



Transmission Spectroscopy of the Earth–Sun System to Inform the Search for Extrasolar Life

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Abstract

Upcoming NASA astrophysics missions such as the James Webb Space Telescope will search for signs of life on planets transiting nearby stars. Doing so will require coadding dozens of transmission spectra to build up sufficient signal to noise while simultaneously accounting for challenging systematic effects such as surface/weather variability, atmospheric refraction, and stellar activity. To determine the magnitude and impacts of both stellar and planet variability on measured transmission spectra, we must assess the feasibility of stacking multiple transmission spectra of exo-Earths around their host stars. Using our own solar system, we can determine if current methodologies are sufficient to detect signs of life in Earth's atmosphere and measure the abundance of habitability indicators, such as H₂O and CO₂, and biosignature pairs, such as O₂ and CH₄. We assess the impact on transmission spectra of Earth transiting across the Sun from solar and planetary variability and identify remaining unknowns for understanding exoplanet transmission spectra. We conclude that a satellite observing Earth transits across the Sun from beyond L2 is necessary to address these long-standing concerns about the reliability of coadding planet spectra at UV, optical, and infrared wavelengths from multiple transits in the face of relatively large astrophysical systematics.

Unified Astronomy Thesaurus concepts: [Exoplanet detection methods \(489\)](#); [Exoplanet atmospheres \(487\)](#); [Exoplanet atmospheric variability \(2020\)](#); [Exoplanet surface characteristics \(496\)](#); [Exoplanet surface variability \(2023\)](#); [Active sun \(18\)](#); [Stellar activity \(1580\)](#); [Stellar granulation \(2102\)](#); [Stellar faculae \(1601\)](#); [Starspots \(1572\)](#); [Transmission spectroscopy \(2133\)](#)

1. Introduction

Over the next decade, users of the James Webb Space Telescope (JWST) will apply transit-based techniques to determine whether rocky planets orbiting M-dwarf stars have tenuous, clear, or cloudy atmospheres. Through secondary eclipse measurements, JWST users will then begin identifying which of those terrestrial planets that reside within the habitable zone (Kopparapu et al. 2013), where planetary surface temperatures are suitable for liquid water, are actually habitable. Finally, and most difficult, astronomers will analyze transmission spectra measurements in the search for signs of life. Studies, like Fauchez et al. (2019), Lustig-Yaeger et al. (2019), Wunderlich et al. (2019), and Tremblay et al. (2020), have shown that JWST will need to accumulate many dozens, if not hundreds, of transits over its 5.5 yr primary mission to build up sufficient signal to potentially confirm the presence of biosignature pairs such as O₂ and CH₄, O₃ and N₂O, and others (see Krissansen-Totton et al. 2016, and references therein). Such an observation campaign will require a substantial investment in telescope time and, thus, will likely only be attempted for a small number of targets within the mission's lifetime. While the search for signs of extrasolar life may also become possible using other techniques, transit observations with

JWST will be our best bet in the search for extrasolar life in the near term.

To build up sufficient signal as to be sensitive to atmospheric biosignatures, it will be necessary to stack dozens of transmission spectra from multiple observational epochs. Doing so will require sufficient understanding of the following concepts:

1. The impact of stellar variability on the measured transmission spectrum from starspots and faculae;
2. The impact of planetary variability on transmission spectra;
3. Whether these varying conditions will invalidate the method of stacking transmission spectra to enhance the detection significance of low signal-to-noise spectroscopic features; and
4. Whether biosignature gases in a planet's atmosphere are detectable through transmission spectroscopy.

The search for life with the transit method hinges on our understanding of the habitable zone around a star. Because current technologies, which make use of the transit method, provide the best data for short-period planets, M-dwarf stars with their small size and close-in habitable zones are the best targets to search for life until space-based direct imaging is realized. Some M dwarfs are more active and flare more frequently than their larger G-dwarf siblings (Günther et al. 2020). The Sun has long been our testing ground for understanding how stellar activity in its various forms impact



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our observations of exoplanets, particularly for radial-velocity observations (Haywood et al. 2020). Naturally, the Earth–Sun system presents the best opportunity for assessing the magnitude of these impacts and testing the feasibility of our techniques to search for life.

The fundamental goal of detecting signs of life on Earth using methods designed for exoplanets is crucial. If we are unable to account for both stellar and planet variability when the ground truth is known, or effectively stack dozens of transmission spectra to derive the presence of weak molecular features, then we should conduct the search for extrasolar life with other techniques. Work by Zellem et al. (2017) demonstrates that this is less of a concern at infrared wavelengths, yet there remains a concern for highly active stars, such as WASP-19, and at optical wavelengths (V. Panwar 2021, in preparation). Additionally, improper correction for stellar surface heterogeneities on transmission spectra in the near-infrared can stymie our interpretation of planetary atmospheric properties (Rackham et al. 2018; Iyer & Line 2020).

There are a number of barriers to detecting the spectra of Earth-like worlds. In this paper, we show how observations of the Earth in transit across the Sun can be used as a unique and high-fidelity proxy, in fact the critical example case, for potentially habitable transiting exoplanets. In Section 2 we break down the unanswered questions around detecting life as we know it using the transit technique with large collecting area, space-based observatories, such as JWST, and suggest the use of small satellites to derive these answers. In Section 3, we simulate the transit light source (TLS) effect (Rackham et al. 2018) on Earth–Sun transmission spectra. In Section 4, we discuss the potential effects of planetary variations and simulate the effects of refraction in transmission observations. In Section 5 we discuss the caveats to stacking transmission spectra and what remains untested about the technique prior to the launch of JWST. Finally, we summarize our conclusions in Section 6.

2. Transiting Exoplanet Geometry and Subsequent Biosignature Detection

A transmission spectrum is a measure of the planet’s apparent change in size as a function of wavelength. Light from the host star passes through the planet’s atmospheric annulus where it interacts with atoms and molecules. The annulus becomes opaque (and the planet appears larger) at wavelengths where these chemical species strongly absorb in the planetary atmosphere. Transmission spectroscopy data are thus sensitive to relative chemical abundances and the presence of cloud or haze particles within the atmosphere. A primary transit occurs when a planet passes in front of its host star, thus blocking part of the star’s light as seen by the observer. Transmission spectra for the Earth have generally been achieved through indirect means, by observing solar occultations (e.g. Macdonald & Cowan 2019) or via a tertiary body reflecting the filtered starlight (e.g. Pallé et al. 2009; Vidal-Madjar et al. 2010; Ugolnikov et al. 2013; Arnold et al. 2014; Yan et al. 2015; Kawauchi et al. 2018; Youngblood et al. 2020). While these clever approaches have allowed us to derive Earth’s transmission spectrum, they are only the first step toward understanding exoplanet transmission spectra in practice. For example, the Sun and the Earth’s temporal variations remain unaccounted for. A satellite-based instrument actually observing an Earth–Sun transit in a geometry analogous to that anticipated

for exoplanet transits is the only way to accomplish a true solar system laboratory test benchmark.

The transit geometry of the Earth–Sun system from within the solar system has slightly different geometry than from beyond. When seen as a point source, the fractional dip in light from the star is determined by the planet-to-star area ratio; however, when the star and planet are spatially resolved, the transit depth is determined by the bodies’ angular radii. Derivations of the equations for transit depth and molecular feature size (Equations (1) and (2)) are identical. However, when the star and planet are spatially resolved, the absolute radii are replaced by the angular radii. Importantly, the absolute scale height is also replaced by the angular scale height.

$$\begin{aligned} \text{Transit Depth} &= \frac{\theta_p^2}{\theta_*^2} = \left(\frac{\tan^{-1}\left(\frac{R_p}{d}\right)}{\tan^{-1}\left(\frac{R_*}{a+d}\right)} \right)^2 \\ &\approx \frac{R_p^2}{R_*^2} \left(\frac{a+d}{d} \right)^2 \\ \text{Feature Size} &\approx \frac{2\theta_H \theta_p}{\theta_*^2} \\ &= \frac{2 \tan^{-1}\left(\frac{H}{d}\right) \tan^{-1}\left(\frac{R_p}{d}\right)}{\tan^{-1}\left(\frac{R_*}{a+d}\right)^2} \\ &\approx \frac{2HR_p}{R_*^2} \left(\frac{a+d}{d} \right)^2. \end{aligned} \quad (1) \quad (2)$$

Here, θ_p is the angular radius of the planet, θ_* is the angular radius of the star, and θ_H is the angular scale height of the planet as seen from a spacecraft within the solar system. These angular radii are calculated in terms of the planet radius R_p , stellar radius R_* , and scale height H using the planet–spacecraft distance d and the planet’s semimajor axis a . The choice of units is arbitrary, but should be consistent across parameters. Equations (1) and (2) can be approximated, as shown, in the limit of small angles as long as $R_p \ll d$, $R_* \ll a+d$, and $H \ll d$. In this form, it is easy to see how resolved solar system transits reduce to the familiar exoplanet-analog forms in the limit where the distance from the observer to the planet greatly exceeds the planet’s semimajor axis, $d \gg a$.

A schematic diagram of the Earth transiting the Sun is shown in Figure 1. The total transit duration, t_{14} , begins at time t_1 and ends at t_4 , while the full transit duration, t_{23} spans t_2 – t_3 . The exact shape of the time-series observation is dictated by a variety of factors. Outside of transit, the baseline flux level is affected by the reflected light from Earth’s atmosphere and surface (such as ocean glint, Robinson et al. 2010, 2014; Lustig-Yaeger et al. 2018 and refracted light, Bétrémieux & Kaltenegger 2014; Misra et al. 2014; Bétrémieux & Swain 2018). These signals are dependent on the precise orientation of Earth in its rotation and, thus, in theory longitudinal variations can be mapped by interpreting the variations in the shape of transit ingress and egress. In transit, each wavelength is affected by a different possible atmospheric absorber. The geometry as pictured in Figure 1 requires the observer to be beyond the Earth–Sun L2 point, where the apparent diameter of the Earth’s disk is substantially smaller than that of the Sun, the Sun and the Earth are the same angular diameter at 0.009 au, L2 is 0.01 au away, and the Moon is roughly 0.002 au away. Spacecraft orbiting other solar system planets may be able

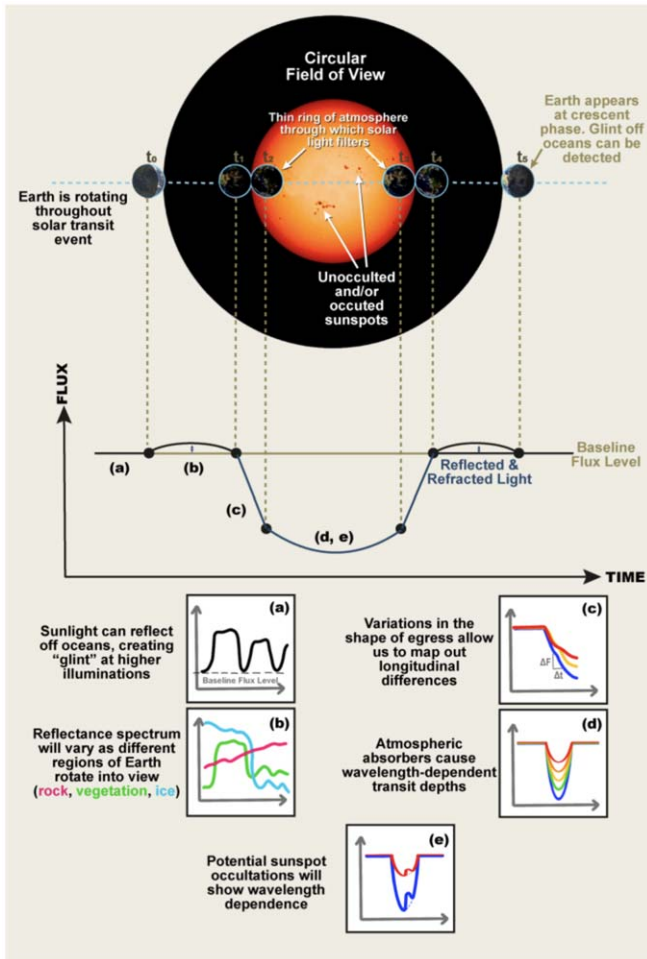


Figure 1. Schematic representation of a single Earth transit across the Sun as observed by a satellite beyond L2. We highlight several features that would be seen in these observations, including (a) ocean glint, outside of transit; (b) wavelength and temporal variability of reflected and refracted light; (c) ingress/egress as different longitudes and clouds rotate into view; (d) wavelength variability in transit depth due to absorption through and scattering by atmospheric constituents; and (e) stellar variability due to occulted and unocculted spots and faculae that can impact the measured transmission spectrum. For all subpanels except (b), axes are flux vs. time and colors represent different wavelengths; for (b) the axes are flux vs. wavelength and the colors represent different contributions.

to achieve the requisite geometry, but these instruments risk damage from imaging an object as bright as the Sun. Also, not all spacecraft orbiting other solar system bodies are able to observe Earth–Sun transits because the planets are not perfectly coplanar.

Figure 2 displays a model transmission spectrum of the Earth at a distance of 0.03 au. The model assumes a present-day, Earth-like atmospheric composition (Robinson 2017, 2018) and was originally generated by T. Robinson for a TRAPPIST-1e-like system configuration, but subsequently was rescaled to meet our needs. Using currently available detectors with modest noise floors, molecules could be clearly detected and their relative abundances readily constrained. The limiting molecular absorber in terms of detectability is methane at $2.3\ \mu\text{m}$. The narrowest molecular feature that needs to be resolved spectrally is O_2 at $0.76\ \mu\text{m}$ (760 nm). The features of O_3 at $0.25\ \mu\text{m}$ and CH_4 at $2.3\ \mu\text{m}$ set the minimum and maximum wavelength range for observations. Additionally, observing wavelengths beyond $2.5\ \mu\text{m}$ would require active cryogenic cooling (Beletic et al. 2008), which can be difficult in a small spacecraft form factor.

The one-scale-height molecular feature size from the Earth’s atmosphere as it transits the Sun would normally be very small (about 0.2 ppm) if viewed from a neighboring stellar system; however, the close proximity of a satellite to the Earth presents a unique opportunity to boost the signal size by three orders of magnitude, because the planet is proportionally larger in comparison to the host star. The top panel of Figure 3 plots the computed transit depth and feature size as a function of distance from the Earth, while the bottom panel does the same for the fraction of total transit duration. Observing transits of the Earth–Sun system from closer than roughly 0.02 au is suboptimal, because the full transit time is very short. At a distance of 0.03 au, the transit depth is 10% and the 1H feature size is 250 ppm. Although a 10% transit depth may not seem analogous to a real terrestrial exoplanet observation, there is fundamentally no difference from a Jupiter-sized planet observation in this case once the planet is fully in transit (t_2 – t_3). Furthermore, these 250 ppm feature sizes are analogous to those seen in hot Jupiter atmospheres (Iyer et al. 2016). Therefore, the lessons learned from a sequence of satellite observations at distances greater than 0.02 au will apply directly to the thousands of known planet-hosting systems. Such observations can be accomplished by a small, single-instrument satellite capable of observing Earth transits across the Sun beyond the Earth–Sun L2 point in the 0.2 – $2.5\ \mu\text{m}$ range with a sufficiently sensitive spectrometer to capture key atmospheric constituents.

3. Stellar Variability

The TLS effect (Rackham et al. 2018, 2019) can have a significant impact on the measured transmission spectrum due to photospheric heterogeneities (i.e., spots and faculae) either inside or outside the transit chord. If a star’s spectrum inside the transit chord differs from the disk-integrated spectrum then that difference is imprinted on the planet’s measured transmission spectrum. For cooler M dwarfs, this contamination can be up to 10 times larger than the expected one-scale-height feature size for a rocky planet atmosphere and can lead to the false detection of chemical species such as water (Rackham et al. 2018).

Figure 4 demonstrates the impact of the TLS effect on a simulated observed transit. Following the formalism of Loudén et al. (2017), we assume the true transit depth is a function of the observed depth and the properties of the spotted regions of the stellar disk. We use a sample blackbody model for the solar, spot, and faculae component spectra. Typical spot and faculae levels for a Sun-like star are taken from Rackham et al. (2019), which broadly agree with the activity levels seen on the Sun, and are used to calculate expected transit depth variations due to unocculted activity features during transit. Following Rackham et al. (2019) we calculate the spot temperature as $T_{\text{spot}} = 0.418 \times T_{\text{phot}} + 1620\ \text{K}$ and the faculae temperature as $T_{\text{fac}} = T_{\text{phot}} + 100\ \text{K}$, assuming a stellar photosphere of 5800 K. Our low, moderate, and high activity levels follow the bounds of 1σ uncertainties on activity levels in Rackham et al. (2019)—specifically faculae covering fractions of 6%, 10%, and 18%, respectively, with 0.6%, 1.1%, and 2.2% covering fractions for the spot component. The baseline transit spectrum is Earth as observed by a satellite at 0.03 au. While the TLS effect is an unconstrained problem in exoplanet observations, our knowledge of spot and faculae distributions on the Sun during observed transits will allow us to remove the TLS effect at high

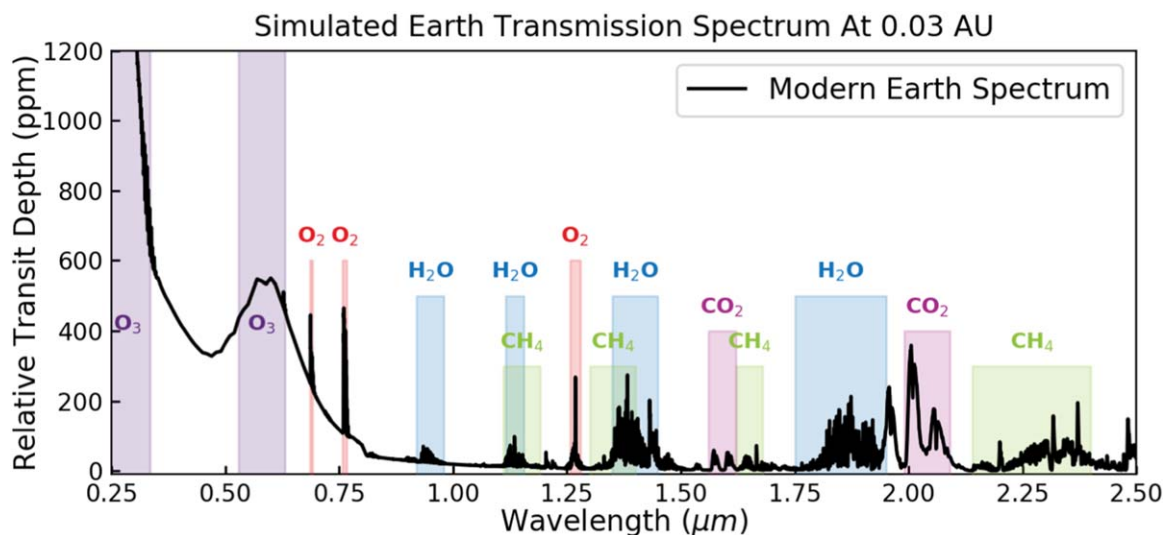


Figure 2. Simulated Earth transmission spectrum at a distance of 0.03 au from the Earth (originally generated by T. Robinson; Robinson 2017, 2018). Colored regions indicate wavelength ranges for the labeled molecular absorption features. A spacecraft with broad wavelength coverage would be sensitive to standard habitability indicators (H_2O , CO_2) and biosignature pairs (O_2 and CH_4 , O_3 and CH_4) that JWST will search for in M-dwarf planet atmospheres.

accuracy, while testing and comparing to methods used to account for stellar activity in exoplanet transits.

Observed variations in the measured exoplanet transit depth and shape are usually associated with unocculted and occulted stellar features, respectively (Pont et al. 2007; Alonso et al. 2008; Knutson et al. 2011; Kreidberg et al. 2014; Bruno et al. 2018). If not properly accounted for, stellar variability can impact derived parameters such as the orbital inclination, stellar density, and limb-darkening coefficients. Numerous strategies exist to correct for stellar variability. For example, in their work deriving the transmission spectrum of TRAPPIST-1g, Wakeford et al. (2019) used out-of-transit TRAPPIST-1 spectra to reconstruct the stellar flux using three model components ($T_{\text{eff}} = 2400$ K, 35% spot coverage at 3000 K, and a 5800 K hot spot covering $<3\%$ of the surface). An observation of an Earth–Sun transit has an advantage over these exoplanet observations, because Earth-based solar observatories are actively monitoring the Sun. To best test common methods of removing stellar variability in exoplanet transits, one can make use of the Sun’s known spot and faculae distribution—regularly monitored and mapped by multiple existing Sun-observing spacecraft and instruments such as space-based observations done with the Solar Dynamics Observatory (Couvidat et al. 2016) and ground-based add-ons to HARPS (Milbourne et al. 2019) and ESPRESSO (N. Santos 2021, private communication)—during each transit, as well as employ traditional exoplanet analysis methods involving photometric monitoring of the Sun outside of transit to derive activity levels.

One such measure of activity is the presence of sunspots on the solar disk. Figure 1(e) demonstrates an example of an occulted sunspot, where a bump in the transit light curve represents the temporary decrease in occulted light as the Earth passes over the dimmer region of the Sun (Miller-Ricci et al. 2008; Rabus et al. 2009). While occulted spots such as this make their presence known and can be useful (Sanchis-Ojeda & Winn 2011; Sanchis-Ojeda et al. 2013), it is the unseen unocculted spots that require assumptions about spot coverage fractions and temperature differences for appropriate corrections in exoplanet observations. Because stellar activity can mimic atmospheric spectral features from the planet (McCullough et al. 2014; Sing et al. 2015;

Bruno et al. 2020), it is crucial that their effects be accurately accounted for in exoplanet observations. An Earth–Sun transit data set facilitates the testing of the corrections applied to exoplanet transits and enables more effective strategies when faced with stellar variability from unresolved sources. For example, a first analysis of the data could be done with no knowledge of the spot locations, but, thanks to long-term studies of the Sun as a resolved star, the locations of solar activity from independent solar observations can be used to aid data analysis. Potential strategies might include improving occulted starspot modeling techniques and/or developing new strategies for transit spectroscopy observations of exoplanets using complementary instruments.

4. Planetary Variability

Exoplanet transmission spectra are limb-averaged observations that contain information from an ensemble of light rays transmitted through all planetary latitudes (Feng et al. 2016, 2020; Pluriel et al. 2020), including contributions from the northern and southern hemispheres, and from clear and cloudy regions. These spatial differences evolve with time as the planet rotates and the seasons change. Thus, planetary variability poses a unique challenge for the accurate interpretation of heterogeneous Earth-like exoplanets that exhibit both spatial variabilities, blended within each transit observation, and temporal variabilities, from one observed transit to the next. A satellite at the proper distance to observe the full annulus of the Earth’s atmosphere can directly probe the magnitude of planetary variability seen in the transmission spectrum of Earth and enable the development of retrieval methods for correctly interpreting exoplanet transit observations.

Earth’s equator-to-pole temperature gradients and active hydrological cycle drive spatial variations that will be integrated together in the exoplanet-analog transmission spectra. Notably, Earth is a partially cloudy planet that permits both clear and cloudy optical paths through its atmosphere. These transit spectroscopy observations are necessary as they will be used to assess whether the combination of clear and cloudy optical paths allows for the accurate retrieval of the cloud-top pressure (Line & Parmentier 2016), and whether the

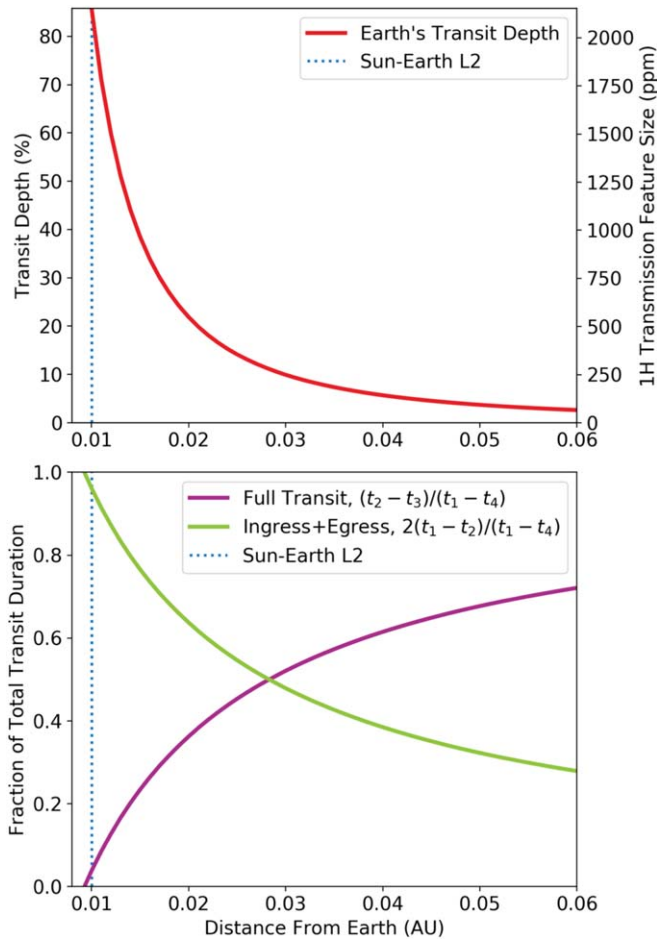


Figure 3. Top: Earth–Sun transit depth and one-scale-height feature size vs. distance from the Earth. The reasonable transit depths and feature sizes just beyond the L2 Lagrange point imply that atmospheric characterization of a transiting Earth is achievable using currently available technology. For example, at a distance of 0.03 au from the Earth, a satellite could observe transit depths of 10% and measure spectroscopic features due to H_2O , O_2 , O_3 , CH_4 , and CO_2 with 1H signal sizes of 250 ppm (see Figure 2). Bottom: the fraction of total transit duration spent in full transit (violet) and ingress/egress (green) vs. distance from the Earth. A spacecraft closer than 0.02 au from the Earth may observe deeper transit depths with larger feature sizes; however, the full transit duration is significantly shorter and, thus, provides less information overall. Distances of 0.02–0.05 au from the Earth provide a good balance of measurable feature sizes and relatively long full transit durations.

inferred abundances of atmospheric gases such as H_2O and O_3 are consistent with the range of spatial and vertical variation seen across the Earth.

Earth’s axial tilt and geography creates measurable seasonal variability in the thermal structure, cloud patchiness, and relative abundance of trace atmospheric species. For example, Earth’s integrated spectra are distinct at winter and summer solstice, due in part to the aggregation of land in its northern hemisphere (Olson et al. 2018; Mettler et al. 2020). While tidally locked planets do not have seasons,⁶ it is important to quantify the magnitude of this effect in relation to other sources of planetary variability.

Earth is expected to exhibit some level of variability due to its rotation during a single transit (Pallé et al. 2008). For example, a full transit lasting six hours in duration will sample half of the

⁶ The timescale over which tidal obliquity erosion occurs predicts that axial tilts are unexpected for synchronously rotating planets (Goldreich 1966; Heller et al. 2011).

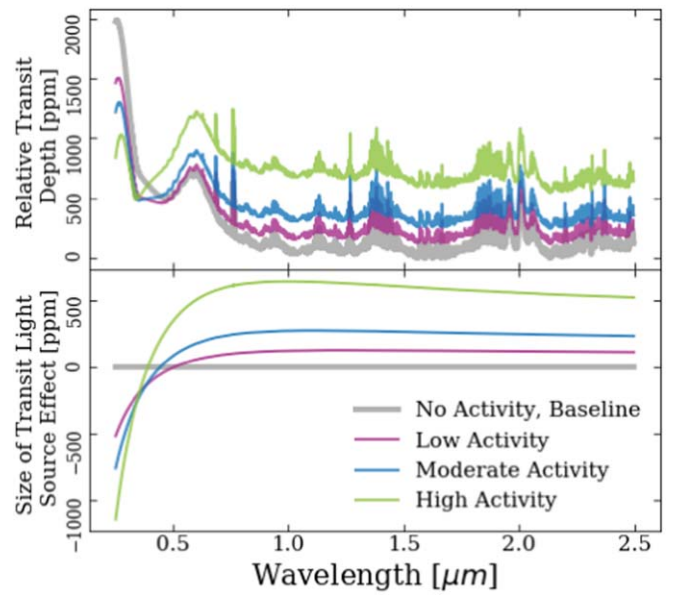


Figure 4. Example of the transit light source effect on an Earth–Sun transit observed by a satellite at a distance of 0.03 au. Top: the relative transit depth of the Earth depends on the amount of activity present on the Sun. Bottom: the size of the transit light source effect for varying levels of solar activity.

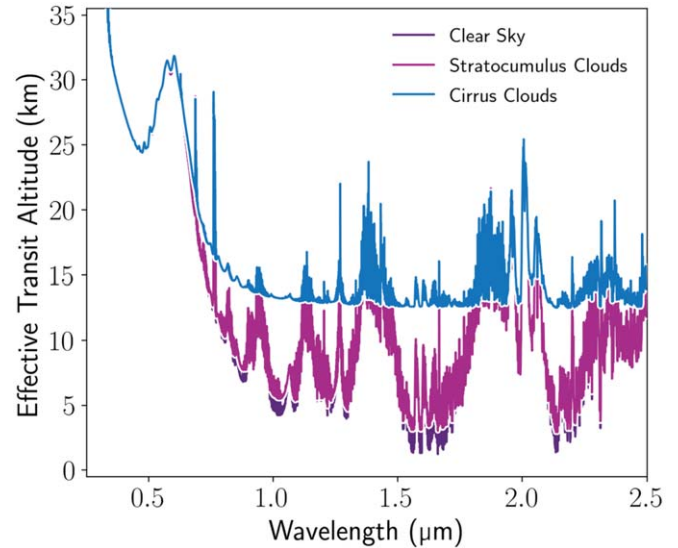


Figure 5. Model transmission spectra of Earth for different end-member cloud cases. Earth without any clouds (purple) probes down to the near surface, while stratocumulus clouds (magenta) and cirrus clouds (blue) raise the observed continuum. An actual observation of Earth would be a combination of these end-members depending on Earth’s orientation and particular weather patterns, such as seasonal developments of large storms.

Earth’s atmosphere. Thus, each segment can provide a semi-independent measurement of planetary variability as if they were subsequent transits from a tidally locked M-dwarf planet. Furthermore, large variations in weather patterns, forest fires, volcanic activity, etc., can have a measurable impact on relatively short timescales. In Figure 5 we show the transmission spectrum of the Earth with different atmospheric constituent end-members, i.e., as if the entire limb was 100% clear, 100% stratocumulus clouds, or 100% cirrus clouds. We model cirrus (ice) and stratocumulus (liquid) water clouds using the same Earth-validated model as Robinson et al. (2011), which was based on

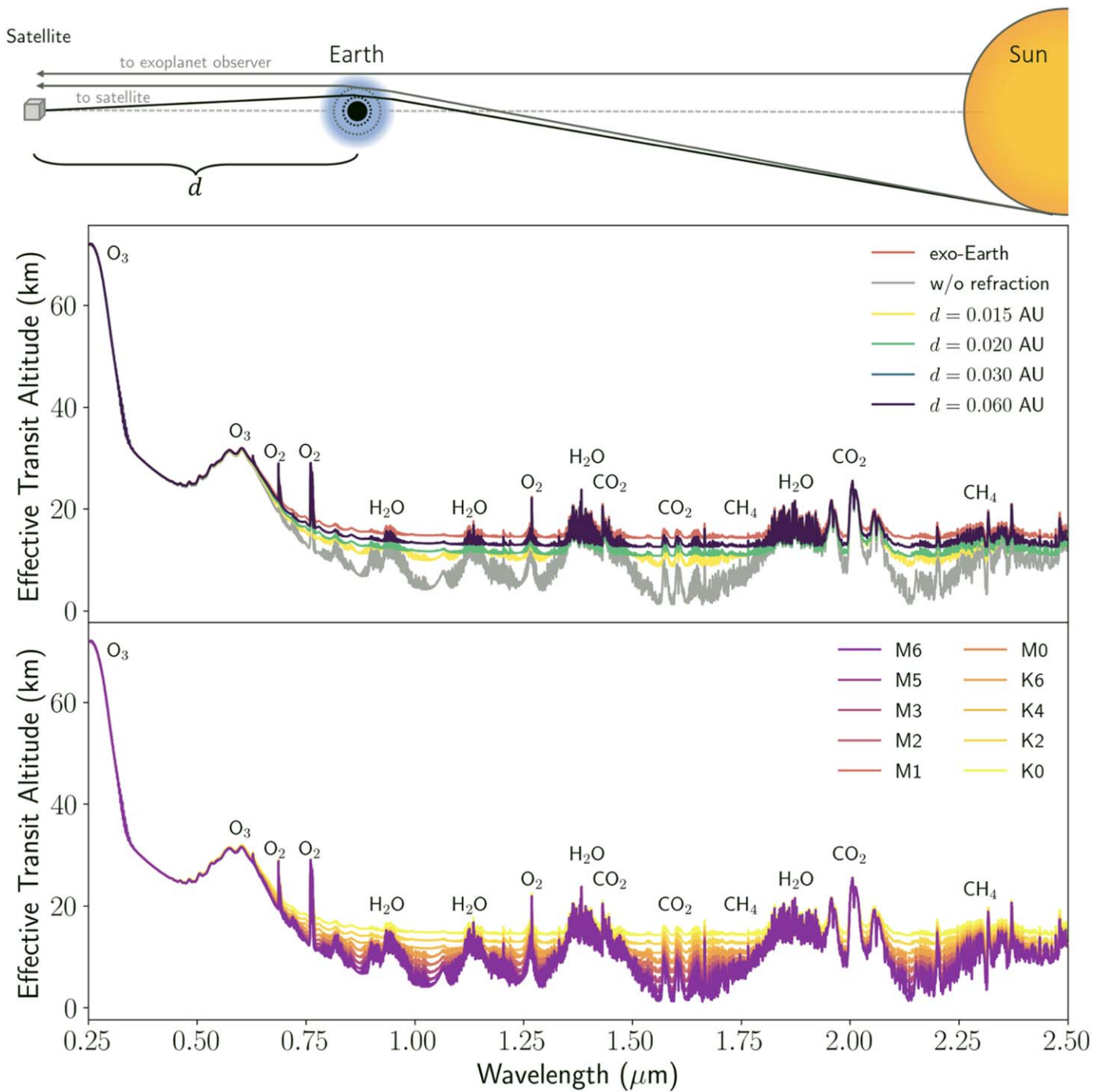


Figure 6. Refraction plays a crucial role in setting the transmission spectrum continuum for Earth and Earth-like exoplanets. Top: solar rays from the limb of the Sun refract as they transmit through the Earth’s atmosphere, creating a critical refraction limit (dotted circles) that restricts the depth probed into the atmosphere. The gray lines show an optical path and the refraction limit for an exo-Earth, while black lines show the same for a potential satellite at various Earth–spacecraft distances (d), which probes deeper due to the converging rays toward the spacecraft. Middle: simulated transmission spectra with and without refraction for an exo-Earth (red and gray lines, respectively), which bracket the observed transmission spectra that vary with distance from Earth. Note that the line for 0.06 au is on top of the line for 0.03 au. Bottom: simulated transmission spectra of possible exo-Earths around various K and M dwarfs. Depending on the distance an Earth transit is observed from, the satellite can access regions of the Earth’s atmosphere that are analogous to Earth-like planets transiting K and M dwarfs; such observations are highly relevant to future exoplanet observations. Note again that the M1 line is on top of M0 and M3 is on M2.

spacecraft measurements, and is described in greater detail in Meadows et al. (2018) and Lincowski et al. (2018). Cirrus clouds assume optical properties from B. Baum’s Cirrus Optical Property Library (Baum et al. 2005), and consist of a variety of particles including 45% solid columns, 35% plates, and 15% 3D bullet rosettes, spanning 2–9500 μm with a cross section weighted mean diameter of 100 μm . Stratocumulus clouds assume Mie scattering

optical properties with a two-parameter gamma distribution ($a=5.3$, $b=1.1$) and mean particle radius 4.07 μm , and use the refractive indices of water from Hale & Querry (1973). Based on approximate Earth averages (Robinson et al. 2011), cirrus clouds are placed near 8.5 km altitude and have an optical depth of 3, while stratocumulus clouds are placed near 1.5 km and have an optical depth of 10. A satellite observation of Earth in transit

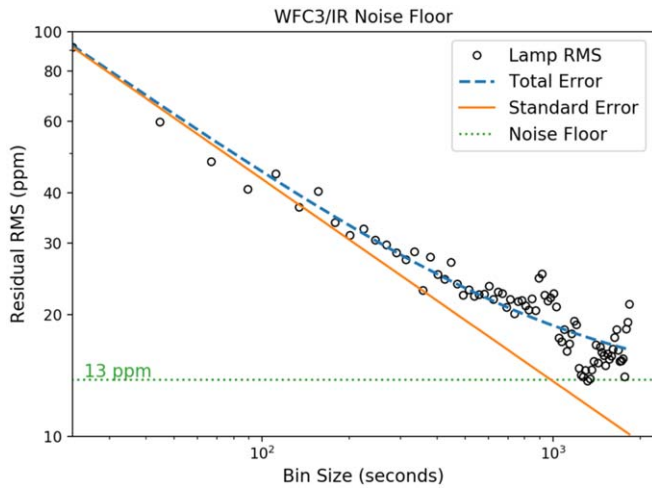


Figure 7. Measured residual rms vs. bin size (black circles) using time-series Tungsten lamp data obtained as part of HST calibration program 15400. At small bin sizes, the measured RMS follows the expected standard error (solid orange line). At larger bin sizes, the RMS begins deviating from the standard error. The best-fit total error (dashed blue line) requires a noise floor of 13 ppm (dotted green line).

would be a combination of these end-members depending on Earth’s physical orientation and temporal variations in weather patterns, such as seasonal developments of large storms. Thus, we find that it is necessary to observe multiple transits of Earth to measure the magnitude of these variations in Earth’s transmission spectra and trace them back to specific sources via comparison with publicly available independent data acquired from Earth-observing satellites. These data will enable us to quantify the impact of planet variability on exoplanet transmission spectra and assess the feasibility of combining transit data over many years.

Recently, Macdonald & Cowan (2019) published an empirical clear sky transmission spectrum of Earth from 2–14 μm using solar occultation measurements. This was a first step to establishing such a data set to assess planetary variability. However, this study was limited by its 2–14 μm wavelength range, missing the UV and optical, and the omission of clouds to produce a clear sky transmission spectrum of Earth. The transmission spectrum of Macdonald & Cowan (2019) was unable to assess crucial regions of Earth’s spectrum, because a much wider window is needed from the UV to the NIR to include strong O_2 absorption (0.76 and 1.27 μm) and Rayleigh scattering. To augment the wavelength range observed previously and detect these important species, the UV and the optical must also be observed. Clouds are well known to play a crucial role in determining an exoplanet’s transmission spectrum continuum, thereby limiting the depth into the atmosphere that can be probed (Fortney 2005; Kreidberg et al. 2014; Sing et al. 2016). Limb-integrated observations of Earth as a true exoplanet transit analog are thus fundamentally necessary and will contain contributions from clear and cloudy optical paths. Because we can independently assess the state (e.g. cloudiness, etc.) of the Earth’s atmosphere in a resolved sense, we can disentangle the contributions of different forms of terrestrial atmospheric variability from a disk-integrated spectrum.

4.1. Refraction

The effects of refraction were not directly observed in the Macdonald & Cowan (2019) transmission spectrum of Earth,

because it was produced from solar occultation measurements from an Earth-orbiting satellite, the Canadian low-Earth orbit satellite SCISAT. As a result, it was able to probe deeper into the Earth’s atmosphere (down to about 4 km) than any distant exoplanet observer would be able to during transit due to the critical refraction limit (Bétrémieux & Kaltenegger 2014; Misra et al. 2014). To mimic the effects of refraction on the spectral continuum analogous to terrestrial exoplanet observations around K and M dwarfs, an observer needs to be some distance from Earth. An exact Earth–Sun analog exoplanet system would have a refraction continuum floor at about 14 km altitude above the solid surface, but for an Earth transiting an M dwarf, the refraction floor can drop below 10 km (Lincowski et al. 2018; Meadows et al. 2018).

In Figure 6, we show the effect of refraction on an exo-Earth (an Earth-like planet at the inner edge of the habitable zone transiting various K and M dwarfs) observation as compared to an observation via satellite of Earth at various distances from 0.015 au to 0.06 au. Transmission spectrum models were produced using the Spectral Mapping Atmospheric Radiative Transfer (SMART) model (Meadows & Crisp 1996) with the ray-tracing refraction transmission spectrum upgrades described by Robinson (2017). We used a spatially averaged 1D atmospheric thermal structure and composition for Earth from Robinson et al. (2011) as an input into the spectral model, and we simulated exo-Earths orbiting various M dwarfs assuming each planet is located at the inner edge of the Kopparapu et al. (2013) habitable zone for each respective star. Since the transmission spectrum code is intended to be used to simulate exoplanet observations at infinity, we modeled the convergence of deflected solar rays toward the satellite (as shown in the upper diagram in Figure 6) by simulating the atmospheric radiative transfer in a rotated reference frame. In the rotated reference frame, light rays may originate from beyond the solar disk and emerge from the Earth’s atmosphere parallel as if heading to an observer at infinity. Thus, by artificially extending the solar radius by the exact amount necessary to preserve the maximum angle of deflection, we are able to simulate identical optical paths to a spacecraft observing the transit of Earth, assuming it lies along the Sun–Earth line to preserve symmetry.

The advantage of a satellite observing a transit at a greater distance from Earth than 0.015 au would thus be to observe refraction continuum floors at slightly different altitudes as shown in Figure 6. These regions of the lower atmosphere are analogous to exoplanets transiting K- and M-dwarf stars. At 0.015 au away from the Earth, the refraction floor will be at an altitude of 8.7 km, while at 0.06 au the floor will rise to 12.5 km. These hard refraction boundaries may limit the ability to characterize the lower atmosphere of Earth; however, this is a necessary test to assess the capabilities of exoplanet transmission spectroscopy and inform future modeling work. Additionally, refraction is also expected to induce subtle time-dependent effects into Earth’s transmission spectrum due to asymmetries in the atmospheric region probed, which changes depending on the exact alignment of the Earth and Sun (Misra et al. 2014). A satellite at an appropriate distance from the Earth ($\gtrsim 0.03$ au) enables a rigorous study into the numerous factors that inject variability into transmission spectrum observations.

5. Stacking Transmission Spectra

With perfect detectors and no noise floor, stacking (or coadding) exoplanet transmission spectra is a viable technique

To improve our understanding of exoplanet biosignature detectability	Science Questions	Science Objectives	Science Observables	Measurements
	1. Is the transit technique an effective means to detect signs of life on Earth-like planets?	Detect habitability indicators in Earth's atmosphere by measuring the abundance of diagnostic chemical absorbers during transit across the Sun	Transmission spectra features of H ₂ O (0.9, 1.1, 1.4, 1.9 μ m), O ₂ (0.69, 0.76, 1.27 μ m), O ₃ (0.25-0.34, 0.4-0.65 μ m), CH ₄ (1.2, 1.4, 1.7, 2.3 μ m), and CO ₂ (1.6, 2.0 μ m) in Earth's atmosphere	0.25-2.5 μ m transmission spectra to detect H ₂ O, O ₂ , O ₃ , CH ₄ , and CO ₂
		Detect signs of life in Earth's atmosphere by measuring the abundance of biosignature gasses during transit across the Sun	Transmission feature biosignature pairs (e.g., CH ₄ and O ₂ , CH ₄ and O ₃)	
	2. What is the impact of stellar variability on interpretability of Earth's own variable transmission spectrum?	Constrain the magnitude of Earth's seasonal, rotational, and meteorological variability by obtaining multiple transmission spectra spanning the relevant timescales	Transmission spectra of Earth spanning ≥ 6 months (≥ 2 seasons) and of sufficient duration (≥ 6 hours) to exhibit rotational/meteorological variability	Transmission variability of Earth across the Sun
		Constrain the impact of stellar inhomogeneity (e.g., spots, faculae) on the measured transmission spectrum	Same as for Earth, compared to contemporary independent observations of solar variability (existing ground-based and orbital)	Spectral variability of Sun behind Earth
	3. What is the efficacy of stacking transmission spectra to enhance the detection significance of low-signal-to-noise spectroscopic features?	Assess the feasibility of stacking multiple Earth transmission spectra when deriving atmospheric abundances of potentially habitable exoplanets transiting nearby M-dwarf stars	At least 40 hours of full-transit spectra spanning at least two seasons and at various separations from the Earth	Variation in spectral feature size, approaching the instrument noise floor

Figure 8. The high-level science questions pertaining to the detection of atmospheric biosignatures using the transit technique. The corresponding science objectives can address these questions using a relatively small and inexpensive spacecraft by inducing Earth transits across the Sun.

to improve precision and increase the detection significance of atmospheric constituents with weak spectroscopic features. The concept has previously been applied in limited capacities for transiting exoplanets. Notably, Kreidberg et al. (2014) obtained 15 transits of the sub-Neptune GJ 1214b using the WFC3 instrument on the Hubble Space Telescope (HST). However, the brightness of GJ 1214 did not push the limits of the detector down to its noise floor and the measured flat spectrum prohibited any constraints on the atmospheric composition. In reality, all detectors have inherent noise floors, and most are poorly understood. For example, Greene et al. (2016) adopt noise floors of 20, 30, and 50 ppm for JWST's NIRISS, NIRCams, and MIRI instruments, respectively, whereas Schlawin et al. (2021) estimate a NIRCams noise floor of only 9 ppm. Similarly for HST, Stevenson & Fowler (2019) analyzed eight years of WFC3 time-series observations to estimate a noise floor of < 21 ppm at a resolving power of ~ 40 . Using Tungsten lamp data from HST calibration program 15400, we estimate a WFC3 noise floor of 13 ppm (see Figure 7). While Zellem et al. (2017) modeled and analyzed stacking transits and the impact of epoch-to-epoch stellar variability, ultimately, stacking transmission spectra from a well calibrated instrument to search for signals that are comparable in magnitude to a detector's noise floor is problematic and unproven.

Better characterization and enhanced detector stability is a long-recognized need. One of the primary challenges is in identifying and maintaining a sufficiently stable light source to conduct tests and calibrate instruments in the lab. The most promising lead is blackbody emission from a source that is temperature controlled at the sub-mK level. The goal is to achieve < 10 ppm stability over several hours (the typical duration of an exoplanet transit observation) and then coadd spectra from individual runs (e.g., Tremblay et al. 2020).

Therefore, the problem is a combination of not enough signal from a planet as well as too much noise that is poorly characterized.

The ideal solution is for a spacecraft to obtain many hours of transmission spectra from an extremely bright and photometrically stable source, such as the Sun. In doing so, we could determine the limits of its detectors (just as exoplanet observations will test JWST's detectors) and validate the process of stacking exoplanet transmission spectra to derive chemical abundances from constituents with spectral feature sizes comparable to the instrument noise floor. Morley et al. (2017) determined that 40 transits are sufficient to characterize planets such as GJ 1132b and TRAPPIST-1b. Observations of the Earth can achieve levels of precision that, to date, no exoplanet observation has achieved. Nor has a single planet's atmosphere yet been characterized using 40+ transit observations. Such exquisite transmission spectra of the Earth would validate the effectiveness of stacking transmission spectra and lead to establishment of best practices for such techniques for future missions. GJ 1132b has a transit duration of ~ 47 minutes (Berta-Thompson et al. 2015), and TRAPPIST-1b has a transit duration of ~ 36 minutes (Gillon et al. 2017); therefore, an assumption of a 1 hr transit is representative of this class of planet. Thus, in order to simulate the stacking of 40 transits, a spacecraft would need to obtain 40+ hr of in-transit data depending on the duration of each transit and how much of the Earth's rotation was captured.

6. Conclusions

In this paper, we have examined the prospects for solving three major challenges in detecting biosignature gases on terrestrial planet atmospheres: stellar variability, planet variability, and how











these affect our fundamental technique of stacking spectra to build up the requisite signal to noise. We outline our resulting science questions and objectives in Figure 8.

We conclude that a true Earth–Sun transit observation is not only needed, but optimal for putting our techniques for exoplanet atmospheric characterization and biosignature detection to the test. An Earth–Sun transit data set collected from a dedicated small satellite would facilitate the testing and development of effective strategies to correct for stellar variability from unresolved sources. We also find that it is necessary to observe multiple transits of Earth to measure the magnitude of planetary variations in Earth’s transmission spectra and trace them back to specific atmospheric and surficial sources. Limb-integrated observations of Earth as a true exoplanet transit analog are necessary to develop techniques to constrain spatial inhomogeneity in cloud cover. The transmission spectra of Earth by Macdonald & Cowan (2019) is able to probe deeper into the Earth’s atmosphere than possible by an exoplanet observer, and thus these critical refraction boundaries have yet to be understood in the context of interpreting exoplanet spectra. We conclude that observing transits of the Earth–Sun system from further than L2 via a satellite capable of imaging the Sun are optimal for detection of biosignature gases on Earth at a similar geometry and signal size to that expected on extrasolar planets and the testing of the effectiveness of stacking transmission spectra.

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