

Global energetics and local physics as drivers of past, present and future monsoons

Michela Biasutti^{1*}, Aiko Voigt^{1,2}, William R. Boos³, Pascale Braconnot⁴, Julia C. Hargreaves^{1,5}, Sandy P. Harrison^{1,6}, Sarah M. Kang⁷, Brian E. Mapes^{1,8}, Jacob Scheff⁹, Courtney Schumacher¹⁰, Adam H. Sobel^{1,11} and Shang-Ping Xie¹²

Global constraints on momentum and energy govern the variability of the rainfall belt in the intertropical convergence zone and the structure of the zonal mean tropical circulation. The continental-scale monsoon systems are also facets of a momentum- and energy-constrained global circulation, but their modern and palaeo variability deviates substantially from that of the intertropical convergence zone. The mechanisms underlying deviations from expectations based on the longitudinal mean budgets are neither fully understood nor simulated accurately. We argue that a framework grounded in global constraints on energy and momentum yet encompassing the complexities of monsoon dynamics is needed to identify the causes of the mismatch between theory, models and observations, and ultimately to improve regional climate projections. In a first step towards this goal, disparate regional processes must be distilled into gross measures of energy flow in and out of continents and between the surface and the tropopause, so that monsoon dynamics may be coherently diagnosed across modern and palaeo observations and across idealized and comprehensive simulations. Accounting for zonal asymmetries in the circulation, land/ocean differences in surface fluxes, and the character of convective systems, such a monsoon framework would integrate our understanding at all relevant scales: from the fine details of how moisture and energy are lifted in the updrafts of thunderclouds, up to the global circulations.

Most tropical precipitation, whether steady rain or intense showers, falls from cloud clusters where individual, small-scale updrafts are organized over a few hundred kilometres in discrete weather systems. These, in turn, are orchestrated by planetary-scale circulation features: monsoons and the intertropical convergence zone (ITCZ; Fig. 1a,b). The clustering of individual clouds is thus the visible signature of an otherwise invisible global stirring. The notion that a link exists between smaller and larger scales underpins our understanding of tropical rain-belt dynamics and is the basis of their representation in global climate models. Our understanding, however, remains incomplete, as evidenced by our inability to achieve reliable predictions of how the ITCZ and monsoons respond to external forcings.

Without the benefit of outdoor controlled experiments, or the possibility of validating predictions of long-term future changes, the reliability of our prediction tools must be tested against past records. Confidence in projected climate responses increases when dynamic theories built on contemporary observations also explain past conditions and when simulations skillfully reproduce the palaeo record.

Good tests of our theories and models are the Holocene waxing and waning of monsoons in response to orbitally driven changes in incoming solar radiation^{1,2} (Fig. 2). During the early-to-mid Holocene (11,000–5,000 years before present), the Northern Hemisphere received more insolation during its summer than today,

while the Southern Hemisphere received less during its summer. In the modern climate, monsoons export energy from regions where the sun delivers most—so we would expect mid-Holocene monsoons to be stronger in the Northern Hemisphere and weaker in the Southern Hemisphere, compared to today. Palaeoenvironmental data from that period indeed indicate increased rainfall in North and Central America^{3,4}, a stronger Indian monsoon and increased inland penetration of the monsoon into China^{5,6}, and a spectacular rainfall expansion in northern Africa^{7,8}, with abundant lake, pollen and archaeological evidence documenting wetting and vegetation increase over much of the Sahara.

While qualitatively consistent with our expectations, the major expansion of the Northern Hemisphere monsoons is widely underestimated in climate model simulations^{1,2}, in particular over the Sahara⁹ (Fig. 2b). This bias can be reduced by modelling Earth-system feedbacks¹⁰ or by imposing the observed changes in vegetation, surface water storage, wetlands, soils and mineral dust as boundary conditions^{11–13}. Yet, simulations that organically produce both the rainfall distribution and the vegetation types that are consistent with the records still elude us². Moist atmospheric dynamics and its coupling with other aspects of the Earth system, including vegetation cover and soil properties, remain prime suspects for the failures. Moreover, the sign of observed Southern Hemisphere changes is not fully consistent with the predictions of reduced

¹Lamont-Doherty Earth Observatory of Columbia University, Palisades, NY, USA. ²Institute of Meteorology and Climate Research - Department of Troposphere Research, Karlsruhe Institute of Technology, Karlsruhe, Germany. ³Department of Earth and Planetary Science, University of California, Berkeley, CA, USA. ⁴Laboratoire des Sciences du Climat et de l'Environnement, Unité mixte CEA-CNRS-UVSQ, Gif-sur-Yvette Cedex, France. ⁵Blue Skies Research, The Old Chapel, Albert Hill, Settle, UK. ⁶School of Archaeology, Geography and Environmental Science, University of Reading, Whiteknights, Reading, UK. ⁷School of Urban and Environmental Engineering, Ulsan National Institute of Science and Technology, Ulsan, South Korea. ⁸University of Miami, Miami, FL, USA. ⁹Department of Geography and Earth Sciences, University of North Carolina, Charlotte, NC, USA. ¹⁰Department of Atmospheric Sciences, Texas A&M University, College Station, TX, USA. ¹¹Department of Earth and Environmental Sciences and Department of Applied Physics and Applied Mathematics, Columbia University, New York, NY, USA. ¹²Scripps Institution of Oceanography, University of California San Diego, La Jolla, CA, USA. *e-mail: biasutti@ldeo.columbia.edu

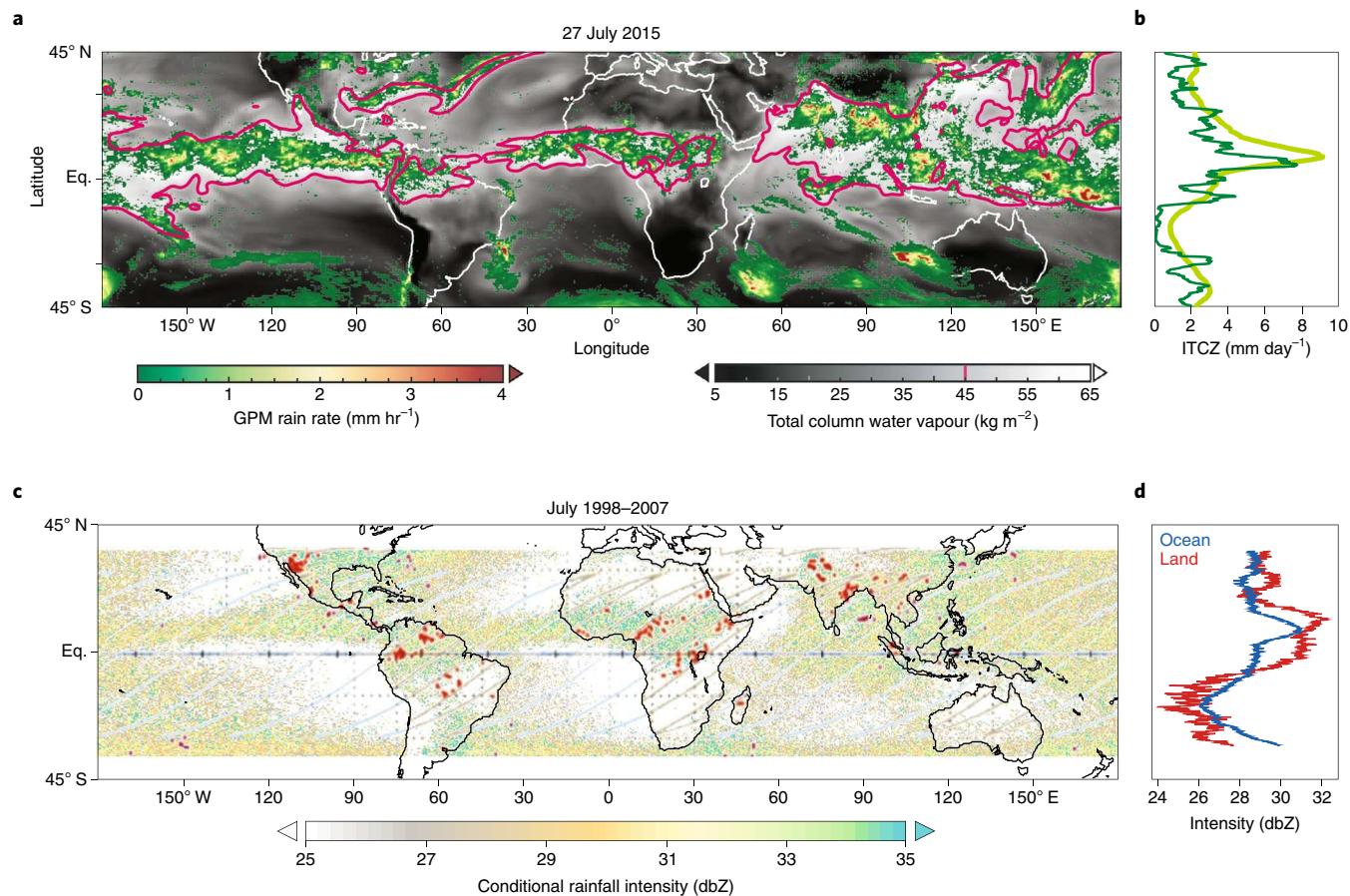


Fig. 1 | Distinct tropical convection systems are organized in a planetary-scale rain belt. **a**, Rainfall (coloured) on 27 July 2015, and the high atmospheric moisture enveloping it (indicated by the 45 mm contour of column-integrated water vapour, the full field is in grey). GPM, global precipitation measurement. **b**, Zonal mean rainfall for the same day (dark green) and climatological values for the same period (light green). **c**, July climatological mean intensity of instantaneous near-surface rainfall (from the Tropical Rainfall Measuring Mission (TRMM) precipitation radar, units of reflectivity (dbZ)) and occurrence of lightning (red dots) on 27 July 2014 (ascending passes of the lightning imaging sensor on TRMM). **d**, Zonal mean rainfall intensity for ocean (blue) and land (red) regions. See Methods for further details.

rainfall from both arguments based on the hemispheric summer-time insolation forcing and complex model simulations. Although many more quantitative estimates of rainfall anomalies from the Southern Hemisphere are needed to paint a full picture, many palaeoenvironmental records from southern Africa imply increased precipitation¹⁴, and palaeoenvironmental records from South America and Australia seem to show (their interpretation being somewhat controversial) mixed wetting and drying signals^{15,16}. Thus, neither complex models nor theoretical intuition are sufficient to explain past monsoon records.

In a traditional dry paradigm, monsoon circulations are akin to continental-scale land–sea breezes driven by surface temperature contrast, their strength increasing with the contrast. But this is too simplistic. Monsoon lands are hottest before the start of the monsoon, but the circulation is strongest in late summer (when increased rain and cloudiness have cooled the land and reduced the contrast). In future climate projections, land–sea temperature contrasts universally strengthen but monsoon circulations generally weaken¹⁷, as does, in some instances, regional early season rainfall^{18,19}.

A better paradigm views monsoons not as giant heat lows for which rainfall is a side effect, or as circulations ‘driven by’ the latent heating of rainfall, but as moist energetically direct circulations tightly coupled to precipitating convection; a facet of the general overturning of the tropical atmosphere inextricably linked to the Hadley circulation and the zonal mean ITCZ²⁰.

At seasonal timescales, convection acts to release any column instability and to bring the free-tropospheric temperature in line with the moist static energy (MSE) of the boundary layer below it²¹. Horizontal atmospheric motions homogenize the free-tropospheric tropical temperatures to one vertical profile that reflects the conditions of the major convective centres. In this view, peak rainfall should coincide with peak low-level MSE; cooler and dryer surfaces can support only shallower or suppressed convection. Land and ocean observations confirm this theoretical prediction²². Meanwhile, the upper-level divergence over these convective regions (maxima of low-level MSE) implies that the circulations associated with peak rainfall export total energy away from the centres of deep convection. In other words, monsoons and the ITCZ are components of a planetary energetically direct circulation that links land and ocean rainfall (similar to a ‘global monsoon’²³) and shows coherent variability from seasonal^{24,25} to geological timescales^{20,26,27}. Such an encompassing view is an important theoretical advance. Its formulation for the zonal mean tropical circulation—known in the literature as the energy-budget framework—underpins our understanding of why the ITCZ moves meridionally in response to forcings that originate well-outside of the tropics^{20,28,29} (see section ‘The explanatory power of energetic constraints’).

Further progress in understanding and simulating regional rainfall anomalies requires us to extend the energy-budget framework

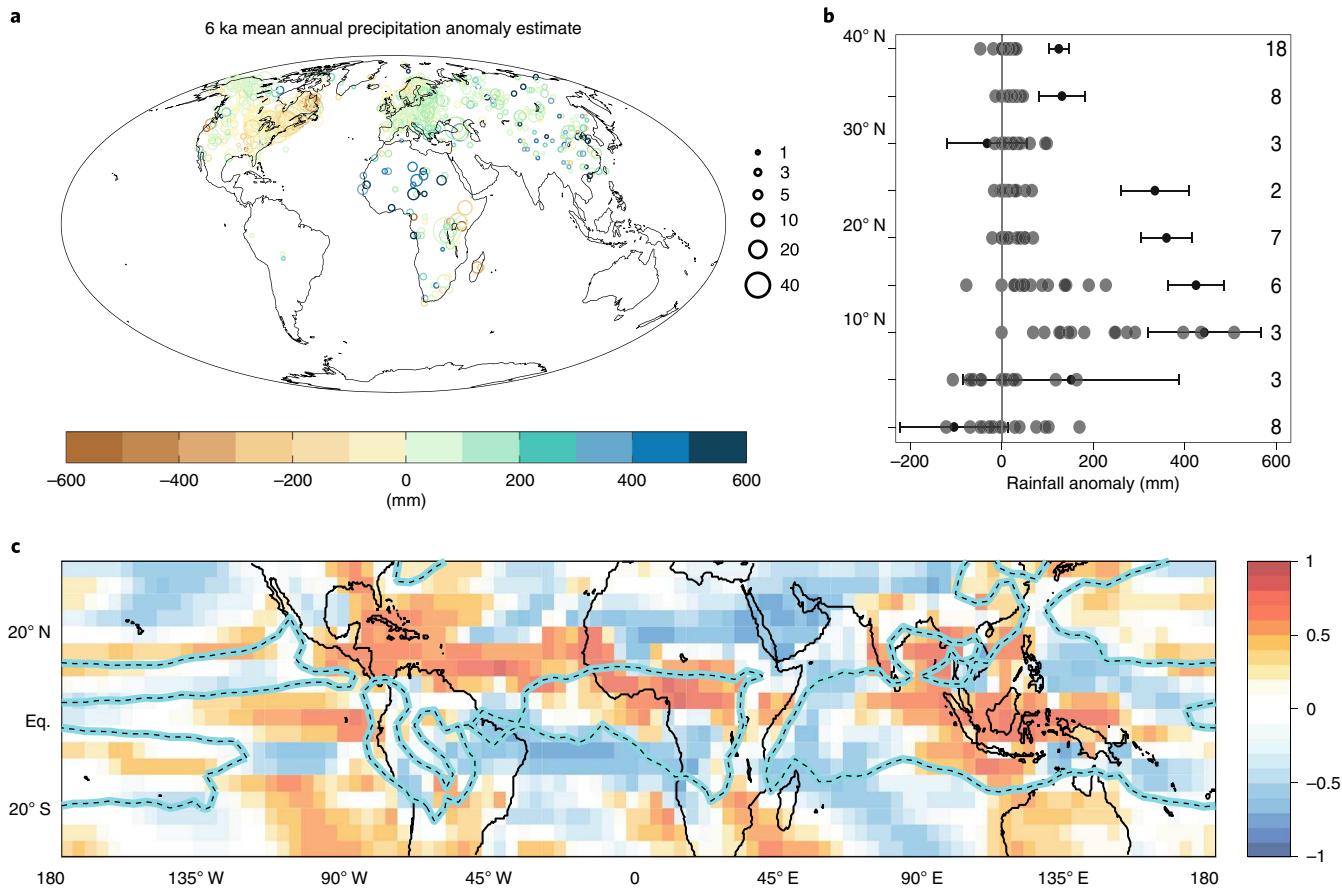


Fig. 2 | Palaeo changes in rainfall test comprehensive models and theoretical simplifications. **a**, Mid-Holocene mean annual precipitation anomalies (coloured circles, sized by the number of used reconstructions). **ka**, thousand years ago. **b**, CMIP5/PMIP3-simulated (grey circles) and reconstructed (black circles with error bars) mid-Holocene mean annual precipitation changes and grid cells contributing to the reconstruction and model means (numbers at right) for northern Africa latitude bands. **c**, Correlation of decadal rainfall variations with the first principal component of the zonal mean precipitation in transient Holocene simulations. See Methods for data references and further details. Panel **b** modified from ref. ⁷, Copernicus Publications.

developed for the Hadley circulation to describe how the full tropical circulation redistributes energy vertically and horizontally, from its poleward boundaries and across the tropics. The extension of the energy-budget framework from a purely zonal formulation to one that includes zonal asymmetries would parallel the progress that has been made in understanding how the momentum budget constrains tropical circulations. The momentum-budget framework has led us from a starting view of the Hadley circulation³⁰ and—to a degree—monsoons³¹ as axisymmetric angular-momentum-conserving circulations, to the recognition that momentum transport by eddies is crucial to maintaining the zonal mean circulation in most cases^{32,33}, to a deepening understanding of the role of the vorticity transport by stationary eddies in both localizing the monsoon circulation and regulating its intensity^{34,35}. An analogous energy-budget framework would invoke zonal asymmetries in surface properties, which introduce horizontal gradients in the distribution of energy, amplify the importance of both stationary and transient eddy transports, and can generate shallow circulations that can be energetically indirect (see section ‘The need for an energy-budget framework’). These asymmetries are expressed in the observable differences in convective weather characteristics between land and ocean (Fig. 1c,d), which aggregate into different ascent profiles and vertical energy transport (see section ‘Interactions between convection and large-scale climate’).

The eventual goal is to combine the momentum-based and energy-based theories of tropical circulations and rainfall into a self-contained model for the tropical climate, one that describes how the interplay between energy and momentum fluxes in a moist atmosphere—where a profusion of weather phenomena, cloud types and scales of motion is organized in mean and eddy effects—drives the seasonal evolution of oceanic and continental rainfall and controls monsoon diversity, variability and response to external forcing. Understanding how the clouds within such circulations modify the energy input into the atmosphere through changes in both radiative fluxes and turbulent surface fluxes would be the next challenge^{36–38}. In the rest of this Review Article, we propose a first step towards our goal: an energy-budget framework suitable for monsoons.

The explanatory power of energetic constraints

Energetic constraints provide a parsimonious explanation of how the zonal mean tropical rainfall shifts its position in response to internal variability and external forcings^{20,28,29}. The ascending branches of the Hadley cells are in the deep tropics, where the solar radiation absorbed by the Earth most greatly exceeds the terrestrial radiation emitted to space. Energy is transferred from the surface to the atmosphere by fluxes of radiation, sensible heat and moisture (latent heat), and a planetary circulation moves this excess energy towards high latitudes (Fig. 3a). Moist surface air converges

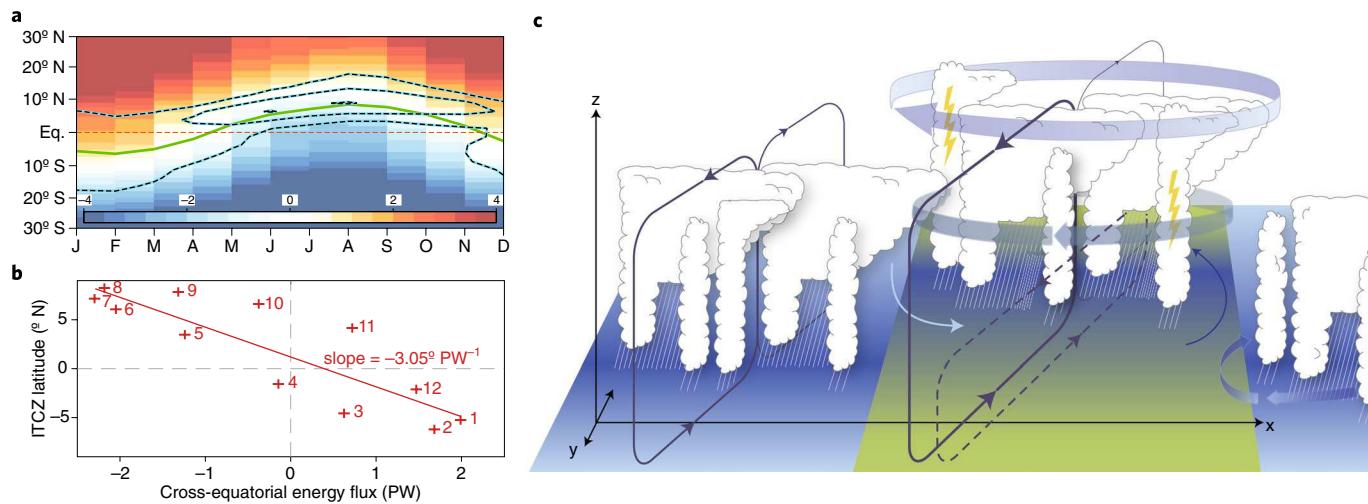


Fig. 3 | Rainfall in the ITCZ and in monsoons is linked to planetary and regional fluxes of energy. **a**, Zonal mean rainfall (dashed lines), ITCZ position (green line) and zonally and vertically integrated atmospheric energy transport (in PW, shaded). Letters on the x-axis represent the calendar month. **b**, Climatological ITCZ latitude as a function of the vertically integrated atmospheric energy transport at the equator. Numbers correspond to calendar month. The ellipse is sloped at 3° ITCZ shift per 1 PW energy flux. **c**, ITCZ (blue background) and monsoon (green background) schematics. The Hadley cells (dark solid lines) meet in the northern tropics, maximum ascent and rainfall occur in the winter cell (equatorward of the cell boundary) close to maximum low-level MSE (darker surface shading). Monsoons are distinguished by the additional shallow meridional circulation (dashed) and ventilation by the rotational horizontal flows (ribbon arrows) and by a different distribution of cloud types. See Methods for further details.

and rises in the ITCZ, thus cooling adiabatically, condensing moisture, and forming precipitating clouds. As a result, maximum rainfall is broadly co-located with the boundary between the Hadley cells (Fig. 3c).

These ideas can be formalized by focusing on the atmospheric energy budget in the annual and zonal means^{20,28}. Averaging in time and longitude disposes of tendency terms and zonal fluxes, so that, under the assumption that eddy terms are unimportant, net energy input into the atmosphere is balanced by divergence of the vertically integrated energy flux by the Hadley circulation (F), and the ITCZ coincides with the ‘energy-flux equator’, that is, the latitude at which the energy flux $F = 0$ and changes direction.

The Coriolis force is weak in the tropics and atmospheric waves are effective in smoothing out pressure gradients. Because the tropical troposphere cannot maintain strong temperature gradients, an extra-tropical eddy heat flux that reaches the tropics will be carried into the tropics and through to the mid-latitudes of the other hemisphere. This argument appears to apply more broadly: any asymmetry in the energy flux across the northern and southern edges of the tropics is felt throughout the tropical band³⁹, making the position of the energy-flux equator sensitive to extra-tropical forcings²⁰. The same argument explains how the inter-hemispheric asymmetry in seasonal insolation drives the north–south annual migration of the ITCZ²⁴. However, because the cross-equatorial Hadley cell is always the strongest throughout the annual cycle, and because maximum ascent and rainfall are concentrated within the cross-equatorial cell and equatorward of the energy-flux equator, seasonal rainfall shifts are less pronounced than those of the energy-flux equator²⁴ (Fig. 3a). In comprehensive models, changes in the annual mean position of the ITCZ, their inter-model spread and the seasonal migration of the rain belt have been shown to follow the same quantitative relationship with energy transports, albeit with some scatter and uncertainty (a 1 petawatt (PW) change in the cross-equatorial energy flux leads to about a 3° shift in the position of the ITCZ, Fig. 3b). This correspondence has been used to suggest that the same dynamics control both the meridional shifts characteristic of the annual cycle and the variability of mean rainfall at palaeo timescales^{20,24,40}.

This theory for the ITCZ position sees the tropical rain belt as the expression of the conservation law that governs energy in the

climate system; it links the occurrence of convection not just to the local environment, but to the planetary adjustments that bring the global atmosphere towards equilibrium; and it subsumes arguments that link the ITCZ position to gradients in tropical sea surface temperature. It has been used to explain the location of the modern ITCZ north of the equator as a consequence of energy transport by the thermohaline ocean circulation⁴¹, the southward shift of tropical rainfall during the Last Glacial Maximum, Dansgaard–Oeschger events as a consequence of northern high-latitude cooling^{29,40}, the effect of Eurasian afforestation on monsoons in the mid-Holocene⁴², the Sahel drought in the 1970s and 1980s as part of a global-scale southward ITCZ shift due to sulfate aerosols⁴³, and the role of Southern Ocean heat uptake in setting up the hemispheric asymmetry in future tropical rainfall changes⁴⁴.

However, the existing framework is limited in important ways. First, it is incomplete as a predictive theory because internal radiative feedbacks from clouds and water vapour can overwhelm the external forcings in setting energy gradients, are not easily predicted from gross energetic constraints, and vary substantially between models^{36,37,45}. Even when the radiative effect of clouds is carefully controlled in models^{46,47}, changes in oceanic heat transport can oppose the inter-hemispheric difference in energy input in the atmosphere and complicate the response of the ITCZ⁴⁸ to external forcings. Indeed, the tight coupling between atmospheric and oceanic heat transport^{20,49–51} suggests an expansion of atmospheric-only arguments. Second, changes in tropical rainfall are often better described as intensifications⁵² or contractions^{19,53–56} of the climatological net rainfall pattern, rather than shifts, and some regional anomalies are strongly driven by localized gradients in surface conditions^{57–59}, including between ocean and land^{60,61}. Third, the assumption that rainfall is the product of deep convection in which ascent extends throughout the troposphere and maximizes at mid-levels is crucial to the portrayal of the Hadley cell as energetically direct. Quantitatively, this requires that the same large-scale circulation that converges moisture into the ITCZ also diverges enough static energy from the ITCZ at upper levels that the vertically integrated result is a net export of MSE—a result that depends heavily on where exactly in the column air is converging and diverging. Yet, rainfall is also produced in circulations with shallow components and bottom-heavy

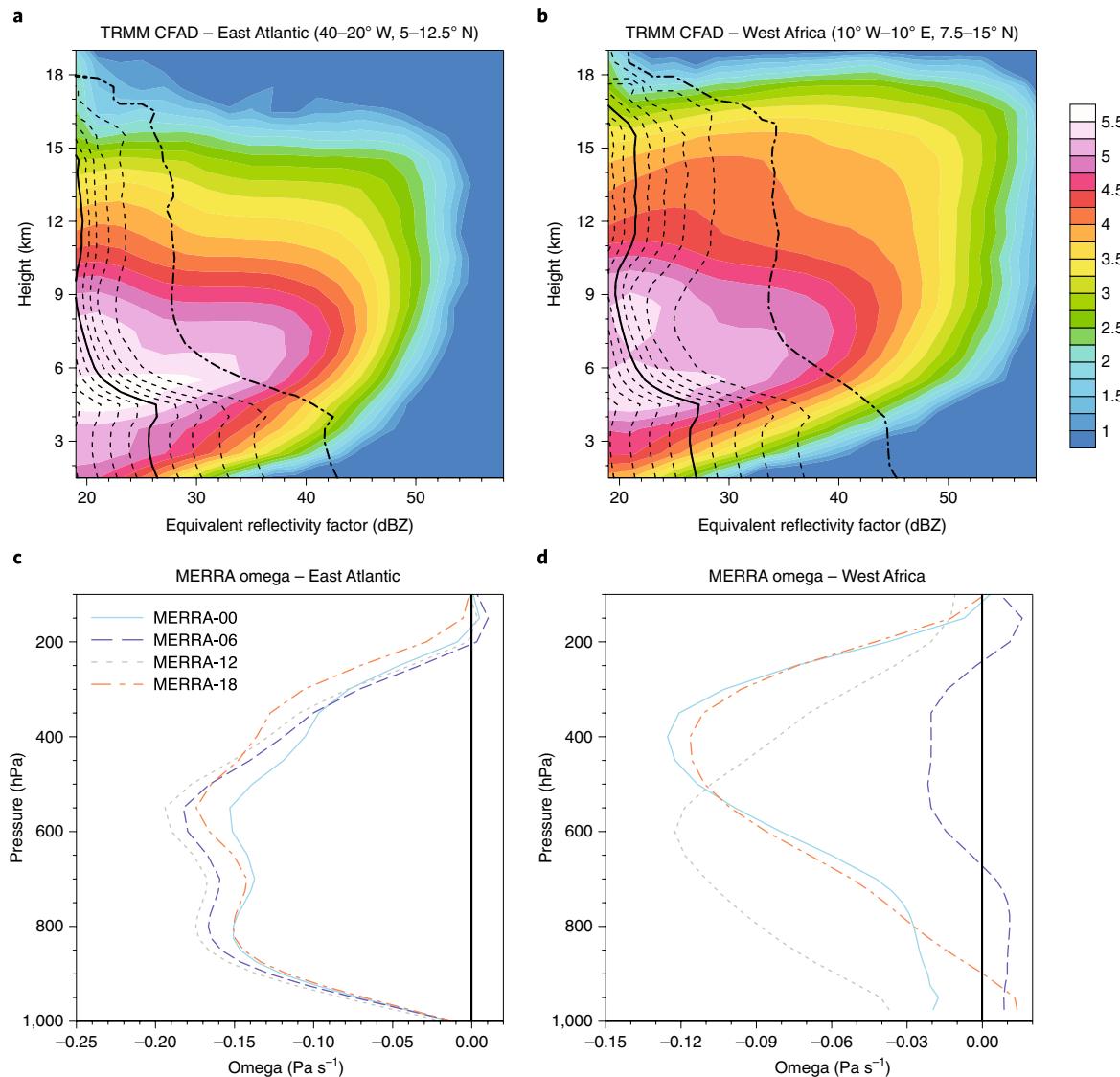


Fig. 4 | The aggregate character of convective systems presents regional differences. Cloud characteristics (top) and ascent profiles (bottom) over the eastern Atlantic ITCZ (left) and western Africa (right) as function of height. **a, b**, The contoured frequency with altitude diagrams (CFAD, filled contour) shows the frequency (logarithmic scale) of storm-top height as a function of rainfall intensity (units of reflectivity (dBZ)). Contours show percentile lines (solid, median; dashed-dotted, 99th percentile) of reflectivity frequency. Frequency is normalized at each level and all rain types are included. **c, d**, Daily evolution of the vertical velocity profiles (pressure coordinates; negative values indicate upward motion) from six-hourly MERRA reanalysis. See Methods for further details.

ascent profiles (see also Fig. 4), and direct calculation of the energy flux out of such regions indicate a net import of energy^{62,63}. Fourth, the assumption in the ITCZ energy-budget framework that eddies are not important is debatable. Even well within tropical latitudes, transient eddies transport latent heat poleward⁶⁴ and thus might be as important as the zonal circulation in transporting energy, especially where the latter is weak. While the bulk effect of transient eddies could be subsumed into the framework developed for the ITCZ in an aquaplanet^{28,65}, the presence of land introduces stationary eddies^{34,66} and localizes transient eddies³¹ in ways that preclude the straightforward application of the zonal mean framework to limited longitude bands (Fig. 2c). This is particularly true in a moist atmosphere, where feedbacks between surface fluxes, clouds and latent and radiative heating can amplify and extend the asymmetries in the forcing⁴⁵. These limitations hamper our ability to use the existing energy-budget framework for the ITCZ to accurately predict the response of regional rainfall to past or future forcings. Yet the fundamental

insight that regional changes are expressions of global conservation laws should not be abandoned.

The need for an energy-budget framework

Zonal asymmetries and contrasts between land and ocean are fundamental to the energy and the momentum budget of monsoon circulations^{13,34,67,68}. To close, the energy budget must include the effect of zonal transports and of the complex vertical structure of the meridional circulation, both of which are the result of inhomogeneities in surface properties. Inhomogeneities do occur over oceans (such as between warm pools and the equatorial cold tongues), but gradients in surface properties are especially strong at coastlines and over continents due to orography, variability in the characteristics of soils and vegetation, and—in a positive feedback—to the response of the land system to differences in precipitation. The conceptual model of the Hadley cell as a simple deep overturning meridional circulation (Fig. 3c)

can broadly capture the main features of the ITCZ, but is insufficient to describe the more complex monsoon circulations. In many regional monsoons^{22,69}, gradients in surface temperature and sensible heat fluxes (for example, between the hot subtropical desert and the cooler oceans and equatorial rainforests) force a shallow circulation with dry ascent poleward of the monsoon rainfall. This circulation includes the low-level monsoon flow that fuels moisture to the rainfall band and the dry return flow above that can cap deep convection, affecting the frequency of rainfall and the occurrence of severe convection (Figs. 1c,d and 3c). Re-evaporation of rainfall in a dryer lower troposphere (Fig. 3c, cloud types) can also affect the vertical distribution of latent heating and, concurrently, ascent (see also the next section). The presence of land thus changes the vertical transports in the meridional divergent circulation. The strong horizontal gradients in temperature and moisture associated with surface type also induce strong non-divergent, rotational flows that transport energy and moisture horizontally (Fig. 3c, broad arrows). Past seminal work⁶⁷ has shown that ventilation, the transport of low-MSE oceanic air by the rotational flow, is key to setting the poleward extent of the monsoon rainfall. Recent studies with more comprehensive models have confirmed the important role of the rotational flow in balancing the energy budget of monsoon regions⁷⁰ and in driving the onset of off-equatorial rainfall³⁴. The annual changes in surface properties that take place during the progression of the rainy season are also reflected in changing flows of energy in the atmosphere. The most obvious change is soil water content, which affects evaporative fluxes and albedo and, because of its high spatial variability, can introduce sharp gradients in surface properties at small spatial scales⁷¹.

There is a vast literature that focuses on the individual regional monsoons and emphasizes the zonally asymmetric regional flow and its interaction with mountains and regional oceans. There is value in formulating such regional analysis within a MSE framework. For example, the extent and intensity of the South Asian monsoon depends on the presence of the Himalayan mountains because the latter shield the MSE maximum over India from low-MSE extra-tropical air⁷². In another example, the African monsoon appears to be particularly susceptible to moisture anomalies entering from the North Atlantic^{73–76} and the Mediterranean⁷⁷ (even though the main moisture source for the monsoonal rains is in the tropical ocean) because they regulate the depth of convection⁷⁸, which is in turn linked to energetic requirements of the large-scale flow⁷⁰. How the profile of convective ascent evolves across the monsoon season modulates such sensitivity, by making changes in either the boundary-layer MSE (such as the horizontal advection or recycling of moisture) or the free-tropospheric dynamics (such as changes in upper-level vertical stability) more or less relevant (a similar dependence of rainfall sensitivity on low-level and upper-level processes is also seen across models^{70,73}).

The broad-brush framework built around the MSE budget has been successful in explaining why and how processes as disparate as aerosol microphysics and the oceanic thermohaline circulation can affect the zonal mean tropical rainfall. Here, we advocate extending the energy-budget framework from its current form, appropriate to the ITCZ/Hadley cell, to one appropriate for monsoons: one that quantifies how the presence of zonally confined land masses modifies both the input and the flow of energy in the atmosphere and thus changes the leading terms of the vertically integrated atmospheric energy budget. In this section, we have emphasized: (1) how land/ocean differences in surface properties change the input of moist and dry energy to the atmosphere; (2) how the continental-scale stationary circulations that develop in response advect MSE gradients horizontally through the geostrophic flow, and vertically through shallow circulations; and (3) how changes in the dominant balance shape both the diversity of regional monsoons and mon-

soon evolution through the development and decay of the rainy season. In the next section, we focus on the role of vertical motion at the scale of clouds, and how it shapes the vertical fluxes of moist energy and its input in the atmosphere.

Interactions between convection and large-scale climate

Recent observations have highlighted the rich diversity of convective systems (Figs. 1 and 4) and allowed fundamental insights into the processes that govern them. Convective precipitation appears controlled by both low-level and deep-column moisture, which set the buoyancy of entraining ascending parcels. Thus, while early theories predict the position of the rainfall maximum from just the boundary-layer MSE (see section ‘The explanatory power of energetic constraints’), full tropospheric water vapour is key to fully account for the spatial and temporal variability of rainfall intensity. In turn, processes such as detrainment from precipitating clouds and re-evaporation of rain make the tropospheric water vapour depend on the occurrence of rainfall. This two-way coupling underpins the correspondence between rainfall amounts and total humidity found over land and ocean (see also Fig. 1a) and is encapsulated in an exponential relationship of daily precipitation intensity on the integrated humidity⁷⁹ (with some variations across convective system⁸⁰ and their drivers⁸¹). However, rainfall characteristics, such as intensity, organization and duration, and the vertical and temporal distribution of clouds depend on factors other than column humidity, including wind shear, the larger-scale flow, and the properties of the surface boundary^{71,82,83}. Continental updrafts are often deeper and more intense than oceanic updrafts (Fig. 4a,b), as evidenced by the preferential occurrence of lightning over land (Fig. 1c), but land convection varies greatly through the day (compare Fig. 4c,d) and season in depth, organization and lifetime—affected by surface inhomogeneity and by stronger triggering and inhibition processes. The growing appreciation of the diversity of convective cloud systems has yet to mature into enough understanding of the interplay between clouds and large-scale dynamics to create much-needed convective parameterizations able to describe such diversity. Parameterizations are still overly reliant on so-called quasi-equilibrium formulations: the gross effects of convection are taken to be nearly instantaneous and deterministic functions of large-scale forcings, thermodynamics drivers are overemphasized over dynamics, and convective organization is typically ignored^{84,85}. The approximations lead to a preference for tall, disorganized convection and the lack of both shallower, developing convection⁸⁶ and more persistent cloud systems, and is likely the source of climate models’ systematic errors in rainfall timing (for example, the too-early peak in the diurnal cycle of land rainfall). It affects the climate at longer timescales as well: model biases in seasonal rainfall are typically established within few days after initialization from observations, and only later amplified in the coupled system, pointing to the dominance of fast atmospheric processes⁸⁷ in setting the bias. Such long-ranging effects are not surprising: the aggregate effect of convection is reflected in the profiles of horizontal convergence and divergence and in cloud and moisture radiative effects and will thus affect both the net vertically integrated energy flux and the total energy input into the atmosphere. The extent and vertical distribution of clouds modify the net energy in the atmosphere via changes in radiative fluxes and, by modifying the structure of the boundary layer, turbulent fluxes. The vertical profile of ascent modulates the energy transport by the circulation, to the point that a predominance of bottom-heavy or top-heavy convection can determine whether the circulation imports or exports MSE to or from a convective region^{62,63}. Thus, shortcomings in the mix of deep convective and stratiform rain production, of warm and cold cloud microphysics, and of mixing and re-evaporation, can translate into global scale biases⁸⁸ and introduce another source of uncertainty in future projections⁷⁰.

High-resolution cloud-resolving dynamical models can portray the full evolution and organization of cloud systems^{88,89} and—when coupled to parameterized large-scale fields⁹⁰ that describe a broad suite of boundary conditions and environments—are being used to investigate how the environment shapes the rich diversity of cloud characteristics and how cloud processes feedback on the environment. These experiments have shown that the specific way in which land convection evolves during the day (such as through morning fog or land–sea breezes) is a key determinant of the large-scale mean and the seasonal evolution of rainfall and other environmental variables^{91,92}. More research on tropical land convection is needed to elucidate whether this cross-scale link is achieved because land processes and clouds modify the energy input into the atmosphere, or because the daily evolution of convection changes the profile of ascent and this effect is rectified into changes in the vertical energy flux.

Synthesis

Simulations of the diurnal and seasonal cycle of rainfall in the current climate share common biases, suggesting similar structural deficiencies across global climate models. Common structural deficiencies are also apparent in the common biases of rainfall anomalies in simulations of the past. These common shortcomings in both the basic state and the response to climate forcings diminish the value of consensus in projections of the climate response to anthropogenic emissions. Convective parameterizations, which produce most tropical rainfall within current climate models, have inherent limitations that are a likely cause of bias, but whether cloud-resolving global models will be able to adequately capture observed rainfall variations in modern or past climate states remains to be seen. Oceanic processes that amplify atmospheric biases in the seasonal cycle⁹³ are likely to play a role in setting long-term trends⁹⁴ and simulations of past climate states highlight the additional importance of Earth-system feedbacks, such as between precipitating atmospheric dynamics, vegetation cover and soil composition. We propose that a modern energy-budget framework, one modelled on past efforts^{20,21,32,95,96} but aware of the advances and challenges outlined in this Review Article, can supply metrics to improve our understanding of the physical mechanisms of tropical rainfall and to evaluate its representation in climate models.

Developing the framework that we are advocating will require a combination of observational, modelling and theoretical approaches, creating extensive modern and palaeoclimate datasets to serve as the ultimate test of our theories.

The mechanisms by which convection becomes organized and monsoon rainfall interacts with its environment on sub-daily-to-monthly timescales will only be elucidated if we gain high-resolution observations of rainfall and vertical atmosphere structure in combination with cloud-resolving simulations⁸⁴. Such observations must also be expanded to encompass a broader cross-section of tropical environments and climate regimes—including a variety of chronically undersampled land regions—to provide a truly useful benchmark. Collection of new palaeo evidence and quantitative reconstructions of mid-Holocene precipitation in the Southern Hemisphere are also required to test model simulations and the theoretical constraints on monsoon dynamics; in turn, forward proxy models and further climate simulations will help morph sparse environmental observations into a coherent portrait of past climates, and help prioritize new data acquisition. Multi-model simulations⁹⁷, simulations with a broad set of forcing⁴⁸, and simulations with models of different vintage are all needed to provide a robust test for a new monsoon framework.

Drawing our lessons from the development of budget-based theories of the ITCZ, we propose that a systematic understanding of monsoons will require a comprehensive hierarchy of model simulations and a common set of energy and momentum diagnostics to compare results across the model hierarchy and with observations. A

key element of such a hierarchy are idealized model setups in which energy is conserved through fluxes at the atmospheric boundaries, and that are designed to highlight the effect of different land characteristics—themselves isolated through ad-hoc idealizations—on the response of the tropical rain belt to a range of external forcings^{98–100}. Accurate diagnostics of how energy flows into and within the tropical atmosphere, and via different elements of the atmospheric circulation (from the zonal mean, to stationary eddies, to synoptic eddies, to much faster variability at the scale of convection) are often difficult to calculate from available model output⁵⁴. Yet, they must underpin any metrics used to elucidate which small- and large-scale processes are essential to the monsoon systems (and by which interactions); to test the coherence of our dynamical theories; and to benchmark model development. A theoretical model that can deduce the behaviour of tropical rainfall from simple indicators based on conservation laws of the physical system with minimal empirical or ad-hoc assumptions would cap the hierarchy and provide the grounds for interpreting both idealized and realistic simulations and for understanding the observed behaviour of the Earth system.

To assess how rainfall might change in the future, we must be able to understand and represent the critical processes that affect tropical rain belts. It is a formidable change to derive such a unified theory of the ITCZ and the regional monsoons from basic conservation laws, but we must meet it to confidently link past to present to future.

Methods

Methods, including statements of data availability and any associated accession codes and references, are available at <https://doi.org/10.1038/s41561-018-0137-1>.

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Author contributions

M.B. led the writing process and produced Fig. 1 (from an idea by B.E.M.), Fig. 2 (from data provided by S.P.H. and P.B.), and Fig. 3 (in collaboration with W.R.B. and A.V.). C.S. produced Fig. 4. All authors collaboratively drafted the outline of the paper and greatly contributed to the writing process.

Competing interests

The authors declare no competing interests.

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Correspondence should be addressed to M.B.

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Methods

Figure 1. Characteristics of observed tropical rainfall. On any given day, rain is produced by cloud systems that range in size from individual convective towers (1–10 km) to mesoscale convective systems and tropical cyclones (>100 km). Despite substantial variability on short timescales and small spatial scales, most disturbances are organized within the large-scale rain belts formed by the monsoons and the inter-tropical convergence zone (>1,000 km). The map (Fig. 1a, colour) and zonal mean (Fig. 1b, dark green) of daily rainfall are obtained from merged GPM satellite measurements calibrated against rain gauges¹⁰¹. Column integrated water vapour (Fig. 1a, grey shading and magenta contour) was obtained from the ERA-Interim reanalysis product¹⁰². The summer climatological zonal mean rainfall (Fig. 1b, light green) is calculated as the 1998–2014 average of the 11-day period centred around 27 July and obtained from the Tropical Rainfall Measuring Mission (TRMM)-3B42 rainfall estimate¹⁰³. Climatological conditional intensity of rainfall (Fig. 1c, coloured shading; Fig. 1d) is obtained from the precipitation radar data on the TRMM satellite¹⁰⁴ and expressed as a reflectivity in units of dBZ (decibels of Z). Reflectivity is the amount of transmitted power returned to the radar receiver; light rain is detected by the TRMM precipitation radar when the dBZ value reaches 18 (corresponding to about 0.4 mm hr⁻¹). The higher the dBZ, the stronger the rain rate. Uncertainty in the methods for translating reflectivity into a quantitative precipitation estimates are detailed in ref. ¹⁰⁵. Lightning flashes for 27 July 2014 are obtained from lightning imaging sensor during ascending passes of the TRMM satellite (<https://lightning.nsstc.nasa.gov/lisib/lisbrowsecal.pl?month=7&year=2014>).

Figure 2. Changes in rainfall in the Holocene indicate a complex behaviour across the monsoon systems not fully captured by climate models, and do not suggest that meridional displacements of the zonal mean ITCZ explain a large fraction of the variance of tropical continental rainfall. Figure 2a: the expectation that Southern Hemisphere monsoons would be weaker in periods of weaker Southern Hemisphere summer insolation is not fully supported by current observations. Quantitative reconstructions of changes in mean annual precipitation between the mid-Holocene (11,000–5,000 years before present) and present day are shown by coloured circles—the colour indicates the size of the anomaly (mm, colourbar), while the size indicates the number of reconstructions used for the estimate, as indicated by the legend inside the map. The data were first published by Bartlein et al.¹⁰⁶. Figure 2b: CMIP5/PMIP3-simulated and reconstructed changes in mean annual precipitation in the mid-Holocene for five latitude bands across northern Africa (longitude 20° W–40° E between 0°–45° N), where the model results are averages for the grid cells with observations and each model is represented by a different grey circle. The mean and standard error of the reconstructions are shown in black and the number of grid cells contributing to the reconstruction is shown for each latitude band. CMIP5/PMIP3 data were available thanks to the World Climate Research Programme's Working Group on Coupled Modelling, which is responsible for CMIP. For CMIP, the US Department of Energy's Program for Climate Model Diagnosis and Intercomparison provides coordinating support and led development of software infrastructure in partnership with the Global Organization for Earth System Science Portals. The figure is modified from ref. ¹⁰⁷. Figure 2c: correlation maps (shading) of decadal rainfall variations at each gridpoint with the first principal component of the zonal mean precipitation in the fully forced TrACE-21000 (ref. ¹⁰⁸) simulation for the period 9.5–0.5 ka before present. The dashed line is a representative contour for the mean precipitation and indicates the climatological position of the rain belt.

Figure 3. The energy-budget framework for the tropical rain belt. Figure 3a: the seasonal evolution of the zonal mean observed climatological rainfall (dashed contours, only the 4, 6 and 8 mm day⁻¹ isolines are shown) and ITCZ position (defined as the centroid of zonal mean rainfall within 20 N/S; green line) are superimposed on a reanalysis-based estimate of the zonally and vertically integrated atmospheric energy transport (shaded, warm and cool colours indicate northward and southward transport, respectively, and the white area indicates the energy flux equator, units of PW). Figure 3b: the seasonal relationship between the vertically integrated atmospheric energy flux at the equator (F_0) and the ITCZ position. The energy flux associated with the net mass movement between the hemispheres is retained here, leading to a phase lag between the two fields. The slope of the relationship, given by the direction of the major axis of the ellipse, indicates a 3° shift for a 1 PW energy flux, consistent with calculations that omit the barotropic circulation.

The climatological rainfall in Fig. 3a,b is calculated as the 1979–2013 average of the Climate Prediction Center (CPC) Merged Analysis of Precipitation (CMAP) dataset¹⁰⁹. The energy fields are from ERA-Interim reanalysis¹⁰² over the same period. The cross-equatorial energy transport is calculated eliminating the mass budget residual before vertically integrating the fluxes, the data were provided by the National Center for Atmospheric Research (retrieved from <https://climatedataguide.ucar.edu/climate-data/era-interim-derived-components>).

Figure 3c: schematic of the ITCZ (depicted over an oceanic surface, blue half on the left) and monsoon (depicted over a continental surface, green half on the right) circulations for Northern Hemisphere summer. The summer and winter Hadley cells (dark solid lines) meet in the northern tropics, close to where

the low-level MSE is maximum (darker shading at the surface), consistent with convective quasi-equilibrium theory. Most upward motion and thus most rainfall occurs in the ascending branch of the stronger, winter cell, so maximum rainfall is slightly equatorward of the Hadley cell boundary and the energy flux equator. As is the case for the ITCZ, rainfall associated with the monsoonal circulation is positioned slightly equatorward of the maximum surface moist energy (dark blue shading) and is associated with large-scale ascent in local meridional overturning cells whose strength is greater when the ascent is further away from the equator. Key distinctions for the monsoons are in the complexity of the circulation and the distribution of cloud types. Notice in the land portion of the diagram the presence of a shallow meridional circulation (dashed lines) with ascent poleward of the rain band and a dry return flow, the rotational circulations associated with the low-level cyclone (light and dark blue ribbon-width arrows indicate negative and positive transport of MSE), mid- and upper-level land anticyclones, and the oceanic anticyclone (anticyclonic circulations are depicted with ribbon-width grey arrows). Notice also the deeper and more intense convection over land (indicated by a distribution of clouds that include more overshoots and fewer clouds lacking an anvil), increased lightning, reduced rain from warm cloud (fewer clouds without anvil), and more re-evaporation of rain (dotted rain from the anvil cloud).

This schematic highlights the aspects of a hypothetical 'essential' monsoon that are addressed in this Review Article. It is not meant to represent any particular monsoon system, as each is highly affected by the geometry of the continent, the location and orientation of orography, the geographical distribution of surface types, including deserts, and oceanic processes unique to each ocean basin.

Figure 4. Cloud characteristics and ascent profiles over ocean and land during the peak of the rainy season. Left: the eastern Atlantic ITCZ (40–20° W, 5–12.5° N). Right: western Africa (10° W–10° E, 7.5–15° N). The contoured frequency with altitude diagrams (CFAD) in the top panels show that western Africa has deeper, more intense convective cells, while the eastern Atlantic has more mid-level rainy cloud. The filled contour shows the log of the frequency of storm top height as a function of radar reflectivity (a measure of rainfall intensity). Both regions show that storms reaching 5 km in height and measuring reflectivities of less than 30 dBZ are the most frequent, but this peak is more pronounced over ocean than land. Conversely, the land region shows more frequent instances of storms that have intensity above 50 dBZ (colours extending to the right of the diagram), and reach 18 km in height (colours extending to the top of the diagram). Overlaid on the colour are percentile lines of reflectivity frequency at each level with the black line equal to the median at each level. The far right line is the 99th percentile line. Notice that the median surface reflectivity is slightly higher over land while the 99th percentile reflectivity is much higher over land, indicating that land convection reaches more extreme values of intensity. The secondary reflectivity maximum in the median line is also more noticeable over land than ocean, suggesting again significant differences in the vertical profile of the cloud systems over land and ocean. Frequency is normalized at each level and all rain types (stratiform, convective, shallow isolated, shallow non-isolated) are included. The bottom panels show how ascent profiles in rain systems have much larger diurnal variations over West Africa and are much more top-heavy than over the adjacent ocean. Profiles of vertical velocity in pressure coordinates (negative values indicate upward motion) were obtained from the MERRA reanalysis.

The TRMM reflectivity data¹⁰ were originally sourced from TRMM orbital files (2A23/2A25) for the 1998–2014 period (August values only). Vertical omega profiles are calculated from 6-hourly MERRA reanalysis¹¹ for August days in 1983–2007. Only samples that contributed the top 50% of rainfall by volume were included (using the MERRA surface precipitation flux data for all rain and rainfall thresholds unique to each domain and six-hourly period). When all times are included, the East Atlantic ascent is even more bottom-heavy than shown in Fig. 4, but the qualitative comparison to land convection is unchanged.

Data availability. No new datasets were generated for this Review Article, and all data used to produce the figures are in the public domain. GPM data are available from <https://pmm.nasa.gov/data-access/downloads/gpm>. ERA Interim reanalysis data are available from <http://apps.ecmwf.int/datasets/data/interim-full-daily/levtype=sfc/>. TRMM instantaneous intensity data are available from http://kage.ldeo.columbia.edu:81/SOURCES/LDEO/.ClimateGroup/DATASETS/TRMM_PR/V2/V2/. Lightning flashes data are available from <https://lightning.nsstc.nasa.gov/lisib/>. PMIP3 and TrACE data are available through the Earth System Grid <https://esgf.llnl.gov/>. TRMM-3B42 and CMAP precipitation data were obtained through the IRI data library <https://irid.ldeo.columbia.edu/>. ERA-interim-derived energy transports were obtained from <https://climatedataguide.ucar.edu/climate-data/era-interim-derived-components>. TRMM orbital files are available from https://disc.gsfc.nasa.gov/datasets/TRMM_2A23_V7/summary?keywords=TRMM. MERRA data are available from <https://disc.gsfc.nasa.gov/datasets?page=1&keywords=merra>.

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