Exploring Tracer Information and Model Framework Trade-offs to Improve 1 **Estimation of Stream Transient Storage Processes** 2

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Key Points: 29

- TSM interpretation improved with analysis of multiple tracers, but results in increased 30 • parameter uncertainty for a more complex model 31
- Nonconservative tracers enabled interpretation of parameters that were highly uncertain 32 when estimated by conservative tracers alone 33
- Achieving reliable parameter estimates depends on choice of tracers and model 34 • framework, and should be coupled with uncertainty assessment 35
- 36

37 Abstract

Novel observation techniques (e.g., "smart" tracers) for characterizing coupled 38 39 hydrological and biogeochemical processes are improving understanding of stream network transport and transformation dynamics. In turn, these observations are thought to enable 40 increasingly sophisticated representations within transient storage models (TSM). However, 41 42 TSM parameter estimation is prone to issues with insensitivity and equifinality, which grow as parameters are added to model formulations. Currently, it is unclear whether (or not) 43 observations from different tracers may lead to greater process inference and reduced parameter 44 uncertainty in the context of TSM. Herein, we aim to unravel the role of in-stream processes 45 alongside metabolically active (MATS) and inactive storage zones (MITS) using variable TSM 46 formulations. Models with one (1SZ) and two storage zones (2SZ) and with and without 47 reactivity were applied to simulate conservative and "smart" tracer observations obtained 48 experimentally for two reaches with differing morphologies. As we show, "smart" tracers are 49 unsurprisingly superior to conservative tracers when it comes to partitioning MITS and MATS. 50 However, when transient storage is lumped within a 1SZ formulation, little improvement in 51 parameter uncertainty is gained by using a "smart" tracer, suggesting the addition of observations 52 should scale with model complexity. Importantly, our work identifies several inconsistencies and 53 open questions related to reconciling timescales of tracer observation with conceptual processes 54 ("parameters") estimated within TSM. Approaching TSM with multiple models and tracer 55 observations may be key to gaining improved insight into transient storage simulation as well as 56 advancing feedback loops between models and observations within hydrologic science. 57

58

59 Plain Language Summary

Solute experiments and transport models, called commonly "tracer experiments," are used to 60 understand the relative importance of different stream processes, especially those that influence 61 water, solutes, and nutrients as they move through a stream network. Within these tracer 62 experiments, there are processes that exchange mass beyond the main stream channel to other 63 parts of the river valley bottom environment. Sometimes there are single or multiple types of 64 tracers used and modeled to try to understand this exchange. There are also multiple models with 65 different equations and structures to simulate these tracers. This study shows that what you can 66 learn about these stream processes depends on experiment choices and which model you use. 67 Hence, refining future multiple tracer experiments and models is needed to determine how we 68 best obtain consistent measurements of key stream processes. 69

70

71 **1 Introduction**

72 The last decade has seen an explosion of novel techniques for collecting data used to characterize dynamic hydrologic systems. Tools and techniques that fall under this umbrella 73 include the burgeoning field of hydrogeophysics (e.g., Ward et al., 2010; St Clair et al., 2015), 74 the use of unmanned aerial vehicles (e.g., Vivoni et al. 2014; Brenner et al. 2017), high space-75 time resolution sensing systems (e.g. Blaen et al., 2016; Khamis et al., 2016) and the growing use 76 of "smart" and conservative tracers in the environment (e.g., Haggerty et al., 2008; González-77 Pinzón et al., 2012; Runkel 2015; Knapp et al., 2017; Blaen et al., 2017). Observational data 78 79 obtained from these techniques has been used to reveal new process dynamics and to refine

80 current understanding of hydrological systems. These techniques have also advanced process-

based mathematical representations within computational models as well as new approaches to

assess whether model frameworks ensure model realism (e.g., Seibert and McDonnell, 2002; Li

- et al., 2015; Clark et al., 2017). Furthermore, research communities have developed approaches
- for testing multiple model frameworks that explore different mathematical representations of hydrological processes (e.g., Clark et al., 2015a, 2015b) as well as approaches for comparing the
- performance of different models applied to a variety of systems (e.g., Butts et al., 2004; Best et
- al., 2015). Our goal herein is to build on recent progress made by these communities to explore
- the relationship between empirical observations, model performance, and model complexity to
- inform the value of new information for addressing historic limitations. We use the example of
- stream solute transport, transient storage, and solute transformation as a study topic.

91 In the field of groundwater-surface water interactions, hyporheic exchange remains one of the most difficult processes to quantify (Orghidan, 1959; Triska et al., 1989; Gooseff, 2010; 92 Boano et al., 2014). The hyporheic zone, although defined in a variety of contexts (Krause et al., 93 2011), is often described as a region both beneath and surrounding the stream channel containing 94 sediments, microbes, and benthic organisms where water and nutrient exchange with the main 95 stream channel occurs (Orghidan, 1959; Gooseff, 2010; Ward, 2016). Identifying this zone and 96 97 characterizing the relative rates and spatial extent of hyporheic exchange with the nearby stream channel has been and continues to be an area of ongoing research (e.g., Triska et al. 1989; Storey 98 et al., 2003; Boano et al., 2014; Caruso et al., 2016; Knapp et al. 2017; Schmadel et al. 2017). 99 While quantifying the role of the hyporheic zone in relation to solute transformations and 100 ecosystem processes has remained elusive, the use of multiple tracers, specifically the emerging 101 "smart" tracer (i.e., resazurin; Raz) technique has shown promise for characterizing stream 102 reactivity and functioning by enabling researchers to quantify the portion of the transient storage 103 that is metabolically-active (Karakashev et al., 2003; Haggerty et al., 2009; Argerich et al., 2011; 104 Knapp et al., 2018). Resazurin decays when in the presence of respiring cells typically found in 105 the hyporheic zone (e.g., González-Pinzón et al., 2012), producing a new chemical, resorufin 106 107 (Rru). Following this transformation, as water is exchanged across the streambed interface, Rru is exchanged back to the main channel and can be detected downstream. Thus, releasing Raz into 108 a stream reach produces two time-series of concentration, referred to as breakthrough curves 109 (BTC), that may be sensitive to different types of either metabolically active (MATS) or inactive 110 storage (MITS). Beyond hyporheic exchange, decay of Raz to Rru is being widely used to 111 characterize MATS, stream reactivity, and ecosystem respiration in many different transient 112 storage zones (Knapp et al., 2018), including biofilms (Haggerty et al., 2014), the benthic zone 113 (Knapp et al., 2017), vegetation beds (Kurz et al., 2017), and around woody debris (Blaen et al., 114 2018). In contrast to MATS, MITS may correspond to portions of a stream reach with a high 115 volume of water and conversely low contact with sediment (e.g., in-stream dead zones). 116

While MATS and MITS are recognized as having two very different effects on stream 117 nutrient exchange, there are few examples of TSM applications to reactive solutes (e.g., Gooseff 118 et al., 2005; Knapp and Cirpka, 2017). Commonly, quantifying reach-scale transient storage has 119 120 drawn upon parameter estimation with TSM representing the lumped effects of transient storage via MATS and MITS alongside advective in-channel processes such as advection and dispersion 121 (Bencala and Walters, 1983; Valett et al., 1996). When combined with field observations of 122 tracers in the form of a BTC, estimates of model parameters representing the temporal (i.e., rate 123 of exchange) and spatial scales (i.e., size) of reach-averaged transient storage zone exchange can 124 be obtained via inverse modeling (Runkel, 1998). This is done by employing one of many 125

methods (e.g., Runkel, 1998; Wagener et al., 2001; Kelleher et al., 2013; Knapp and Cirpka,

- 127 2017) to search the parameter space for a parameter set that produces the simulation with lowest
- 128 model error, assessed between a simulated and observed BTC for a given stream reach (Runkel,
- 129 1998; Ward et al., 2017). There are several recognized limitations to the timescales of transient 130 storage zone exchange that may be assessed using TSM (e.g., Harvey et al., 1996) as well as
- storage zone exchange that may be assessed using TSM (e.g., Harvey et al., 1996) as well as accounting for spatial heterogeneity that exists at sub-reach scales (Harvey et al., 1996; Knapp et
- al., 2017). Despite these limitations, TSM remains a popular approach that can provide an
- assessment of the relative roles of different in- and near-stream processes.

The most commonly applied TSM (known as the One-Dimensional Transport with 134 Inflow and Storage model, or "OTIS"; Runkel, 1998) uses four parameters to simulate transport 135 of a conservative tracer (five for a non-conservative tracer, e.g., Raz transforming to Rru) with a 136 single transient storage zone, and is available as open-source software from the United States 137 Geological Survey. However, this single-storage zone representation is inconsistent with current 138 understanding of transient storage (Briggs et al., 2009), as there are multiple dominant domains 139 of transient storage that may alter the flow of water and nutrients in different ways. As a result, 140 several other structural forms of the solute transport equations have been proposed (e.g., Lees et 141 al., 2000; Bencala and Walters, 1983; Runkel, 1998; Lees et al., 2000; Marion et al., 2008; 142 Briggs et al., 2009; Liao and Cirpka, 2011; Ward et al., 2015). One recent iteration of this model 143 separates the effects of transient storage into two zones described by parameters in terms of size 144 and exchange rate with the main channel (Briggs et al. 2009). Generally, we desire models with 145 process representations that most closely match our understanding of streams (e.g., Briggs et al., 146 2009). This desire often results in the addition of model parameters, with the tradeoff of 147 introducing additional uncertainty and equifinality due to parameter interactions (Beven 1993; 148 Butts et al. 2004; Beven 2006). This must also be balanced against the addition of observations 149 to vet simulations and constrain realistic parameter estimates. For instance, as shown by Briggs 150 et al. (2009), the addition of model parameters to segment transient storage was accompanied by 151 additional BTC observations from in-channel dead zones. Though numerous TSM studies exist, 152 153 there is a broad need to better understand the tradeoffs between parameter uncertainty and choices that determine the number of estimated model parameters (i.e., model complexity, 154 motivated by more realistic representation of dominant processes) alongside the addition of field 155 observations for estimating parameter values (i.e., "smart" tracers). 156

An added challenge to quantifying stream-reach transient storage is the growing body of 157 158 evidence that has shown that TSMs are susceptible to parameter equifinality (e.g., Choi et al., 2000; Kelleher et al., 2013), such that parameter determinations may be uncertain and therefore 159 uninformative for assessing the role of transient storage in physical and ecological river 160 161 processes (Wagner and Harvey, 1997; Wagener et al., 2002; Ward et al., 2017). Existing studies suffering from equifinality issues have typically assessed parameter estimates and uncertainty 162 through inverse modeling of a single conservative tracer. When used, as shown in a small but 163 growing number of studies, "smart" tracers provided different estimates for dispersion and 164 transient storage parameter values (Lemke et al., 2013). Adding observations to constrain models 165 is often viewed as an approach for reducing parameter uncertainty (e.g. Nearing and Gupta, 166 2015; Nearing et al., 2016). In practice, this requires that the observations in question contain 167 non-redundant information. If new or more observations lead to the same parameter estimates, or 168 similar levels of parameter uncertainty, the added information is not useful in reducing parameter 169 uncertainty. As parameter estimates are often used to characterize systems, collecting datasets 170 that can reduce this uncertainty is a common goal. Consequently, there is a need to explore how 171

equifinality and process inference may vary with multiple tracer observations, across different
 types of stream reach morphologies and across model formulations of varying complexity.

Within the context of TSM, we explore whether conservative and non-conservative 174 "smart" tracers may better constrain different TSM parameters, providing alternative but 175 potentially complimentary information. Furthermore, we offer a unique comparison of 176 177 constraints on parameter uncertainty arising from estimation of parameters by fitting to different tracer BTCs across model frameworks of varying complexity (e.g., conservative and non-178 conservative tracers, single versus multiple storage zone models). We aim to address the 179 following two fundamental questions in the context of TSM parameter estimation: (1) when are 180 multiple tracers useful?, and (2) when is increasing model complexity beneficial? Drawing from 181 several growing efforts in the land-surface and watershed modeling communities, we take a 182 183 model intercomparison-based approach (e.g., Best et al., 2015; Clark et al., 2015a, b; Clark et al., 2017), treating these TSM model formulations as different testable hypotheses, comparing the 184 performance and parameter uncertainty associated with each unique model formulation. We test 185 this approach using data from conservative (uranine; Ura) and reactive (Raz, Rru) solute tracer 186 experiments performed in two lowland stream reaches with distinct morphological settings 187 located in the Hammer Stream, in West Sussex, UK. To evaluate the addition of observations 188 alongside changes to model complexity, we test four different formulations of the TSM, ranging 189 190 in complexity from four to seven parameters. While we expect that inverse models constrained by Raz and Rru will reduce uncertainty and yield similar parameter estimates for active and 191 inactive storage zone parameters, we speculate that uncertainty in main channel parameters may 192 grow in the more complex model framework associated with two storage zones. 193

194 2 Study Area

Field experiments were conducted in the Hammer Stream (West Sussex, UK; 51°0' N 195 0°47' W: Figure S1). The 2.640 ha catchment drains mixed land use and is primarily underlain 196 197 by sandstone and mudstone. We identified two study reaches located upstream and downstream of Hammer Lake. Upstream of the lake the streambed material is sandy (hereafter the sand 198 reach), whereas the reach downstream of the lake is armored gravel (hereafter the gravel reach) 199 as result of the sand and other fine sediment having been trapped in Hammer Lake. Both reaches 200 201 include large woody debris in the stream channel (Blaen et al., 2018; Shelley et al., 2017). For the sand reach, we established a study reach along 760 m of channel (mean width 5.3 m, mean 202 203 depth 0.42 m). For the gravel reach, we established a study reach of 683 m immediately downstream of Hammer Lake (mean width 6.35 m, mean depth 0.28 m). In each reach, a 204 combination of Ura and Raz was injected as an instantaneous pulse about 150-m upstream of the 205 206 start of the study reach to ensure complete mixing, even at the start of the study reach. All 207 injections occurred in late afternoon/evening to minimize the effect of tracer mass photodegredation. In-situ field fluorometers (GGUN-FL30, Albillia Sàrl, Switzerland) were 208 used to monitor the fluorescence signals of all three tracers at 10-s time intervals at each end of 209 the study reach. Discharge was $73.2 \text{ L} \text{ s}^{-1}$ at the upstream end of the sand reach and 86.5 L s⁻¹ at 210 the upstream end of the gravel reach, and calculated using dilution gaging with Uranine at the 211 upstream end of the study site. Additional details regarding the sand reach and injection are 212 provided by Blaen et al. 2018; the gravel reach injection replicated the same experimental 213 methods. 214

215 3 Transient Storage Modeling

2163.1 Model Formulation

We derive models representing transport and transformation of solute tracers following 217 closely after the TSM (Thackston and Schnelle 1970; Bencala and Walters 1983) and integrate 218 several subsequent extensions such as multiple storage zones (e.g., Briggs et al. 2009; Kerr et al., 219 2013), reactivity (e.g., Haggerty et al. 2009; Lemke et al. 2013), and transport of multiple 220 221 interacting solutes (Keefe et al., 2004; Ward et al., 2015). While many TSM formulations have partitioned transient storage using location (e.g., surface and sub-surface transient storage; 222 Briggs et al. 2009; Kerr et al. 2013), MATS and MITS formulations separate transient storage 223 based on the apparent presence or absence of metabolic activity (Argerich et al. 2011). 224

In this TSM formulation, we simulate concentration in the main channel, the MITS domain, and the MATS domain via three equations with flexibility to vary up to seven different parameters. Concentrations in the stream domain are described according to:

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229
$$\frac{\partial C}{\partial t} = -\frac{Q}{A}\frac{\partial C}{\partial x} + D\frac{\partial^2 C}{\partial x^2} + \frac{q_{L,in}C_L}{A} - \frac{q_{L,out}C}{A} + \alpha_{MATS}(C_{MATS} - C) + \alpha_{MITS}(C_{MITS} - C)$$
(1a)

230

where C is solute concentration (g m⁻³), t is time (s), Q is discharge (m³ s⁻¹), A is cross-sectional 231 area of the stream (m²), D is the longitudinal dispersion coefficient (m² s⁻¹), $q_{L,in}$ is the lateral 232 inflow per meter of stream (m² s⁻¹), C_L is the concentration of the solute in the lateral inflow (g 233 m⁻³), $q_{L,out}$ is the lateral outflow per meter of stream (m² s⁻¹), and α describes the exchange rate 234 between the stream and transient storage zones (s^{-1}) . Inflows and outflows to the system 235 simulated as occurring in the stream instead of the hyporheic zone, consistent with TSM 236 conceptualization (Bencala and Walters, 1983). For the purposes of this experiment, both q_{Lin} 237 and q_{Lout} were set to zero on the basis of minimal changes in discharge and the absence of known 238 surface outflows along the study reach and to minimize the number of free parameters. 239 Furthermore, surface inflows do not contain the tracers Raz, Rru or Ura (i.e. C_L is also zero). 240

Within the MATS domain (denoted by subscript MATS), the mass balances for Raz, Rru, and Ura (denoted by subscripts) are calculated as:

243
$$\frac{\partial C_{MATS,Raz}}{\partial t} = \alpha_{MATS} \frac{A}{A_{MATS}} (C_{Raz} - C_{MATS,Raz}) - kC_{MATS,Raz}$$
(1b)

244
$$\frac{\partial C_{MATS,Rru}}{\partial t} = \alpha_{MATS} \frac{A}{A_{MATS}} (C_{Rru} - C_{MATS,Rru}) + k C_{MATS,Raz}$$
(1c)

245
$$\frac{\partial C_{MATS,Ura}}{\partial t} = \alpha_{MATS} \frac{A}{A_{MATS}} (C_{Ura} - C_{MATS,Ura})$$
(1d)

where k (s⁻¹) is a reactive rate constant that describes the transformation of the parent (Raz in our study) to product (Rru in our study), and C_{Ura} , C_{Raz} , and C_{Rru} are the in-stream concentrations of Uranine, Resazurin, and Resorufin based on solving equation 1a for each solute. Similarly, in the MITS domain (denoted by subscript MITS) the concentrations of Raz, Rru, and Ura are calculated as:

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$$\frac{\partial C_{MITS}}{\partial t} = \alpha_{MITS} \frac{A}{A_{MITS}} \left(C_{Raz} - C_{MITS,Raz} \right)$$
(1e)

252
$$\frac{\partial C_{MITS}}{\partial t} = \alpha_{MITS} \frac{A}{A_{MITS}} \left(C_{Rru} - C_{MITS,Rru} \right)$$
(1f)

253
$$\frac{\partial C_{MITS,Ura}}{\partial t} = \alpha_{MITS} \frac{A}{A_{MITS}} (C_{Ura} - C_{MITS,Ura})$$
(1g)

254 Simulations are performed through forward modeling based on the stream (1a), MATS (1b-1d), and MITS (1e-1g) equations for Ura, Raz, and Rru. Critical to this study is that all three 255 equations are solved using the same physical transport parameters (A, D, A_{MATS}, A_{MITS}, α_{MITS} , α_{MITS} , 256 α_{MATS}), allowing for simulation of dynamic parent-to-product transformations. This solution 257 allows the simultaneous transport and interaction of both conservative and interacting reactive 258 259 solutes (after Ward et al. 2015). Model equations for all solutes were solved simultaneously using a Crank-Nicolson solution scheme, common for TSM applications (e.g., Runkel 1998; 260 Ward et al., 2015). Models were implemented using measured discharge at the upstream end of 261 each study reach, with observed breakthrough curves used as upstream boundary conditions. 262 Spatial and temporal steps for the simulations were fixed at 5-m and 10-s, respectively. 263 Important and commonly used assumptions of the model include laterally and vertically well-264 mixed domains, exponential residence time distributions within transient storage zones, temporal 265 constancy for transient storage zone model parameters, perfect conversion of Raz to Rru, no 266 retardation (sorption), and no additional transformation pathways for any solutes. 267

As derived, the model is flexible in that it can represent both one storage zone (1SZ) and two storage zone (2SZ) realizations of the TSM (Figure 1a). Within this framework, we test the following model-tracer combinations (Figure 1b):

- 271 (1) A one storage zone model fit to a conservative tracer (Ura), where transient 272 storage combines MATS and MITS (four parameters; A, D, A_s , α_S)
- 273 (2) A two storage zone model fit to a conservative tracer (Ura), where MITS and 274 MATS do not distinguish active storage, but instead represent two different 275 storage zones (six parameters; A, D, A_{MATS} , α_{MATS} , α_{MITS} , α_{MITS})
- 276 (3) A one storage zone model fit to "smart" tracer (Raz) and biproduct (Rru), where 277 transient storage refers to MATS, and MITS is effectively incorporated into the 278 dispersion term (D) (five parameters; A, D, A_{MATS} , α_{MATS} , k)
- 279 (4) A two storage zone model fit to "smart" tracer (Raz) and biproduct (Rru) (seven parameters; *A*, *D*, *A*_{MATS}, *a*_{MATS}, *a*_{MITS}, *a*_{MITS}, *k*).

While comparison (2) is included in this study to assist with interpretation, this combination of a 281 two-storage zone model fit with a conservative tracer is not expected to yield useful storage zone 282 283 parameter estimates. Each tracer was independently tested as a source of parameter information for each tracer-model combination listed above. Importantly, MATS and MITS, as visualized in 284 Figure 1a, are integrations of transient storage along the channel that may either be reactive or 285 286 inactive, respectively. MATS and MITS, in our formulation, do not represent physical zones such as the hyporheic zone, as many of these physical locations may include zones of both active 287 and inactive storage. For 1SZ models, the exchange coefficient α_{MTS} is set to zero, which 288 eliminates any exchange between the stream and MITS. For Ura, this represents a 1SZ model 289 identical in formulation and implementation to the broadly used USGS OTIS model (Runkel 290 1998). For Raz and Rru, the MATS storage retains the ability to simulate transformations in the 291 292 storage zone and assumes that all transient storage is metabolically active (i.e., with $\alpha_{MTS} = 0$, A_{MITS} cannot affect concentrations in the model). We do assume Ura to be a conservative tracer 293

(see Supporting Information), though it may decay in direct sunlight. Notably, Rhodamine WT
 is also non-conservative (Runkel, 2015), highlighting that there is no single perfect conservative
 tracer.

Conceptually, the combinations of observations and model formulations listed in Figure 297 1b also represent four different scenarios for gaining insights regarding parameter importance. 298 299 Though we assess parameter sensitivity and uncertainty per tracer and per model, we do not expect all parameter estimates to be sensitive to all tracers, and seek to test these potential 300 relationships. In this same vein, certain tracers are likely to be more or less informative for 301 different parameters. We expect that Ura, as a conservative tracer, will yield the most physically 302 representative parameter distributions for A and D. Similarly, we do not expect that Ura is 303 capable of separating the influence of MATS and MITS, and are uncertain as to whether A and D 304 305 are sensitive to Raz or Rru, given these tracers are nonconservative.

Importantly, parameters estimated via different tracers represent different processes.
 Within the 1SZ formulation, storage zone parameters estimated via Ura assume the transient storage zone combines both MATS and MITS, while storage zone parameters estimated via Raz estimate transient storage zone size for MATS, assuming MITS is incorporated into the dispersion term. Thus, we may not necessarily expect distributions of *D* or storage zone

parameters to be similar when fitting to Ura versus Raz with the 1D model.

312 3.2 Computational Experiments

We performed several computational experiments with inputs (parameters) to and outputs (errors and simulations) from models of varying complexity. Simulations and parameter sets were constrained to match different observations, including conservative (Ura) and nonconservative (Raz, Rru) BTCs. Model formulations used in these experiments are outlined in Figure 1b. To interrogate parameter uncertainty and equifinality, we used a Latin Hypercube approach to sample the model parameter space (N = 27,000 runs; e.g., Pianosi et al., 2015). All parameters and associated ranges are listed in Table 1.

Within the 2SZ model formulation, we sampled total area (A_{TOT}), representing the combined area of the advective channel and the area of MITS. For this formulation,

322	$A = A_{TOT} \cdot (1 - F_{MITS})$	2a
323	$A_{MITS} = A_{TOT} \cdot F_{MITS}$	2b

where F_{MITS} describes the fraction of the stream channel that is metabolically inactive. To enable comparisons across model formulations, all results are presented in terms of *A* and A_{MITS} . This does result in slightly wider bounds for *A* for the 2SZ model and narrower bounds for *A* for the 1SZ model, but otherwise is purely a function of model formulation.

Model complexity, defined by the number of parameters, ranged from four to seven 328 parameters. We tested a 1SZ model (Fig. 1b, 1) and 2SZ model (Fig. 1b, 2) constrained by 329 observations from only Ura. We also tested a 1SZ (Fig. 1b, 3) and a 2SZ (Fig. 1b, 4) with an 330 added parameter representing reactive decay constrained by BTCs for Raz and for Rru. All 331 computational experiments were performed using the same structural model equations 332 (Equations 1a, 1b, 1c, and 1d). For model formulations 1 and 3, we use a model formulation that 333 has a single transient storage zone (i.e., $\alpha_{MITS} = 0$). To model this, we sampled the first five 334 parameters, setting values for the fraction of stream area as MITS and the MITS exchange rate to 335

small non-zero values (10^{-10}) . For model formulations 2 and 4, we sampled all seven parameters across feasible ranges.

338 3.3 Model Performance

For each of the 27,000 runs, we calculated model fits in terms of a normalized Root Mean Squared Error *(nRMSE)* for each BTC (Ura, Raz, Rru) independently, according to:

341
$$nRMSE = \frac{1}{c_p} \left(\frac{\sum_{t=1}^n (o_t - s_t)^2}{n} \right)^{0.5}$$
(2)

342 where O_t and S_t correspond to observations and simulations at a given time step, n is the total number of observations, and C_p is peak concentration for each tracer (employed to normalize 343 RMSE values across tracers; g m⁻³). RMSE (and close variants) remains one of the key objective 344 functions used to assess BTC errors (Runkel, 1998; Ward et al., 2017). We also calculated a log-345 transformed Root Mean Squared Error (LRMSE; similar to a weighted RMSE), where the 346 observed and simulated time-series were log-transformed prior to applying equation (2) above. 347 348 Past work has shown log-transformed error metrics can be particularly useful for obtaining reliable TSM parameter estimates (e.g., Wörman and Wachniew, 2007; Ward et al., 2017). 349

Our analysis relies on the use of behavioral thresholds to segment populations of error 350 and parameter estimates (e.g., Hornberger and Spear, 1980, 1981; Spear and Hornberger, 1980). 351 We employ a behavioral threshold to identify a subset of simulations and parameter sets that 352 closely match BTCs by achieving low errors. Instead of selecting a single best simulation and 353 parameter set, the use of behavioral thresholds allows us to identify a distribution of these values. 354 Behavioral thresholds may be implemented by identifying parameter sets and simulations below 355 a certain error value, or by identifying those with errors below some percentage criterion (i.e., 356 top 10% of errors). We use the latter (thresholds of 10% and 1%) to compare error, simulations, 357 and parameter estimates across different tracers and models. 358

359 3.4 Parameter Sensitivity and Uncertainty

For an ideal TSM inverse modeling exercise there is a unique 'best' estimate for each 360 parameter, such that behavioral parameter values occupy a defined and narrow area of the 361 parameter space (Wagener et al., 2001; Kelleher et al., 2013). However, parameters are often 362 insensitive or uncertain. This may be represented as wide distributions of behavioral parameter 363 values spanning the entire parameter range, or by no difference between parameter distributions 364 between the best and worst simulations. The former may also occur when a parameter is largely 365 unimportant, or due to interactions with other parameters. To understand the influence of model 366 complexity and different tracer observations on parameter estimates, we assessed parameter 367 sensitivity, parameter uncertainty, and parameter interactions for all model-tracer combinations. 368 While some studies have estimated parameters by first fitting parameters associated with 369 conservative transport, and then fitting nonconservative parameters (e.g., Gooseff et al., 2005), 370 we treat all BTCs as independent sources of information for assessing parameter sensitivity and 371 uncertainty. Our goal is to avoid making assumptions about which BTCs may contain 372 information regarding certain parameters, and instead to use the analysis presented here to more 373 374 thoroughly assess how parameter estimates are impacted by fitting to different BTCs.

Approaches to obtain parameter estimates include use of optimization algorithms (e.g., Runkel, 1998; Briggs et al., 2009; Kerr et al., 2013), Markov Chain Monte Carlo approaches

(e.g., Lemke et al., 2013; Knapp and Cirpka, 2017), and Monte Carlo approaches coupled with 377 378 behavioral thresholds (e.g., Wagener et al., 2001; Kelleher et al., 2013), as well as a broad literature on approaches to parameter sensitivity (see Pianosi et al., 2016). In this study, we 379 380 employed approaches based on Monte Carlo methods to enable a tiered assessment that draws on both the very best and worst simulations and corresponding parameter estimates. To provide a 381 global assessment of parameter sensitivity, we generated regional sensitivity analysis (RSA) 382 plots for each parameter based on errors associated with *nRMSE* calculated from simulations and 383 observations of Ura, Raz, and Rru (Fig. 2a). RSA is a useful technique for mapping portions of 384 the parameter space corresponding to either best or worst errors (e.g., Freer et al., 1996; Pianosi 385 et al., 2016) and has been commonly applied to assess TSM parameter sensitivity (e.g., Wagener 386 et al., 2002; Wlostowski et al., 2013). To apply RSA, we identified the top (best) 10% of errors 387 and the bottom (worst) 10% of errors for Ura, Raz and Rru across all simulations. Parameter 388 values corresponding to these best and worst 10% of simulations were transformed into marginal 389 empirical cumulative distribution functions (CDF; Freer et al., 1996; Wagener et al., 2001; 390 Pianosi et al., 2016). Sensitive parameters satisfied two criteria: parameter CDFs corresponding 391 392 to the top 10% of all error values (1) deviated from the 1:1 line (representing a purely uniform distribution), assessed by visual inspection, and (2) deviated from parameter CDFs 393 corresponding to the worst 10% of all errors (Fig. 2a, b). 394

395 We assessed parameter uncertainty and model performance comparing the top 1% of all simulations per error metric. To test whether the parameter values corresponding to the lowest 396 model errors converged, we applied a visualization based on the widely used dotty plot (e.g., 397 Wagener and Kollat, 2007). Dotty plots visualize model error plotted against model parameter 398 values for all simulations meeting a given behavioral threshold (Fig. 2c, d). To summarize the 399 distribution of optimal parameter values (those corresponding to the lowest error) across each 400 dotty plot, we identified the single best parameter value (with lowest error) within a moving 401 window $(1/20^{\text{th}} \text{ the width of parameter range})$ incremented across each parameter range $(1/40^{\text{th}})$ 402 the width of the parameter range). This distribution of optimal errors was then normalized to a 403 404 cumulative value of one (Fig. 2c, d). We report all dotty plots in Supporting Information (Figs. S1-S4). Optimal parameter values and 90% confidence intervals are also reported. Finally, we 405 also investigated parameter interactions via scatter plots of parameter values to assess the 406 dependency between parameters and how this changes for subsets of the very best simulations. 407 Together, these assessments yield transferable approaches for assessing parameter sensitivity and 408 uncertainty within environmental models, and for comparing these outcomes across 1SZ and 409

410 2SZ models and error metrics.

411 **4 Results**

412 4.1 Model errors and simulations

Tracer observations obtained from the two reaches are shown in Figure 3. While peak concentrations for Ura and Raz are coincident, peak concentrations for Rru occur at a later time, representing a temporal lag as Raz is converted to Rru in the presence of aerobic respiration. All tracers are capably simulated by one and two-storage zone models (Figure 4). As we show, a behavioral threshold of 1% yielded envelopes of simulations that bracketed observations for all tracers. Upper and lower bounds, representing the range of the 270 simulations with lowest error, were nearly identical for the 1SZ and 2SZ models. Information on mass recovery is included in 420 Supporting Information (Text S2) and demonstrated acceptable levels of mass recovery

421 conforming to model assumptions.

Model errors corresponding to the top 1% of all model simulations are visualized for 422 nRMSE in Figure 5. Comparing distributions of error across tracers, nRMSE_{Raz} had the lowest 423 overall error across both reaches, though medians and ranges of error between tracers were 424 425 similar for all but $nRMSE_{Rru}$. Minimum and median errors for Rru were larger for the 2SZ as opposed to 1SZ model, and $nRMSE_{Rru}$ had larger errors than $nRMSE_{Raz}$ and $nRMSE_{Ura}$. 426 Comparing errors across models, we found that median errors for 2SZ models were slightly 427 higher than median errors for 1SZ models for nearly all reach-tracer combinations. Though the 428 ranges of error were found to be wider for 2SZ as opposed to 1SZ models, 2SZ models still 429 yielded the simulation with the single lowest error across all parameter sets for $nRMSE_{Raz}$ for 430 both reaches and $nRMSE_{Ura}$ for the gravel reach (Figure 5a). 431

432 4.2 Parameter sensitivities

Interpretation of global parameter sensitivities assessed via RSA are shown in Figure 6. A 433 434 select number of RSA plots are included for the 2SZ models, with all plots included in Supporting Information (see Figure S8). In general, D, A, and α_{MATS} were globally sensitive 435 across tracers. Distributions for D differed between tracers. For the gravel reach, lower errors for 436 $nRMSE_{Ura}$ corresponded to larger values for D, but smaller values of D for $nRMSE_{Raz}$ and 437 $nRMSE_{Rru}$. A_S (1SZ) and A_{MATS} (2SZ) were both sensitive, the latter to $nRMSE_{Ura}$ and $nRMSE_{Raz}$ 438 and the former to $nRMSE_{Raz}$ and $nRMSE_{Rru}$. Estimates for A_{MITS} and α_{MITS} were difficult to 439 interpret, in part, because CDFs corresponding to both best and worst performing parameter 440 values were similar, likely indicating that these parameters are influenced by interactions with 441 other parameters. Finally, k was globally insensitive across all models and performance metrics. 442

443 4.3 Parameter uncertainties

Parameter estimates corresponding to the top 1% of *nRMSE* values for each tracer are 444 summarized as distributions in Figure 7 (parameter values and confidence intervals are reported 445 in Table S2). In general, flatter distributions indicate that all values across the parameter range 446 produce equal model errors, while the presence of peaks indicates certain areas of the parameter 447 space produce higher or lower errors, suggesting that there are optimal values that better simulate 448 observations. Comparisons of distributions were informative for testing whether conservative 449 versus "smart" tracer errors vielded differences in parameter uncertainty as well as whether 450 regions of the parameter space corresponding to the best simulations and therefore minimum 451 error were similar across tracers. 452

Across models and reaches, parameter estimates were peaked for *A* and *D*, and narrower for the 1SZ (vs. 2SZ) errors. Within the 1SZ models, parameter estimates were uncertain for lumped transient storage size (A_S) for all tracers (note that parameter values with lowest error were distributed across the entire parameter range, spanning two orders of magnitude). In contrast, PDFs for α_S were peaky for $nRMSE_{Ura}$ and $nRMSE_{Rru}$ (Fig. 7).

458 Separating transient storage into two storage zones (two indiscernible zones for Ura; 459 MATS and MITS for Raz, Rru) introduced different patterns of parameter uncertainty. PDFs for 460 A_{MATS} were wide for all reach-tracer combinations. Empirical PDFs of α_{MATS} suggest better 461 actimates for this parameter correspond to lower values when fitting to Ura and Paz, and better

estimates for this parameter correspond to lower values when fitting to Ura and Raz, and better

estimates correspond to higher values when fitting to Rru. MITS parameters (A_{MITS} , α_{MITS}) were best constrained by $nRMSE_{Rru}$ (Fig. 7).

Employing a non-conservative tracer (e.g., Raz) introduced an additional parameter k to 464 describe the rate of transformation from Raz to Rru within the 1SZ and 2SZ models. This implies 465 that fitting to one or both of these "smart" tracers should reduce uncertainty in this value. Global 466 sensitivity analyses suggest that k is less sensitive than some storage zone parameters (e.g., 467 Figure 6, Figure 7). However, dotty plots of k for both reaches (Figure S7) do suggest that this 468 parameter is both sensitive (i.e., errors vary across the parameter range) and unique (such that a 469 single best value exists within the parameter range). Within these dotty plots, values for k470 appeared insensitive to Raz, suggesting that Rru may contain more information for estimating 471 this parameter. 472

Our results also underscore the importance of considering alternative objective functions. 473 While not the primary focus of this manuscript, we include an additional assessment of 474 parameter uncertainty with respect to log-transformed *nRMSE* (*LRMSE*) for both reaches. PDFs 475 for the gravel reach were generally wide across parameters, suggesting that *nRMSE* was a more 476 informative error metric in this reach (Figure S4). In contrast, empirical PDFs for the top 1% of 477 errors by *LRMSE* were peaked for nearly all sand reach parameter values (Figure S6; Figure S7). 478 These results demonstrate potential value added by considering alternative error formulations in 479 assessments of parameter uncertainty. 480

481 4.4 Joint Distributions

Given past work suggesting that TSM parameters are influenced by interactions, we 482 examined joint distributions of parameter values to explore how the interactive nature of TSM 483 may impact parameter estimates and parameter uncertainty (Figs. 8, 9). While some parameters 484 may be globally insensitive (Fig. 6) or exhibit flat parameter distributions (Fig. 7), visualizing 485 joint distributions can reveal the presence of more complex relationships as well as the value of 486 different tracers to discern these relationships. Figure 9 displays how parameter estimates and 487 joint distributions varied with model complexity. Joint distributions of A and D were bimodal, 488 and widened for the 2SZ model. Similar patterns were also observed between $A_{S}(1SZ)$ and 489 A_{MATS} (2SZ). In particular, these plots display that estimates for α_s were best constrained by Ura 490 491 and Rru for the 1SZ model, but Rru for the 2SZ model, shown by the shrinking 2D boundary of highest performing parameter combinations. While PDFs of parameter estimates for k did not 492 reveal any strong patterns, joint distributions suggest lower errors are concentrated in a distinct 493 494 portion of the parameter range for k (Raz, 1SZ; Rru, 2SZ). Last, we explored joint distributions between parameters only present in the 2SZ model, A_{MITS} and α_{MITS} (Fig. 9). In particular, these 495 joint distributions display the importance of Rru for refining estimates of MITS parameters. We 496 497 note that Figures 9 and 10 display results for the gravel reach, with visualizations for the sand reach included in Supporting Information (Fig. S8 and S9), as the patterns of these joint 498 distributions were similar between the reaches. 499

500 **5 Discussion**

501 5.1 Model complexity and conceptualization, simulations, and errors

502 Behavioral simulation bounds (Fig. 4; Fig. S2) and model errors (Fig. 5) indicate that, 503 regardless of the tracer error or model framework used to constrain behavioral simulations, observations of all tracers were simulated to a reasonable degree of accuracy. Average errors for

- behavioral simulations were similar between 1SZ and 2SZ models. However, it is important to
- note that accurate simulation from an inverse model do not necessarily indicate meaningful
- 507 information was gained from parameter values. Parameters that are not identifiable may provide 508 a good inverse model fit without characterizing system processes and should not be over-
- a good inverse model fit without characterizing system processes and should not be overinterpreted (e.g., by comparing or interpreting values of insensitive or uncertain parameters).
- 510 Thus, we echo calls for assessment of model parameter uncertainties, interactions, and
- 511 identifiability as a requisite step prior to their interpretation (Wagener and Harvey, 1997;
- 512 Wagener et al., 2002; Kelleher et al., 2013; Ward et al., 2017).

For the stream reaches analyzed in this study, we found that employing a more complex 513 model did not necessarily yield simulations that better approximated tracer observations. Given 514 515 the increased degrees of freedom in a 2SZ (as opposed to 1SZ) model, we expected 2SZ models to display smaller magnitude and range of errors than the 1SZ formulation errors calculated 516 between measured and simulated BTCs. Instead, 1SZ versus 2SZ model errors were similar (and 517 even notably larger for Rru), though the simulation with the lowest error was almost always 518 generated with a 2SZ model (Fig. 5). We do not believe that these similarities in error indicate 519 that the model is a poor representation of reality, as simulations well approximated observations 520 (Fig. 4). Instead, we postulate that this shows that adding additional parameters introduces 521 further uncertainty in addition to degrees of freedom, yielding similar model fits. It is also 522 possible that a more complex or alternative model formulation could lead to improvements in 523 error, and potentially a better representation of reality. Given the many iterations of TSM 524 formulations (e.g., Gooseff et al., 2003; Gooseff et al., 2007; Kerr et al., 2013), we advise future 525 work is needed to perform TSM model intercomparison with respect to both conservative and 526 "smart" tracer BTCs. 527

While we sought to compare parameter inference through uncertainty assessment across 528 multiple models and tracers, this introduces some challenges in interpretation. This is because 529 530 transient storage parameters conceptually represent different processes when inverse modeling is performed with respect to different tracers. Transient storage zones cannot be partitioned into 531 MATS or MITS through inverse modeling to simulate Uranine. Instead, transient storage zone 532 parameters estimated via fitting a 2SZ model to Uranine assumed these parameters represent two 533 534 independent storage zones with no association with MATS or MITS. Therefore, parameter distributions for storage zone parameters represent fundamentally different processes when 535 536 fitting to Ura versus Raz and Rru, and as such, are not expected to be comparable. For this reason, we do not recommend estimating 2SZ MATS and MITS parameters by fitting to Ura, but 537 include this comparison to emphasize that combining different model formulations and tracers 538 539 can lead to fundamentally different conceptual representations of a system. Likewise, in the 1SZ formulation, storage parameters are assumed to represent MATS processes when fitting to Raz, 540 with MITS lumped with dispersion. Therefore, we did not expect empirical PDFs for these 541 542 parameter values to be similar. Indeed, these differences likely explain why fitting to Ura versus Raz yields such different empirical PDFs for $\alpha_{\rm S}$ (Figure 7). These differences also show that 543 parameter estimates obtained by fitting a 1SZ model to Ura are not comparable to parameter 544 estimates obtained by fitting a 1SZ model to Raz. 545

546

5.2 How do conservative versus nonconservative tracers affect parameter uncertainty?

547 In contrast to the expectation that conservative tracers may not always provide 548 meaningful parameter estimates, our results show that conservative tracer BTCs do contain useful information for estimating TSM parameter values. In support of this, we found parameter

- uncertainty tended to be lower for parameters fit to conservative tracer BTCs (i.e., black
- distributions are narrower than blue or red distributions in Fig. 7). Encouragingly, we found that
- relatively narrow estimates for α could be achieved using a conservative tracer with either a 1SZ or a 2SZ model. This is in contrast to studies that have concluded α is typically highly uncertain
- or a 2SZ model. This is in contrast to studies that have concluded α is typically highly uncertain (Wagener et al., 2002; Kelleher et al., 2013; Wlostowski et al., 2013). Thus, estimates of α with
- low uncertainty can be achieved, but this model result may be dependent upon the system and
- tracers. Overall, using multiple tracers allowed us to estimate and evaluate BTC parameters to a
- 557 higher degree than could be achieved by using a single tracer, with consistency in findings across
- both reaches. We therefore recommend TSM parameter estimates and subsequent process-based
- interpretation should be based on the combination of conservative and non-conservative tracers.

The parameter we found most problematic to estimate was k, which describes the 560 transformation of Raz to Rru and effectively determines mass balance. As shown in Figures 8 561 and 9, we found k to be highly interactive, which may explain apparent insensitivity and 562 uncertainty for this parameter (Figures 6 and 7, Table S1). Furthermore, dotty plots (Fig. S7) 563 between k and objective functions ($nRMSE_{Rru}$, $LRMSE_{Rru}$) show that k is indeed sensitive, it is 564 just less sensitive than other model parameters. While only a few studies exist that employ 565 formulations of MITS and MATS alongside "smart" tracer observations, some have concluded, 566 similar to our study, that k may be highly uncertain (Yakerivich et al., 2017). Others have found 567 low uncertainty for k through joint fitting of multiple tracers (Lemke et al., 2013). As this value 568 is of particular interest to biogeochemists, future research with paired conservative and 569 nonconservative tracer experiments will be needed to identify conditions that may lead to more 570 (or less) uncertain k estimates. 571

In a similar vein, parameters A_S (1SZ) and A_{MATS} (2SZ) were also uncertain across study 572 reaches (Fig. 7). Though we observed some organization between the structure of first order 573 parameter interactions and model errors, our work suggests that these processes were difficult to 574 575 estimate in this particular system. While not performed here, other analyses of parameter sensitivity and uncertainty (e.g., Kelleher et al., 2013) have shown that sometimes nested 576 sampling schemes (narrowing bounds on certain parameters before completing additional 577 analysis) can improve estimates of parameter values and associated uncertainty. This is because 578 579 fitting to all BTCs is likely to be dominated by first finding best estimates for A and D. Fixing these values to narrow ranges, thereby reducing degrees of freedom, enables the importance of 580 581 other parameters less sensitive than A and D to be identified, and may be an approach for obtaining more reliable estimates of problematically uncertain parameters. 582

583 Consistent with several recent studies using reactive tracer systems and TSM, we broadly found improved parameter constraints for some, but not all parameters associated with inclusion 584 of reactive tracers (e.g., Lemke et al., 2013; Yakerivich et al., 2017) or additional experimental 585 observations (e.g., Briggs et al., 2009; Neilson et al., 2010). Transient storage parameter 586 uncertainties were minimized when a more complex model was used, most likely because this 587 leads to greater degrees of freedom for fitting observations. For researchers who wish to separate 588 589 the relative influences of transient storage between MITS and MATS, a 2SZ model simulating both conservative and "smart" tracer BTCs was capable of narrowing nearly all parameter 590 estimates. We did find variations in parameter sensitivity and uncertainty across reaches. This is 591 592 not surprising, given that the relative importance of different processes varies at the reach scale, and will determine parameter sensitivity and uncertainty within TSM applications. Though 593

⁵⁹⁴ "smart" tracers are unsurprisingly superior to conservative tracers when it comes to partitioning

595 MITS and MATS, little improvement in parameter uncertainty was gained for 1SZ model 596 formulations by using a "smart" tracer

596 formulations by using a "smart" tracer.

597 5.3 Is information obtained from conservative and "smart" tracers complimentary or598 redundant?

For 1SZ models of conservative and "smart" tracers, a similar number of sensitive 599 600 parameters were identified, illustrating that both tracer types contain valuable and potentially complimentary information. Furthermore, parameter estimates obtained with respect to all tracers 601 were similar, but differed in some cases. On one hand, some tracers are likely to be more 602 sensitive to main channel (e.g., Ura) versus storage zone (e.g., Rru) parameters and 603 corresponding processes. This is a likely explanation for the difference in the empirical PDFs 604 obtained for A and D fitting to Ura and Rru. As we would not expect fitting to Rru would contain 605 606 information about main channel processes, this is unsurprising. A further explanation for the non-ideal estimation of A and D may be its sorption behavior in the subsurface (e.g., Lemke et 607 al., 2014). Conversely, Raz and Ura may both provide similar information regarding A and D. 608 609 Therefore, our work shows that even non-conservative tracers like Raz may still be useful for estimating parameters conceptualizing main channel processes. 610

In contrast, we also found differences in parameter estimates for transient storage 611 exchange rate, α_s , when fitting to different tracers. This outcome was also mirrored within the 612 2SZ formulation for MATS exchange rates, and similar to findings from Lemke et al. (2013). 613 These differences in estimates of transient storage parameters indicate that conservative and 614 "smart" tracers may be sensitive to different timescales of transient storage. It is not clear why 615 Raz and Rru would lead to different empirical PDFs and therefore different parameter estimates, 616 but merits future work to explore why this may arise. As we only consider one objective 617 function in this analysis, and we do not combine and propagate these parameter estimates back 618 into the observation space, we can only speculate on how these findings may lead to improved 619 calibration strategies. We do note that our findings challenge a common approach where some 620 model parameters are constrained first using a conservative tracer, then fixed and others 621 constrained in a second step using a reactive tracer (e.g., Keefe et al., 2004, Claessens et al., 622 2010, Yakirevich et al., 2017). Lemke et al. (2013) also found differences in optimized 623 parameters for transport when a conservative tracer was fitted alone or jointly with Raz. Thus, 624 our results demonstrate that improved interpretation of BTCs may be aided by fitting 625 conservative and nonconservative tracers separately and comparing parameter estimates, instead 626 of using conservative tracers to constrain parameters associated with nonconservative behavior. 627

628 Within our exercise, the tracer that provided the least redundant information was Rru, 629 which contained unique information regarding MITS processes (α_{MITS}, A_{MITS}). While we 630 anticipated differences between empirical PDFs fit to conservative versus non-conservative 631 tracers, differences were especially pronounced between empirical PDFs for Raz versus Rru. 632 This difference suggests that "smart" tracers may be more useful than conservative tracers for 633 separating the hydrological and biogeochemical impacts of transient storage.

While our study suggests that Raz and Ura provide in part redundant information, we caution that this may not be the case for all systems. Making such a claim of redundancy based on a modeling exercise considering two stream reaches is unrealistic; more studies are needed to resolve questions of redundancy between tracers and parameter information content. Future work, especially experimental observations of transient storage processes (e.g., Knapp et al.,

639 2017), is needed to clarify and investigate timescales of MATS and MITS, and whether these

tracers are truly redundant when it comes to estimating parameter values. This ultimately relies

on improved reconciliation by the TSM community of what is captured by a tracer versus what is

represented within a given TSM formulation. At this stage, we do not have enough information

to assess whether tracer observations may provide complementary or redundant information, as
 such an assessment should be based on numerous paired conservative and non-conservative

645 tracer observations coupled with TSM.

646

5.4 How does model complexity impact parameter estimates and uncertainty?

Regardless of model complexity, the goals of tracer experiments are often to obtain 647 reliable estimates with low uncertainty for parameters describing the influence of transient 648 storage. Our results demonstrate that achieving this objective will ultimately be affected by the 649 650 choice of tracer(s) (e.g., Abbot et al., 2016) and the choice of model framework, including the level of process representation. Increasing model complexity through the addition of model 651 parameters may allow more realistic representation of in- and near-stream processes, but also has 652 653 important implications for parameter uncertainty. In our analysis, we found that parameters typically well-estimated by TSM, A and D, saw wider uncertainty bounds moving from a 1SZ to 654 2SZ formulation (Fig. 6). This is likely due to increased degrees of freedom and interactions 655 with added parameters in the 2SZ formulation (Fig. 8). As with our analysis, other studies have 656 found A and D to be the most sensitive parameters with narrow ranges of uncertainty across 657 658 many TSM applications (Wagener et al., 2002; Kelleher et al., 2013; Ward et al., 2017). These studies have also found strong interactions between A and D, likely the cause of the bimodal 659 behavior observed in Figure 8. Our work adds to this existing body of literature by 660 demonstrating how uncertainty in these well-estimated parameters changes alongside model 661 complexity. When considering these uncertainty bounds in the context of uncertainty for other 662 parameters, differences in these uncertainty bounds were still relatively small, leading us to 663 conclude that only minor inference was lost with increased model complexity. Regardless, this 664 outcome is a good reminder that as parameters are added to a model framework, uncertainty for 665 666 some parameter estimates is likely to grow, even with additional information in the form of added tracer observations. 667

Our study offers cautious optimism regarding use of 2SZ models to infer process-based 668 understanding of solute transport. As we show, 2SZ models, while more complex than 1SZ 669 counterparts, produced narrow estimates of transient storage parameters and showed promise for 670 separating the effects of MITS and MATS. Though parameters were highly interactive within the 671 2SZ model formulation (Figs. 8 and 9), we encouragingly found that we could obtain consistent 672 and precise estimates of transient storage zone parameters (e.g., α_{MITS} , α_{MATS}) that are 673 traditionally dominated by interactions and therefore have proved difficult to estimate in past 674 studies (Wagner and Harvey, 1997; Wagener et al., 2002; Kelleher et al., 2013; Ward et al., 675 2017). However, our results also demonstrate that with increased complexity comes increased 676 uncertainty with respect to other model parameters. Studies utilizing 2SZ models, or any TSM 677 for that matter, should ultimately evaluate the uncertainty associated with parameter estimates 678 (echoing past recommendations; Wagener et al., 2002; Kelleher et al., 2013; Ward et al., 2017). 679 This need for uncertainty evaluation is especially clear in our analysis, in that we demonstrate 680 that while this uncertainty may be reduced for 2SZ as compared to 1SZ models for some 681 scenarios and parameters, uncertainty can still increase for other scenarios and parameters. 682

683 6 Conclusions

While researchers may wish to estimate size and exchange rates associated with transient 684 storage in streams, and further to separate the effects of different transient storage zones, these 685 goals rely on parameter estimation within a TSM framework. Within this context, we explored 686 the tradeoffs between model complexity and utility of novel observations to estimate the effects 687 of transient storage within stream reaches. Our results were consistent across two stream reaches 688 with distinct morphologies; they suggest that model complexity and the necessity for new tracer 689 observations are highly connected. For a 1D TSM, we found that parameter estimates were well-690 constrained by conservative tracer BTCs, but that fitting TSM simulations to a nonconservative 691 tracer (Raz) yielded minimal additional gains in parameter inference. Thus, if using only a 692 conservative tracer, a simpler model may yield more informative parameter estimates. In 693 contrast, estimating parameters within a more complex 2SZ formulation from both conservative 694 and "smart" tracer BTC error metrics produced complimentary insights, suggesting that (if the 695 goal of a given study is to characterize both MITS and MATS) conservative and "smart" tracers 696 should be used in tandem. Our findings suggest cautious optimism that nearly all parameters in 697 2SZ TSM formulations may be capably estimated by jointly fitting simulations to both 698 conservative and "smart" tracer observations. Though our study represents a first step towards 699 this goal, future work is needed to translate evaluations of parameter sensitivity and uncertainty 700 701 into robust approaches to fitting multiple BTCs.

Though we show "smart" tracers have value for improving TSM approaches, we must 702 ultimately reconcile how different process representations within TSM and tracer observations 703 can be used to better quantify and understand specific stream transport processes. This is 704 highlighted by the fact that experiments conducted with "smart" tracers, compared to 705 conservative single tracer studies, require additional instrumentation, consumable costs, field 706 time, and expertise. It remains to be seen whether "smart" tracers provide enough extra 707 information to warrant their use within TSM, given our study solely demonstrates this outcome 708 709 for two reaches with data collected at a single flow state. This detailed model assessment of multiple tracer types from two morphologically distinct stream reaches gives future stream 710 investigators some insights, but, more importantly, quantitatively demonstrates that there are 711 difficult tradeoffs each researcher will face (e.g., tradeoffs between tracer observations and 712 713 model process representation efforts) when conducting stream tracer experiments. Furthermore, if unique information from tracers does not improve our current modeling tools, this may also 714 715 suggest we need to interrogate and refine our perceptual models of these processes with the goal of improving numerical modeling tools. 716

717 The caution we offer, and are even prone to in this work, is that so many TSM analyses are treated as case studies, and there are few TSM synthesis efforts that have examined model 718 frameworks, approaches, and outcomes across multiple sites, flow states, and physical 719 720 representations of transient storage, let alone streams with different types of MATS and MITS. We note that our conclusions are specific to stream setting and flow state, and that there are 721 likely other settings where these findings may differ. Continued discussion and evaluation of 722 723 TSM formulations applied to conservative and nonconservative BTCs is therefore needed to refine the inference we can gain from tracer experiments across different environments, and to 724 deliver a set of defensible recommendations regarding what can be achieved via TSM to the 725 community of ecologists, hydrologists, and biogeochemists that apply these models. 726

Overall, our results validate that novel techniques for hydrologic data collection can help constrain parameter estimates within more complex and potentially more physically realistic models. This progress moves us toward improved process inference within hydrologic modeling of streams. More broadly, the approach we have taken of using gradients of both model

complexity and observations is one that could be adapted and utilized for other hydrological

model-based investigations. By continuing to interrogate the relationships between observations

- and model outcomes, we ultimately have great potential to improve our understanding of
- reactivity and transport within streams, especially when and where disconnects between modeled
- processes and observed processes occur.

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763 **References**

Abbott, B.W., Baranov, V., Mendoza-Lera, C., Nikolakopoulou, M., Harjung, A., Kolbe, T.,

- Balasubramanian, M.N., Vaessen, T.N., Ciocca, F., Campeau, A. & Wallin, M.B. (2016). Using
- 766 multi-tracer inference to move beyond single-catchment ecohydrology. *Earth-Science*
- 767 *Reviews*, *160*, 19-42. https://doi.org/ 10.1016/j.earscirev.2016.06.014.
- Argerich, A., Haggerty, R., Martí, E., Sabater, F., & Zarnetske, J. (2011). Quantification of
- metabolically active transient storage (MATS) in two reaches with contrasting transient storage

- and ecosystem respiration. *Journal of Geophysical Research: Biogeosciences, 116,* G03034.
- 771 https://doi.org/10.1029/2010JG001379.
- 772 Bencala, K.E. & Walters, R.A. (1983). Simulation of solute transport in a mountain pool-and-
- riffle stream: A transient storage model. Water Resources Research, 19, 718–724.
- 774 https://doi.org/10.1029/WR019i003p00718
- Best, M.J., Abramowitz, G., Johnson, H.R., Pitman, A.J., Balsamo, G., Boone, A., Cuntz, M.,
- Decharme, B., Dirmeyer, P.A., Dong, J. & Ek, M. (2015). The plumbing of land surface models:
- benchmarking model performance. *Journal of Hydrometeorology*, *16*(3), 1425-1442.
- 778 https://doi.org/ 10.1175/JHM-D-14-0158.1.
- Beven, K. (2006). A manifesto for the equifinality thesis. *Journal of Hydrology, 320,* 18–36.
 https://doi.org/10.1016/j.jhydrol.2005.07.007
- 781 Beven, K. (1993). Prophecy, reality and uncertainty in distributed hydrological modelling.
- 782 Advances in Water Resources, 16, 41–51. https://doi.org/10.1016/0309-1708(93)90028-E
- 783 Blaen P.J., Khamis K., Lloyd C.E.M., Bradley C., Hannah D. & Krause S. (2016). Real-time
- monitoring of nutrients and dissolved organic matter in rivers: Capturing event dynamics,
- technological opportunities and future directions, *Science of the Total Environment*, 569-570,
 647-660, https://doi.org/10.1016/j.scitotenv.2016.06.116.
- 787 Blaen, P. J., Brekenfeld, N., Comer-Warner, S., & Krause, S. (2017). Multitracer Field
- 788 Fluorometry: Accounting for Temperature and Turbidity Variability During Stream Tracer
- 789 Tests. Water Resources Research, 53, 9118–9126. <u>https://doi.org/10.1002/2017WR020815</u>.
- Blaen, P.J., Kurz, M.J., Drummond, J.D., Knapp, J.L.A., Mendoza-Lera, C., Schmadel, N.M.,
- Klaar, M.J., Jager, A., Folegot, S., Lee-Cullin, J., Ward, A.S., Zarnetske, J.P., Datry, T., Milner,
- A.M., Lewandowski, J., Hannah, D.M., & Krause, S. (2018). Woody debris is related to reach-
- scale hotspots of lowland stream ecosystem respiration under baseflow conditions,
- *Ecohydrology*, Accepted 13 February 2018. https://doi.org/ 10.1002/eco.1952.
- Blaen, P., M. Kurz, J. Knapp, S. Krause, A. Ward, C. Kelleher (2018). 2015 Hammer Solute
- 796 Tracer Injections, HydroShare,
- 797 http://www.hydroshare.org/resource/7c63809428cf495faceb67c0278ed036.
- Boano, F., Harvey, J.W., Marion, A., Packman, A.I., Revelli, R., Ridolfi, L., & Wörman, A.
- (2014). Hyporheic flow and transport processes: Mechanisms, models, and biogeochemical
- 800 implications. *Reviews of Geophysics*, *52*, 2012RG000417.
- 801 https://doi.org/10.1002/2012RG000417
- 802 Brenner, C., Thiem, C.E., Wizemann, H.-D., Bernhardt, M., & Schulz, K. (2017). Estimating
- spatially distributed turbulent heat fluxes from high-resolution thermal imagery acquired with a
- UAV system. International Journal of Remote Sensing, 38, 3003–3026.
- 805 https://doi.org/10.1080/01431161.2017.1280202
- Briggs, M.A., Gooseff, M.N., Arp, C.D., & Baker, M.A. (2009). A method for estimating surface
- transient storage parameters for streams with concurrent hyporheic storage. *Water Resources Research*, 45, W00D27. https://doi.org/10.1029/2008WR006959

- 809 Butts, M.B., Payne, J.T., Kristensen, M., & Madsen, H. (2004). An evaluation of the impact of
- model structure on hydrological modelling uncertainty for streamflow simulation. *Journal of*
- 811 *Hydrology, 298,* 242–266. https://doi.org/10.1016/j.jhydrol.2004.03.042
- 812 Caruso, A., Ridolfi, L., & Boano, F. (2016). Impact of watershed topography on hyporheic
- exchange. *Advances in Water Resources, 94,* 400–411.
- 814 https://doi.org/10.1016/j.advwatres.2016.06.005
- 815 Choi, J., Harvey, J.W., & Conklin, M.H. (2000). Characterizing multiple timescales of stream
- and storage zone interaction that affect solute fate and transport in streams. *Water Resources*
- 817 Research, 36, 1511–1518. https://doi.org/10.1029/2000WR900051
- 818 Clark, M.P., Nijssen, B., Lundquist, J.D., Kavetski, D., Rupp, D.E., Woods, R.A., Freer, J.E.,
- Gutmann, E.D., Wood, A.W., Brekke, L.D. & Arnold, J.R. (2015a). A unified approach for
- process-based hydrologic modeling: 1. Modeling concept. *Water Resources Research*, 51(4),
- 821 2498-2514. https://doi.org/ 10.1002/2015WR017198.
- Clark, M.P., Nijssen, B., Lundquist, J.D., Kavetski, D., Rupp, D.E., Woods, R.A., Freer, J.E.,
- Gutmann, E.D., Wood, A.W., Gochis, D.J. & Rasmussen, R.M. (2015b). A unified approach for
- 824 process-based hydrologic modeling: 2. Model implementation and case studies. *Water Resources*
- 825 *Research*, *51*(4), 2515-2542. https://doi.org/ 10.1002/2015WR017198.
- Clark, M.P., Bierkens, M.F., Samaniego, L., Woods, R.A., Uijlenhoet, R., Bennett, K.E.,
- Pauwels, V.R., Cai, X., Wood, A.W. & Peters-Lidard, C.D. (2017). The evolution of process-
- based hydrologic models: historical challenges and the collective quest for physical
- realism. *Hydrology and Earth Systems Sciences, 21*(7), 3427-3440, https://doi.org/ 10.5194/hess-21-3427-2017.
- Claessens, L., Tague, C.L., Groffman, P.M. and Melack, J.M. (2010). Longitudinal assessment
- of the effect of concentration on stream N uptake rates in an urbanizing
- watershed. *Biogeochemistry*, 98(1-3), 63-74. https://doi.org/10.1002/2012RG000417.
- ⁸³⁴ Damköhler, G. (1936). Einflüsse der Strömung, Diffusion und des Wärmeüberganges auf die
- 835 Leistung von Reaktionsöfen.: I. Allgemeine Gesichtspunkte für die Übertragung eines
- chemischen Prozesses aus dem Kleinen ins Große. Zeitschrift für Elektrochemie und angewandte
- 837 physikalische Chemie, 42, 846–862. <u>https://doi.org/10.1002/bbpc.19360421203</u>.
- 838 Freer, J. Beven, K., & Ambroise, B. (1996). Bayesian estimation of uncertainty in runoff
- prediction and the value of data: an application of the GLUE approach, *Water Resources*
- 840 Research, 32, 2161-2173, https://doi.org/ 10.1029/95WR03723.
- González-Pinzón, R., Haggerty, R., & Myrold, D.D. (2012). Measuring aerobic respiration in
- stream ecosystems using the resazurin-resorufin system. *Journal of Geophysical Research:*
- 843 Biogeosciences, 117, G00N06. <u>https://doi.org/10.1029/2012JG001965</u>.
- Gooseff, M. N., Wondzell, S. M., Haggerty, R., & Anderson, J. (2003). Comparing transient
- storage modeling and residence time distribution (RTD) analysis in geomorphically varied
- reaches in the Lookout Creek basin, Oregon, USA. *Advances in Water Resources*, *26*(9), 925937. Doi: 10.1016/S0309-1708(03)00105-2.
- Gooseff, M. N., R. O. Hall Jr., and J. L. Tank (2007), Relating transient storage to channel
- complexity in streams of varying land use in Jackson Hole, Wyoming, *Water Resour. Res.*, 43,
- 850 W01417, doi: 10.1029/2005WR004626.

- 851 Gooseff, M.N. (2010). Defining Hyporheic Zones Advancing Our Conceptual and Operational
- Definitions of Where Stream Water and Groundwater Meet. *Geography Compass, 4,* 945–955.
- ktps://doi.org/10.1111/j.1749-8198.2010.00364.x
- Haggerty, R., Argerich, A., & Martí, E. (2008). Development of a "smart" tracer for the
- assessment of microbiological activity and sediment-water interaction in natural waters: The
- resazurin-resorufin system. *Water Resources Research, 44*, W00D01.
- 857 https://doi.org/10.1029/2007WR006670
- Haggerty, R., Martí, E., Argerich, A., von Schiller, D., & Grimm, N.B. (2009). Resazurin as a
- ⁸⁵⁹ "smart" tracer for quantifying metabolically active transient storage in stream ecosystems.
- *Journal of Geophysical Research Letters*, *114*, G03014. https://doi.org/10.1029/2008JG000942
- Harvey, J.W., Wagner, B.J., & Bencala, K.E. (1996). Evaluating the Reliability of the Stream
- Tracer Approach to Characterize Stream-Subsurface Water Exchange. *Water Resources Research*, 32, 2441–2451. https://doi.org/10.1029/96WR01268
- 864 Hrachowitz, M. & Clark, M.P. (2017). HESS Opinions: The complementary merits of competing
- modelling philosophies in hydrology. *Hydrology and Earth Systems Sciences*, *21*, 3953–3973.
 https://doi.org/10.5194/hess-21-3953-2017
- Karakashev, D., Galabova, D., & Simeonov, I. (2003). A simple and rapid test for differentiation
- of aerobic from anaerobic bacteria. *World Journal of Microbiology and Biotechnology, 19,* 233–
- 869 238. https://doi.org/10.1023/A:1023674315047
- 870 Keefe, S. H., Barber, L. B., Runkel, R. L., Ryan, J. N., McKnight, D. M., & Wass, R.
- D. (2004). Conservative and reactive solute transport in constructed wetlands. *Water Resources Research*, 40, W01201, https://doi.org/10.1029/2003WR002130.
- Kelleher, C., Wagener, T., McGlynn, B., Ward, A.S., Gooseff, M.N., & Payn, R.A. (2013).
- Identifiability of transient storage model parameters along a mountain stream. *Water Resources Research*, *49*, 5290–5306. https://doi.org/10.1002/wrcr.20413
- Kerr, P.C., Gooseff, M.N., & Bolster, D. (2013). The significance of model structure in one-
- dimensional stream solute transport models with multiple transient storage zones competing vs.
- nested arrangements. *Journal of Hydrology*, 497, 133–144.
- 879 https://doi.org/10.1016/j.jhydrol.2013.05.013
- Knapp, J.L.A., González-Pinzón, R., Drummond, J.D., Larsen, L.G., Cirpka, O.A., & Harvey,
- J.W. (2017). Tracer-based characterization of hyporheic exchange and benthic biolayers in
- streams. Water Resources Research, 53, 1575–1594. <u>https://doi.org/10.1002/2016WR019393</u>.
- Knapp, J.L.A., & Cirpka, O.A. (2017). Determination of hyporheic travel time distributions and
- other parameters from concurrent conservative and reactive tracer tests by local-in-global
- optimization, Water Resources Research, 53, 4984-5001,
- 886 <u>https://doi.org/10.1002/2017WR020734</u>.
- Knapp, J. L. A., González-Pinzón, R., & Haggerty, R. (2018). The resazurin-resorufin system:
- 888 Insights from a decade of "smart" tracer development for hydrologic applications. *Water*
- 889 Resources Research, 54, 6877–6889. <u>https://doi.org/10.1029/2018WR023103</u>.
- Khamis K., Bradley C., Stevens R., & Hannah D.M. (2016). Continuous field estimation of
- dissolved organic carbon concentration and biochemical oxygen demand using dual-wavelength

- fluorescence, *Hydrological Processes Scientific Briefings*, 31, 540-555.
- 893 https:/10.1002/hyp.11040.
- Krause S., Hannah D.M., Fleckenstein J.H., Heppell C.M., Pickup R., Pinay G., Robertson A.L.
- & Wood P.J. (2011). Interdisciplinary perspectives on processes in the hyporheic zone,
- *Ecohydrology*, *4*, 481–499. https://10.1002/eco.176
- Lees, M.J., Camacho, L.A., & Chapra, S. (2000). On the relationship of transient storage and
- aggregated dead zone models of longitudinal solute transport in streams. *Water Resources*
- *Research, 36*, 213–224. https://doi.org/10.1029/1999WR900265
- 900 Lemke, D., Liao, Z., Wöhling, T., Osenbrück, K., & Cirpka, O.A. (2013). Concurrent
- 901 conservative and reactive tracer tests in a stream undergoing hyporheic exchange. *Water*
- 902 *Resources Research, 49,* 3024–3037. https://doi.org/10.1002/wrcr.20277
- Lemke, D., González-Pinzón, R., Liao, Z., Wöhling, T., Osenbrück, K., Haggerty, R., and
- ⁹⁰⁵ Cirpka, O. A. (2014). Sorption and transformation of the reactive tracers resazurin and resorufin
- 906 in natural river sediments, *Hydrol. Earth Syst. Sci.*, 18, 3151-3163, https://doi.org/10.5194/hess-18-3151-2014.
- ⁹⁰⁷ Li, H., Xu, C.-Y., & Beldring, S. (2015). How much can we gain with increasing model
- complexity with the same model concepts? *Journal of Hydrology*, *527*, 858–871.
- 909 https://doi.org/10.1016/j.jhydrol.2015.05.044
- Liao, Z., & Cirpka, O.A. (2011). Shape-free inference of hyporheic traveltime distributions from
- synthetic conservative and "smart" tracer tests in streams. *Water Resources Research, 47,*
- 912 W07510. https://doi.org/10.1029/2010WR009927
- 913 Marion, A., Zaramella, M., & Bottacin-Busolin, A. (2008). Solute transport in rivers with
- multiple storage zones: The STIR model. *Water Resources Research, 44*, W10406.
- 915 https://doi.org/10.1029/2008WR007037
- Nearing, G. S., & Gupta, H. V. (2015). The quantity and quality of information in hydrologic
 models. *Water Resources Research 51*, 524–538. https://doi.org/10.1002/2014WR015895.
- 918 Nearing, G. S., Mocko, D. M., Peters-Lidard, C. D., Kumar, S. V., & Xia, Y. L. (2016).
- 919 Benchmarking NLDAS-2 Soil Moisture and Evapotranspiration to Separate Uncertainty
- Contributions. *Journal of Hydrometeorology*, *17*(3), 745-759. <u>https://doi.org/10.1175/jhm-d-15-</u>
 0063.1.
- Orghidan, T., (1959). Ein neuer Lebensraum des unterirdischen Wasser: der hyporheische
 Biotop. *Archiv für Hydrobilogie*, *55*, 392–414.
- 924 Pianosi, F., SarRazin, F., & Wagener, T. (2015). A Matlab toolbox for global sensitivity
- analysis. Environmental Modelling & Software, 70, 80-85, https://doi.org/
- 926 10.1016/j.envsoft.2015.04.009.
- 927 Runkel, R. L. (1998). One-Dimensional Transport with Inflow and Storage (OTIS): A Solute
- Transport Model for Streams and Rivers (No. Water-Resources Investigations Report 98-4018).
 United States Geological Survey.
- Runkel, R.L. (2015). On the use of rhodamine WT for the characterization of stream
- hydrodynamics and transient storage. *Water Resources Research*, 51, 6125–6142.
- 932 https://doi.org/10.1002/2015WR017201

- 933 Schmadel, N.M., Ward, A.S., & Wondzell, S.M. (2017). Hydrologic controls on hyporheic
- exchange in a headwater mountain stream. *Water Resources Research 53*, 6260–6278.
- 935 https://doi.org/10.1002/2017WR020576
- 936 Schoups, G., van de Giesen, N.C., & Savenije, H.H.G. (2008). Model complexity control for
- hydrologic prediction. *Water Resources Research, 44*, W00B03.
- 938 https://doi.org/10.1029/2008WR006836.
- 939 Seibert, J., & J. J. McDonnell (2002). On the dialog between experimentalist and modeler in
- 940 catchment hydrology: Use of soft data for multicriteria model calibration. *Water Resources*
- 941 *Research*, 38(11), 1241, https://doi.org/10.1029/2001WR000978.
- 942 Shelley, F., Klaar, M., Krause, S., & Trimmer, M. (2017). Enhanced hyporheic exchange flow
- around woody debris does not increase nitrate reduction in a sandy streambed.
- 944 Biogeochemistry, 136, 353-372, https://doi.org/ 10.1007/s10533-017-0401-2.
- 945 St Clair, J., Moon, S., Holbrook, W.S., Perron, J.T., Riebe, C.S., Martel, S.J., Carr, B., Harman,
- C., Singha, K., & Richter, D. deB (2015). Geophysical imaging reveals topographic stress
- ontrol of bedrock weathering. *Science*, *350*, 534–538. https://doi.org/10.1126/science.aab2210
- 948 Storey, R.G., Howard, K.W.F., & Williams, D.D. (2003). Factors controlling riffle-scale
- 949 hyporheic exchange flows and their seasonal changes in a gaining stream: A three-dimensional
- groundwater flow model. *Water Resources Research*, 39, 1034.
- 951 https://doi.org/10.1029/2002WR001367
- Thackston, E.L., & Schnelle, K.B. (1970). Predicting Effects of Dead Zones on Stream Mixing.
 Journal of the Sanitary Engineering Division, *96*, 319–331.
- Triska, F.J., Kennedy, V.C., Avanzino, R.J., Zellweger, G.W., & Bencala, K.E. (1989).
- 955 Retention and Transport of Nutrients in a Third-Order Stream in Northwestern California:
- 956 Hyporheic Processes. *Ecology*, *70*, 1893–1905. https://doi.org/10.2307/1938120
- 957 Valett, H.M., Morrice, J.A., Dahm, C.N., & Campana, M.E. (1996). Parent lithology, surface-
- groundwater exchange, and nitrate retention in headwater streams. *Limnology and*
- 959 Oceanography, 41, 333–345. https://doi.org/10.4319/lo.1996.41.2.0333
- 960 Vivoni, E.R., Rango, A., Anderson, C.A., Pierini, N.A., Schreiner-McGraw, A.P., Saripalli, S., &
- Laliberte, A.S. (2014). Ecohydrology with unmanned aerial vehicles. *Ecosphere 5*, 1–14.
 https://doi.org/10.1890/ES14-00217.1
- Wagener, T., Camacho, L.A., & Wheater, H.S. (2002). Dynamic identifiability analysis of the transient storage model for solute transport in rivers. *Journal of Hydroinformatics 4*, 199–211.
- 965 Wagener, T., & Kollat, J. (2007). Numerical and visual evaluation of hydrological and
- 966 environmental models using the Monte Carlo analysis toolbox. *Environmental Modelling &*
- 967 Software, 22, 1021–1033. https://doi.org/10.1016/j.envsoft.2006.06.017.
- 968 Wagner, B.J., & Harvey, J.W. (1997). Experimental design for estimating parameters of rate-
- limited mass transfer: Analysis of stream tracer studies. *Water Resources Research 33*, 1731–
- 970 1741. https://doi.org/10.1029/97WR01067.
- Ward, A.S. (2016). The evolution and state of interdisciplinary hyporheic research. *WIREs*
- 972 *Water, 3,* 83–103. https://doi.org/10.1002/wat2.1120.

- Ward, A.S., Gooseff, M.N., & Singha, K. (2010). Imaging hyporheic zone solute transport using electrical resistivity. *Hydrologic Processes, 24,* 948–953. https://doi.org/10.1002/hyp.7672.
- 975 Ward, A.S., Cwiertny, D.M., Kolodziej, E.P., & Brehm, C.C. (2015). Coupled reversion and
- stream-hyporheic exchange processes increase environmental persistence of trenbolone
- 977 metabolites. *Nature Communications, 6,* 7067. https://doi.org/10.1038/ncomms8067.
- Ward, A.S., Kelleher, C.A., Mason, S.J.K., Wagener, T., McIntyre, N., McGlynn, B., Runkel,
- 879 R.L., & Payn, R.A. (2017). A software tool to assess uncertainty in transient-storage model
- parameters using Monte Carlo simulations. *Freshwater Science*, *36*, 195–217.
- 981 https://doi.org/10.1086/690444.
- Wlostowski, A.N., Gooseff, M.N., & Wagener, T. (2013). Influence of constant rate versus slug
 injection experiment type on parameter identifiability in a 1-D transient storage model for stream
- solute transport. *Water Resources Research*, 49, doi: 10.1002/wrcr.20103.
- Wörman, A., & Wachniew, P. (2007). Reach scale and evaluation methods as limitations for
- transient storage properties in streams and rivers. *Water Resources Research*, 43, W10405.
 https://doi.org/10.1029/2006WR005808.
- 988 Yakirevich, A., Shelton, D., Hill, R., Kiefer, L., Stocker, M., Blaustein, R., Kuznetsov, M.,
- 989 McCarty, G., & Pachepsky, Y. (2017). Transport of Conservative and "smart" Tracers in a First-
- 990 Order Creek: Role of Transient Storage Type. *Water*, *9*, 485. <u>https://doi.org/10.3390/w9070485</u>.

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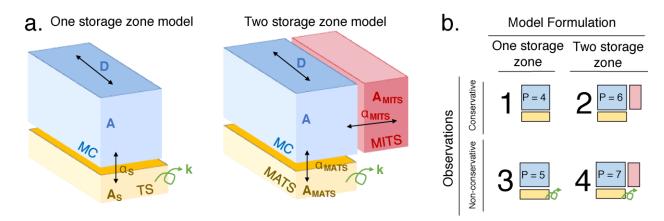
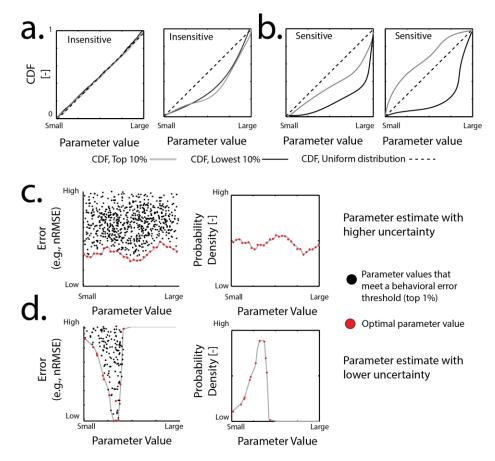
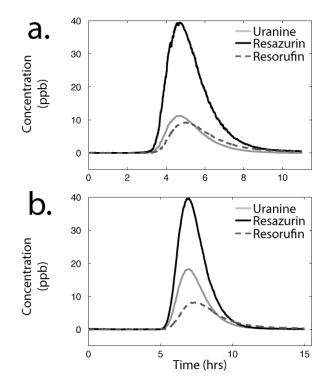


Figure 1. Model framework displaying (a) model parameters and the hypothetical compartments 995 within the stream reach they are associated with (MC = main channel) for both one storage zone 996 (1SZ) and two storage zone (2SZ) models and (b) the multiple model formulations utilized 997 998 within this study (and corresponding numbers of parameters). In particular, we compare across the number of transient storage (TS) zones (one vs. two), as well as parameter estimates with 999 respect to both conservative (Ura) or nonconservative (Raz, Rru) tracer dynamics. By combining 1000 these formulations and observations, we tested four different models ranging from four to seven 1001 model parameters (P). Additional figure abbreviations include: metabolically inactive storage 1002 (MITS), metabolically active storage (MATS), and parameters main-channel area (A), dispersive 1003 1004 coefficient (D), transient storage zone exchange (α_s), transient storage zone size (A_s), conversion of Raz to Rru (k), MATS cross-sectional area (A_{MATS}), MATS exchange rate (α_{MATS}), MITS 1005 cross-sectional area (F_{MITS}), and MITS exchange rate (α_{MITS}). 1006



1009 **Figure 2.** Interpretation and approaches for (a, b) sensitivity analysis and (c, d) uncertainty 1010 analysis. Regional sensitivity analysis is used to assess parameter sensitivity for parameter values with the best (top 10%) and worst (lowest 10%) errors, compared to a uniform distribution (1:1) 1011 1012 line. Conceptual examples of cumulative distribution functions can be used to interpret whether parameters are insensitive (Fig. 2a), due to either falling along the 1:1 line or CDFs 1013 indistinguishable between parameter values corresponding to the best and worst error values, or 1014 sensitive (Fig. 2b), where the CDF of parameter values corresponding to the best errors are 1015 clearly distinguishable from the 1:1 line and the CDF corresponding to the worst errors. To 1016 compliment RSA, uncertainty is assessed by translating dotty plots to empirical probability 1017 density functions (PDFs) of optimal model errors across feasible parameter ranges. Optimal 1018 parameters (red) represent those with the lowest error for a narrow moving window along the 1019 parameter space. We display two hypothetical examples for a parameter with high uncertainty 1020 (Fig. 2c) and low uncertainty (Fig. 2d). Peaky distributions, found for a parameter with low 1021 uncertainty, indicate that certain regions of the parameter space yield better performance, while a 1022 flat distribution, corresponding to the parameter with greater uncertainty (Fig. 2d), suggests that 1023 1024 all parameter values yield similar model performance. 1025



- **Figure 3.** Observed breakthrough curves (concentration through time) for a conservative tracer
- 1028 (Ura) and nonconservative tracer Raz and biproduct Rru for (a) sand and (b) gravel reaches.

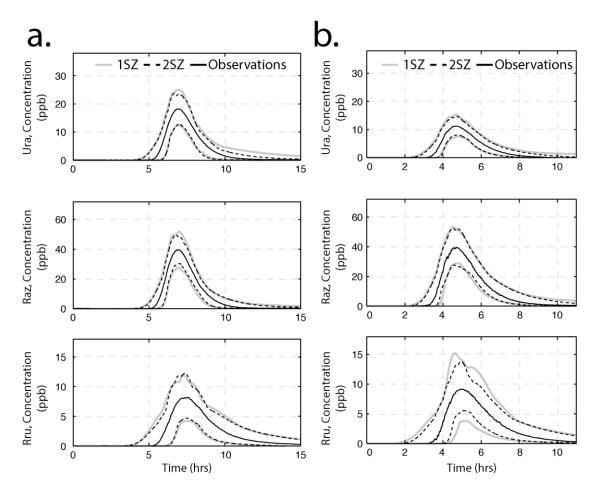
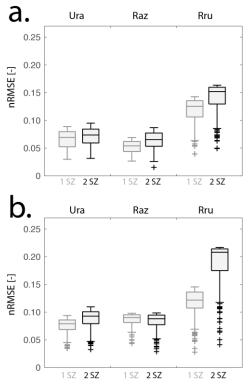


Figure 4. Upper and lower bounds for the 270 simulations corresponding to the minimum 1% of
RMSE values for the (a) gravel and (b) sand reaches. These bounds represent the envelope
encompassing the range of all simulations corresponding to the top 1% of values by *nRMSE*, per
tracer and per model. Bounds are shown relative to observations. All simulations are included as
ensemble averages in Figure S2.



1039 Figure 5. Distributions of model error shown for the top 1% of nRMSE values for all tracers and

- 1040 for a combined tracer metric for the (a) sand and (b) gravel reaches. Results are shown for the (107) (107) (107)
- 1041 one (1SZ) and two storage zone models (2SZ).

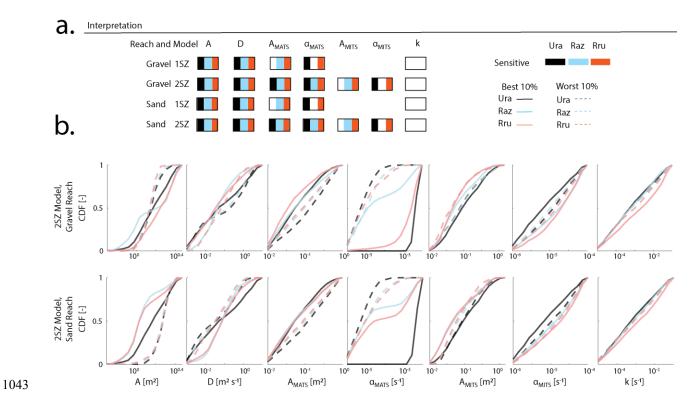


Figure 6. Analysis of parameter sensitivities including (a) interpretation of sensitivities across 1045 1SZ and 2SZ models, reaches, and tracers, and (b) select RSA plots for 2SZ gravel and sand 1046 reaches for D, A_{MATS} , α_{MATS} , A_{MITS} , and α_{MITS} . Interpretation of (a) is based on Fig. 2a, with 1047 sensitive parameters deviating from a uniform CDF and from the CDF corresponding to the 1048 "worst 10%" of error values. A color shown in (a) indicates interpretation based on (b) that a 1049 parameter is sensitive. RSA plots compare empirical CDFs corresponding to the top 10% and 1050 worst 10% of all model simulations per tracer error metric.

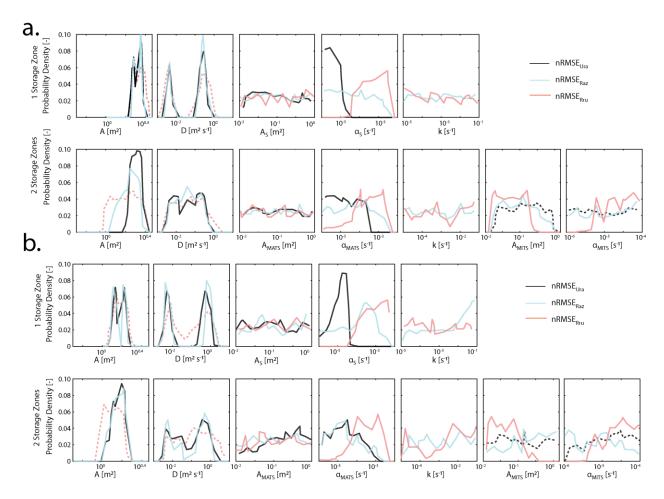


Figure 7. PDFs of the top 1% of RMSE values plotted across log-transformed parameter values for the 1SZ and 2SZ models of the (a) sand and (b) gravel reaches. Results were independently generated for each of three tracers (Fig. 2). Dotted lines indicate parameters that we do not expect to be physically related to or informed by a given tracer.

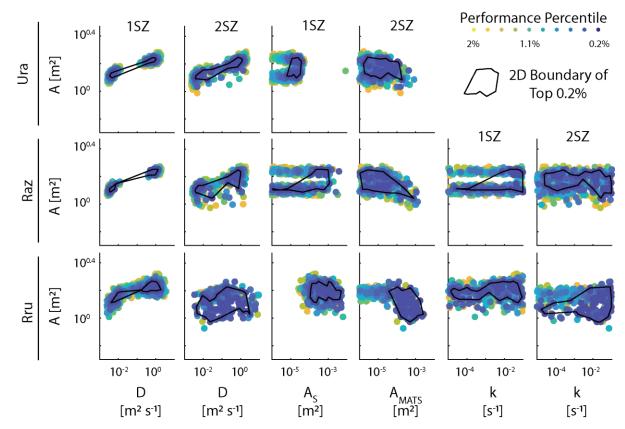
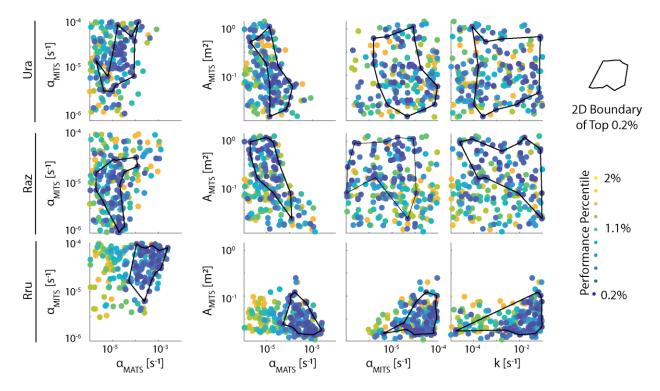


Figure 8. Joint distributions of 1SZ and 2SZ model parameters for the gravel reach. Black lines
 indicate the boundary of the top 0.2% of parameter sets (by *nRMSE* per tracer). Colors indicate
 different percentiles of performance corresponding to the top 2% of all parameter sets.



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Figure 9. Joint distributions for the 2SZ gravel reach parameter sets. Black lines indicate the boundary of the top 0.2% of parameter sets (by nRMSE per tracer). Colors indicate different

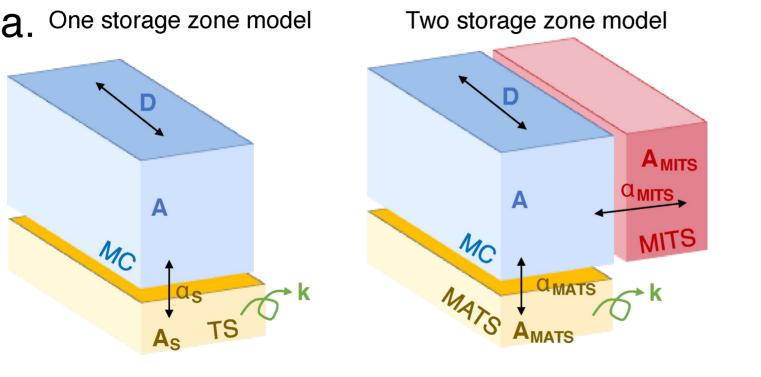
1066 percentiles of performance corresponding to the top 2% of all parameter sets.

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1069	Table 1. Parameter names, abbreviations, and ranges for sensitivity and uncertainty analysis
1070	applied to variable TSM formulations (Figure 1).

Abbrev.	Parameter	Model	Tracer	Units	Lower Bound	Upper Bound
D	Dispersion coefficient	1SZ, 2SZ	All	$m^2 s^{-1}$	0.001	10
А	Advective channel cross- sectional area	1SZ	All	m ²	1	3
$\mathbf{A}_{\mathrm{TOT}}$	Total area	2SZ	All	m ²	1	3
A _S	Transient storage cross- sectional area	1SZ	All	m ²	0.01	1
$\alpha_{\rm S}$	Transient storage exchange rate	1SZ	All	s ⁻¹	10-6	10 ⁻²
k	Conversion, Raz to Rru	2SZ	Raz, Rru	s^{-1}	10 ⁻⁵	10 ⁻¹
A _{MATS}	MATS cross-sectional area	2SZ	All	m^2	0.01	1
α_{MATS}	MATS exchange rate	2SZ	All	s^{-1}	10-6	10 ⁻²
F _{MITS}	Fraction of stream area as MITS	2SZ	All	-	0.01	0.5
α_{MITS}	MITS exchange rate	2SZ	All	s ⁻¹	10 ⁻⁵	10-1

Figure 1.



Observations

D.

Non-conservative

Conservative

Model Formulation One storage Two storage zone zone P = 6P = 4 P = 5P = 7

Figure 2.

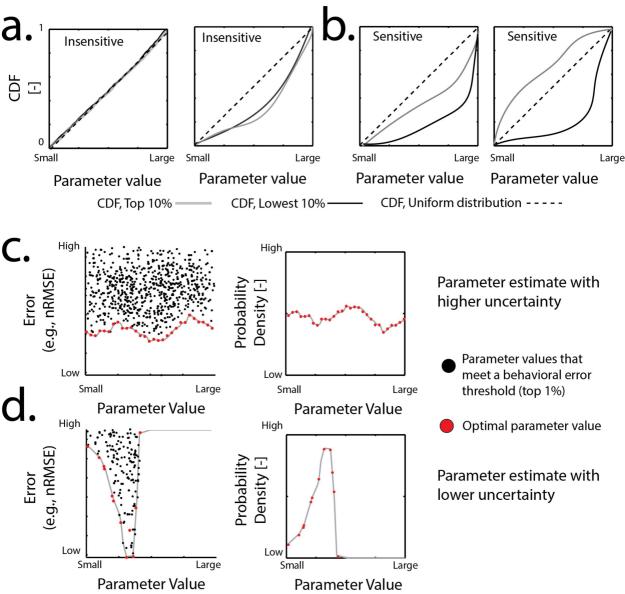


Figure 3.

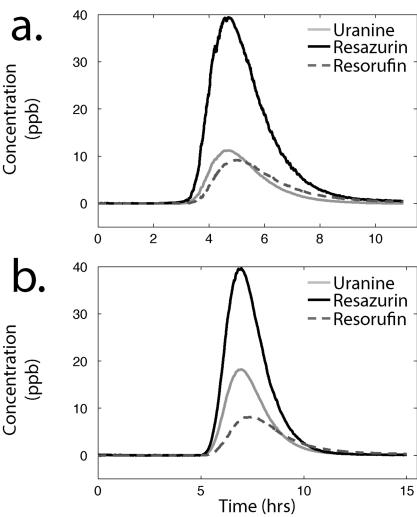


Figure 4.

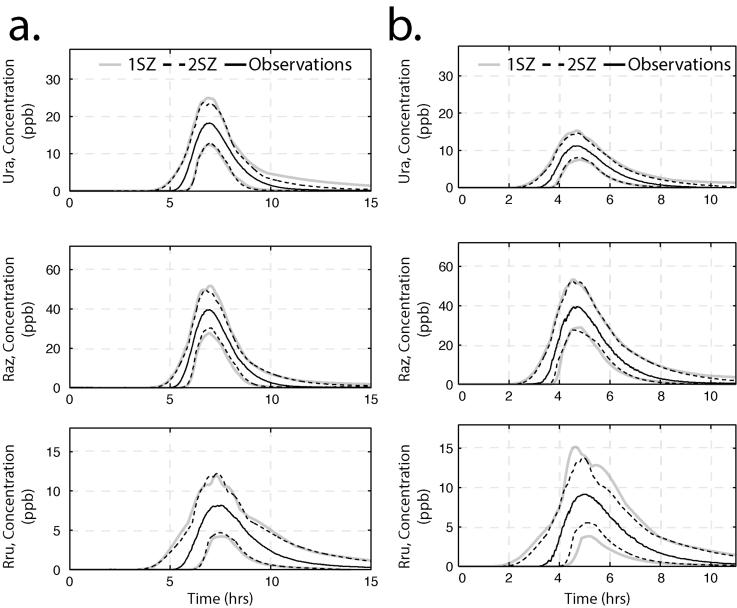


Figure 5.

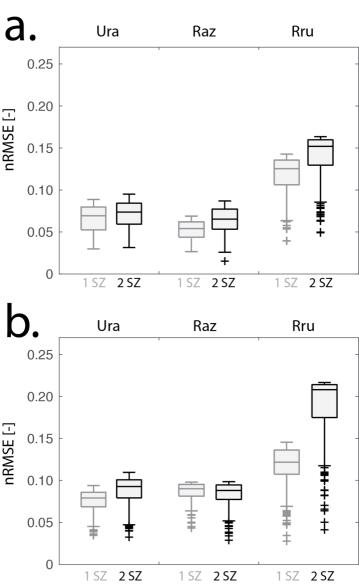


Figure 6.

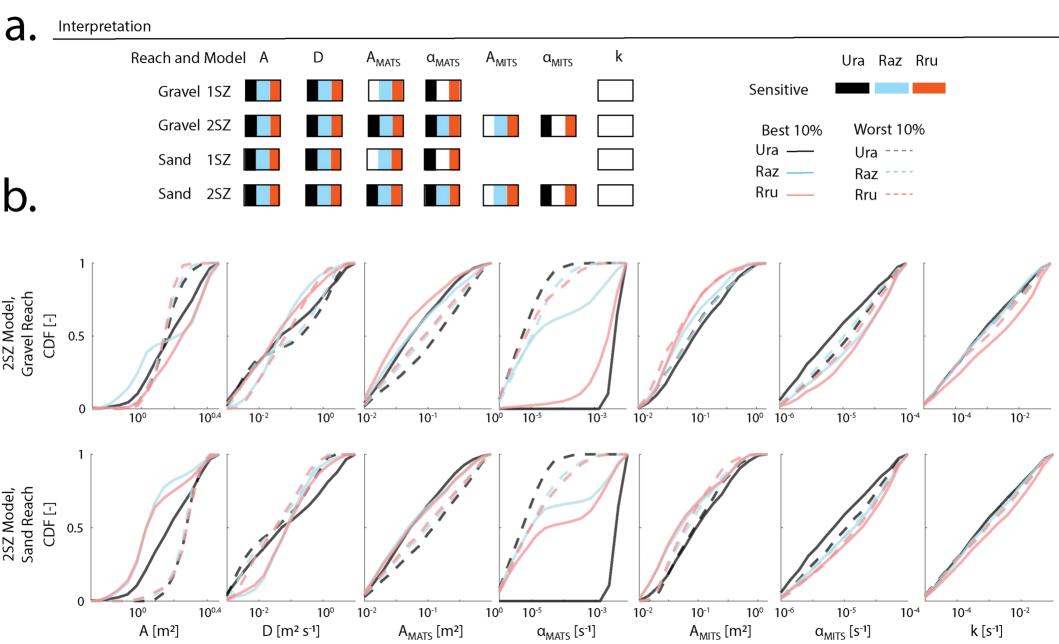


Figure 7.

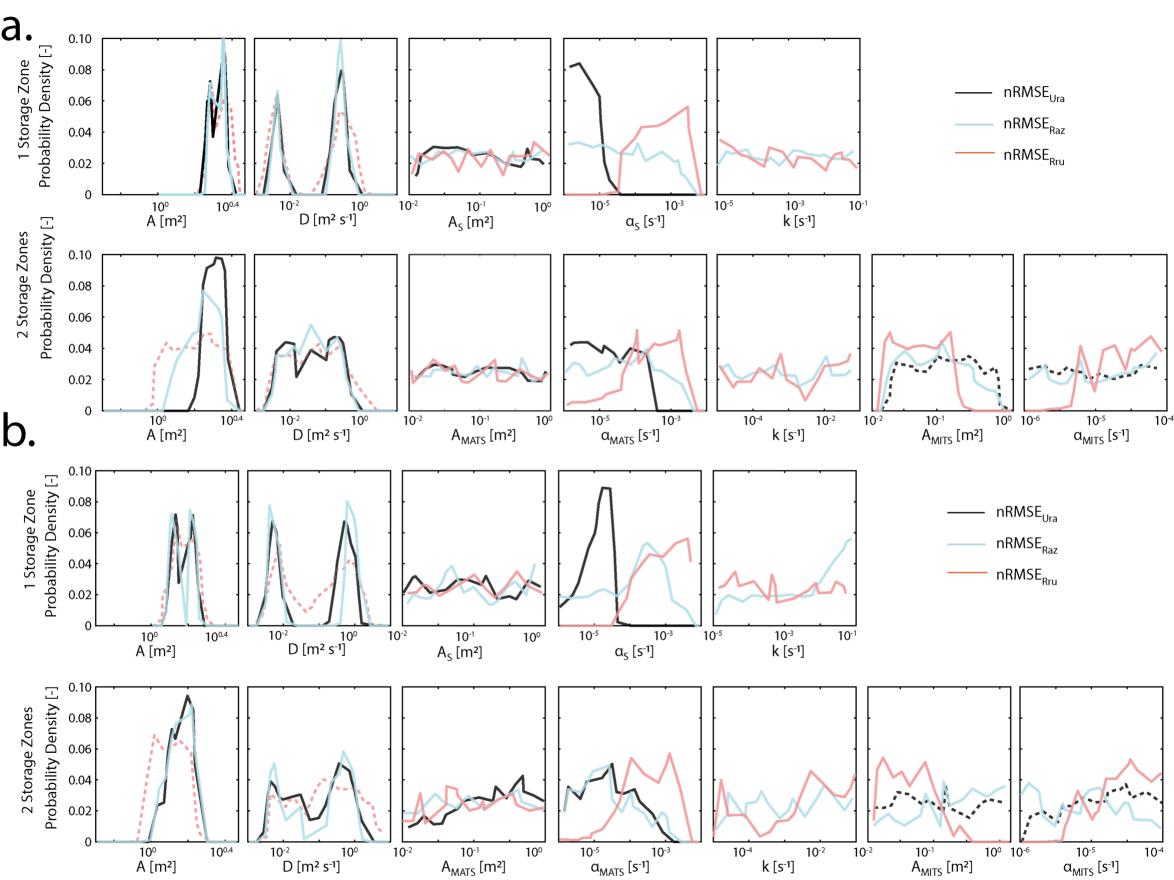


Figure 8.

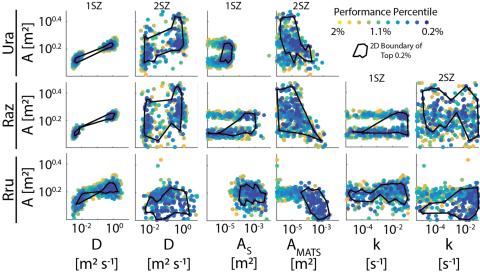


Figure 9.

