

1 **Emulsified and liquid-liquid phase separated states of α -pinene secondary organic
2 aerosol determined using aerosol optical tweezers**

3 Kyle Gorkowski,¹ Neil M. Donahue,¹ Ryan C. Sullivan^{1,*}

4 ¹Center for Atmospheric Particle Studies, Carnegie Mellon University, Pittsburgh, PA,
5 15213

6 *Corresponding Author: Ryan C. Sullivan; rsullivan@cmu.edu

7

8 **Abstract**

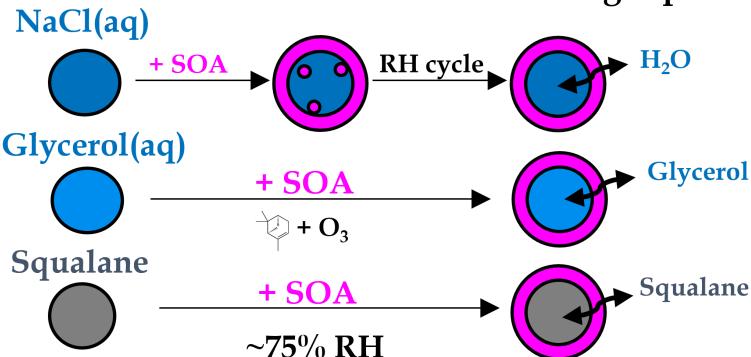
9 We demonstrate the first capture and analysis of secondary organic aerosol (SOA)
10 on a droplet suspended in an aerosol optical tweezers (AOT). We examine three initial
11 chemical systems of aqueous NaCl, aqueous glycerol, and squalane at \sim 75% relative
12 humidity. For each system we added α -pinene SOA – generated directly in the AOT
13 chamber – to the trapped droplet. The resulting morphology was always observed to be
14 a core of the original droplet phase surrounded by a shell of the added SOA. We also
15 observed a stable emulsion of SOA particles when added to an aqueous NaCl core phase,
16 in addition to the shell of SOA. The persistence of the emulsified SOA particles
17 suspended in the aqueous core suggests that this metastable state may persist for a
18 significant fraction of the aerosol lifecycle for mixed SOA/aqueous particle systems. We
19 conclude that the α -pinene SOA shell creates no major diffusion limitations for water,
20 glycerol, and squalane core phases under humid conditions. These experimental results
21 support the current prompt-partitioning framework used to describe organic aerosol in
22 most atmospheric chemical transport models, and highlight the prominence of core-shell
23 morphologies for SOA on a range of core chemical phases.

24

25

26 **Abstract Art**

SOA shell does not limit core equilibration
with gas phase.



27

28

29 KEYWORDS: atmospheric chemistry, organic aerosol, Raman spectroscopy, optical
30 tweezers, whispering gallery modes, phase separations, particulate matter

31

32 **Introduction**

33 Atmospheric particulate matter is a complex mixture of organic carbon, elemental
34 carbon, and inorganic, metallic, and mineral components, all mixed with a variable
35 amount of water^{1,2}. Organic aerosols arise from direct emissions of primary organic
36 aerosol (POA) and from oxidation of precursors – VOCs and evaporated POA – that
37 generates less-volatile secondary organic aerosol (SOA)^{3,4}. Complex SOA is a major
38 component of atmospheric organic aerosol, constituting between 30-70% of the total
39 aerosol mass^{4,5}. It is composed of thousands of individual organic compounds⁶⁻⁸. SOA is
40 a major component of particulate matter, which affects human health and climate
41 change^{9,10}. Aerosol optical tweezers (AOT) provide a new analytical technique to explore
42 the composition and the thermodynamic properties of this complex material using
43 isolated suspended droplets and cavity-enhanced Raman spectroscopy.

44 Liquid particles are common under atmospheric conditions. Depending on the
45 constituents, they can undergo phase separation into distinct chemical phases¹¹⁻¹⁷. We
46 broadly categorize these liquid particles as either homogeneously mixed single-phase
47 structures or liquid-liquid phase separated (LLPS) structures^{15,18-21}. For a phase-separated
48 particle, the two predominant equilibrium states are a core-shell or partially-engulfed
49 arrangement of the phases^{19,22-24}. There is also visual evidence from optical microscopy
50 that phase-separation can result in multiple inclusions forming on or in the aqueous
51 phase, producing a complex morphology^{25,26}. Recent work by Altaf and Freedman²⁷ has
52 shown that under atmospheric conditions there is a size dependent transition such that
53 particles remain in a homogeneous morphology when they are <50 nm in diameter²⁷⁻²⁹.
54 Their observation supports applying the results of phase-separation studies, like this one,
55 to atmospheric particles > 50 nm. The cavity-enhanced Raman spectrum retrieved from
56 a droplet trapped in an AOT instrument provides a unique and real-time morphology
57 measurement that can distinguish between these three structures in mixed droplets³⁰⁻³².

58 Particle morphology is important. The composition at the air-particle interface
59 dictates the particle's ability to undergo chemical reactions with gas-phase reactants and
60 controls the rates and mechanisms of those reactions^{33,34}. N₂O₅ is a key reservoir of
61 nitrogen oxides that demonstrates the importance of this air-particle interface. The
62 reactive uptake coefficient of N₂O₅, $\gamma(N_2O_5)$, governs the N₂O₅ nocturnal removal rate
63 from the atmosphere *via* hydrolysis in aqueous aerosol phases, and a hydrophobic shell
64 on an aqueous core substantially reduces the probability of the reactive uptake^{35,36}.
65 Reactive uptake also increases aerosol mass loadings through aqueous chemistry, for
66 example with isoprene-derived SOA^{37,38}.

67 Using smog-chamber experiments, Ye et al.³⁹ showed that the Hansen solubility
68 framework could inform the miscibility of α -pinene SOA into pure component seeds. The
69 general result from these smog-chamber mixing experiments is that highly polar and
70 nonpolar phases do not mix with SOA, while SOA does dissolve into phases of
71 intermediate or mixed polarity. These techniques used in smog-chamber experiments can

72 infer phase separation from changes in the aerosol mass yield, aerosol size distribution,
73 or the mixing state of components measured using single-particle analysis, but are unable
74 to determine the critical information regarding the resulting particle morphology^{7,39-41}.
75 For a core-shell morphology, diffusion through the shell could dictate the equilibration
76 rates of molecules partitioning into or out of the particle-phase⁴²⁻⁴⁶. This shell effect
77 would not be present in a partially-engulfed morphology.

78 Here we report the first AOT experiments on complex SOA, extending our current
79 understanding regarding the phase separation of SOA using the AOT's ability to
80 distinguish between the three equilibrium morphologies: phase-separated core-shell and
81 partially-engulfed structures, and homogeneous single-phase structures^{15,20,24,30,32,47}.
82 Furthermore, we can observe in real-time any changes in the droplet's morphology while
83 its composition changes^{32,48,49}. The AOT isolates individual micron-sized droplets freely
84 suspended in air and avoids interference from contact with a surface. It does not require
85 collection of the aerosol onto a substrate followed by solvent extraction required by other
86 methods^{12,32,50,51}.

87

88 **Experimental Methods**

89 ***In situ* Ozonolysis of α -pinene in the Aerosol Optical Tweezers**

90 Our AOT system was previously described and characterized by Gorkowski et al.
91 ³². Briefly, a 532 nm laser passes through a set of expansion lenses overfilling the back
92 aperture of a 100x (NA 1.25) oil immersion objective. The objective presses against the
93 bottom of a thin glass coverslip, and an optical gradient force trap forms approximately
94 40 microns above the coverslip inside the AOT chamber. To generate glycerol and
95 squalane droplets we use a Condensational Monodisperse Aerosol Generator (CMAG;
96 TSI, Inc.) and to generate NaCl(aq) droplets we use a medical nebulizer (PARI TREK S).
97 We control the relative humidity (RH) in the chamber by varying the flow ratio of dry to
98 humidified air and we measure RH at the inlet using a hygrometer (Vaisala).

99 Previous AOT work by Dennis-Smither et al.⁴⁷ used direct ozonolysis of a trapped
100 oleic acid droplet, a proxy for primary organic aerosol; we extended the AOT technique
101 to determine the properties and chemistry of SOA produced from the oxidation of
102 precursor vapors. We used dark ozonolysis of α -pinene vapor to produce SOA directly
103 inside the AOT chamber from the formation of less-volatile oxidation products, some of
104 which homogeneously nucleated to form particles. The gas and particle-phase products
105 then collided with the trapped particle, modifying its composition. A diagram of the
106 experimental setup and a detailed description of the procedural methods are provided in
107 the Supplemental Information.

108 During a typical ozonolysis experiment, the ozone concentration in the AOT is 55
109 ppm_v, and the α -pinene vapor concentration is 500 ppm_v, based on the saturated vapor
110 pressure of α -pinene and volumetric flow rates. The ozonolysis of α -pinene proceeds by
111 addition of O₃ to the lone double bond and the reaction is ozone limited under these
112 conditions⁵². This decomposition of the resulting Criegee intermediate goes on to
113 produce ·OH radicals with a yield of ~0.8; under our conditions essentially all of the ·OH
114 reacts with α -pinene⁵². Consequently, the SOA consists of an approximately 0.8:1 mixture
115 of ·OH and ozone oxidation products. We occasionally observe nucleation of detectable
116 particles from the α -pinene oxidation in the form of small-particle light scattering events
117 in the tweezed droplet's visual image. The SOA vapors and particles are produced in the
118 upper section of the AOT chamber and then flow down to the tweezed droplet at the
119 bottom of the AOT chamber. The metal chamber walls do adsorb some of the less volatile
120 SOA components, but these components by necessity are also present in the nucleated
121 particles that coalesce with the tweezed droplet⁵³.

122 **Raman Spectrum Analysis of Core-Shell Droplets**

123 We use a spectrograph to record Raman scattering from the tweezed droplet,
124 including resonant whispering gallery modes (WGMs). At the wavelength of each WGM,
125 the droplet acts as a high finesse optical cavity where the WGMs form a standing wave
126 near the droplet surface, amplifying the Raman scattering^{54,55}. Any radial

127 inhomogeneities in the droplet will weaken the amplification of the WGMs. For partially-
128 engulfed morphologies the WGM amplification process is completely quenched^{20,47}. The
129 WGM amplification process will remain for core-shell morphologies as it is radially
130 symmetric^{15,32}.

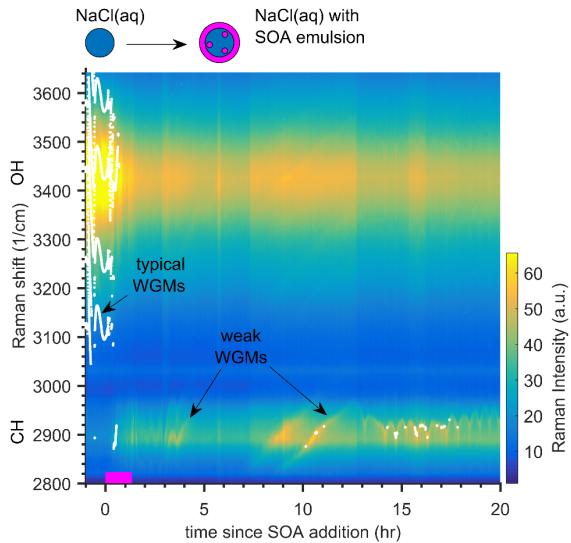
131 To distinguish the morphology of a core-shell versus a homogeneous droplet, we
132 use a Mie scattering algorithm based off the one developed by Preston and Reid^{56,57}.
133 When applied to a core-shell droplet we use the plausibility of the retrieved effective
134 refractive index and diameter parameters as well as an abnormally high fit error to
135 identify a droplet with a core-shell morphology^{15,32}. Fitting WGMs of a core-shell droplet
136 to a homogeneous Mie model results in a consistently high fit error. To assess whether
137 the core-shell morphology is valid we have developed a fitting algorithm for the WGMs
138 from a core-shell morphology that retrieves the diameter and effective refractive index of
139 both the core and shell phases, separately. Both the homogeneous and core-shell fit only
140 retrieve the real part of the effective refractive index. We provide an accuracy
141 characterization for the core-shell fitting in the Supplemental Information and will fully
142 describe the fitting algorithm in a subsequent manuscript.

143

144 **Results**

145 **Emulsions and Coatings of SOA on an Aqueous Inorganic Droplet**

146 SOA and SOA mimics have been observed to phase separate from an aqueous
147 inorganic phase in large droplets examined visually on microscope slides^{14,17,58}. We
148 performed experiments in the AOT on similar systems composed of SOA mixed with
149 aqueous NaCl to determine the conditions under which SOA phase separates from
150 aqueous phases. We determined the resulting droplet morphology using the cavity-
151 enhanced Raman spectrum retrieved by the AOT and thus avoided interference through
152 contact with any surface.



153

154 **Figure 1.** Raman spectral time series for a tweezed aqueous NaCl droplet before, during
 155 (pink bar at the bottom), and after α -pinene SOA was generated *in situ*, at 78% RH. The
 156 C-H hydrocarbon Raman mode center is at 2900 cm^{-1} , the O-H water mode is at 3450 cm^{-1} ,
 157 and the color scale indicates the intensity of the Raman signal at that Raman shift
 158 position. The bright white points indicate the positions of detectable WGMs. The SOA
 159 flow reaction started at 0 min and lasted for 1.3 hr, indicated by the pink bar. The weak
 160 WGMs following the SOA addition are visible by eye but remain below our peak
 161 detection limit for 48 hours; only the first 20 hours are shown here. The small waves (with
 162 a period of $\sim 0.5\text{ hr}$) in the WGMs after 14 hr are caused by small temperature/humidity
 163 fluctuations ($\sim 1\%$ RH) due to the laboratory's HVAC system.

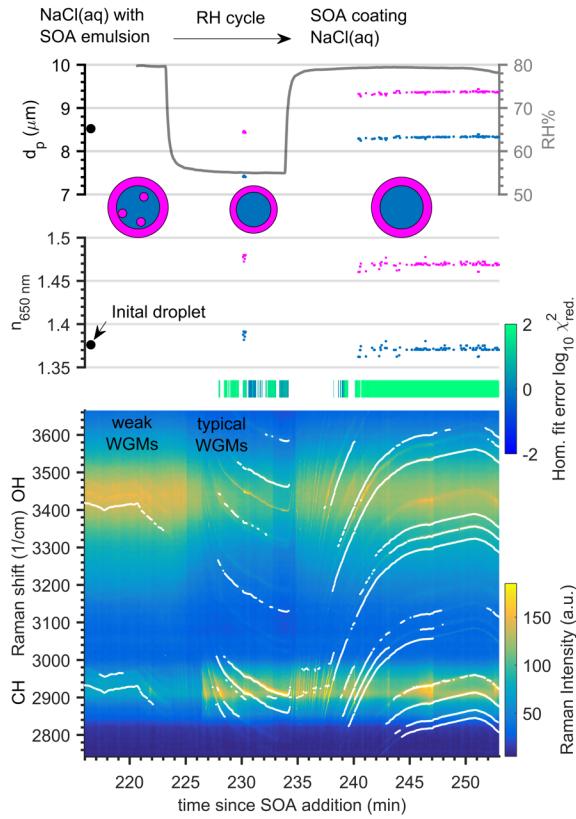
164 In Figure 1 we show data from an experiment in which we added SOA generated
 165 *in situ* to a tweezed NaCl(aq) droplet beginning at 0 minutes. We observe typical WGM
 166 resonances in the Raman spectrum in the broad O-H water stretch region before the SOA
 167 addition ($t \leq 0\text{ min}$); these correspond to a homogeneous morphology. The droplet's size
 168 was $7.308 \pm 0.001\text{ }\mu\text{m}$, and its effective refractive index ($n_{650\text{nm}}$) was 1.380 ± 0.0015 just
 169 before SOA addition. During and after the SOA addition, Raman scattering appears in
 170 the C-H hydrocarbon region due to the uptake of organic carbon, and weak WGMs are
 171 present in both the O-H and C-H Raman regions. The WGMs are weaker following SOA
 172 addition but still faintly visible in the Raman spectrum's time series; their intensity is
 173 diminished but not completely quenched.

174 The persistence of weak WGMs following addition of SOA to an aqueous NaCl
175 drop means we cannot conclude that the mixed droplet formed a partially-engulfed
176 structure, because that would completely quench the WGMs. Theoretical modeling of a
177 spherical droplet with an emulsion showed that as the number of emulsified particles
178 increased the WGMs are weakened and become broader⁵⁹. Experiments using an
179 electrodynamic balance showed that adding polystyrene latex spheres (d_p of 30, 64, or
180 105 nm) to a supermicron glycerol droplet weakens the WGM signal⁶⁰. The
181 heterogeneities in the radial path of the resonating wavelengths of light would scatter the
182 resonating WGM and intermittently quench the WGM signal. Some possible
183 configurations of these heterogeneities are shown in the supplemental information; they
184 are all classified as a type of emulsion. As we observe some WGM amplification, we
185 propose that the droplet consists of a homogeneous SOA shell coating a supermicron
186 aqueous NaCl core containing emulsified SOA particles retained during the coalescence
187 of SOA particles onto the tweezed droplet. Along with the emulsion, the more polar
188 compounds in the SOA are water soluble, so a small fraction of the SOA should also
189 dissolve into the aqueous phase. Two-dimensional microscopy images in the literature
190 support this proposed shell phase with an emulsion of SOA in an aqueous core^{14,17,58}.
191 However, it is difficult to resolve the 3-dimensional morphology from those microscopy
192 images because they only show one viewing angle. This is not a limitation of the AOT. In
193 the Supplemental Information, we demonstrate that radial heterogeneities can weaken
194 the WGM amplification by coagulating NaCl crystals onto a trapped squalane droplet.
195 This further supports our SOA emulsion interpretation of the Raman spectrum in Figure
196 1.

197 To further test our emulsion hypothesis, we waited for 48 hours (the first 20 hours
198 are shown in Figure 1) to see if the emulsion would relax (Ostwald ripening) since
199 emulsified particles represent a metastable state⁶¹. We continued the flow of humidified
200 clean air to maintain the aqueous phase, so semi-volatile SOA should have evaporated to
some extent from the droplet. The relatively stable WGM positions with time (Figure 1,

202 10 to 20 hours) shows that the droplet diameter remained nearly constant, indicating the
203 remaining SOA did not evaporate to a noticeable extent.

204 In a second experiment on the same SOA + NaCl(aq) aerosol system, we used an
205 RH cycle on the emulsified core-shell droplet to test the effect of dehydrating the aqueous
206 core (which presumably contained the emulsified particles), as this might promote
207 demulsification. The Raman spectrum and extracted WGMs positions measured in this
208 experiment are shown in the bottom of panel of Figure 2 and the effective refractive
209 indexes and diameters fit from the WGM positions are shown in the top two panels. We
210 use the homogeneous fit error (blue/green bar in Figure 2) and inconsistent fit
211 parameters to indicate an invalid homogeneous fit result. The initial aqueous NaCl
212 droplet was $8.522 \pm 0.001 \mu\text{m}$ with an $n_{650\text{nm}}$ of 1.376 ± 0.0015 , after which we added the
213 SOA. We subsequently cycled the RH from 80% down to 55% and then back up to 79%
214 (Figure 2). After the RH decreased at 224 min, 4 or 6 relatively strong WGMs appeared in
215 the spectrum. The spectra during the drier period also exhibit more typical WGM
216 amplification. After we increased the RH back up to 79% at 235 min, the strong WGMs
217 typical of a homogeneous spherical geometry intensified, with 6 or more strong WGMs
218 visible in a 2-second frame. Isolated Raman spectra covering this transition period are
219 provided in the Supplemental Information.



220

221 **Figure 2.** Raman spectra time series for a tweezed aqueous NaCl droplet after
 222 coagulation with α -pinene SOA, undergoing a relative humidity (RH) cycle. The bottom
 223 graph is the time series of the Raman spectra, the middle is the refractive index at 650 nm,
 224 and the top is the droplet diameter (left axis) and conditioned air's relative humidity
 225 (right axis). The time series started after the α -pinene SOA addition which happened at 0
 226 min. and lasted for 65 min. The initial homogeneous NaCl droplet fit results are shown
 227 as black circles. WGMs that meet the peak criteria are shown as white points/lines in the
 228 Raman spectral time series. The fit error for the homogeneous morphology model is
 229 shown in the green/blue bar, where consistently green indicates an invalid fit to the
 230 homogeneous model. The fitted core-shell droplet diameter and refractive index are
 231 shown above for both the core (blue dots) and shell (pink dots) phases. The retrieved
 232 $n_{650\text{nm}}$ values are consistent with a core of NaCl(aq) and a shell of α -pinene SOA. The
 233 morphologies we infer during this experiment are displayed where an emulsion of SOA
 234 in the aqueous core is relaxed to a core-shell morphology via the RH cycle.

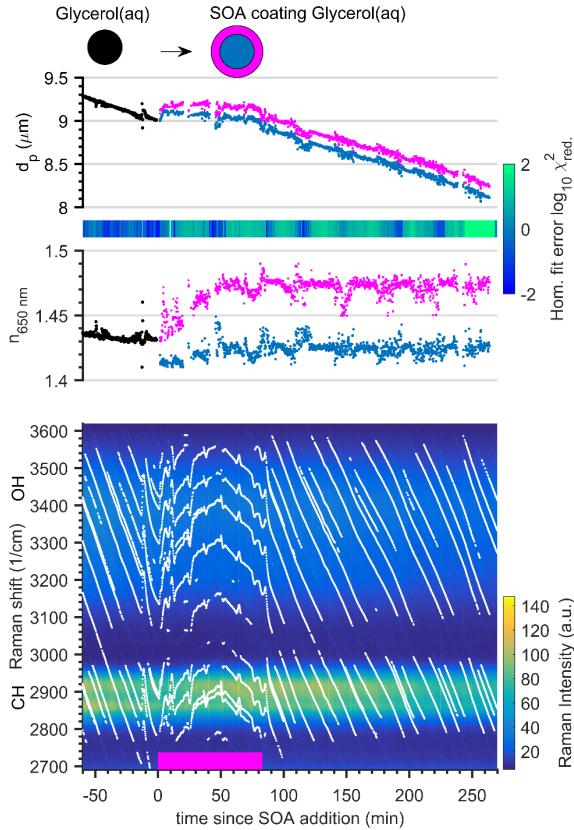
235

236 When we fit the WGMs after the RH cycle with the homogeneous Mie model
 237 appropriate for a single phase, the results are inconsistent and have high fit error,

238 $\log_{10}(\chi^2_{red}) > 0$ (blue/green bar in Fig. 2). When we fit the WGMs using the core-shell
239 Mie model the results are consistent with what we would expect based on the initial
240 conditions, the RH cycling, and the starting droplet size, shown in Figure 2. The refractive
241 index of the α -pinene SOA shell ($n_{650\text{nm}} = 1.469 \pm 0.004$) is well within the range reported
242 in the literature ($n_{532\text{nm}} = 1.39\text{-}1.52$); see the Supplemental Information for a comparative
243 table⁶²⁻⁶⁶. The core refractive index, $n_{650\text{nm}}$, of 1.372 ± 0.005 is within the uncertainty of the
244 initial droplet ($n_{650\text{nm}} = 1.376 \pm 0.0015$), suggesting that no measurable SOA dissolved into
245 the core. The solution has an NaCl mole fraction of 0.16, calculated via E-AIM^{8,67}. By using
246 a volume fraction mixing rule we estimated the maximum amount of SOA that could
247 dissolve and still be within the error of our fit. For the core refractive index this means
248 the volume fraction of SOA in the core is less than 4.3%. The droplet size is quite stable,
249 as shown in the figure; though the WGM positions after the RH cycle at 235 min do
250 change with time, 1 cm^{-1} corresponds to roughly 2 nm in diameter, so the growth is a
251 fraction of a micrometer. The core-shell fit results support a morphology where α -pinene
252 SOA coats the NaCl(aq) core. Based on these results we conclude that coalescence of
253 submicron SOA particles onto a high ionic strength aqueous NaCl droplet results in a
254 droplet coated with a spherical shell but containing emulsified SOA particles as well. The
255 subsequent RH cycle causes the droplet to assume an equilibrium core-shell morphology
256 with an SOA shell, no residual emulsified SOA, and little dissolved SOA.

257 **Morphology of an Aqueous-Organic Mixed Droplet**

258 A natural extension of SOA coalescence to an aqueous inorganic core is SOA
259 coalescence to an aqueous-organic core. In this way, we can isolate the role of water in
260 the phase separation and remove the salting out effect of the relatively high ionic strength
261 aqueous inorganic solution. We use glycerol as it is hygroscopic and is active in both the
262 C-H and O-H Raman modes.



263

264 **Figure 3.** Raman spectral time series for a tweezed aqueous glycerol droplet with the
 265 addition of α -pinene SOA, at 73% RH. Raman spectra are shown in the bottom graph,
 266 and the white points indicate the positions of detectable whispering gallery modes
 267 (WGMs). The decreasing slope of the WGM traces indicates the droplet is evaporating.
 268 The α -pinene SOA flow reaction started at 0 min. and lasted for 80 min., indicated by the
 269 pink bar. The droplet's refractive index and diameter retrieved from the WGM analysis
 270 are shown in the middle and top panels. The fit results of the WGM positions to a
 271 homogeneous Mie model are in the top two plots as black dots. The fit error for the
 272 homogeneous model is shown in the green/blue bar, where consistently blue indicates a
 273 valid fit and green indicates an invalid fit to the model. The core-shell fitting is then used
 274 to retrieve the properties of the glycerol(aq) core (blue points) and the α -pinene SOA shell
 275 after SOA generation started at 0 min.

276

277 We trapped a glycerol droplet at 73% RH, finding $n_{650\text{nm}}$ of 1.4314 ± 0.0013 with a
 278 homogeneous Mie model. In Figure 3 it is apparent that, before the α -pinene SOA
 279 addition at $t = 0$, the droplet was shrinking rapidly. The retrieved diameter and the

negative slope of the WGM positions indicate evaporation of the glycerol (water volume fraction should remain constant because of the constant RH). When we added α -pinene SOA to the droplet, the homogeneous Mie fitting of the WGMs failed immediately, as indicated by the increased fit error of the homogeneous model. However, the core-shell Mie model could successfully fit the WGMs both during and after the SOA addition. The core-shell refractive index results indicate that a shell of α -pinene SOA ($n_{650\text{nm}} = 1.474 \pm 0.005$) coated the glycerol(aq) core ($n_{650\text{nm}} = 1.4325 \pm 0.007$). The refractive index of the SOA shell is consistent with the NaCl(aq) experiment described above and within the literature range (summarized in the Supplemental Information). As with the high ionic strength NaCl(aq) droplet, we observed no measurable dissolution of the SOA into the aqueous organic core; the volume fraction of SOA in the aqueous glycerol solution was < 17% based on the measured $n_{650\text{nm}}$. In contrast to the NaCl(aq) system, we do not observe evidence of an emulsion, as strong WGMs persisted throughout the experiment.

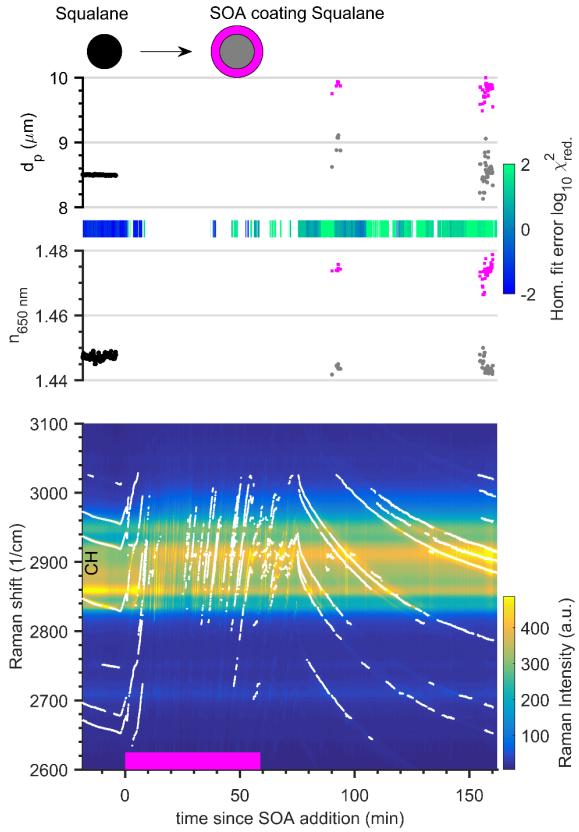
The continued evaporation of the core indicates that the α -pinene SOA did not prevent or retard evaporation of glycerol from the core. The measured evaporation rate for the glycerol(aq) core decreased by 22% after we added the α -pinene SOA. This decrease may be due to Raoult's Law of SOA mixing in the glycerol(aq) core depressing the overall evaporation rate of the core. That decrease could also be due to a diffusion limitation of glycerol through the SOA shell. If the SOA dissolved in the core was at our upper limit of 17% by volume, Raoult's Law could explain the majority of the decrease in evaporation rate. We will systematically explore these confounding effects of mixing and diffusion limitations in the glycerol plus SOA system in future experiments by varying the amount of SOA added to the droplet. One thing that is clear from this experiment is that the SOA shell does not trap glycerol in the SOA phase and does not prevent glycerol from partitioning to the gas phase.

We can also use the evolution of the shell thickness, as shown in Figure 3, to calculate the evaporation rate and thus the effective vapor pressure of the SOA shell. To do this, we calculated the volume loss of the shell for the period between 160 min. and

260 min. and assumed an SOA molecular weight of 300 g/mol and a density of 1.4 g/cm³
68-70. There was very little evaporation, and so the effective vapor pressure of the SOA
shell was 5.2×10^{-6} Pa or equivalently a C^* of 0.64 $\mu\text{g}/\text{m}^3$ ($10^{-0.193}$ $\mu\text{g}/\text{m}^3$)^{48,53,71,72}. This
retrieved volatility is higher due to our high precursor conditions than the measurements
made during the α -pinene SOA evaporation experiment described by Yli-Juuti et al.⁷³,
with $C^* \approx 10^{-3}$ $\mu\text{g}/\text{m}^3$ at 80% RH. Lastly, in contrast to typical SOA evaporation
experiments, our shell phase does not appear to contain a large portion of intermediate-
volatile organic carbon with higher C^* values as the evaporation rate we observed was
constant. This could be due to additional aqueous chemistry of the SOA at the aqueous-
core interface, leading to lower volatility SOA in the shell phase⁷⁴.

318 Morphology of SOA Added to a Hydrophobic Organic Droplet

Having studied two aqueous cores, next we studied the mixing of SOA with a
tweezed hydrophobic squalane ($\text{C}_{30}\text{H}_{62}$) droplet. Our objective was to replicate the
chemical system studied in a smog-chamber mixing experiment conducted by Robinson
et al.⁷. They studied the same particle system by condensing α -pinene ozonolysis SOA
onto squalane seed particles and also condensing squalane vapors from a heated injector
onto SOA particles. After increasing the chamber temperature by 22 °C, they observed
rapid evaporation of the squalane and little evaporation of the SOA. They also found that
squalane and α -pinene SOA remained phase-separated during a particle mixing
experiment. Based on these experiments they concluded that the mixed particles formed
a two-phase morphology, noting that if the particles had a core-shell morphology with
the squalane forming the core, the evaporating squalane would have to diffuse relatively
rapidly through the SOA shell. That single-particle smog-chamber experiment could not
fully constrain the phase-separated particle morphology, leaving the hypothesis that
squalane can diffuse rapidly through SOA unproven. If the morphology were partially-
engulfed, then squalane would have direct access to the gas phase. The aerosol optical
tweezers provide a unique way to determine the morphology of this complex system
directly.



336

337 **Figure 4.** Raman spectral time series for a tweezed squalane droplet and addition of α-
338 pinene SOA, at 80% RH, following the same format as in Fig. 3. A decreasing trend of the
339 WGM traces indicates the droplet was evaporating whereas an increase indicates droplet
340 growth. The α-pinene SOA addition started at 0 min. and lasted for 60 min., indicated by
341 the pink bar. The fit results of the WGM positions to a homogeneous model are shown in
342 the top two plots as black dots. The core-shell fitting model is then used to retrieve the
343 properties of the squalane core (gray) and α-pinene SOA shell (pink) following SOA
344 addition. Few WGMs are retrieved as the C-H hydrocarbon band is too narrow to
345 consistently contain at least 6 WGMs that are required for the core-shell Mie fit.

346 In Figure 4 we show spectra from an experiment starting with a pure squalane
347 droplet, with an initial diameter of $8.492 \pm 0.004 \mu\text{m}$ and a refractive index, $n_{650\text{nm}}$, of 1.447 ± 0.001 , to which we added α-pinene SOA by coagulation. The WGMs were initially
348 almost horizontal, with a slight negative slope due to the slow evaporation of squalane,
349 which has an estimated vapor pressure of $1.93 \times 10^{-6} \text{ Pa}$ ⁷⁵. When we began to add α-pinene
350 SOA, the WGM slopes became strongly positive, indicating that the particle was growing.
351 The persistence of the WGMs also indicates that the SOA spread rapidly across the
352

353 squalane surface as it condensed and coagulated, excluding a partially-engulfed
354 morphology. After the α -pinene SOA addition at 70 min., we held the chamber static with
355 all airflows off before turning on clean airflow at 80 min.; the droplet then began to shrink,
356 indicated by the negative slope of the WGMs. The fit results for the core-shell droplet are
357 sparse because the droplet contains mostly C-H bonds, which have a narrow Raman band
358 reducing the number of WGMs observed. However, the few fits are consistent with a
359 squalane core coated by SOA. Moreover, the slope of the WGMs at the end of the
360 experiment is similar to the slope at the beginning, suggesting that slow squalane
361 evaporation (few nm per minute) dominates the particle evolution. The slightly higher
362 slope soon after the SOA addition also suggests that some SOA evaporated from the shell.

363 Consistently high homogeneous Mie fit error following the addition of α -pinene
364 SOA indicates that the droplet was not homogeneous, while the persistence of the WGMs
365 rules out a partially-engulfed morphology. Based on this evidence we conclude that the
366 droplet was phase separated, which is consistent with the smog-chamber studies^{7,39}. Our
367 AOT results add the additional insight that the particle morphology is core-shell with α -
368 pinene SOA coating squalane. This direct measurement of the morphology is inaccessible
369 from traditional smog-chamber observations of a large ensemble of particles. The core
370 refractive index, $n_{650\text{nm}}$, of 1.444 ± 0.003 and diameter of $8.577 \pm 0.186 \mu\text{m}$ provide evidence
371 that no measurable amount of SOA was dissolved into the core (volume fraction < 11%).
372 The retrieved refractive index of the α -pinene SOA shell ($n_{650\text{nm}} = 1.473 \pm 0.003$) is
373 consistent with the previous two experiments on aqueous NaCl and glycerol.

374 Note that squalane lacks any unsaturated bonds for ozone to attack. Squalane can
375 be oxidized by the $\cdot\text{OH}$ radicals produced by the ozonolysis of α -pinene, but this
376 heterogeneous oxidation would be much slower than the gas-phase oxidation of α -pinene
377 by $\cdot\text{OH}$ due to mass transfer limitations from the gas phase to particle interface⁷⁶. The
378 squalane core also quickly became protected by the SOA coating.

379

380 **Discussion**

381 Our extension of the AOT technique to generate and study complex α -pinene SOA
382 *in situ* enables real-time investigations of complex aerosol systems at the individual
383 particle level that more realistically mimic atmospheric aerosol compositions. Fitting of
384 the WGMs provides highly accurate measurements of the droplet's size and refractive
385 index that allow us to determine how different droplet compositions respond to the
386 addition of SOA. These AOT experiments on *in situ* generated α -pinene SOA avoid
387 limitations intrinsic to other measurements of phase separation and morphology using
388 SOA filter extracts, where the solvent can alter the composition and only recovers the
389 soluble portion, such as the water-soluble fraction of SOA^{25,77-79}.

390 The mixing experiments on aqueous organic, aqueous inorganic, or hydrophobic
391 organic particle phases with SOA all resulted in a liquid-liquid phase separation,
392 demonstrating that phase separation is common for the aerosol mimics probed here. We
393 also add further evidence to indicate that core-shell morphologies are common structures
394 for complex mixed atmospheric aerosol mimics^{21,80}. Our work here focused on high
395 humidity ranges of 70% to 80% RH to isolate the effect of the starting droplet
396 composition; in future work, we will explore the effect of RH on the resulting phase-
397 separated morphology.

398 The high precursor concentrations used here are not representative of the ambient
399 troposphere, and as such may affect our resulting SOA oxidative state and volatility. This
400 high precursor effect appears to have shifted our SOA volatility higher as our measured
401 SOA volatility was greater than observed in an evaporation experiment by Yli-Juuti et al.
402 ⁷³. The higher volatility SOA products that partitioned to the particle phase would likely
403 result in a lower viscosity⁸¹. More volatile SOA products are typically less oxidized, and
404 the less polar nature of these molecules reduces the strength of the inter-molecular
405 interactions they experience, reducing their viscosity while increasing their volatility.

406 Through the core-shell fitting of our different SOA systems, we also retrieve the
407 refractive index to a 0.4% accuracy. Our retrieved refractive index is valuable as previous
408 optical closure studies on SOA systems yielded a wide range of refractive indices Table
409 S2). The variability is in part due to the methods used where a new volume of particles is
410 measured at each data point. In the AOT the environment is more controlled, and we
411 measure the properties of the same SOA mass in real-time, which reduces experimental
412 variability between measurements.

413 Synthesizing the results from the different RH stages of the aqueous NaCl droplet
414 plus SOA experiment strengthens our proposal for the emulsion of SOA particles in the
415 aqueous core. Dehydration of the aqueous core demulsified the SOA and produced a
416 stable irreversible core-shell morphology of SOA coating the NaCl(aq). From this, we can
417 conclude that the core-shell morphology observed is the thermodynamically stable
418 condition for this mixed system, and the emulsion was metastable. Therefore, before the
419 RH cycle, it is likely that a core-shell morphology was thermodynamically favored then
420 as well. Finally, a droplet morphology composed of a homogenous SOA shell and the
421 inhomogeneous core of emulsified SOA particles in aqueous NaCl would also explain the
422 diminished WGM amplification observed. This homogeneous shell and inhomogeneous
423 core would account for the resonance of WGMs in the surface layer of the SOA shell,
424 while not permitting the full amplification of the WGMs due to interference by the
425 emulsion in the core. The prompt collapse of the emulsified state following the decrease
426 in RH suggests a chemical rather than physical explanation for the rapid demulsification.
427 The increase in the ionic strength of the NaCl(aq) core caused by its partial dehydration
428 (Fig. 2) would significantly change the interfacial energies between the SOA and
429 NaCl(aq) phases. This likely caused the emulsified SOA particles to destabilize and the
430 organic material to largely partition to the SOA shell around the NaCl(aq) core.

431 It is also notable that the SOA appears to be relatively insoluble in this aqueous
432 solution, even though isolated α -pinene SOA is known to be reasonably hygroscopic⁸². If
433 the solubility of SOA is greatly diminished in mixed composition particles (which

434 pervade the atmosphere) this has significant implications for the CCN activity of these
435 particles; the aqueous solution will be much more dilute at the Köhler maximum in the
436 droplet growth curve under supersaturated conditions than at 80% RH.

437 The glycerol(aq) plus SOA experiment is especially interesting when taken into
438 the context of the NaCl(aq) plus SOA experiment. First, it shows that in both experiments
439 there is limited solubility of the SOA in the initial droplet phase, leading to the phase
440 separation. Second, both experiments added SOA to an aqueous droplet, but the
441 NaCl(aq) system formed an emulsion in the core whereas the glycerol(aq) core did not.
442 The physical formation of the emulsion is due to SOA particles coagulating onto the
443 surface of the droplet, but that does not explain why it is a stable emulsion as shown in
444 Figure 1. The same coagulation of homogeneously nucleated SOA is occurring with the
445 glycerol(aq) droplet yet for that system SOA promptly starts to form a coating without
446 creating an emulsion. Differences in the ionic strength of the NaCl(aq) versus glycerol(aq)
447 phases and the interfacial energies between the SOA and the two aqueous phases likely
448 explains why a stable emulsion only exists in the NaCl(aq) phase. In the atmosphere, this
449 emulsified state would increase exposed SOA surface area to the aqueous medium and
450 thus give more SOA the opportunity to undergo aqueous chemistry. The prevalence of
451 the SOA emulsion may be limited to high number concentration regimes where
452 coagulational loss on to larger particles is significant, such as new particle formation
453 events that frequently occur, particularly near forested regions with high biogenic VOC
454 emissions^{78,83}. It would also delay the formation of a stable core-shell morphology where
455 all the undissolved SOA resides at the air-particle interface. Considering our observations
456 of a stable emulsified state for 48 hours, RH cycling in the atmosphere may play an
457 important role in determining the morphology of mixed SOA and aqueous inorganic
458 particles, and the prevalent atmospheric particle mixing state.

459 The AOT experiment on the α -pinene SOA coating squalane system refines the
460 interpretation of smog-chamber experiments reported by Robinson et al.⁷. They also
461 observed phase separation of squalane and SOA, which our AOT experiments confirm.

462 Our experiment provides the additional observation that the phase-separated particle is
463 a core-shell morphology, an important property that was unobservable in the smog
464 chamber experiments. We have determined that a shell of α -pinene SOA surrounding the
465 squalane core is the thermodynamically stable state for this mixture. Now we can
466 definitively interpret the results from the aerosol evaporation experiments performed by
467 Robinson et al.⁷. They observed prompt evaporation of squalane (which is more volatile
468 than the oxidized SOA) which could indicate that the diffusion coefficient of the squalane
469 in SOA phase is sufficiently high to promote the transport of squalane through the SOA
470 as it evaporates. This interpretation requires the SOA to be coating the squalane phase,
471 which our AOT experiments now confirm is the equilibrium morphology. Combining the
472 results from our AOT experiments and their chamber experiments we can conclude that
473 squalane has no major diffusion limitation through α -pinene SOA, at moderate humidity.

474 In addition to squalane, we also observe glycerol evaporating through an SOA
475 shell in Figure 3. From that experiment, we conclude that glycerol has no major diffusion
476 limitation through α -pinene SOA, at moderate humidity. Finally, from the rapid WGM
477 response to the RH cycle of the SOA coating an NaCl(aq) droplet, we conclude that water
478 also has no major diffusion limitation through α -pinene SOA. The shell thickness of the
479 α -pinene SOA on the glycerol(aq) and NaCl(aq) droplets was 65 nm and 519 nm,
480 respectively, thus there would be minimal diffusion limitations for particles of similar
481 sizes composed only of α -pinene SOA. The smog-chamber mixing experiments of
482 squalane and α -pinene SOA directly probed atmospherically relevant sizes at 5% RH and
483 observed no diffusion limitations through the α -pinene SOA⁷. If these results hold for
484 other OA systems then assuming prompt partitioning under moderate humidities in
485 atmospheric chemical transport models is valid when the model time steps are on the
486 order of minutes^{46,53,84-86}. A forthcoming paper will explore in detail the physical
487 properties of the SOA observed from these and additional AOT experiments, such as the
488 volatility, diffusivity, viscosity, and surface tension of the SOA.

489

490 **Acknowledgements.** Kyle Gorkowski was supported in part by a Bertucci Fellowship
491 from Carnegie Mellon University's College of Engineering. This research was supported
492 by the National Science Foundation (Awards CHE-1213718 & CHE-1554941). We thank
493 Jonathan Reid and James Walker at the University of Bristol for valuable discussions
494 regarding the analysis and interpretation of the AOT Raman spectra.

495 **Supporting Information.** Includes a detailed schematic of the aerosol optical tweezers
496 experimental system, assessment of the core-shell fitting algorithm, details of the
497 evaporation rate calculation, Raman spectra from a squalane droplet containing solid
498 NaCl crystals, isolated frames of the Raman spectra from the NaCl(aq) plus SOA
499 experiment, and a tabulation of reported refractive index values for α -pinene SOA.

500

501 **References**

502 (1) Kolb, C. E.; Worsnop, D. R. Chemistry and Composition of Atmospheric Aerosol
503 Particles. *Annu. Rev. Phys. Chem.* **2012**, *63* (1), 471–491.

504 (2) Pöschl, U. Atmospheric aerosols: composition, transformation, climate and health
505 effects. *Angew. Chem. Int. Ed. Engl.* **2005**, *44* (46), 7520–7540.

506 (3) Gentner, D. R.; Isaacman, G.; Worton, D. R.; Chan, A. W. H.; Dallmann, T. R.;
507 Davis, L.; Liu, S.; Day, D. A.; Russell, L. M.; Wilson, K. R.; et al. Elucidating
508 secondary organic aerosol from diesel and gasoline vehicles through detailed
509 characterization of organic carbon emissions. *Proc. Natl. Acad. Sci. U. S. A.* **2012**.

510 (4) Hallquist, M.; Wenger, J. C.; Baltensperger, U.; Rudich, Y.; Simpson, D.; Claeys,
511 M.; Dommen, J.; Donahue, N. M.; George, C.; Goldstein, A. H.; et al. The
512 formation, properties and impact of secondary organic aerosol: current and
513 emerging issues. *Atmos. Chem. Phys.* **2009**, *9* (14), 5155–5236.

514 (5) Zhang, Q.; Jimenez, J. L.; Canagaratna, M. R.; Allan, J. D.; Coe, H.; Ulbrich, I.;
515 Alfarra, M. R.; Takami, A.; Middlebrook, A. M.; Sun, Y. L.; et al. Ubiquity and
516 dominance of oxygenated species in organic aerosols in anthropogenically-
517 influenced Northern Hemisphere midlatitudes. *Geophys. Res. Lett.* **2007**, *34* (13),
518 L13801.

519 (6) Surratt, J. D.; Murphy, S. M.; Kroll, J. H.; Ng, N. L.; Hildebrandt, L.; Sorooshian,
520 A.; Szmigielski, R.; Vermeylen, R.; Maenhaut, W.; Claeys, M.; et al. Chemical

521 Composition of Secondary Organic Aerosol Formed from the Photooxidation of
522 Isoprene. *J. Phys. Chem. A* **2006**, *110* (31), 9665–9690.

523 (7) Robinson, E. S.; Saleh, R.; Donahue, N. M. Probing the evaporation dynamics of
524 mixed {SOA}/squalane particles using size-resolved composition and single-
525 particle measurements. *Environ. Sci. Technol.* **2015**, *49*, 9724–9732.

526 (8) Zhao, Y.; Wingen, L. M.; Perraud, V.; Finlayson-Pitts, B. J. Phase, composition,
527 and growth mechanism for secondary organic aerosol from the ozonolysis of
528 alpha-cedrene. *Atmos. Chem. Phys.* **2016**, *16* (5), 3245–3264.

529 (9) IPCC. *Climate Change 2013: The Physical Science Basis. Contribution of Working Group*
530 *I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*;
531 Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels,
532 Y. Xia, V. Bex, P. M. M., Ed.; Cambridge University Press: Cambridge, 2013.

533 (10) Pope, C. A.; Ezzati, M.; Dockery, D. W. Fine-particulate air pollution and life
534 expectancy in the United States. *N. Engl. J. Med.* **2009**, *360* (4), 376–386.

535 (11) Hekayati, J.; Roosta, A.; Javanmardi, J. Liquid–liquid equilibria in the quinary
536 aqueous two-phase system of poly(ethylene glycol) 6000+sodium sulfate+water
537 in the presence of glucose and ethanol: Experimental investigation and
538 thermodynamic modeling. *Thermochim. Acta* **2016**, *625*, 47–52.

539 (12) Bateman, A. P.; Gong, Z.; Liu, P.; Sato, B.; Cirino, G.; Zhang, Y.; Artaxo, P.;
540 Bertram, A. K.; Manzi, A. O.; Rizzo, L. V.; et al. Sub-micrometre particulate matter
541 is primarily in liquid form over Amazon rainforest. *Nat. Geosci.* **2015**, *9* (1), 34–37.

542 (13) Song, M.; Marcolli, C.; Krieger, U. K.; Zuend, A.; Peter, T. Liquid-liquid phase
543 separation in aerosol particles: Dependence on O:C, organic functionalities, and
544 compositional complexity. *Geophys. Res. Lett.* **2012**, *39* (19).

545 (14) You, Y.; Smith, M. L.; Song, M.; Martin, S. T.; Bertram, A. K. Liquid-liquid phase
546 separation in atmospherically relevant particles consisting of organic species and
547 inorganic salts. *Int. Rev. Phys. Chem.* **2014**, *33* (1), 43–77.

548 (15) Stewart, D. J.; Cai, C.; Nayler, J.; Preston, T. C.; Reid, J. P.; Krieger, U. K.; Marcolli,
549 C.; Zhang, Y.-H. H. Liquid–Liquid Phase Separation in Mixed Organic/Inorganic
550 Single Aqueous Aerosol Droplets. *J. Phys. Chem. A* **2015**, *119* (18), 4177–4190.

551 (16) Schill, G. P.; Tolbert, M. A. Heterogeneous ice nucleation on phase-separated
552 organic-sulfate particles: Effect of liquid vs. glassy coatings. *Atmos. Chem. Phys.*
553 **2013**, *13* (9), 4681–4695.

554 (17) Renbaum-Wolff, L.; Song, M.; Marcolli, C.; Zhang, Y.; Liu, P. F.; Grayson, J. W.;

555 Geiger, F. M.; Martin, S. T.; Bertram, A. K. Observations and implications of
556 liquid-liquid phase separation at high relative humidities in secondary organic
557 material produced by *α*-pinene ozonolysis without inorganic salts. *Atmos. Chem.*
558 *Phys.* **2016**, *16* (12), 7969–7979.

559 (18) You, Y.; Bertram, A. K. Effects of molecular weight and temperature on liquid-
560 liquid phase separation in particles containing organic species and inorganic salts.
561 *Atmos. Chem. Phys.* **2015**, *15* (3), 1351–1365.

562 (19) Song, M.; Marcolli, C.; Krieger, U. K.; Lienhard, D. M.; Peter, T. Morphologies of
563 mixed organic/inorganic/aqueous aerosol droplets. *Faraday Discuss.* **2013**, *165*,
564 289–316.

565 (20) Kwamena, N.-O. A.; Buajarern, J.; Reid, J. P. Equilibrium morphology of mixed
566 organic/inorganic/aqueous aerosol droplets: investigating the effect of relative
567 humidity and surfactants. *J. Phys. Chem. A* **2010**, *114* (18), 5787–5795.

568 (21) Metcalf, A. R.; Boyer, H. C.; Dutcher, C. S. Interfacial Tensions of Aged Organic
569 Aerosol Particle Mimics Using a Biphasic Microfluidic Platform. *Environ. Sci.*
570 *Technol.* **2016**, *50* (3), 1251–1259.

571 (22) Buajarern, J.; Mitchem, L.; Reid, J. P. Characterizing multiphase
572 Organic/Inorganic/Aqueous aerosol droplets. *J. Phys. Chem. A* **2007**, *111* (37),
573 9054–9061.

574 (23) O'Brien, R. E.; Wang, B.; Kelly, S. T.; Lundt, N.; You, Y.; Bertram, A. K.; Leone, S.
575 R.; Laskin, A.; Gilles, M. K.; O'Brien, R. E.; et al. Liquid-Liquid Phase Separation
576 in Aerosol Particles: Imaging at the Nanometer Scale. *Environ. Sci. Technol.* **2015**,
577 *49* (8), 4995–5002.

578 (24) Reid, J. P.; Dennis-Smith, B. J.; Kwamena, N.-O. A.; Miles, R. E. H.; Hanford, K.
579 L.; Homer, C. J. The morphology of aerosol particles consisting of hydrophobic
580 and hydrophilic phases: hydrocarbons, alcohols and fatty acids as the
581 hydrophobic component. *Phys. Chem. Chem. Phys.* **2011**, *13* (34), 15559–15572.

582 (25) Smith, M. L.; You, Y.; Kuwata, M.; Bertram, A. K.; Martin, S. T. Phase transitions
583 and phase miscibility of mixed particles of ammonium sulfate, toluene-derived
584 secondary organic material, and water. *J. Phys. Chem. A* **2013**, *117* (36), 8895–8906.

585 (26) You, Y.; Renbaum-Wolff, L.; Carreras-Sospedra, M.; Hanna, S. J.; Hiranuma, N.;
586 Kamal, S.; Smith, M. L.; Zhang, X.; Weber, R. J.; Shilling, J. E.; et al. Images reveal
587 that atmospheric particles can undergo liquid-liquid phase separations. *Proc. Natl.*
588 *Acad. Sci. U. S. A.* **2012**, *109* (33), 13188–13193.

589 (27) Altaf, M. B.; Freedman, M. A. Effect of Drying Rate on Aerosol Particle

590 Morphology. *J. Phys. Chem. Lett.* **2017**, *8* (15), 3613–3618.

591 (28) Veghte, D. P.; Altaf, M. B.; Freedman, M. A. Size dependence of the structure of
592 organic aerosol. *J. Am. Chem. Soc.* **2013**, *135* (43), 16046–16049.

593 (29) Altaf, M. B.; Zuend, A.; Freedman, M. A. Role of nucleation mechanism on the
594 size dependent morphology of organic aerosol. *Chem. Commun.* **2016**, *52* (59),
595 9220–9223.

596 (30) Cai, C.; Tan, S.; Chen, H.; Ma, J.; Wang, Y.; Reid, J. P.; Zhang, Y. Slow water
597 transport in MgSO₄ aerosol droplets at gel-forming relative humidities. *Phys.*
598 *Chem. Chem. Phys.* **2015**, *17* (44), 29753–29763.

599 (31) Meresman, H.; Hudson, A. J.; Reid, J. P. Spectroscopic characterization of aqueous
600 microdroplets containing inorganic salts. *Analyst* **2011**, *136* (17), 3487–3495.

601 (32) Gorkowski, K.; Beydoun, H.; Aboff, M.; Walker, J. S.; Reid, J. P.; Sullivan, R. C.
602 Advanced aerosol optical tweezers chamber design to facilitate phase-separation
603 and equilibration timescale experiments on complex droplets. *Aerosol Sci. Technol.*
604 **2016**, *50* (12), 1327–1341.

605 (33) Abbatt, J. P. D.; Lee, A. K. Y.; Thornton, J. A. Quantifying trace gas uptake to
606 tropospheric aerosol: recent advances and remaining challenges. *Chem. Soc. Rev.*
607 **2012**, *41* (19), 6555–6581.

608 (34) Alexander, B.; Hastings, M. G.; Allman, D. J.; Dachs, J.; Thornton, J. A.; Kunasek,
609 S. A. Quantifying atmospheric nitrate formation pathways based on a global
610 model of the oxygen isotopic composition (D¹⁷O) of atmospheric nitrate. *Atmos.*
611 *Chem. Phys.* **2009**, *9* (14), 5043–5056.

612 (35) Gaston, C. J.; Thornton, J. A. Reacto-Diffusive Length of N₂O₅ in Aqueous
613 Sulfate- and Chloride-Containing Aerosol Particles. *J. Phys. Chem. A* **2016**, *120* (7),
614 1039–1045.

615 (36) Park, S.-C.; Burden, D. K.; Nathanson, G. M. The Inhibition of N₂O₅ Hydrolysis
616 in Sulfuric Acid by 1-Butanol and 1-Hexanol Surfactant Coatings. *J. Phys. Chem. A*
617 **2007**, *111* (15), 2921–2929.

618 (37) Surratt, J. D.; Kroll, J. H.; Kleindienst, T. E.; Edney, E. O.; Claeys, M.; Sorooshian,
619 A.; Ng, N. L.; Offenberg, J. H.; Lewandowski, M.; Jaoui, M.; et al. Evidence for
620 Organosulfates in Secondary Organic Aerosol. *Environ. Sci. Technol.* **2007**, *41* (2),
621 517–527.

622 (38) Gaston, C. J.; Riedel, T. P.; Zhang, Z.; Gold, A.; Surratt, J. D.; Thornton, J. A.
623 Reactive uptake of an isoprene-derived epoxydiol to submicron aerosol particles.

624 625 (39) Ye, J.; Gordon, C. A.; Chan, A. W. H. Enhancement in Secondary Organic Aerosol
626 Formation in the Presence of Preexisting Organic Particle. *Environ. Sci. Technol.*
627 **2016**, *50* (7), 3572–3579.

628 (40) Song, C.; Zaveri, R. A.; Shilling, J. E.; Alexander, M. L.; Newburn, M. Effect of
629 hydrophilic organic seed aerosols on secondary organic aerosol formation from
630 ozonolysis of α -pinene. *Environ. Sci. Technol.* **2011**, *45* (17), 7323–7329.

631 (41) Gordon, C. A.; Ye, J.; Chan, A. W. H. Secondary Organic Aerosol Formation
632 Enhanced by Organic Seeds of Similar Polarity at Atmospherically Relative
633 Humidity. *STEM Fellowsh. J.* **2015**, *1* (2), 6–10.

634 (42) Lu, J. W.; Rickards, A. M. J.; Walker, J. S.; Knox, K. J.; Miles, R. E. H.; Reid, J. P.;
635 Signorell, R. Timescales of water transport in viscous aerosol: measurements on
636 sub-micron particles and dependence on conditioning history. *Phys. Chem. Chem.
637 Phys.* **2014**, *16* (21), 9819–9830.

638 (43) Davies, J. F.; Haddrell, A. E.; Miles, R. E. H.; Bull, C.; Reid, J. P. Bulk, Surface and
639 Gas-Phase Limited Water Transport in Aerosol. *J. Phys. Chem. A* **2012**, *116* (45),
640 10987–10998.

641 (44) Vaden, T. D.; Imre, D.; Beránek, J.; Shrivastava, M.; Zelenyuk, A. Evaporation
642 kinetics and phase of laboratory and ambient secondary organic aerosol. *Proc.
643 Natl. Acad. Sci. U. S. A.* **2011**, *108* (6), 2190–2195.

644 (45) Abramson, E.; Imre, D.; Beránek, J.; Wilson, J.; Zelenyuk, A. Experimental
645 determination of chemical diffusion within secondary organic aerosol particles.
646 *Phys. Chem. Chem. Phys.* **2013**, *15* (8), 2983–2991.

647 (46) Liu, P.; Li, Y. J.; Wang, Y.; Gilles, M. K.; Zaveri, R. A.; Bertram, A. K.; Martin, S. T.
648 Lability of secondary organic particulate matter. *Proc. Natl. Acad. Sci.* **2016**, *113*
649 (45), 12643–12648.

650 (47) Dennis-Smith, B. J.; Hanford, K. L.; Kwamena, N. O. A.; Miles, R. E. H.; Reid, J.
651 P. Phase, morphology, and hygroscopicity of mixed oleic acid/sodium
652 chloride/water aerosol particles before and after ozonolysis. *J. Phys. Chem. A* **2012**,
653 *116* (24), 6159–6168.

654 (48) Cai, C.; Stewart, D. J.; Preston, T. C.; Walker, J. S.; Zhang, Y.-H.; Reid, J. P. A new
655 approach to determine vapour pressures and hygroscopicities of aqueous
656 aerosols containing semi-volatile organic compounds. *Phys. Chem. Chem. Phys.*
657 **2014**, *16* (7), 3162–3172.

658 (49) Rickards, A. M. J.; Miles, R. E. H.; Davies, J. F.; Marshall, F. H.; Reid, J. P.
659 Measurements of the sensitivity of aerosol hygroscopicity and the κ parameter to
660 the O/C ratio. *J. Phys. Chem. A* **2013**, 117 (51), 14120–14131.

661 (50) Davis, R. D.; Lance, S.; Gordon, J. A.; Ushijima, S. B.; Tolbert, M. A. Contact
662 efflorescence as a pathway for crystallization of atmospherically relevant
663 particles. *Proc. Natl. Acad. Sci.* **2015**, 112 (52), 15815–15820.

664 (51) Shiraiwa, M.; Zuend, A.; Bertram, A. K.; Seinfeld, J. H. Gas-particle partitioning of
665 atmospheric aerosols: interplay of physical state, non-ideal mixing and
666 morphology. *Phys. Chem. Chem. Phys.* **2013**, 15 (27), 11441–11453.

667 (52) Presto, A. A.; Donahue, N. M. Investigation of α -Pinene + Ozone Secondary
668 Organic Aerosol Formation at Low Total Aerosol Mass. *Environ. Sci. Technol.* **2006**,
669 40 (11), 3536–3543.

670 (53) Donahue, N. M.; Robinson, A. L.; Trump, E. R.; Riipinen, I.; Kroll, J. H.; McNeill,
671 V. F.; Ariya, P. A. Volatility and Aging of Atmospheric Organic Aerosol. *Atmos.*
672 *Aerosol Chem.* **2014**, 339, 97–143.

673 (54) Wills, J. B.; Knox, K. J.; Reid, J. P. Optical control and characterisation of aerosol.
674 *Chem. Phys. Lett.* **2009**, 481 (4–6), 153–165.

675 (55) Lan, J. J. Y. and M. H. and J. Y. and Y. Z. Surface whispering-gallery mode. *EPL*
676 (*Europhysics Lett.* **2011**, 96 (5), 57003.

677 (56) Preston, T. C.; Reid, J. P. Accurate and efficient determination of the radius,
678 refractive index, and dispersion of weakly absorbing spherical particle using
679 whispering gallery modes. *J. Opt. Soc. Am. B-Optical Phys.* **2013**, 30 (8), 2113–2122.

680 (57) Preston, T. C.; Reid, J. P. Determining the size and refractive index of
681 microspheres using the mode assignments from Mie resonances. *J. Opt. Soc. Am. A*
682 **2015**, 32 (11), 2210–2217.

683 (58) Song, M.; Marcolli, C.; Krieger, U. K.; Zuend, a.; Peter, T. Liquid-liquid phase
684 separation and morphology of internally mixed dicarboxylic acids/ammonium
685 sulfate/water particles. *Atmos. Chem. Phys.* **2012**, 12 (5), 2691–2712.

686 (59) Mishchenko, M. I.; Liu, L.; Mackowski, D. W. Morphology-dependent resonances
687 of spherical droplets with numerous microscopic inclusions. *Opt. Lett.* **2014**, 39 (6),
688 1701.

689 (60) Ngo, D.; Pinnick, R. G. Suppression of scattering resonances in inhomogeneous
690 microdroplets. *J. Opt. Soc. Am. A* **1994**, 11 (4), 1352.

691 (61) Voorhees, P. W. The theory of Ostwald ripening. *J. Stat. Phys.* **1985**, *38* (1-2), 231-
692 252.

693 (62) Nakayama, T.; Sato, K.; Matsumi, Y.; Imamura, T.; Yamazaki, A.; Uchiyama, A.
694 Wavelength Dependence of Refractive Index of Secondary Organic Aerosols
695 Generated during the Ozonolysis and Photooxidation of alpha-Pinene. *Sci. Online*
696 *Lett. Atmos.* **2012**, *8* (0), 119-123.

697 (63) Kim, H.; Liu, S.; Russell, L. M.; Paulson, S. E. Dependence of Real Refractive
698 Indices on O:C, H:C and Mass Fragments of Secondary Organic Aerosol
699 Generated from Ozonolysis and Photooxidation of Limonene and α -Pinene.
700 *Aerosol Sci. Technol.* **2014**, *48* (5), 498-507.

701 (64) Kim, H.; Barkey, B.; Paulson, S. E. Real refractive indices of α - And β -pinene and
702 toluene secondary organic aerosols generated from ozonolysis and photo-
703 oxidation. *J. Geophys. Res. Atmos.* **2010**, *115* (23), 1-10.

704 (65) Kim, H.; Paulson, S. E. Real refractive indices and volatility of secondary organic
705 aerosol generated from photooxidation and ozonolysis of limonene, alpha-pinene
706 and toluene. *Atmos. Chem. Phys.* **2013**, *13* (15), 7711-7723.

707 (66) Redmond, H.; Thompson, J. E. Evaluation of a quantitative structure-property
708 relationship (QSPR) for predicting mid-visible refractive index of secondary
709 organic aerosol (SOA). *Phys. Chem. Chem. Phys.* **2011**, *13* (15), 6872-6882.

710 (67) Wexler, A. S. Atmospheric aerosol models for systems including the ions H+,
711 NH4+, Na+, SO₄ 2-, NO₃, Cl-, Br-, and H₂O. *J. Geophys. Res.* **2002**, *107* (D14),
712 4207.

713 (68) Tröstl, J.; Chuang, W. K.; Gordon, H.; Heinritzi, M.; Yan, C.; Molteni, U.; Ahlm, L.;
714 Frege, C.; Bianchi, F.; Wagner, R.; et al. The role of low-volatility organic
715 compounds in initial particle growth in the atmosphere. *Nature* **2016**, *533* (7604),
716 527-531.

717 (69) Kroll, J. H.; Ng, N. L.; Murphy, S. M.; Flagan, R. C.; Seinfeld, J. H. Secondary
718 Organic Aerosol Formation from Isoprene Photooxidation. *Environ. Sci. Technol.*
719 **2006**, *40* (6), 1869-1877.

720 (70) Dommen, J.; Metzger, A.; Duplissy, J.; Kalberer, M.; Alfarra, M. R.; Gascho, A.;
721 Weingartner, E.; Prevot, A. S. H.; Verheggen, B.; Baltensperger, U. Laboratory
722 observation of oligomers in the aerosol from isoprene/NO_x photooxidation.
723 *Geophys. Res. Lett.* **2006**, *33* (13), 1-5.

724 (71) Cai, C.; Stewart, D. J.; Reid, J. P.; Zhang, Y.; Ohm, P.; Dutcher, C. S.; Clegg, S. L.
725 Organic component vapor pressures and hygroscopicities of aqueous aerosol

726 measured by optical tweezers. *J. Phys. Chem. A* **2015**, *119* (4), 704–718.

727 (72) Ray, A. K.; Davis, E. J.; Ravindran, P. Determination of ultra-low vapor pressures
728 by submicron droplet evaporation. *J. Chem. Phys.* **1979**, *71* (2), 582–587.

729 (73) Yli-Juuti, T.; Pajunoja, A.; Tikkainen, O.-P.; Buchholz, A.; Faiola, C.; Väistönen, O.;
730 Hao, L.; Kari, E.; Peräkylä, O.; Garmash, O.; et al. Factors controlling the
731 evaporation of secondary organic aerosol from alpha-pinene ozonolysis. *Geophys.
732 Res. Lett.* **2017**, *44* (5), 1–9.

733 (74) Kampf, C. J.; Waxman, E. M.; Slowik, J. G.; Dommen, J.; Pfaffenberger, L.;
734 Praplan, A. P.; Prévôt, A. S. H.; Baltensperger, U.; Hoffmann, T.; Volkamer, R.
735 Effective Henry's Law Partitioning and the Salting Constant of Glyoxal in
736 Aerosols Containing Sulfate. *Environ. Sci. Technol.* **2013**, *47* (9), 4236–4244.

737 (75) EPA, U. Estimation Programs Interface Suite™ for Microsoft® Windows. United
738 States Environmental Protection Agency 2012.

739 (76) Kroll, J. H.; Lim, C. Y.; Kessler, S. H.; Wilson, K. R. Heterogeneous Oxidation of
740 Atmospheric Organic Aerosol: Kinetics of Changes to the Amount and Oxidation
741 State of Particle-Phase Organic Carbon. *J. Phys. Chem. A* **2015**, *119* (44), 10767–
742 10783.

743 (77) Bateman, A. P.; Nizkorodov, S. a; Laskin, J.; Laskin, A. High-Resolution
744 Electrospray Ionization Mass Spectrometry Analysis of Water-Soluble Organic
745 Aerosols Collected with a Particle into Liquid Sampler. *Anal. Chem.* **2010**, *82* (19),
746 8010–8016.

747 (78) Zhang, X.; Dalleska, N. F.; Huang, D. D.; Bates, K. H.; Sorooshian, A.; Flagan, R.
748 C.; Seinfeld, J. H. Time-resolved molecular characterization of organic aerosols by
749 PILS + UPLC/ESI-Q-TOFMS. *Atmos. Environ.* **2016**, *130*, 180–189.

750 (79) Xu, L.; Guo, H.; Weber, R. J.; Ng, N. L. Chemical Characterization of Water-
751 Soluble Organic Aerosol in Contrasting Rural and Urban Environments in the
752 Southeastern United States. *Environ. Sci. Technol.* **2017**, *51* (1), 78–88.

753 (80) Stewart, D. J.; Cai, C.; Nayler, J.; Preston, T. C.; Reid, J. P.; Krieger, U. K.; Marcolli,
754 C.; Zhang, Y. H. Liquid-Liquid Phase Separation in Mixed Organic/Inorganic
755 Single Aqueous Aerosol Droplets. *J. Phys. Chem. A* **2015**, *119* (18), 4177–4190.

756 (81) Davies, J. F.; Miles, R. E. H.; Haddrell, A. E.; Reid, J. P. Influence of organic films
757 on the evaporation and condensation of water in aerosol. *Proc. Natl. Acad. Sci. U.
758 S. A.* **2013**, *110* (22), 8807–8812.

759 (82) Rothfuss, N. E.; Petters, M. D. Influence of Functional Groups on the Viscosity of

