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# PROVENANCE OF SANDSTONE CLASTS FROM CONGLOMERATE OF THE PALEOCENE-EOCENE ORCA GROUP IN PRINCE WILLIAM SOUND, ALASKA

MOLLIE D. POPE, Union College

Research Advisors: John Garver, Cameron Davidson

# **INTRODUCTION**

The thick flysch facies of the Cretaceous to Eocene Chugach-Prince William terrane (CPW) represents a thick accretionary complex that extends approximately 2200 km across southern Alaska, and in the central area is comprised mainly of the Valdez Group and the Orca Group (Fig. 1) (Garver and Davidson, 2015; Davidson and Garver, 2017). The Valdez Group is traditionally viewed as a Campanian to Maastrichtian turbidite deposit with mafic volcanic rocks that have experienced lower greenschist facies metamorphism (Dusel-Bacon, 1991; Gasser et al., 2012). The Orca Group is Paleocene to Eocene turbidite and volcanic deposit that, in most places, has undergone prehnite-pumpellyite facies metamorphism (Dusel-Bacon, 1991; Wilson et al., 2012).

The relationship between the Valdez Group and the Orca Group is poorly understood (Moffit, 1954). A common hypothesis suggested long ago is that they are stratigraphically related and are a continuous sequence (Capps and Johnson, 1915). Given recent zircon dating, the Valdez Group appears to have maximum depositional ages (MDA) of 75-65 Ma and the deposition of the Orca Group is between 60-50 Ma (Davidson and Garver, 2017). In this case, deformation of the Valdez Group may have occurred 65-60 Ma, just before the deposition of the oldest Orca Group turbidites began. Thus, the youngest strata of the Valdez Group must be older than the oldest strata of the Orca Group. An alternative hypothesis is that the Orca Group formed in a different location and was translated to its current position along strike slip faults after the deformation of the Valdez Group (cf. Plafker et al., 1994). This idea would mean that the

ages of the two groups may overlap in age, and the time of juxtaposition of the Orca Group to the Valdez Group is unknown but important. After the deposition of the bulk of the Orca Group was completed, the CPW experienced plutonism by the near-trench Sanak-Baranof Belt (SBB) and the Eshmay plutons (Cowan, 2003; Davidson and Garver, 2017).

If a pluton crosscuts two terranes then the age of that pluton is the minimum age that the two terranes were juxtaposed (Coney et al., 1980). The SBB plutons intruded the CPW from 63-47 Ma, with a distinct age progression from 63 Ma to the west to 50-47 Ma to the east (Davidson and Garver, 2017). In Prince William Sound the CPW terrane is also intruded by the Eshmay Suite Plutons (ESP) around 37-41 Ma (Fig. 1) (Johnson, 2012; Davidson and Garver, 2012; Garcia et al., 2019). The Eshamy suite plutons could be explained by high heat flow that melted Orca Group sediments and these melts then mixed in with mantlederived basalts (Johnson, 2012). The ESP stitch the two terranes, as they occur on both sides of the Contact Fault System (Fig. 1) (Davidson and Garver, 2017).

A key link between the Orca and Valdez Groups may be conglomerates that occur in the Orca Group. There are five main localities of conglomerates in PWS, and some of the most significant exposures are in eastern and northern PWS. These conglomerates were described by Grant and Higgins (1910) as being near the bottom of the Orca Group stratigraphy, specifically at the basal unconformity. However, Capps and Johnson (1915) described the conglomerates as being at the top of the Orca Group, occurring after and interleaved with basaltic volcanic rocks (cf. Tysdal and

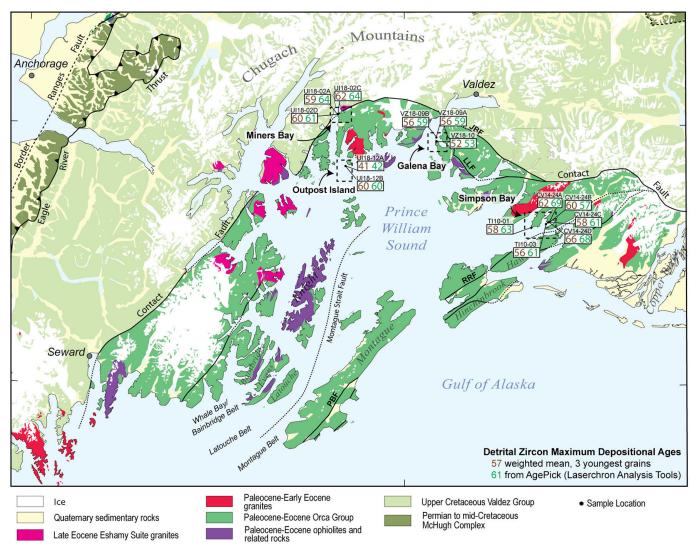


Figure 1. Map of study region in Prince William Sound. On this map the age of the the Sanak Baranof Belt and Eshamy Suite Plutons are indicated in red. Conglomerate matrix and clasts MDAs are all labeled at their sample site. The MDAs used in this paper are in brown, determined through the weighted mean on the youngest three grains (modified from Davidson and Garver 2017).

Case, 1979). If the Valdez Group is the source of the Orca Group conglomerate clasts, then the two terranes were adjacent at a time earlier than previously known (38-39 Ma) (Davidson and Garver, 2017).

Capps and Johnson (1915) proposed that the matrix of the conglomerates and the majority of the clasts were derived from the Valdez Group. They also suggest that a few clasts could be derived from the greenstones of the Orca Group. The provenance of the Orca Group conglomerates is important in our understanding of the relationship between the Valdez and Orca Groups as well as our overall understanding of the Cordilleran tectonics. This study will focus on understanding the Valdez Group and the Orca Group through the study of detrital zircons from sandstone clasts from the Orca Group Conglomerates and the host strata to those

conglomerates.

# **GEOLOGIC SETTING**

In the PWS area, the Valdez Group, which makes up about 90% of the Chugach terrane, is generally viewed as being Campanian-Maastrichtian (75-65 Ma), deformed, metasedimentary graywacke, siltstone, and shale interbedded with meta-volcanic tholeitic basalts and meta-intrusive rocks, that have undergone greenschist grade metamorphism (Dusel-Bacon, 1991; Plafker et al., 1994; Wilson and Hults, 2012). The Valdez Group is correlative with the Shumagin and Kodiak formations to the west, and parts of the Sitka Graywacke to the east, and together they extend for 2000 km along southern Alaska (Wilson and Hults, 2012). The Eagle River thrust bounds the flysch of the

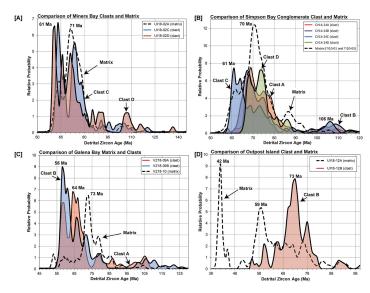


Figure 2. The probability density function (PDF) for each sample taken at the four conglomerate locations. [A] Probability density plot of Miners Bay conglomerate. The matrix UI18-02A (blacked dashed line) has a major population at 71 Ma, the clast UI18-02C (blue) has a major population at 63 Ma, and the clast UI18-02D (green) has a major population at 61 Ma. [B] PDF plot of the Simpson Bay conglomerate. The samples TI10-01 and TI10-03 (blacked dashed line) represent the matrix of the conglomerate. The clast has CV14-24C (blue) has a major population at 61 Ma and the matrix has a major age population at 70Ma. [C] PDF plot of Galena Bay conglomerate clasts VZ18-09A (red) and VZ18-09B (blue) and matrix (black dashed line). The clasts age distribution has major population at 58 and 64 Ma. The matrix lacks these populations and instead has a large population at 73 Ma. [D] Probability density plot of Outpost Island conglomerate. The matrix UI18-12A (blacked dashed line) has a major population at 40 Ma and 59 Ma, and the clast UI18-12B (red) has a major population at 73 Ma.

Chugach terrane to the north, separating it from the older mélange and the McHugh Complex and Potter Creek Assemblage (Amato et al., 2013; Davidson and Garver, 2017). The Border Ranges fault is the northern border of the Chugach terrane and it separates the terrane from the older more inland terranes of the Wrangellia composite terranes, and the southern boundary of the terrane is the Contact Fault and its equivalents (Davidson and Garver, 2017; Malik et al., 2019).

The Contact Fault system is thought to separate the Valdez Group to the North and the Orca Group to the south and it has an arcuate map trace across northern PWS (Malik et al., 2019). Across the fault, the Valdez Group and the Orca Group contain turbidites that have similar sandstone compositions, but overall the Valdez Group has undergone much more deformation (Plafker et al., 1994). In the east the Contact fault

system is primarily composed of steeply north-dipping dextral strike-slip faults, with the system in the west being very difficult to define (Bol and Gibbons, 1992; Davidson and Garver, 2017). South of the main arc of the Contact fault is the Prince William terrane, composed of the Orca Group (Davidson and Garver, 2017).

The Paleocene to Eocene Orca Group is a thick sequence of turbidities and conglomerates with ultramafic to mafic volcanic sequences, that are metamorphosed to prehnite-pumpellyite facies (Dusel-Bacon, 1991; Wilson and Hults, 2012). In PWS there are four locations of the Orca Group Conglomerates that were studied for this project: Miners Bay, Simpson Bay, Galena Bay, and Outpost Island (Fig. 1). Miners Bay is a poorly sorted, gravel to cobble (5-30 cm), clast-supported conglomerate with outsized clast (40 cm). The conglomerate contains primarily coarse-grained sandstone with some volcanic clasts. The Simpson Bay conglomerate is a clast supported, disorganized, sandstone-rich (80-90%) conglomerate with minor light-gray, felsic volcanic rocks. The conglomerate at Galena Bay is a massive, clast supported, cobble conglomerate with outsize clasts, a sandstone matrix, and no obvious bedding or grading. This conglomerate contains approximately 65% sandstone (of which 60% are veined), 15% shale, 20% volcanics, and <1% carbonate clasts. The conglomerate at Outpost Island is a cobble conglomerate with a coarse-grained sandstone matrix with approximately 80% of clast being fine-grained

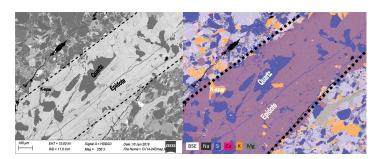


Figure 3. Mineral assemblage of veins from sample CV14-24D (MDA~66 Ma), a clast from Simpson Bay. In both images the vein boundary is defined with a black dashed line and an example of each mineral phase (Quartz, K-spar, and Epidote) is labeled. In the EDS image minerals that are orange correspond to K-rich feldspar, purple is indicative of quartz, and pink is indicative of epidote. Similar veins can be observed in Simpson Bay clast CV14-24A (MDA~62 Ma).

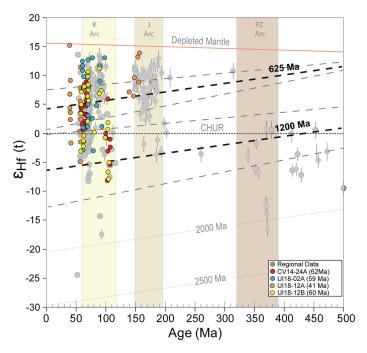


Figure 4. The  $\varepsilon$ Hf(t) for 67 new detrital zircon samples compared to other locations in southeastern Alaska (Shumagin Islands, Kodiak Island, and other parts of Prince William Sound). Extraction age of the zircons with a crystallization age of  $\sim$ 70 Ma is 625 Ma and for those with a crystallization age of  $\sim$ 100 Ma it is 1200 Ma. This figure is adapted from Davidson and Garver (2017).

sandstone, 15-20% mudstone, and less than 5% porphyry volcanic.

#### **METHODS**

Sandstone clasts and matrix were sampled from each of the four conglomerate locations in PWS. Whole rock samples were sent to GeoSep Services in Idaho or Carleton College in Minnesota to undergo a mineral separation process where detrital zircons were isolated. This process involves a large piece of the whole rock (2-4 kg) crushed to sand sized particles that are then separated by density, first on a water table and then using heavy liquids. Magnetic separation was then used to isolate zircons. U/Pb geochronology and Hf isotope analysis was conducted on detrital zircons by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) at the University of Arizona LaserChron Center. The maximum depositional age (MDA) for each sample was determined by calculating the weighted mean of the three youngest zircon ages of each sample. Thin sections of sandstone clasts were imaged using a Zeiss EVO-MA15 scanning electron microscope (SEM)

and mapped using QUANTX Electron Dispersive Spectroscopy (EDS) at Union College.

### **RESULTS**

The conglomerate at Miner's Bay (UI18-02) is a poorly sorted, gravel to cobble (5-30 cm), clast-supported conglomerate with outsized clasts (up to 40 cm). The conglomerate contains primarily coarse-grained sandstone with some volcanic clasts. The matrix of this sample (UI18-02A) yielded an MDA of  $59.4 \pm 1.1$  Ma, and the clast, UI18-02C and UI18-02D, yielded MDAs of  $61.6 \pm 1.1$  and  $59.5 \pm 1.0$  Ma with major grain-age populations at 61 and 63 Ma, respectively.

The Simpson Bay conglomerate (CV14-24) near Cordova is a clast-supported, disorganized, sandstonerich (80-90%) conglomerate with minor light-gray, felsic volcanic rocks. For this site, samples TI10-01 and TI10-3, from a location close to Simpson Bay, were used as matrix proxies (Izykowski, 2011), and were re-dated for this project. The clasts for this sample yielded MDAs of  $61.5 \pm 1.0$  (CV14-24A),  $60.3 \pm 1.3$  (CV14-24B),  $58.4 \pm 0.8$  (CV14-24C), and  $66.3 \pm 1.1$  Ma (CV14-24D), with major populations at 74 Ma, 69 Ma, 68 Ma, and 61 Ma (Fig. 2B). The matrix samples yielding a combined MDA of  $57.4 \pm 1.2$  Ma. Clasts CV14-24A and CV14-24D have clear veins with a mineral assemblage consisting of zoned epidote, quartz, and potassium-rich feldspar (Fig. 3).

The conglomerate at Galena Bay Island is a massive, clast-supported, conglomerate with a fine-grained sandstone matrix. Clasts consist of sandstone, shale, volcanic, and carbonate. At this location three sandstone clasts were collected (VZ18-09A, VZ18-09B, and VZ18-09C), however, the matrix at this location was not easily sampled, so the matrix was collected from a nearby exposure on the mainland  $\sim$ 750 meters to the SSW (VZ18-10). An MDA of 56.4  $\pm$  1.0 and 56.1  $\pm$  1.0 was determined for VZ18-09A and VZ18-09B with major age populations at 58 Ma and 64 Ma, respectively (Fig. 2C). The nearby sandstone for this site (VZ18-10) has an MDA of 52.4  $\pm$  0.9.

The conglomerate at Outpost Island is a cobble conglomerate with a coarse-grained sandstone matrix

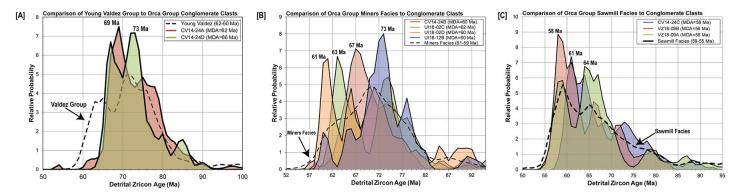


Figure 5. Probability density comparisons of sandstone conglomerate clasts to regional data. [A] Probability density plot of clasts CV14-24A (MDA ~62 Ma) and CV14-24D (MDA ~66 Ma) to the youngest Valdez Group samples (Gross Almonte et al., 2019). [B] Probability density plot of clasts CV14-24B (MDA ~60 Ma), U118-02C (MDA ~62 Ma), U118-02D (MDA ~60 Ma), and U118-12B (MDA ~60 Ma) to the Miners Facies of the Orca Group (Fisher et al., 2019). [C] Probability density plot of clasts CV14-24C (MDA ~58 Ma), VZ18-09A (MDA ~56 Ma), and VZ18-09B (MDA ~56 Ma) to the Sawmill Facies of the Orca Group (Fisher et al., 2019).

dominated by volcanic and less common sandstone clasts. In this regard this unit is distinct and different from other conglomerates because it has an abundance of volcanic clasts. The matrix of this conglomerate (UI18-12A) yielded an MDA of  $40.9 \pm 0.7$  Ma, and the clast (UI18-12B) yielded an age of  $60.1 \pm 0.8$  Ma with a major grain age population of 73 Ma (Fig. 2D).

The data for this project includes Hf isotope values for 193 detrital zircon throughout PWS. Hafnium isotope measurements when combined with zircon crystallization ages can be used to understand crustal evolution. Measured  $^{176}$ Hf/ $^{177}$ Hf values are normalized to the chondritic uniform reservoir (CHUR) and reported in epsilon notation [ $\epsilon$ Hf (t)] (Garver and Davidson, 2015). The  $\epsilon$ Hf (t) values for this data set range from -8.3 to 15.1 with two exceptions that have values of -14.3 and -24.4 (Fig. 4).

#### **DISCUSSION**

The MDAs from this project provide insight into the position of the conglomerates in the stratigraphic column of the Orca Group (~ 61 to 35 Ma) (Davidson and Garver, 2017). The MDAs of the Miners Bay and Simpson Bay conglomerates fall within the earliest time frame of the Orca Group (57 to 60 Ma), which would suggest that they sit at or near the base of the Orca Group (at ~60 Ma; Fisher et al., 2019). Recent dates of the Valdez Group indicate that it is as young as ~61 Ma meaning that there is no gap in age between the deposition of the Valdez Group and Orca Group (Fisher et al., 2019; Gross Almonte et al., 2019). The conglomerate at Galena Bay has an MDA of 52

Ma, this falls between the middle and upper part of the Orca Group. The matrix of the conglomerate at Outpost Island (UI18-12A) has an MDA of 41 Ma, which corresponds to the youngest of the Orca Group and thus sits at the upper part of the stratigraphic column.

The ages of the Galena Bay and Outpost Island conglomerate correspond to the emplacement of plutons in this area. The nearby Sheep Bay Pluton of the SBB was emplaced at ~54 Ma, so the deposition of the Galena Bay conglomerate (~52 Ma) occurred just after the emplacement of the pluton (Davidson and Garver, 2017). The Eshamy Suite plutons were emplaced from ~38 to 41 Ma, so The MDA of Outpost Island (~41 Ma) overlaps with the emplacement of the Eshamy Suite plutons (~38 to 41 Ma) (Davidson and Garver, 2017). Thus, two conglomerates (Galena Bay and Outpost Island) appear to coincide, or nearly so, with pluton intrusion which raises the question of conglomerate sedimentation being a result of pluton emplacement, and perhaps disruption of the basin due to uplift and inflation.

Almost all conglomerate clasts dated yield ages that correspond to depositional ages associated with the Orca Group and probability density functions (PDF) that pass the K-S similarity test when compared to various Orca Group samples (Fig. 5) (Fisher et al., 2019). This result suggest that within a few million years of deposition the Orca Group was being uplifted and eroded. The veined sandstone clasts are mainly derived from the Orca Group cannibalization during this period. Two clasts from Simpson Bay, CV14-

24A (~62 Ma) and CV14-24D (~66 Ma), have MDAs that can be tied to the depositional age of the Valdez Group (Figure 5A). These clasts also contain distinct veins composed of epidote, quartz, and potassium-rich feldspar, indicating that it has undergone greenschist grade metamorphism (Fig. 3). The Orca Group as a whole did not experience this high of a grade of metamorphism so it is likely that these two clasts have a Valdez affinity. Thus, this result means that the Orca Group must have been adjacent to the Valdez Group by 57 Ma.

The majority of  $\epsilon$ Hf (t) data for clasts CV14-24A and UI18-12B are positive and corresponds to other regional data from the Shumagin Islands, Kodak Island, and other parts of Prince William Sound (Fig. 4). The  $\epsilon$ Hf (t) values collected for zircons with crystallization ages of  $\sim$ 70 Ma and  $\sim$ 100 Ma suggest the existence of two distinct juvenile arcs with mantle extraction ages of  $\sim$ 625 Ma and  $\sim$ 1200 Ma, respectively (Fig. 4). Overall, this suggest that during its formation, there was a large component of juvenile crust in the source region. The similarity between the clasts  $\epsilon$ Hf (t) data to the regional data set (largely Valdez and equivalents) implies that the provenance of the two clasts are consistent with a local source, such as the Valdez Group.

#### CONCLUSIONS

The new U/Pb detrital zircon data from the Orca Group Conglomerates shows that the MDAs of each conglomerate occur throughout the deposition of the Orca Group, so the conglomerates are located throughout the stratigraphic column. The MDAs and PDFs for the majority of the sandstone clasts correspond to the Orca Group, which implies that within a few years of its deposition the Orca Group was being uplifted and eroded. Two clasts from Simpson Bay have experienced greenschist metamorphism, and thus provide a link between the Orca Group and the Valdez Group. This means that the two groups must have been together by 57 Ma, much earlier than previously thought.

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