

1 Title. Critical questions about CSEP, in the spirit of Dave, Yan, and Ilya.

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8

9 Abstract. In honor of our dear departed friends Yan Kagan, Dave Jackson and Ilya Zaliapin,

10 we propose a selection of broad questions regarding earthquake forecasting and especially the

11 Collaboratory for the Study of Earthquake Predictability (CSEP) in particular, and give our

12 thoughts on their answers. This article reflects our opinions, not necessarily those of Yan

13 Kagan, Dave Jackson, and Ilya Zaliapin, and not necessarily those of the seismological

14 community at large. Rather than to provide definitive answers, we hope to provoke the

15 reader to think further about these important topics. We feel that Dave Jackson in particular

16 might have liked this approach and may have seen this as an appropriate goal.

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20 assessment, statistical seismology.

21 1. Is the Collaboratory for the Study of Earthquake Predictability (CSEP) worthwhile? How
22 can it be improved?

23

24 While our response to the first question is unequivocally in the affirmative, we must first
25 admit that CSEP has not yet come close to achieving all that was initially hoped of it. When
26 CSEP was formed, many anticipated that the experiments would lead to clear and decisive
27 improvements in earthquake forecasting, would indicate which models are superior and
28 which are inferior, would highlight ways to improve models, and ultimately would lead to
29 marked improvements in our ability to forecast large earthquakes. Perhaps some of these
30 goals will be fulfilled in the future, but the steps so far in these directions have been very
31 small.

32

33 It can, at times, seem unclear if anyone is actually making any use of CSEP results in
34 practice, and how much seismic hazard models have been improved using CSEP results.

35 While the Uniform California Earthquake Rupture Forecast 3 (UCERF3) model (Field et al.
36 2017) has included the model of Helmstetter et al. (2007), which performed best in the 5-
37 year RELM experiment (Schorlemmer et al. 2010, Zechar et al. 2013, Bayona et al. 2022), as
38 one branch for the background seismicity part, UCERF3 remains driven by the fault-based

forecasts of the large earthquakes (Field et al. 2017). CSEP has never, to our knowledge, influenced the fault-based parts of hazard models that provide the bulk of the hazard.

On the other hand, from a statistical standpoint, CSEP is a most remarkable scientific achievement. It is a kind of gold standard that other areas of statistical application can only dream of achieving. In wildfire forecasting, for instance, many of the proposed models are not even well-defined and would result in 0 likelihood given data. In epidemiology, the models most often used to forecast the spread of diseases like Ebola or Covid-19 are at least 50 years old and scant attention is given to their goodness-of-fit (Kresin et al. 2021). Further, one cannot reasonably expect government officials and industrial practitioners instantly to make use of scientific advances. Scientific progress has almost always been painfully slow and methodical. There are so many possible examples here, but a recent one is Dave Jackson and Yan Kagan debunking the characteristic earthquake hypothesis (Kagan et al. 2012). It has typically taken decades at least for most scientists and professionals adequately to accept and employ the results of careful scientific work. It might be a bit naive to expect practitioners and other researchers to adjust quickly to reports that a given model does not fit well to data. Such testing is definitely progress nonetheless.

Analyses of CSEP results have pointed out that Epidemic-Type Aftershock Sequence (ETAS) models tend to fit well, at least on relatively short-term scales. ETAS models were initially proposed by Ogata (1988) to describe the times and magnitudes of earthquakes, and were

60 subsequently extended to model spatial-temporal-magnitude catalogs in Ogata (1998).
61 Subsequently, a host of slight modifications have been proposed (e.g. Sornette and Sornette
62 1999, Helmstetter and Sornette 2002, Console et al. 2003, Ogata et al. 2003, Ogata 2004,
63 Ogata and Zhuang 2006, Marzocchi and Lombardi 2008, Ogata 2011, Zhuang 2012, Nandan
64 et al. 2017, Grimm et al. 2022, Iacoletti et al. 2022, Li and Pu 2022, Aso and Terai 2023). One
65 major problem, however, is that models such as ETAS typically do not help much in
66 forecasting the earthquakes we are most interested in. As their name indicates, ETAS models
67 are mostly useful for describing the frequency and spatial-temporal distribution of
68 aftershocks one expects to see following large earthquakes, or perhaps as a null model to
69 which alternative models might be compared. However, for purposes of planning, public
70 safety, building codes, and most other purposes, what is really sought is the accurate
71 forecasting of the very largest events, or at least the estimation of their long-term frequency,
72 and when it comes to these tasks, most versions of ETAS seem to be scarcely better than a
73 simple homogeneous Poisson model. Indeed, most formulations of the ETAS model assume a
74 Gutenberg-Richter distribution of earthquake magnitudes with the magnitude of each
75 earthquake drawn independently of what occurred previously, and thus essentially the
76 model makes no effort to pinpoint where or when the rate of the largest earthquakes may be
77 higher relative to the rate of smaller events.

78

79 This may be an area where more statistical work can be of assistance. Current methods for
80 assessing the fit of earthquake forecast models emphasize overall measures of fit such as the

81 log-likelihood, or total number of events, or other summaries that do not adequately take
82 into account what aspects of the model we care most about. If, for instance, we care
83 exclusively about the model's ability to forecast the largest events, then it may be appropriate
84 to choose a goodness-of-fit measure that properly emphasizes this feature. For instance,
85 suppose one is given data on the times, locations, and magnitudes of n events, (t_i, x_i, y_i, m_i)
86 for $i = 1, \dots, n$, and let $\lambda(t, x, y, m)$ denote the modeled conditional intensity at spatial-temporal-
87 magnitude (t, x, y, m) , with λ_i representing the conditional intensity at point i . The log-
88 likelihood,

89

$$90 \quad L = \sum_{i=1}^n \log \lambda_i - \int \lambda(t, x, y, m) dt dx dy dm,$$

91

92 has a first term that properly rewards the model for accurately forecasting earthquakes
93 (where by "accurately forecasting", we mean positing a high value of λ where an earthquake
94 ends up occurring), and a second term that punishes the model for having high values of λ
95 elsewhere. However, the log-likelihood essentially rewards the model equivalently for
96 accurately forecasting a magnitude 3 event or a magnitude 7 event, despite the fact that
97 forecasting the latter event is so much more of interest.

98

99 Since forecasting the largest 5% of magnitudes are of most interest, one could instead use, as
100 a measure of fit, a summary such as the quotient

101
$$Q = \frac{\text{mean}\{\lambda_i: m_i > m^{[.95]}\}}{\text{mean}\{\lambda(t,x,y,m)\}},$$

102 for example, where $m^{[.95]}$ is the 95th percentile of the magnitude distribution, perhaps
 103 estimated based on prior seismicity, and the denominator $\text{mean}\{\lambda(t,x,y,m)\}$ may be estimated
 104 e.g. using several thousand locations selected at random from the space-time-magnitude
 105 observation region. The higher the value of Q the better the model appears to be forecasting
 106 the spatial-temporal locations of the largest 5% of events. For a homogeneous Poisson model
 107 with uniform magnitude density, Q will be close to 1. Any model that adequately accounts
 108 for the spatial inhomogeneity of seismicity will have $Q > 1$, as it should since such a model
 109 will tend to vastly outperform a homogeneous Poisson model at forecasting the largest
 110 events. Among competing models, the model that forecasts the larger events more accurately
 111 will tend to be the model with higher Q especially if all the models are similarly calibrated
 112 overall [i.e. $\text{mean}\{\lambda(t,x,y,m)\}$ is close to the overall rate of seismicity] which can readily be
 113 checked via other methods, such as the N -test.

114

115 Forecasting large earthquakes as point sources as CSEP defines them for the purpose of
 116 testing bears its problems. Modelers can employ the knowledge of faults and distribute the
 117 hypocenter probabilities along the fault but this is making a model fit to the test design and
 118 not the (preferred) other way around. This gets us back to the problem of whether CSEP can
 119 influence the fault-based forecasts of hazard models. Clearly, adequate testing procedures for
 120 such models are needed and were discussed many times within the CSEP community.

121 However, the unambiguous identification of fault segments ruptured in earthquakes is
122 already a problem, not to mention the incompleteness of fault models. This is an unfortunate
123 disconnect between common practice in hazard modeling and testing possibilities.

124

125 Another problem with some CSEP results is that the observation that model A fits better
126 than model B does not necessarily directly tell us how to improve either of the models.
127 Further, if model A offers superior fit to model B over a 5-year period, it is unclear whether
128 this means model A is likely to outperform model B in the future. If not, then what does
129 testing tell us? We have often observed that models use any seemingly suitable statistical
130 distributions in fitting and then use such fitted models for forecasting. In one paper about
131 testing intensity-prediction equations, it was shown that the models fitted to functions
132 reproducing basic physical principles of wave propagation have higher forecasting
133 capabilities [Mak et al. 2015]. But is CSEP able to discriminate between overfitting and
134 physics-based fitting without very long forecast experiments in which the former are likely
135 to fail compared to the latter? Furthermore, there is a general tendency to increase the
136 complexity of models with more and more parameters. Given that in CSEP all models'
137 forecasts are fully specified with zero degrees of freedom, the number of input parameters
138 does not count into the result. But should it maybe? Again, an overfit model temporarily
139 might fit extremely well or a model might simply have been lucky to fit well in the short
140 term (Sornette et al. 2019), but in the long run the fit will likely deteriorate.

141

While this criticism may be valid for many analyses of goodness-of-fit, residual methods could perhaps help here, if they can provide useful graphics highlighting where exactly model A seems to outperform model B. Voronoi deviance residuals and super-thinned residuals (Clements et al. 2012, Bray and Schoenberg 2013, Bray et al. 2014, Gordon et al. 2015) seem potentially useful in this regard. Learning how to improve a complex model is no easy task, so even a suggestion of a rather minor improvement should perhaps be seen as a major achievement from a statistical procedure. As for the possibility of overfitting, or for model A to outperform model B in CSEP over several years but not in the long term for whatever reason, while such possibilities may be inevitable, CSEP seems ideally suited to handle these types of problems. Overfitting and lack of reproducibility are enormous problems in conventional research involving retrospective analyses of data. With CSEP and its strict insistence on truly prospective testing, these problems may still exist but they are minimized as much as possible.

2. Are ETAS models valuable?

ETAS models seem frequently to offer the best fit among the proposed models for earthquake occurrences. While so many other models have been suggested based on retrospective analyses, the fact that ETAS performs well in the sense of offering satisfactory fit to data even when used prospectively is extremely impressive, and is solid evidence of its value.

163 However, as mentioned previously, most versions of ETAS have very little value for
164 forecasting the largest events. (There are exceptions, however, that treat large earthquakes
165 fundamentally differently from smaller ones, such as Nandan et al. 2019, 2022).
166 Furthermore, what exactly is ETAS telling us that goes beyond the heuristic notion that large
167 earthquakes are followed by aftershocks and possibly even larger events? Reasenbergs & Jones
168 (1989) have decades ago quantified the chance of a larger earthquake following an already
169 large shock. Does ETAS tell us significantly more despite the much heavier computational
170 load and its complexity? ETAS provides higher resolution, but in what sense is this really
171 useful? Does it change decisions in risk mitigation? Does it allow us to call an area "safe"
172 earlier?

173

174 An important topic that has been insufficiently explored is how to quantify a model's *value*,
175 for forecasting. Much attention has been paid to the quantification of a model's goodness-of-
176 fit to data, and this is certainly a component of a model's value, since the better a model fits,
177 the more confidence one has in its forecasting ability. However, goodness-of-fit does not tell
178 the whole story. While ETAS models, for instance, may fit very well to catalogs of
179 earthquakes including aftershocks, and while forecasts can be obtained using ETAS via
180 simulations (Omi et al. 2014, Shcherbakov et al. 2019, Petrillo and Zhuang 2024) or multi-
181 element probability formulas and other techniques (Ogata 2017a, Ogata 2022, Ogata2024),
182 most versions of the ETAS model typically have little value for forecasting the largest
183 earthquakes, which happen to be the ones we care most about (in synthetic tests,

184 Helmstetter and Sornette 2003 quantify that ETAS allows one to forecast about 20% of the
185 largest events). We should explore alternative measures, such as the measure Q proposed
186 above, or variants of the Brier score or information gain restricted to the subset of
187 earthquakes of most concern, perhaps weighting earthquakes differentially based on their
188 energy release, damage potential, or other measures. In particular, it would be of great
189 interest to agree on a measure of a model's forecasting value as a function of time horizon. It
190 may be, for instance, that ETAS has excellent forecasting value for forecasting seismicity
191 several hours or days into the future, but practically no value at forecasting several months or
192 years into the future. Other models possessing the opposite qualities might exhibit worse fit
193 to data overall, yet have more forecasting value in many situations. We believe that
194 including better quantifiers of a model's value into CSEP would be a great step toward
195 discovering and raising awareness about alternative models that are potentially more useful.
196

197 The ETAS model should perhaps be seen as the null model, to which alternatives could be
198 proposed and compared. Something similar was proposed by Stark (1997). Indeed, in the
199 original paper proposing ETAS, Ogata (1988) actually used ETAS as a kind of null model,
200 rather than a model to be directly used for forecasting. He identified times of quiescence,
201 which were essentially times when the model appeared to fit poorly, as potential indicators
202 of an impending future large event (even though the quiescence hypothesis has been
203 debunked by van Stiphout et al. 2011 employing a declustering algorithm using Southern
204 California data).

205

206 There still seems to be some work to do at constructing a suitable null model, however.

207 Numerous models have been proposed, starting from the silly spatially uniform model (easy

208 to beat), via simple Poisson smoothed seismicity models, all the way to time-varying ETAS

209 models. Each model needs to be somehow calibrated and this creates a plethora of different

210 model flavors for each basic concept. There is presently not a single ETAS model, but rather

211 a host of different varieties, parameterizations, and implementations. Therefore, more

212 thought should go into how to create a suitable version of a null model, and what the

213 requirements should be for such a model.

214

215 3. Is ETAS the end?

216

217 On the one hand, looking at the collection of recent publications, the answer seems to be

218 yes. Few genuinely new model classes have recently been introduced to capture the time

219 dependence of earthquakes. Instead, more and more flavors of ETAS models have been

220 developed, mainly by attempting to fit the ETAS concept to local, regional, or sequence-

221 specific datasets, perhaps because of the 'because we can' effect: people doing what they

222 know they can achieve even if it is unlikely to lead anywhere substantial. Out of these many

223 ETAS flavors, no consensus model has been selected by the wider community. ETAS seems

224 to reproduce itself constantly without adding significant improvements to a solution of the

225 earthquake forecasting problem. Furthermore, it may be that earthquakes are a natural

226 phenomena too complex to be modeled or forecast accurately, in CSEP or any type of
227 forecasting experiment. No predictability in the pattern of earthquake magnitudes has been
228 discovered, or at least none that has been consistently reproduced. It is possible, as Yan, Dave
229 and their collaborators posited, that earthquakes may be inherently unpredictable, and
230 simply cannot be predicted (Geller et al. 1997, Kagan 1997a).

231

232 On the other hand, there may be room for optimism. Some promising recently proposed
233 versions of ETAS take into account focal depth and rupture geometry, for instance (Guo et al.
234 2015, 2018, 2021, 2024), as well as other physical components such as the magnitude-
235 dependent Omori law (Sornette and Ouillon 2005, Ouillon and Sornette 2005, Ouillon 2009,
236 Tsai et al. 2012) or a two-branched Gutenberg-Richter distribution (Saichev and Sornette
237 2005, Nandan et al. 2019), but see also Petrillo and Zhuang (2023) for the opposite opinion.
238 Also, while ETAS may fit best among existing models, it has not been shown that ETAS fits
239 better than any *possible* alternative model. Certainly any simple model forecasting the
240 precise occurrence of future seismicity is highly elusive, but this does not mean
241 improvements will not be forthcoming. Scientific progress has often been slow and laborious,
242 and seismology is no exception. Perhaps increased emphasis on *value*-based goodness-of-fit
243 statistics, as described above, may reasonably be anticipated to yield improved models that
244 optimize these more sensible and practical criteria, and CSEP has the potential to be an
245 important leader in this aim. If we increase our focus on models that improve forecasting in

246 useful ways, and reward models for improving forecasting in useful ways, such models are
247 more likely to be discovered. As they say in *Field of Dreams*, "If you build it, he will come."

248

249 4. What are the best ways forward for earthquake forecasting?

250

251 This may be divided into practical, technical approaches as well as the overall big picture,
252 and we start with the former. While there may be advantages to looking at local results in
253 some situations, testing generally should be done on a global scale to increase the power of
254 the tests so we can have some confidence in the results. Only globally can we test a
255 sufficiently large number of significantly large earthquakes that matter. Yan Kagan
256 consistently advocated this (Kagan 2003, Kagan and Jackson 1991, Kagan and Knopoff 1981,
257 1987, Kagan et al. 2012). CSEP has, after its inception, expanded from testing earthquake
258 forecasts in California to further testing regions in Japan, Italy, New Zealand, the western
259 Pacific and finally to global tests. However, the majority of tested models were developed for
260 the regional testing regions and only a few are being tested globally.

261

262 All testing regions have been defined on 0.1 degree longitude/latitude grid cells
263 (Schorlemmer and Gerstenberger 2007). While this resolution has been chosen to match the
264 location uncertainty of local events, it backfired in the global experiment in which the
265 testing area consists of 6.4+M cells and 100+M bins. However, the grid size should be scaled
266 according to how much data one has. With millions of cells but just a handful of

267 earthquakes, the test has little to no power, so the grid cells should be chosen adaptively in
268 order to maximize the power of the tests. One pathway to solve this problem is to use a
269 multi-resolution grid, e.g. the Quadtree approach taken by Ogata et al. (1996) and Asim et al.
270 (2023), or using Delaunay tessellations (Ogata et al. 2003, Ogata et al. 2019). The forecasts
271 can have high spatial resolution (small cells) in areas with many earthquakes and low
272 resolution in areas with few earthquakes (large cells). This way, meaningful global forecasts
273 can be provided with a few ten thousand of cells, matching the number of observations in
274 the Global CMT catalog.

275

276 As far as the big picture and where and whether scientific progress is likely in the future,
277 when we first got interested in earthquakes, we imagined that the historical and modern,
278 high-resolution earthquake record might contain hidden information about when the next
279 big one would occur, and that if we just looked hard enough, we could find some pattern or
280 signal and successfully predict the next major event. Now we realize that was naive, and the
281 Earth tends not to resemble a cartoon villain leaving obvious clues about future calamities.
282 Many natural phenomena are exceedingly complex, and earthquakes seem to fit this mold, so
283 there seems little hope at finding some simple pattern that predicts when the next big event
284 will occur. On the other hand, there is something rather mysterious about large earthquakes.
285 They are, after all, major ruptures of the Earth, and sudden releases of enormous amounts of
286 tectonic energy. And they must be triggered by *something* and must experience a
287 preparation phase that should leave some hints to be observed. Does it not stand to reason

288 that, as our knowledge of the Earth's structure deepens, we ought to be able to figure out and
289 possibly anticipate, or at least see some warning signs, of what is triggering these gigantic
290 outbursts? Naive as it may be, we still believe that a more precise understanding of the
291 Earth's structure should yield an improved ability to forecast major events.

292

293 This highlights another possibly important way that CSEP could be improved going forward:
294 by incorporating other types of signals apart from earthquake occurrences. If the information
295 content of earthquake catalogs does not allow for powerful tests and/or useful models, CSEP
296 should reach out to model developments that include further signals that can be observed;
297 potentially important efforts have been made in incorporating such signals recently (Zhuang
298 et al. 2005, Han et al. 2016, Kumazawa et al. 2016, Freund et al. 2021). CSEP has already
299 included models that use strain data as input, however these data have not been included as
300 an authoritative data source to ensure all models use the same strain data. Of course,
301 modelers can technically include any type of data in their models, but only if these data are
302 authoritative and provided by CSEP can comparative testing become really meaningful and
303 ensure full reproducibility of all forecasts generated by the model.

304

305 5. Can we use CSEP test results to improve models? Does CSEP provide a useful feedback
306 loop?

307

308 The L-test and similar results typically assess the performance of the entire model. However,
309 models are typically compiled of different ingredients and their interplay can be complex.
310 We usually do not know if all components work well and contribute correctly; maybe one
311 component is not calibrated well and lowering the overall performance of the model, but test
312 results rarely indicate this. Instead, each metric typically tests a different feature of a
313 forecast, not a component of the underlying model. But do we know what part of the model
314 causes the feature to perform well or not? Do we know how to translate results of a specific
315 metric into model improvement? The answer is often no!

316

317 On the other hand, it definitely seems that, if any substantial progress is ever made at
318 forecasting seismicity, it is going to be largely the result of very careful and intricate model
319 evaluation. Put another way, it seems very unlikely we will ever have substantial
320 improvement without excellent model evaluation experiments like CSEP. Without this kind
321 of rigorous look at the models, seismology would be doomed to keep repeating the mistakes
322 of the 20th century, where model after model was proposed, based on retrospective analysis,
323 only to find out later that the models did not do well prospectively (Kagan and Jackson 1991,
324 Geller et al. 1997, Kagan 1997a, Jackson and Kagan 2006).

325

326 The situation is somewhat analogous to examples from statistics in medicine, where before
327 the proper emphasis was placed on randomized controlled experiments, it was almost
328 impossible to tell whether a drug or procedure was effective, and the literature was full of

329 reports promoting procedures like the portacaval shunt, when later experiments clearly
330 showed them to be ineffective (Freedman et al. 1998 very nicely summarizes such examples).

331

332 In seismology, with CSEP already firmly in place, the field is poised to move in a positive
333 direction. Even though progress might be slow, we at least have a system in place that could
334 identify improved models once they are proposed, and we have a mechanism for more
335 efficiently sifting through and debunking poorer models. However, the power in medical
336 tests are often much higher, and the success criteria clearer, compared to earthquake
337 forecasting tests.

338 Will CSEP be as successful for earthquake forecasting as double-blind tests have been for
339 medical research? That remains to be seen.

340

341 As mentioned previously, statistical methods for model evaluation still need to be improved
342 so they can more readily lead to model improvements. Graphical methods, such as the
343 smoothed residual field (discussed by Baddeley et al. 2005 for the purely spatial case),
344 Voronoi residuals (Bray et al. 2014) and superthinned residuals (Clements et al. 2012), rather
345 than numerical summary statistics, seem to be the most promising for suggesting model
346 improvements. With superthinned residuals, points are added or removed at random
347 according to the model, so that in the end the residual points should be uniformly scattered if
348 the model fits well, and departures from uniformity indicate places where the model fit
349 poorly. With Voronoi residuals, the domain is divided into cells, one cell for each

350 earthquake, such that each cell consists of all locations closer to the corresponding
351 earthquake than to any other earthquake. The fact that residuals on such an adaptive grid
352 allow one to pinpoint around which earthquakes a model fits well or poorly can sometimes
353 lead directly to ideas for model improvement (Clements et al. 2011, Gordon et al. 2015).
354 However, these methods have their problems as well. For one thing, given a host of
355 competing models, it is cumbersome and difficult to examine a collection of plots, and so for
356 comparing many models, sometimes the simplicity of a numerical summary is desirable.
357 Second, Voronoi residuals might be good for 2-dimensional data, but it remains unclear how
358 exactly to use them for the 3-d case, or even just for 2 spatial dimensions and time.
359 Currently, one typically just ignores time and depth and carves out Voronoi cells using just
360 the epicenters of earthquakes, but this should be improved. Superthinned residuals are easier
361 to implement and do not have the dimensional problem of Voronoi residuals, but because of
362 the randomness introduced in the generation of these residuals, often it is difficult to discern
363 clear patterns or ways to improve models from the superthinned residuals alone. Perhaps we
364 still need more improvements in the realm of visual summaries of goodness-of-fit for point
365 process models.

366

367 Some of the current tests in CSEP really are not useful for comparing competing models, and
368 should probably be removed from CSEP. If you are evaluating just one model, then maybe
369 the N-test, L-test, S-test, etc. might be useful (Schorlemmer et al. 2007, Zechar et al. 2010),
370 but when comparing 2 or more models, these tests can be very misleading, since a poorly-

371 fitting but highly variable model will often have a higher p-value than a competing model
372 that actually fits better. So, since CSEP is mainly for model comparison, perhaps going
373 forward we should focus on methods that are better for this purpose. If an overall numerical
374 measure is desired, we think the log-likelihood score is probably best (Ogata 2017b, 2024), at
375 least among previously proposed measures, although the statistic Q mentioned in Section 1
376 seems potentially more informative. We should probably be content with a summary of the
377 overall fit and not be too concerned with p-values, for model comparison purposes.

378

379 6. How much information is in the system and how well is CSEP prepared to harvest it?

380

381 This is a tough question to answer. Dave Jackson believed that if strike angle estimates could
382 be improved in the future using more accurate seismometers, we could hope for some
383 substantial information to be gained in that area, since geophysical theory seems to suggest
384 such moment tensors should be extremely useful, though presently their use seems to be
385 somewhat limited. The main information we have seems to be the seismic record, which of
386 course also has errors and missing data, but this has quite thoroughly been studied especially
387 by Kagan (2003, 2004), and obviously is a wealth of information. Certainly the paleoseismic
388 record provides some useful information, and potentially geodetic information should be
389 useful to improve our models as well. While we might be somewhat pessimistic, we must
390 admire the positivity of Ilya Zaliapin, who never seemed to have any doubt that accurate
391 earthquake forecasting should be possible, and whose excellent paper with Ben-Zion

392 highlights some promising avenues to explore, such as localization of seismicity before large
393 events (Ben-Zion and Zaliapin 2020).

394

395 The community has used the seismic record since decades to forecasts earthquakes but we
396 have not moved beyond rough statistical methods that are the more successful the smaller
397 the magnitude. Having said this, it seems that the overall number of events per magnitude is
398 the key criterion for successful forecasting. A consequence of this statement would be that
399 we have to wait very very long until we better forecast large events and that the forecast
400 horizons for these events will also be very long.

401

402 CSEP seems to be ideally prepared to harvest this type of information. CSEP offers
403 researchers a unique way to use their models to generate actual prospective forecasts in real
404 time, and to evaluate those forecasts as accurately as possible. One thing CSEP really needs
405 going forward is easy public access to these forecasts; currently they seem difficult to obtain
406 for some reason. In addition, we still face the issue of the meaning of test results and what
407 they tell us about how to improve a model or which part of the model is right or wrong.

408

409 7. Can CSEP save us from bogus Artificial Intelligence (AI) forecasts?

410

411 It is to be expected that some AI researchers will attempt to forecast earthquakes by feeding
412 whatever they find into their AI algorithms. In addition, we pessimistically anticipate many

413 papers where research simply average over multiple models, possibly in a Bayesian way,
414 forming complex ensemble models that perhaps fit well retrospectively but rarely
415 prospectively, and we do not foresee this being a useful or productive enterprise because the
416 increasingly complex ensemble models become increasingly difficult to improve using
417 residual analyses. In fact, in some sense the idea behind Bayesian ensemble models runs
418 counter to the central idea behind CSEP, which is to evaluate models prospectively and
419 rigorously in order to improve them and to distinguish the ones that seem to fit well from
420 those that do not. Rather than attempting to modify or improve individual aspects of a
421 model, Bayesian model averaging seems to make problems with the models more obscure
422 rather than clearer, and models that should be discarded become instead incorporated into
423 the ensembles, like rotten fruit in a milkshake. Even worse, these kind of models do not
424 contribute to our understanding of the physical processes and how they can be modeled in a
425 useful way. They are merely statistical stunts that are likely doomed to fail sooner rather
426 than later in prospective tests. And even if not, they do not provide any insight that let us
427 improve models that we understand and that mimic the physics we have discovered. CSEP
428 protocols might be the only guardian preventing us from false declarations of success, as
429 many researchers do not fully appreciate the necessity for truly prospective tests.

430

431 But what if the CSEP tests cannot provide any conclusive results? How can the wider CSEP
432 community argue against the plethora of AI forecasts that are expected? Serious thinking

433 about the power and value of CSEP tests is needed and needs to be codified somehow in a
434 community process.

435

436 On the other hand, CSEP forecasts should be made easily publicly available to welcome
437 further analysis of the results, even if this may spur faulty analyses of the results using AI,
438 deep learning, Bayesian ensemble modeling, etc. Perhaps if CSEP eventually has so many
439 models that it becomes overwhelming, this could be a problem. But otherwise, the fact that
440 CSEP is fully prospective seems to guard against overfitting which is a main problem with AI
441 and deep learning models, and machine learning algorithms could be very useful to
442 incorporate information from various sources, as mentioned above. CSEP forecasts need to be
443 readily available online to enable easy statistical assessment because the statistical analysis of
444 the results will likely be very useful.

445

446 8. Are CSEP tests powerless?

447

448 On the one hand, we have seen that several of the tests, including the S-test, are not reliably
449 able to reject uninformative models on the global high-resolution grid with 6.5M cells and
450 many more bins because even the longest catalogs contain too few events (Asim et al. 2013,
451 Khawaja et al. 2013). In addition, it is unclear how to measure the power of a test if the true
452 distribution is unknown. It seems that CSEP is caught in the dilemma of either producing
453 powerful tests with low resolution that provide rather meaningless results or have powerless

454 tests at high resolution, again resulting in meaningless results. How can CSEP find the sweet
455 spot and what does this sweet spot tell us about the overall information content of
456 earthquake occurrences and their testing? Much could potentially be gained by finding a
457 formulation that describes the overall information content and the resulting best-possible
458 power of CSEP tests and translate this into uncertainties of the forecasts or the test results.
459 Currently, it is unclear how to use CSEP results and to what extent to trust them given the
460 open question of the power of these tests. Quantifying the power/information content could
461 also help planning for more powerful tests by estimating what amount of input data is
462 needed or how to otherwise compensate with other types of data.

463

464 On the other hand, if we include data on lower magnitude events, we actually have plenty of
465 data already and in this context power is not a serious problem. This gets at one of the key
466 philosophical issues in seismology: do the largest earthquakes obey the same fundamental
467 properties as the moderate and small events?

468

469 If not, then it might be centuries before we can forecast large earthquakes in any meaningful
470 way. And there are some reasons to be skeptical, especially since so few large events have
471 been observed, their behavior seems difficult to characterize simply, and while small events
472 can reasonably be approximated as point sources, large events can rupture hundreds of
473 kilometers, so reducing them to point sources as done in CSEP may be unrealistic.

474 Incidentally, CSEP's attempts to introduce fault-based testing have so far badly failed as no

475 earthquake-to-fault segment association has ever been developed that could reasonably run
476 without case-by-case human intervention and decision.

477

478 But if so, then if we can find models to forecast the moderate earthquakes accurately using
479 the small to moderate events, these same models should be able to forecast the largest
480 earthquakes as well using moderate events. This seems likely and promising, especially since
481 we already have so much data on the moderate and smaller events. Yan Kagan and Dave
482 Jackson were critical of those who kept constantly claiming that we had insufficient data to
483 reject the characteristic earthquake hypothesis, at least for the largest events (Kagan 1997b,
484 Jackson and Kagan 2006). We should not make this same mistake going forward. We have
485 plenty of data. We just need to be scientific about how we use it.

486

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493

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495

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