

CRYSTALLIZATION AGES AND GEOCHEMISTRY OF THE MINERS BAY AND CEDAR BAY PLUTONS, PRINCE WILLIAM SOUND, ALASKA

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INTRODUCTION

The Late Eocene Eshamy Suite plutons (ESP) intrude turbidites of the Chugach-Prince William (CPW) terrane in western Prince William Sound. The Prince William Sound region of South-Central Alaska consists of the Chugach-Prince William Terrance (CPW), containing the Valdez and Orca Group separated by the Contact fault. The Valdez Group is Late Cretaceous in age and is characterized by a thick sequence of interbedded siltstone, greywacke, and pebble conglomerate likely deposited as turbidites on submarine fans (Tysdal and Plafker, 1978). South of the Valdez is the Orca Group of Paleocene to Eocene age, characterized by folded and faulted rocks that have been intruded by younger plutons (Davidson and Garver, 2017). The CPW has been traditionally interpreted as a Late Cretaceous to Paleocene accretionary wedge complex that was either formed in situ or was deposited farther south and subsequently transported a significant distance along the continental margin (cf. Cowan, 2003; Haeussler et al., 2003).

The ESP is a bimodal suite of granites dominated by biotite granites and leucogranites with subordinate gabbro. Questions for the ESP include the crystallization ages for the plutons, their relationship to igneous rocks found farther inboard, and the nature of the tectonic setting. For this study, we examined a set of granites from Miners Bay and Cedar Bay to compare with the rocks from the Nellie Juan and Eshamy Bay plutons to the southwest (Fig. 1), and the Caribou Creek volcanics (CCV) that occur 200 km inland. Our results show that the Miners and Cedar Bay plutons fall within the age range of the CCV and are marginally older than the Nellie Juan (NJP)

and Eshamy Bay plutons (EBP) (Cole et al., 2006; Johnson, 2012). However, preliminary geochemical data suggest that the ESP may not be directly related to the CCV and therefore may have formed in a different tectonic setting.

FIELD AND PETROGRAPHIC DESCRIPTIONS

This study focuses on the Miners Bay diorite and leucogranite which intrude the Valdez turbidites, and the Cedar Bay pluton which intrudes the Orca Group in Prince William Sound (Fig. 1). The Miners Bay diorite is truncated along its southern margin by the Miners Bay fault and the Miners Bay leucogranite occurs as a small intrusive body along the northern margin of the diorite (Fig. 1B). The diorite is black and white, equigranular hypidiomorphic containing plagioclase, clinopyroxene, and hornblende with minor biotite and ilmenite (Fig. 2A). The Miners Bay leucogranite is white, equigranular hypidiomorphic containing plagioclase, perthitic orthoclase, and quartz with trace biotite (Fig. 2B). The Cedar Bay pluton is dominated by light grey leucogranite with rare mafic enclaves and crops out south of Miners Bay (Fig. 1B). It is primarily equigranular hypidiomorphic containing plagioclase, orthoclase and quartz with trace biotite. In most samples the feldspars are partially to completely replaced by fine grained sericite and epidote (Fig. 2C).

ANALYTICAL METHODS & RESULTS

Fourteen whole-rock samples from the Miners Bay and Cedar Bay plutons including three mafic enclaves

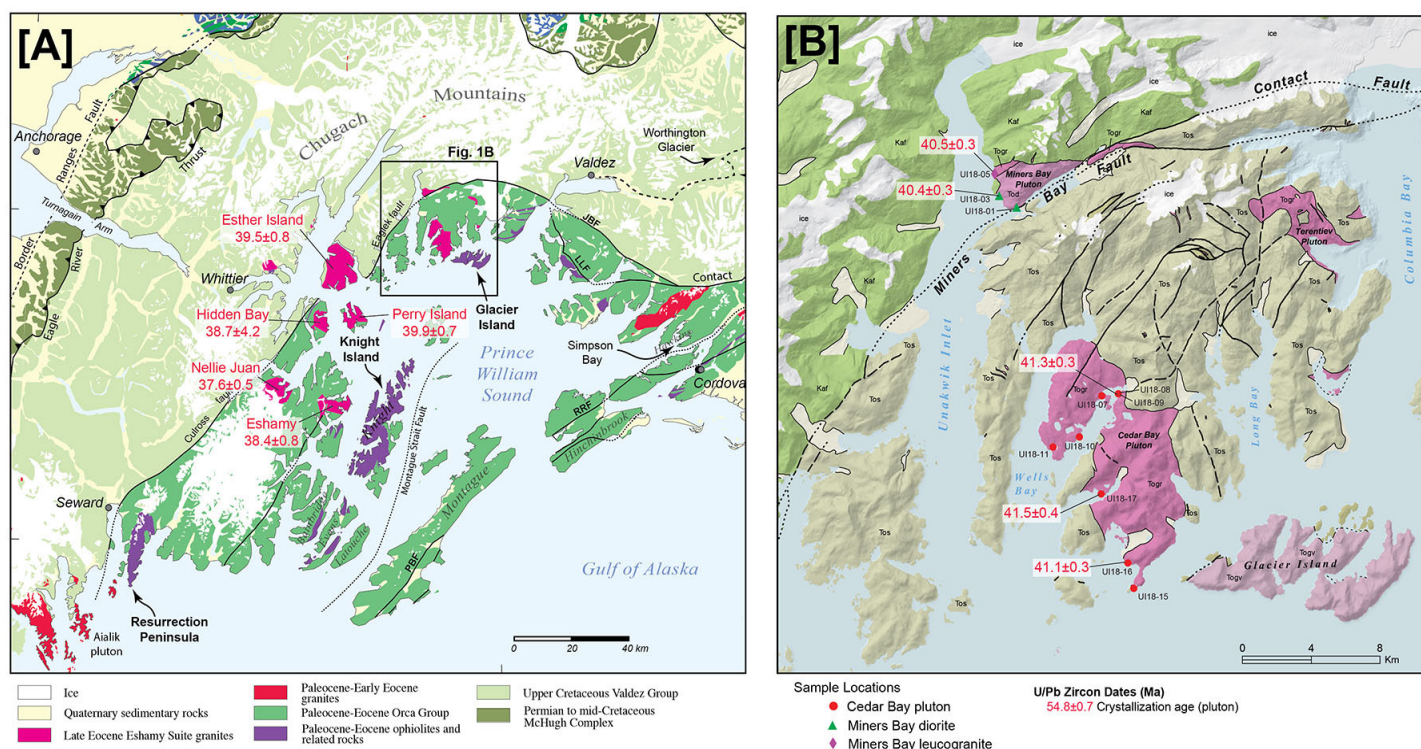


Figure 1. A) Geologic map of the CPW terrane (modified from Bradley et al., 2003). U/Pb ages of the Eshamy Suite plutons from Davidson and Garver (2017) and this study. Box shows approximate location of Fig. 1B. B) Geologic Map of northern Prince William Sound showing the Miners Bay and Cedar Bay plutons (modified from Wilson et al., 2015).

were analyzed at the Hamilton Analytical Laboratory using XRF for major elements and LA-ICP-MS for trace elements. In addition, zircon from the Miners Bay diorite (UI18-03), leucogranite (UI18-05), and from three samples from the Cedar Bay pluton (UI18-09, 16, 17) were extracted and U/Pb dated using LA-ICP-MS at the University of Arizona Laserchron Center. Approximately 35 zircons were analyzed for each sample, and BSE and CL images were used to discriminate between inherited cores (rare) and magmatic rims.

Geochemistry

The Miners Bay diorite contains ~53 wt% SiO₂ and plots as a gabbroic diorite on a plutonic TAS classification diagram (Fig. 3A). The Miners Bay leucogranite and all host-rock samples from the Cedar Bay pluton contain 72-77 wt% SiO₂ and plot in the granite field, and the three mafic enclaves contain 66-72 wt% SiO₂. All samples are calc-alkalic to calcic on the modified alkali index of Frost (2008) and are magnesian with the exception of the leucogranite and mafic enclave from the Miners Bay leucogranite which plot in the ferroan field (Fig. 3). All of the

rocks are peraluminous except for the Miners Bay diorite, the enclave from the Miners Bay leucogranite, and two samples from the Cedar Bay pluton (UI18-15&17; Fig. 3C).

Selected trace elements in the Miners Bay and Cedar Bay plutons are shown in Figure 4. REE concentrations from the Cedar Bay leucogranites and their xenoliths and the Miners Bay leucogranite are remarkably similar and are enriched in the LREE's and show a pronounced negative Eu anomaly (Fig. 4A). The xenolith from the Miners Bay leucogranite is enriched in HREE's relative to its host, has a pronounced negative Eu anomaly, and is slightly depleted in the LREE's compared to the HREE's. The Miners Bay diorite had a flatter REE profile, is depleted relative to the granites, and shows a small positive Eu anomaly relative to chondrites. All samples have low Sr/Y ratios except for the Miners Bay diorite which has low Y and high Sr/Y ratios (Fig. 4B).

U/Pb Geochronology and Hf Isotope Geochemistry

The weighted mean ages of the five samples dated

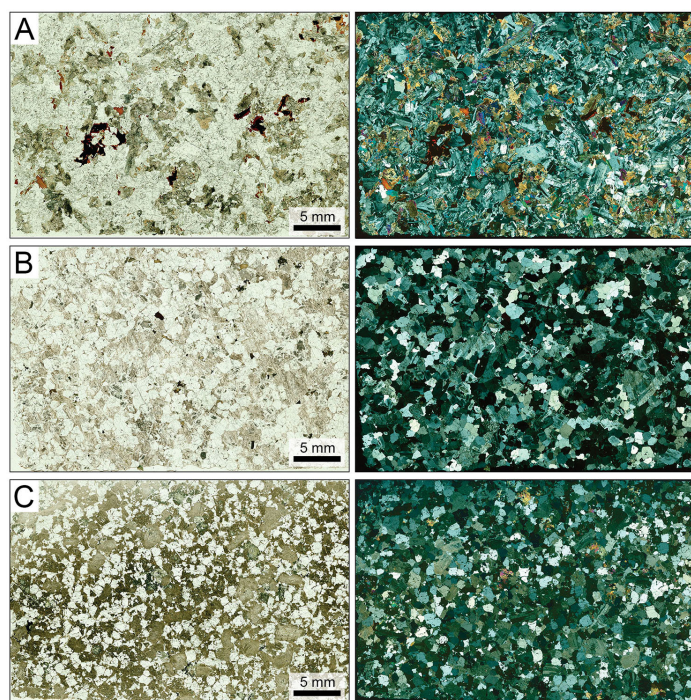


Figure 2. Photomicrographs (PPL and XPL) from the Miners Bay and Cedar Bay plutons. A) Miners Bay diorite (UI18-03). B) Miners Bay leucogranite (UI18-05A). C) Cedar Bay leucogranite (UI18-16) showing near complete replacement of feldspars with sericite and fine-grained epidote. High interference colors in XPL are coarse-grained (secondary) epidote.

in this study are shown in Table 1. The Miners Bay diorite (UI18-03) and leucogranite (UI18-05A) yielded ages of 40.41 ± 0.32 Ma and 40.45 ± 0.36 Ma, respectively. Cedar Bay pluton samples (UI18-09, 16, 17) yielded ages of 41.27 ± 0.37 Ma, 41.53 ± 0.35 Ma, and 41.06 ± 0.37 Ma, respectively. Most of the samples show no inheritance with the exception of a 61.5 Ma zircon in the Miners Bay leucogranite. Zircons from the Miners Bay diorite (UI18-03), leucogranite (UI18-05A), and Cedar Bay leucogranite (UI18-16) were also analyzed for Hf isotope ratios (Fig. 4; Table 1). Most zircons yield juvenile signatures ($\epsilon_{\text{Hf}} \sim +4$ to $+10$), but the Miners Bay leucogranite contains zircon with ϵ_{Hf} as low as -14.7 (Fig. 4).

DISCUSSION

The Miners Bay and Cedar Bay leucogranites are aluminum-rich, calc-alkalic to calcic plutonic rocks and appear to be similar to the Eshamy and Nellie Juan plutons of the Eshamy suite that crop out to the southwest (Figs. 1&3). However, the leucogranites have fewer mafic minerals and appear to be more

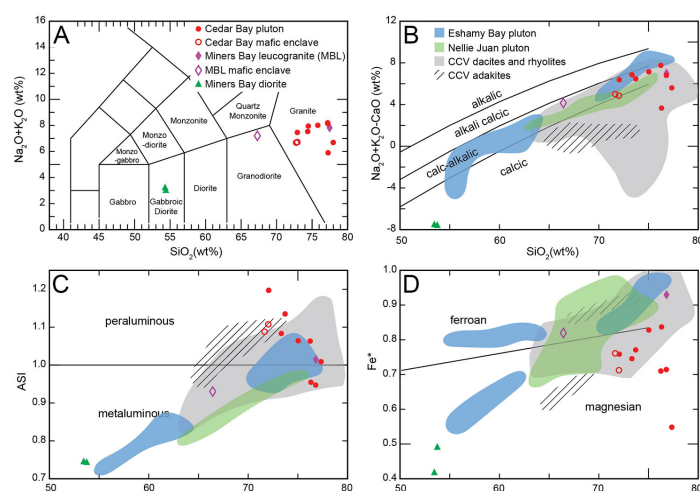


Figure 3. Major element geochemistry of the Miners Bay and Cedar Bay plutons. A) TAS classification for plutonic rocks (Middlemost, 1994). B) Modified alkali-lime index, C) aluminum-saturation index (ASI), and D) Fe index from Frost and Frost (2008). Data for the Eshamy Bay and Nellie Juan plutons from Johnson (2012) and data from the Caribou Creek volcanics (CCV) from Cole et al. (2006).

magnesian than the rest of the Eshamy suite (Fig. 3D).

The ESP are similar to the CCV in terms of major elements. While samples from the Miners Bay and Cedar Bay pluton are magnesian, most of the ESP are ferroan (Fig. 3D). Although some of the CCV plots onto the magnesian field, most of the CCV samples are ferroan aligning with the ESP. The EBP/ NJP and CCV also plot in the metaluminous field along with some of the Miners Bay and Cedar Bay pluton plotting in the peraluminous field (Fig. 3C).

The negative Eu anomaly associated with samples from the Miners Bay and Cedar Bay pluton suggests that plagioclase is fractionating from the melt, whereas a positive anomaly suggests plagioclase accumulation. The REE signatures from the dacites and rhyolites of the CCV are similar to the ESP which both share a negative Eu anomaly, but with just a flatter profile and slightly enriched HREEs (Fig. 4A). The Miners Bay diorite has a positive anomaly indicating the possibility that these are the cumulates of the Miners Bay leucogranites. The MB diorite also compares to the CCV in terms of REE's and Sr/Y ratio, indicating that these diorites may have undergone the same magmatic processes or come from a similar source region (Fig. 4B).

Previously sampled granitoid intrusions from the EBP and NJP were emplaced between 37.6 ± 0.5 and 39.9

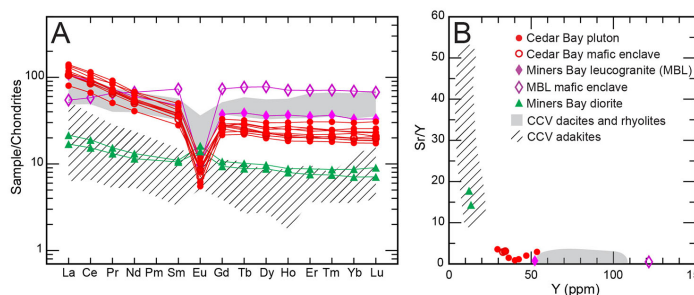


Figure 4. Trace element geochemistry of the Miners Bay and Cedar Bay plutons. A) REE's normalized to chondrites (Sun and McDonough, 1989). B) Sr-Y diagram. Data for the Caribou Creek volcanics (CCV) from Cole et al. (2006).

Table1: U/Pb and Hf isotope results of Miners Bay and Cedar Bay plutons from Prince William Sound

Sample	Location	Latitude	Longitude	U/Pb Age (Ma)	\pm	MWSD	No. of zircons analyzed	Mean ϵ_{Hf}	\pm
UI18-03 (gabbroic diorite)	Head of Miners Bay	61.0773	-147.5052	40.41	0.32	1.4	34	+8.0	0.6
UI18-05A (granite)	Miners Bay	61.0887	-147.5089	40.45	0.33	1.8	32	+4.7	0.6
UI18-09 (granite)	Cedar Bay	60.5827	-147.2340	41.27	0.33	1.9	33	N/A	N/A
UI18-16 (granite)	Cedar Bay (Fairmont Bay)	60.8876	-147.3836	41.53	0.35	0.99	30	+7.4	0.5
UI18-17 (granite)	Cedar Bay (Granite Bay)	60.9233	-147.4008	41.06	0.33	1.8	26	N/A	N/A

± 0.7 Ma (Johnson, 2012). The age range of 40.41 ± 0.32 Ma to 41.53 ± 0.35 Ma of the Miners Bay and Cedar Bay plutons confirm that they are part of the Eshamy Suite and also suggests an eastward age progression. In addition, a sandstone sample in UI18-12A in the Orca Group on Outpost Island (south of Unakwik Inlet) has an age of 40.9 ± 0.7 Ma suggesting it formed at the same time as the Eshamy Suite plutons and deposition occurred during intrusion. The CCV ranges from 52 to 36 Ma and fall within the age range of the ESP however may differ in source region and tectonic environment as they are located 200 km to the north of the ESP (Fig. 1; Cole et al., 2006). The Miners Bay and Cedar Bay plutons share a mostly juvenile Hf signature with the EBP and NJP however,

two samples from the Miners Bay pluton have a zircons with distinctly negative values suggesting that the source region for some of the plutons must involve partial melting and an incorporation of evolved crustal material (Fig 5).

CONCLUSIONS

The Miners Bay and Cedar Bay plutons are part of the Eshamy suite based on age and geochemistry. Although the CCV and ESP are similar by being ferroan and having comparable REE signatures, they differ by being primarily calcic and metaluminous with a high Sr/Y ratio in the CCV adakites. A similar tectonic environment may attribute to a possible linkage between the CCV and ESP, but it is unclear if they come from the same magmatic system and origin.

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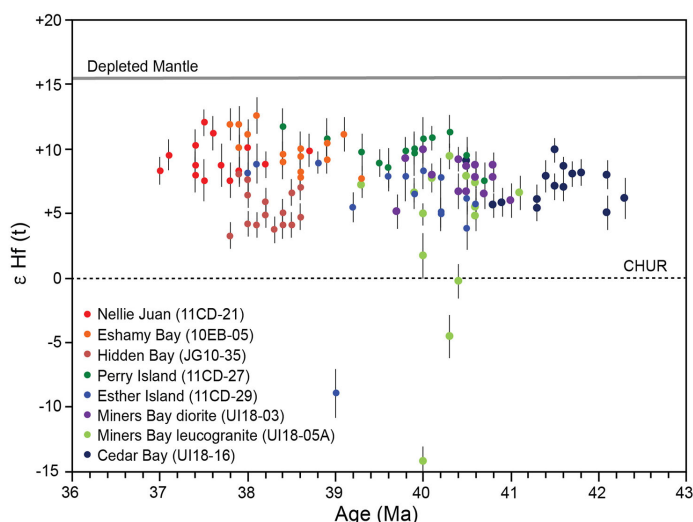


Figure 5. Zircon ϵ_{Hf} vs U/Pb age from the Eshamy Suite plutons (Fig. 1). Most zircons have a juvenile signature; however, the Esther Island pluton and Miners Bay leucogranite have negative ϵ_{Hf} values suggesting the incorporation of more evolved crustal material in the source region.

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continuity of the Valdez Group, southern Alaska, in *Changes in Stratigraphic Nomenclature* by the U.S. Geological Survey, 1977: Geological Survey Bulletin 1457-A, v. 1422-A, p. A120–A132.

REFERENCES

Cole, R.B., Nelson, S.W., Layer, P.W., and Oswald, P.J., 2006, Eocene volcanism above a depleted mantle slab window in southern Alaska: *Bulletin of the Geological Society of America*, v. 118, p. 140–158, doi:10.1130/B25658.1.

Cowan, D.S., 2003, Revisiting the Baranof-Leech River hypothesis for early Tertiary coastwise transport of the Chugach-Prince William terrane: *Earth and Planetary Science Letters*, v. 213, p. 463–475, doi:10.1016/S0012-821X(03)00300-5.

Davidson, C., and Garver, J.I., 2017, Age and Origin of the Resurrection Ophiolite and Associated Turbidites of the Chugach–Prince William Terrane, Kenai Peninsula, Alaska: *The Journal of Geology*, v. 125, p. 681–700, doi:10.1086/693926.

Haeussler, P.J., Bradley, D.C., Wells, R.E., and Miller, M.L., 2003, Life and death of the resurrection plate: Evidence for its existence and subduction in the northeastern Pacific in Paleocene-Eocene time: *Bulletin of the Geological Society of America*, v. 115, p. 867–880, doi:10.1130/0016-7606(2003)115<0867:LADOTR>2.0.CO;2.

Johnson, E., 2012, Origin of the Late Eocene Eshamy Suite granitoids in Western Prince William Sound, Alaska: 25th Annual Keck Symposium: 2012, p. 33–39.

Sun, S. S. and McDonough, W. F., 1989, Chemical and isotopic systematics of oceanic basalts; implications for mantle composition and processes (in *Magmatism in the ocean basins*), Geological Society Special Publications, vol. 42 p. 313–345.

Tysdal, R.G., and Plafker, G., 1978, Age and