

The density of Li_2CO_3 - Na_2CO_3 - K_2CO_3 - Rb_2CO_3 - Cs_2CO_3 - CaCO_3 - SrCO_3 - BaCO_3 liquids: New measurements, ideal mixing, and systematic trends with composition

Sean M. Hurt^{*}, Rebecca A. Lange

Department of Earth and Environmental Sciences, University of Michigan, 2534 C.C. Little Bldg., 1100 N. University Ave., Ann Arbor, MI 48109-1005, USA

Received 19 October 2017; accepted in revised form 20 December 2018; available online 4 January 2019

Abstract

The densities of 15 liquids in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 - Rb_2CO_3 - Cs_2CO_3 - CaCO_3 - SrCO_3 - BaCO_3 system were measured at 1 bar with the Pt double-bob Archimedean method between 758 and 1455 K. Melt compositions contain ≤ 50 mol% CaCO_3 , SrCO_3 and BaCO_3 , and $\leq 100\%$ Rb_2CO_3 and Cs_2CO_3 . These results were combined with double-bob density measurements from the literature on nine liquids in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 - CaCO_3 system to calibrate a new linear volume equation for multicomponent carbonate liquids. The regression led to the following fitted partial molar volumes ($\pm 1\sigma$ cm³/mol) at 1100 K for Li_2CO_3 (41.22 ± 0.09), Na_2CO_3 (53.27 ± 0.11), K_2CO_3 (71.59 ± 0.13), Rb_2CO_3 (80.78 ± 0.11), Cs_2CO_3 (94.00 ± 0.09), CaCO_3 (40.18 ± 0.16), SrCO_3 (44.33 ± 0.22) and BaCO_3 (50.99 ± 0.19). At 1100 K, the thermal expansion coefficients of all alkali carbonate liquid components are indistinguishable within 1-sigma error ($22.07 \pm 1.66 \times 10^{-5} \text{ K}^{-1}$), but differ from the thermal expansion coefficients of all alkaline-earth carbonate liquid components, which are also indistinguishable within 1-sigma error ($16.40 \pm 2.85 \times 10^{-5} \text{ K}^{-1}$). The linear volume equation recovers the measurements within analytical error ($\pm 0.3\%$). The partial molar volumes of all eight carbonate components increase linearly along two different trends, one for the alkali carbonates ($R^2 = 0.999$) and another for the alkaline earth carbonates ($R^2 = 0.999$) as a function of cation volumes, where metal cation-oxygen coordination numbers are obtained from molecular dynamic simulations in the literature (ranging from 4- to 6-fold among the alkali metals and 7- to 8-fold among the alkaline-earth metals). The linear fits lead to two different partial molar volumes (~ 38 and ~ 31 cm³/mole) for the carbonate ion (CO_3^{2-}) at 1100 K, which reflects a more open topological arrangement of the carbonate ions when in 4-fold coordination with the alkali metals vs. 6-fold coordination with alkaline earth metals. The results permit the partial molar volume of MgCO_3 and FeCO_3 in multicomponent carbonate liquids to be calculated if the oxygen and carbonate ion coordination with Mg^{2+} and Fe^{2+} are known. For example, in the case where Mg^{2+} and Fe^{2+} are in 6-fold coordination with both oxygen and carbonate, the estimated partial molar volumes at 1100 K are $34.4 (\pm 0.6)$, and $35.1 (\pm 0.6)$ cm³/mol, respectively, with a thermal expansion coefficient of $16.40 (\pm 2.85) 10^{-5} \text{ K}^{-1}$. In contrast, if Mg^{2+} and Fe^{2+} are in 4-fold coordination with both oxygen and carbonate, the estimated partial molar volumes at 1100 K are $40.0 (\pm 0.6)$, and $40.4 (\pm 0.6)$ cm³/mol, respectively, with a thermal expansion coefficient of $22.07 (\pm 1.66) 10^{-5} \text{ K}^{-1}$. To resolve these different estimates of molar volume, molecular dynamic simulations are needed to determine the respective coordination environments for Mg^{2+} and Fe^{2+} in carbonate liquids.

© 2018 Elsevier Ltd. All rights reserved.

^{*} Corresponding author.

E-mail address: seanhurt@umich.edu (S.M. Hurt).

1. INTRODUCTION

Despite its relatively low abundance in the mantle, carbonate plays a key role in the global carbon cycle and in the formation of mantle melts in various tectonic settings (e.g., Dasgupta and Hirschmann, 2010). Carbonate may be a major factor in deep (240–300 km) partial melting of the mantle beneath mid-ocean spreading ridges and intra-plate settings (e.g., Dasgupta and Hirschmann, 2006), and it plays a critical role in the genesis of silica-undersaturated alkaline magmas (e.g., Dasgupta et al., 2006). Moreover, because carbonate melts are efficient scavengers of various incompatible elements (e.g., Blundy and Dalton, 2000), they are also effective agents of mantle metasomatism (e.g., Dalton and Wood, 1993). Therefore, understanding the stability of carbonate melts under mantle conditions and incorporating them into thermodynamic models (e.g., p-MELTS; Ghiorso et al., 2002) is a major imperative. In order to extend thermodynamic models to include carbonate melts it is necessary to constrain their equation of state (P-V-T relation). The objective of this study is to expand the compositional range of carbonate liquids for which direct density measurements are known at 1-bar, with a focus on alkaline-earth components.

The three most abundant carbonate components in the mantle are MgCO_3 , CaCO_3 and FeCO_3 ; however these three pure end-member liquids are not stable at one bar and decompose at temperatures below their melting temperatures. In previous studies, Liu and Lange (2003) and O'Leary et al. (2015) showed that the CaCO_3 component could be added to low-temperature alkali carbonate liquids to form CaCO_3 -bearing liquids that are stable at one bar. The same strategy was employed in this study, where CaCO_3 , SrCO_3 , and BaCO_3 were added in various proportions (≤ 50 mol%) to the eutectic in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 ternary to permit density measurements of alkaline earth carbonate-bearing liquids to be obtained over as wide a range of composition and temperature as possible. In addition, the density of melts with 50–100% Rb_2CO_3 and Cs_2CO_3 were also measured. By expanding the compositional data set for which double-bob liquid density measurements are currently available from four to eight

components (Li_2CO_3 - Na_2CO_3 - K_2CO_3 - Rb_2CO_3 - Cs_2CO_3 - CaCO_3 - SrCO_3 - BaCO_3), systematic variations with composition and temperature can be evaluated. This in turn allows constraints to be placed on the partial molar volume and thermal expansion of the MgCO_3 and FeCO_3 components in multicomponent carbonate liquids at 1 bar. Standard state thermodynamic property estimates on these critical components are necessary to extend thermodynamic models such as p-MELTS to carbonated-bearing lithologies relevant to the mantle.

2. MATERIALS AND METHODS

2.1. Sample synthesis

Fifteen samples in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 - Rb_2CO_3 - Cs_2CO_3 - CaCO_3 - SrCO_3 - BaCO_3 system were synthesized in this study (Table 1). For 13 of these samples, the eutectic composition in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 ternary (referred to as T in Table 1) was combined with varying amounts of CaCO_3 , SrCO_3 , BaCO_3 , Cs_2CO_3 and Rb_2CO_3 . The remaining two samples are pure Rb_2CO_3 and Cs_2CO_3 . The samples were made in 200 g batches from powdered or granular reagents with $>99.5\%$ purity. Before synthesis, the loss on ignition (LOI) of each reagent was determined and used to calculate the appropriate weight of each reagent to be poured directly into the glass jar of an electric blender. The reagents were mixed at high speeds for several minutes to make a powdered sample with a homogenous composition.

2.2. Double-bob Archimedean method

The double-bob Archimedean method was used to determine the density of each experimental liquid over a 697° temperature interval (758–1455 K), following the methods described in detail in Lange and Carmichael (1987) and Liu and Lange (2003). Sufficient amounts of each sample (~ 75 g) were loaded into a Pt crucible to ensure a liquid depth of ~ 2.5 cm. The sample and Pt crucible were loaded into a Deltech vertical tube furnace and brought up to temperature under a constant stream of pure CO_2 gas.

Table 1
Molar composition and gram formula weight (g.f.w.) of experimental liquids.

Sample	Li_2CO_3	Na_2CO_3	K_2CO_3	Rb_2CO_3	Cs_2CO_3	CaCO_3	SrCO_3	BaCO_3	g.f.w.
T-50%Rb				0.500					124.22
Rb_2CO_3				1.000					230.95
T-50%Cs	0.218	0.158	0.125		0.500				212.95
Cs_2CO_3					1.000				325.82
T-25% Ca	0.326	0.236	0.188			0.250			100.08
T-35%Ca	0.283	0.205	0.163			0.350			100.08
T-50%Ca	0.218	0.158	0.125			0.500			100.08
T-10%Sr	0.392	0.284	0.225				0.100		104.84
T-20%Sr	0.348	0.252	0.200				0.200		109.59
T-30%Sr	0.305	0.221	0.175				0.300		114.35
T-50%Sr	0.218	0.158	0.125				0.500		123.79
T-15%Ba	0.370	0.268	0.213					0.150	114.67
T-35%Ba	0.283	0.205	0.163					0.350	134.12
T-50%Ba	0.218	0.158	0.125					0.500	148.71
T-50%CaSrBa	0.218	0.158	0.125			0.167	0.167	0.167	111.26

A density measurement was made by suspending a Pt bob from a mass balance (with a precision of ± 0.0001 g) above the surface of the sample liquid. The bob and sample were allowed at least thirty minutes to come to thermal equilibrium. The bob was weighed and then immersed in the sample liquid and weighed again to obtain buoyancy, $B(T)$. The density of the liquid is obtained from the relation:

$$\rho^{liq}(T) = \frac{B(T) + S(T)}{V(T)} \quad (1)$$

where $\rho^{liq}(T)$ is the density of the liquid at the experimental temperature, $B(T)$ is the buoyancy, $S(T)$ is the surface tension effect from the liquid surface pulling on the stem of the bob, and $V(T)$ is the volume of the bob immersed in the liquid. The procedure was repeated with a second bob of different size, which allows the $S(T)$ term to be eliminated in the calculation of density with the following equation:

$$\rho^{liq}(T) = \frac{B_L(T) - B_S(T)}{V_L(T) - V_S(T)} \quad (2)$$

where the L and S subscripts denote a large and small bob, respectively. The buoyancy measurements were repeated with two different large bobs (26.4 and 25.9 g) and two different small bobs (11.7 and 9.8 g), which allowed four different density measurements to be obtained at each experimental temperature, using Eq. (2). To further assess reproducibility, replicate density measurements were made on two different aliquots from the same synthesized sample batch in two cases (T-10%Sr and T-50%Ba), and the T-35%Ba sample was re-synthesized from scratch and its density re-measured. The BaCO_3 -bearing liquids posed experimental difficulties not found for the other sample liquids, where the platinum bobs were prone to adhering to the sides of the crucible as they were immersed during a buoyancy measurement.

Density measurements were made at three to four different temperatures, with the temperature interval bounded by the liquidus and decomposition temperature of each sample. For one sample (T-35%Ca; Table 1), liquid density was measured at only one temperature because the supply of CO_2 ran out and there was insufficient CO_2 in the furnace tube to prevent decarbonation (decomposition) of the sample. For the alkaline earth carbonate components, measurements were made over temperature intervals that ranged from 444 (CaCO_3), 697 (SrCO_3) and 407 (BaCO_3) K. For the Rb_2CO_3 and Cs_2CO_3 components, density measurements were made over temperature intervals of 392 and 355 K, respectively.

2.2.1. Temperature control

Determination of the experimental temperature during each density measurement was obtained with the following calibration. The Pt crucible was filled with alumina (Al_2O_3) powder to the same depth as that used for the experimental liquids and placed in the furnace. A thermocouple was buried two cm deep into the powder, a depth corresponding to that of immersed bobs, and the recorded temperature was used to locate the furnace hotspot and isothermal zone. The crucible was situated so that the sample was near the bottom of the isothermal zone to ensure that the liquid at

the bottom of the crucible was not hotter than the top, which prevented convection. The furnace was then set to various temperatures, spanning the range used in the density experiments (800–1500 K). The thermocouple temperature was recorded at each set point and used to create a temperature calibration curve. The thermocouple used in temperature calibration was also tested for accuracy against the melting point of gold; the gold melted within five degrees of the accepted value (1337 K).

2.2.2. Density measurements of liquid NaCl as a standard

To verify the accuracy of the Pt double-bob Archimedeian method and the temperature calibration curve employed in this study, the density of liquid NaCl was measured eleven different times between 1102 and 1455 K. These measurements were made before, during and after the density measurements on the sample liquids in Table 1 were made.

3. RESULTS

3.1. NaCl liquid density

The NaCl liquid density results are listed in Table 2 at each temperature of measurement, together with molar volume, which is calculated by dividing the molar mass of NaCl (58.44 g/mol) by density. In Fig. 1, the liquid molar volume data from this study are plotted against temperature and compared to those reported in Stein et al. (1986), Lange and Carmichael (1987), Liu and Lange (2001, 2006) and Guo et al. (2013), all of which are based on the double-bob Archimedeian method. The results from this study expand the temperature interval for which density data are available for NaCl liquid from 194 to 353°, which improves estimates of its thermal expansion. A linear fit to the data in Fig. 1 gives the molar volume of NaCl liquid as a function of temperature: $V(T) = 22.11 + 0.0143 \cdot T$ (K) cm^3/mol . This linear fit recovers the measurements from the literature within 0.35%, on average, and those from this study within 0.25%, on average.

3.2. Carbonate liquid densities

The density results for the 15 experimental liquids at each temperature of measurement are reported in Table 3.

Table 2
NaCl liquid density measurements.

Temp. (K)	Density (g/cm^3)	Volume (cm^3/mol)
1102	1.535	38.07
1102	1.540	37.95
1152	1.513	38.63
1152	1.516	38.55
1152	1.516	38.55
1204	1.488	39.28
1244	1.467	39.84
1310	1.429	40.88
1360	1.399	41.79
1455	1.364	42.86
1455	1.372	42.60

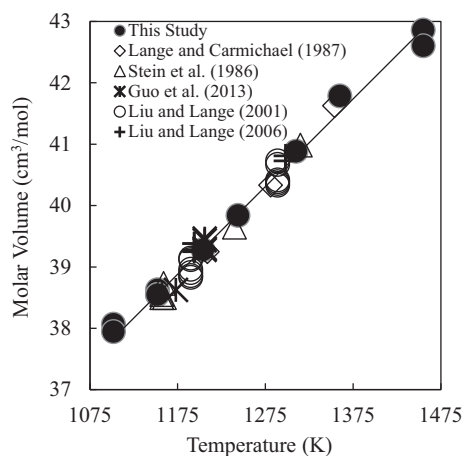


Fig. 1. Liquid NaCl molar volume data with temperature from this study and other double-bob Archimedean results from the literature.

Molar volumes were calculated by dividing the measured density into the gram formula weight (g.f.w.; defined in Table 1) of each sample. Molar volumes of all liquids are plotted as a function of temperature in Fig. 2. Ideally, each density result at each temperature should be based on the average of four ($n = 4$; Table 3) density determinations obtained from four bobs (two large and two small). However, during some experiments, one or two of the four bob measurements failed, often because the bob adhered to the crucible wall during immersion. In some cases, many more bob buoyancy measurements were made, enabling more density determinations, and in Table 3, the number of different density measurements at each experimental temperature ranges from one to 13. The standard deviation of the measured density for each sample ranges from 0.04 to 0.57%, and the average is 0.26%, which is identical to the measurement error of Liu and Lange (2003).

A linear fit of the molar volume with temperature was made for each sample liquid with the following equation:

$$V(T) = a_{1100\text{K}} + b(T - 1100\text{ K}) \quad (3)$$

where $V(T)$ is the molar volume of the liquid, a is the fitted molar volume of the liquid at 1100 K, b is the thermal expansivity $\frac{\partial V}{\partial T}$ and T is temperature in Kelvin. Table 4 gives the fitted values of a and b for all sample liquids (except T-35%Ca, for which molar volume was obtained at only one temperature; Table 3).

3.3. A linear model equation for molar volume

The molar volume data for the 15 liquids in Table 3 were combined with those from Liu and Lange (2003) on nine Li_2CO_3 - Na_2CO_3 - K_2CO_3 - CaCO_3 liquids in a fit to the following linear volume equation:

$$V^{liq}(X, T) = \sum X_i \left(\bar{V}_{i,1100\text{K}} + \frac{\partial \bar{V}_i}{\partial T} [T - 1100\text{ K}] \right). \quad (4)$$

In Eq. (4), V^{liq} is the measured molar volume, X_i is the mole fraction of each carbonate component i in each experimental liquid, $\bar{V}_{i,1100\text{K}}$ is the partial molar volume of each carbonate component at 1100 K, and $\frac{\partial \bar{V}_i}{\partial T}$ is the partial molar thermal expansivity of each carbonate component. Fitted values of \bar{V}_i and $\frac{\partial \bar{V}_i}{\partial T}$ for each carbonate component were obtained by multiple linear regression of Eq. (4) and reported in Table 5. The final adjusted R^2 of the fit is 0.98. The average residual is $\pm 0.32\%$, which is close to the average analytical error for density of 0.26%. Most importantly, the residuals show no systematic variation with temperature (Fig. 3) or composition (Fig. 4). The maximum residual of 0.89% is for a BaCO_3 -bearing liquid, which is not surprising given the experimental challenges with these liquids. The results from this study show that the molar volume of multicomponent carbonate liquids, including those with more than one alkaline earth carbonate component (e.g. T-50CaSrBa; Table 1), mix ideally with respect to composition in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 - Rb_2CO_3 - CaCO_3 - SrCO_3 - BaCO_3 system. Moreover, for the alkali carbonate components, (Rb_2CO_3 and Cs_2CO_3 in this study; Li_2CO_3 , Na_2CO_3 and K_2CO_3 in Liu and Lange, 2003), the fitted partial molar volumes (Table 5) match those for the pure end-members (Table 4).

3.4. Constant thermal expansion coefficients for alkali vs. alkaline earth carbonate liquids

A plot of the thermal expansion coefficient, $\alpha = (1/\bar{V}_i) (\partial \bar{V}_i / \partial T)$, for each liquid carbonate component at 1100 K is shown in Fig. 5 as a function of cation volume. The thermal expansion coefficient appears to be independent of composition, within error, among all five alkali carbonate liquids, with an average value of $22.5 (\pm 2.1) 10^{-5} \text{ K}^{-1}$, and among all three alkaline earth carbonate liquid components, with an average value of $16.4 (\pm 2.3) 10^{-5} \text{ K}^{-1}$.

3.5. Comparison to results based on the maximum bubble pressure method

The liquid density results from this study are compared to those on Li_2CO_3 - Na_2CO_3 - K_2CO_3 - Rb_2CO_3 - Cs_2CO_3 - CaCO_3 - SrCO_3 - BaCO_3 liquids from the literature based on the maximum bubble pressure technique (Kojima et al., 1999, 2003, 2008), with a reported error of 0.9%. Although the Kojima et al. density measurements were made on carbonate liquids in the same eight-component system, only two compositions from those three studies match those from this study, namely pure Rb_2CO_3 and Cs_2CO_3 . A direct comparison of the density results is made in Fig. 6; there is general agreement for Rb_2CO_3 liquid, but the density results for Cs_2CO_3 liquid from this study are $\sim 1.8\%$ higher than those measured by Kojima et al. (1999).

A global comparison of all the density results on the eight-component carbonate liquids from the maximum bubble pressure method, converted to molar volume, to those from the double-bob Archimedean method is made

Table 3
Experimental liquid density measurements and molar volumes.

Sample	Temp (K)	ρ (g/cm ³)	^a SD	^b N	V _{meas.} (cm ³ /g.f.w.)	^c V _{calc.} (cm ³ /g.f.w.)	^d % diff (meas-calc)
T-25%Ca	813	2.138	0.30	4	46.81	46.52	0.63
T-25%Ca	859	2.123	0.23	4	47.14	47.00	0.31
T-25%Ca	912	2.112	0.30	4	47.39	47.54	−0.32
T-25%Ca	1152	1.992	0.17	4	50.24	50.05	0.39
T-25%Ca	1257	1.948	0.25	4	51.38	51.14	0.47
T-35%Ca	1007	2.110	0.43	4	47.43	47.34	0.20
T-50%Ca	1007	2.201	0.21	2	45.47	45.54	−0.16
T-50%Ca	1054	2.181	0.13	4	45.89	45.98	−0.19
T-50%Ca	1102	2.164	0.25	1	46.25	46.42	−0.36
T-10%Sr	758	2.170	0.08	4	48.31	47.98	0.69
T-10%Sr	859	2.123	0.31	2	49.38	49.11	0.55
T-10%Sr	1257	1.966	0.29	10	53.32	53.52	−0.37
T-20%Sr	1102	2.141	0.16	2	51.19	50.98	0.41
T-20%Sr	1257	2.071	0.30	4	52.92	52.61	0.58
T-20%Sr	1455	2.002	0.20	4	54.74	54.71	0.05
T-30%Sr	960	2.351	0.31	4	48.64	48.72	−0.17
T-30%Sr	1054	2.305	0.42	4	49.61	49.66	−0.12
T-30%Sr	1152	2.248	0.32	2	50.87	50.65	0.42
T-50%Sr	1152	2.524	0.29	2	49.04	48.97	0.15
T-50%Sr	1152	2.540	0.10	4	48.74	48.97	−0.48
T-50%Sr	1257	2.489	0.17	4	49.73	49.91	−0.36
T-50%Sr	1362	2.433	0.47	4	50.88	50.87	0.02
T-50%Sr	1455	2.390	0.09	4	51.79	51.71	0.17
T-15%Ba	850	2.316	0.16	4	49.51	49.56	−0.11
T-15%Ba	960	2.253	0.36	4	50.90	50.80	0.18
T-35%Ba	960	2.643	0.24	7	50.75	50.56	0.36
T-35%Ba	1102	2.553	0.46	7	52.53	52.07	0.89
T-35%Ba	1152	2.546	0.57	13	52.68	52.60	0.15
T-35%Ba	1204	2.534	0.24	4	52.93	53.14	−0.41
T-35%Ba	1257	2.503	0.56	4	53.58	53.70	−0.22
T-50%Ba	1102	2.887	0.05	4	51.51	51.83	−0.61
T-50%Ba	1127	2.852	0.47	4	52.14	52.08	0.13
T-50%Ba	1152	2.855	0.07	2	52.09	52.33	−0.47
T-50%Ba	1257	2.769	0.08	4	53.71	53.39	0.59
T-50%Rb	865	2.618	0.08	4	63.22	63.22	0.01
T-50%Rb	885	2.621	0.32	4	63.15	63.51	−0.57
T-50%Rb	982	2.537	0.17	4	65.24	64.95	0.45
T-50%Rb	1078	2.491	0.07	4	66.44	66.37	0.11
T-50%Rb	1152	2.453	0.42	4	67.47	67.47	0.01
T-50%Rb	1112	2.476	0.20	4	66.85	66.87	−0.04
Rb ₂ CO ₃	1152	2.819	0.23	2	81.92	81.72	0.25
Rb ₂ CO ₃	1257	2.769	0.23	3	83.40	83.60	−0.24
T-50%CaSrBa	1030	2.587	0.20	4	48.02	48.23	−0.44
T-50%CaSrBa	1080	2.556	0.00	1	48.60	48.70	−0.21
T-50%Cs	859	3.085	0.45	4	69.03	69.22	−0.28
T-50%Cs	960	3.009	0.56	4	70.77	70.93	−0.23
T-50%Cs	1112	2.881	0.54	2	73.92	73.51	0.55
Cs ₂ CO ₃	1112	3.461	0.07	4	94.14	94.27	−0.14
Cs ₂ CO ₃	1122	3.438	0.30	4	94.77	94.49	0.29
Cs ₂ CO ₃	1127	3.433	0.39	4	94.91	94.60	0.32
Cs ₂ CO ₃	1163	3.419	0.21	4	95.30	95.41	−0.11
Cs ₂ CO ₃	1214	3.388	0.04	4	96.17	96.54	−0.39

^a SD = 1 – sigma standard deviation (%) on each density determination.

^b N = number of separate density measurements.

^c V_{calc} = calculated from Eq. (4) and fitted terms in Table 5.

^d % diff = 100 * (V_{meas} – V_{calc})/V_{meas}.

by comparing the results of a fit to the model equation in Eq. (4). In Table 6, the fitted values for the partial molar volume and thermal expansivity of each of the carbonate components, based on the density data from Kojima et al. (1999, 2003, 2008), are shown. A comparison of the

results in Tables 5 and 6 shows that the largest differences are seen for the alkaline earth carbonate components, where the errors on the fitted terms based on the maximum bubble pressure method (Table 6) are larger by a factor of two than those based on the double-bob method (Table 5).

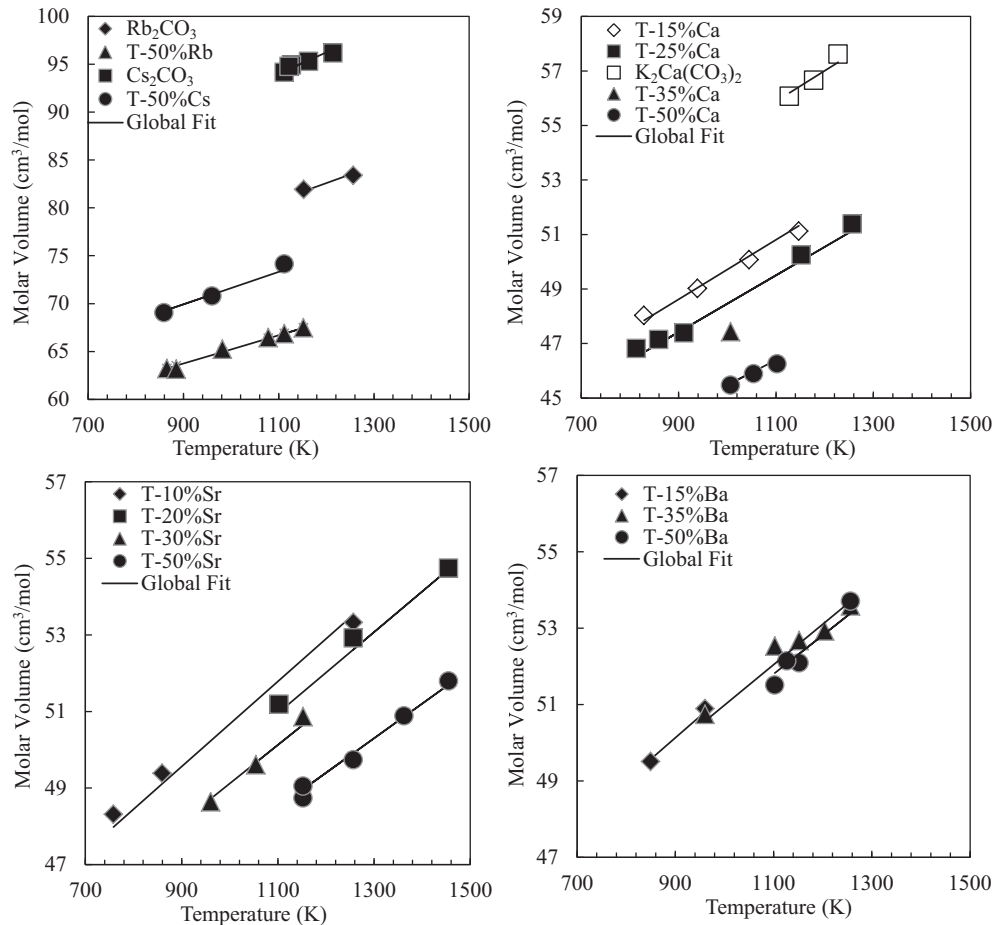


Fig. 2. Molar volumes of experimental liquids with temperature from this study (solid symbols) and Liu and Lange (2003) (open symbols). The lines are the global fit to the entire data set (Eq. (4) and Table 5).

Table 4
Linear fit to volume of each experimental liquid.

$V = a_{1100K} + b(T - 1100 \text{ K}) \text{ cm}^3/\text{g.f.w.}$			
Sample	a_{1100K}	$b \cdot 10^3$	Temp. range (K)
T-25%Ca	49.67	10.62	813–1257
T-35%Ca	—	—	1007
T-50%Ca	46.24	8.16	1007–1102
T-10%Sr	51.76	10.01	758–1257
T-20%Sr	51.23	10.04	1102–1455
T-30%Sr	50.22	11.61	960–1152
T-50%Sr	48.35	9.61	1152–1455
T-15%Ba	52.64	12.50	850–960
T-35%Ba	52.17	9.13	960–1257
T-50%Ba	51.55	13.63	1102–1257
T-50%Rb	66.74	15.36	1152–1257
Rb ₂ CO ₃	81.18	14.16	865–1152
T-50%CaSrBa	48.83	11.65	1030–1080
T-50%Cs	73.63	19.43	859–1112
Cs ₂ CO ₃	94.24	17.24	1112–1214

This is consistent with the larger error on density for the maximum bubble pressure method ($\sim 0.9\%$; Kojima et al., 1999, 2003, 2008) compared to that for the double-bob

Table 5
Fitted partial molar volumes (Eq. (4)).

$V = V_{i(1100K)} + \delta V_i/\delta T (T - 1100 \text{ K})$		
Component	$V_{i(1100K)} (\text{cm}^3/\text{mol})$	$\delta V_i/\delta T (10^{-3} \text{ cm}^3/\text{mol K})$
Li ₂ CO ₃	41.22 ± 0.09	8.83 ± 1.20
Na ₂ CO ₃	53.27 ± 0.11	12.30 ± 1.11
K ₂ CO ₃	71.59 ± 0.13	15.68 ± 1.13
Rb ₂ CO ₃	80.78 ± 0.11	17.98 ± 0.90
Cs ₂ CO ₃	94.00 ± 0.09	22.28 ± 1.14
CaCO ₃	40.18 ± 0.16	6.75 ± 1.52
SrCO ₃	44.33 ± 0.22	6.40 ± 1.20
BaCO ₃	50.99 ± 0.19	8.65 ± 1.90

Calibration data set: this study and Liu and Lange (2003).

method ($\sim 0.3\%$; Liu and Lange, 2003; this study). Therefore, although the two sets of density measurements obtained by different techniques are broadly consistent, they were not combined into a final data set. The recommended values to use in the molar volume equation (Eq. (4)) are those in Table 5 and based solely on Pt double-bob Archimedean density measurements.

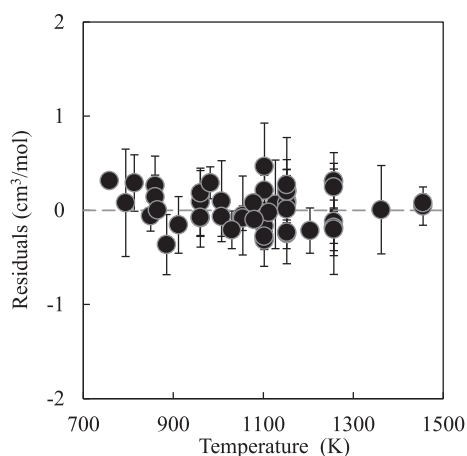


Fig. 3. Residuals from fit to Eq. (4) as a function of temperature. Two-sigma error bars are shown.

4. DISCUSSION

4.1. A comparison of crystalline and liquid molar volumes with temperature

The newly calibrated volume equation (Eq. (4)), with the fitted parameters in Table 5, can be used to calculate the molar volume of each liquid carbonate component with temperature and compare it to their crystalline molar volume with temperature, for phases where these data are available from X-ray and neutron diffraction studies. This allows a direct comparison to be made between crystal and liquid volumes over a range of temperatures, and it shows how changes in crystal structure and cation coordination affect molar volumes. In Table 7, crystalline volumes for each of the eight carbonates (and their polymorphs) at both 298 and 1100 K have been collected from the literature and compared to their respective liquid volumes at 1100 K.

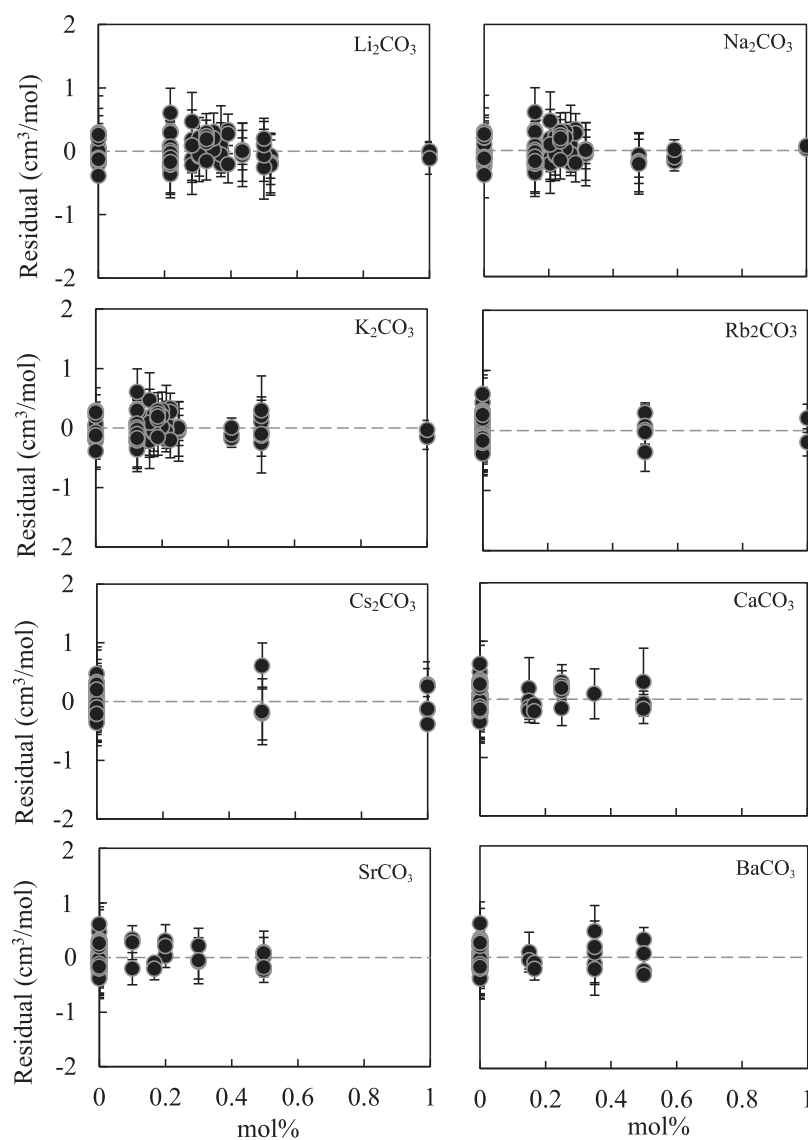


Fig. 4. Residuals from fit to Eq. (4) as a function of composition. Two-sigma error bars are shown.

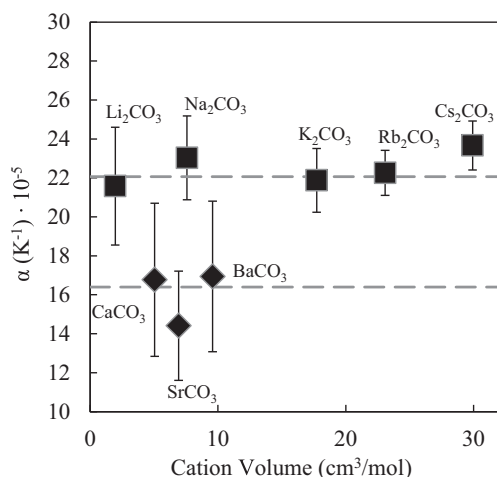


Fig. 5. Thermal expansion coefficient derived for each carbonate liquid component at 1100 K versus cation volume. Dashed lines are average for alkali carbonate liquids ($22.07 \cdot 10^{-5} \text{ K}^{-1}$) and the alkaline earth carbonate liquids ($16.40 \cdot 10^{-5} \text{ K}^{-1}$).

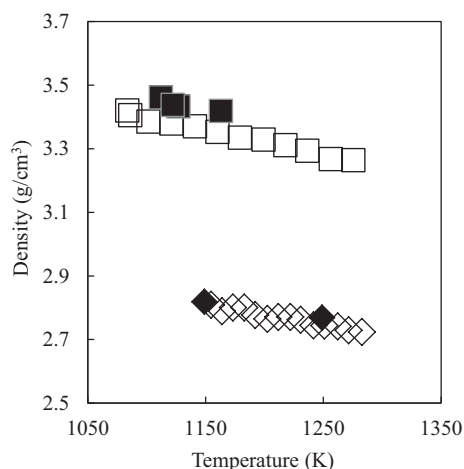


Fig. 6. Liquid density measurements for Rb_2CO_3 (diamonds) and Cs_2CO_3 (squares) with temperature (solid symbols, this study; open symbols, Kojima et al., 1999).

Table 6

Fit of Eq. (4) on data from Kojima et al. (1999, 2003, 2008).

$V = V_{\ell(1100\text{K})} + \delta V_{\ell}/\delta T (T - 1100 \text{ K})$		
Component	$V_{\ell(1100\text{K})} (\text{cm}^3/\text{mol})$	$\delta V_{\ell}/\delta T (10^{-3} \text{ cm}^3/\text{mol K})$
Li_2CO_3	41.58 ± 0.11	8.41 ± 0.83
Na_2CO_3	53.25 ± 0.14	12.48 ± 1.14
K_2CO_3	72.01 ± 0.22	16.55 ± 1.53
Rb_2CO_3	81.47 ± 0.17	19.33 ± 1.22
Cs_2CO_3	95.69 ± 0.17	22.61 ± 1.17
CaCO_3	39.14 ± 0.40	7.35 ± 3.41
SrCO_3	43.12 ± 0.39	11.84 ± 3.84
BaCO_3	48.40 ± 0.40	9.09 ± 3.34

For crystalline Li_2CO_3 , a volume equation from Liu and Lange (2003) (citing data from A. Kirfel and B. Barbier) has been used (T is in Celsius):

$$V(T)_{\text{Li}_2\text{CO}_3} = 35.2804 + 1.510 \times 10^{-3} T + 1.440 \times 10^{-6} T^2 \quad \text{cm}^3/\text{mol} \quad (5)$$

Among the alkali carbonates, a comparison of crystalline volumes at 298 K to the liquid at 1100 K reveals a systematic decrease in the liquid to crystal volume ratio ($V_{\text{liq}}/V_{\text{xtl}}$) from Na_2CO_3 (1.28) to Cs_2CO_3 (1.22); the exception to this trend is Li_2CO_3 with a $V_{\text{liq}}/V_{\text{xtl}}$ ratio of 1.17. Among the alkaline earth carbonates, the similar trend is observed for the orthorhombic $Pm\bar{c}n$ phases, namely a systematic decline in $V_{\text{liq}}/V_{\text{xtl}}$ from CaCO_3 (1.18) to BaCO_3 (1.12).

At 1100 K, the crystalline alkali carbonates have all transitioned to a hexagonal structure (usually $P6_3/mmc$) and in all cases the $V_{\text{liq}}/V_{\text{xtl}}$ ratio is ~ 1.10 . For CaCO_3 , SrCO_3 and BaCO_3 , the $R3m$ or $R\bar{3}m$ phases are stable near 1100 K. CaCO_3 undergoes a phase transition from $R\bar{3}c$ to $R\bar{3}m$ at ~ 1240 K (e.g. Antao and Hassan, 2009; Ishizawa et al., 2013) whereas SrCO_3 and BaCO_3 undergo a phase change to a hexagonal structure ($R3m$ or $R\bar{3}m$ at ~ 1173 and ~ 1073 K, respectively). The orthorhombic to hexagonal phase change is accompanied by a marked increase in volume (4.7%, 3.8% and 2.7% for CaCO_3 , SrCO_3 and BaCO_3 , respectively). The alkaline earth carbonates range to lower $V_{\text{liq}}/V_{\text{xtl}}$ ratios at 1100 K (≥ 1.04) than the 1.10 ratio for all the alkali carbonates. This suggests that the hexagonal alkaline earth carbonate structures (i.e. $R3m$ or $R\bar{3}m$) are more similar to the average liquid structure than are the hexagonal alkali carbonate structures (i.e. $P6_3/mmc$).

Using the General Lattice Utility Program (Gale and Rohl, 2003), pair distribution functions (Cope and Dove, 2007) were calculated in this study for a subset of the crystalline structures for Li_2CO_3 , Na_2CO_3 , K_2CO_3 , CaCO_3 and BaCO_3 . The objective was to obtain coordination numbers for carbonate-carbonate (C-C), metal cation-carbonate (M-C) and metal cation-oxygen (M-O) pairs. These coordination numbers are listed in Table 8 and compared with those calculated for the respective liquid carbonates (Tissen and Janssen, 1990; Vuilleumier et al., 2014).

The coordination of the metal cation with both oxygen and carbonate for all eight crystalline structures are similar to those for the corresponding liquids (Table 8). Instead, the primary structural difference between the crystals and the corresponding liquids is in the carbonate-carbonate (C-C) coordination number, which is 12-fold in all of the carbonate liquids (Tissen and Janssen, 1990; Vuilleumier et al., 2014). The low-temperature monoclinic structures for the crystalline alkali carbonates and orthorhombic structures for the crystalline alkaline-earth carbonates all have C-C coordination numbers of 1 or 2. At high temperature the alkali carbonate structure (e.g., $P6_3/mmc$ for Na_2CO_3) has a C-C coordination number of ~ 6 whereas the alkaline-earth carbonate structure (e.g., $R\bar{3}m$) has a C-C coordination number of ~ 12 , similar to that of the liquid. At even higher temperature (>1250 K), the structure for BaCO_3 becomes cubic ($Fm\bar{3}m$; Strømme, 1975; Nie et al., 2017) and the C-C coordination of 12 is maintained. Thus with increasing temperature, the crystalline phase of BaCO_3

Table 7
Partial molar liquid volumes and crystalline volumes (with space groups).

	V_{1100K}^{liq} (cm ³ /mol)	V_{298K}^{xtl} (cm ³ /mol)	V_{liq}/V_{xtl} 298 K	V_{1100K}^{xtl} (cm ³ /mol)	V_{liq}/V_{xtl} 1100 K
Li ₂ CO ₃	41.22	35.20 C2/c ^a	1.17	37.03 C2/c ^b	1.11
Na ₂ CO ₃	53.27	41.78 C2/m ^c	1.28	48.24 P6 ₃ /mmc ^{d,e}	1.10
K ₂ CO ₃	71.59	56.58 P2 ₁ /c ^f	1.27	65.0 P6 ₃ /mmc ^{g,h,i}	1.10
Rb ₂ CO ₃	80.78	65.04 P2 ₁ /c ^j	1.24	73.7 P6 ₃ /mmc ^{k,*}	1.10
Cs ₂ CO ₃	94.00	76.92 P2 ₁ /c ^j	1.22	85.8 P6 ₃ /mmc ^{k,*}	1.10
CaCO ₃	40.18	34.18 Pmcn ^k	1.18	36.1 Pmcn ^k	1.11
		36.90 R 3 m ^l	1.09	37.75 R 3 m ^{m,n}	1.06
SrCO ₃	44.33	38.98 Pmcn ^o	1.14	41.04 Pmcn ^o	1.08
				42.75 R 3 m ^{p,q}	1.04
BaCO ₃	50.99	45.73 Pmcn ^{o,r}	1.12	48.22 Pmcn ^{o,r}	1.06
				49.66 R 3 m ^{o,p}	1.03
				49.70 Fm3m ^o	1.03

^a Effenberger and Zemmann (1979).

^b Liu and Lange (2003).

^c Dusek et al. (2003).

^d Swaison et al. (1995).

^e Brouns and Visser (1964).

^f Gatehouse and Lloyd (1973).

^g Schneider and Levin (1973).

^h Becht and Struikmans (1976).

ⁱ Dinnebier et al. (2005).

^j Ehrhardt et al. (1980).

^k Antao and Hassan (2010).

^l Markgraf and Reeder (1985).

^m Dove and Powell (1989).

ⁿ Antao and Hassan (2009).

^o Nie et al. (2017).

^p Lander (1949).

^q Moreau (1964).

^r Antao and Hassan (2007).

* Value extrapolated up to 1100 K from high-temperature data.

has a structure that is increasingly similar to the average structure of its corresponding liquid. The fact that all of the alkali and alkaline-earth carbonate liquids, irrespective of their M-O and M-C coordination numbers, have a shared and constant C-C coordination number of 12 may explain why the alkali and alkaline-earth liquid carbonate components all mix ideally with each other with respect to volume.

4.2. Systematic trends of carbonate liquid volume with cation coordination

The partial molar volumes (at 1100 K) of the liquid carbonate components are plotted as a function of cation field strength in Fig. 7. Cation field strength is defined as $z/(0.126 + r)^2$, where z is the formal cation charge, r is the crystalline ionic radius of the cation (Shannon and Prewitt, 1969; Shannon, 1976) and the 0.126 term is the ionic radius of the oxygen anion. The ionic radius of the cation depends on its coordination number with oxygen (M-O coordination number). For example, for alkali carbonate liquids, molecular dynamic simulations of Li₂CO₃, Na₂CO₃ and K₂CO₃ point to an oxygen coordination of ~4, ~5 and ~6 for Li⁺, Na⁺ and K⁺, respectively (e.g. Habasak, 1990; Tissen and Janssen, 1990; Roest et al., 2017). The ~4-fold coordination of Li-O in Li₂CO₃ liquid

has been corroborated by neutron diffraction and XRD studies (Kohara et al., 1999). For alkaline earth carbonate liquids, molecular dynamic simulations show an average oxygen coordination of 7-fold for Ca²⁺ and Sr²⁺ and 8-fold for Ba²⁺ at 1 bar (Hurt, 2018). These results are consistent with first principle molecular dynamics simulations of CaCO₃ melt (Vuilleumier et al., 2014), which shows Ca²⁺ cations are 7- to 8-fold coordinated by oxygen at the higher pressure of ~0.5 GPa. Collectively, these 1-bar M-O coordination numbers for the alkali and alkaline earth metals are used to calculate their respective ionic radii and cation field strength (Table 9).

Plots of the liquid partial molar volume of each carbonate component as a function of field strength (Fig. 7) show two distinct trends, one for the alkali carbonate components (Li₂CO₃, Na₂CO₃, K₂CO₃, Rb₂CO₃ and Cs₂CO₃) and another for the alkaline earth carbonate components (CaCO₃, SrCO₃ and BaCO₃). The trend is curved for the alkali carbonate liquids, which cautions against a linear extrapolation of the alkaline earth carbonate trend to estimate the liquid partial molar volume for the MgCO₃ and FeCO₃ components.

In a second plot (Fig. 8), the partial molar volumes of all carbonate components are plotted as a function of cation volume (cm³/mol), which is defined as $AvN/4\pi r^3/3$ where Av is Avogadro's number, N is the stoichiometric number

Table 8

Coordination numbers of metal cations (i.e. Li, Na, K, Rb, Cs, Ca, Sr, Ba) with oxygen atoms (M-O), metal cations with carbonate groups (M-C) and carbonate groups with other carbonate groups (C-C) for low- and high-temperature crystal phases are compared to coordination numbers in their respective liquids.

	C-C	M-O	M-C
Li ₂ CO ₃ C2/c ^a	2	4	4
Li ₂ CO ₃ liquid ^b	12.5	3.7	3.8
Na ₂ CO ₃ C2/m ^c	1	4.5	4
Na ₂ CO ₃ P6 ₃ /mmc ^d	5.9	6	5.5
Na ₂ CO ₃ liquid ^b	12.6	5.2	4.2
K ₂ CO ₃ P2 ₁ /c ^e	2	4.5	5
K ₂ CO ₃ Liquid ^b	12.0	6.0	4.4
CaCO ₃ Pmcn ^f	2	9	6
CaCO ₃ R3c ^g	6	6	6
CaCO ₃ R3m ^h	12	–	6
CaCO ₃ liquid ⁱ	12	7.8	6
SrCO ₃ R3m ^j	12	–	6
BaCO ₃ R3m ^k	12	9	6
BaCO ₃ Fm3m ^j	12	–	6

^a Effenberger and Zemmann (1979).

^b Tissen and Janssen (1990).

^c Arakcheeva et al. (2010).

^d Swaison et al. (1995).

^e Idemoto et al. (1998).

^f Antao and Hassan (2010).

^g Markgraf and Reeder (1985).

^h Ishizawa et al. (2013).

ⁱ Vuilleumier et al. (2014).

^j Strømme (1975).

^k Antao and Hassan (2007).

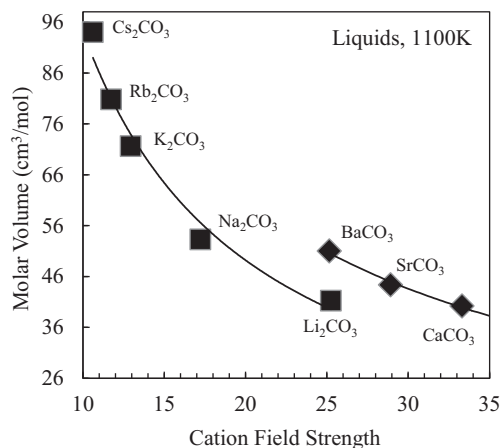


Fig. 7. Fitted liquid partial molar volume of each carbonate component at 1100 K (Table 5) as a function of cation field strength.

of cations per formula unit (2 for alkalis and 1 for alkaline earths) and r is the crystalline ionic radius. The results lead to two sub-parallel linear trends, with the alkali carbonate liquids falling on one line, and the alkaline-earth carbonate liquids on the other line. A linear fit to each set of liquids (Fig. 8) leads to the following equations, both with $R^2 = 0.999$:

$$V_{\text{carbonate(alkali)}} = 1.86(\pm 0.03) \cdot V_{\text{cation}} + 38.3(\pm 0.6) \text{ cm}^3/\text{mol} \quad (6a)$$

$$V_{\text{carbonate(alkalineearth)}} = 2.08(\pm 0.01) \cdot V_{\text{cation}} + 31.0(\pm 0.1) \text{ cm}^3/\text{mol} \quad (6b)$$

In both Eqs. (6a) and (6b), $V_{\text{carbonate}}$ is the molar volume of the carbonate liquid and V_{cation} is the cation volume as defined above. Table 9 lists the calculated cation volumes and field strength for each component in addition to the M-O coordination number and ionic radii. The slope intercepts of $38.3(\pm 0.6)$ and $31.0(\pm 0.1)$ cm³/mol are fitted values ($\pm 1\sigma$) for the partial molar volume of the carbonate ion, (CO₃)²⁻, at 1100 K.

Importantly, the larger partial molar volume of (CO₃)²⁻ in the alkali vs. alkaline earth carbonate liquids is not caused by a larger size for the carbonate ion. This is demonstrably not the case, as C–O distances (a proxy for the size of the carbonate molecule) in alkali carbonate liquids are not appreciably larger than in alkaline earth carbonate liquids. For example, Vuilleumier et al. (2014) and Hurt (2018) report C–O bond lengths in alkaline earth carbonate liquids of 1.29 Å (which, as Vuilleumier points out, is essentially the same as in calcite and aragonite). Roest et al. (2017), in their MD simulations, fixed C–O distances at 1.27 Å for alkali carbonate liquids. Instead, the difference in the partial molar volume of the carbonate ion reflects a systematic difference in the coordination of the various metal cations to the carbonate ions (M-C coordination number). For example, in the alkali carbonate liquids, the alkali metals are in a four-fold coordination with carbonate ions (e.g., Habasaki, 1990; Tissen and Janssen, 1990; Kohara et al., 1999; Roest et al., 2017), whereas in the alkaline earth carbonate liquids, the alkaline earth metals are all in a six-fold coordination with carbonate ions (Vuilleumier et al., 2014; Hurt, 2018). When the M-C coordination is four, this leads to a more open arrangement of the carbonate ions, and thus a larger partial molar volume (~ 38.3 cm³/mole at 1100 K), in comparison to the case when the M-C coordination is six, leading to a more compact topology, and thus a smaller partial molar volume (~ 31.0 cm³/mole at 1100 K).

An analogy can be drawn with 1-bar crystalline polymorphs of SiO₂, all of which are based on silica tetrahedra (Si⁴⁺ that is 4-fold coordinated by oxygen). The size of the silica tetrahedra does not change among the crystalline polymorphs (e.g., Stixrude and Bukowinski, 1988), but the way the tetrahedra are arranged into different ring structures (i.e., network topology) varies significantly (e.g., 6-membered rings in tridymite; 6- and 3-membered rings in quartz), leading to very different molar volumes at similar T-P conditions (e.g., Stixrude and Bukowinski, 1990).

4.3. Estimates of the partial molar volume of the liquid MgCO₃ and FeCO₃ components

If it is assumed that both Mg²⁺ and Fe²⁺ are six-fold coordinated by oxygen (slightly lower than the 7-fold coordination inferred for Ca²⁺, given their smaller size) and six-

Table 9

Crystal radius, cation volume and cation field strength for each carbonate component.

Component	M-O coordination	Crystal radius (pm)	Cation volume (cm ³ /mol)	Cation field strength
Li ₂ CO ₃	4 ^a	73	1.96	25.3
Na ₂ CO ₃	5 ^a	115	7.58	17.2
K ₂ CO ₃	6 ^a	152	17.7	12.9
Rb ₂ CO ₃	6	166	23.1	11.7
Cs ₂ CO ₃	6	181	29.9	10.6
MgCO ₃ ^c	6	86	1.60	44.5
	4	71	0.90	51.5
CaCO ₃	7 ^b	119	4.39	33.3
SrCO ₃	7	137	6.36	28.9
BaCO ₃	8	156	9.58	25.2
FeCO ₃ ^c	6	92	1.96	42.1
	4	77	1.15	48.5

^a Tissen and Janssen (1990).

^b Vuilleumier et al. (2014).

^c Two sets of values are shown depending on an M-O coordination of 6 or 4.

fold coordinated by carbonate (similar to the case for Ca²⁺, Sr²⁺ and Ba²⁺), their cation volumes can be calculated and input into Eq. (6b), resulting in partial molar volumes of 34.4 ± 0.1 and 35.1 ± 0.1 cm³/mol (Fig. 9), respectively, for the liquid MgCO₃ and FeCO₃ components. Note that if Mg²⁺ and/or Fe²⁺ are 4-fold (vs. 6-fold) coordinated by oxygen (similar to the case for Li⁺ relative to K⁺) this leads to smaller ionic radii, smaller cation volumes and smaller partial molar carbonate volumes. However, if Mg²⁺ and/or Fe²⁺ are 4-fold (vs. 6-fold) coordinated by carbonate, then the equation to use is Eq. (6a) (not (6b)), which leads to larger calculated partial molar volumes (40.1 ± 0.1 and 40.4 ± 0.1 cm³/mol for MgCO₃ and FeCO₃ liquid respectively).

4.4. Comparison to estimates of liquid V_{MgCO_3} and V_{FeCO_3} from the literature

Dobson et al. (1996) measured the density of K₂Mg(CO₃)₂ liquid density at 773 and 837 K, using the single double-bob Archimedean method, using a corundum bob (~0.5 g) with a Pd wire stem (0.1 mm). They used the same approach to measure the density of K₂Ca(CO₃)₂ liquid between 1115 and 1249 K, which was directly compared to Pt double-bob density measurements on the same liquid in Liu and Lange (2003). There is a difference of ~3% in the results, with the corundum single-bob method leading to lower density determinations than those from the Pt double-bob method (Liu and Lange, 2003). Error analysis shows that because of the small mass of the corundum bob (~0.5 g) relative to that for Pt bobs (≤26 g), the estimated uncertainty in density using the corundum bob is ~2.5% (Liu and Lange, 2003). With this caveat in mind, an estimate of the partial molar volume of MgCO₃ obtained from the two Dobson et al. (1996) density measurements on K₂Mg(CO₃)₂ liquid can be made by incorporating them one by one into a fit of Eq. (4), combined with

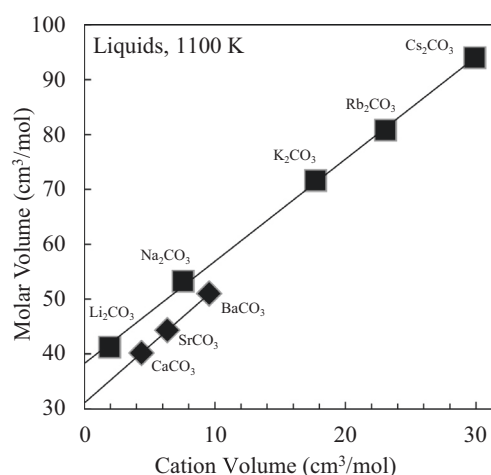


Fig. 8. Fitted liquid partial molar volume of each carbonate component at 1100 K (Table 5) as a function of cation field strength as a function of cation volume.

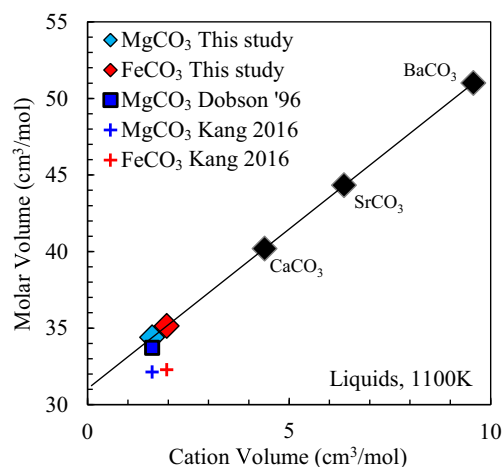


Fig. 9. Fitted liquid partial molar volumes of alkaline earth carbonate components at 1100 K (Table 5) as a function of cation volume together with estimates of liquid partial molar volume of MgCO₃ from data of Dobson et al. (1996) and Kang et al. (2016) and FeCO₃ from data of Kang et al. (2016). The cation volumes for Mg²⁺ and Fe²⁺ assume 6-fold M-O coordination.

all the Pt double-bob density results from this study and Liu and Lange (2003). The results lead to estimated \bar{V}_{MgCO_3} values of 31.87 and 32.31 cm³/mole at 773 and 837 K, respectively. Use of the thermal expansion coefficient of $16.4 (\pm 2.3) \cdot 10^{-5} \text{ K}^{-1}$ (i.e., that for the alkaline-earth carbonate liquids; Fig. 5) leads to a molar volume of 33.7 cm³/mol at 1100 K which overlaps within $\pm 2.5\%$ uncertainty the estimated value for V_{MgCO_3} on the basis of Eq. (6b) (based on 6-fold M-O and M-C coordination). In contrast, this relatively low V_{MgCO_3} (33.7 cm³/mol) derived from Dobson's experiments is inconsistent with estimates calculated by Eq. (6a) (40.1 ± 0.1 cm³/mol, which assumes a 4-fold M-O and M-C coordination). However, such an inconsistency does not rule out the possibility of 4-fold M-O and M-C coordination in MgCO₃ liquid since the composition explored by Dobson, K₂Mg(CO₃)₂, is

highly unusual. It is both the only MgCO_3 -bearing liquid that is stable at 1 bar (e.g., Ragone et al., 1966) and the only purely carbonate liquid that quenches to a glass (e.g., Datta et al. 1964; Ragone et al. 1966). The unique properties of $\text{K}_2\text{Mg}(\text{CO}_3)_2$ may be a reflection of a unique liquid structure that is generally unrepresentative of the MgCO_3 liquid component in natural systems. This is exemplified in Mg-bearing silicate liquids where the average oxygen coordination of Mg^{2+} depends on the field strength of competing modifier cations (e.g., Ca^{2+} , Li^+ , Na^+ , K^+) in the melt (Shimoda et al., 2008).

The molar volumes and thermal expansion coefficients of liquid MgCO_3 and FeCO_3 have also been estimated by Kang et al. (2015, 2016) respectively, by fitting thermodynamic properties to recover the fusion curves of FeCO_3 and MgCO_3 . Their fitted values for the molar volume of liquid MgCO_3 and FeCO_3 (calculated at 1100 K) are 32.1 and 32.3 cm^3/mol , respectively, which are lower than those estimated from Eq. (6b) and may be indirect evidence that the Mg^{2+} and Fe^{2+} cations are in a lower than 6-fold M-O coordination in carbonate liquids. However, the thermal expansion coefficients for the MgCO_3 and FeCO_3 liquid components in Kang et al. (2016) are those of the crystalline magnesite and siderite phases, which are nearly three times smaller than the average ($\sim 16.4 \cdot 10^{-5} \text{ K}^{-1}$) obtained in this study at 1100 K for the alkaline earth carbonate liquid components. The liquid phase is expected to have a higher thermal expansion coefficient than corresponding mineral phases, because of configurational contributions to thermal expansion that is related to the fact that liquids undergo continual breaking and re-forming of bonds, which permits access to different topologies and modes of expansion that are unavailable to solids (e.g., Richet and Neuville, 1992; Lange, 1997; Toplis and Richet, 2000). To resolve this issue, high-quality computational simulations of MgCO_3 and FeCO_3 liquids are needed to constrain their M-O and M-C coordination numbers.

4.5. Importance of accurate liquid \bar{V}_{MgCO_3} and \bar{V}_{FeCO_3} values for applications to the Earth

As illustrated in the above two sections, estimates of \bar{V}_{MgCO_3} and \bar{V}_{FeCO_3} are highly dependent on M-C and M-O coordination. If M-O and M-C are in 4-fold coordination the estimated molar volumes of MgCO_3 and FeCO_3 liquid are ~ 16 – 17% higher than if the M-O and M-C pairs are in 6-fold (40.1 versus 34.4 cm^3/mol for MgCO_3 liquid and 40.4 versus 35.1 cm^3/mol for FeCO_3 liquid, at 1 bar and 1100 K). This corresponds to a 14% decrease in MgCO_3 liquid density (2.45 versus 2.10 g/cm^3) and a 13% decrease in FeCO_3 liquid density (3.30 versus 2.87 g/cm^3) if M-O and M-C are in four-fold coordination compared to six-fold. Such a density contrast could hold important implications for the buoyancy of Mg and Fe-rich carbonatites in the mantle, and thus for long term storage of carbon in the mantle. This is especially true for ferrocarnatites that, according to some workers (i.e., Kang and Schmidt, 2017), could be dense enough to stagnate or sink in some parts of the upper mantle. While the magnitude of the 1-bar density contrast (13%) in FeCO_3 liquid with Fe^{2+} in 4-fold vs. 6-fold coor-

dination probably diminishes with pressure, it may nonetheless be sufficiently large that it leads to a significant barrier to FeCO_3 -rich liquids from achieving neutral or negative buoyancy in the upper-mantle. Knowledge of the liquid compressibility of these two components, therefore, and how it changes with pressure, is required before this issue can be fully explored and resolved.

It is also possible that the actual coordination of Mg^{2+} and Fe^{2+} is a non-integer value intermediate between 4- and 6-fold. The key point is that there is a need for high-quality computational simulations of MgCO_3 and FeCO_3 liquids to constrain their M-O and M-C coordination numbers, which will allow more accurate estimates of their respective liquid partial molar volumes from the linear trends in Fig. 8.

5. CONCLUSIONS AND FUTURE WORK

The results from this study show that multicomponent carbonate liquids mix ideally with respect to volume in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 - Rb_2CO_3 - Cs_2CO_3 - CaCO_3 - SrCO_3 - BaCO_3 system at 1 bar. Tight constraints on the partial molar volumes of the alkaline-earth carbonate components (CaCO_3 , SrCO_3 and BaCO_3), which are not stable as pure liquids at one bar, are obtained by adding up to 50 mol% of each to the eutectic in the Li_2CO_3 - Na_2CO_3 - K_2CO_3 system, leading to multicomponent liquids that are stable at one bar over a wide range of temperature (758–1455 K). An area of future work is to extend this strategy to MgCO_3 - and FeCO_3 -bearing alkali carbonate liquids and to other thermodynamic property measurements, including 1-bar sound speeds to obtain melt compressibility (e.g., O'Leary et al., 2015) and 1-bar drop calorimetry to obtain the partial molar liquid enthalpies, heats of mixing, and partial molar liquid heat capacities for multicomponent carbonate liquids that include the alkaline-earth carbonate components.

Another key conclusion is that the partial molar volume of the carbonate ion depends on its coordination with the various alkali- and alkaline-earth metals. In this study, molecular dynamic simulations from the literature were used to constrain the oxygen coordination around the alkali and alkaline-earth metals in carbonate liquids, which controls the size of the ionic radius and thus the cation volume for each metal. When the liquid partial molar volumes of each respective carbonate component are plotted against their cation volumes, two distinct linear trends ($R^2 = 0.999$) result. The slope intercepts reflect two different partial molar volumes (and thermal expansion coefficients) for the carbonate ion, depending on its coordination with the respective alkali and alkaline-earth metals (i.e., the M-C coordination number). When the M-C is four-fold, both the partial molar volume ($\sim 38 \text{ cm}^3/\text{mol}$) and its thermal expansion coefficient ($\sim 22.1 \cdot 10^{-5} \text{ K}^{-1}$) are larger than when the M-C is six-fold ($\sim 31 \text{ cm}^3/\text{mol}$ and $\sim 16.4 \cdot 10^{-5} \text{ K}^{-1}$, respectively). Fortunately, these differences do not preclude linear (ideal) mixing with respect to volume as long as the model equation employs carbonate components (and not oxide components with the carbonate ion). The reason why the liquid partial molar volumes mix ideally

at one bar is because all of the liquid carbonate components share the same carbonate-carbonate coordination number of 12, irrespective of their M-C and M-O coordination numbers.

The insights obtained in this study, when computational constraints on M-O and M-C coordination numbers are combined with macroscopic density measurements, show that accurate estimates of the partial molar volume and thermal expansion coefficient of the MgCO_3 and FeCO_3 components are possible if the M-C and M-O coordination numbers for Mg^{2+} and Fe^{2+} in carbonate melts can be obtained. This underscores the need for high-quality molecular dynamic simulations on MgCO_3 - and FeCO_3 -bearing carbonate melts in order to constrain the volumetric properties of these two relatively abundant carbonate components in the Earth's mantle. Knowledge of the standard state thermodynamic properties of these two important carbonate liquid components will allow thermodynamic models, such as p-MELTS (Ghiorso et al.

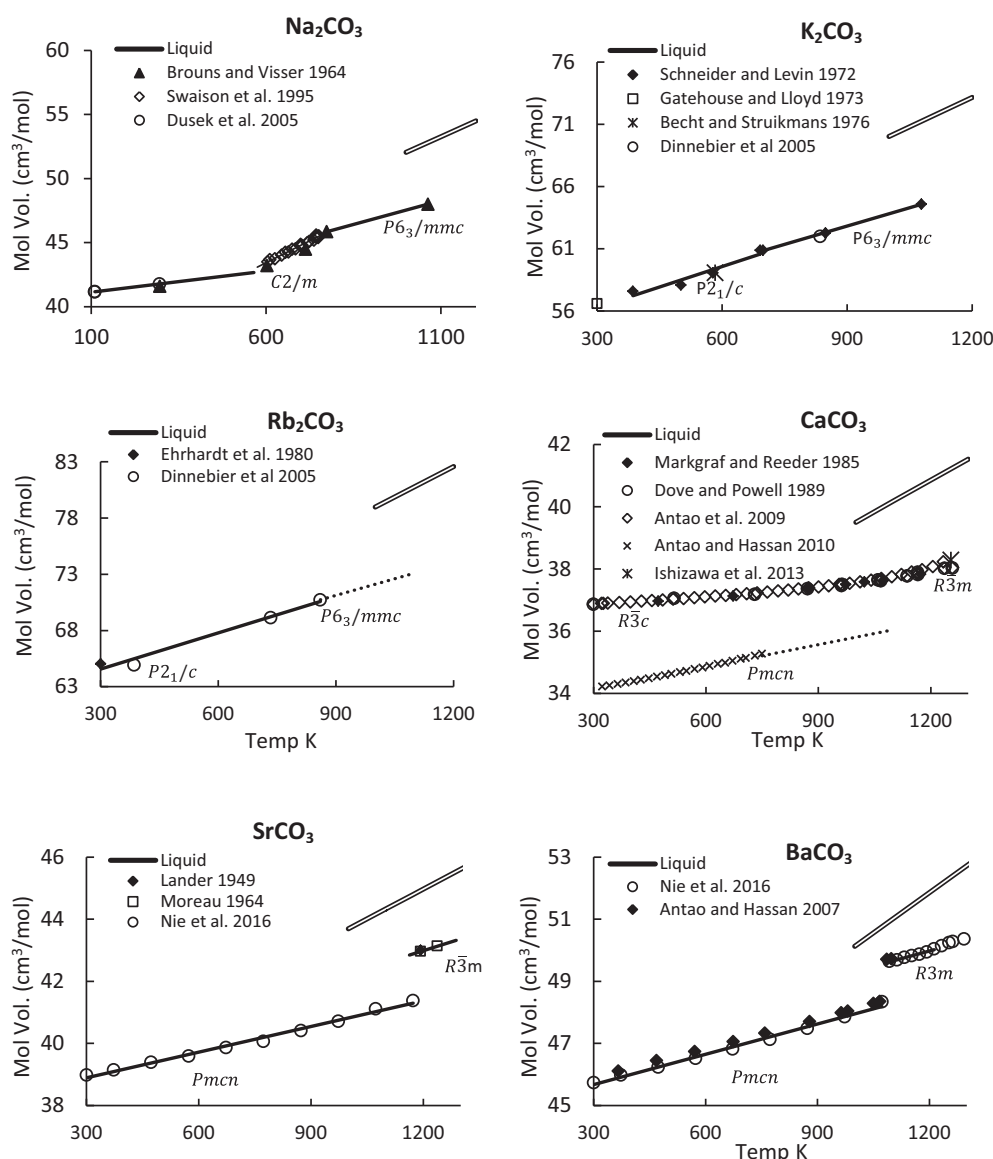
2002), to model partial melting of carbonate-bearing peridotite.

ACKNOWLEDGEMENTS

We thank the reviewers of this manuscript for their helpful and constructive feedback. We additionally thank Youxue Zhang, Aaron Wolf and Will Bender for their useful insights while writing this manuscript. Partial support for this research was provided from the National Science Foundation (NSF EAR-1763189), as well as the University of Michigan (Rackham Merit Fellowship to S.M. Hurt).

APPENDIX A

The molar volumes of crystal phases of each component have been collected from the literature and are shown as a function of temperature. The molar volume of the liquid is also shown as predicted by this study.



APPENDIX B. SUPPLEMENTARY MATERIAL

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.gca.2018.12.031>.

REFERENCES

- Antao S. M. and Hassan I. (2007) BaCO₃: high-temperature crystal structures and the Pmcn→R3m phase transition at 811 °C. *Phys. Chem. Miner.* **34**, 573–580.
- Antao M. and Hassan I. (2009) The orthorhombic structure of CaCO₃, SrCO₃, PbCO₃ and BaCO₃: linear structural trends. *Can. Mineral.* **47**, 1245–1255.
- Antao S. M. and Hassan I. (2010) Temperature dependence of the structural parameters in the transformation of aragonite to calcite, as determined from in situ synchrotron powder X-ray diffraction data. *Can. Mineral.* **48**, 1225–1236.
- Arakcheeva A., Bindi L., Pattison P., Meisser N., Chapuis G. and Pekov I. (2010) The incommensurately modulated structures of natural natrite at 120 and 293 K from synchrotron X-ray data. *Am. Min.* **95**, 574–581.
- Becht H. Y. and Struikmans R. (1976) A monoclinic high-temperature modification of potassium carbonate. *Acta Cryst.* **B32**, 3344–3346.
- Blundy J. and Dalton J. (2000) Experimental comparison of trace element partitioning between clinopyroxene and melt in carbonate and silicate systems, and implications for mantle metasomatism. *Contrib. Mineral. Petrol.* **139**, 356–371.
- Brouns E. and Visser J. W. (1964) An anomaly in the crystal structure of Na₂CO₃. *Acta Cryst.* **17**, 614.
- Cope E. R. and Dove M. T. (2007) Pair distribution functions calculated from interatomic potential models using the general utility lattice program. *J. Appl. Cryst.* **40**, 589–594.
- Dalton J. A. and Wood B. J. (1993) The compositions of primary carbonate melts and their evolution through wallrock reaction in the mantle. *Earth Planet. Sci. Lett.* **119**, 511–525.
- Dasgupta R. and Hirschmann M. M. (2006) Melting in the Earth's deep upper mantle caused by carbon dioxide. *Nature* **440**, 659–662.
- Dasgupta R. and Hirschmann M. M. (2010) The deep carbon cycle and melting in Earth's interior. *Earth Planet. Sci. Lett.* **298**, 1–13.
- Dasgupta R., Hirschmann M. M. and Stalker K. (2006) Immiscible transition from carbonate-rich to silicate-rich melts in the 3 GPa melting interval of eclogite + CO₂ and genesis of silica-undersaturated ocean island lavas. *J. Petrol.* **47**, 647–671.
- Datta R. K., Roy D. M., Faile S. P. and Tuttle O. F. (1964) Glass formation in carbonate systems. *J. Am. Ceram. Soc.* **47**, 153.
- Dinnebier R. E., Vensky S., Jansen M. and Hanson J. C. (2005) Crystal structures and topological aspects of the high-temperature phases and decomposition products of the alkali-metal oxalates M₂[C₂O₄](M=K, Rb, Cs). *Chem. Eur. J.* **11**, 1119–1129.
- Dobson D. P., Jones A. P., Rabe R., Sekine T., Kurita K., Taniguchi T., Kondo T., Kato T., Shimomura O. and Urakawa S. (1996) In-situ measurement of viscosity and density of carbonate melts at high pressure. *Earth Planet. Sci. Lett.* **143**, 207–215.
- Dove M. T. and Powell B. M. (1989) Neutron diffraction study of the tricritical orientational order/disorder phase transition in calcite at 1260 K. *Phys. Chem. Miner.* **16**, 503–507.
- Dusek M., Chapuis G., Meyer M. and Petricek V. (2003) Sodium carbonate revisited. *Acta Cryst.* **B59**, 337–352.
- Effenberger H. and Zemmann J. (1979) Verfeinerung der kristallstruktur des lithiumkarbonates, Li₂CO₃. *Z. Kristallogr. Cryst. Mater.* **150**, 133–138.
- Ehrhardt V. H., Schweer H. and Seidel H. (1980) Hochdrucksynthesen einiger carbonat mit überkritischem CO₂. *Z. Anorg. Allg. Chem.* **462**, 185–198.
- Gale J. D. and Rohl A. L. (2003) The General Utility Lattice Program (GULP). *Mol. Simul.* **29**, 291–341.
- Gatehouse B. M. and Lloyd D. J. (1973) Crystal structure of anhydrous potassium carbonate. *J. Chem. Soc. Dalton Trans.*, 70–72.
- Ghiorso M. S., Hirschmann M. M., Reiners P. W. and Kress V. C. (2002) The pMELTS: a revision of MELTS for improved calculation of phase relations and major element partitioning related to partial melting of the mantle to 3 GPa. *Geochem. Geophys. Geosystems* **3**, 36.
- Guo X., Lange R. A. and Ai Y. (2013) The density and compressibility of CaO-FeO-SiO₂ liquids at one bar: Evidence for four-coordinated Fe²⁺ in the CaFeO₂ component. *Geochim. Cosmochim. Acta* **120**, 206–219.
- Habasaki J. (1990) Molecular dynamics simulation of molten Li₂CO₃ and Na₂CO₃. *Mol. Phys.* **69**, 115–128.
- Hurt S. (2018) *The thermodynamic properties and structure and alkali and alkaline earth carbonate melts* Ph. D. thesis. Univ. Michigan.
- Idemoto Y., Richardson J. W., Koura N., Kohara S. and Loong C. K. (1998) Crystal structure of (Li_xK_{1-x})₂CO₃ (x = 0, 0.43, 0.5, 0.62, 1) by neutron powder diffraction analysis. *J. Chem. Phys. Solids* **59**, 363–376.
- Ishizawa N., Setoguchi H. and Yanagisawa K. (2013) Structural evolution of calcite at high temperatures: Phase V unveiled. *Sci. Rep.* **3**, 2832.
- Kang N., Schmidt M. W., Poli S., Franzolin E. and Connolly J. A. D. (2015) Melting of siderite to 20 GPa and thermodynamic properties of FeCO₃-melt. *Chem. Geol.* **400**, 34–43.
- Kang N., Schmidt M. W., Poli S., Connolly J. A. D. and Franzolin E. (2016) Melting relations in the system FeCO₃-MgCO₃ and thermodynamic modelling of Fe-Mg carbonate melts. *Contrib. Mineral. Petrol.* **171**, 1–16.
- Kohara S., Suzuya K. and Ohno H. (1999) A reverse Monte Carlo study of molten lithium carbonate. *Plasma Ions* **2**, 79–83.
- Kojima T., Yanagida M., Tanimoto K., Tamiya Y., Matsumoto H. and Miyazaki Y. (1999) The surface tension and the density of molten binary alkali carbonate systems. *Electrochemistry* **67**, 593–602.
- Kojima T., Miyazaki Y., Nomura K. and Tanimoto K. (2003) Density, molar volume, and surface tension of molten Li₂CO₃-Na₂CO₃ and Li₂CO₃-K₂CO₃ containing alkaline earth (Ca, Sr, and Ba) carbonates. *J. Electrochem. Soc.* **150**, E535–E542.
- Kojima T., Miyazaki Y., Nomura K. and Tanimoto K. (2008) Density, surface tension, and electrical conductivity of ternary molten carbonate system Li₂CO₃-Na₂CO₃-K₂CO₃ and methods for their estimation. *J. Electrochem. Soc.* **155**, F150.
- Lander J. J. (1949) Polymorphism and anion rotational disorder in the alkaline earth carbonates. *J. Chem. Phys.* **17**, 892–901.
- Lange R. A. (1997) A revised model for the density and thermal expansivity of K₂O-Na₂O-CaO-MgO-Al₂O₃-SiO₂ liquids from 700 to 1900 K: extension to crustal magmatic temperatures. *Contrib. Mineral. Petrol.* **130**, 1–11.
- Lange R. A. and Carmichael I. S. E. (1987) Densities of Na₂O-K₂O-CaO-MgO-FeO-Fe₂O₃-Al₂O₃-TiO₂-SiO₂ liquids: New measurements and derived partial molar properties. *Geochim. Cosmochim. Acta* **51**, 2931–2946.
- Liu Q. and Lange R. A. (2001) The partial molar volume and thermal expansivity of TiO₂ in alkali silicate melts: Systematic variation with Ti coordination. *Geochim. Cosmochim. Acta* **65**, 2379–2393.

- Liu Q. and Lange R. A. (2003) New density measurements on carbonate liquids and the partial molar volume of the CaCO_3 component. *Contrib. Mineral. Petrol.* **146**, 370–381.
- Liu Q. and Lange R. A. (2006) The partial molar volume of Fe_2O_3 in alkali silicate melts: evidence for an average Fe^{3+} coordination number near five. *Am. Mineral* **91**, 385–393.
- Markgraf S. and Reeder R. (1985) High-temperature structure refinements of calcite and magnesite. *Am. Miner.* **70**, 590–600.
- Moreau R. (1964) Etude diffractometrique du carbonate de strontium et du carbonate de baryum entre 20 et 1000°C. *Bull. Soc. Franc. Miner. Crist.* **87**, 547–552.
- Nie S., Liu Y., Liu Q., Wang M. and Wang H. (2017) Phase transitions and thermal expansion of BaCO_3 and SrCO_3 up to 1413 K. *Eur. J. Mineral.* **29**, 1–11.
- O’Leary M. C., Lange R. A. and Ai Y. (2015) The compressibility of CaCO_3 - Li_2CO_3 - Na_2CO_3 - K_2CO_3 liquids: Application to natrocarbonatite and CO_2 -bearing nephelinite liquids from Oldoinyo Lengai. *Contrib. Mineral. Petrol.* **170**, 1–18.
- Ragone S. E., Datta R. K., Roy D. M. and Tuttle O. F. (1966) The system potassium carbonate-magnesium carbonate. *J. Phys. Chem.* **70**, 3360–3361.
- Richet P. and Neuville D. R. (1992) Thermodynamics of silicate melts: configurational properties. *Adv. Phys. Geochem.* **10**, 132–160.
- Roest D. L., Ballone P., Bedeaux D. and Kjelstrup S. (2017) Molecular dynamics simulations of metal/molten alkali carbonate interfaces. *J. Phys. Chem. C* **121**, 17827–17847.
- Schneider S. J. and Levin E. M. (1973) Polymorphism of K_2CO_3 . *J. Am. Ceram. Soc.* **56**, 218–219.
- Shannon R. D. (1976) Revised effective ionic radii and systematic studies of interatomic distances in halides and chalcogenides. *Acta Cryst.* **A32**, 751–767.
- Shannon R. D. and Prewitt C. T. (1969) Effective ionic radii in oxides and fluorides. *Acta Cryst.* **B25**, 925–946.
- Shimoda K., Takahiro N. and Saito K. (2008) Local structure of magnesium in silicate glasses: a ^{25}Mg 3QMAS NMR study. *J. Phys. Chem. B* **112**, 6747–6752.
- Stein D. J., Stebbins J. F. and Carmichael I. S. E. (1986) Density of molten sodium aluminosilicates. *J. Am. Ceram. Soc.* **69**, 396–399.
- Stixrude L. and Bukowinski M. S. T. (1990) Rings, topology, and the density of tectosilicates. *Am. Mineral* **75**, 1159–1169.
- Stixrude L. and Bukowinski M. S. T. (1988) Simple covalent potential models of tetrahedral SiO_2 : Applications to α -quartz and coesite at pressure. *Phys. Chem. Minerals* **16**, 199–206.
- Strømme K. O. (1975) On the crystal structures of the high-temperature forms of strontium and barium carbonate and structurally related compounds. *Acta Chem. Scand.* **29**, 105–110.
- Swaison I. P., Dove M. T. and Harris M. J. (1995) Neutron powder diffraction study of the ferroelastic phase transition and lattice melting in sodium carbonate, Na_2CO_3 . *J. Phys.: Condens Matter* **7**, 4395–4417.
- Tissen J. T. W. M. and Janssen G. J. M. (1990) Molecular-dynamics simulation of molten alkali carbonates. *Mol. Phys.* **71**, 413–426.
- Toplis M. J. and Richet P. (2000) Equilibrium density and expansivity of silicate melts in the glass transition range. *Contrib. Mineral. Petrol.* **139**, 672–683.
- Vuilleumier R., Seitsonen A., Sator N. and Guillot B. (2014) Structure, equation of state and transport properties of molten calcium carbonate (CaCO_3) by atomistic simulations. *Geochim. Cosmochim. Acta* **141**, 547–566.

Associate editor: Michael J. Toplis