

1 **THE EFFECTS OF PRESSURE ON pH OF TRIS BUFFER IN SYNTHETIC SEAWATER**

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Abstract

Equimolar Tris (2-amino-2-hydroxymethyl-propane-1,3-diol) buffer prepared in artificial seawater media is a widely accepted pH standard for oceanographic pH measurements, though its change in pH over pressure is largely unknown. The change in volume (ΔV) of dissociation reactions can be used to estimate the effects of pressure on the dissociation constant of weak acid and bases. The ΔV of Tris in seawater media of salinity 35 (ΔV_{Tris}^*) was determined between 10 and 30 °C using potentiometry. The potentiometric cell consisted of a modified high pressure tolerant Ion Sensitive Field Effect Transistor pH sensor and a Chloride-Ion Selective Electrode directly exposed to solution. The effects of pressure on the potentiometric cell were quantified in aqueous HCl solution prior to measurements in Tris buffer. The experimentally determined ΔV_{Tris}^* were fitted to the equation $\Delta V_{\text{Tris}}^* = 4.528 + 0.04912t$ where t is temperature in Celsius; the resultant fit agreed to experimental data within uncertainty of the measurements, which was estimated to be $0.9 \text{ cm}^3 \text{ mol}^{-1}$. Using the results presented here, change in pH of Tris buffer due to pressure can be constrained to better than 0.003 at 200 bar, and can be expressed as:

$$\Delta \text{pH}_{\text{Tris}} = \frac{-(4.528 + 0.04912t)P}{\ln(10)RT}$$

where T is temperature in Kelvin, R is the universal gas constant ($83.145 \text{ cm}^3 \text{ bar K}^{-1} \text{ mol}^{-1}$), and P is gauge pressure in bar. On average, pH of Tris buffer changes by approximately -0.02 at 200 bar.

1 Introduction

Open ocean pH is declining at rates between -0.001 to -0.002 pH yr⁻¹ as atmospheric CO₂ increases and the surface ocean equilibrates with larger partial pressures of CO₂ [Dore *et al.*, 2009; Byrne *et al.*, 2010; Bates *et al.*, 2014]. Decreases in pH may have large impacts on ocean ecosystems, particularly on organisms that make calcium carbonate skeletons and shells [Doney *et al.*, 2009; Barton *et al.*, 2012; Bednaršek *et al.*, 2012]. It is expected that rates of pH change will not be uniform [Feely *et al.*, 2008; Wootton *et al.*, 2008; Takeshita *et al.*, 2015], and different regions will experience varying degrees of impact. The rate of decrease has been observed directly at a few long-term time series stations [Bates *et al.*, 2014] using highly reproducible spectrophotometric pH measurements [Clayton and Byrne, 1993; Liu *et al.*, 2011]. The expense and expertise required to sustain such observations prevent them from being scaled to large areas of the ocean. Alternatively, a monitoring system that utilizes autonomous chemical sensors may alleviate such problems and greatly improve our understanding of the spatial and temporal variability of the rate of pH decline in the ocean [Johnson *et al.*, 2007]. However, in order to establish such a chemical sensor network, the development and implementation of stringent calibration protocols are necessary.

The assignment of proper pH values to a suitable buffer solution is essential in obtaining accurate pH measurements [Bates, 1973]. Equimolar Tris (2-amino-2-hydroxymethyl-propane-1,3-diol) buffer prepared in artificial seawater (referred to as just Tris buffer hereafter) has been widely accepted as the primary pH standard for oceanographic pH measurements [Dickson, 1993; DelValls and Dickson, 1998], and has been used to calibrate potentiometric pH measurements [Millero *et al.*, 1993; Martz *et al.*, 2010], and to characterize indicator dyes used in spectrophotometric pH measurements [Clayton and Byrne, 1993; Liu *et al.*, 2011]. Although the

temperature and salinity dependence of the dissociation constant of Tris has been quantified at 1.01 bar ($K_{\text{Tris}}^{1.01}$) [DelValls and Dickson, 1998], its pressure dependence has yet to be determined in seawater media. As the number of in situ pH measurements at high pressures is expected to increase rapidly in the next decade due to improvements in robust pH sensor technology [Johnson *et al.*, 2016] and development of the deep-sea Free Ocean Carbon Enrichment system [Barry *et al.*, 2014], characterizing Tris buffer in seawater under high pressures will meet a crucial need in sensor calibration and traceability.

The pH of a buffer solution can be quantified by the Henderson-Hasselbach equation:

$$\text{pH} = \text{p}K_a + \log_{10} \left(\frac{[A^-]}{[HA]} \right) \quad (1)$$

In equimolar buffer solutions, where the deprotonated and protonated forms of the buffer are at equal concentrations, pH is equivalent to the $\text{p}K_a$. Therefore the change in $\text{p}K_a$ is equivalent to the change in pH of the solution. The effect of pressure on K_a can be expressed as

$$RT \left(\frac{\partial \ln K_a}{\partial P} \right) = -\Delta V \quad (2)$$

where ΔV [$\text{cm}^3 \text{mol}^{-1}$] is the difference in partial molal volume of the acid dissociation reaction, R is the universal gas constant ($83.145 \text{ cm}^3 \text{ bar K}^{-1} \text{mol}^{-1}$), T is temperature in Kelvin, and P is pressure in bar [Byrne and Laurie, 1999; Millero, 2001]. This can be rearranged to give the relationship between the dissociation constant at in situ gauge pressure (K_a^P) and at atmospheric pressure (K_a^0):

$$\log \left(\frac{K_a^P}{K_a^0} \right) = \frac{-\Delta V P}{\ln(10) RT} \quad (3)$$

Note that this equation does not include the partial molal compressibility term, as this only becomes significant at higher pressures than we investigated in this study [Millero, 2001]. P refers

to gauge pressure hereafter. The ΔV_{Tris} has been measured at infinite dilution and in low ionic strength solutions ($\leq 0.1 \text{ mol dm}^{-3} \text{ NaCl}$) using dilatometry [Katz and Miller, 1971; Kitamura and Itoh, 1987] and spectrophotometry [Neuman et al., 1973]. More recently, the ΔV_{Tris} in $0.725 \text{ mol kg}^{-1} \text{ NaCl}$ was reported between 5 and 25 °C by measuring density and sound speed, and using Pitzer equations to interpret the results [Rodriguez et al., 2015]. These results can be extended to seawater media, as agreement of ΔV in artificial seawater media (salinity = 35) and $0.725 \text{ mol kg}^{-1} \text{ NaCl}$ solutions have been shown for various weak acid-base species [Millero, 2001]. However, since ΔV is dependent on ionic strength and solution composition [Byrne and Laurie, 1999], ΔV_{Tris} should be quantified in seawater media to validate the extension of ΔV_{Tris} measured in NaCl solution to seawater.

Here, we report ΔV_{Tris}^* at 10 to 30 °C, where the * symbol refers to seawater media of salinity 35. A potentiometric cell consisting of an Ion Sensitive Field Effect Transistor (ISFET) pH sensor and a Chloride-Ion Selective Electrode (Cl-ISE) was utilized to quantify the effects of pressure on K_{Tris} up to 200 bar. Excellent agreement with ΔV_{Tris} in $0.725 \text{ mol kg}^{-1} \text{ NaCl}$ solution was observed [Rodriguez et al., 2015], validating the extension of the reported ΔV_{Tris} in NaCl solution to seawater media. Sources of uncertainty for the reported values are explored.

2 Background and Theory

The pressure dependence of K_{Tris} can be calculated by monitoring the pH of equimolar Tris buffer solutions over a range of pressures. The left hand side of equation (1) is equivalent to the difference in solution pH at 0 and experimental P ($\text{pH}^0 - \text{pH}^P$) for an equimolar solution, given that the concentration of H^+ and OH^- is negligible compared to the buffer concentration:

$$(\text{pH}^0 - \text{pH}^P) = \frac{-\Delta V_i P}{\ln(10)RT} \quad (4)$$

Therefore ΔV_{Tris}^* can be quantified from a linear regression between $RT\ln(\text{pH}^0 - \text{pH}^P)$ versus gauge pressure. We ignore the effects of partial molal compressibility, because this effect becomes only significant at pressures above 600 bar [Rodriguez *et al.*, 2015]. Furthermore, we make the assumption that the presence of Tris buffer in the artificial seawater media (0.04 mol kg^{-1}) has negligible influence on ΔV_{Tris}^* .

The pH of the solution was quantified using the cell ISFET|Tris Buffer|Cl-ISE. The operating principle of ISFET pH sensors is based on potentiometry, though significantly differs from conventional potentiometric electrodes such as the glass electrode [Bergveld, 2003]. The ISFET is an active electronic circuit, in which the potential between the reference electrode and the source of the transistor follows a Nernstian response when the gate insulator material has sufficient buffer capacity [Bergveld, 2003; Takeshita *et al.*, 2014]. The ISFET utilized in the Honeywell Durafet line of pH sensors (referred to as ISFET hereafter) with a Cl-ISE directly exposed to solution as the reference electrode have been shown to produce extremely stable measurements in seawater for months to years at 1 atm [Martz *et al.*, 2010]. The Cl-ISE is especially well suited for high pressure applications, as it seems to exhibit minimal pressure hysteresis [Shitashima *et al.*, 2002]. These traits make the ISFET and Cl-ISE a good candidate for characterizing Tris buffer over periods of weeks to months.

Detailed description of the calibration protocol and calculation of pH at experimental temperature and pressure has been presented previously [Johnson *et al.*, 2016], but will be summarized here. The electromotive force of the cell ISFET|Tris Buffer|Cl-ISE (E) at experimental temperature and pressure range between 0 and 200 bar can be explained by the Nernst relation:

$$E = E_{\text{ref}}(T, P) - k\log(a_{\text{H}}a_{\text{Cl}}) \quad (5)$$

$$= E_{\text{ref}}(T, P) - k \left[\log(\gamma_{\pm\text{HCl}}^2 m_{\text{H}} m_{\text{Cl}}) + \frac{\bar{V}_{\text{HCl}} P}{\ln(10)RT} \right]$$

121 where $E_{\text{ref}}(T, P)$ is the reference potential at experimental T and P , k is the Nernst slope, a and m
 122 are the activity and molality of the respective ions, $\gamma_{\pm\text{HCl}}$ is the mean activity coefficient of HCl
 123 ($\gamma_{\pm\text{HCl}} = (\gamma_{\text{H}}\gamma_{\text{Cl}})^{1/2}$), and $\bar{V}_{\text{HCl}}$ is the partial molal volume of HCl in seawater; the Nernst slope is
 124 defined as $k = RT\ln(10)/F$, where R is the universal gas constant ($8.3145 \text{ J K}^{-1} \text{ mol}^{-1}$) and F is the
 125 Faraday constant (96485 C mol^{-1}). Since partial molal volume of H^+ is defined as zero, $\bar{V}_{\text{HCl}}$ is
 126 equivalent to $\bar{V}_{\text{Cl}}$. The pressure effects on the activity coefficients are encapsulated in the difference
 127 between the molal volume of HCl in seawater and water. Rearranging equation (5) for pH, defined
 128 as the $-\log(m_{\text{H}})$, gives

$$\text{pH} = \frac{(E - E_{\text{ref}}(T, P)) + k \left[\log(\gamma_{\pm\text{HCl}}^2 m_{\text{Cl}}) + \frac{\bar{V}_{\text{Cl}} P}{\ln(10)RT} \right]}{k} \quad (6)$$

129 In order to calculate pH at experimental conditions, $E_{\text{ref}}(T, P)$ must be quantified.

130 $E_{\text{ref}}(T, P)$ can be expressed as the sum of the reference potential at a given temperature at
 131 0 bar ($E_{\text{ref}}(T, 0)$), and the pressure coefficient of the sensor ($f(P)$):

$$E_{\text{ref}}(T, P) = E_{\text{ref}}(T, 0) + f(P) \quad (7)$$

132 The $E_{\text{ref}}(T, 0)$ can be quantified by measuring E in a solution with known activity of HCl at
 133 atmospheric pressure (e.g. aqueous HCl solution or Tris buffer) [Martz *et al.*, 2010]. The pressure
 134 coefficient must be determined for every sensor individually [Johnson *et al.*, 2016]. The $f(P)$ can
 135 be quantified by measuring E in a solution where the activity of HCl and $\bar{V}_{\text{Cl}}$ are known over the
 136 desired range of T and P . Combining equations (5) and (7) gives

$$f(P) = E - E_{\text{ref}}(T, 0) + k \left[\log(\gamma_{\pm\text{HCl}}^2 m_{\text{H}} m_{\text{Cl}}) - \frac{-\bar{V}_{\text{Cl}} P}{\ln(10)RT} \right] \quad (8)$$

Aqueous HCl is an ideal solution for this approach since HCl is fully dissociated, thus m_{H} and m_{Cl} are independent of T and P , $\gamma_{\pm\text{HCl}}$ is well characterized over a wide range of temperature [Harned and Ehlers, 1932; Harned and Owen, 1963], and the partial molal quantities of Cl^- are known [Millero, 1982]. Once $f(P)$ is quantified for a given ISFET Cl-ISE pair, then it can be used to measure pH of Tris buffer at experimental T and P (equation (6)) to quantify ΔV_{Tris}^* (equation (4)). Note that these partial molal quantities are solution dependent, and the $\bar{V}_{\text{Cl}}$ in seawater media ($\bar{V}_{\text{Cl}}^*$) must be used for Tris buffer [Millero, 1982].

3 Materials and Methods

Temperature and pressure cycles were carried out in a custom system capable of reproducing a T - P range of 0 to 40 °C, and 0 to 200 bar. The ISFET and Cl-ISE was placed into a pressure vessel consisting of a titanium housing with an inert PEEK insert as the wetted material [Johnson *et al.*, 2016]. The T and P of the housing is controlled by a temperature bath (Thermo Scientific, RTE-7) and an ISCO 260D Syringe pump, respectively. The pressure chamber was placed in an air bath (controlled by the same water bath) in order to improve temperature stability. Temperature was measured using a YSI 4600 thermometer with an Omega thermistor (P/N ON-410-PP), and pressure was measured using an Omega analog pressure sensor (P/N: PX409-5.0KG5V-XL). The thermistor was calibrated between 0 and 40 °C using a Hart 5611T thermistor (estimated accuracy of 0.01 °C). The factory calibration was used for the pressure sensor (accuracy of 0.03%).

Modifications were made for the supporting structure of the ISFET to withstand high pressures; details on the modifications are documented elsewhere [Johnson *et al.*, 2016]. Cl-ISE

pellets were purchased from Van London, and packaged in-house. The ISFET and Cl-ISE was operated in either a $0.0100 \pm 0.0001 \text{ mol kg}^{-1}$ HCl or Tris buffer solution; HCl was prepared gravimetrically using a 1 mol dm^{-3} HCl standard solution (Fisher, SA48-1), and a certified Tris buffer distributed by the Dickson lab (batch 8) was used. A 24-bit Analog to Digital Converter (NI-9219) was used to measure E every 45 seconds. All electronics except for the water bath were powered through an isolation transformer to eliminate potential ground loop issues.

The experimental apparatus was controlled through a LabView interface, and pressure was cycled between 0 – 200 bar at a rate of approximately $0.83 \text{ bar min}^{-1}$, taking 8 hours for a full compression/decompression cycle. This resulted in > 600 discrete measurements per pressure cycle. Pressure cycles were performed at 7 different temperatures at $5 \text{ }^{\circ}\text{C}$ intervals between 0 and $30 \text{ }^{\circ}\text{C}$ ($n = 6$ for HCl, $n = 9$ for Tris buffer at each temperature). Temperature was stable to better than $0.2 \text{ }^{\circ}\text{C}$ during the pressure cycles; changes were due to adiabatic heating and cooling of the solution. The first pressure cycle at a new temperature was discarded due to incomplete thermal equilibration of the housing. Only the data collected during increasing pressure were used in the analysis due to the small pressure hysteresis in $f(P)$. This was repeated twice, to serve as replicate estimates for $f(P)$. The ΔV_{Tris}^* was quantified by averaging results from 4 to 8 pressure cycles at each temperature, and then performing a linear regression between $-RT\ln(10)(\text{pH}^0 - \text{pH}^P)$ and P (equation (4)). Data for $P > 100 \text{ bar}$ was used at $10 \text{ }^{\circ}\text{C}$ due to increased noise at lower pressures. Due to a sensor malfunction during the final set of measurements in Tris, results are not reported for (0, 5, and $15 \text{ }^{\circ}\text{C}$). Experimental data is provided in the supplementary materials.

4 Results and Discussion

A repeatable $f(P)$ was observed at all temperatures, and decreased by $\sim 6 \text{ mV}$ over 200 bar (Figure 1). There was a small, but noticeable effect of temperature on $f(P)$, and generally was lower

at lower temperatures. However, this effect was very repeatable at all temperatures. Agreement between replicate assessments of $f(P)$ at all temperatures performed roughly a month apart was better than 180 μV (maximum difference), and typically was better than 100 μV . This demonstrates the high reproducibility and long term stability of the established pressure coefficient for the sensor. We assumed that $f(P)$ remained consistent between HCl and seawater media [Johnson *et al.*, 2016].

The ΔV_{Tris}^* was calculated as the slope of a regression for $-RT\ln(10)(\text{pH}^0 - \text{pH}^P)$ versus P (equation (4)) between 30 and 200 bar (Figure 2) over the range 10 - 30 °C. Sensor noise often increased with a full depressurization of the sensors, theorized to be due to nucleation of an air bubble in the open gate of the FET. Therefore, regression was limited to data above 30 bar. ΔV_{Tris}^* increased with temperature, ranging from 5.1 $\text{cm}^3 \text{mol}^{-1}$ at 10 °C to 6.0 $\text{cm}^3 \text{mol}^{-1}$ at 30 °C (Table 1). The ΔV_{Tris}^* was fitted to functions of temperature (t , °C):

$$\Delta V_{\text{Tris}}^* = 4.528 + 0.04912t \quad (9)$$

Combining this with equation (4) allows us to calculate the change in pH of Tris buffer due to pressure relative to atmospheric conditions over a range of temperatures:

$$\Delta \text{pH}_{\text{Tris}} = \frac{-(4.528 + 0.04912t)P}{\ln(10)RT} \quad (10)$$

Note that t and T refers to temperature in Celsius and Kelvin, respectively. The effect of temperature on $\Delta \text{pH}_{\text{Tris}}$ is relatively small, ranging from -0.017 to -0.021 between 0 and 30 °C, respectively at 200 bar. This translates to a $\Delta \text{pH}_{\text{Tris}}/\Delta P$ of approximately -0.0001 bar^{-1} .

The ΔV_{Tris}^* determined in this study were in excellent agreement with previously reported ΔV_{Tris}^* in 0.725 mol kg^{-1} NaCl solution [Rodriguez *et al.*, 2015]. At 25 °C, ΔV_{Tris}^* was 5.8 $\text{cm}^3 \text{mol}^{-1}$ compared to 5.6 $\text{cm}^3 \text{mol}^{-1}$ in NaCl solution [Rodriguez *et al.*, 2015]. They were in agreement to

201 better than $0.2 \text{ cm}^3 \text{ mol}^{-1}$ between 10 and 30 °C, leading to a difference in $\Delta\text{pH}_{\text{Tris}}$ of < 0.001 at
202 200 bar. Quantifying ΔV_{Tris}^* at temperatures $< 5 \text{ °C}$ would be most useful for validating pH
203 measuring systems in situ, as the vast majority of the deep sea falls into this temperature range.
204 The agreement between the two studies becomes slightly worse at 0 °C ($0.35 \text{ cm}^3 \text{ mol}^{-1}$, calculated
205 by extending the linear regression from our dataset to 0 °C), which leads to a difference in $\Delta\text{pH}_{\text{Tris}}$
206 of 0.0013, approaching precision of many autonomous pH sensing systems [Martz *et al.*, 2003;
207 Bresnahan *et al.*, 2014; Johnson *et al.*, 2016]. However, at this temperature, the pH of equimolar
208 Tris buffer is greater than 8.9, making it non-ideal for validating in situ pH measurements,
209 especially for those using m-cresol purple indicator dye. Tris buffer pH can be lowered by
210 adjusting the ratio of the acidic and basic form of the buffer. Standardization of Tris buffer in
211 varying mole ratios in artificial seawater for oceanographic pH measurements are currently
212 underway [Pratt, 2014]. Use of such buffers are recommended for validating in situ pH
213 measurements in the deep ocean.

214 The uncertainty in the determined ΔV_{Tris}^* is directly associated to our ability to quantify the
215 relative change in pH between atmospheric and experimental pressure (equation (4)). The sources
216 of error and their contribution to the uncertainty at 200 bar based on standard propagation of
217 uncertainty calculations [EURACHEM/CITAC, 2000], are listed in Table 2. These results indicate
218 that the change in pH of Tris due to change in pressure of 200 bar can be constrained to 0.0031,
219 which translates to an uncertainty in ΔV_{Tris}^* of $0.9 \text{ cm}^3 \text{ mol}^{-1}$; the effect of temperature on this value
220 is small, and changes by $< 0.1 \text{ cm}^3 \text{ mol}^{-1}$ between 10 and 30 °C. This uncertainty likely represents
221 a lower bound estimate, as sources of error such as the accuracy of the analog to digital converter
222 were not included. However, it is interesting to note that the uncertainty in the regression using the

combined dataset from this study and from Rodriguez et al. 2015 is significantly smaller (RMSE = 0.1 cm³ mol⁻¹) than that calculated from Table 2.

5 Conclusion

The ΔV_{Tris}^* was quantified between 10 and 30 °C by measuring the change in pH of the buffer solution between 1.01 and 200 bar. pH of the solution was measured using a modified high pressure tolerant ISFET pH sensor and a Cl-ISE as a reference electrode. A custom system was utilized to control experimental temperature and pressure. The pressure coefficient $f(P)$ of the cell ISFET|Tris Buffer|Cl-ISE was quantified in aqueous HCl solution prior to measurements in a certified Tris buffer solution. Our measured values were fitted to the equation: $\Delta V_{\text{Tris}}^* = 4.528 + 4.912E-02t$. We estimate our uncertainty to be 0.9 cm³ mol⁻¹. Our measurements for ΔV_{Tris}^* were in excellent agreement with those measured in 0.725 mol kg⁻¹ NaCl. This agreement is very encouraging, and demonstrates the validity of extending partial molal volume data in NaCl solutions to seawater media. Potential uses of this information include benchtop and in situ validation of pH measuring systems at elevated pressures in a widely accepted seawater pH standard (Tris in artificial seawater).

6 Acknowledgements

This work was funded by the National Oceanographic Partnership Program award N00014-10-1-0206 and by the David and Lucile Packard Foundation.

242 Table 1: ΔV_{Tris}^* [cm³ mol⁻¹] between 10 and 30 °C

Temperature	ΔV_{Tris}^*
10	5.1
20	5.4
25	5.8
30	6.0

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244 Table 2: Summary of Propagation of Error for Calculated pH at Experimental Conditions

Source of Error	Error	Uncertainty in pH ^a
$f(P)$	180 μV ^b	0.0030
$\bar{V}^*_{Cl}$	0.1 ^c $cm^3 mol^{-1}$	0.0005
$\bar{V}^o_{Cl}$	0.1 ^d $cm^3 mol^{-1}$	0.0007
Total		0.0031

245 ^a pH calculated at 200 bar and 25 °C

246 ^b Error in $f(P)$ was calculated using the worst case scenario

247 ^c Error from [*Poisson and Chanu*, 1976]

248 ^d Error from [*Millero*, 1982], at 25 °C.

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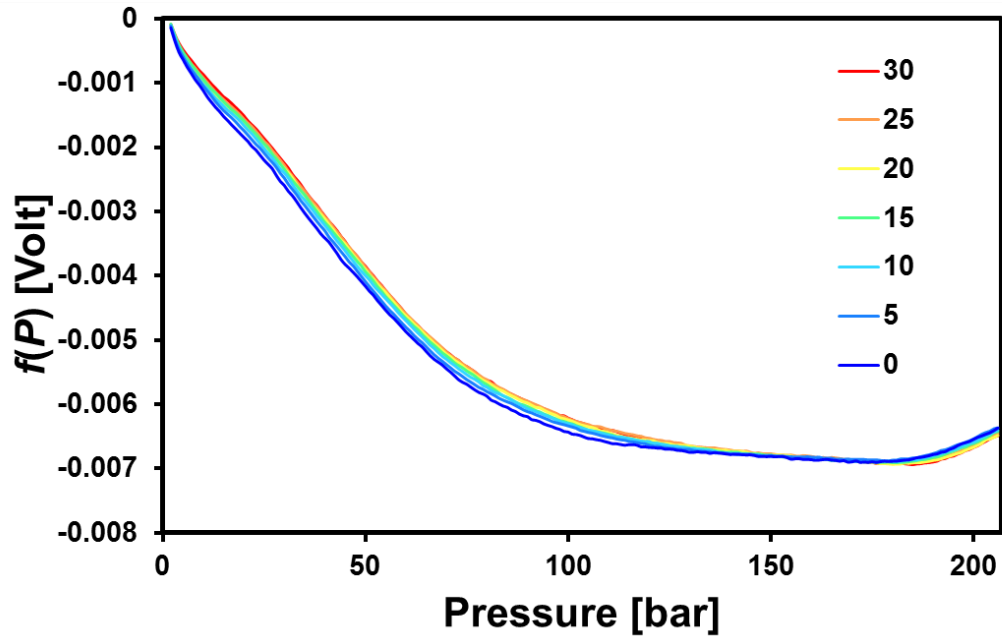
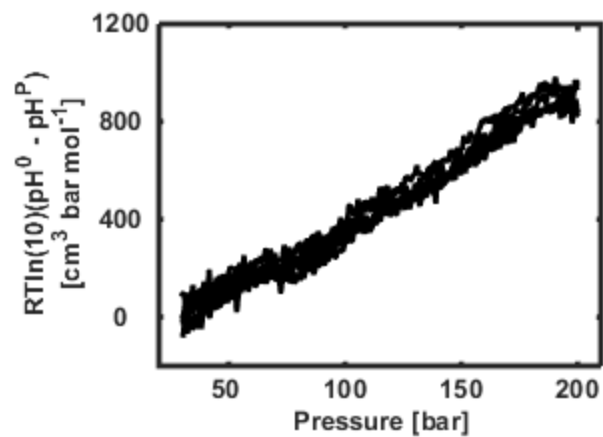
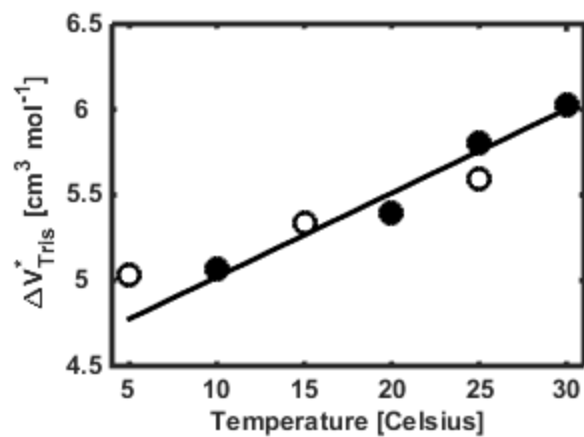


Figure 1: The pressure coefficient, $f(P)$, of an ISFET Cl-ISE pair between 0 and 30 °C.



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254 Figure 2: Typical results of a set of pressure cycles at a constant temperature used to determine
 255 ΔV_{Tris}^* (data shown for 8 pressure cycles at 30 °C).
 256



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258 Figure 3: ΔV_{Tris}^* versus temperature from this study (closed circle) and *Rodriguez et al. 2015*
 259 (open circle). The black line represents the linear regression for the data in this study.

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