

Direct oxygen isotope effect identifies the rate-determining step of electrocatalytic OER at an oxidic surface

Sandra Haschke,^a Michael Mader,^b Stefanie Schlicht,^a André M. Roberts,^b
Alfredo M. Angeles-Boza,^{c,*} Johannes A.C. Barth,^{b,*} Julien Bachmann^{a,d,*}

^a *Friedrich-Alexander-Universität Erlangen-Nürnberg, Department of Chemistry and Pharmacy, Chemistry of Thin Film Materials, Cauerstr. 4, 91058 Erlangen, Germany*

^b *Friedrich-Alexander-Universität Erlangen-Nürnberg, Department für Geographie und Geowissenschaften, GeoZentrum NordBayern, Applied Geology, Schlossgarten 5, 91054 Erlangen, Germany*

^c *University of Connecticut, Department of Chemistry and Institute of Materials Science, 55 North Eagleville Rd., Storrs, CT 06269, USA*

^d *Saint-Petersburg State University, Institute of Chemistry, Universitetskii pr. 26, 198504 St. Petersburg, Russia*

** Authors to whom correspondence should be addressed:*

alfredo.angeles-boza@uconn.edu

johannes.barth@fau.de

julien.bachmann@fau.de

Abstract

Understanding the mechanism of water oxidation to dioxygen represents the bottleneck towards the design of efficient energy storage schemes based on water splitting. The investigation of kinetic isotope effects has long been established for mechanistic studies of various such reactions. However, so far natural isotope abundance determination of O₂ produced at solid electrode surfaces has not been applied. Here, we demonstrate that such measurements are possible. Moreover, they are experimentally simple and sufficiently accurate to observe significant effects. Our measured kinetic isotope effects depend strongly on the electrode material and on the applied electrode potential. They suggest that in the case of iron oxide as the electrode material, the oxygen evolution reaction occurs via a rate-determining O—O bond formation via nucleophilic water attack on a ferryl unit.

Introduction

The direct electrochemical splitting of water into its constituent elements represents a non-polluting method of producing storable energy carrier from renewable sources, such as wind, hydroelectric or solar power. This power-to-fuel conversion involves two half-reactions: the four-electron, four-proton oxidation of water to dioxygen and the reduction of protons to dihydrogen.^{1,2} The mechanistically more demanding and kinetically limiting half-reaction here is the generation of O₂, since it involves a proton and electron currency of four as well as the formation of the oxygen–oxygen double bond.^{3,4} Thus, the overall water splitting activation energy requirement is defined, beyond the thermodynamic potential difference between the redox pairs ($\Delta E = +1.23$ V), mostly by the overpotential of the oxygen evolution reaction (OER,

whereas the contribution associated to the overpotential of the hydrogen evolution reaction is significantly smaller).² The list of materials evaluated for their potential use as OER catalysts is promising for the future and includes, rare noble metal oxides,⁵ earth-abundant first-row transition metal oxides and (oxy)hydroxides.⁶⁻⁸ However, de novo designs of completely original future catalysts must rely on the fundamental comprehension of possible OER mechanisms, including the unambiguous identification of their key (rate-determining) step. Thus, the elucidation of individual steps, intermediates and transition states represents one of the central challenges towards a broadly applicable power-to-fuel scheme via water electrolysis.

So far, studies on the identification of reaction intermediates at solid anode and photo-anode surfaces have relied on a variety of indirect techniques, such as electrochemical impedance spectroscopy (IS),⁹ Tafel plot analysis,^{1,10} photoinduced absorption spectroscopy,¹¹ time-resolved Fourier-transform infrared (FT-IR) spectroscopy,¹² attenuated total reflection infrared (ATR-IR) spectroscopy,¹³ Raman spectroscopy,¹⁴ and transient absorption spectroscopy,^{15,16} complemented by density functional theory (DFT) calculations.^{17,18}

Recently, iron oxide (Fe_2O_3) electrodes have received much attention by Peter,^{19,20} Hamann,^{9,13} Durrant,^{11,15,16} and their co-workers. In general, they have agreed on the first step of water oxidation involving the formation of an $\text{Fe}^{\text{IV}}=\text{O}$ species (ferryl) via oxidation of the initially hydroxylated Fe^{III} surface (see the mechanistic overview in **Figure 1**). This is in part based on the similarity to other molecular systems amenable to investigation in a homogeneous (dissolved) phase. The steps following the ferryl formation are, however, still under discussion and probably vary depending on several parameters such as temperature, pH of electrolyte solution and nature of the electrolyte.^{11,13,18,21}

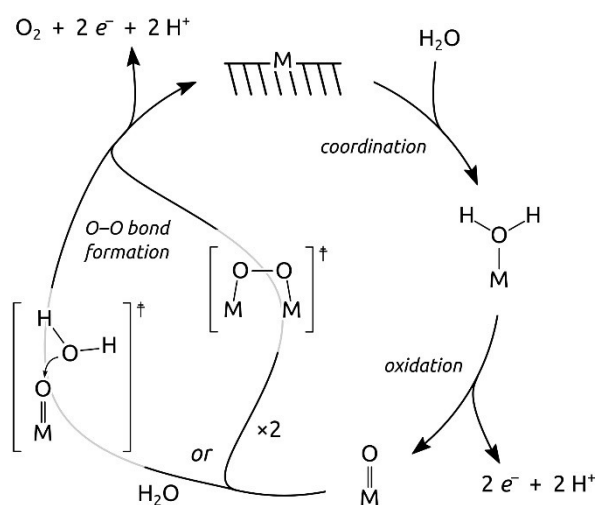


Figure 1. Minimal mechanism of the OER at a solid metal oxide electrode. While coordination of water is fast, the oxidation to a high-valent metal oxo (which occurs in several steps) or the formation of an O—O bond (which can take place on either of two distinct pathways) can, depending on the system and conditions considered, represent the slowest and rate-determining step.

A direct approach to the identification of rate-determining steps (RDS) in molecular catalysis (such as the photosystem II complex, its models, or artificial OER catalysts) has been the analysis of kinetic isotope effects (KIEs):²²⁻²⁹ The mass of an atom affects the overall reaction kinetics if a bond in its vicinity is made or broken in the RDS. KIEs are largest, and have been exploited most extensively, in the case of $^2\text{H}/^1\text{H}$.³⁰⁻³² Isotope effects also exist for $^{18}\text{O}/^{16}\text{O}$. They directly probe the formation of the O—O bond and are therefore particularly well suited to the study of OER. However, they have been determined in a significantly smaller number of cases.^{25-27,33} Determinations of ^{18}O KIE have shown that contrasting reaction steps are rate-

determining in the photosynthetic activity of various microorganisms, which feature near-unity values of ^{18}O KIE (0.993 to 1.001) and in documented synthetic OER catalysts.^{22,24,34-36} KIE determinations have been pivotal in refuting the original assumption that the O—O bond formation via symmetric coupling of two vicinal high-valent metal oxos is the RDS in catalysis by the blue dimer family of molecular ruthenium complexes.^{28,29,37-40}

Thus, the ^{18}O KIE could crucially contribute to the design of more suitable OER catalysts. However, the spectrum of its practical applications has been limited by three significant experimental constraints. First, the number of ^{18}O KIE studies on heterogeneous systems is limited to a handful of examples with the most recent dating back to 1987.⁴¹⁻⁴⁵ Second, ^{18}O KIEs have mostly been determined via density measurements,^{41,42} later via mass spectrometric analysis of CO_2 after combustion⁴³ and rarely on the O_2 molecule itself. Third, isotopic labeling as usually performed for the determination of ^2H KIEs, is significantly less feasible as commercial sources are expensive.⁴⁶

The direct determination of small isotope composition changes on a natural abundance level between a substrate and its product has been performed only in a very small number of cases in the field of photosynthetic water oxidation.^{22,34,47} Two studies reported natural-abundance measurements by isotope ratio mass spectrometry (IRMS) directly of the O_2 product in heterogeneous (electrocatalytic) systems.^{44,45} However, their experimental limitations have prevented the investigation of reaction mechanisms and/or RDS at solid OER electrocatalyst surfaces. (In one case, BrF_3 as the oxidant is described as the source of an unknown systematic error,⁴⁴ whereas in the other study the KIE shows error bars of ± 0.007 for variations from unity of ± 0.025 .)⁴⁵

Here, we present a direct IRMS $^{18}\text{O}/^{16}\text{O}$ effect study of the oxygen evolution reaction at nanoporous Fe_2O_3 and Ir electrode surfaces under steady-state conditions in an aqueous electrolyte. These samples are chosen for their relatively high electrocatalytic activity (caused by the geometric surface area increase), which is necessary for the isotope analyses. We work at natural abundance levels in continuous flow mode at $m/z = 32$ from small sample volumes (12 mL),⁴⁸ without preliminary conversion to CO_2 .^{49,50} The instrumental error on repeat standard measurements is ± 0.0002 . Repeat measurements of real samples yield a maximum standard deviation of ± 0.0015 . Experimentally determined changes of natural abundance $^{18}\text{O}/^{16}\text{O}$ ratios of dissolved O_2 (DO) formed from water yield kinetic isotope effects (KIE) significant enough to differentiate between various catalysts and draw mechanistic conclusions. The data hint at a nucleophilic attack of water on a high-valent surface ferryl as the rate-determining step of water oxidation at iron(III) oxide surfaces.

Results

Electrodes and electrolysis

Nanoporous iron oxide and oxide-covered iridium electrodes were prepared by coating a highly ordered array of anodic alumina pores with a thin layer of approximately 10 nm Fe_2O_3 or 20 nm Ir via atomic layer deposition (ALD).⁵¹⁻⁵³ These electrodes were used in steady-state electrolyses at different applied potentials, that is, different driving forces in pH 7 phosphate buffer (for Fe_2O_3) or acidic sulphuric acid electrolyte (for Ir). Let us focus on the Fe_2O_3 system first. Care was taken to exclude any contamination by atmospheric O_2 . The electrochemical cell was degassed prior to the experiment and properly sealed as confirmed by a control experiment in which the working electrode was left in open circuit (**Supplementary Fig. 1** in the Supplementary Information, dashed line). Upon electrolysis O_2 evolved continuously, quantified by an optical oxygen sensor. **Supplementary Fig. 1** shows how the evolved O_2

concentration follows the Faradaic current for an applied potential E of +1.30 V vs. Ag/AgCl (overpotential $\eta = 0.68$ V).

For subsequent isotope ratio mass spectrometric (IRMS) analyses of the DO, 12 mL of electrolyte were extracted from the electrochemical cell and filled into helium-flushed vials with butyl rubber caps. Depending on the overpotential ($0.48 \leq \eta \leq 0.78$ V), the concentrations of DO evolved upon electrolysis obtained by IRMS lie between $5 \mu\text{mol L}^{-1}$ and $50 \mu\text{mol L}^{-1}$. They agree, within experimental uncertainty, with the values determined by an optical sensor (**Supplementary Fig. 2**). The measured c_{DO} curve reveals the typical exponential dependence of electrocatalytic current density J on applied overpotential η (**Supplementary Fig. 3**). The values lie in the expected range based on the current passed and the partitioning of O_2 into the gas phase (see **Supplementary Note 1**).

Isotope composition of evolved oxygen

The isotope composition of DO in the aqueous electrolyte is expressed by ratios of the stable oxygen isotopes ^{18}O and ^{16}O in the standard δ -notation in per mille (‰) normalised to Vienna Standard Mean Ocean Water.³⁶ As controls, IRMS analyses of aerobic O_2 dissolved in the pH 7 phosphate electrolyte and of the associated H_2O yielded $\delta^{18}\text{O}_{\text{DO,atm}}$ and $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ values of +24.6‰ and -9.8‰, respectively (+24.7‰ and -9.6‰ in the sulphuric acid electrolyte). These values agree with the literature.^{36,54-56} **Figure 2** shows the influence of applied overpotentials on the isotopic compositions of the evolved DO after 3 h of electrolysis. From a starting value of $\delta^{18}\text{O}_{\text{DO}} = 21.0$ ‰ at 0.48 V, a significant decrease in $\delta^{18}\text{O}_{\text{DO}}$ is observed down to -9.6‰ at $\eta = 0.78$ V. Thus, the applied overpotential (i.e., the external activation energy) has a profound influence on $\delta^{18}\text{O}_{\text{DO}}$.

In control experiments, we checked that the electrolysis duration has negligible influence on the isotopic composition of the product (**Supplementary Fig. 4**), proving again the lack of adventitious atmospheric O_2 . Furthermore, the fractionation effect worth 0.7‰ of physical O_2 transfer from the dissolved to gaseous phase⁵⁷ is not significant with respect to the effects measured experimentally. Thirdly, taking into account the O_2 contamination visible in **Supplementary Fig. 1** (DO concentration of $3 \mu\text{mol L}^{-1}$ or 0.1 ppm after degassing) would affect our $\delta^{18}\text{O}_{\text{DO}}$ value by a maximum of -1.9‰ (the detailed calculation and values are presented in the Electronic Supplementary Information, included in **Supplementary Table 1** and **Supplementary Note 2**). This number is significantly smaller than the values of $\delta^{18}\text{O}_{\text{DO}}$ and differences between them determined in this work. Thus, measuring the DO is representative of the whole O_2 generated at the electrode. The fourth and final possible source of error that we considered is the possibility that O_2 produced at the working electrode be reduced back to water at the counter-electrode. This would affect the isotopic composition of DO in an uncontrolled manner. This possibility is excluded by our experimental measurement of the potential difference between working electrode and counter-electrode: since the potential difference is larger than +1.10 V in all conditions of bulk electrolysis used here, we can positively conclude that the reaction occurring at the counter-electrode is the reduction of water to H_2 .

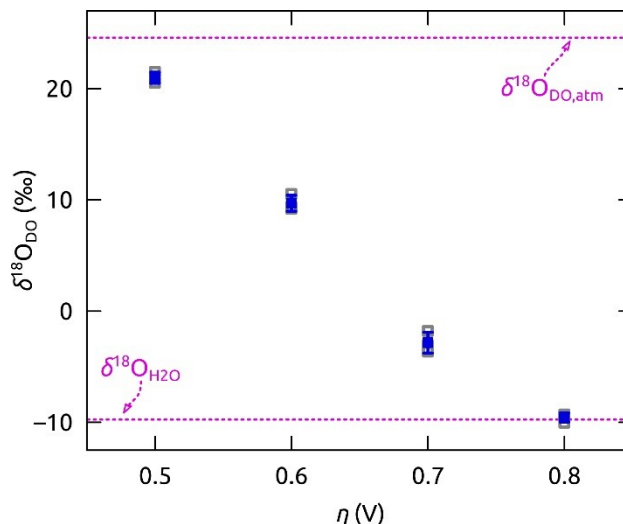


Figure 2. Potential-dependent $\delta^{18}\text{O}_{\text{DO}}$ composition of O_2 evolved at nanoporous Fe_2O_3 electrodes. The DO compositions are studied for O_2 evolving upon 3 h of steady-state electrolysis in a pH 7 aqueous KH_2PO_4 electrolyte at different applied overpotentials $0.48 \leq \eta \leq 0.78$ V (potentials E vs. Ag/AgCl : $+1.10 \leq E \leq +1.40$ V). $\delta^{18}\text{O}_{\text{DO}}$ values are presented for three nominally identical individual samples (grey empty squares) and the corresponding average values with standard deviations (blue squares and bars). Isotopic compositions measured for atmospheric O_2 dissolved in the aqueous electrolyte and electrolyte H_2O are given as control values (dashed lines): $\delta^{18}\text{O}_{\text{DO,atm}} = +24.6\text{‰}$, $\delta^{18}\text{O}_{\text{H}_2\text{O}} = -9.8\text{‰}$.

The experiment performed on iridium electrodes shows a qualitatively similar trend of $\delta^{18}\text{O}_{\text{DO}}$ decline with η increase as observed for the Fe_2O_3 electrode (**Supplementary Fig. 5**). At low overpotentials, the $\delta^{18}\text{O}_{\text{DO}}$ values are more positive than the substrate (water). Thus, both catalysts preferentially consume H_2^{18}O . At large applied η , both catalysts converge to a $\delta^{18}\text{O}_{\text{DO}}$ almost identical to that of the substrate H_2O consumed in the reaction. In other words, we observe that the discrimination of the catalyst for H_2^{18}O is reduced with increasing overpotential. This is typical of electrochemical reactions, where the applied overpotential not only affects the thermodynamic driving force of the reaction, but also its activation energy. In fact, this effect furnishes the basis of the Butler-Volmer equation, the fundament of all electrochemical kinetics. When two reactions have distinct activation energies at equilibrium ($\eta = 0$, **Figure 3** left), applying an overpotential ($\eta > 0$) reduces both of their activation energies, which simultaneously and necessarily also reduces the difference between their activation energies (**Figure 3** center and right). In other words, the selectivity of electrochemical reactions decreases with increasing overpotential without a change in the reaction mechanism (a quantitative treatment of this effect in the framework of Marcus' theory of electron transfer has been published).⁵⁸ Thus, for our case, the properties of electrochemical reactions cause the discrimination between ^{18}O and ^{16}O to decrease as η is increased: Accordingly, the approach towards the $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ -value observed at high overpotentials is plausible, because after a critical point the driving force has become so large that any activation energy is insignificant. In these conditions of large overpotentials, the electrochemical turnover at the surface is large and not selective for stable isotopes and diffusion becomes rate-limiting. Thus, the isotopic composition of H_2O is directly transferred to its product O_2 .

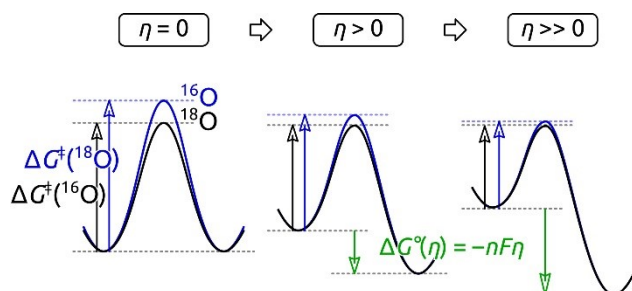


Figure 3. Schematic explanation for the decrease of selectivity upon increase in overpotential in electrochemical reactions. A change in mechanism is not required to explain a decrease in selectivity. The scheme includes the free enthalpy of activation ΔG^\ddagger , and the free enthalpy of reaction ΔG° , which depends on the number of electrons n exchanged in the rate-determining step, the Faraday constant F and the overpotential η .

Mechanistic interpretation

For further interpretation, we convert the abundance ratios R_{O_2} of O_2 produced from natural abundance water to ^{18}O KIEs, specifically, the ratio between the rate constants k for conversion of $H_2^{16}O$ and $H_2^{18}O$.^{59,60} The ^{18}O KIEs determined for Fe_2O_3 electrodes (blue data points) at overpotentials $0.48 \leq \eta \leq 0.78$ V (**Figure 4**) range between 0.970 and 1.000. When translated from the isotope value above, the unity KIE value at large η corresponds to the absence of isotopic discrimination when the reaction becomes diffusion-controlled. More importantly, the ^{18}O KIEs determined at $\eta \leq +0.68$ V are all KIE < 1.000 (the values lie between 0.970 and 0.993). This is usually referred to as an ‘inverse’ KIE.⁵⁹ Taking into account the fractionation effect of 0.7‰ mentioned above for the physical transfer of O_2 from the dissolved to gaseous phase causes a maximal variation of ^{18}O KIEs by 0.0007. The effect of adventitious DO (also mentioned above) represents 0.0018 in KIE units. It is possible that this could affect the absolute values somewhat, yet the trend remains unambiguous.

The completely different electrocatalyst iridium also yields a pronounced inverse KIE with ^{18}O KIEs ranging between 0.977 and 1.001 for overpotentials $0.20 \leq \eta \leq 0.50$ V (green data points). However, a significant difference is that the noble metal electrodes approach ^{18}O KIE unity much faster than their ‘rust’ counterparts. This behaviour is likely due to the higher catalytic activity of Ir. For this catalyst the activation energy is lower to start with, the case of negligible discrimination is reached much earlier. The shift of the curves is about 0.4 V, or 39 kJ mol⁻¹. This difference can be interpreted as the additional activation energy decrease that could be achieved by the Ir surface with respect to Fe_2O_3 .

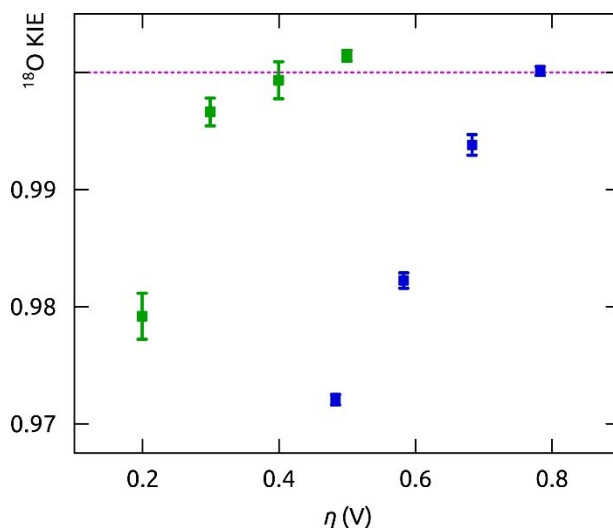


Figure 4. Potential-dependent ^{18}O KIEs for the OER at nanoporous Fe_2O_3 or Ir electrodes. ^{18}O KIEs of nanoporous Fe_2O_3 (blue data points) or Ir (green data points) are presented for different applied overpotentials η . Steady-state electrolyses were performed in a 0.1 M KH_2PO_4 solution (pH neutral) for 3 h or in a 0.1 M H_2SO_4 solution for 0.5 h at Fe_2O_3 and Ir electrodes, respectively. The ^{18}O KIEs presented are average values calculated from three nominally identical individual samples. Their corresponding standard deviations are shown as error bars. The unity KIE value is indicated by the dashed line.

For a mechanistic interpretation of the isotope effect, let us now focus on the Fe_2O_3 electrodes, for which ample literature material can be exploited.^{9,11,13,16,19} Additional information is provided by 2H isotope effects.⁴⁶ The use of deuterated water significantly decreases the OER activity (**Supplementary Fig. 6** and **Supplementary Table 2**). The ratio of exchange current densities J_0 for OER in H_2O and 2H_2O is determined from the quasi-steady-state current-

potential curve (Tafel slopes, see Methods section and **Supplementary Fig.6**). This ratio is equal to the ratio of rate constants k_0 and defines the ^2H KIE. Using the Tafel slope over the range $0.45\text{ V} \leq \eta \leq 0.55\text{ V}$, the calculated ^2H KIE is 0.56 (varying the η -range for the fit yields KIE values between 0.56 and 0.96, always inverse).

Discussion

What is the mechanistic origin of the inverse ^{18}O KIE determined during turnover conditions for OER? This question can be answered by the competitive ^{18}O KIEs, which yield information about the steps from the encounter of the substrate (H_2O) with the catalyst up to the first irreversible step.⁶¹ Regardless of the specific mechanism in OER for the catalysts under study, more than one step is expected to be isotopically sensitive. Interestingly, ^{18}O KIEs have been shown to report on O—O bond formation steps as well as rate-limiting steps that occur before O—O bond formation in water oxidation reactions catalysed by metal ions.^{25,27,62} Both the O—O bond formation and its preliminary steps have been hypothesised as rate-limiting steps for various systems.^{1,11,13,17,21,63}

In order to test these two hypotheses in our iron oxide system, we analyse the isotope effect data considering the simplest possible mechanism for O_2 evolution during the catalytic cycle (**Figure 5**). Binding of water molecules is known to be a reversible process and therefore we do not consider it as a potential rate-limiting step in our analysis.⁴⁴ The following discussion will be based on the closest possible systems for which isotope effect values are available in the literature.

First, we consider the hypothesis that the oxidation / deprotonation is the primary irreversible step. In this case, the overall ^{18}O KIE determined (KIE_{obs}) represents the product of the equilibrium isotope effect (EIE) for H_2O -binding ($\text{EIE}_{\text{coord}}$) and the intrinsic KIE for the oxidation / deprotonation step (KIE_{ox}) shown in equation 1:

$$\text{KIE}_{\text{obs}}^{\text{ferryl}} = \text{EIE}_{\text{coord}} \cdot \text{KIE}_{\text{ox}} \quad (1)$$

Given the preference of metal ions for H_2^{18}O over H_2^{16}O , the binding of water to the surface is expected to produce an inverse $\text{EIE}_{\text{coord}}$.^{64,65} This effect has been estimated by Taube to be in the range of 0.95 - 0.97.⁴⁴ Given our experimentally determined KIE_{obs} of 0.97 at low overpotential, the expected ^{18}O KIE for the oxidation / deprotonation process must be in the range of $1.00 \leq \text{KIE}_{\text{ox}} \leq 1.02$. This is inconsistent with KIE_{ox} -values determined for the oxidation of aquo Fe(II) ions, which vary from 1.04 ± 0.01 to 1.09 ± 0.02 as supported by a nonadiabatic multiphonon quantum mechanical approach.⁶⁶ Furthermore, rate-determining oxidation / deprotonation results in small normal ^2H KIE values (1.03 - 1.07),²⁵⁻²⁷ which contradicts the observations for the Fe_2O_3 system being studied. Thus, our experimental data allow us to exclude conclusively that oxidation / deprotonation is the rate-determining step.

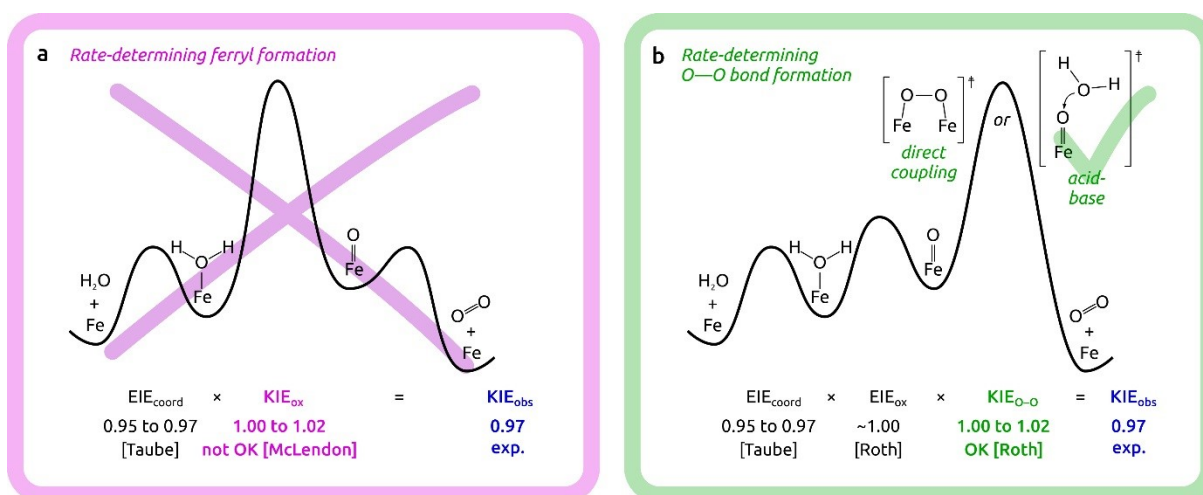


Figure 5. Qualitative energy diagrams for the mechanistic interpretation of the observed ^{18}O KIE. The mechanistic interpretation is provided in the framework of the minimal reaction mechanism sketched in **Figure 1**. It shows two distinct reaction steps as potentially rate-determining:^{1,11,13,17,21,63} either **a** the local oxidation and deprotonation to a reactive metal-oxo or **b** the O—O bond formation with two distinct possible mechanisms. For the first scenario, the ^{18}O KIE determined experimentally (KIE_{obs}) is not consistent with the range of values expected from literature data, whereas for the latter case KIE_{obs} is in agreement with the data available so far, indicating an acid-base mechanism rather than direct oxo coupling.^{25,27,44,66} The overall ^{18}O KIE determined (KIE_{obs}) represents the product of the equilibrium isotope effect (EIE) for H_2O -binding ($\text{EIE}_{\text{coord}}$) and the intrinsic KIE for the oxidation / deprotonation step (KIE_{ox}).

In the alternative scenario with the O—O bond formation as the rate-limiting step, the observed KIE results from equation 2,

$$\text{KIE}_{\text{obs}}^{\text{O}-\text{O}} = \text{EIE}_{\text{coord}} \cdot \text{EIE}_{\text{ox}} \cdot \text{KIE}_{\text{O}-\text{O}} \quad (2)$$

where EIE_{ox} represents the EIE for the oxidation / deprotonation step and $\text{KIE}_{\text{O}-\text{O}}$ is the KIE of the O—O bond formation step. To our knowledge, EIE_{ox} has not been directly determined for iron-containing systems, but it can be estimated to be ~ 1.00 based on computational calculations of three different ruthenium systems that catalyse the oxidation of water.²⁷ Thus, given $0.95 \leq \text{EIE}_{\text{coord}} \leq 0.97$ and $\text{KIE}_{\text{obs}} = 0.97$, the $\text{KIE}_{\text{O}-\text{O}}$ is in the 1.00 - 1.02 range, equating to a normal KIE. Normal ^{18}O KIEs have been determined to originate from O—O bond formation steps in four different homogeneous catalysts, including ferrate,²⁵ which agrees with this estimation. Further credence to this model is provided by the inverse ^2H KIE values (0.10 to 0.95) determined for these systems,²⁵ which coincide with our ^2H KIE values of 0.56 to 0.96. Additionally, experimental evidence by Hamann *et al.* identified a surface ferryl as an intermediate species in line with our interpretation.¹³

Van Voorhis *et al.* have recently predicted the O—O bond formation as the rate-limiting step for iron oxide, based on density functional theory calculations in combination with Sabatier analysis.¹⁷ In these calculations, both the acid-base and direct coupling of metal-oxo species seem to be equally plausible. In the four systems in which ^{18}O KIEs were used to study O—O bond formation experimentally, it has been reported that direct coupling of metal-oxo species leads to ^{18}O KIE values larger than 1.030.²⁷ In contrast, acid-base mechanisms range from 1.017 to 1.031.²⁷ Based on the range of possible ^{18}O $\text{KIE}_{\text{O}-\text{O}}$ for the system studied here (1.00 - 1.02) we suggest that the rate-limiting step in OER catalysed by Fe_2O_3 is the O—O bond formation via an acid-base mechanism rather than a direct coupling mechanism.

In conclusion, direct natural-abundance determinations by IRMS of product isotope compositions ($\delta^{18}\text{O}_{\text{DO}}$) are experimentally facile to implement. Macroscopic $\delta^{18}\text{O}_{\text{DO}}$ values can easily be converted to kinetic isotope effects. The overall kinetic isotope effects (KIE_{obs}) determined in this manner are sufficiently accurate to allow for microscopic, mechanistic interpretation. The application of this method to heterogeneous electrocatalytic water oxidation

systems yields clear dependencies on the applied overpotential and distinct profiles for two different electrode materials. The KIE_{obs} in the iron oxide case is quantitatively incompatible with the high-valent ferryl formation as the rate-determining step. Instead, it favours the hypothesis of a nucleophilic water attack on the ferryl as the mechanism of the rate-determining step over the direct coupling between two adjacent ferryl units. The methods demonstrated in this paper render systematic investigations of various heterogeneous or heterogenised electrocatalysts (and photoelectrocatalysts) possible. We envision such systematic studies will lead to the identification of distinct catalyst classes governed by distinct reaction mechanisms. Such information has not been available so far. The insight generated will lay the foundation for a more efficient search of improved or alternative water oxidation catalysts.

Methods

Materials

Chemicals were purchased from Sigma–Aldrich, Alfa Aesar, ABCR, or VWR and used as received. Water was purified in a Millipore Direct-Q system for application in electrolytes. Aluminium plates (99.99%) and Si (100) wafers covered with an oxide layer were supplied by Smart Membranes and Silicon Materials Inc., respectively. Ozone was generated with a BMT 803N ozone generator from oxygen purchased from Air Liquide.

Preparation of iron oxide electrodes

Nanoporous iron oxide electrodes were prepared according to previously published work on porous Fe_2O_3 electrodes.⁵² Anodic aluminium oxide (AAO) of approximately 11 μm length (obtained with 4 h of anodization duration during the second anodization step) and 370 nm diameter were used as porous templates. One side of the nanostructured substrate was closed with a several micrometer thick nickel layer, serving as the electrical backside contact in electrochemical measurements. The porous templates were then coated with Fe_2O_3 via atomic layer deposition in a commercial Gemstar-6 XT ALD reactor from Arradiance operating with a N_2 carrier gas. The deposition was performed at 200 °C with ferrocene ($FeCp_2$, kept in a stainless steel bottle maintained at 45 °C) and ozone as reactants. In the case of ferrocene, a 2-s pulse with 40 s of subsequent exposure time was repeated two times before the chamber was purged with N_2 for 90 s. For O_3 , single pulses were carried out with pulse, exposure and purge times of 0.5, 40, 90 s, respectively. 800 ALD cycles yielded a thickness of approximately 12 nm of Fe_2O_3 . A spectroscopic ellipsometer SENpro by Sentech was used to determine the Fe_2O_3 thickness on an Al_2O_3 coated silicon wafer, which was added to the reaction chamber. A total of 50 data points were recorded for wavelengths between 380 and 1050 nm under an angle of incidence of 70° with a tungsten halogen lamp as light source. The data orientation Θ and ellipticity ϵ were then fitted with a fixed optical model for ALD- Fe_2O_3 with SpectraRay/3.

Preparation of iridium electrodes

Nanoporous iridium electrodes were produced by the multistep preparation procedure by Schlicht et al.⁵³ The pore length and diameter of the AAO template used here were approximately 17 μm (obtained with 8 h of anodization duration) and 370 nm, respectively. The porous substrates, also equipped with an electrical nickel backside contact, were coated with a thin iridium layer in the same commercial Gemstar-6 XT ALD reactor. The deposition process was carried out at 220 °C with ethylcyclopentadienyl-1,3-cyclohexadiene-iridium (I) (($EtCp$)Ir(CHD) from Strem, kept at 90 °C in a stainless steel bottle) and ozone as precursors. Four consecutive microcycles consisting of a 0.5 s ($EtCp$)Ir(CHD) pulse and 40 s of exposure time were performed before the chamber was purged with N_2 for 90 s. For O_3 , single pulses were carried out with pulse, exposure and purge times of 0.5, 40, 90 s, respectively.

Approximately 20 nm of iridium were deposited with 240 ALD cycles. An indium-doped tin oxide (ITO) coated silicon wafer was used as a reference for the layer thickness determination via spectroscopic ellipsometry.

Electrochemical studies

The nanoporous electrodes were laser-cut into half with a GCC LaserPro Spirit LS Laser and glued with the nickel contact on smaller copper plates ($A \approx 4 \text{ cm}^2$) with double-sided conductive copper foil (**Supplementary Figure 7a**). A chemically resistant and electrically insulating polyimide (Kapton) adhesive tape, featuring a laser-cut oval ($a = 0.3 \text{ cm}$ and $b = 0.5 \text{ cm}$), was used to define the macroscopic sample area $A = 0.471 \text{ cm}^2$ exposed to the electrolyte also protecting the copper plate from the electrolyte. All electrochemical studies were performed at room temperature in a custom-made glovebox continuously flushed with N_2 using a Gamry Interface 1000 potentiostat. **Supplementary Figure 7b** shows nanoporous working electrodes (WE) adjusted into a double-walled cell (EuroCell from Gamry instruments) that was additionally equipped with a graphite rod as the counter-electrode (CE), a $\text{Ag}/\text{AgCl}/\text{KCl}(\text{sat.})$ reference electrode (RE, shifted by $+0.20 \text{ V}$ relative to the normal hydrogen electrode, NHE), and an optode (model Visiform DO 160 by Hamilton) for direct DO quantification with an accuracy $\pm 0.05 \text{ vol.}\%$. The cell was filled with 130 mL of fresh electrolyte, properly sealed and degassed with N_2 for 1 h via a septum prior to each measurement. A 0.1 M aqueous KH_2PO_4 solution (adjusted to pH 7) and a 0.1 M H_2SO_4 solution (pH 0.5) were used as electrolytes for nanoporous Fe_2O_3 or iridium electrodes, respectively.

Electrochemical studies of the ^2H kinetic isotope effect

The effect of protons on the OER was investigated at nanoporous Fe_2O_3 electrodes. Cyclic voltammograms (CV) were therefore measured at a scan rate of 5 mV s^{-1} in protio and deuterio ($^2\text{H}_2\text{O}$, 99.9 at.%) solutions containing 0.1 M KH_2PO_4 . Both phosphate electrolytes were adjusted to pH 7 with 5 M NaOH dissolved in H_2O or $^2\text{H}_2\text{O}$, respectively.

Preparation of water samples for KIE analysis

The oxygen isotope fractionation and the corresponding ^{18}O KIE of the water oxidation reaction was investigated for Fe_2O_3 and iridium as catalyst materials. Steady-state electrolyses were performed under stirring for various reaction conditions. In the case of water oxidation at nanoporous Fe_2O_3 electrodes, either the applied potential or the duration was varied between $+1.10 \leq E \leq +1.40 \text{ V}$ (overpotentials $0.48 \leq \eta \leq 0.78 \text{ V}$) and $0.50 \leq t \leq 5.25 \text{ h}$, respectively. For iridium electrodes, steady-state electrolyses were conducted for a fixed duration ($t = 0.5 \text{ h}$) in a potential range of $+1.20 \leq E \leq +1.50 \text{ V}$ (overpotentials $0.20 \leq \eta \leq 0.50 \text{ V}$). After termination of each experiment, aqueous electrolyte was directly extracted with a syringe from the electrochemical cell and transferred into helium-flushed vials ($V = 12.1 \text{ mL}$). For proper sealing, each vial was completely filled with electrolyte solution and immediately capped with a butyl rubber septum held in place with a plastic screw cap. These aqueous samples were then analysed within 24 h for their $\delta^{18}\text{O}_{\text{DO}}$. One concern when conducting this experiment was leakage of atmospheric O_2 into the sample vials that would have masked the oxygen produced by the OER. Any contributions of atmospheric oxygen due to improper sealing could be excluded via a series of blank samples that revealed negligible incoming O_2 . In addition, samples for stable oxygen isotope determinations of H_2O (expressed as $\delta^{18}\text{O}_{\text{H}_2\text{O}}$) in the non-converted (fresh) electrolyte were also studied as reference values.

Mass spectrometric analysis of oxygen isotopes

All values are reported in the standard delta notation in per mille (‰) versus the Vienna Standard Mean Ocean (VSMOW) according to equation 3,

$$\delta^{18}\text{O} = \left[\frac{R_{\text{sample}}}{R_{\text{standard}}} - 1 \right] \cdot 1000 \text{‰} \quad (3)$$

where R is the molar ratio of the of the heavy and light isotope of an element ($n(^{18}\text{O})/n(^{16}\text{O})$) in the sample and the standard (VSMOW), with the latter having a value of 2.0052×10^{-3} .^{67,68} Stable oxygen isotopes of DO (expressed as $\delta^{18}\text{O}_{\text{DO}}$) were measured with an automated continuous flow approach by Barth et al.⁴⁸ that was adapted to a Gasbench II auto sampler coupled to a Thermo Fischer Delta V Advantage IRMS. The isolation of DO into a headspace relies on a helium extraction technique by Wassenaar and Koehler.⁶⁹ In the adapted technique, a headspace of 3 to 6 mL was generated by removing the same aliquot of water via a double syringe system. Variable headspaces were necessary to adapt to various DO concentrations. The DO was then transferred to the headspace by shaking the vials for 30 min on an orbital shaker. Then the headspace was extracted into a sampling loop on the Gasbench II system. From there it was passed on to the IRMS in a helium flow via a capillary gas chromatographic column that separated O_2 from N_2 and accessory gases (e.g. Ar). The GC column was a Varian CP molecular sieve column of 25 m length with an inner diameter of 0.53 mm. It had an inner film of 0.5 μm thickness that acted as a 5 Å molecular sieve. The entry pressure of the helium carrier gas was 0.6 bar, yielding a flow of 5.6 mL min^{-1} . All operations were performed at 25 °C. This process yields sufficient O_2 gas from 12-mL water samples for reliable IRMS analyses. The additional tuning of the mass spectrometer to a m/z ratio of 32 enables favorable direct determination of molecular $\text{O}_2(\text{g})$ without combustion to CO_2 .^{49,50} The standard deviation of this method was determined with a value better than $\pm 0.2\text{‰}$ vs. VSMOW by multiple measurements of $\delta^{18}\text{O}_{\text{DO}}$ controls and standards.

The $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ compositions were analysed by an automated equilibration unit (Gasbench II) that was coupled in continuous flow mode to a Delta plus XP IRMS by Thermo Fisher Scientific. External reproducibility based on repeat analyses of laboratory internal controls was better than 0.1‰. All data for stable isotope measurements were corrected for a blank signal, linearity (i.e. detector-related shifts of isotope ratios in response to peak size) and instrumental drift during the run.⁷⁰

All $\delta^{18}\text{O}_{\text{DO}}$ values were further converted into abundance ratios R in order to define ^{18}O KIEs.^{59,60} If the substrate H_2O can be treated as an infinitely large reservoir with insignificantly small molar changes upon product (O_2) formation,⁶⁰ then equation 4 is valid:

$$\text{KIE}(^{18}\text{O}) = \frac{k_{\text{light}}}{k_{\text{heavy}}} = \frac{k_{^{16}\text{O}}}{k_{^{18}\text{O}}} = \frac{R_{\text{substrate}}}{R_{\text{product}}} = \frac{R_{\text{H}_2\text{O}}}{R_{\text{O}_2}} \quad (4)$$

with $R_{\text{H}_2\text{O}} = 1.986 \times 10^{-3}$ for the aqueous electrolyte.

Data availability

Any data generated and analysed for this study that are not included in this article and its Supplementary Information are available from the authors upon request.

References

1. Surendranath Y., Kanan M. W., Nocera D. G. Mechanistic studies of the oxygen evolution reaction by a cobalt-phosphate catalyst at neutral pH. *J. Am. Chem. Soc.* **132**. 16501-16509 (2010).

- 430 2. Cook T. R., Dogutan D. K., Reece S. Y., Surendranath Y., Teets T. S., Nocera D. G. Solar energy supply and storage for the legacy and nonlegacy worlds. *Chem. Rev.* **110**. 6474-6502 (2010).
3. Eisenberg R., Gray H. B. Preface on making oxygen. *Inorg. Chem.* **47**. 1697-1699 (2008).
4. Betley T. A., Wu Q., Van Voorhis T., Nocera D. G. Electronic design criteria for O–O bond
435 formation via metal–oxo complexes. *Inorg. Chem.* **47**. 1849-1861 (2008).
5. Rasiyah P., Tseung A. C. C. The role of the lower metal oxide/higher metal oxide couple in oxygen evolution reactions. *J. Electrochem. Soc.* **131**. 803-808 (1984).
6. Fujishima A., Honda K. Electrochemical photolysis of water at a semiconductor electrode. *Nature*. **238**. 37-38 (1972).
- 440 7. Osterloh F. E. Inorganic nanostructures for photoelectrochemical and photocatalytic water splitting. *Chem. Soc. Rev.* **42**. 2294-2320 (2013).
8. Burke M. S., Enman L. J., Batchellor A. S., Zou S., Boettcher S. W. Oxygen evolution reaction electrocatalysis on transition metal oxides and (oxy)hydroxides: Activity trends and design principles. *Chem. Mater.* **27**. 7549-7558 (2015).
- 445 9. Klahr B., Gimenez S., Fabregat-Santiago F., Bisquert J., Hamann T. W. Electrochemical and photoelectrochemical investigation of water oxidation with hematite electrodes. *Energy Environ. Sci.* **5**. 7626-7636 (2012).
10. Bediako D. K., Surendranath Y., Nocera D. G. Mechanistic studies of the oxygen evolution reaction mediated by a nickel–borate thin film electrocatalyst. *J. Am. Chem. Soc.* **135**. 3662-3674 (2013).
- 450 11. Le Formal F., Pastor E., Tilley S. D., Mesa C. A., Pendlebury S. R., Grätzel M., Durrant J. R. Rate law analysis of water oxidation on a hematite surface. *J. Am. Chem. Soc.* **137**. 6629-6637 (2015).
12. Zhang M., de Respinis M., Frei H. Time-resolved observations of water oxidation intermediates on a cobalt oxide nanoparticle catalyst. *Nat. Chem.* **6**. 362-367 (2014).
- 455 13. Zandi O., Hamann T. W. Determination of photoelectrochemical water oxidation intermediates on haematite electrode surfaces using operando infrared spectroscopy. *Nat. Chem.* **8**. 778-783 (2016).
14. Diaz-Morales O., Ferrus-Suspedra D., Koper M. T. M. The importance of nickel oxyhydroxide deprotonation on its activity towards electrochemical water oxidation. *Chem. Sci.* **7**. 2639-2645 (2016).
- 460 15. Pendlebury S. R., Cowan A. J., Barroso M., Sivula K., Ye J., Gratzel M., Klug D. R., Tang J., Durrant J. R. Correlating long-lived photogenerated hole populations with photocurrent densities in hematite water oxidation photoanodes. *Energy Environ. Sci.* **5**. 6304-6312 (2012).
- 465 16. Barroso M., Pendlebury S. R., Cowan A. J., Durrant J. R. Charge carrier trapping, recombination and transfer in hematite (α-Fe₂O₃) water splitting photoanodes. *Chem. Sci.* **4**. 2724-2734 (2013).
17. Mavros M. G., Shepherd J. J., Tsuchimochi T., McIsaac A. R., Van Voorhis T. Computational design principles of two-center first-row transition metal oxide oxygen evolution catalysts. *J. Phys. Chem. C*. **121**. 15665-15674 (2017).
- 470 18. Yatom N., Neufeld O., Caspary Toroker M. Toward settling the debate on the role of Fe₂O₃ surface states for water splitting. *J. Phys. Chem. C*. **119**. 24789-24795 (2015).
19. Peter L. M., Wijayantha K. G. U., Tahir A. A. Kinetics of light-driven oxygen evolution at α-Fe₂O₃ electrodes. *Faraday Discuss.* **155**. 309-322 (2012).
20. Cummings C. Y., Marken F., Peter L. M., Upul Wijayantha K. G., Tahir A. A. New insights into water splitting at mesoporous α-Fe₂O₃ films: A study by modulated transmittance and impedance spectroscopies. *J. Am. Chem. Soc.* **134**. 1228-1234 (2012).
- 475 21. Zhang Y., Zhang H., Liu A., Chen C., Song W., Zhao J. Rate-limiting O–O bond formation pathways for water oxidation on hematite photoanode. *J. Am. Chem. Soc.* **140**. 3264-3269 (2018).
- 480 22. Guy R. D., Fogel M. L., Berry J. A. Photosynthetic fractionation of the stable isotopes of oxygen and carbon. *Plant Physiol.* **101**. 37-47 (1993).
23. Burda K., Bader K. P., Schmid G. H. ¹⁸O isotope effect in the photosynthetic water splitting process. *Biochim. Biophys. Acta, Bioenerg.* **1557**. 77-82 (2003).
24. Eisenstadt D., Barkan E., Luz B., Kaplan A. Enrichment of oxygen heavy isotopes during photosynthesis in phytoplankton. *Photosynth. Res.* **103**. 97-103 (2010).
- 485 25. Sarma R., Angeles-Boza A. M., Brinkley D. W., Roth J. P. Studies of the di-iron(VI) intermediate in ferrate-dependent oxygen evolution from water. *J. Am. Chem. Soc.* **134**. 15371-15386 (2012).
26. Angeles-Boza A. M., Roth J. P. Oxygen kinetic isotope effects upon catalytic water oxidation by a monomeric ruthenium complex. *Inorg. Chem.* **51**. 4722-4729 (2012).
- 490 27. Angeles-Boza A. M., Ertem M. Z., Sarma R., Ibanez C. H., Maji S., Llobet A., Cramer C. J., Roth J. P. Competitive oxygen-18 kinetic isotope effects expose O–O bond formation in water

- oxidation catalysis by monomeric and dimeric ruthenium complexes. *Chem. Sci.* **5**. 1141-1152 (2014).
28. Song N., Concepcion J. J., Binstead R. A., Rudd J. A., Vannucci A. K., Dares C. J., Coggins M. K., Meyer T. J. Base-enhanced catalytic water oxidation by a carboxylate–bipyridine Ru(II) complex. *Proc. Natl. Acad. Sci.* **112**. 4935-4940 (2015).
 29. Shaffer D. W., Xie Y., Szalda D. J., Concepcion J. J. Manipulating the rate-limiting step in water oxidation catalysis by ruthenium bipyridine–dicarboxylate complexes. *Inorg. Chem.* **55**. 12024-12035 (2016).
 30. Wiberg K. B. The deuterium isotope effect. *Chem. Rev.* **55**. 713-743 (1955).
 31. Westheimer F. H. The magnitude of the primary kinetic isotope effect for compounds of hydrogen and deuterium. *Chem. Rev.* **61**. 265-273 (1961).
 32. Bigeleisen J., Wolfsberg M. *Theoretical and Experimental Aspects of Isotope Effects in Chemical Kinetics*. In: *Advances in Chemical Physics*, vol. 1. (John Wiley & Sons. Inc., New York, 2007).
 33. Zhu H., Peck S. C., Bonnot F., van der Donk W. A., Klinman J. P. Oxygen-18 kinetic isotope effects of nonheme iron enzymes HEPD and MPnS support iron(III) superoxide as the hydrogen abstraction species. *J. Am. Chem. Soc.* **137**. 10448-10451 (2015).
 34. Stevens C. L. R., Schultz D., Van Baalen C., Parker P. L. Oxygen isotope fractionation during photosynthesis in a blue-green and a green alga. *Plant Physiol.* **56**. 126-129 (1975).
 35. Mader M., Schwerna P., Buchholz R., van Geldern R., Barth J. A. C. A new approach to quantify system efficiency with dissolved oxygen isotopes during engineered growth of *Galdieria sulphuraria*. *Algal Research*. **26**. 294-301 (2017).
 36. Mader M., Schmidt C., van Geldern R., Barth J. A. C. Dissolved oxygen in water and its stable isotope effects: A review. *Chem. Geol.* **473**. 10-21 (2017).
 37. Gersten S. W., Samuels G. J., Meyer T. J. Catalytic oxidation of water by an oxo-bridged ruthenium dimer. *J. Am. Chem. Soc.* **104**. 4029-4030 (1982).
 38. Gilbert J. A., Eggleston D. S., Murphy W. R., Geselowitz D. A., Gersten S. W., Hodgson D. J., Meyer T. J. Structure and redox properties of the water-oxidation catalyst $[(bpy)_2(OH_2)RuORu(OH_2)(bpy)_2]^{4+}$. *J. Am. Chem. Soc.* **107**. 3855-3864 (1985).
 39. Geselowitz D., Meyer T. J. Water oxidation by $[(bpy)_2(O)Ru^VORu^V(O)(bpy)_2]^{4+}$. An oxygen-labeling study. *Inorg. Chem.* **29**. 3894-3896 (1990).
 40. Concepcion J. J., Zhong D. K., Szalda D. J., Muckerman J. T., Fujita E. Mechanism of water oxidation by $[Ru(bda)(L)_2]$: the return of the "blue dimer". *Chem. Commun.* **51**. 4105-4108 (2015).
 41. Washburn E. W., Smith E. R., Smith F. A. Fractionation of the isotopes of hydrogen and of oxygen in a commercial electrolyzer. *J. Res. Natl. Bur. Stand.* **13**. 599-608 (1934).
 42. Smith E. R., Wojciechowski M. Fractionation of the isotopes of oxygen in a commercial electrolyzer. A correction. *J. Res. Natl. Bur. Stand.* **15**. 187-188 (1935).
 43. Hall W. H., Johnston H. L. Concentration of the heavier isotopes of oxygen in commercial electrolytic cells. *J. Am. Chem. Soc.* **57**. 1515-1517 (1935).
 44. Anbar M., Taube H. Oxygen isotope effects at anodes. *J. Am. Chem. Soc.* **78**. 3252-3255 (1956).
 45. Gulens J., Olmstead W. J., Longhurst T. H., Gale K. L., Rolston J. H. Electrolytic separation factors for oxygen isotopes in light and heavy water solutions. *J. Electrochem. Soc.* **134**. 1467-1470 (1987).
 46. Tse E. C. M., Hoang T. T. H., Varnell J. A., Gewirth A. A. Observation of an inverse kinetic isotope effect in oxygen evolution electrochemistry. *ACS Catalysis*. **6**. 5706-5714 (2016).
 47. Ruben S., Randall M., Kamen M., Hyde J. L. Heavy oxygen (O^{18}) as a tracer in the study of photosynthesis. *J. Am. Chem. Soc.* **63**. 877-879 (1941).
 48. Barth J. A. C., Tait A., Bolshaw M. Automated analyses of $^{18}O/^{16}O$ ratios in dissolved oxygen from 12-mL water samples. *Limnol. Oceanogr. Methods*. **2**. 35-41 (2004).
 49. Wang X., Veizer J. Erratum to Xuefeng Wang and Jan Veizer (2000), respiration-photosynthesis balance of terrestrial aquatic ecosystems, Ottawa area, Canada. *Geochimica et Cosmochimica Acta* **64**(22), 3775–3786. *Geochim. Cosmochim. Acta*. **68**. 933-934 (2004).
 50. Wang X., Veizer J. Respiration–photosynthesis balance of terrestrial aquatic ecosystems, Ottawa area, Canada. *Geochim. Cosmochim. Acta*. **64**. 3775-3786 (2000).
 51. Haschke S., Wu Y., Bashouti M., Christiansen S., Bachmann J. Engineering nanoporous iron(III) oxide into an effective water oxidation electrode. *ChemCatChem*. **7**. 2455-2459 (2015).
 52. Haschke S., Pankin D., Petrov Y., Bochmann S., Manshina A., Bachmann J. Design rules for oxygen evolution catalysis at porous iron oxide electrodes: A 1000-fold current density increase. *ChemSusChem*. **10**. 3644-3651 (2017).

53. Schlicht S., Haschke S., Mikhailovskii V., Manshina A., Bachmann J. Highly reversible water oxidation at ordered nanoporous iridium electrodes based on an original atomic layer deposition. *ChemElectroChem*. **5**. 1259-1264 (2018).
54. Benson B. B., Krause Jr. D. The concentration and isotopic fractionation of gases dissolved in freshwater in equilibrium with the atmosphere. 1. Oxygen. *Limnol. Oceanogr.* **25**. 662-671 (1980).
55. Coplen T. B., Hopple J. A., Bohlke J. K., Peiser H. S., Rieder S. E., Krouse H. R., Rosman K. J. R., Ding T., Vocke R. D., Jr., Revesz K. M., Lamberty A., Taylor P., De Bièvre P. *Compilation of minimum and maximum isotope ratios of selected elements in naturally occurring terrestrial materials and reagents*. (U.S. Department of the Interior, U.S. Geological Survey, Reston, Water-Resources Investigations Report 01-4222; 2002).
56. van Geldern R., Baier A., Subert H. L., Kowol S., Balk L., Barth J. A. C. Pleistocene paleo-groundwater as a pristine fresh water resource in southern Germany – evidence from stable and radiogenic isotopes. *Sci. Total Environ.* **496**. 107-115 (2014).
57. Benson B. B., Krause Jr. D., Peterson M. A. The solubility and isotopic fractionation of gases in dilute aqueous solution. I. Oxygen. *J. Solution Chem.* **8**. 655-690 (1979).
58. Kavner A., Bonet F., Shahar A., Simon J., Young E. The isotopic effects of electron transfer: An explanation for Fe isotope fractionation in nature. *Geochim. Cosmochim. Acta*. **69**. 2971-2979 (2005).
59. Cleland W. W. The use of isotope effects in the detailed analysis of catalytic mechanisms of enzymes. *Bioorg. Chem.* **15**. 283-302 (1987).
60. O'Leary M. H. Carbon isotope fractionation in plants. *Phytochemistry*. **20**. 553-567 (1981).
61. O'Leary M. H., Richards D. T., Hendrickson D. W. Carbon isotope effects on the enzymic decarboxylation of glutamic acid. *J. Am. Chem. Soc.* **92**. 4435-4440 (1970).
62. Khan S., Yang K. R., Ertem M. Z., Batista V. S., Brudvig G. W. Mechanism of manganese-catalyzed oxygen evolution from experimental and theoretical analyses of ^{18}O kinetic isotope effects. *ACS Catalysis*. **5**. 7104-7113 (2015).
63. Cowan A. J. Intermediate identification. *Nat. Chem.* **8**. 740-741 (2016).
64. Hunt J. P., Taube H. The exchange of water between hydrated cations and solvent. *J. Chem. Phys.* **19**. 602-609 (1951).
65. Feder H. M., Taube H. Ionic hydration: An isotopic fractionation technique. *J. Chem. Phys.* **20**. 1335-1336 (1952).
66. Guarr T., Buhks E., McLendon G. Quantum effects of high-frequency modes in inorganic electron transfer: kinetic isotope effects in redox reactions of $[\text{Fe}(\text{H}_2\text{O})_6]^{2+}$, $[\text{Fe}(\text{D}_2\text{O})_6]^{2+}$, and $[\text{Fe}(^{18}\text{OH}_2)_6]^{2+}$. *J. Am. Chem. Soc.* **105**. 3763-3767 (1983).
67. Baertschi P. Absolute ^{18}O content of standard mean ocean water. *Earth. Planet. Sci. Lett.* **31**. 341-344 (1976).
68. Kendall C., McDonnell J. J. *Fundamentals of Isotope Geochemistry*. In: *Isotope Tracers in Catchment Hydrology*. (Elsevier Science B.V., Amsterdam, 1998).
69. Wassenaar L. I., Koehler G. An on-line technique for the determination of the $\delta^{18}\text{O}$ and $\delta^{17}\text{O}$ of gaseous and dissolved oxygen. *Anal. Chem.* **71**. 4965-4968 (1999).
70. Brand Willi A., Coplen Tyler B., Vogl J., Rosner M., Prohaska T. Assessment of international reference materials for isotope-ratio analysis (IUPAC Technical Report). *Pure Appl. Chem.* **86**. 425-467 (2014).

Acknowledgements

We thank the Friedrich-Alexander-Universität Erlangen-Nürnberg for travel support in the Visiting Professor program. A.M. A.-B. is grateful for support from the National Science Foundation CAREER grant (CHE-1652606). The work was also supported by the DFG grant IsoDO (BA 2207/15-1).

Author contributions

The data were interpreted and the manuscript written with contributions of all authors. S. H. and S. S. prepared the samples and performed the O_2 sampling upon electrolysis. M. M. analysed the dissolved oxygen by IRMS and determined the isotope

compositions. A.M. R. contributed to discussions and editing. A.M. A.-B., J.A.C. B., and J. B. devised the project.

Additional information

610 **Supplementary Information** accompanies this paper at

Competing interests: The authors declare no competing interests.