



## Article

## Megadrought and cultural exchange along the proto-silk road

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## ABSTRACT

Arid Central Asia (ACA), with its diverse landscapes of high mountains, oases, and deserts, hosted the central routes of the Silk Roads that linked trade centers from East Asia to the eastern Mediterranean. Ecological pockets and ecotones in ACA are largely determined by local precipitation. However, little research has gone into the effects of hydroclimatic changes on *trans*-Eurasian cultural exchange. Here, we reconstruct precipitation changes in ACA, covering the mid-late Holocene with a U-Th dated, ~3 a resolution, multi-proxy time series of replicated stalagmites from the southeastern Fergana Valley, Kyrgyzstan. Our data reveal a 640-a megadrought between 5820 and 5180 a BP, which likely impacted cultural development in ACA and impeded the expansion of cultural traits along oasis routes. Instead, it may have diverted the earliest transcontinental exchange along the Eurasian steppe during the 5th millennium BP. With gradually increasing precipitation after the megadrought, settlement of peoples in the oases and river valleys may have facilitated the opening of the oasis routes, “prehistoric Silk Roads”, of *trans*-Eurasian exchange. By the 4th millennium BP, this process may have reshaped cultures across the two continents, laying the foundation for the organized Silk Roads.

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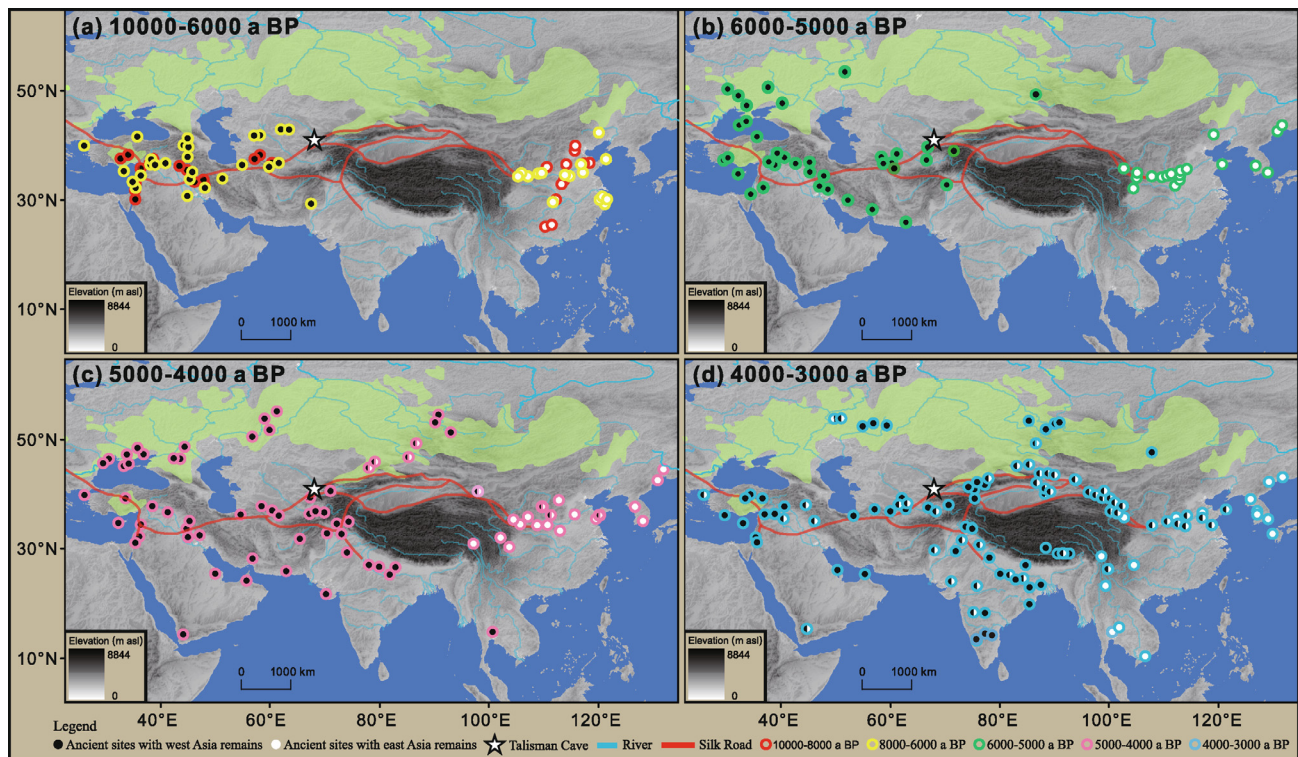
## 1. Introduction

Early agricultural populations in East and West Asia followed distinct trajectories of cultural development (Fig. 1a), leading to differing languages, socio-political systems, rituals, culinary traditions, cultivation practices, and domestication pathways for crops and animals [1–3]. Many scholars have noted the existence of a cultural barrier that roughly followed the mountains and deserts of ACA. Differences are visible in the types of crops and

the cultivation practices used on either side of this divide [4]. *Trans*-Eurasian exchange [5] led to the gradual dispersal of cultural traits across Eurasia, starting several millennia before the formation of organized and taxed commercial trade. Archeological studies indicate that wheat, barley, sheep, goats, and cattle (West Asian cultural traits) spread to East Asia during the late fifth and fourth millennium BP. Likewise, broomcorn and foxtail millet evolved under cultivation in what is now northern China and dispersed to West Asia during the fourth millennium BP [6,7] (Fig. 1c, d). Some scholars argue that the early *trans*-Eurasian exchange crossed the Eurasian Steppe from Europe to Asia, a model often referred to as the steppe highway [8]. However,

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**Fig. 1.** Maps showing the distribution of ancient sites in ACA and the Eurasian Steppe during different millennial increments over the past 10,000 a. Representative sites in other regions of Eurasia are also shown. (a) 10,000–8000 a BP (red) and 8000–6000 a BP (yellow), (b) 6000–5000 a BP, (c) 5000–4000 a BP, and (d) 4000–3000 a BP. The black dots represent the ancient sites with West Asian cultural elements, specifically wheat, barley, sheep/goat, and cattle. The white dots represent the ancient sites with East Asian cultural elements, including foxtail and broomcorn millet. The red lines represent prominent historical routes of the Silk Roads and green shaded area represents the steppe ecoregion.

increasingly more archaeobotanical, zooarchaeological, and isotopic data suggest that the main passage for the transcontinental exchange followed the rich river valleys and alluvial slopes of the mountains to the beaded oasis passage, a continual route of communication for more than 4000 a [4,9,10] (Fig. 1b, c), designating the pre-Silk Roads (Dataset S1 online). This spatial–temporal transformation significantly influenced the development of cultures across Europe and Asia [1,2,11], yet the underlying mechanisms that drove or constrained population expansions remain unclear.

Extreme climate events, especially megadroughts lasting for decades or centuries, are thought to be a natural forcings that can contribute to the collapse of ancient imperial systems [12,13]. Shifts in surface water have been implemented in the abandonment of many major urban centers in ACA, such as those along the peripheries of the Bukhara Oasis in Uzbekistan or across the northern edge of the Kopet Dag Mountains during the third millennium BP [14]. However, whether extreme climate events in ACA played an important role in prehistoric demographic shifts and the trans-Eurasian exchange remains unexplored, due to a dearth of high-resolution climate records with accurate chronologies covering the middle to late Holocene in this region. Existing paleoclimate records reveal Holocene climate trends in eastern ACA [15–17]. These records have issues with dating uncertainty, temporal resolution, and proxy interpretation [17]. To address these uncertainties, we have produced a climate record from two stalagmites collected from Talisman Cave in Kyrgyzstan. This high-resolution (~3 a), precisely  $^{230}\text{Th}$  dated (dating errors are ~6‰) record allows us to correlate regional climate changes with the archaeological record, and test potential cause-effect hypotheses [18].

## 2. Materials and methods

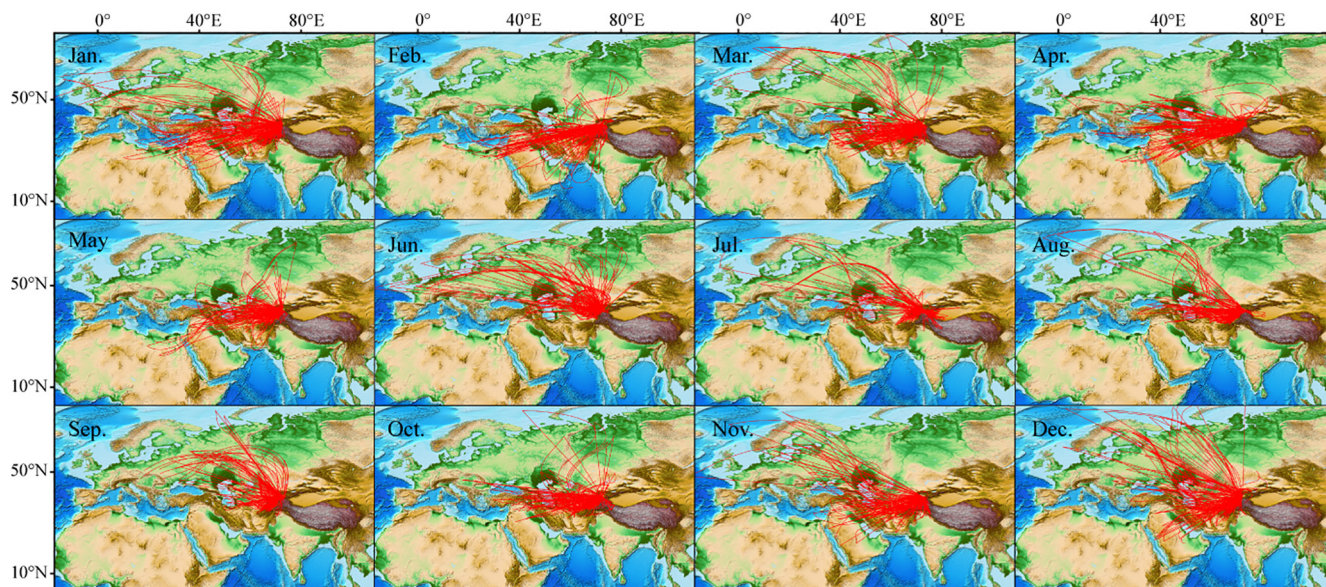
### 2.1. Cave and stalagmites

Talisman Cave is located in the southeastern Fergana Valley in ACA (40.39°N, 72.35°E, 1486 m a.s.l.), near the crossroads of the historical Silk Roads (Figs. 1 and S1 online). Mean annual precipitation in this region is ~300 mm, with more than 70% occurring during winter and spring (Figs. S2 and S3 online). The westerlies bring moisture from the Mediterranean, Black Sea, and Caspian Basin, as well as from the North Atlantic [19] (Figs. 2 and S4 online). These moisture sources also feed high-elevation snowfall and summer glacial-melt streams, which support agriculture and grassland farming in the region. Two columnar-shaped calcite stalagmites, F2 and F11, with lengths of 15.3 and 20 cm, respectively, were collected from the northern end of the cave chamber. When halved and polished, both stalagmites show clear growth layers (Fig. 3). A hiatus was determined in 114.5 mm from the top of the stalagmite F2.

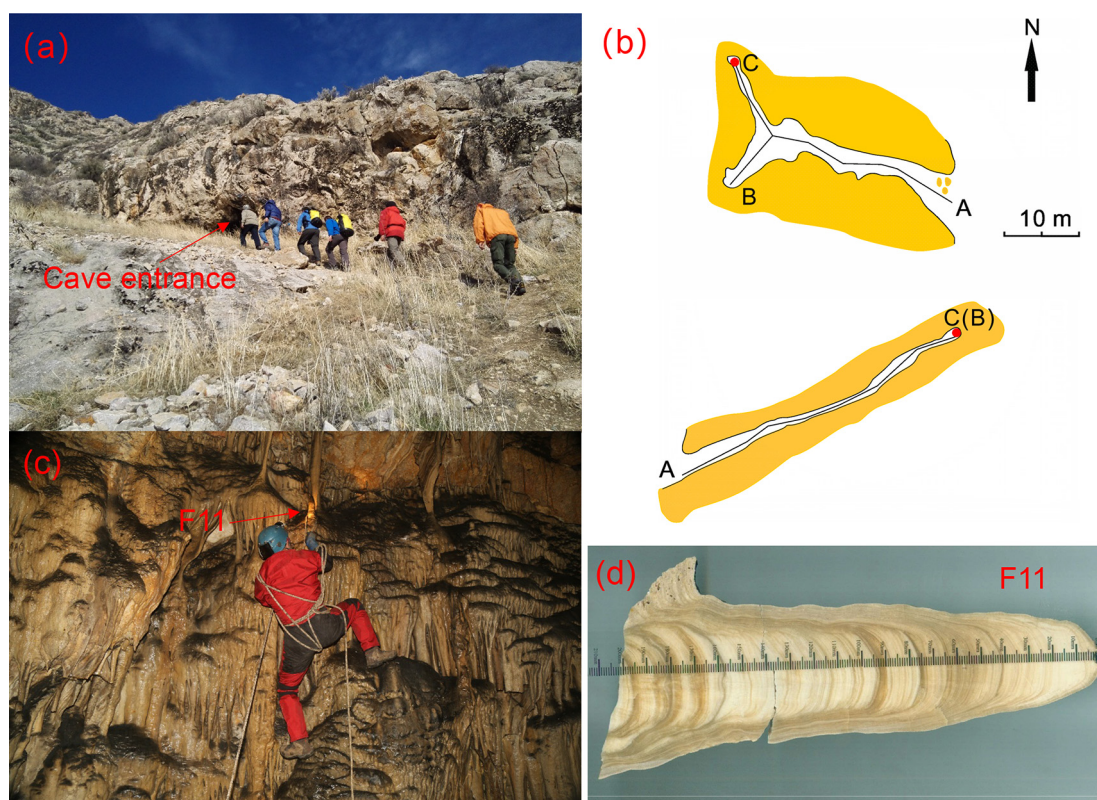
### 2.2. Methods

Subsamples of 50–100 mg of 33 (F11) and 16 (F2) layers were drilled along the growth axis on the polished surface for U–Th dating. We followed the chemical procedure described by Edwards et al. [21] and Shen et al. [22] to separate uranium and thorium. U–Th isotopic composition and  $^{230}\text{Th}$  dates were determined by a multi-collector inductively coupled plasma mass spectrometer (MC-ICPMS), Thermo Fisher Neptune, at the Isotope Laboratory, Xi'an Jiaotong University [23]. Age models were established by





**Fig. 2.** HYSPLIT [20] backward trajectory analyses for moisture of Talisman Cave during different months of 2015 (data from <http://ready.arl.noaa.gov/HYSPLIT.php>). For each month, we computed ~120 isentropic back-trajectories (168 h) (sampled four times daily at UTC 00, 06, 12 and 18) initialized at 2000 m (above ground level).



**Fig. 3.** Outlook of Talisman Cave (a) and plane and vertical views of Talisman Cave (b), with A denoting the entrance, B and C denoting the bottoms of two chambers, respectively. (c) and (d) are polished sections of F11 and F2, respectively. Both stalagmites were collected from a chamber at the bottom of the cave (location C in (b)).

using 5000 Monte-Carlo simulations and a polynomial interpolation procedure in the COPRA routine [24].

Stalagmite subsamples for stable isotopes (oxygen and carbon) analyses were contiguously drilled at intervals of 0.1 and 0.15 mm for F11 and F2, respectively, by using a Micromill device. Subsamples were analyzed on an IsoPrime100 gas source stable isotope ratio mass spectrometer equipped with a MultiPrep system at

the Institute of Earth Environment, Chinese Academy of Sciences (IEECAS), Xi'an, China. Reported  $\delta^{18}\text{O}$  and  $\delta^{13}\text{C}$  values were calculated with respect to the Vienna Pee Dee Belemnite (VPDB). An international standard NBS 19 was analyzed every 10–15 samples to monitor instrumentation and reproducibility. The replicates showed that the external error for  $\delta^{18}\text{O}$  and  $\delta^{13}\text{C}$  are better than 0.1‰ ( $2\sigma$ ), respectively.

Sr and Ca measurements were scanned along the growth axis of the ethanol-cleaned surface of F11. The measurements were carried out using the 4th generation Avaatech X-ray fluorescence (XRF) core scanner equipped with the latest variable optical XRF technology at the IEECAS [25]. The analytical settings were as follows: 30 kV and 0.03 mA with a Pd-tube, 10 s scanning time at 0.1 mm resolution, and with an irradiated area of 4 mm<sup>2</sup>. Results are reported in counts per second (cps). A total of 1457 Sr and Ca measurements were scanned, respectively.

### 3. Results

<sup>230</sup>Th dating results (Table S1 online and Fig. 4) indicate that F11 continuously deposited calcite from 7774 to 656 a BP. F2 formed between 4943 and 211 a BP, with a 316-a hiatus between 3628 and 3312 a BP. This paper focuses on the prehistoric period between 7774 and 3000 a BP. F11's average dating uncertainty is ~6‰, and average temporal resolution of proxies is ~3 a, making it one of the most precisely dated and highest-resolved mid-Holocene records from ACA.

The F11  $\delta^{13}\text{C}$  and  $\delta^{18}\text{O}$  values range from −8.36‰ to −1.16‰ and −14.99‰ to −9.73‰, respectively. Significant positive correla-

tions are observed between  $\delta^{13}\text{C}$ ,  $\delta^{18}\text{O}$ , and Sr/Ca timeseries during the period from 7774 to 3000 a BP (Dataset S2 and Table S2 online). “Hendy test” results for the F11 subsamples show no evidence of large kinetic disequilibrium effects on the  $\delta^{18}\text{O}$  and  $\delta^{13}\text{C}$  values of the stalagmite (Fig. S5 online). Broad similarities between the  $\delta^{18}\text{O}$  profiles of F11 and F2, within the margin of dating errors, during the overlapping period are consistent with deposition at or near isotopic equilibrium (Fig. S6 online).

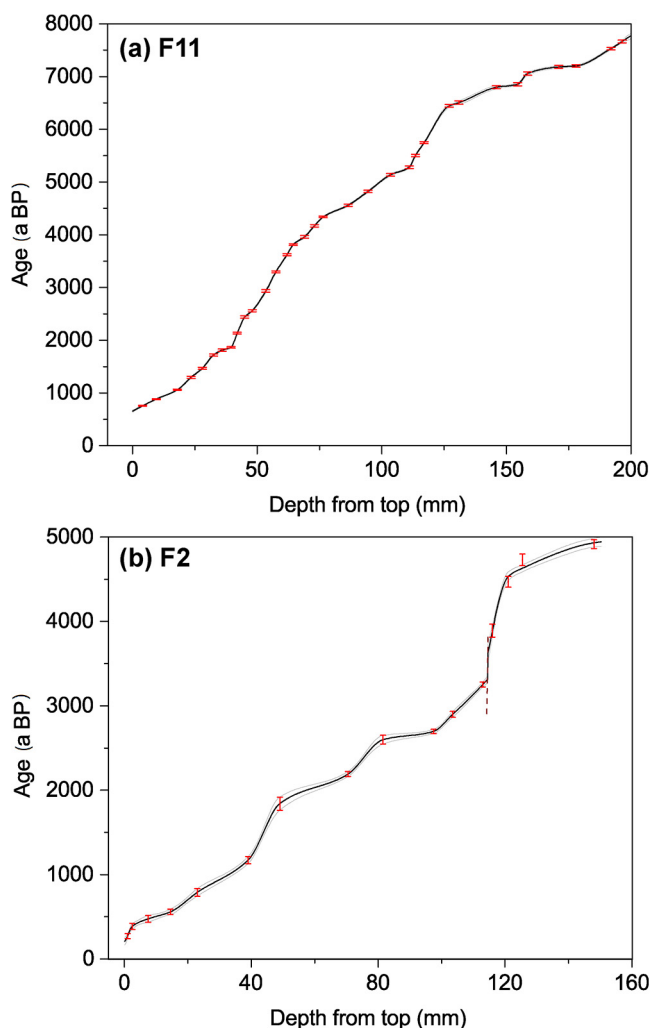
### 4. Discussion

#### 4.1. Interpretation of proxies

Previous studies suggested vegetation type and density, soil microbial productivity, and infiltration rates could affect speleothem  $\delta^{13}\text{C}$  values [26,27]. A dry climate would reduce the vegetation cover and density and soil microbial activity on one hand, and increase residence time of infiltrating water and allow more  $\delta^{13}\text{C}$ -enriched bedrock to be dissolved on the other hand. Furthermore, such a lack of water enhances prior calcite precipitation in the epikarst and reduces drip rates in the cave, which in turn allows for prolonged CO<sub>2</sub> degassing from the drip-water. All these factors combined together, show an increasing  $\delta^{13}\text{C}$  evident in the speleothem. On the contrary, a wetter climate would contribute to dense vegetation cover and microbial productivity, decreases the water-rock interaction and CO<sub>2</sub> degassing, resulting in more negative  $\delta^{13}\text{C}$  values [26–28]. Therefore, The F11 stalagmite  $\delta^{13}\text{C}$  could be viewed as reflecting local effective precipitation or precipitation-evaporation condition, with higher  $\delta^{13}\text{C}$  value reflect lower precipitation-evaporation condition (mainly in winter/spring time), and vice versa [29,30]. Prior calcite precipitation during dry conditions could also lead to higher Sr/Ca ratio in the speleothem due to preferential removal of Ca in seepage water [31], although lower growth rates caused by insufficient seepage water during dry conditions might reduce the Sr/Ca ratio at the same time [32].

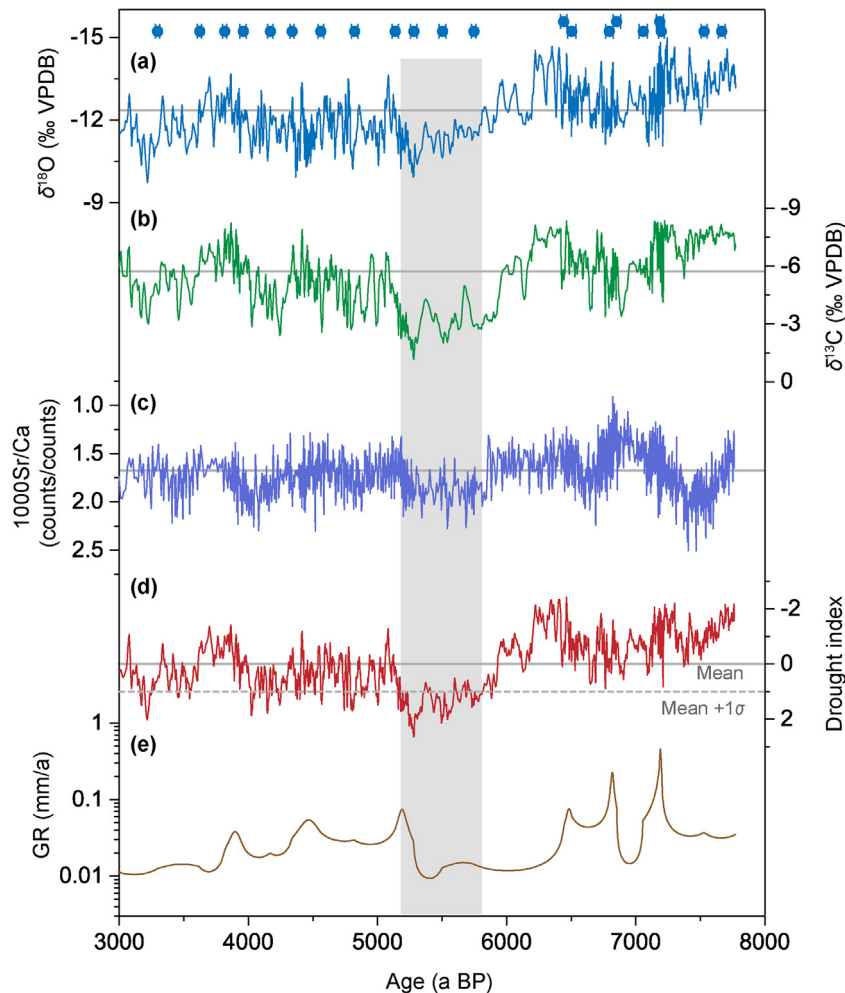
Under isotopic equilibrium fractionation conditions, stalagmite  $\delta^{18}\text{O}$  changes are controlled by precipitation  $\delta^{18}\text{O}$  and cave temperature [33]. Modern observations reveal lower precipitation  $\delta^{18}\text{O}$  during the winter half year, but higher  $\delta^{18}\text{O}$  during the summer half year (i.e., temperature effect) in the westerly controlled Eastern Mediterranean and ACA [34]. However, temperature effects cannot explain the large range of  $\delta^{18}\text{O}$  values in F11 (~5‰). Winter/spring precipitation was employed to explain the stalagmite  $\delta^{18}\text{O}$  variations in this region [30]. Increased winter/spring precipitation (rainfall/snowfall) brought by an enhanced Mediterranean storm track will result in lower annual mean precipitation  $\delta^{18}\text{O}$  values, and cause depleted  $\delta^{18}\text{O}$  values in speleothems from this region. In contrast, decreased winter/spring precipitation will enrich  $\delta^{18}\text{O}$  values in a speleothem. In addition to winter/spring precipitation amount, the  $\delta^{18}\text{O}$  changes in moisture sources, the so-called “source effect” could also affect the stalagmite  $\delta^{18}\text{O}$  [34]. However, the significantly positive correlation between the  $\delta^{18}\text{O}$  and  $\delta^{13}\text{C}$  of F11 suggests precipitation as the primary control factor.

Significant positive correlations between  $\delta^{13}\text{C}$ ,  $\delta^{18}\text{O}$ , and Sr/Ca suggest local precipitation exerts a common control on all. As a result, we applied principal components analysis (PCA) to the three proxy records (Table S3 online). PC1, which explains 62% of the total variance, was then used as a drought index for this region, with higher values representing less precipitation, and lower values representing more precipitation. Considering high evaporation and low precipitation in summer and autumn, we infer that our record largely reflects changes in winter/spring precipitation in



**Fig. 4.** (Color online) <sup>230</sup>Th age-depth plots of stalagmite F11 (a) and F2 (b). Age models were established using 5000 Monte-Carlo simulations [24]. Gray lines represent the 95% confidence intervals and black lines are the median ages, respectively. Error bars are 2σ error. A hiatus was determined in 114.5 mm from the top of the stalagmite F2.





**Fig. 5.** Proxy records of stalagmite F11 during the mid-late Holocene. (a)  $\delta^{18}\text{O}$  record; (b)  $\delta^{13}\text{C}$  record; (c) Sr/Ca record; (d) drought index record represented by the PCA results of  $\delta^{18}\text{O}$ ,  $\delta^{13}\text{C}$  and Sr/Ca records; (e) growth rate record. The grey lines in (a), (b), (c) and (d) represent the mean values of the  $\delta^{18}\text{O}$  ( $-12.37\text{‰}$ ),  $\delta^{13}\text{C}$  ( $-5.72\text{‰}$ ), Sr/Ca ratios ( $1.68 \times 10^{-3}$ ) and drought index (0) for the entire series, respectively. The grey bar marks the period with the drought index exceed  $1\sigma$  during 5820 and 5180 a BP. Blue dots with error bars represent  $^{230}\text{Th}$  dates.

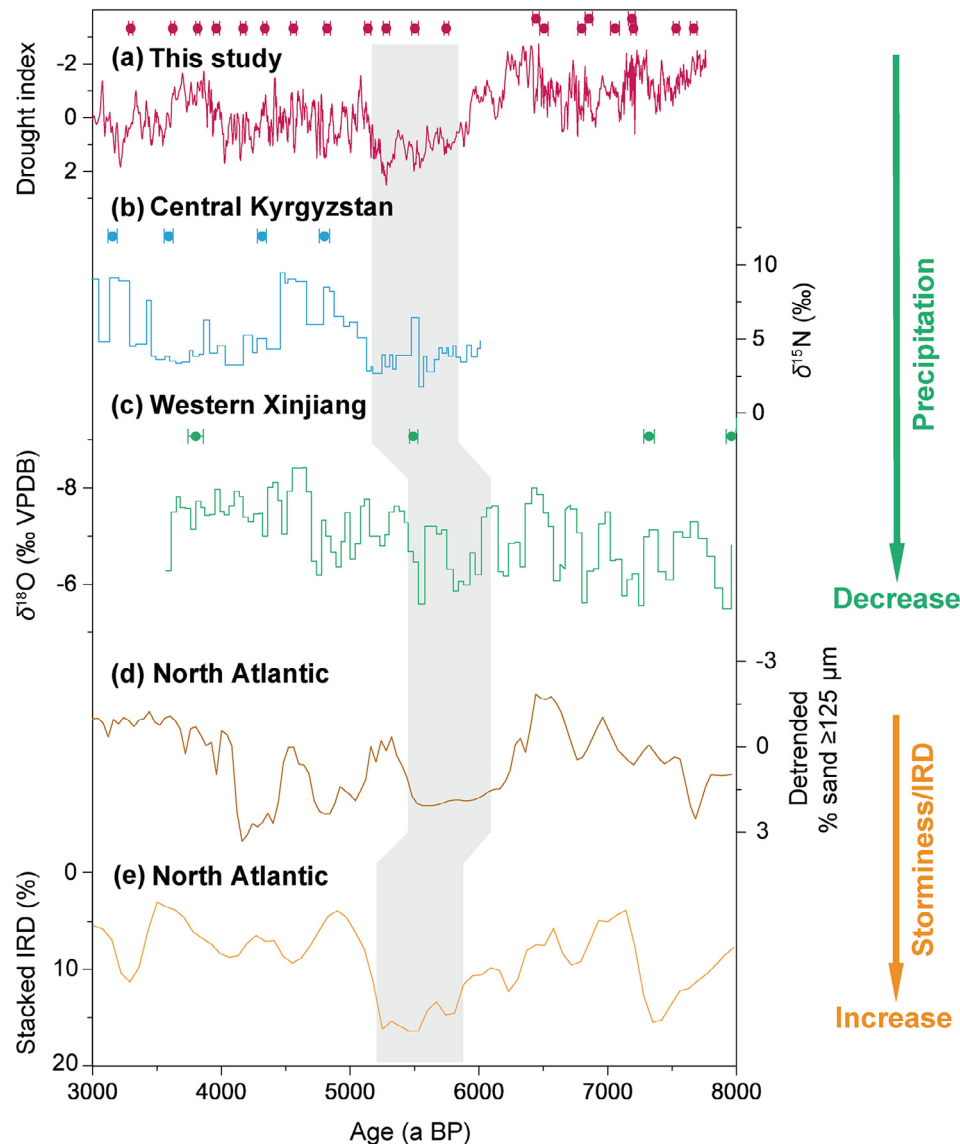
this region. The index is consistent with changes in stalagmite growth rate (Fig. 5), supporting this general interpretation.

#### 4.2. Megadrought in ACA during the mid-Holocene

Our result suggests the precipitation in ACA was relatively high before 6000 a BP, and then decreased to the lowest values at ~5280 a BP. After that, precipitation increased with substantial centennial- to decadal-scale fluctuations, until to another dry interval between 4300 and 4000 a BP. Notably, this second period of dry conditions broadly correlates with other evidence of drought across the region [35]. Precipitation increased again starting at 4000 a BP, and rainfall was relatively high from 3930 to 3610 a BP. Precipitation then decreased, before increasing to relatively high values between 3220 and 3000 a BP (Fig. 5). The most remarkable feature of the record is the interval of abnormally positive shift between 5820 and 5180 a BP, exceeding one standard deviation ( $1\sigma$ ) shift relative to data from the full record (Fig. 5), indicating a prolonged period of aridity or megadrought in ACA, lasting 640 a.

A precise comparison of our records to existing Holocene stalagmite records from western ACA [29,30] is difficult because the latter are typically characterized by substantially coarser temporal resolution and less precise chronologic constraints (Fig. S7 online).

Nonetheless, some lower-resolution records (Fig. S1 online) are consistent with the megadrought idea (Fig. 6a). A 6000 a sediment sequence from Lake Son Kol in central Kyrgyzstan recorded anomalously low  $\delta^{15}\text{N}$  values during this period [36] (Fig. 6b). Decreased  $\delta^{15}\text{N}$  indicates a reduction of input of terrestrial organic material, plausibly resulting from decreased precipitation and/or meltwater runoff during winter/spring, although summer moisture increase may have played a role [36]. During the peak of the megadrought (~5280 a BP), the level of Lake Balkhash was at least 20 m lower than at present, as indicated by abundant gypsum crystals in lake sediments [37]. In addition, multiple proxy records from Lake Issyk-Kul in northeastern Kyrgyzstan reveal a change from a fresh water open-basin to a closed-basin in the late mid-Holocene [38]. Organic  $\delta^{13}\text{C}$  in loess from southern Kazakhstan reached its highest value of the entire Holocene during this period, suggesting limited moisture availability [39]. A 30000-a record of eolian deposition in the northwestern Pacific Ocean revealed the greatest dust flux ~6000 a ago, indicating severe aridity in the Asian interior [40]. The megadrought induced dust probably caused a rapid increase in coarse-grained detrital magnetite particles in the downwind region of Dali lake [41]. Although it does not cover the entire mid-Holocene, a recently published stalagmite  $\delta^{13}\text{C}$  record from northern Iran [35] shows good agreement with our F11  $\delta^{13}\text{C}$  record during overlapping segments between 3770 and 4920 a BP



**Fig. 6.** Comparison of climate records in ACA, the eastern Mediterranean, and the North Atlantic. (a) Drought index record (this study); (b)  $\delta^{15}\text{N}$  record from Lake Son Kol in central Kyrgyzstan [36]; (c) stalagmite  $\delta^{13}\text{C}$  record from Kesang Cave in western Xinjiang, China [34]; (d) storminess activities in North Atlantic recorded by detrended coarse sand ( $\geq 125\ \mu\text{m}$ ) percentage [47]; (e) hematite-stained grains record of Ice-Rafted Debris (IRD) in North Atlantic [50]. The grey bar marks the megadrought that occurred during 5820 and 5180 a BP. Dating points with errors of the records are also shown.

( $r = 0.436$ ,  $P < 0.001$ ), implying similar precipitation variations over the western ACA.

Although long-term Holocene moisture evolution in eastern ACA, where precipitation mainly occurs in summer and autumn [42], is strongly debated [15–17,34,42,43], there is some evidence consistent with a late mid-Holocene megadrought. Substantially increased  $\delta^{13}\text{C}$  in a stalagmite from Kesang Cave in the eastern Tianshan Mountains, western Xinjiang, China [29,34] is observed during this period (Fig. 6c). A megadrought might explain the slow accumulation or even cessation of peat development in the eastern Tianshan [44] at this time. A notably dry climate was also inferred for the eastern Mediterranean region, though perhaps not as severe as that in ACA. A stacked record integrating the  $\delta^{18}\text{O}$  records from six lakes in the eastern Mediterranean, reveals continuously decreasing wetness during the late mid-Holocene [45]. This drying trend is supported by high  $\delta^{13}\text{C}$  and  $\delta^{18}\text{O}$  values in a stalagmite from Lebanon [46].

#### 4.3. Driving force of the megadrought

A northward shift of the westerly jet may have played a role in causing the megadrought. During this period, storminess in the North Atlantic dramatically increased (Fig. 6d), indicating a northward shift of the westerly jet, akin to a present-day positive North Atlantic Oscillation (NAO) [47]. Modern observations indicate that north-shifted westerlies during a positive NAO winter correlates with reduced frequency and intensity of Mediterranean storms, decreasing precipitation in central and southern Europe, the Mediterranean and parts of southwest Asia [19]. Previous studies reveal dry climatic conditions in southern Europe [48], but wetter conditions in northwestern Europe [49] during this period, consistent with this pattern. As precipitation in ACA is mainly derived from recycled moisture from the Mediterranean, Black, and Caspian Seas [19,42] (Figs. S4 and S5 online), a weakened westerly jet and dry conditions in the Mediterranean and Caspian basin

would reduce the moisture transfer to ACA, and result in drier conditions in this region [42]. In addition, strengthened and northward shifted westerlies and storm-tracks are consistent with observations of Ice-Rafted Debris in the northwestern Atlantic during this period [47,50] (Fig. 6e). This would further decrease the sea surface temperature, and reduce the evaporated moisture transported ultimately from the North Atlantic to ACA.

#### 4.4. Climate change and trans-Eurasian exchange

Modern observations suggest that winter/spring precipitation has a strong influence on grassland productivity in the Tianshan mountains [51]. The megadrought may have reduced productivity in mountain grasslands, at the same time reducing runoff to inland rivers, thereby affecting the oases. This may have limited prehistoric cultural development in ACA, and impeded the eastward dispersal of West Asian elements along the pre-Silk Roads (Fig. 1a, b). At the beginning of the megadrought, peoples ascribed to the Keltiminar Culture (8000–6000 a BP) migrated from oases between the Amu and Sir Rivers to the southern Siberian steppe [52,53]. In the mountain foothills of Central Asia, agropastoral populations do not migrate north or east until the fifth millennium BP. At this time, farming populations across the southern Kara Kum Desert reduced their ranges, concentrating in oases and along river valleys, such as the Murgab and Gyokser Oases and upper Amu River, as well as some locations along the northern foothills of the Kopet Dag range, where gravity fed irrigation was implemented [14,54] (Fig. 1b). While cereal cultivation is well known in parts of western ACA dating back to 6000 a BP, the eastward expansion into the desert oases, northern mountains, and eventually East Asia, did not occur until after 5000 a BP [7,55] (Fig. 1c, d). The megadrought would have hindered human movement and effectively reduced or blocked overland travel between eastern and western Central Asia along the pre-Silk Roads for these 6-plus centuries between the 6th to 5th millennia BP (Fig. 1b).

In contrast, the northward shift of the westerlies may have made conditions in the Eurasian Steppe more hospitable [19,56], enhancing pasturelands [57]. Indeed, relative light  $\delta^{13}\text{C}$  values in a stalagmite record from the southern Ural Mountains during this period suggest warm/humid conditions [58]. The expansion of pastoralist populations into new grasslands that were previously sparsely populated by hunter-gatherers, may explain the archaeologically complicated phenomenon that some scholars refer to as the Yamnaya migrations during the 6th and 5th millennium BP [1–3]. Humans that utilized a similar economic strategy and may have shared similarities in material culture and genetics appear to have spread along a northern steppe/forest ecotone (Fig. 1b) when these northern steppe regions were experiencing more humid conditions [3]. Steppe-derived technologies, such as metallurgy, wheat cultivation and sheep/goat herding, started to trickle into East Asia during the 5th millennium BP (Fig. 1c) [6].

At the end of the megadrought, precipitation gradually increased in ACA, although with fluctuations (Fig. 6a). It may have resulted in greater glaciation and more prominent summer-melt streams, which could have enhanced agropastoral activities and irrigation-based farming in oases and along the alluvial fans. Starting in the fourth millennium BP, the greater investment in irrigated farming across the desert oases facilitated demographic expansions in southern ACA, what many archaeologists refer to as the Bactria–Margiana Archaeological Complex (4250–3650 a BP) [59]. Population genomics illustrate that there were greater rates of intermixing between regional populations, including the first movements of people from the steppe into southern ACA (Fig. 1c, d) [60]. Small-scale agropastoral communities developed across the foothills of ACA, and the demographic expansion resulted in cultural dispersal across ACA. This cultural boom corresponded to

the wettest period (3930–3610 a BP) in our record during the last 6000 a (Figs. 5 and S6 online). These low-investment farmers further facilitated the spread of cereal crops and herd animals into East Asia by four millennia ago. At the same time, increased precipitation [15] may have fostered the rapid rise of farming settlements in Xinjiang, China, during the early 4th millennium BP.

While further research is needed to correlate human cultural responses to shifts in climatic events, the apparent large-scale linkages are enticing. Demographic shifts, including population movement and population increase provide support for models aimed at elucidating processes affecting cultural development in ACA. After the megadrought, the mobility of agropastoral groups likely facilitated the interconnection of different oases in ACA [7,61], opening oasis routes of trans-Eurasian exchange during the 4th millennium BP (Fig. 1d), laying the foundation for the subsequent organized Silk Roads.

## 5. Conclusion

We reconstructed, so far, the most precisely dated and highest-resolved mid-late Holocene precipitation record for ACA, by using two replicated stalagmites from the southeastern Fergana Valley, Kyrgyzstan. Our data reveal a 640-a, previously unknown megadrought between 5820 and 5180 a BP. The megadrought likely impeded the expansion of cultural traits along the oasis routes in ACA, and diverted the earliest transcontinental exchange along the Eurasian Steppe during the 5th millennium BP. Increased precipitation after the megadrought have made conditions for the flourishing of Bactria–Margiana Archaeological Complex, facilitating the interconnection of different oases, further opened the “pre-historic Silk Roads”, of trans-Eurasian exchange during the 4th millennium BP.

## Conflict of interest

The authors declare that they have no conflict of interest.

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## Author contributions

Liangcheng Tan directed this project. Liangcheng Tan, Guanghui Dong, Zhisheng An, and R. Lawrence Edwards formulated research questions. Liangcheng Tan, Guanghui Dong, and Fahu Chen wrote the paper. Liangcheng Tan, Hai Cheng, and R. Lawrence Edwards contributed to the  $^{230}\text{Th}$  dating. Liangcheng Tan, Yanjun Cai, and Dong Li performed the stable isotopic and trace elemental

analyses. Liangcheng Tan and Dong Li conducted the age model. Liangcheng Tan, Dong Li, and Rustam Orozbaev did the field work. Guanghui Dong, Haiming Li, Robert Spengler, and Ruiliang Liu analyzed the archeological sites. All authors discussed the results, edited, and commented on the manuscript.

## Appendix A. Supplementary materials

Supplementary materials to this article can be found online at <https://doi.org/10.1016/j.scib.2020.10.011>.

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