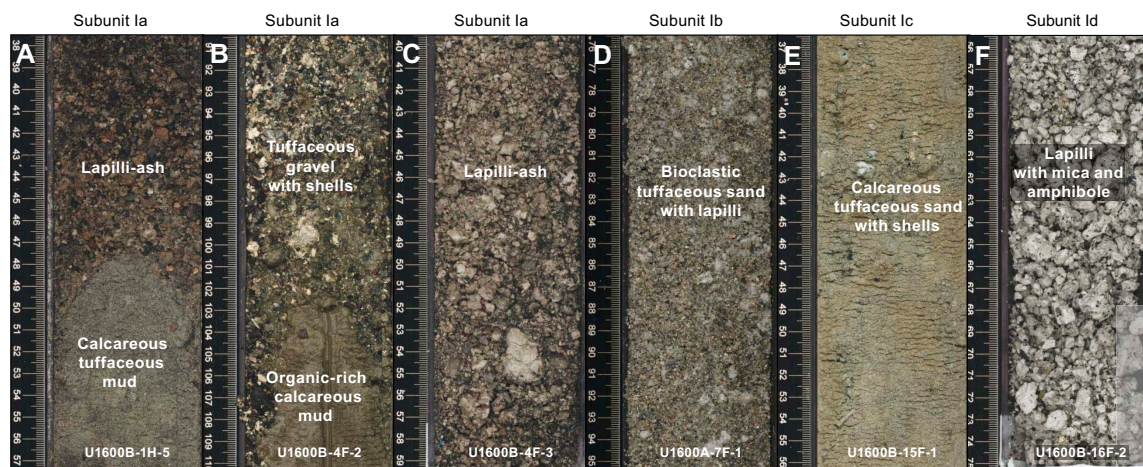


Figure F7. Common lithologies from Unit I, Site U1600. A. Lapilli-ash with underlying calcareous tuffaceous mud. B. Tuffaceous gravel with shells with underlying organic-rich calcareous mud. C. Lapilli-ash. D. Bioclastic tuffaceous sand with lapilli. E. Calcareous tuffaceous sand with shells. F. Mica and amphibole-rich lapilli.

Nonvolcanic sediment intervals are less abundant but occur throughout both holes. They include calcareous/bioclastic mud and sand and their organic-rich counterparts.

Shells and shell fragments are abundant in Unit I, even appearing in abundances >25% in ash and lapilli intervals. Naming protocols thus allowed for unusual lithology names such as bioclastic ash or bioclastic lapilli-ash as well as bioclastic tuffaceous sand with lapilli.

The subunits are described as follows:

- Subunit Ia (Sections 398-U1600A-1H-1, 0 cm, through 5F-CC, 8 cm [0–37 mbsf], and 398-U1600B-1H-1, 0 cm, through 8F-2, 117 cm [0–37.67 mbsf]): volcanic-dominated succession of ash, lapilli-ash, and lapilli with tuffaceous ooze/mud intercalated with a minor amount of calcareous mud.
- Subunit Ib (Sections 398-U1600A-6F-1, 0 cm, through 8F-1, 14 cm [37.2–46.74 mbsf], and 398-U1600B-10F-1, 0 cm, through 10F-CC, 18 cm [44.7–45.28 mbsf]): bioclastic ash and lapilli.
- Subunit Ic (Sections 398-U1600A-8F-1, 14 cm, through 10F-CC, 10 cm [46.74–60.61 mbsf], and 398-U1600B-11F-1, 0 cm, through 15F-3, 113 cm [49.1–72.03 mbsf]): calcareous tuffaceous sand with shells.
- Subunit Id (Sections 398-U1600A-13X-CC, 0 cm, through 13X-CC, 3 cm [74.7–74.73 mbsf], and 398-U1600B-15F-3, 113 cm, through 16F-CC, 9 cm [72.03–74.46 mbsf]): lapilli.

3.2.1.1. Subunit Ia

Subunit Ia extends from 0 to 37 mbsf in Hole U1600A and from 0 to 37.67 mbsf in Hole U1600B (Figure F3; Table T2). This subunit is distinguished by the predominance of volcanic lithologies (ash, lapilli-ash, and lapilli) intercalated with tuffaceous sediment (calcareous tuffaceous mud) and a small amount of nonvolcanic sediments (calcareous mud and organic-rich calcareous mud). Subunit Ia begins in both holes with ~60 cm of brown organic-rich tuffaceous mud and gray tuffaceous mud with an intercalated centimetric gray ash layer. This thin ash layer is well sorted and contains biotite-bearing, subrounded white pumice. It is followed by several distinct intervals of dark-colored ash with lapilli and lapilli-ash. The first interval consists of poorly sorted ash and lapilli-ash with multiple normal grading, subrounded white and gray pumice, and plagioclase and pyroxene crystals (intervals 398-U1600A-1H-1, 56 cm, through 1H-CC, 13 cm, and 398-U1600B-1H-1, 63 cm, through 1H-3, 97 cm). Another notable interval, only observed in Hole U1600B (1H-4, 5 cm, through 1H-5, 48 cm), consists of dark grayish brown ash and lapilli-ash with subrounded pink, white, and tan pumice clasts that contain plagioclase and pyroxene crystals. Shorter intervals (1–25 cm) of ash, lapilli-ash, and lapilli within calcareous tuffaceous mud are observed throughout Subunit Ia, often occurring in repetitive packages within the same core or core section (Figure F7A–F7C). Microscopic smear slide observations confirmed the presence of ash and volcanic crystals within tuffaceous lithologies and dominantly pumiceous and blocky glass in ash (Figure F8A, F8B). Calcareous tuffaceous mud and sand observed in Subunit Ia are commonly greenish gray to light grayish brown and often contain lithics, pumice/scoria, and shell fragments (Figure F7A, F7B). Observed biogenic clasts include shells, brachiopods, gastropods, bivalves, corals, and scaphopods. The bottom boundary of Subunit Ia is marked by the cessation of frequent intercalated dark volcanic lithologies and the appearance of bioclastic sands, ashes, and lapilli. Correlation between Subunit Ia in Holes U1600A and U1600B is generally good, although the volcanic lithologies observed in Hole U1600B tend to be finer grained and less extensive than their equivalent units in Hole U1600A, potentially due to variable recovery in each hole.

3.2.1.2. Subunit Ib

Subunit Ib extends from 37.2 to 46.74 mbsf in Hole U1600A and from 44.7 to 45.28 mbsf in Hole U1600B (Sections 398-U1600A-6F-1, 0 cm, through 8F-1, 14 cm, and 398-U1600B-10F-1, 0 cm, through 10F-CC, 18 cm) (Figure F3; Table T2). Subunit Ib consists of bioclastic volcanic lithologies, including ash, lapilli-ash, and lapilli (Figure F7D) intercalated with bioclastic tuffaceous sand with lapilli, bioclastic sand with lapilli, and gravel with lapilli. The boundary between the bottom of Subunit Ia and the top of Subunit Ib is characterized by a distinct color change from dark gray and olive to light gray and greenish gray. Bioclastic volcanic lithologies are typically light greenish gray and characterized by grain size ranging from fine ash to coarse lapilli. The interval is moderately to poorly sorted and contains subangular fibrous white pumice and shell fragments. Some

normal grading is observed. Smear slides taken in Subunit Ib show sedimentary lithics, foraminifera, biotite, and feldspar. Clay minerals, calcite, and nannofossils are also observed (Figure F8C, F8D).

3.2.1.3. Subunit Ic

Subunit Ic extends from 46.74 to 60.61 mbsf in Hole U1600A (Sections 8F-1, 14 cm, through 10F-CC, 10 cm) and from 49.1 to 72.03 mbsf in Hole U1600B (Sections 11F-1, 0 cm, through 15F-3, 113 cm) (Figure F3; Table T2). The maximum thickness of this subunit (22.93 m) is observed in Hole U1600B. The subunit is dominated by tuffaceous lithologies including bioclastic tuffaceous sand, calcareous tuffaceous sand with shells, gravel (granule), and organic-rich calcareous mud (Figure F7E). Intervals are mainly greenish to light greenish gray in color and well sorted with a maximum grain size of coarse sand. Clasts are lithics, biogenic, and crystals and consist of volcanic rocks, shells, biotite, pyroxene, and plagioclase. Some intervals contain organic material with moderate bioturbation and are more brownish in color. Shell fragments are scaphopoda (Sections 398-U1600A-9F-2 and 398-U1600B-14F-CC) and echinoderm spines (Sections 398-U1600A-8F-2, 85F-3, and 8F-CC). Subunit Ic ends with a calcareous tuffaceous sand with lapilli containing hornblende- and biotite-bearing pumice.

3.2.1.4. Subunit Id

Subunit Id extends from 74.7 to 74.73 mbsf in Hole U1600A (Sections 13X-CC, 0 cm, through 13X-CC, 3 cm [bottom of the hole]), and from 72.03 to 74.46 mbsf in Hole U1600B (Sections 15F-3, 113 cm, through 16F-CC, 9 cm) (Figure F3; Table T2). The total thickness of this subunit is 2.43 m in Hole U1600B; only 3 cm of this subunit was recovered in Hole U1600A. Subunit Id consists entirely of light gray to white lapilli (Figure F7F). The subunit is characterized by moderately to well-sorted pumice clasts that are rich in biotite and hornblende. This lapilli layer has a grain size range of fine lapilli to coarse lapilli (2–64 mm). The whole of Core 398-U1600B-16F shows normal grading from the bottom (core catcher) to the top (Section 16F-1). Each section in that core is also itself normally graded.

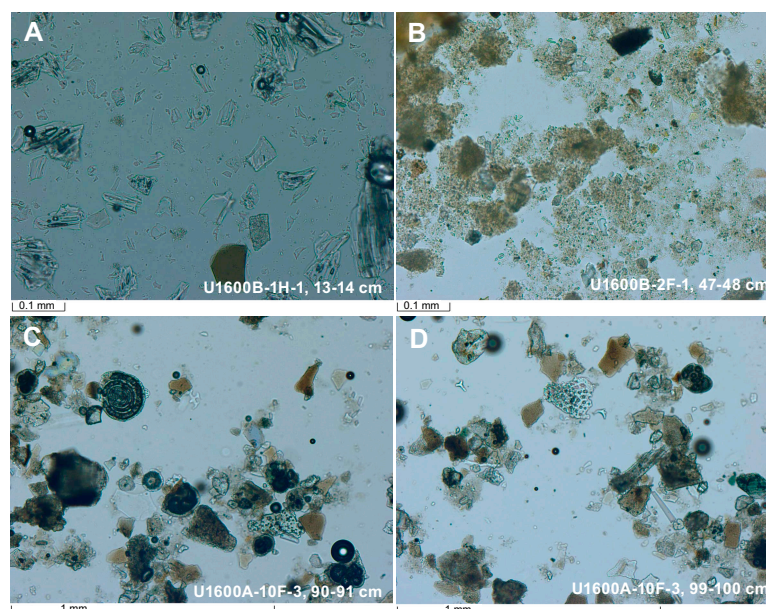


Figure F8. Common lithologies, Holes U1600A and U1600B. A. Ash (Subunit Ia); note the biotite, bottom center. B. Calcareous tuffaceous mud with organic material (Subunit Ia). C. Calcareous tuffaceous sand with shells (Subunit Ic). D. Calcareous tuffaceous sand with shells (Subunit Ic).

3.2.2. Unit II

Intervals: 398-U1600B-17F-1, 0 cm, to 19F-CC, 16 cm (bottom of the hole); and 398-U1600C-3R-CC, 0 cm, to 12R-1, 10 cm

Depths: Hole U1600B = 77.3–91.63 mbsf; Hole U1600C = 84.70–172.1 mbsf

Thickness: 92.1 m (calculating from the top of Unit II in Hole U1600B to the bottom of Unit II in Hole U1600C)

Age: Early Pleistocene to Pliocene

Lithology: primarily organic-rich calcareous mud/sand with shells with lesser calcareous mud/sand, sandstone, and rare ash pods

Unit II begins with the first calcareous sand with shells beneath a lapilli interval in Hole U1600B. It continues in Hole U1600C and is dominated by interbedded organic-rich calcareous mud and sand with shells (Figures F3, F9A; Table T2). The unit starts at 77.3 mbsf in Hole U1600B with cyclical variations between calcareous sands with shells and organic-rich calcareous muds through 91.63 m (bottom of the hole). The unit continues with these cyclic variations in Hole U1600C from 84.7 to 105.39 mbsf, at which point all sand and mud intervals become organic rich until 166.01 mbsf, where there is an unrecovered transition to sandstone.

A brachiopod was recovered in the upper part of this unit (interval 398-U1600B-19F-3, 21–25 cm) (Figure F10).

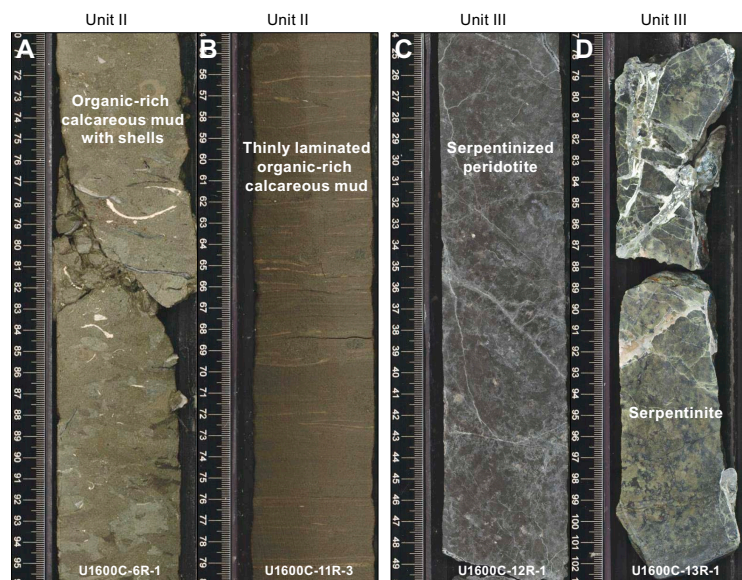


Figure F9. Representative lithologies from Units II and III, Hole U1600C. A. Organic-rich calcareous mud with shells. B. Thinly laminated organic-rich calcareous mud. C, D. Serpentinized peridotite.



Figure F10. Close-up image of a brachiopod from Subunit Ie, Site U1600.

Organic-rich calcareous muds are olive-gray to dark olive-gray with or without thin (<0.3 cm) laminations (Figure F9A, F9B). Some intervals of mud are normally graded from fine sand at the base to mud. They are typically highly bioturbated with *Zoophycos* and *Chondrites* traces and contain burrows infilled with ash. Grain size ranges from mud to medium sand. Calcareous sand and organic-rich calcareous sand are well sorted, with colors ranging from greenish gray in the nonorganic-rich intervals to olive-gray in the organic-rich intervals.

The sands and muds are interbedded on scales of 25–50 cm, but with depth, mud intervals become thicker such that by 150.15 mbsf organic-rich calcareous mud is the only lithology until the bottom of the unit.

3.2.3. Unit III

Interval: 398-U1600C-12R-1, 10 cm, to 14R-1, 54 cm (bottom of the hole)

Depth: 172.1–184.24 mbsf

Thickness: >12.14 m

Age: Uncertain

Lithology: ultramafic rocks

The rocks recovered from the base of Hole U1600C are summarized in one unit (Unit III) (Figures F3, F9C, F9D; Table T2). Unit III extends from 172.1 to 184.24 mbsf in Hole U1600C and consists of serpentinized ultramafic rocks (peridotite) (Table T2). Figure F9C and F9D displays core images of the main lithology found in Unit III.

Macroscopically, two distinct types of peridotite were observed. Dark gray serpentinized peridotite in Section 398-U1600C-12R-1 (Figure F9C) grades into grayish green, heavily serpentinized peridotite (serpentinite) in Sections 13R-1 and 14R-1 (Figure F9D). The dark gray variety is characterized by olivine replaced by serpentine, coarse orthopyroxene, and white crosscutting talc veins. The grayish green variety is partly brecciated, mesh-textured, completely serpentinized, and also contains white crosscutting talc veins. These rocks represent in situ basement rocks that underlie the sediments of Unit II.

Thin sections of two samples from Unit III are displayed in Figure F11. The dark gray peridotite from Section 398-U1600C-12R-1 (Figure F11A–F11C) shows orthopyroxene porphyroblasts often partly or almost completely altered to opaques (magnetite), chlorite, and replacement of

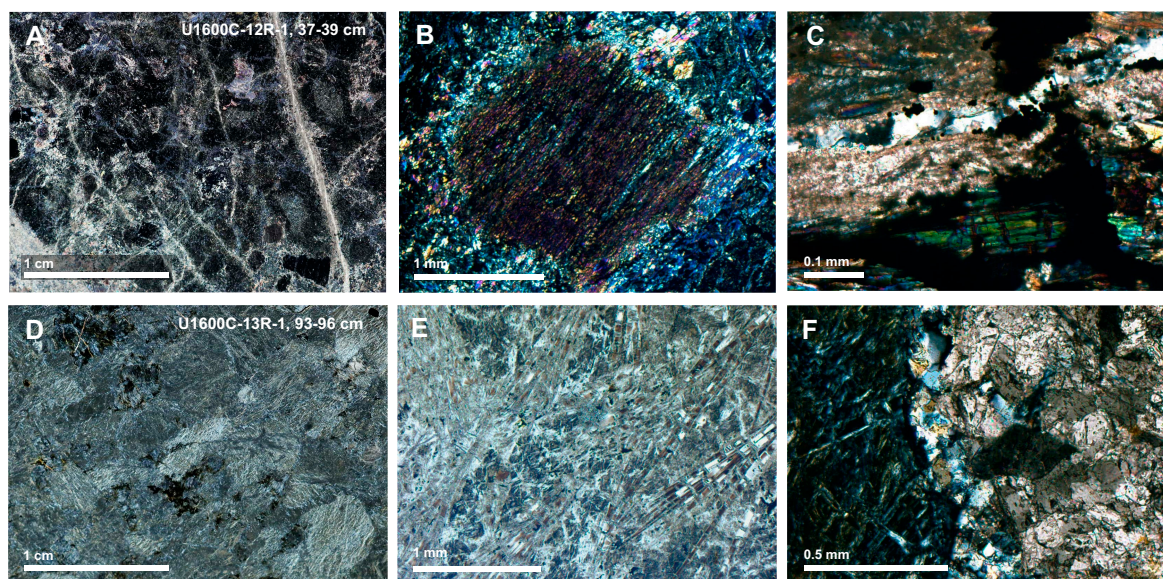


Figure F11. Ultramafic rocks from Unit III, Hole U1600C (cross-polarized light [XPL]). A–C. Dark gray serpentinized peridotite: (A) entire sample; (B) orthopyroxene porphyroblast; and (C) close-up of quartz vein with calcite, serpentine minerals, talc, and opaques. D–F. Grayish green heavily serpentinized peridotite (serpentinite): (D) entire sample, (E) heavily serpentinized peridotite with olivine completely replaced by serpentine minerals, and (F) serpentinite and carbonate parts of the sample with a quartz vein at the contact.

olivine by serpentine, with only small relics of olivine preserved. Crosscutting veins are characterized by serpentine minerals, calcite, talc, and quartz. The grayish green serpentinite from Sections 13R-1 and 14R-1 (Figure F11D–F11F) is characterized by serpentine minerals, chlorite, spinel, and opaques (magnetite). Olivine is completely replaced by serpentine, producing a mesh texture of pseudomorphic serpentine after olivine. Brecciation is indicated by the presence of a coarse-grained part of the thin section with interlocking carbonate crystals. Crosscutting veins consist of serpentine minerals, talc, and quartz, among others.

3.3. Correlations between holes

Holes U1600A and U1600B were drilled to overlap so that recovered stratigraphy could be correlated. Hole U1600A extends to 74.73 mbsf, and Hole U1600B extends to 91.65 mbsf. The overlap in the uppermost 75 m allows correlation of sediment-specific features. In both holes, recovery was poor between 40 and 46 mbsf. Although the base of Hole U1600B overlaps ~7 m with the top of Hole U1600C, the limited recovery within these core sections does not allow correlation between these two holes.

3.3.1. Notable correlated deposits

- Mud with lapilli overlying a light gray ash layer in the uppermost 17 cm of Holes U1600A and U1600B (Figure F12A).
- Lithic lapilli overlying calcareous tuffaceous mud in Subunit Ia (9.27 mbsf in Hole U1600A and 8.79 mbsf in Hole U1600B) (Figure F12B).

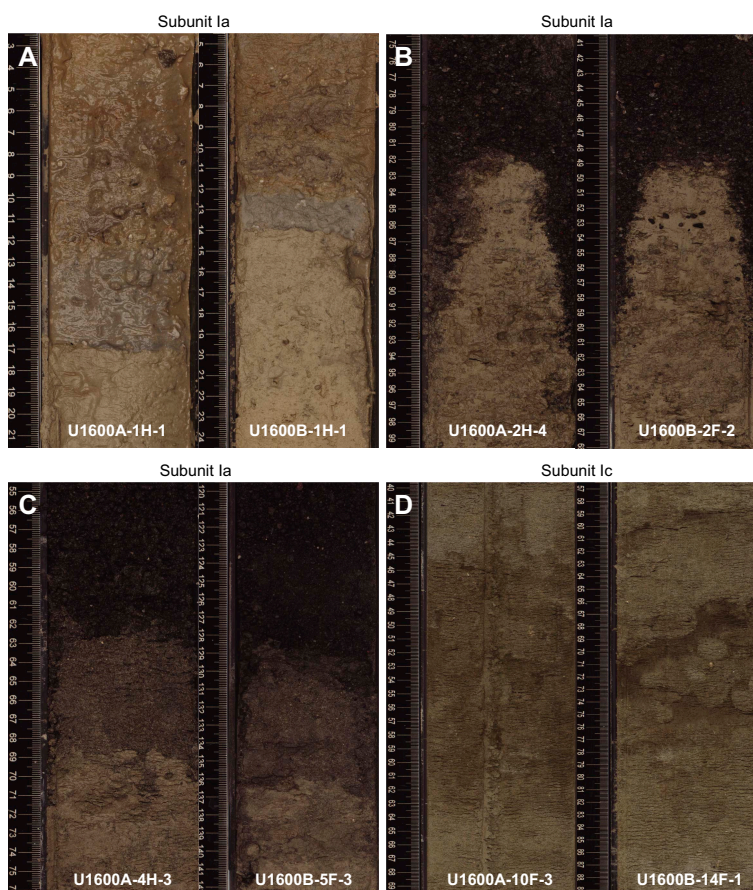


Figure F12. Correlations between Holes U1600A and U1600B. A. Gray ash layer near the top of both holes (~0.17 mbsf). B. Lithic lapilli overlying calcareous tuffaceous mud (Hole U1600A: 9.27 mbsf; Hole U1600B: 8.79 mbsf). C. Double ash layer (Hole U1600A: 26.11 mbsf; Hole U1600B: 25.19 mbsf). D. Rarely observed organic-rich tuffaceous ooze (Hole U1600A: 59.47 mbsf; Hole U1600B: 63.86 mbsf).

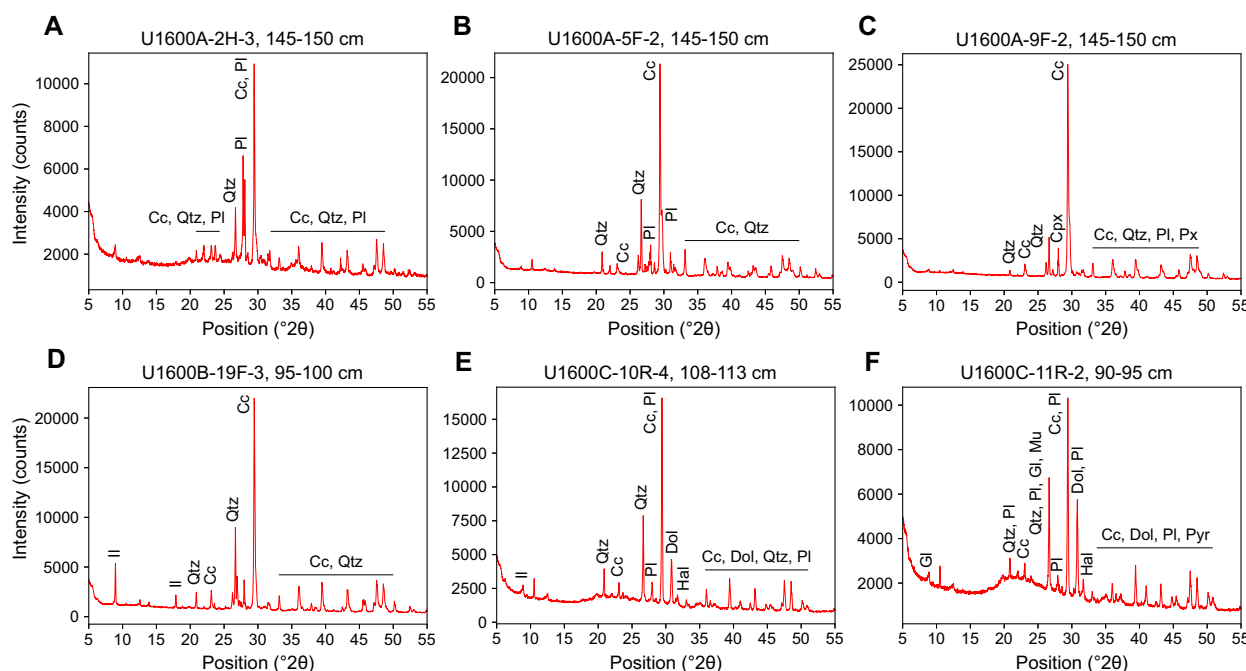


Figure F13. XRD spectra of representative lithologies, Site U1600. A. Organic-rich tuffaceous ooze (Subunit Ia). B. Bioclastic tuffaceous sand with organic material (Subunit Ia). C. Calcareous tuffaceous sand with shells (Subunit Ic). D. Calcareous sand with shells (Subunit Ie). E. Organic-rich dolomitic mud (Unit II). F. Calcareous sand with shells (Unit II). Cc = calcium carbonate (calcite, aragonite), Qtz = quartz, Pl = Ca-rich or Na-rich plagioclase, Cpx = clinopyroxene, Px = pyroxene, Il = illite, Dol = dolomite, Hal = halite, Gl = glauconite, Mu = muscovite, Pyr = pyrite.

- Double ash layer at 26.11 mbsf in Hole U1600A and 25.19 mbsf in Hole U1600B showing a finer grained, gray ash layer at the bottom and a coarser grained, black ash layer at the top (Figure F12C).
- Rarely observed organic-rich interval in tuffaceous ooze (59.47 mbsf in Hole U1600A; 63.86 mbsf in Hole U1600B) (Figure F12D).

3.4. X-ray diffraction

XRD data were collected from three IW squeeze cake sediment residues from Hole U1600A and five squeeze cake sediment residues each from Holes U1600B and U1600C. The analyzed samples comprise lithologies ranging from organic-rich tuffaceous ooze to bioclastic tuffaceous sand with organic material, calcareous tuffaceous sand with shells, and calcareous mud and sand with shells. Muds and sands in Hole U1600C are calcareous, with some also containing significant amounts of dolomite. Spectra of representative lithologies from Subunits Ia, Ic, and Ie and Unit II are shown in Figure F13. Major mineral phases identified in these lithologies include calcium carbonate (calcite and/or aragonite), quartz, and clay minerals of the illite group. In addition, Ca-rich or Na-rich plagioclase and clinopyroxene (diopside) are identified in tuffaceous lithologies, which may also show the characteristic amorphous hump at low 2θ , if larger amounts of volcanic glass are present. Samples of mud or mud with shells analyzed in Hole U1600C are often dolomitic, with abundances ranging 6.6%–43.2%, as identified using the Rietveld method.

4. Stratigraphic correlation

Three holes were drilled at Site U1600. Hole U1600A was cored to 74.76 mbsf (CSF-A scale) using the APC (0–31.74 mbsf), HLAPC (32.5–60.66 mbsf), and XCB (60.7–74.76 mbsf) systems. Hole U1600B was cored to 91.65 mbsf using the APC (0–6.83 mbsf) and HLAPC (6.8–91.65 mbsf) systems. Hole U1600C was drilled using the RCB system to 184.24 mbsf; the uppermost interval from 0 to 75 mbsf was drilled without coring.

No correlation during coring was required due to the prevailing challenging drilling condition recognized in Hole U1600A.

4.1. Correlation for establishing CCSF-A depth scale and construction of the splice

To establish the core composite depth below seafloor, Method A (CCSF-A), depth scale, we analyzed cores from Holes U1600A and U1600B for their physical properties using the Whole-Round Multisensor Logger (WRMSL) for magnetic susceptibility (MS) and gamma ray attenuation (GRA) density; the Natural Gamma Radiation Logger (NGRL) for NGR intensity; and photos after the cores were split into working and archive halves (see [Physical properties](#); also see [Physical properties](#) in the Expedition 398 methods chapter [Kutterolf et al., 2024]). The MS data proved to be the most reliable physical parameter for correlations. Because Hole U1600C had no overlap with Holes U1600A and U1600B, no correlation was possible.

Both Holes U1600A and U1600B preserved the mudline; we used Core 398-U1600A-1H as the anchor for stratigraphic correlation. Using this anchor core, we attempted to determine the relative depth offset of each core by establishing affine ties between the holes based on the maximum correlation of all measured physical properties. These depth offsets are reported in Table T3. Figure F14 shows the MS of the three holes on the resulting CCSF-A scale, in which individual cores are shifted according to the identified correlations. This figure highlights the vertical alignment of characteristic events in the adjacent holes.

In general, we were able to identify reliable correlations between most cores of Holes U1600A and U1600B. Interruptions in the correlation chain occurred in ashy and tuffaceous segments, where no reliable correlations were possible. For each of these interruptions, we used the relative offset between untied (uncorrelated) cores derived from the CSF-A scale in order to keep the composite depth scale as close to the original CSF-A scale as possible.

After we established the composite depth scale, we spliced selected sequences from Holes U1600A and U1600B and appended Hole U1600C to create the most complete and representative section possible. The end product of this process is shown in Table T4 and Figure F15. To ~70 mbsf (CCSF, Method D [CCSF-D], scale), the splice contains only small gaps less than 2 m. Farther below, larger gaps occur, which are the result of the generally lower recovery.

Site U1600 cores showed no significant expansion. Figure F16 plots the CSF-A and CCSF-A scales for Holes U1600A–U1600C and shows a linear regression curve calculated for all holes together. The deviation from the dotted line indicates an affine growth factor of approximately 2%.

Table T3. Affine ties, Site U1600. [Download table in CSV format.](#)

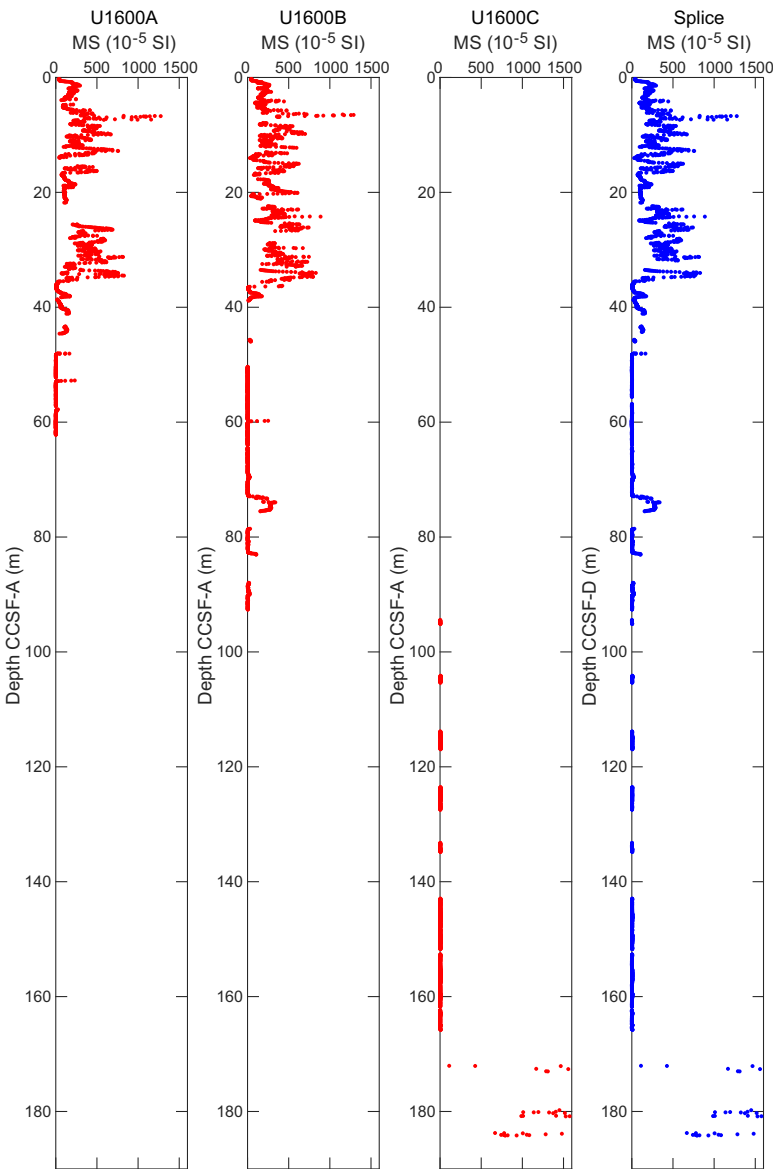


Figure F14. WRMSL-derived MS data, Holes U1600A–U1600C. MS data are on the CCSF-A depth scale, and the splice (right) is on the CCSF-D scale.

Table T4. Splice ties, Site U1600. [Download table in CSV format.](#)

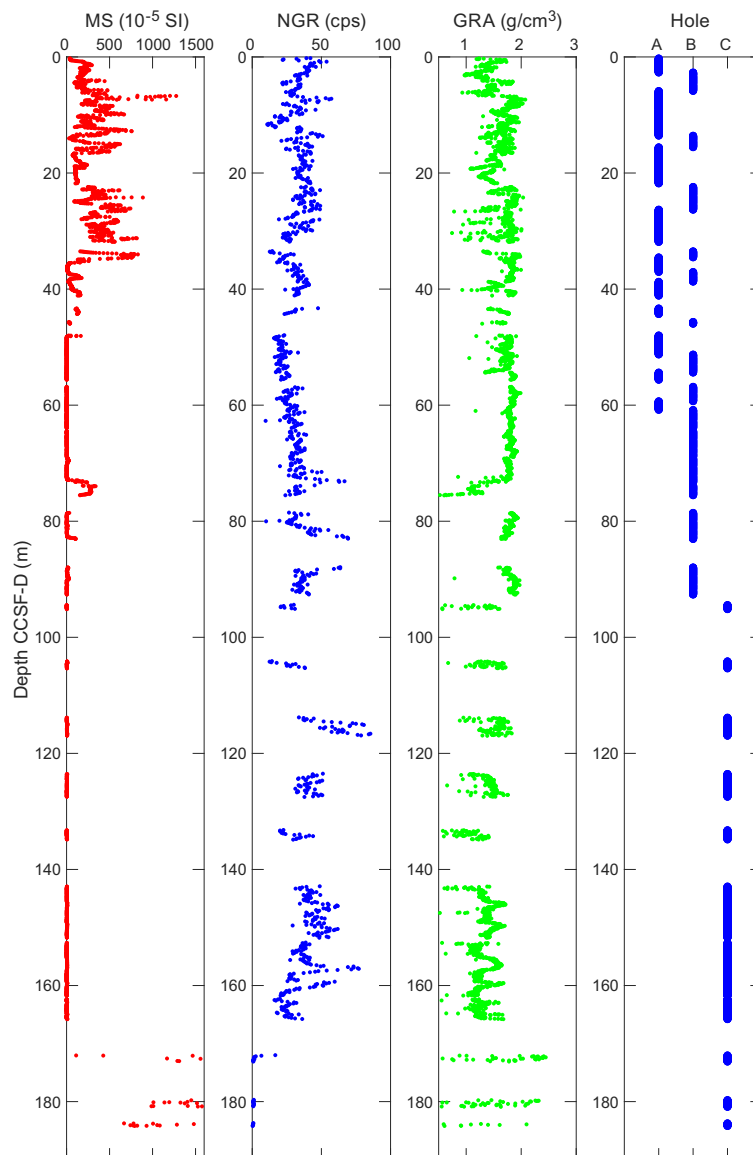


Figure F15. Splice, Site U1600. cps = counts per second.

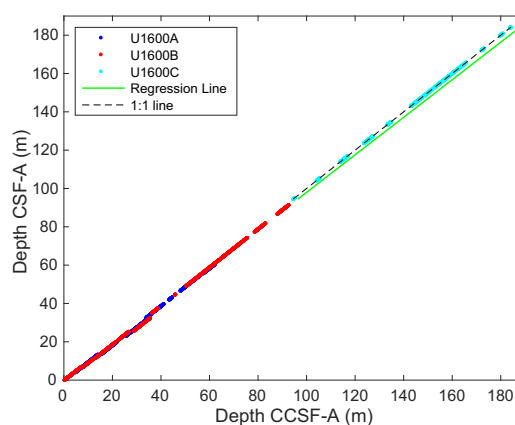


Figure F16. CCSF-A versus CSF-A core top depths, Holes U1600A–U1600C. Lines fit through the core top depths of all holes give an estimate of the core expansion. At Site U1600, this is estimated to be 2%.

5. Structural geology

Structural geology analyses at Site U1600 included description of cores retrieved from Holes U1600A–U1600C. A total of 70 structures were measured, and most of those measurements derived from relatively consolidated intervals. Observed and measured structures in cores are bedding, which includes lamination, faults, deformation bands, breccias, and mineral veins. The distributions and dip angles of planar structures (measured to an estimated accuracy of $\pm 2^\circ$) are shown in Figure F17.

Deformation related to drilling and core recovery was noted but not recorded. Here, we describe and provide examples of the features that were recorded.

5.1. Bedding

Bedding planes ($n = 48$), measured mainly on thin sand beds/laminae, organic-rich mud, and ooze boundaries, mostly exhibit horizontal to subhorizontal dips ranging from 1° to 15° (mean dip = 4°) (Figure F17). However, steeper inclinations develop at 72–82 mbsf within Lithostratigraphic Subunit Id where dips range 4° – 15° (mean = 11°) (Figure F18).

To evaluate the bedding dip variations in the ~190 m deep hole, we divided bedding measurements to correspond to the lithostratigraphic units/subunits defined by core description (see [Lithostratigraphy](#)) and used a box plot to indicate the distribution of bedding dips (Figure F18). Measurements were mostly concentrated in Subunits Ia and Id and Unit II (Figure F17). Thus, a total of three bedding measurement units were described. In the box plot, each box indicates the data distribution range between the twenty-fifth and seventy-fifth percentiles, which is statistically meaningful. Median values (fiftieth percentile of all respective measurements in a unit), which are representative of the data distribution, were used for the evaluation of dip angles. Major increases in bedding dips were identified from Unit II through Subunit Id, where median dips increase from 4° to 11° . In contrast, in Subunits Id–Ia, median dip decreases from 11° to 2° (Figure F18).

5.2. Mineral veins

Mineral veins are abundant in the interval below 172 mbsf in Hole U1600C (Figure F17), which corresponds to Lithostratigraphic Unit III (see [Lithostratigraphy](#)), which is composed of ser-

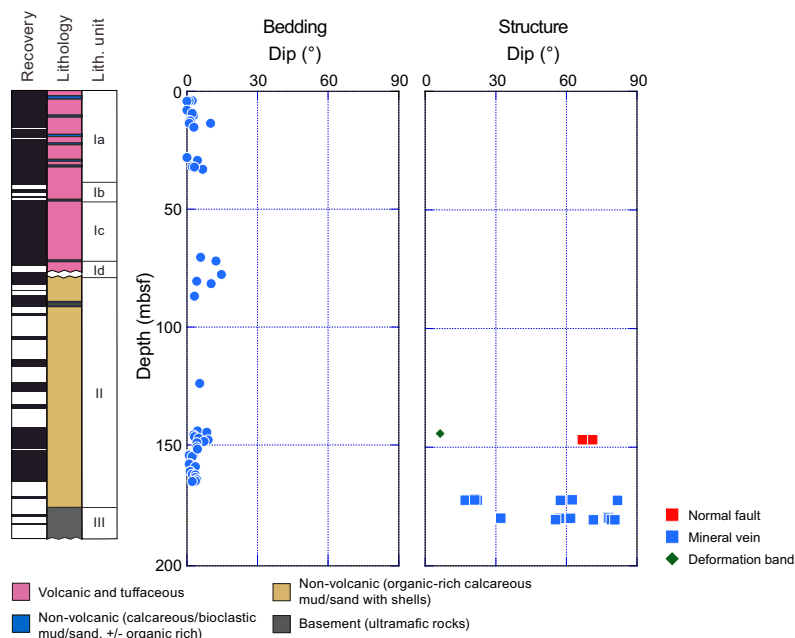


Figure F17. Dip data, Site U1600. Lithostratigraphic subunits are described in [Lithostratigraphy](#).

pentinized peridotite. Crystal growth of quartz was identified in the gap along fractures. Sometimes it represented brecciated textures composed of angular-shaped host rock clasts (Figure F19). A thin talc film developed along the contact between the host rock and breccia and along brecciated clasts. Also, in situ brecciation of talc was frequently identified along the thick mineral veins. These textures apparently indicate that the talc was formed from the reaction between serpentine and SiO_2 -saturated fluid during brecciation.

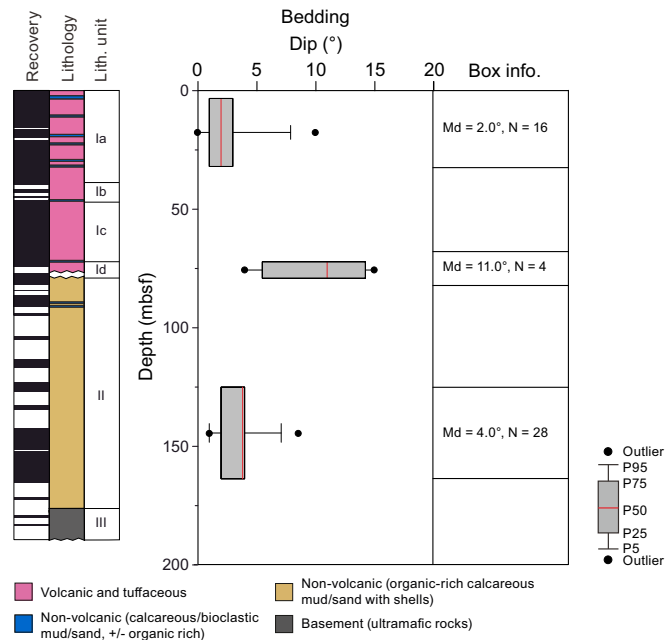


Figure F18. Box plots of bedding dip distribution, Site U1600. The minimum (P5), first quartile (P25), and median value (P50), third quartile (P75), and maximum (P95) are shown. Only the first outlier smaller than P5 and larger than P95 is plotted as a dot. Md = median dip. N = number of samples. P = percentile. Lithostratigraphic subunits are described in Lithostratigraphy.

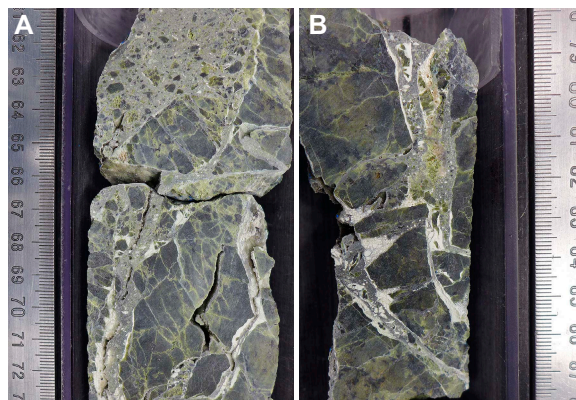


Figure F19. A, B. Mineral veins and breccia (A: 398-U1600C-13R-1, 61–73 cm; B: 13R-1, 78–88 cm). Quartz-crystal growth was identified in the hollow. Talc coatings along the brecciated surfaces and in situ brecciation of these coatings were identified.

6. Biostratigraphy

Planktonic and benthic foraminifera and calcareous nannofossils were examined from core catcher samples and additional split core samples from Holes U1600A–U1600C to develop a shipboard biostratigraphic framework for Site U1600. Additionally, planktonic and benthic foraminifera provided data on paleowater depths, downslope reworking, and possible dissolution.

Site U1600 drilled the sedimentary sequence atop the tectonic horst situated between the Anafi and Anhydros Basins and recovered a 184.24 m thick Holocene to Upper Pliocene sequence comprising volcanogenic sediments, calcareous and tuffaceous sands, and ultramafic basement. Calcareous nannofossils and planktonic foraminifera provided good resolution in the Holocene through Upper Pliocene sediments. Ages provided by benthic foraminifera were also consistent with those of planktonic foraminifera and calcareous nannofossils. Biostratigraphic datums recognized at Site U1600 are given in Tables T5 and T6, and an age-depth plot is shown in Figure F20.

Hole U1600A recovered Holocene to Early Pleistocene-aged sediments (0.00–0.91 Ma) from Samples 1H-CC, 30–31 cm, to 6F-CC, 0–3 cm (4.08–39.79 mbsf). Hole U1600B recovered Holocene to Early Pleistocene-aged sediments (0.00–1.99 Ma) from Samples 1H-CC, 8–11 cm, to 19F-CC, 16–18 cm (6.83–91.65 mbsf). Hole U1600C recovered Early Pleistocene to Late Pliocene-aged sediments (1.95–3.19 Ma) from Samples 3R-CC, 14–17 cm, to 11R-CC, 0–5 cm (84.87–165.89 mbsf).

Table T5. Calcareous nannofossil events and absolute ages, Site U1600. [Download table in CSV format.](#)

Calcareous nannofossil event	Age (Ma)	Top core, section, interval (cm)	Bottom core, section, interval (cm)	Top depth CSF-A (m)	Mid-depth CSF-A (m)	Bottom depth CSF-A (m)
Base <i>Emiliania huxleyi</i>	0.265	398-U1600A-2H-CC, 97–100	398-U1600A-3H-CC, 11–14	12.46	16.27	20.08
Base <i>Emiliania huxleyi</i>	0.265	398-U1600B-3H-CC, 0–4	398-U1600B-5H-CC, 23–27	15.60	20.585	25.57
Base <i>Reticulofenestra asanoi</i>	1.078	11F-CC, 13–15	12F-CC, 13–15	53.94	56.22	58.50
Base (>5.5 µm) <i>Gephyrocapsa</i> spp. large	1.617	13F-CC, 12–14	14F-CC, 11–13	62.94	65.41	67.88
Base (>4 µm) <i>Gephyrocapsa</i> spp. medium	1.73	17F-CC, 16–18	18F-CC, 14–16	82.06	82.1	82.14
Top <i>Discoaster brouweri</i>	1.95	398-U1600C-4R-CC, 14–17	398-U1600C-5R-CC, 11–14	95.33	100.415	105.50
Acme base <i>Discoaster triradiatus</i>	2.216	6R-CC, 13–18	7R-CC, 13–18	117.19	122.42	127.65
Top <i>Discoaster tamalis</i>	2.80	10R-CC, 13–15	11R-CC, 0–5	161.97	163.905	165.84

Table T6. Foraminiferal events and absolute ages, Site U1600. [Download table in CSV format.](#)

Foraminifer event	Age (Ma)	Top hole, core, section, interval (cm)	Bottom hole, core, section, interval (cm)	Top depth CSF-A (m)	Mid-depth CSF-A (m)	Bottom depth CSF-A (m)
		398-	398-			
Base <i>Globigerinoides ruber</i> pink	0.33	U1600A, 3H-CC, 11–14	U1600A, 4H-CC, 83–85	20.08	25.91	31.74
Paracme top <i>Neogloboquadrina</i> spp. (sin)	0.51	U1600B, 8F-CC, 0–2	U1600A, 6H-CC, 0–3	37.69	38.74	39.79
Paracme base <i>Neogloboquadrina</i> spp. (sin)	0.91	U1600A, 7F-CC, 0–3	U1600A, 8F-CC, 8–11	43.28	47.155	51.03
Paracme top <i>Neogloboquadrina</i> spp. (sin)	1.21	U1600B, 11F-CC, 13–15	U1600B, 12F-CC, 13–15	53.94	56.22	58.5
Base common <i>Hyalinea balthica</i>	1.492	U1600B, 13F-CC, 12–14	U1600B, 14F-CC, 11–13	62.94	65.41	67.88
Base common <i>Neogloboquadrina</i> spp. (sin)	1.79	U1600B, 17F-CC, 16–18	U1600B, 18F-CC, 14–16	82.06	82.1	82.14
Base common <i>Globoconella inflata</i>	1.99	U1600C, 3R-CC, 14–17	U1600C, 4R-CC, 14–17	84.87	90.045	95.22
Influx <i>Sphaeroidinella dehiscens</i>	1.995	U1600B, 19F-CC, 16–18	U1600B, 19F-CC, 16–18	91.65	91.65	91.65
Base <i>Globoconella inflata</i>	2.09	U1600C, 4R-CC, 14–17	U1600C, 5R-CC, 11–14	95.22	100.375	105.53
Top common <i>Globorotalia bononiensis</i>	2.41	U1600C, 6R-CC, 13–18	U1600B, 7R-CC, 13–18	117.19	122.42	127.65
Base <i>Neogloboquadrina atlantica</i> (sin)	2.72	U1600C, 9R-CC, 11–14	U1600C, 11R-CC, 0–5	151.93	158.885	165.84
Top <i>Dentoglobigerina altispira</i>	3.17	U1600C, 9R-CC, 11–14	U1600C, 11R-CC, 0–5	151.93	158.885	165.84
Top <i>Sphaeroidinellopsis</i> s.l.	3.19	U1600C, 9R-CC, 11–14	U1600C, 11R-CC, 0–5	151.93	158.885	165.84

6.1. Calcareous nannofossils

Calcareous nannofossil biostratigraphy in Holes U1600A–U1600C was established through analysis of core catcher samples and additional split core samples. Nannofossils are rare to abundant in samples from the Pleistocene to Upper Pliocene sequence (Hole U1600A: 4.08–74.76 mbsf; Hole U1600B: 6.83–91.65 mbsf; Hole U1600C: 84.855–165.865 mbsf). Preservation is generally good to moderate with sporadically poor intervals throughout the sequence (Figure F20); however, there is significant reworking of older material in most of the Pleistocene samples. The assemblages are characterized by the occurrence of warm-water species such as *Rhabdosphaera clavigera* and *Umbilicosphaera sibogae* throughout the sequence in Holes U1600A and U1600B.

Seven nannofossil biostratigraphic datums are recognized at Site U1600, representing a continuous Late Pliocene to Late Pleistocene sedimentary sequence (Figure F21). The distribution of calcareous nannofossil taxa is shown in Table T7, and biostratigraphic datums are given in Table T5.

The presence of *Emiliana huxleyi* in Samples 398-U1600A-1H-CC, 30–31 cm, to 2H-6, 97–100 cm (4.08–12.46 mbsf), and 398-U1600B-1H-CC, 8–11 cm, to 3F-CC, 0–4 cm (6.83–15.6 mbsf), indicates a Holocene to Middle Pleistocene age (≤ 0.265 Ma) within Zones MNN21a and 21b

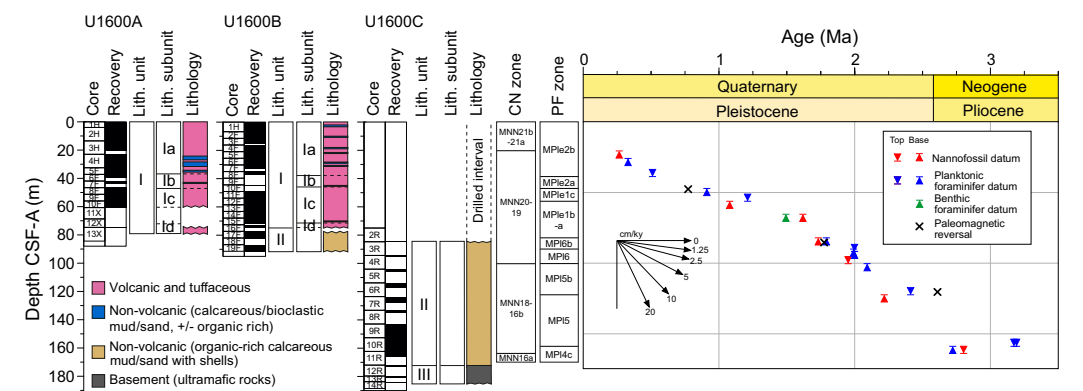


Figure F20. Age-depth plot, Hole U1600A. Integrated biochronology and magnetostratigraphy are shown. CN = calcareous nannofossil. PF = planktonic foraminifer. Biohorizons correspond with those in Tables T5 and T6. Biohorizons denoted with a question mark are not fully constrained and are interpolated.

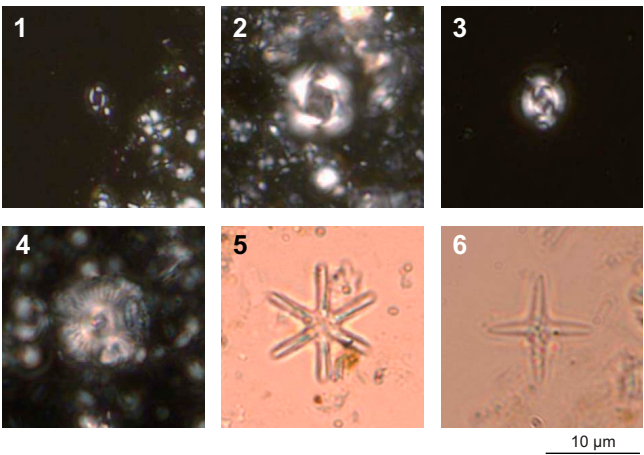


Figure F21. Calcareous nannofossils. 1. *Emiliana huxleyi* (Lohmann) Hay and Mohler (398-U1600A-2H-CC, 30–31 cm). 2. *Reticulofenestra asanoi* Sato and Takayama (398-U1600B-11F-CC, 13–15 cm). 3. *Gephyrocapsa* spp. large form (>5.5 μm) (13F-CC, 12–14 cm). 4. *Calcidiscus macintyre* (Bukry and Bramlette) Loeblich and Tappan (19F-CC, 16–18 cm). 5. *Discoaster brouweri* Tan Sin Hok (398-U1600C-6R-CC, 13–18 cm). 6. *Discoaster tamalis* Kamptner (11R-CC, 0–5 cm).

Table T7. Distribution of calcareous nannofossils, Site U1600. [Download table in CSV format.](#)

(CNPL11: Backman et al., 2012; NN21: Martini, 1971; CN15: Okada and Bukry, 1980) of Rio et al. (1990) and Di Stefano and Sturiale (2010). The basal occurrence of *Reticulofenestra asanoi* (1.078 Ma), which lies just below the Jaramillo Subchron of the Matuyama Chron in Zone MNN19e, is found between Samples 11F-CC, 13–15 cm, and 12F-CC, 13–15 cm (53.94–58.5 mbsf). The large form of *Gephyrocapsa* spp. (>5.5 μm) that appears between 1.245 and 1.617 Ma (Zones MNN19d, CNPL8, NN19, and CN13b) occurs in Samples 12F-CC, 13–15 cm, and 13F-CC, 12–14 cm (58.5–62.94 mbsf). The basal occurrence of medium-form *Gephyrocapsa* spp. (>4.0 μm), which is situated just above Olduvai Subchron of the Matuyama Chron within Zone MNN19b (CNPL8, NN19, and CN13a), is recognized between Samples 17F-CC, 16–18 cm, and 18F-CC, 14–16 cm. The assemblages in Samples 18R-CC, 14–16 cm, and 19F-CC, 16–18 cm (82.14–91.64 mbsf), are characterized by occurrences of *Calcidiscus macintyrei*, *Gephyrocapsa* spp. (small form), *Pseudoemiliania lacunosa*, and *Reticulofenestra* spp. without Discoasteraceae. Thus, Samples 18F-CC, 14–16 cm, and 19F-CC, 16–18 cm, are tentatively placed in Zone MNN19a (CNPL7, NN19, and CN13a; 1.73–1.95 Ma). The Gelasian (Early Pleistocene) assemblages found at this site are characterized by the occurrence of discoasters, which include marker species of Miocene to Pliocene successions. The last occurrence of *Discoaster brouweri*, which indicates the boundary of Zones MNN18/19a (NN18/19 and CN12d/CN13a), occurs between Samples 398-U1600C-4R-CC, 8–12 cm, and 5R-CC, 11–14 cm (100.4 mbsf). The acme base of *Discoaster triradiatus*, which characterizes the middle MNN18 Zone (2.216 Ma), appears between Samples 6R-CC, 13–18 cm, and 7R-CC, 13–18 cm (122.4 mbsf). The *Discoaster tamalis* disappearance at 2.8 Ma occurs in the upper part of the Gauss Chron, and this specimen last occurs between Samples 10R-CC, 13–15 cm (161.96 mbsf), and 11R-CC, 0–5 cm (165.865 mbsf).

6.2. Foraminifera

Planktonic and benthic foraminifera were examined from core catcher samples and split core samples from Holes U1600A–U1600C (Tables T8, T9). Absolute ages assigned to biostratigraphic datums follow those listed in Table T6 in the Expedition 398 methods chapter (Kutterolf et al., 2024). Planktonic and benthic foraminifer datums for Site U1600 are given in Table T6. Planktonic foraminifer abundances and indications of oceanicity (e.g., Hayward et al., 1999) and benthic foraminifer paleowater depth estimations are given in Figures F22 and F23.

Because of the volcanogenic and siliciclastic nature of sections of the cored sedimentary sequence, residues (>125 μm) from washed samples were often commonly composed of volcanoclastic particles such as pumice, scoria, and ash and siliciclastic grains that often diluted the microfossil component of residues. Foraminifera often dominate the biogenic component of residues where age markers are present in sufficient numbers to date most samples reliably; however, a significant component of other fossil material are present, including shells and fragments (Bivalvia and Gastropoda), Pteropoda, Scaphopoda, Bryozoa, Cnidaria (coral), echinoid spines and plate fragments, radiolarians, and ostracods, as well as fish teeth. In addition, reworked Lower Pleistocene and Pliocene foraminifera are often present.

Throughout the section, foraminifera with very good to poor preservation are present in sediments. Foraminifer abundances are also variable and are notably rare in tuffaceous oozes, likely due to sedimentary dilution, whereas coarser bioclastic intervals are sometimes barren of planktonic foraminifera.

6.2.1. Holocene to Pleistocene foraminiferal biostratigraphy

Because of explosive volcanic events and rapid deposition of the upper sedimentary section, the base of the Holocene is not possible to assign accurately. Planktonic foraminifer assemblages from the Holocene–Pleistocene section of Holes U1600A–U1600C are mostly well preserved, where specimens are rarely broken or exhibiting partially dissolved shell walls.

Table T8. Distribution of planktonic foraminifera, Site U1600. [Download table in CSV format.](#)

Table T9. Distribution of benthic foraminifera, Site U1600. [Download table in CSV format.](#)

Hole, core, section, interval (cm)	Bottom depth CSF-A (m)	AW Planktonic (%)	OK Planktonic (%)	Planktonic mean (%)	Oceanicity	Paleowater depth (m)
398-						
U1600A-1H-CC	4.08	78	64	71	Sub-oceanic	
U1600B-1H-CC	6.83	50		50	Extra-neritic	
U1600B-2F-CC	11.46	94		94	Oceanic	
U1600A-2H-CC	13.64	52	55	53.5	Extra-neritic	Upper bathyal (200-600 m)
U1600B-3F-CC	15.60	83		83		
U1600A-3H-CC	20.08	75	68	71.5	Sub-oceanic	
U1600B-5F-CC	25.61	68		68		Reworked benthics
U1600B-6F-CC	29.88	32		32	Extra-neritic	
U1600A-4H-CC	31.74	56	43	49.5		
U1600B-7F-CC	32.56	22		22	Outer neritic	
U1600A-5F-CC	37.02	60	70	65	Sub-oceanic	
U1600B-8F-CC	37.69	68	49	58.5	Extra-neritic	
U1600A-6F-CC	39.79	4	8	6	Inner neritic	
U1600A-7F-CC	43.28	0	0	0		Mid-lower bathyal (700-1000 m)
U1600A-8F-CC	51.03	43	32	37.5		
U1600B-11F-CC	53.94	53	53	53		
U1600A-9F-CC	56.00	36	30	33		
U1600B-12F-CC	58.52	30	31	30.5	Extra-neritic	
U1600A-10F-CC	60.66	28	33	30.5		Reworked benthics
U1600B-13F-CC	62.94	40	41	40.5		
U1600B-14F-CC	67.90	30	35	32.5		
U1600B-15F-CC	72.37	27	32	29.5	Outer neritic	
U1600A-13X-CC	74.76	52	54	53		
U1600B-17F-CC	82.06	38	46	42		
U1600B-18F-CC	82.16	50	33	41.5		
U1600C-3R-CC	84.87	35	37	36	Extra-neritic	
U1600B-19F-CC	91.65	31	34	32.5		
U1600C-4R-CC	95.22	34		34		
U1600C-5R-CC	105.53	39		39		
U1600C-6R-CC	117.19	24		24	Outer neritic	
U1600C-7R-CC	127.70	13		13		
U1600C-8R-CC	134.94	14		14		
U1600C-9R-CC	151.93	2		2	Inner neritic	
U1600C-10R-CC	161.97	1		1		
U1600C-11R-CC	165.89	39		39	Extra-neritic	

Figure F22. Foraminiferal oceanicity and paleowater depth estimates, Site U1600. Blue colors show relationship between oceanicity index and paleowater depth. Observers: AW = Adam Woodhouse, OK = Olga Koukousioura.

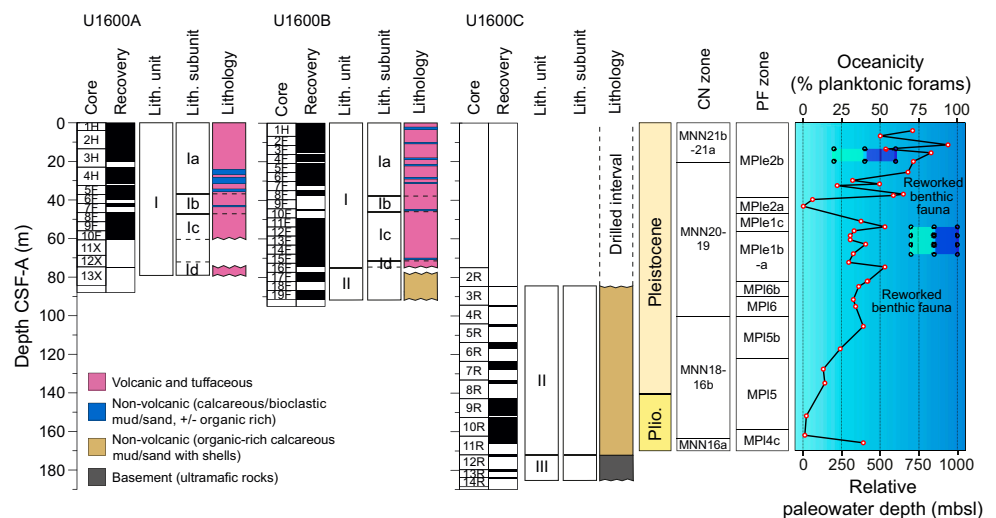


Figure F23. Biostratigraphic summary, Site U1600. CN = calcareous nannofossil, PF = planktonic foraminifer. Interpreted oceanicity: solid line/red points = interpreted oceanicity of Hayward et al. (1999), dashed line = extrapolation through barren/unreliable sample data. Interpreted paleowater depths: light blue points/shading = shallower paleowater depth interpretation, dark blue points/shading = deeper paleowater depth interpretation.

Holocene–Pleistocene foraminifer faunas indicate large fluctuations in relative paleowater depth and oceanicity, with highly variable planktonic abundances that range 0%–94% where foraminifera are present (Figure F23). The fauna are typical of Pleistocene Mediterranean biostratigraphic zones, primarily composed of *Neogloboquadrina incompta*, *Globigerina bulloides*, *Globigerina falconensis*, *Globigerina umbilicata*, *Globigerinella siphonifera*, *Globigerinita glutinata*, *Globigerinoides elongatus*, *Globigerinoides pyramidalis*, *Trilobatus trilobus* group, *Globigerinoides ruber* var. white, *Globigerinoides ruber* var. pink, *Globoconella inflata*, *Hirsutella scitula*, *Neogloboquadrina pachyderma*, *Orbulina universa*, *Turborotalita quinqueloba*, and *Globigerinella calida* through the Middle Pleistocene–Holocene, and the addition of species such as *Globoturbotalita woodi*, *Globorotalia bononiensis*, *Globigerinoides obliquus*, *Globigerinoides extremus*, and *Neogloboquadrina atlantica* (sinistral) through the Lower Pleistocene (Figure F24). Reworked specimens are common throughout the section.

Foraminiferal faunas are sufficiently common to biostratigraphically divide the Pleistocene into the majority of Pleistocene Mediterranean planktonic foraminiferal biostratigraphic zones determined via both primary and secondary marker species (Lirer et al., 2019):

- Zone MPle2b (0.00–0.53 Ma), Hole U1600A: 4.08–37.02 mbsf.
- Zone MPle2a (0.53–0.94 Ma), Hole U1600A: 39.79–43.28 mbsf.
- Zone MPle1c (0.94–1.21 Ma), Hole U1600B: 53.94 mbsf.
- Zones MPle1b–1a (1.21–1.79 Ma), Hole U1600B: 58.52–82.06 mbsf.
- Zone MPle6b (1.79–2.00 Ma), Holes U1600B and U1600C: 82.16–84.87 mbsf.
- Zone MPle6 (2.00–2.09 Ma), Holes U1600B and U1600C: 91.65–95.22 mbsf.
- Zone MPle5b (2.09–2.41 Ma), Hole U1600C: 105.53–117.19 mbsf.

The faunal criteria on which these age assignments are based are given in the sections below.

6.2.1.1. Zone MPle2b (0.00–0.53 Ma)

Samples 398-U1600A-1H-CC, 30–31 cm, to 5F-CC, 8–10 cm (4.08–37.02 mbsf), are assigned to Zone MPle2b based on the paracme top of sinistrally coiled *Neogloboquadrina* spp. (<0.51 Ma) occurring between Samples 5F-CC, 8–10 cm, and 6F-CC, 0–3 cm (37.02–39.79 mbsf), above which this species is more common.

6.2.1.2. Zone MPle2a (0.53–0.94 Ma)

Samples 398-U1600A-6F-CC, 0–3 cm, to 7F-CC, 0–3 cm (39.79–43.28 mbsf), are assigned to Zone MPle2a based on the absence of sinistrally coiled *Neogloboquadrina* spp. (0.51–0.91 Ma) throughout this interval.

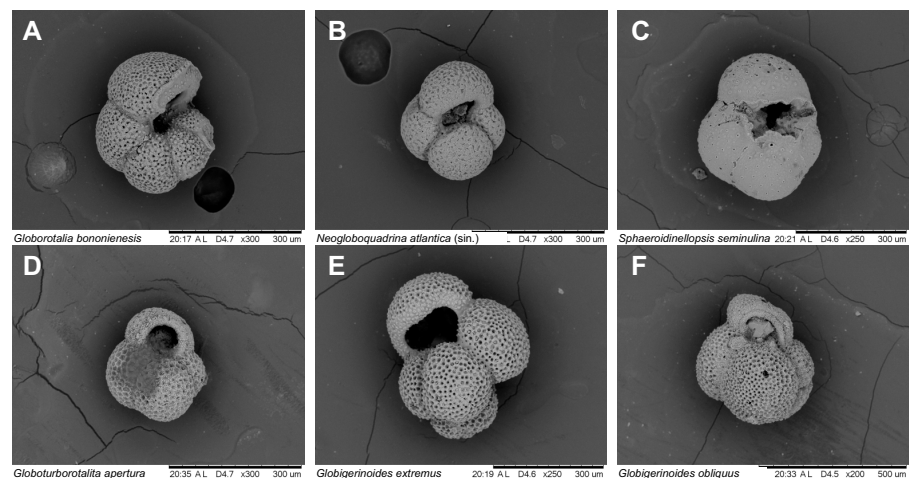


Figure F24. Planktonic foraminifera. A. *Globorotalia bononiensis* (398-U1600C-6R-CC, 13–18 cm). B. *Neogloboquadrina atlantica* (sinistral) (9R-CC, 11–14 cm). C. *Sphaeroidinellopsis seminulina* (398-U1600A-11R-CC, 0–5 cm). D. *Globoturbotalita apertura* (398-U1600C-7R-CC, 13–18 cm). E. *Globigerinoides extremus* (6R-CC, 13–18 cm). F. *Globigerinoides obliquus* (13–18 cm).

6.2.1.3. Zone MPle1c (0.94–1.21 Ma)

Sample 398-U1600B-11F-CC, 13–15 cm (53.94 mbsf), is assigned to Zone MPle1c based on the high abundances of sinistrally coiled *Neogloboquadrina* spp. (0.91–1.21 Ma) throughout this interval.

6.2.1.4. Zones MPle1b–1a (1.21–1.79 Ma)

Samples 398-U1600B-12F-CC, 13–15 cm, to 17F-CC, 16–18 cm (58.52–82.06 mbsf), are assigned to Zones MPle1b–1a based on the first common occurrence of sinistrally coiled *Neogloboquadrina* spp. (1.79 Ma) in Sample 17F-CC, 16–18 cm (82.06 mbsf). Additionally, Sample 13F-CC, 12–14 cm (62.94 mbsf), marks the first common occurrence of the benthic foraminifer *Hyalinea balthica* (1.492 Ma; Lourens et al. 1998).

6.2.1.5. Zone MPI6b (1.79–2.00 Ma)

Samples 398-U1600B-18F-CC, 14–16 cm, to 398-U1600C-3R-CC, 14–17 cm (82.16–84.87 mbsf), are assigned to Zone MPI6b based on the absence or rare occurrences of sinistrally coiled *Neogloboquadrina* spp. (>1.79 Ma) and the first common occurrence of *Globorotalia inflata* (1.99 Ma) in Sample 3R-CC, 14–17 cm (84.87 mbsf).

6.2.1.6. Zone MPI6 (2.00–2.09 Ma)

Samples 398-U1600B-19F-CC, 16–18 cm, to 4R-CC, 8–12 cm (91.65–95.33 mbsf), are assigned to Zone MPI6 because it represents the interval between the first occurrence of *G. inflata* (2.09 Ma) and the first common occurrence of *G. inflata* (2.00 Ma).

6.2.1.7. Zone MPI5b (2.09–2.41 Ma)

Samples 398-U1600C-5R-CC, 11–14 cm, to 6R-CC, 13–18 cm (105.53–117.19 mbsf), are assigned to Zone MPI5b based on the absence of *G. inflata* and the last common occurrence of *G. bononiensis* (2.41 Ma) between Samples 6R-CC, 13–18 cm, and 7R-CC, 13–18 cm (117.19–127.7 mbsf).

6.2.2. Pliocene foraminiferal biostratigraphy

Planktonic foraminifer assemblages from the Pliocene section of Hole U1600C are mostly well preserved, where specimens are rarely broken or exhibiting partially dissolved shell walls.

Pliocene planktonic foraminifer faunas indicate relatively shallow paleowater depths and oceanicity values that range 1%–39% when foraminifera are present (Figure F23). The fauna is typical of Mediterranean biostratigraphic Zones MPI5–MPI4c, primarily composed of *N. incompta*, *G. bulloides*, *G. glutinata*, *G. elongatus*, *T. trilobus* group, *G. ruber* var. white, *O. universa*, and more typical Upper Pliocene species such as *G. bononiensis*, *N. atlantica* (sinistral), *G. obliquus*, *Globoturborotalita decoraperta*, *G. woodi*, *Globoturborotalita apertura*, *Sphaeroidinellopsis* s.l., and *Dentoglobigerina altispira* (Figure F24).

Foraminiferal faunas are sufficiently common to biostratigraphically divide the Pliocene into two Upper Pliocene Mediterranean planktonic foraminiferal biostratigraphic zones determined via both primary and secondary marker species (Lirer et al., 2019):

- Zone MPI5 (2.41–3.19 Ma), Hole U1600C: 127.7–161.97 mbsf.
- Zone MPI4c (3.19–3.31 Ma), Hole U1600C: 165.89 mbsf.

The faunal criteria on which these age assignments are based are given in the sections below.

6.2.2.1. Zone MPI5 (2.41–3.19 Ma)

Samples 398-U1600C-7R-CC, 13–18 cm, to 10R-CC, 13–15 cm (127.7–161.97 mbsf), are assigned to Zone MPI5 based on the consistent common occurrences of *G. bononiensis* (last common occurrence = 2.41 Ma), the first occurrence of *N. atlantica* (sinistral) (first occurrence = 2.72 Ma), and the absence of *Sphaeroidinellopsis* s.l. (last occurrence = 3.19 Ma).

6.2.2.2. Zone MPI4c (3.19–3.31 Ma)

Sample 398-U1600C-11R-CC, 0–5 cm (165.89 mbsf), has been assigned to biostratigraphic Zone MPI4c based on the last occurrences of *D. altispira* (3.17 Ma) and *Sphaeroidinellopsis* s.l. (3.19 Ma) in this sample.

6.2.3. Planktonic foraminiferal oceanicity

Planktonic foraminifer abundances are highly variable, ranging 0%–94% when foraminifera are present (Figures F22, F23). Oceanicity values are rarely in good agreement with benthic foraminiferal paleowater depth indicators (Figure F24), indicating an overall general decrease in oceanicity from suboceanic (200–1000 mbsl) to inner neritic (0–50 mbsl) conditions downcore. Despite this trend, there is a notable interval of reduced oceanicity from Sample 398-U1600A-6F-CC, 0–3 cm, to 7F-CC, 0–3 cm (39.79–43.28 mbsf), where foraminifer faunas indicate inner neritic (0–50 mbsl) oceanicity values (Figure F23).

6.2.4. Benthic foraminifera paleowater depths

Benthic foraminiferal distributions are highly variable, indicating paleowater depths from uppermost bathyal (200 m) to lower bathyal (>1000 m) (Figure F23). The low abundances of benthic foraminiferal faunas in some samples (e.g., 398-U1600B-2F-CC, 20–22 cm, to 8F-CC, 0–2 cm [11.46–37.69 mbsf]) are possibly correlated with rapid emplacement of volcanoclastic sediments and/or inhospitable environmental conditions. Samples 398-U1600A-2H-CC, 24–27 cm, to 3H-CC, 11–14 cm (13.64–20.08 mbsf), record an uppermost to upper bathyal (200–600 m) paleowater depth, as indicated by the presence of species such as *Cassidulina* spp., *Bolivina* spp., *Bulimina* spp., and *Cibicides pachyderma*. Samples 398-U1600B-11F-CC, 13–15 cm, to 14F-CC, 11–13 cm (53.94–67.9 mbsf), indicate mid- to lower bathyal (700 to >1000 m) paleowater depths due to the presence of lower bathyal markers such as *Cibicidoides wuellerstorfi*, signifying deposition in paleowater depths >1000 mbsl (Figure F23). Commonly reworked benthic foraminifera in Samples 398-U1600A-4H-CC, 83–85 cm, to 6F-CC, 0–3 cm (31.74–39.79 mbsf); 398-U1600B-5F-CC, 23–27 cm, to 8F-CC, 0–2 cm (25.61–37.69 mbsf); and 15F-CC, 18–19 cm, to 18F-CC, 14–16 cm (72.37–82.16 mbsf), likely represent downslope reworking.

7. Paleomagnetism

Paleomagnetic analysis at Site U1600 focused on measurement and demagnetization of archive-half sections to determine magnetostratigraphic age controls, together with demagnetization of 17 discrete samples. Results are discussed below using depths on the CCSF scale to allow data from all three holes at this site to be combined on a common depth scale.

The uppermost 40 m of the sequence sampled at this site carries normal polarity remanences acquired during the Brunhes Chron (C1n) (Figures F25, F26), with a mean inclination of statistically acceptable principal component directions of $+51.2^\circ$ ($k = 62.0$; $\alpha_{95} = 1.2$; $n = 163$), in good agreement with the expected geocentric axial dipole inclination of $+56^\circ$ at this site.

The interval 55.3–83.0 mbsf is marked by wide variations in inclinations estimated using automated principal component analysis; calculation of Fisher mean directions of magnetization using the data from the 15, 20, and 25 mT demagnetization steps; and inclinations of the 25 mT step (see **Paleomagnetism** in the Site U1591 chapter [Druitt et al., 2024b]). This can only be explained by variable partial to complete diagenetic overprinting of a reversed polarity depositional remanent magnetization acquired during the Matuyama Chron by a chemical remanence acquired during a subsequent normal polarity chron. Some samples in this interval do show clear reversed polarity, however (Figure F26). Biostratigraphic constraints indicate that the top of this interval postdates the end of the Jaramillo Subchron (C1r.1n), so we assign it to reversed Chron C1r.3r (Gradstein et al., 2020). On this basis, we assign the Brunhes/Matuyama boundary (0.773 Ma; Gradstein et al., 2020) to the interval between 40.0 and 55.3 mbsf, where no data are available (Figure F25), and this interval must also span Chrons C1r.1r and C1r.1n.

There is a clear change in magnetic polarity on either side of the interval 83.0–87.7 mbsf, with reversed and normal polarity principal component directions clustering around the expected values of $\pm 56^\circ$ (Figures F25, F26). Biostratigraphic markers for interval 1.73–1.79 Ma indicate this reversal must correspond to the boundary between Chrons C1r.3r and C2n (Olduvai Subchron) at 1.775 Ma (Gradstein et al., 2020). The base of Chron C2n is undefined.

Discrete samples on either side of the interval 117.0–123.8 mbsf (where no data are available) indicate another change in polarity (Figure F25), with normal polarity inclinations seen in both archive-half section and discrete sample data from 123.8 to 165.8 mbsf (Figures F25, F26). Biostratigraphic markers indicate that this normal polarity interval corresponds to Chron C2An.1n. This places the top of this chron (2.610 Ma; Gradstein et al., 2020) in the interval 117.0–123.8 mbsf.

Very limited data are available below 165.8 mbsf. One discrete sample showing clear reversed polarity (Figure F25, F26) was collected from the ultramafic rocks of the basement, but this has no magnetostratigraphic significance because the age of this unit significantly predates the tectonic development of the volcanic rift zone.

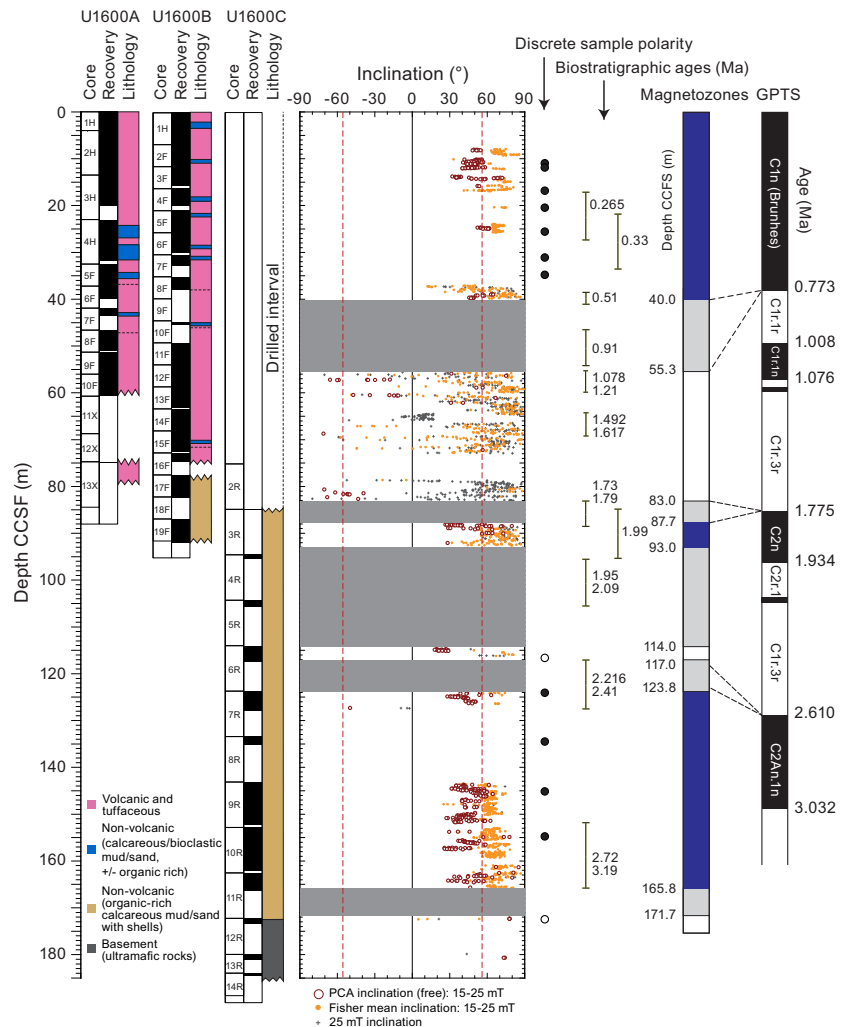


Figure F25. Archive-half section magnetic inclinations, Site U1600. Red dashed lines = geocentric axial dipole inclinations expected at this site. Solid/open circles = normal/reversed polarity of discrete samples, dark blue/white = normal/reversed polarity magnetozones, gray shading = no available paleomagnetic data. GPTS = geomagnetic polarity timescale.

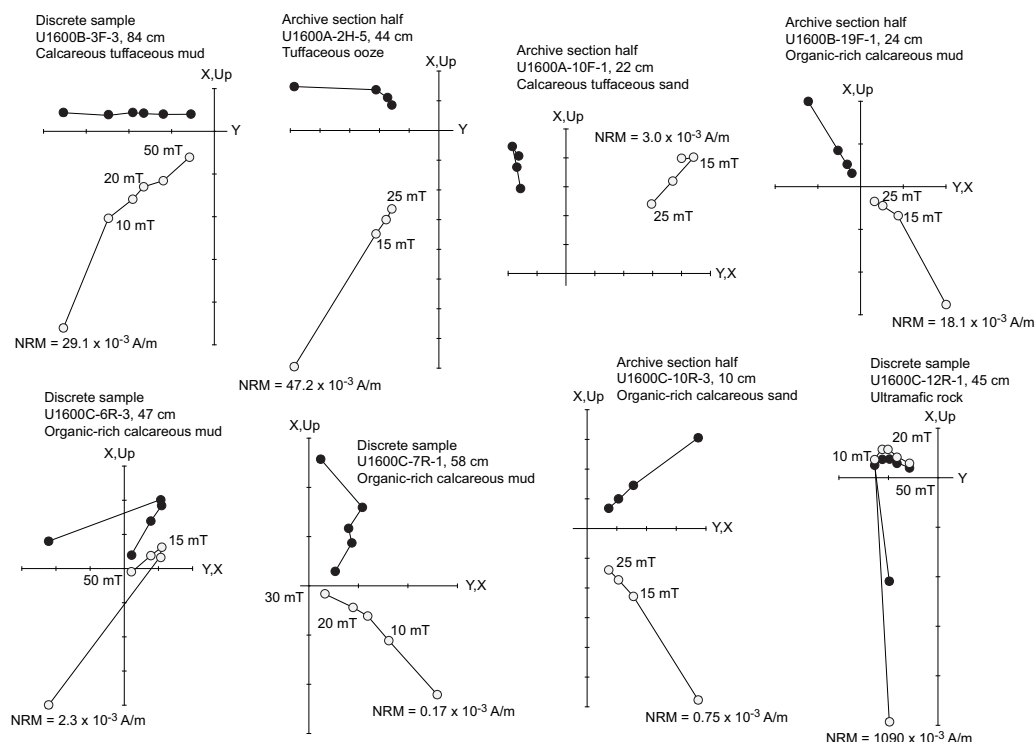


Figure F26. AF demagnetization of discrete samples and archive-half sections, Site U1600. Solid circles = projection onto horizontal plane, open circles = projection onto vertical plane. NRM = natural remanent magnetization.

8. Physical properties

The typical increases of bulk density and P -wave velocity with increasing depth are not clearly documented at Site U1600. MS can be high and highly variable in volcanoclastic layers, is low in the tuffaceous and ooze-dominated layers, and is very high in the basement rocks. Grain density in volcanic and tuffaceous layers is occasionally $<2.0 \text{ g/cm}^3$.

8.1. Whole-round GRA, MS, P -wave velocity, and NGR

Figure F27 summarizes data collected on whole-round cores measured on the GRA densitometer, MS loop, and P -wave logger on the WRMSL, as well as natural gamma radiation (NGR). All cited depths are on the CSF-A scale. A few possible systematic sources of error in these data should be considered when interpreting absolute values and trends.

- Redistribution of unconsolidated volcanoclastic materials in core liners during coring and in the core receiving facility can lead to sorting by particle size and density.
- APC cores often contained large amounts of water. As a result, WRMSL measurements of bulk density are systematically lower than discrete measurements at this site.
- Whole-round measurements in sections that match discrete measurement values on split cores are thus most reliable.

MS is high and highly variable in the volcanoclastic deposits in Sections 398-U1600A-1H-1 through 5H-1 (0–34 mbsf) and 398-U1600B-1H-1 through 7F-1 (0–32 mbsf). MS is exceptionally high in the basement below 165 mbsf ($>10,000 \times 10^{-5} \text{ SI}$). Two high P -wave velocity layers at ~17 and ~71 mbsf correspond to lapilli-ash layers.

Overall, above the basement there is no clear systematic increase in bulk density and P -wave velocity with increasing depth. Whether the lack of a signature of compaction is real, a consequence of the limited depth of the holes, or an artifact of drilling and recovery disturbances cannot be readily determined.

8.2. Discrete measurements

8.2.1. Material strength

Sediment strength was measured with a pocket penetrometer (PP) on the catwalk immediately after section splitting. Automated vane shear (AVS) strength measurements were made on working-half sections. Totals of 4 and 11 AVS measurements were made on working-half sections from Holes U1600A and U1600B, respectively (Figure F28; Table T10). Totals of 7 and 13 PP measurements were made (Table T10) on fine-grained materials that remained intact in core liners upon recovery on the catwalk in Holes U1600A and U1600B, respectively. There is no strong trend in

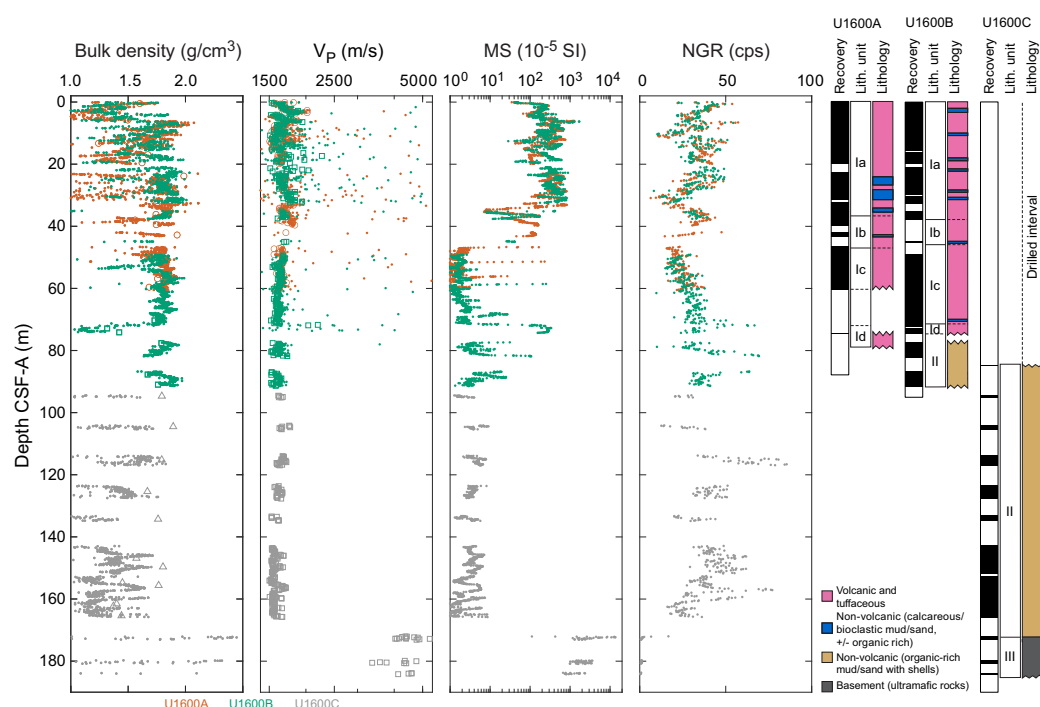


Figure F27. Physical properties, Site U1600. Dots = whole-round measurements, open symbols = discrete measurements. cps = counts per second.

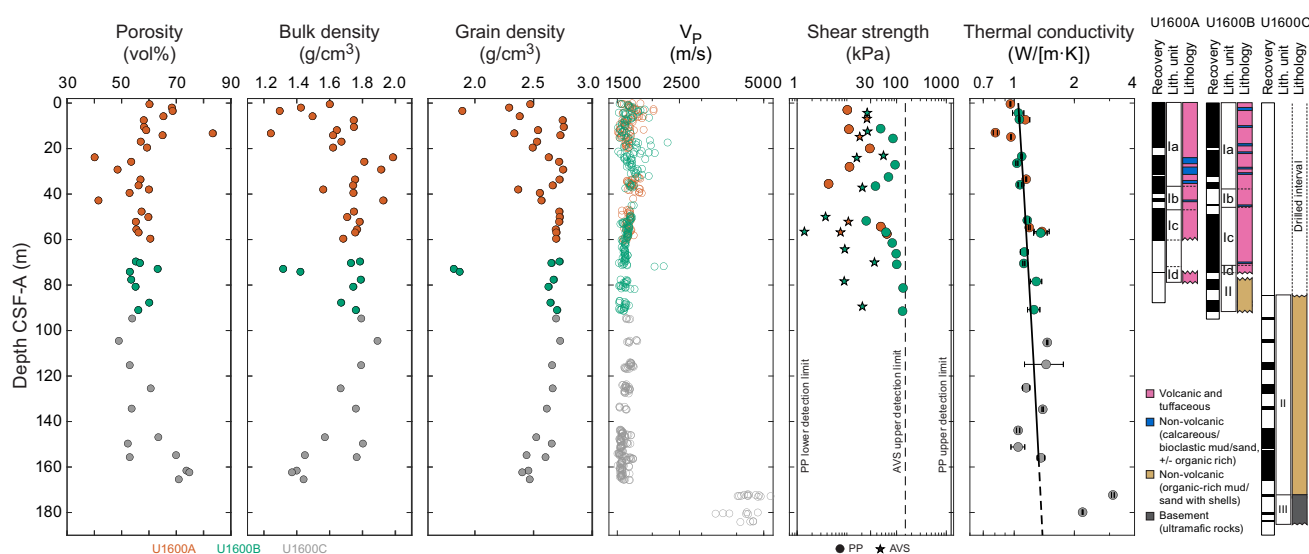


Figure F28. Discrete physical properties measurements, Site U1600. Shear strength: dashed/solid lines = AVS and PP measurement limits, respectively. Thermal conductivity: solid line = linear fit over the depth range of included data, dashed line = greater depths. P -wave velocity and MS are plotted on logarithmic scales.

shear strength with increasing depth except for slight decreases in shear strengths above and below intervals of no recovery at 40–47 and 38–49 mbsf in Holes U1600A and U1600B, respectively.

8.2.2. *P*-wave velocity

Totals of 149, 221, and 223 discrete *P*-wave velocity measurements were conducted on Hole U1600A, U1600B, and U1600C working-half sections, respectively (Figure F28; Table T11). Above the basement in the uppermost 166 m, *P*-wave velocity ranges 1.51–2.27 km/s (mean = 1.63 km/s). Discrete measurements of *P*-wave velocity on working-half core sections show similar downhole variations as those measured using the WRMSL on whole-round cores (Figure F27). There is no clear trend of increasing velocity with increasing depth. Rather, the highest velocities, and largest variability, are found in the volcanoclastic materials at ~3, 17, and 72 mbsf.

Discrete measurements were sometimes challenging to make in coarse, pumice-rich deposits because signals were especially noisy. Furthermore, particle size was a substantial fraction of the distance between the two calipers used to make the measurements; hence, the measurements may not be representative of true bulk properties (i.e., sample size is not much larger than a representative elementary volume).

In basement rocks below 166 mbsf, *P*-wave velocity ranges 3.37–5.29 km/s (mean = 4.27 km/s).

8.2.3. Moisture and density

Totals of 25, 8, and 13 discrete samples were collected from Holes U1600A, U1600B, and U1600C, respectively, to conduct moisture and density (MAD) measurements (Figure F28; Table T12). Bulk density derived using MAD measurements on discrete samples should be more reliable than GRA bulk density data from WRMSL measurements on whole-round cores, although in both cases, coring and recovery disturbances may have impacted measured values.

Above the basement in the uppermost 166 m, porosity ranges 40–83 vol% (mean = 59 vol%; standard deviation = 8 vol%). Bulk density ranges 1.24–2.00 g/cm³ (mean = 1.67 g/cm³). Bulk density is lower in volcanoclastic materials than ooze-dominated sediment at the same depth; nonvolcanic muds and sands in the lower part of Lithostratigraphic Unit II also have low bulk density values. Grain density ranges 1.82–2.77 g/cm³ (mean = 2.56 g/cm³; median = 2.65 g/cm³). The lowest values in volcanoclastic subunits average 2.56 g/cm³ (median = 2.65 g/cm³). The lowest grain densities are found in lapilli layers at 3.4 mbsf in Hole U1600A and 72.6–74.4 mbsf in Hole U1600B. In Unit II, organic-rich calcareous muds have lower grain densities than organic-rich calcareous sands. The usual trends of decreasing porosity and increasing bulk density with increasing depth are not seen in Units I and II at Site U1600.

One sample from basement Unit III was collected from Hole U1600C to conduct MAD measurements (not shown on Figure F28). Porosity is 7%, which is significantly lower than the sediments of Units I and II. Bulk density is 2.66 g/cm³, and grain density is 2.79 g/cm³.

8.2.4. Thermal conductivity

A total of 27 thermal conductivity measurements were made on selected working-half sections (Table T13; Figure F28). Measurements could not be made on coarse volcanoclastic materials; hence, the values we report are not representative of the full range of recovered lithologies. In the

Table T10. Shear strength, Site U1600. [Download table in CSV format.](#)

Table T11. *P*-wave velocity, Site U1600. [Download table in CSV format.](#)

Table T12. MAD measurements, Site U1600. [Download table in CSV format.](#)

Table T13. Thermal conductivity, Site U1600. [Download table in CSV format.](#)

uppermost 166 m above basement rocks, mean thermal conductivity is 1.16 W/(m·K), and a linear fit to these data is

$$k(z) = (1.048 \pm 0.048) + (1.78 \pm 0.58)z \text{ W/(m·K)},$$

where z is in kilometers and reported $\pm$ values are standard errors. The two basement samples from Unit III have thermal conductivities markedly higher, at 3.1 and 2.2 W/(m·K).

9. Geochemistry

9.1. Volcaniclastic bulk geochemistry

To determine the geochemistry of the volcanic and tuffaceous materials, seven tephra samples were handpicked from various layers within Hole U1600A. Following cleaning, grinding, fusion, and dissolution, the materials were analyzed shipboard for major (Si, Al, Fe, Mg, and Ca), minor (Ti, Mn, Na, K, and P), and trace (Sc, V, Cr, Co, Ni, Cu, Zn, Rb, Sr, Y, Zr, Nb, Ba, Ce, and Nd) elements using inductively coupled plasma–atomic emission spectroscopy (ICP-AES) (see [Geochemistry](#) in the Expedition 398 methods chapter [Kutterolf et al., 2024]). Several unknown samples were run multiple times to determine analytical reproducibility.

9.1.1. ICP-AES: major, minor, and trace elements

Of the volcaniclastic units sampled, one was classified as a basalt, one as a basaltic andesite, two as trachyandesites, and two as dacites or trachydacites (Table [T14](#); Figure [F29](#)). Bulk chemistry values are less evolved than glass chemistry reported in Kutterolf et al. (2021), as expected due to bulk analyses including both minerals and glass.

Concentrations are reported for all analyzed trace elements, but Ce, Cr, Cu, Nb, Ni, P, Rb, S, and V were below detection limits in the majority of samples and are not shown for respective samples in Table [T14](#); volcaniclastic analytical errors are $\pm 1\%$ for major elements and $\pm 5\%$ – 10% for trace

Table T14. ICP-AES data for major, minor, and trace elements, Site U1600. [Download table in CSV format.](#)

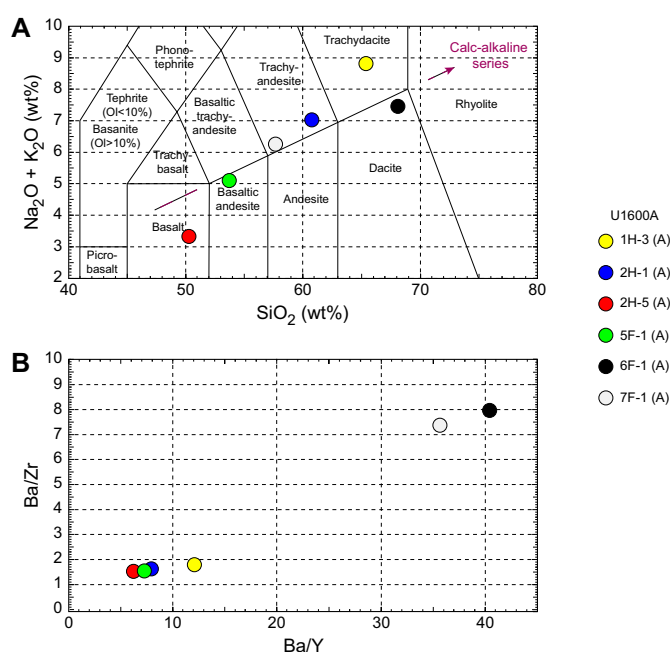


Figure F29. ICP-AES analyses of selected volcaniclastic units used to discriminate between potential volcanic sources, Hole U1600A. A. Total alkali vs. SiO_2 plot with the rock nomenclature of Le Maitre et al. (2002) overlain used for sample naming. Ol = olivine. B. Ba/Y vs. Ba/Zr plot used to correlate samples.

elements (see [Geochemistry](#) in the Expedition 398 methods chapter [Kutterolf et al., 2024]). Trace element ratios were used to broadly discriminate between the volcanic centers of Kolumbo, Santorini, and Christiana.

9.2. Interstitial water geochemistry

To determine the inorganic constituents of IW, a total of 14 water samples were taken from the mudline and whole-round squeezing of sediment intervals from Holes U1600A (4 samples), U1600B (5 samples), and U1600C (5 samples). Aliquots of IW were used for shipboard analyses, and the remaining water was taken for shore-based analysis following protocols specified by individual scientists. The retrieved pore waters were analyzed shipboard for salinity, alkalinity, pH, major anions (Cl^- , SO_4^{2-} , and Br^-), major cations (Ca^{2+} , Na^+ , Mg^{2+} , and K^+), and major (S, Ca, Mg, K, and Na) and minor (B, Ba, Fe, Li, Mn, P, Si, and Sr) elements using the methods described in [Geochemistry](#) in the Expedition 398 methods chapter (Kutterolf et al., 2024).

9.2.1. Salinity, alkalinity, and pH

Salinity ranges 38–40 (Table T15; Figure F30). The highest salinity value of 40 was recorded at the mudline (Section 398-U1600A-1H-1) and at 61.5 mbsf (Section 398-U1600B-13F-2).

Alkalinity ranges from 2.5 mM at the mudline to 3.3 mM at 54 mbsf at the base of Hole U1600A (Table T15; Figure F30). An alkalinity high of 3.5 mM was measured at 69 mbsf in Hole U1600B. Below 115.3 mbsf, alkalinity decreases with depth.

Values for pH exhibit a narrow range, 7.5–7.9, and generally decrease downhole (Table T15; Figure F30).

9.2.2. Bromide, chloride, boron, sodium, potassium, magnesium, calcium, and sulfate

IW concentrations of Br^- , Cl^- , B, Na^+ , K^+ , Mg^{2+} , Ca^{2+} , and SO_4^{2-} are provided in Table T16 and plotted in Figure F31. Bromide, Cl^- , and Na^+ show similar trends and generally decrease with depth with some scatter. Calcium concentrations increase with depth (11–16 mM). Potassium ranges 11–13 mM. Boron increases with depth but with some variation, including minima at 35.5 and 148.1 mbsf and maxima at 90.5 and 164.6 mbsf. Concentrations of Mg^{2+} and SO_4^{2-} decrease from their highest values at 8.4 mbsf to a minimum at 35.5 mbsf before increasing slightly to 61.5

Table T15. IW alkalinity, pH, and salinity, Site U1600. Download table in CSV format.

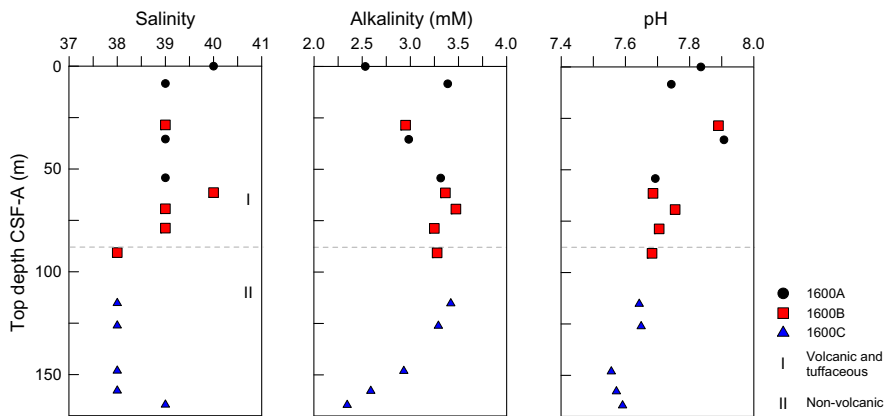


Figure F30. IW salinity, alkalinity, and pH, Site U1600. Lithostratigraphic Units I and II are described in Lithostratigraphy.

Table T16. IW data for major anions and cations measured using ion chromatography and major and minor elements measured using ICP-AES, Site U1600. Download table in CSV format.

mbsf. Below 61.5 mbsf, both Mg^{2+} and SO_4^{2-} decrease to minimum values observed at 115.3 mbsf before increasing again downhole.

9.2.3. Lithium, strontium, manganese, barium, silicon, iron, and phosphorus

IW concentrations of Li, Sr, Mn, Ba, and Si are provided in Table T16 and Figure F32. Lithium shows little variation with depth aside from one elevated value at 8.4 mbsf that corresponds to the maximum Mn value. Strontium increases initially with depth to a maximum of $284.7 \mu\text{M}$ at 90.7

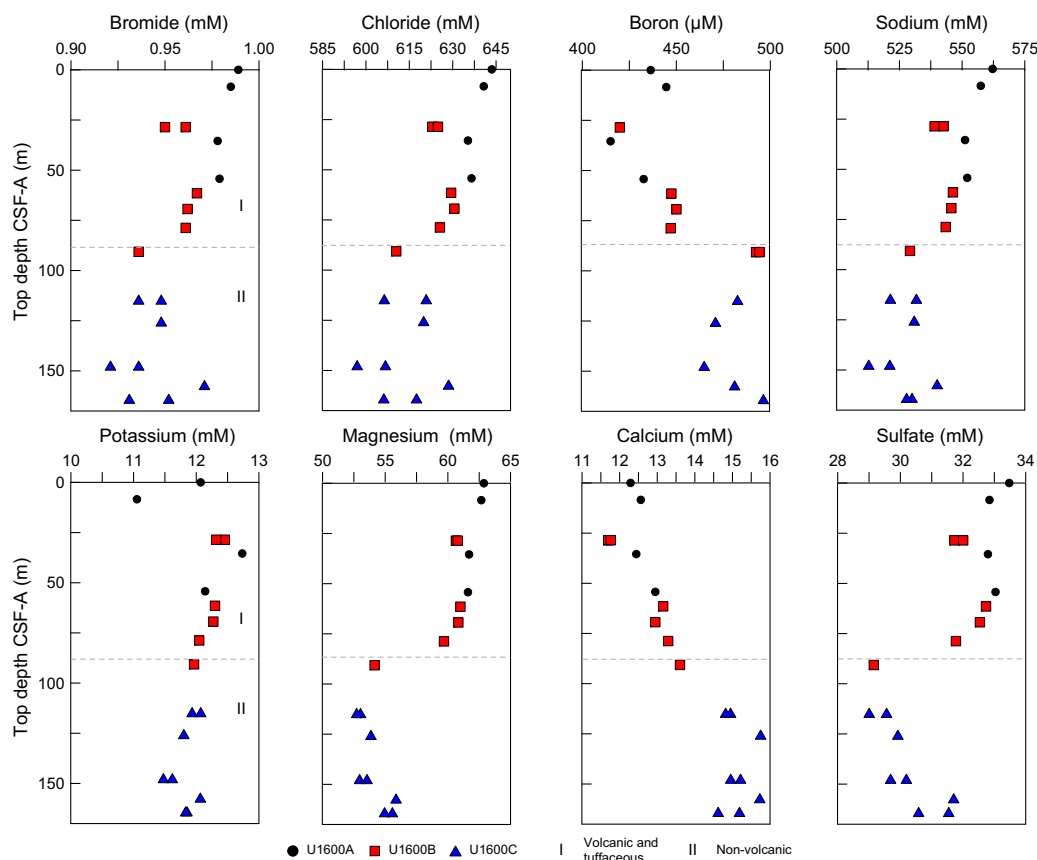


Figure F31. IC and ICP-AES concentrations of Br, Cl, B, Na, K, Mg, Ca, and SO_4^{2-} in IW samples, Site U1600. Lithostratigraphic Units I and II are described in Lithostratigraphy.

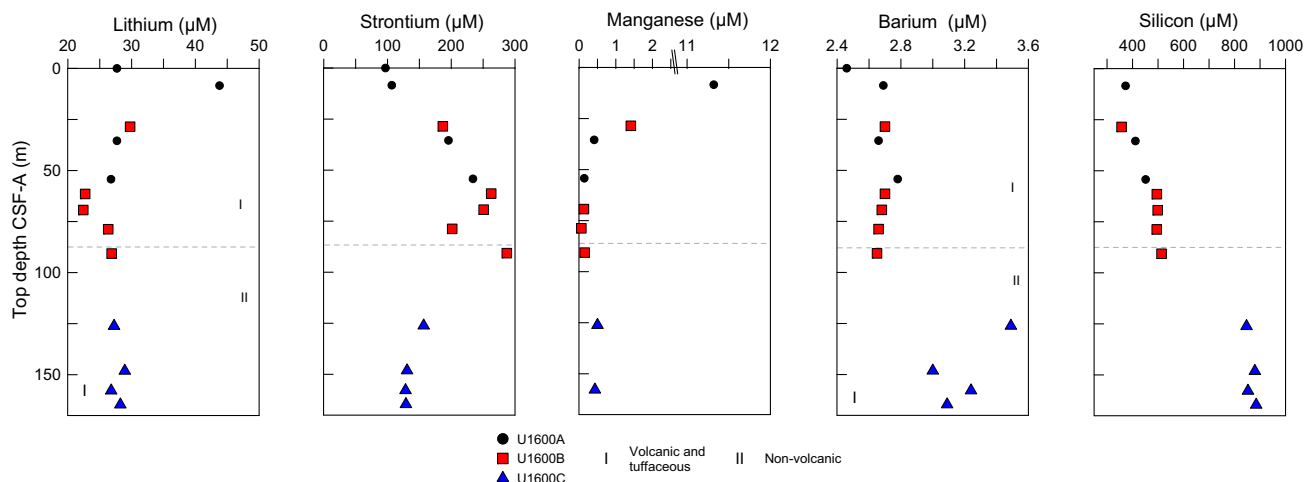


Figure F32. ICP-AES concentrations of Li, Sr, Mn, Ba, and Si in IW samples, Site U1600. Lithostratigraphic Units I and II are described in Lithostratigraphy.

mbsf and decreases below this depth. Barium and Si both remain relatively constant until they increase in concentration below 115.3 mbsf. Iron and P are both below the detection limit in a majority of samples and are not plotted.

9.3. Sediment bulk geochemistry

A total of 13 sediment samples were analyzed for bulk geochemistry (Table T17). All were analyzed for total carbon, hydrogen, and nitrogen (CHN) and for inorganic carbon and carbonate content following the protocols explained in **Geochemistry** in the Expedition 398 methods chapter (Kutterolf et al., 2024). For CHN analysis, duplicates were run for ~10% of the samples to determine analytical reproducibility (standard deviations: N = 0.01; C = 0.03). Carbonate values are highest in Lithostratigraphic Unit I (see **Lithostratigraphy**), peaking at 82 wt% at 54 mbsf, which is characterized by volcanic and tuffaceous material but contains some calcareous ooze lithologies (Figure F33). Notably, carbonate values decrease and total organic carbon (TOC) increases in Lithostratigraphic Unit II, which is characterized by organic-rich calcareous mud/sand.

9.3.1. Sapropel identification

TOC was calculated using total carbon and inorganic carbon values. Following the convention set forth by Kidd et al. (1978), units with TOC >2.0 wt% were identified as sapropels, and units with TOC of 0.5–2.0 wt% were identified as sapropelitic. Using these values, four units were identified as sapropels and seven units were identified as sapropelitic (Figure F33).

9.4. Headspace gas analysis

Headspace gas analyses were performed at a resolution of one sample per full-length core (9.5 m advance) or one sample every other core for half-length cores (4.7 m advance) throughout Hole U1600A. Analyses were resumed on Hole U1600B when depth of penetration exceeded recovery in Hole U1600A. Analyses were resumed on Hole U1600C when depth of penetration exceeded recovery in Hole U1600B. The aim was to monitor the presence and abundance of C₁–C₃ hydrocarbons as part of the standard IODP safety protocol (Pimmel and Claypool, 2001). A total of 19 headspace gas samples were analyzed using gas chromatography (GC) (see **Geochemistry** in the Expedition 398 methods chapter [Kutterolf et al., 2024]). Methane, ethane, and propane concentrations were below the detection limit in all measured samples.

Table T17. Total inorganic carbon, percent CaCO₃, total carbon, total nitrogen, and TOC data, Site U1600. [Download table in CSV format.](#)

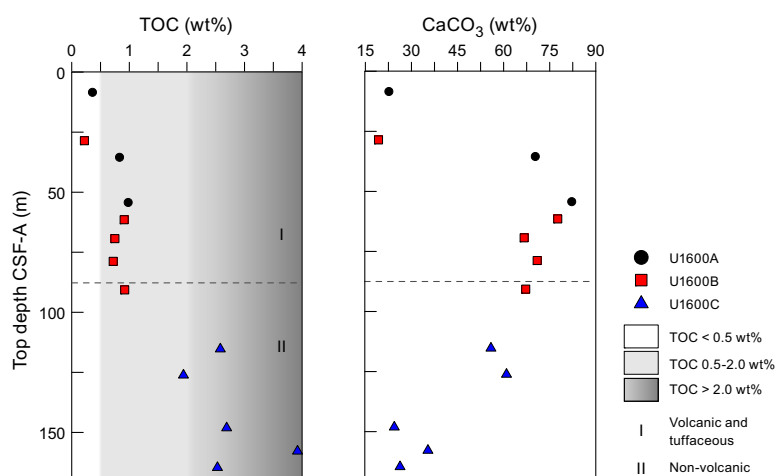


Figure F33. TOC and carbonate, Site U1600. Lithostratigraphic Units I and II are described in **Lithostratigraphy**. Sapropel units follow Kidd et al. (1978).

Table T18. Whole-round core samples, Site U1600. [Download table in CSV format.](#)

10. Microbiology

At Site U1600, three whole-round core samples were collected. After the cores were transferred to the core receiving platform, whole-round samples were collected based on sediment color. Reddish to orange coloration was used as an indicator for iron-rich layers that often host the microbial communities we sought. To choose the appropriate incubation temperature, the core temperature was measured in the section above the sample.

The whole-round samples were transferred to the laboratory, and the following subsamples for different analyses were collected: POLYmeta, POLYmicro, POLYanae, POLYcell, POLYFe, POLY-microscope, and POLYFe-agar. In total, 21 subsamples were collected for microbiological analysis from the whole-round core samples.

A list of all samples, including duplicates, from this site is shown in Table [T18](#).

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