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Data report: moisture and density analysis and ultrasonic velocity, Hole U1513E, IODP Expedition 369¹

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Keywords: International Ocean Discovery Program, IODP, *JOIDES Resolution*, Expedition 369, Australia Cretaceous Climate and Tectonics, Naturaliste Plateau, Site U1513, Hole U1513E, volcanic sequence, moisture and density, MAD, *S*-wave velocity, dynamic Poisson's ratio

Abstract

In this data report, we present postcruise petrophysical measurement results for Hole U1513E. During International Ocean Discovery Program (IODP) Expedition 369, five holes were drilled at Site U1513 on the Naturaliste Plateau offshore southwest Australia. The last and deepest hole, U1513E, recovered a volcanic sequence (Lithostratigraphic Unit VI) consisting of basalt flows, dolerite dikes, and volcanoclastic beds. Because of time constraints, moisture and density (MAD) measurements were not possible on board for Hole U1513E. To obtain bulk (wet), dry, and grain density and porosity data, we performed the MAD analysis on 25 core samples collected from Hole U1513E after the expedition. Among these samples, five were selected to measure ultrasonic velocity and dynamic Poisson's ratio. Six additional samples from Hole U1513D were analyzed to compare with shipboard data to validate the postcruise measurements. The results are compatible with shipboard data in individual lithologic units. Samples of relatively fresh rocks show bulk and dry density values near 2.5 g/cm³ and porosity near 10%, whereas altered basalts and volcanoclastics exhibit lower values of bulk and dry density and higher values of porosity. Grain density varies between 2.6 and 3.3 g/cm³. *S*-wave velocity ranges from 934 to 3135 m/s, which accompanies variable dynamic Poisson's ratio between 0.1 and 0.35.

Introduction

During International Ocean Discovery Program (IODP) Expedition 369, five holes (U1513A–U1513E) were drilled at Site U1513 on the eastern flank of the Naturaliste Plateau offshore southwest Australia (Figure F1A). The site is near Deep Sea Drilling Project

(DSDP) Site 258, which was drilled during Leg 26 (Davies et al., 1974). The deepest hole, U1513E (33°47.6190'S, 112°29.1204'E), reached 770.27 m core depth below seafloor, Method A (CSF-A) (Huber et al., 2019b). Holes U1513E and U1513D penetrated through a volcanoclastic-rich sedimentary sequence (Lithostratigraphic Unit V) and recovered a volcanic sequence (Lithostratigraphic Unit VI) (Figure F1B). The volcanoclastic-rich sequence (235 m thick) consists of sandstone, siltstone, and silty claystone layers containing abundant volcanic clasts and lithic fragments (Lee et al., 2020). The volcanic sequence (cored thickness = 82.2 m) is composed of basaltic flows (Lithologic Units 1, 3, 5, 7, and 9) and volcanoclastic beds (Lithologic Units 2, 4, 6, and 8) cut by multiple dolerite dikes (Tejada et al., 2020).

Because of time constraints during Expedition 369, it was not possible to finish shipboard core descriptions and the moisture and density (MAD) analysis for Hole U1513E. After the expedition, Cores 369-U1513E-2R through 9R were reexamined at the Kochi Core Center, Japan. A modified volcanic stratigraphy based on descriptions of petrological and mineralogical characteristics and lithologic boundaries was established for the volcanic sequence at Site U1513 (Figure F1B). The data were reported and studied in Lee et al. (2020) and Tejada et al. (2020). The studies helped us to select core samples that were used for the postcruise MAD analysis to obtain bulk (wet), dry, and grain density and porosity data for Hole U1513E. Among the samples, five were selected to measure ultrasonic velocity and dynamic Poisson's ratio. Additional samples from Hole U1513D were also analyzed to compare with shipboard data to validate the postcruise measurements. In this data report, we present our postcruise petrophysical measurement results of the volcanic sequence in Hole U1513E.

¹ Lee, E.Y., and Song, I., 2021. Data report: moisture and density analysis and ultrasonic velocity, Hole U1513E, IODP Expedition 369. In Hobbs, R.W., Huber, B.T., Bogus, K.A., and the Expedition 369 Scientists, *Australia Cretaceous Climate and Tectonics*. Proceedings of the International Ocean Discovery Program, 369: College Station, TX (International Ocean Discovery Program). <https://doi.org/10.14379/iodp.proc.369.204.2021>

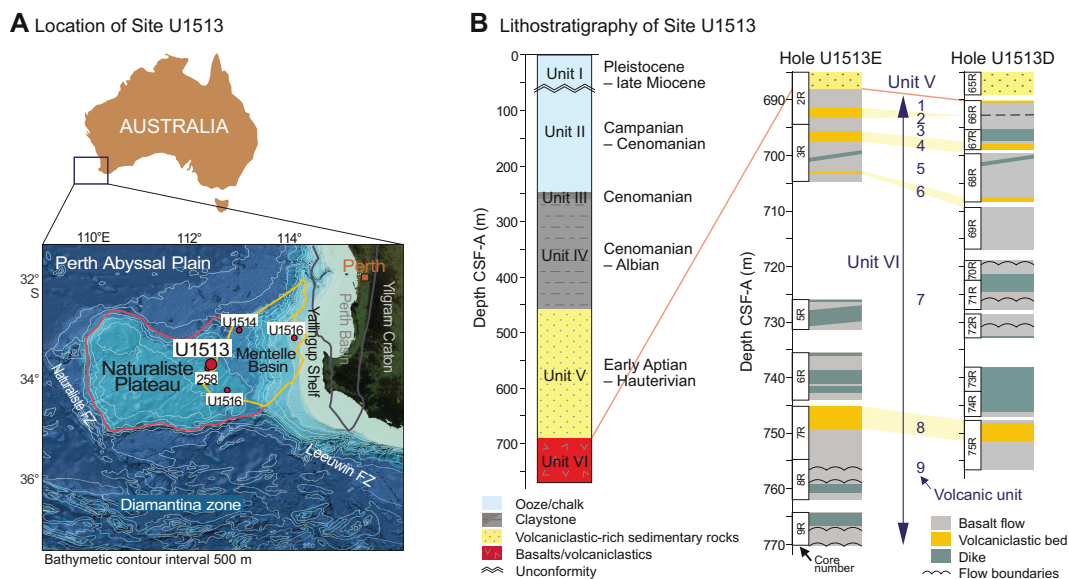
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Figure F1. A. Map of submarine Naturaliste Plateau and Mentelle Basin offshore southwest Australia (modified from Borissova, 2002; Huber et al., 2019b) showing location of Site U1513, adjacent DSDP Site 258, and IODP Sites U1514–U1516. FZ = fracture zone. B. Lithostratigraphy at Site U1513 with Lithostratigraphic Units I–VI. Lithologic (volcanic) units of Lithostratigraphic Unit VI in Holes U1513E and U1513D are shown (revised from Huber et al., 2019b; Tejada et al., 2020).



Materials and methods

Moisture and density measurements

During Expedition 369, shipboard MAD measurements for the volcanic sequence (Lithostratigraphic Unit VI) in Hole U1513D were performed on 16 core samples collected from the working-half sections of Cores 369-U1513D-66R through 75R. The standard IODP MAD procedures (ASTM International, 1990; Blum, 1997) were applied to analyze the hard rock samples from Hole U1513D (Huber et al., 2019a). The measured wet mass, dry mass, and dry volume of each sample were used to determine bulk (wet) density, dry density, grain density, and porosity. Salt precipitation in pore space during the drying process was included in the wet mass and dry volume values. The shipboard MAD data were described in Huber et al. (2019b) and are available at <https://web.iodp.tamu.edu/LORE>.

The postcruise MAD analysis was completed at the Korea Institute of Geoscience and Mineral Resources (KIGAM), Republic of Korea. MAD instrumentation, measurements, and analysis followed shipboard procedures as much as possible, except for a few minor procedural differences. For saturation, we placed the samples in individual plastic vials that were filled with distilled water (density = ~ 1.0 g/cm³) instead of seawater (density = ~ 1.024 g/cm³). Thus, the mass and volume of salt were not evaluated and are not considered for MAD calculations in this report. Because of their very low values, however, only slight differences are observed, which is supported by a good comparison between shipboard and postcruise MAD data using samples from Hole U1513D (Table T1).

The vacuum chamber (Model OV-12; Figure F2A) was used to saturate pore spaces. The vacuum pump removed air from the chamber to a pressure of ~ 50 kPa below atmospheric pressure. These samples were kept under saturation for ~ 24 h with the vacuum maintained in the chamber. After removal from the saturator, the samples were patted dry with a paper towel, and their wet masses were immediately determined using the Mettler-Toledo electronic balance (Model XS404SV; Figure F2B). The samples were

Table T1. Moisture and density analysis and data for 25 samples, Hole U1513E. [Download table in CSV format.](#)

then dried in an oven for ~ 24 h at 100° – 110° C and measured for dry mass using the electronic balance after cooling in a desiccator. Dry sample volume was determined with a five-celled helium gas pycnometer (Model Pentapyc 5200e; Figure F2C). Volume measurement was preceded by five purges of the sample chamber with helium. The data from the five measurement cycles of each sample were averaged to determine dry volume and grain density (see MAD_DATA in [Supplementary material](#)). The measured values were used to calculate bulk density, dry density, porosity, and other parameters (Table T1) following the standard IODP MAD calculations (Blum, 1997; Huber et al., 2019a).

To conduct the postcruise MAD analysis for Hole U1513E, dozens of core samples were collected from Cores 396-U1513E-2R through 9R. Among them, 25 samples were carefully selected from what were deemed suitable conditions because shipping induced fragmentation of some samples. A total of 12 samples were from dolerite dikes, 9 samples were from basalt flows, 3 samples were from volcaniclastic beds in the volcanic sequence, and 1 sample was from the volcaniclastic-rich sequence (Table T1). Six samples from Hole U1513D were also analyzed alongside to compare with shipboard data.

Ultrasonic measurements

During Expedition 369, *P*-wave velocity was determined by measuring the time required for a compressional wave to travel a specific distance through a core section. The velocity measurement was performed using Panametrics-NDT Microscan delay line transducers for selected positions on the working-half sections (Huber et al., 2019a). A total of 27 data sets of *P*-wave velocity were acquired for the volcanic sequence in Hole U1513D (Figure F3E), and 99 data sets were acquired for Hole U1513E (Figure F4E), all of which can be accessed at <https://web.iodp.tamu.edu/LORE>.

Figure F2. Instruments used for moisture and density measurements at Korea Institute of Geoscience and Mineral Resources (KIGAM). A. Vacuum chamber (Model OV-12, Jeio Tech) to saturate pore spaces of core samples. B. Electronic balance (Model XS404SV, Mettler-Toledo) to measure wet and dry mass. C. Five-celled helium gas pycnometer (Model Pentapyc 5200e, Quantachrome Instruments) to measure volume of solids.

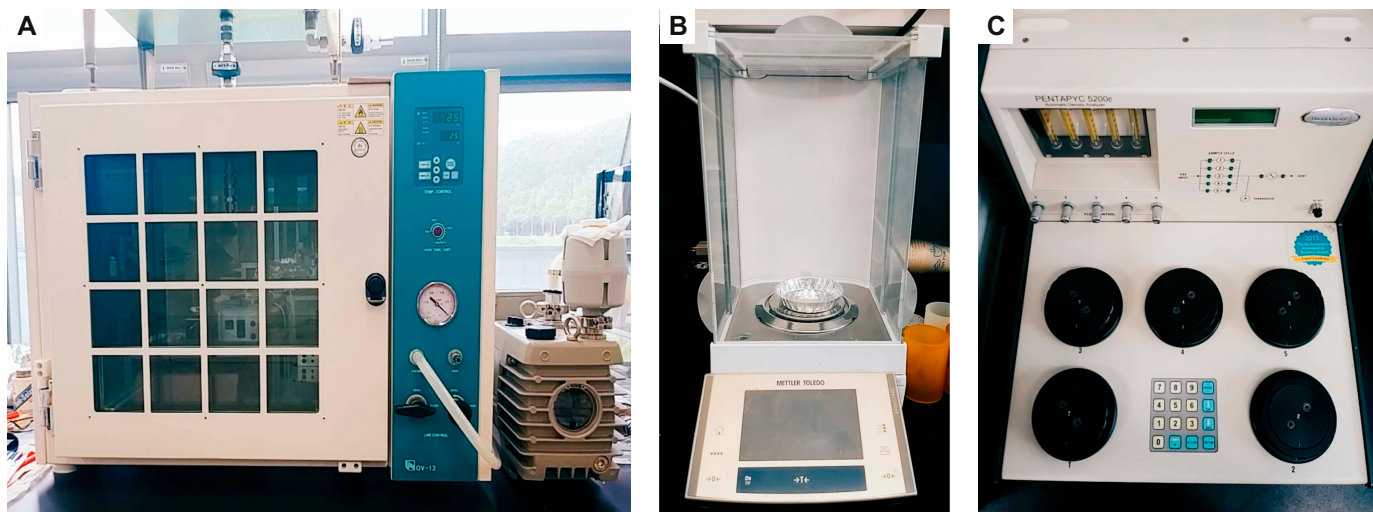
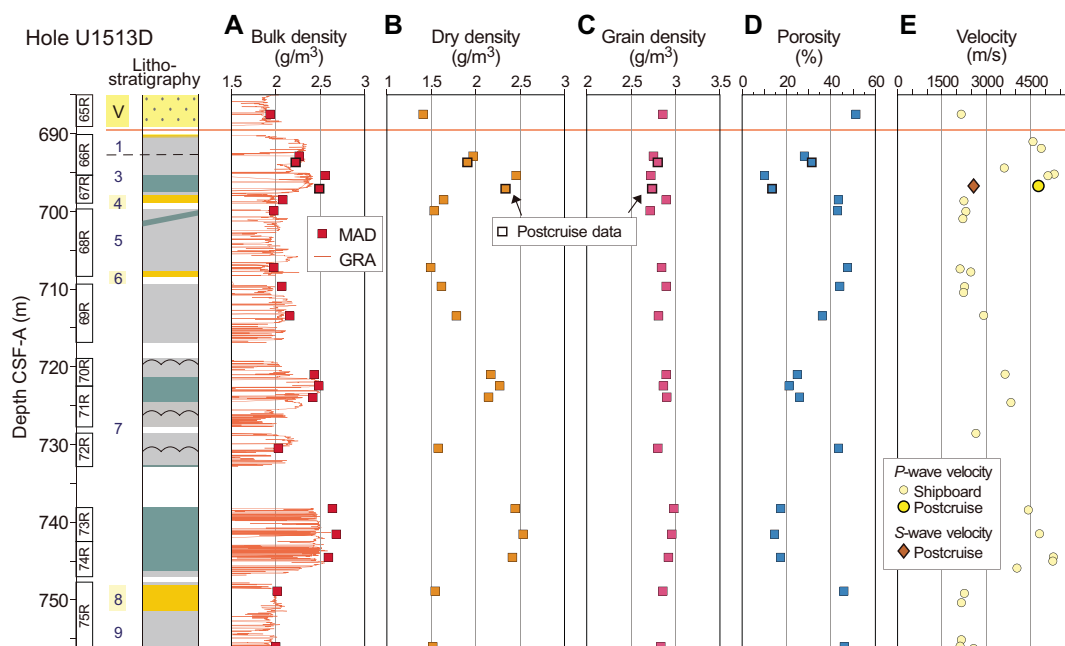


Figure F3. Density, porosity, and velocity data, Cores 369-U1513D-65R through 75R. Lithostratigraphy, lithologic unit, and core number are from Figure F1B. A. Bulk (wet) density data calculated using moisture and density (MAD) analysis on discrete samples and gamma ray attenuation (GRA) on whole-round sections. (B) Dry density, (C) grain density, and (D) porosity data calculated by MAD analysis on discrete samples. Squares with thicker outlines = postcruise data from two discrete samples. E. Shipboard *P*-wave velocity data measured on working-half sections, with postcruise *P*-wave and *S*-wave velocity data measured on a discrete sample.



The postcruise measurement of *P*-wave and *S*-wave velocity was conducted using a 5058PR high voltage ultrasonic pulser-receiver and a 12-bit 60 MHz digitizer at the KIGAM. Six core samples that were the least fractured and a suitable size were carefully selected; five samples were from Hole U1513E, and one sample was from

Hole U1513D. Four samples were from basalt flows, one sample was from a dolerite dike, and one sample was from a volcanoclastic bed. Based on *P*-wave and *S*-wave velocity, dynamic Poisson's ratio was evaluated for each sample (Table T2).

Figure F4. Density, porosity, and velocity data, Cores 369-U1513E-2R through 9R. Lithostratigraphy, lithologic unit, and core number are from Figure F1B. A. Bulk (wet) density data calculated using moisture and density (MAD) analysis on discrete samples and gamma ray attenuation (GRA) on whole-round sections. (B) Dry density, (C) grain density, and (D) porosity data calculated by MAD analysis on discrete samples. E. Shipboard *P*-wave velocity data measured on working-half sections, with postcruise *P*-wave and *S*-wave velocity data measured on five discrete samples.

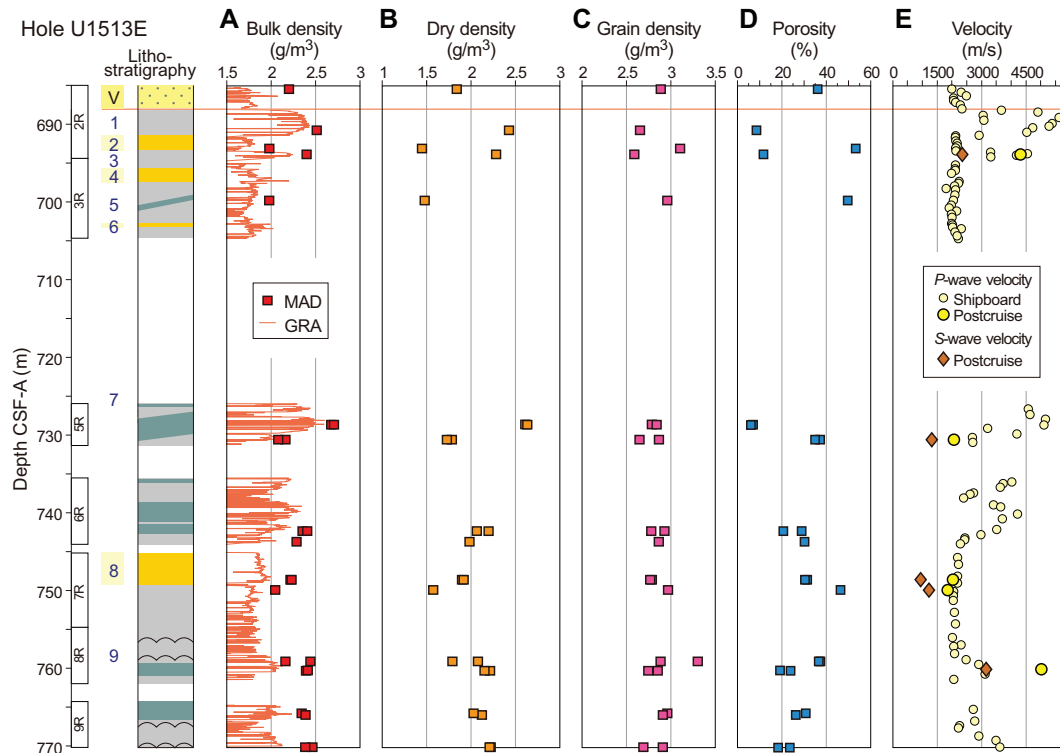


Table T2. *P*-wave and *S*-wave velocity data and dynamic Poisson's ratio for five samples from Hole U1513E and one sample from Hole U1513D. [Download table in CSV format.](#)

Results

Density and porosity

To compare postcruise data to shipboard data, six core samples from Hole U1513D were selected for MAD analysis, one of which was also measured for *P*-wave and *S*-wave velocity. The postcruise data show values similar to those from adjacent samples in individual basalt flow, dike, and sedimentary layers (Figure F3; Table T1). These comparisons indicate that the difference in measuring equipment had little impact on the results.

Bulk (wet) density for Hole U1513E ranges between 1.98 and 2.70 g/cm³. Dry density ranges between 1.45 and 2.63 g/cm³ (Figure F4A, F4B; Table T1). These ranges are similar to those from the volcanic sequence in Hole U1513D (Figure F3A, F3B). Variations in bulk and dry density values are attributed to lithologic changes and alteration degree (Tejada et al., 2020). Less altered basalt flows (e.g., Lithologic Unit 1) and dikes show relatively high bulk and dry density values near 2.5 g/cm³, whereas strongly altered basalt flows (e.g., Lithologic Unit 5) and volcanoclastic beds show lower values near 2.0 and 1.5 g/cm³, respectively. MAD bulk density values were generally higher than shipboard bulk density data from gamma ray attenuation (GRA) measurements (2.5 cm interval) on whole-round sections (Figures F3A, F4A). We assume that the difference results from a gap between the core section and liner as well as GRA measurement on fractured/fragmented sections (e.g., Planke et al., 1999).

Grain density for Hole U1513E ranges between 2.6 and 3.3 g/cm³ (Figure F4C; Table T1). Because the average grain density of fresh basalts is ~2.9 g/cm³ at other drilling sites (Broglia and Moos, 1998; Hamano, 1980; Hyndman and Drury, 1977), we infer that the grain density variation in the volcanic sequence at Site U1513 is likely due to the heterogeneous distribution of primary igneous and secondary alteration minerals (Tejada et al., 2020).

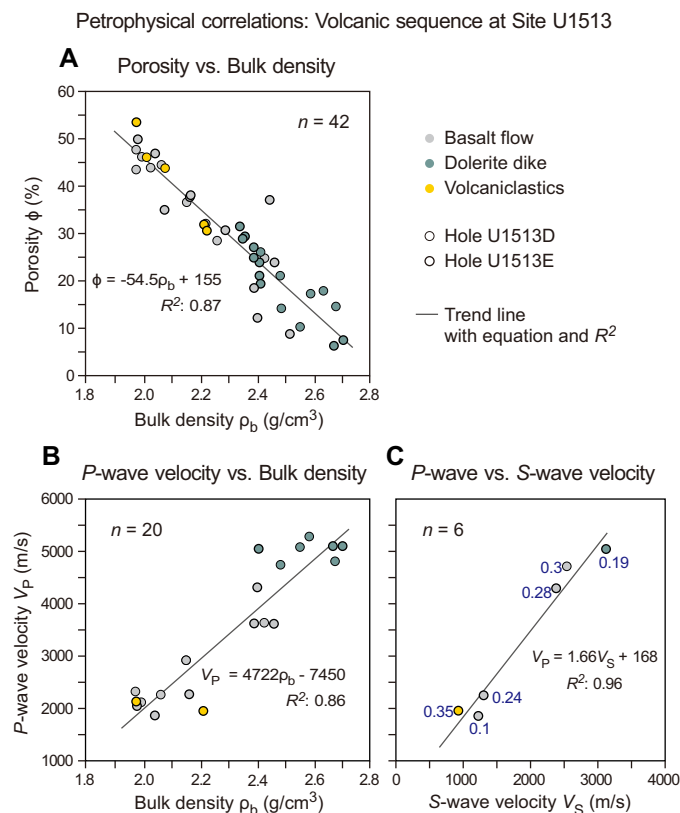
Porosity shows a wide variation in Hole U1513E, ranging from 6.3% to 53.5% (Figure F4D; Table T1). The scattered distribution is also observed in the volcanic sequence in Hole U1513D (Figure F3D). Low porosity values are observed in relatively fresh and less altered intervals of basalt flows and dolerite dikes, which correspond to fresh basalts' porosity of 5%–10% and higher in vesicular condition (Broglia and Moos, 1998; Hamano, 1980; Hyndman and Drury, 1977; Planke et al., 1999). Altered basalts and volcanoclastics show higher porosity. The increase in porosity correlates well with the bulk density decrease (Figure F5A).

One sample from Hole U1513E was collected from the volcanoclastic-rich siltstone (685.4 m CSF-A) of Lithostratigraphic Unit V, and it has 2.21 g/cm³ bulk density, 1.85 g/cm³ dry density, 2.9 g/cm³ grain density, and 36.5% porosity (Figure F4A–F4D; Table T1).

Ultrasonic velocity and dynamic Poisson's ratio

The postcruise measurements of *P*-wave velocity exhibit results similar to shipboard data (Figures F3E, F4E). Variations in *P*-wave velocity correspond closely to lithologic changes and alteration degree through the section (Tejada et al., 2020). Less-altered basalts (e.g., Lithologic Unit 1) and dikes show a high velocity of ~5000 m/s, whereas the velocities of strongly altered basalts (e.g., Lithologic

Figure F5. Petrophysical data correlations for the volcanic sequence, Site U1513. A. Porosity vs. bulk density. B. *P*-wave velocity vs. bulk density. C. *P*-wave vs. *S*-wave velocity and dynamic Poisson's ratio (dark blue). Trend lines with equation and determination coefficient (R^2) are shown. n = number of data points. Offset between positions of *P*-wave velocity and moisture and density bulk density data is ≤ 10 cm.



Unit 5) and volcaniclastics are ~ 2000 m/s. The velocity decrease is associated with the bulk density decrease (Figure F5B). *S*-wave velocity ranges from 934 to 3135 m/s (Table T2). The lowest two values of *S*-wave velocity are from strongly altered samples of basalt and volcaniclastics, whereas the highest value is from a relatively fresh dike sample. Dynamic Poisson's ratio varies between 0.1 and 0.35 without a clear relation with lithology (Figure F5C).

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