

¹ A Probabilistic Assessment of the Next Geomagnetic Reversal

Bruce Buffett,¹ William Davis¹

Corresponding author: Bruce Buffett, Department of Earth and Planetary Science, University of California, Berkeley, CA, USA. (bbuffett@berkeley.edu)

¹Department of Earth and Planetary Science, University of California, Berkeley, CA, 94720-4767 USA.

3 Deterministic forecasts for the next geomagnetic reversal are not feasible
4 due to large uncertainties in the present-day state of the Earth's core. A more
5 practical approach relies on probabilistic assessments using paleomagnetic
6 observations to characterize the amplitude of fluctuations in the geomagnetic
7 dipole. We use paleomagnetic observations for the past two million years to
8 construct a stochastic model for the axial dipole field and apply well-established
9 methods to evaluate the probability of the next geomagnetic reversal as a
10 function of time. For a present-day axial dipole moment of 7.6×10^{22} A
11 m^2 , the probability of the dipole entering a reversed state is less than 2% af-
12 ter 20 kyr. This probability rises to 11% after 50 kyr. An imminent geomag-
13 netic reversal is not supported by paleomagnetic observations. The current
14 rate of decline in the dipole moment is unusual, but within the natural vari-
15 ability predicted by the stochastic model.

1. Introduction

16 The strength of the geomagnetic dipole has been declining since the first direct mea-
17 surements were made in 1832 [Malin, 1987]. The rate of decline greatly exceeds the rate of
18 decay by diffusion [Olson, 2002], prompting speculations that the Earth may be entering
19 the early stage of a geomagnetic reversal [Hulot *et al.*, 2002]. Much of the present-day
20 decline is attributed to transport of magnetic flux from high latitudes toward the equator
21 [Finlay *et al.*, 2016]. However, a rearrangement of magnetic flux exclusively at the core-
22 mantle boundary is probably insufficient to produce a reversal because the axial dipole
23 outside the core reflects the volume-averaged axial magnetic field in the core [Davidson,
24 2001, p. 173]. On the other hand, a synchronous reorganization of the internal magnetic
25 field cannot be ruled out [Metman *et al.*, 2018]. We currently lack sufficient information
26 about the internal configuration of the core to make reliable forecasts using numerical geo-
27 dynamo models [Kuang *et al.*, 2009; Aubert and Fournier, 2011; Aubert, 2014]. **Forecasts**
28 **are also affected by fundamental limitations in the predictability of complex dynamical**
29 **systems [e.g. Hulot *et al.*, 2010].**

30 Paleomagnetic observations offer an alternative way to address the question of the next
31 geomagnetic reversal because records of past behavior can serve as a basis for making infer-
32 ences about the future. In particular, Constable and Korte [2006] found several intervals
33 in the last 7 kyr with rates of dipole decline comparable to the present-day observations.
34 None of these intervals of rapid decline led to a reversal, suggesting that the present trend
35 is simply part of the natural variability of the geomagnetic field. A different approach
36 was taken by Morzfeld *et al.* [2017], who assimilated paleomagnetic data into several low-

37 degree dynamical models. They demonstrated skill in forecasting reversals 4 kyr into the
38 future and concluded that a reversal in the next few kyr was unlikely. Extending these
39 predictions in time was limited by the use of dynamical models in which one or more of
40 the state variables were not directly constrained by observations. Uncertainties in the
41 unobserved states were found to adversely affect the reliability of the predictions. In this
42 study we extend the time frame for predictions by using several compilations of paleo-
43 magnetic observations to construct a stochastic model for fluctuations in the dipole field.
44 Once the stochastic model is established, well-known methods offer a rigorous assessment
45 of the reversal probability for arbitrary times into the future. It is important to point out
46 that these predictions are inherently statistical in that we cannot forecast the outcome
47 of a single realization. Instead, we predict the probability of ending up with a magnetic
48 field in the opposite polarity after a specified period time.

49 Our predictions are contingent on the choice of stochastic model. This dependence
50 might be perceived as a drawback, but an advantage of the approach is that the basic
51 form of the stochastic model is largely determined by paleomagnetic observations [e.g. *Buf-*
52 *fett et al.*, 2013; *Meduri and Wicht*, 2016]. More importantly, it is possible to incorporate
53 different types of observations into these models to improve the statistical description of
54 dipole fluctuations, including properties like the mean rate of geomagnetic reversals [*Gee*
55 *and Kent*, 2015]. This means that stochastic models offer an extendable strategy for com-
56 bining diverse observations into a single quantitative framework for making probabilistic
57 assessments of the field behavior.

58 Our assessment of the reversal probability should be interpreted as the average or ex-
 59 pected number of realizations that enter the reversed state after a specified time, although
 60 we do not explicitly restore to realizations to assess this probability. Instead, the prob-
 61 lem is formulated in terms of the adjoint to the Fokker-Planck equation, which is some-
 62 times called the backward Fokker-Planck equation. Numerical solutions of the backward
 63 Fokker-Planck equation yield simple and direct estimates for the probability of reversal.
 64 We proceed with a brief description of the stochastic model used in this study and outline
 65 the formulation of the backward Fokker-Planck equation. Quantitative estimates for the
 66 probability of reversal are given for the next few tens of thousands of years.

2. Stochastic Model for the Dipole Moment

Time variations in the axial dipole moment, $x(t)$, are governed by a stochastic differen-
 tial equation [e.g. Van Kampen, 2007]

$$\frac{dx}{dt} = v(x) + \sqrt{D(x)} \Gamma(t) \quad (1)$$

where the drift term, $v(x)$, describes the deterministic evolution of the dipole moment and the noise term, $D(x)$, defines the amplitude of random variations. The time dependence of the random process, $\Gamma(t)$, is represented by uncorrelated Gaussian noise. The time average of $\Gamma(t)$ vanishes

$$\langle \Gamma(t) \rangle = 0 \quad (2)$$

and the autocorrelation function is defined by a delta function

$$\langle \Gamma(t_1) \Gamma(t_2) \rangle = 2\delta(t_1 - t_2), \quad (3)$$

67 where the factor of two is a common convention. The idealization of the autocorrelation in
 68 terms of a delta function is intended to approximate a noise source with a short correlation
 69 time relative to the sampling of $x(t)$. Figure 1 shows estimates for the drift and noise
 70 terms recovered by *Buffett and Puranum* [2017] from stacks of relative paleointensity
 71 measurements from the past two million years [Ziegler *et al.*, 2011], supplemented with
 72 higher resolution measurements from the past 10 kyr [Constable *et al.*, 2016]. **The noise**
 73 **term is approximated as a constant, while the drift term is defined by the extended model**
 74 **in equation (26) of *Buffett and Puranum* [2017]**

The random component of the stochastic model in (1) introduces uncertainty into $x(t)$ as the solution evolves forward in time. The probability of $x(t)$ is commonly specified in terms of a transitional probability, $p(x, t|x', t')$, which describes the probability of the dipole moment evolving from x' at some initial time t' to x at some later time $t > t'$. When the initial state is also uncertain (denoted by $p(x', t')$), the probability of the subsequent state is

$$p(x, t) = \int p(x, t|x', t')p(x', t') dx', \quad (4)$$

where we follow the usual custom of using $p(x, t)$ to define the probability density of x at time t . When the initial state is known (e.g. $p(x', t') = \delta(x' - x_0)$), the probability of $x(t)$ is given by

$$p(x, t) = \int p(x, t|x', t')\delta(x' - x_0)dx' = p(x, t|x_0, t'). \quad (5)$$

75 In other words, the probability of the later state is equivalent to the transitional probability
 76 for a known initial state.

Both the transitional probability $p(x, t|x', t')$ and the more general expression for $p(x, t)$ from (4) evolve with time according to the Fokker-Planck equation [Risken, 1989]

$$\frac{\partial p}{\partial t} = \mathcal{L} p, \quad (6)$$

where

$$\mathcal{L} = \left(-\frac{\partial}{\partial x} v(x) + \frac{\partial^2}{\partial x^2} D(x) \right). \quad (7)$$

77 Here the drift and noise terms are assumed to be independent of time, although this
 78 restriction is not essential. In many situations the Fokker-Planck equation provides a
 79 powerful tool for evaluating the probability of a stochastic process. We adopt this general
 80 approach here, although similar information can also be inferred from a large number
 81 of realizations of the stochastic model. We might use these realizations to evaluate the
 82 mean or variance at some later time. Both of these approaches are commonly used to
 83 characterize the statistics of the later state, given a description of the initial state.

84 The problem of determining the likelihood of a future polarity reversal is different
 85 because we know the end state (e.g. the dipole moment has reversed) and we seek to
 86 determine the probability of that event given a particular starting condition. For present
 87 purposes we denote the reversed state by $x(t) < 0$ and refer to this subset of states as
 88 the target set. The machinery needed to estimate probability of entering the target set
 89 is based on the adjoint equation to the Fokker-Planck equation, which is often called the
 90 backward Fokker-Planck equation [Gardiner, 2002].

3. Application of the Adjoint Equation

The adjoint equation is used to quantify the probability of $x(t)$ being in the target set (e.g. $x < 0$), subject to the condition that the dipole moment starts in a prescribed initial state x' at time t' . We define a function, $f(x)$, to specify whether $x(t)$ is in the target set. We let $f(x) = 1$ when $x < 0$ and set $f(x) = 0$ otherwise. The probability that $x(t)$ ends in the target state at time t is given by

$$u(x', t') = \int f(x)p(x, t|x', t') dx, \quad (8)$$

which is equivalent to the expected value of $f(x)$ once $x(t)$ is evolved from the initial state x' at t' . Any given realization for $x(t)$ will give either 0 or 1 for the value of $f(x)$, but the average (or expected) value defines the probability. The future time t of the reversed state is specified, so the probability u depends solely on x' and t' . An equivalent estimate of the expected value could also be obtained from realizations of the stochastic model: each realization is started at the initial condition $x(t') = x'$ and the mean value of $f(x)$ would be computed from the ensemble of solutions at t . When $t' \rightarrow t$ and $p(x, t|x', t') \rightarrow \delta(x - x')$ in (8), we find that

$$u(x', t) = f(x') \quad (9)$$

⁹¹ which serves as an *initial* condition when the evolution of $u(x', t')$ is integrated backward
⁹² in time from $t' = t$, the final state of the system.

The time derivative of $u(x', t')$ for backward integration is defined by

$$-\frac{\partial u(x', t')}{\partial t'} = - \int f(x) \frac{\partial}{\partial t'} p(x, t, |x', t') dx \quad (10)$$

where the change in the transition probability is taken with respect to the initial time t' rather than the final time t (compare with the Fokker-Planck equation in (6)). The required derivative of the transition probability is defined by the backward Fokker-Planck equation [Gardiner, 2002]

$$-\frac{\partial p(x, t|x', t')}{\partial t'} = \mathcal{L}^\dagger p(x, t|x', t') \quad (11)$$

where

$$\mathcal{L}^\dagger = \left(v(x') \frac{\partial}{\partial x'} + D(x') \frac{\partial^2}{\partial x'^2} \right) \quad (12)$$

is the adjoint operator of \mathcal{L} in (7), subject to the condition $p \rightarrow 0$ as $x \rightarrow \pm\infty$. Substituting (11) into (10) and interchanging the order of differentiation and integration gives

$$-\frac{\partial u(x', t')}{\partial t'} = \mathcal{L}^\dagger \int f(x) p(x, t|x', t') dx = \mathcal{L}^\dagger u(x', t') . \quad (13)$$

We use (13) to integrate $u(x', t')$ backward in time using (9) as an initial condition.

Reflecting boundary conditions

$$D(x') \frac{\partial u}{\partial x'} = 0 \quad (14)$$

are adopted at the upper and lower limits of the domain in x' . If we view the calculation as the outcome of an infinite number of realizations then the boundary condition are equivalent to ensuring that no realizations leave the domain. We also adopt a sufficiently broad domain ($|x| < 12 \times 10^{22} \text{ Am}^2$) so that the probability of reaching the upper or lower limits is effectively zero. Additional calculations confirm that an increase the size of the domain does not change the result. In summary, numerical solutions of (13) are used to predict the probability of a reversed field at t , given the magnitude and sign of the dipole field at t' .

4. Probability of the Next Polarity Reversal

Figure 2 shows solutions for $u(x', t')$ at two time differences $\tau = t - t' = 10$ kyr and 20 kyr. When the initial state, x' , is already inside the target set, it is quite likely that the dipole remains in the reversed state at 10 kyr and 20 kyr. There is a much smaller probability for a field with a positive initial polarity to reverse after 10 or 20 kyr. The details depend on the amplitude of the dipole at t' , and there is a general trend for $u(x', t')$ to decrease at larger values for the initial field. For the present-day axial dipole moment, $x' = 7.6 \times 10^{22}$ A m² [Gillet *et al.*, 2013], a reversal is expected to occur after 20 kyr with a probability of 0.019. Longer integrations of $u(x', t')$ for larger time differences τ eventually yield a constant probability of 0.5 for all values of x' . This is expected as the probability of finding the field in the reversed state at some distant time in the future is 50%, independent of the initial amplitude of the dipole moment, subject to the assumption that the field spends equal time in both polarities. (This assumption is implicit in the construction of the stochastic model because the drift term is required to be an odd function of x .)

Shorter time integrations offer more meaningful insights. Figure 3 shows a prediction for the probability of reversal as a function of the time difference τ when $x' = 7.6 \times 10^{22}$ A m². Initially, the chance of reversal is very low because the noise term is unable to drive the dipole moment into the opposite polarity over a few thousand years. The root-mean-square change in the dipole moment due solely to noise is

$$\langle \Delta x^2 \rangle^{1/2} = \sqrt{2D\Delta t}. \quad (15)$$

115 The expected change (or standard deviation) after 20 kyr is roughly $3.7 \times 10^{22} \text{ A m}^2$ using
 116 $D = 340 \times 10^{44} \text{ A}^2 \text{ m}^4 \text{ Myr}^{-1}$ [Buffett and Puranum, 2017]. This change in $x(t)$ is less than
 117 half the distance to the origin $x = 0$, so the noise term is not expected to drive the solution
 118 to the point of reversal after 20 kyr. Larger changes are possible, but these are less likely.
 119 For example, the change in dipole moment required to reach the onset of reversal ($x = 0$)
 120 is slightly in excess of two standard deviations. Consequently, we require an unusual
 121 sequence of events to move the present-day dipole moment to zero. A steady increase in
 122 probability $u(x', t')$ occurs once τ exceeds 15 kyr, although the predicted probability only
 123 reaches 11% after 50 kyr. Thus the prospects of an imminent geomagnetic reversal is low
 124 given the paleomagnetic record from PADM2M and CALS10k.2.

5. Discussion

125 The stochastic model offers a useful perspective for interpreting the recent historical
 126 decline in the dipole field. Over a 150-year interval between 1860 and 2010 the dipole
 127 moment has decreased by $\Delta x = 0.68 \times 10^{22} \text{ A m}^2$ [Gillet et al., 2013], compared with the
 128 standard deviation $\langle \Delta x^2 \rangle^{1/2} = 0.32 \times 10^{22} \text{ A m}^2$ from (15). Consequently, the observed
 129 change exceeds the standard deviation by a factor of 2.13, corresponding to a probability
 130 of 1.7% for Gaussian distributed noise. While the recent historical decline is unusual, it
 131 is within the allowable variation of the model. Extending the current linear trend over
 132 the next 1700 years would bring the dipole moment to zero (i.e. $\Delta x = 7.6 \times 10^{22} \text{ A m}^2$),
 133 whereas the standard deviation for this time interval is only $\langle \Delta x^2 \rangle^{1/2} = 1.08 \times 10^{22}$
 134 A m^2 . Changing $x(t)$ by seven times the standard deviation is highly unlikely, and this
 135 assessment is overly optimistic. A full description of the dipole evolution includes the drift

¹³⁶ term, which is positive once $x(t)$ drops below the time-averaged dipole ($\langle x \rangle = 5.3 \times 10^{22}$
¹³⁷ A m²). A positive drift opposes the steady decline and makes the prospect of a reversal in
¹³⁸ the next 1700 years even more unlikely. By comparison, the drift would contribute to the
¹³⁹ recent historical trend because we expect a negative drift with the current dipole moment
¹⁴⁰ (see Fig. 1).

Another point of reference for appraising our prediction is obtained by assuming that reversals obey a Poisson process [Cox, 1969]. Such a process lacks memory, so the cumulative probability of a reversal within a specified time interval, Δt , is independent of the prior history. Instead, the probability is specified by the mean reversal rate, which is approximately $r = 4.4 \text{ Myr}^{-1}$ over the past 5 Myr [Ogg, 2012]. The cumulative probability of a reversal occurring within the time interval Δt is given by

$$P(t < \Delta t) = 1 - e^{-r\Delta t} \quad (16)$$

¹⁴¹ which yields $P = 0.084$ for $\Delta t = 20$ kyr. By comparison, the prediction from our stochastic
¹⁴² model ($P = 0.019$) is much lower because the noise term limits the rate at which the dipole
¹⁴³ approaches a reversed state. Extending the time interval to $\Delta t = 50$ kyr, increases the
¹⁴⁴ reversal probability for the Poisson process to 0.197. A similar increase in probability
¹⁴⁵ ($P = 0.11$) is predicted using the stochastic model. These two estimates have been
¹⁴⁶ brought into slightly closer agreement, which can be interpreted as greater agreement in
¹⁴⁷ the underlying assumptions. A larger value for Δt lowers the dependence of the stochastic
¹⁴⁸ model on initial conditions, which makes the statistical description more like a process
¹⁴⁹ with no memory. However, our predictions over the next 10^4 years are still strongly

₁₅₀ dependent on the initial state. On the other hand, there is no dependence on the current
₁₅₁ rate of change in the dipole.

₁₅₂ Strictly speaking, the cumulative probability for a Poisson process differs conceptually
₁₅₃ from the prediction of the backwards Fokker-Planck equation. In former case we predict
₁₅₄ the probability of a chron duration and in the latter we predict the probability of ending
₁₅₅ up in the reversed state. For very large Δt the Poisson process predicts $P \rightarrow 1$, whereas
₁₅₆ the probability of being in the reversed state is $P \rightarrow 0.5$. In principle, a realization can
₁₅₇ enter the reversed state and subsequently revert back to the normal state by the end of
₁₅₈ the prescribed time interval. Such a path would be recorded as a realization in the normal
₁₅₉ state (e.g. not in the target set). However, the probability of this event is greatly reduced
₁₆₀ at short Δt because the actual reversal would need to occur within a shorter time frame.
₁₆₁ As long as the probability of this earlier event is low, it should not affect our interpretation
₁₆₂ reversal probabilities.

6. Concluding Remarks

₁₆₃ Quantitative predictions for the probability of reversal are dependent on the underlying
₁₆₄ stochastic model. The stochastic model used here is constructed from two compilations
₁₆₅ of paleomagnetic measurements. The noise term is inferred from the high-resolution
₁₆₆ CALS10k.2 model [*Constable et al.*, 2016], and the drift term is recovered from the longer
₁₆₇ PADM2M model [*Ziegler et al.*, 2011]. It is encouraging that our stochastic model also
₁₆₈ gives reasonable predictions for the mean reversal rate ($r = 4.2 \text{ Myr}^{-1}$), as well as the
₁₆₉ duration of polarity transitions [*Buffett and Puranum*, 2017], even though this information
₁₇₀ was not used in constructing the model. This consistency lends support to the model.

171 On the other hand, the model is not capable of reproducing systematic features prior
172 to a reversal [e.g. *Valet et al.*, 2012] because the evolution of the dipole into a reversal is
173 driven by random noise. Extensions of the model are required to account for any recurrent
174 behavior prior to a reversal. Other sources of information could also be incorporated into
175 the model to improve the predictions. A notable example is the evidence for asymmetry
176 in the growth and decay of the geomagnetic dipole [*Avery et al.*, 2017], which could
177 be related to the shape of the drift term. In summary, the stochastic model provides
178 a framework for combining a variety of paleomagnetic observations into a quantitative
179 statistical description of the field behavior. Current estimates of the stochastic model
180 suggest that the next geomagnetic reversal is improbable in the next 10 to 20 kyr.

181 **Acknowledgments.** This work is partially supported by the National Science Foun-
182 dation (EAR-1644644). We thank Phil Livermore and Matti Morzfeld for thoughtful
183 comments and suggestions, which improved the paper. All software used to generate the
184 results in this paper are included in an electronic supplement.

References

185 Aubert, J. (2014). Earth's core internal dynamics 1840-2000 imaged by inverse dynamo
186 modelling, *Geophys. J. Int.*, 197, 119-134.

187 Aubert, J. and A. Fournier (2011). Inferring internal properties of Earth's core dynam-
188 ics and their evolution from surface observations and a numerical geodynamo model,
189 *Nonlin. Processes Geophys.*, 18, 657-674.

190 Avery, M., J.S. Gee and C.G. Constable (2017) Asymmetry in growth and decay of the
191 geomagnetic dipole revealed in seafloor magnetization, *Earth Planet. Sci. Lett.*, 467,
192 79-88.

193 Buffett, B., L. Ziegler and C.G. Constable (2013). A stochastic model for palaeomagnetic
194 field variations, *Geophys. J. Int.*, 195, 86-97.

195 Buffett, B. and A. Puranum (2017). Constructing stochastic models for dipole fluctuations
196 from paleomagnetic observations, *Phys. Earth Planet. Inter.*, 272, 68-77.

197 Cox, A. (1969), Geomagnetic reversals, *Science*, 163, 237-245.

198 Constable, C. and M. Korte (2004). Is Earth's magnetic field reversing? *Earth Planet.
199 Sci. Lett.*, 246, 1-16.

200 Constable, C., M. Korte and S. Panovska (2016). Persistent high paleosecular variation
201 activity in southern hemisphere for at least 10,000 years, The complexity of reversals,
202 *Earth Planet. Sci. Lett.*, 453, 78-86.

203 Davidson, P.A. (2016). *An Introduction to Magnetohydrodynamics*, Cambridge University
204 Press, Cambridge.

205 Finlay, C.C., J. Aubert and N. Gillet (2016). Gyre-driven decay of the Earth's dipole,
206 *Nature Comm.*, 7, 10422.

207 Gardiner, C.W. (2002). *Handbook of Stochastic Methods for Physics, Chemistry and the
208 Natural Sciences*, 2nd Edition, Springer, New York.

209 Gee, J.S. and D.V. Kent (2015). Sources of oceanic magnetic anomalies and the geomag-
210 netic polarity timescale, in *Treatise on Geophysics*, 2nd Edition, ed. G. Schubert, 5,
211 419-460.

212 Gillet, N., D. Jault, C.C. Finlay and N. Olsen (2013). Stochastic modeling of the Earth's
213 magnetic field: Inversion for covariances over the observatory era, *Geochem. Geophys.*
214 *Geosys.*, 14, 766-786.

215 Gillet, N., D. Jault and C.C. Finlay (2015). Planetary gyre, time-dependent eddies tor-
216 sional waves, and equatorial jets at the Earth's core surface, *J. Geophys. Res.*, 120,
217 3991-4013.

218 Hulot, G., C. Eymin, B. Langlais, M. Miorari and N. Olsen (2002). Small-scale structure
219 of the geodynamo inferred from Oersted and Magsat, *Nature*, 416, 620-623..

220 Hulot, G., F. Lhuillier, and J. Aubert (2010). Earth's dynamo limit of predictability,
221 *Geophys Res. Lett.*, 37, L06305.

222 Kuang, W., A. Tangborn, Z. Wei and T. Sabaka (2009). Constraining a numerical geody-
223 namo model with 100 years of surface observations, *Geophys. J. Int.*, 179, 1458-1468.

224 Malin, S.R.C. (1996). Historical introduction to geomagnetism, in *Geomagnetism*, vol. 1,
225 Academic Press, London, 1-49.

226 Metman, M.C., P.W. Livermore, and J.E. Mound (2018). The reversed and normal flux
227 contributions to axial dipole decay for 1880-2015, *Phys. Earth Planet. Inter.*, *in press*.

228 Morozfeld, M., A. Fournier and G. Hulot (2017). Coarse predictions of dipole reversals by
229 low-dimensional modeling and data assimilation, *Phys. Earth Planet. Inter.*, 262, 8-17.

230 Meduri, D.G. and J. Wicht (2016). A simple stochastic model for dipole fluctuations in
231 numerical dynamo models, *Front. Earth Sys.*, 4:38, doi: 10.3389/feart2016.00038.

232 Ogg, J.G. (2012), Geomagnetic polarity time scale, in *The Geological Time Scale 2012*,
233 ed. F.M. Gradstein et al., Elsevier Science, Chap. 5, 85-113.

²³⁴ Olson, P.L. (2002). Geophysics: The disappearing dipole, *Nature*, *416*, 591-593.

²³⁵ Olson, P.L., G.A. Glatzmaier, G.A. and R.S. Coe (2011). Complex polarity reversals in a

²³⁶ geodynamo model, *Earth Planet. Sci. Lett.*, *304*, 168-179.

²³⁷ Risken, H., (1989). *The Fokker-Planck Equation*, Springer, New York.

²³⁸ Poletti, W., A.J. Biggin, R.I.F. Trindale, G.A. Hartmann, and F. Terra-Nova (2018).

²³⁹ Continuous millennial decrease of the Earth's magnetic axial dipole, *Phys. Earth Planet.*

²⁴⁰ *Inter.*, *274*, 72-86.

²⁴¹ Valet, J.P., A. Fournier, V. Courtillot, and E. Herrero-Bevera (2012). Dynamical similarity

²⁴² of geomagnetic field reversals, *Nature*, *490*, 89-93.

²⁴³ Van Kampen, N.G., *Stochastic Methods in Physics and Chemistry*, North-Holland, Ams-

²⁴⁴ terdam.

²⁴⁵ Ziegler, L.B., C.G. Constable, C.L. Johnson, and L. Tauxe (2011). PADM2M: A penalized

²⁴⁶ maximum likelihood model of the 0-2 Ma palaeomagnetic axial dipole moment, *Geophys.*

²⁴⁷ *J. Int.*, *184*, 1069-1089.

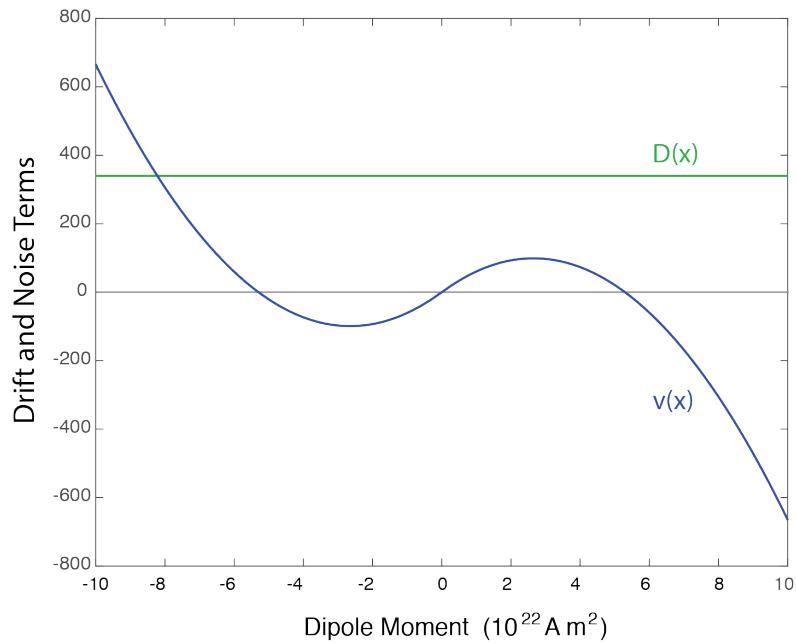


Figure 1. Drift, $v(x)$, and noise, $D(x)$, terms recovered from paleomagnetic observations (see text). The units of $v(x)$ and $D(x)$ are $10^{22} \text{ Am}^2 \text{ Myr}^{-1}$ and $10^{44} \text{ A}^2 \text{ m}^4 \text{ Myr}^{-1}$, respectively.

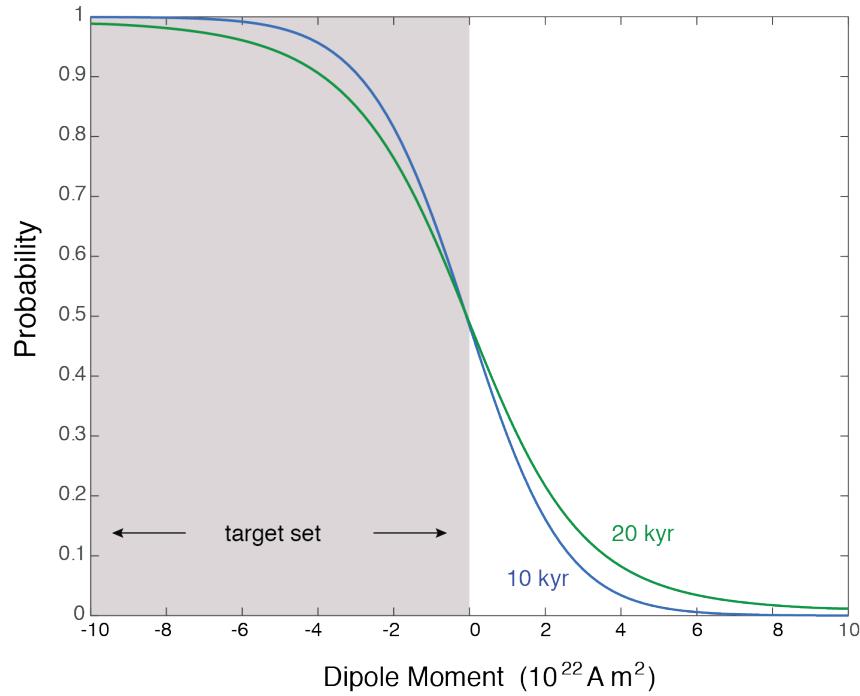


Figure 2. Reversal probability, $u(x', t')$, as a function of x' for two different values of the time difference $\tau = t - t'$. The probability of moving from positive x' to the reversed state $x < 0$ increases with larger time differences.

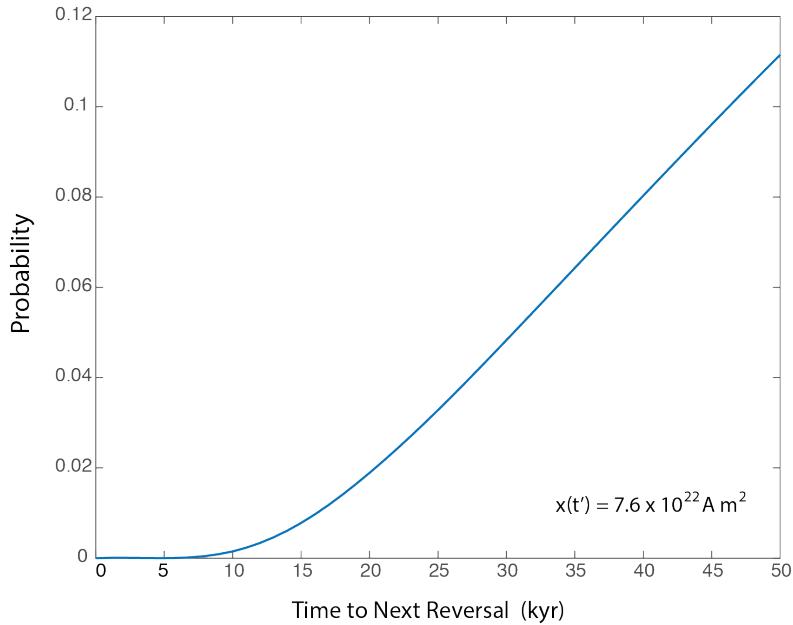


Figure 3. Reversal probability as a function of time difference $\tau = t - t'$ for an initial dipole moment of $x' = 7.6 \times 10^{22} \text{ A m}^2$. The time difference can be interpreted as the time to the next reversal from the present-day state. The predicted probability is low during the first 10 kyr, but increases linearly with τ after 15 kyr and reaches 11% after 50 kyr.