



# **International Ocean Discovery Program**

## **Expedition 396 Preliminary Report**

### **Mid-Norwegian Margin Magmatism and**

### **Paleoclimate Implications**

**6 August–5 October 2021**

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This publication was prepared by the *JOIDES Resolution* Science Operator (JRSO) at Texas A&M University (TAMU) as an account of work performed under the International Ocean Discovery Program (IODP). This material is based upon work supported by the JRSO, which is a major facility funded by the National Science Foundation Cooperative Agreement Number OCE1326927. Funding for IODP is provided by the following international partners:

National Science Foundation (NSF), United States  
Ministry of Education, Culture, Sports, Science and Technology (MEXT), Japan  
European Consortium for Ocean Research Drilling (ECORD)  
Ministry of Science and Technology (MOST), People's Republic of China  
Korea Institute of Geoscience and Mineral Resources (KIGAM)  
Australia-New Zealand IODP Consortium (ANZIC)  
Ministry of Earth Sciences (MoES), India

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## Citation

Planke, S., Berndt, C., Alvarez Zarikian, C.A., and the Expedition 396 Scientists, 2022. *Expedition 396 Preliminary Report: Mid-Norwegian Continental Margin Magmatism*. International Ocean Discovery Program. <https://doi.org/10.14379/iodp.pr.396.2022>

## ISSN

World Wide Web: 2372-9562

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## Abstract

The opening of the North Atlantic about 56 My ago was associated with the emplacement of the North Atlantic Igneous Province, including the deposition of voluminous extrusive basaltic successions and intrusion of magma into the surrounding sedimentary basins. The mid-Norwegian Margin is a global type example of such volcanic rifted margins and is well suited for scientific drilling with its thin sediment cover and good data coverage. During International Ocean Discovery Program Expedition 396, 21 boreholes were drilled at 10 sites in five different geological settings on this volcanic margin. The boreholes sampled a multitude of igneous and sedimentary settings ranging from lava flow fields to hydrothermal vent complexes, along with thick successions of upper Paleocene and lower Eocene strata. A comprehensive suite of wireline logs was collected in eight boreholes. The main goals of the expedition were to provide constraints for geodynamic models to test different hypotheses that can explain the rapid emplacement of large igneous provinces and the hypothesis that the associated Paleocene/Eocene Thermal Maximum was caused by hydrothermal release of carbon in response to magmatic intrusions. Successful drilling, combined with high core recovery of target intervals of all nine primary sites and one additional alternate site, should allow us to achieve these goals during postcruise work.

## 1. Introduction

Continental extension, breakup, and the formation of new mid-oceanic spreading centers are fundamental parts of the plate tectonic cycle and have wide implications for the global environment (Berndt et al., 2019). Passive rifted margin studies have been at the core of the International Ocean Discovery Program (IODP) and its predecessors since the 1960s. Deep Sea Drilling Project (DSDP) and Ocean Drilling Program (ODP) drilling, along with extensive seismic surveying of the northeast Atlantic conjugate margins, demonstrated anomalously high volcanic activity during continental breakup. This led to the classification of these margins as “volcanic rifted margins” (Talwani and Eldholm, 1977; Eldholm et al., 1987, 1989; Saunders et al., 1998; Sengör and Burke, 1978; Ziegler and Cloetingh, 2004; Abdelmalak et al., 2016a, 2016b) (Figure F1).

Expedition 396 builds on previous successful drilling campaigns in the northeast Atlantic (DSDP Leg 38 [1974] and ODP Legs 104 [1985], 152 [1993], and 163 [1995]) that were instrumental in developing the concepts of volcanic rifted margins and large igneous provinces (LIPs) (see Mahoney and Coffin [1997], Ernst [2014], and Svensen et al. [2019] for a summary). There are, however, many unresolved scientific questions related to the formation and environmental implications of breakup volcanism that can only be addressed by scientific drilling.

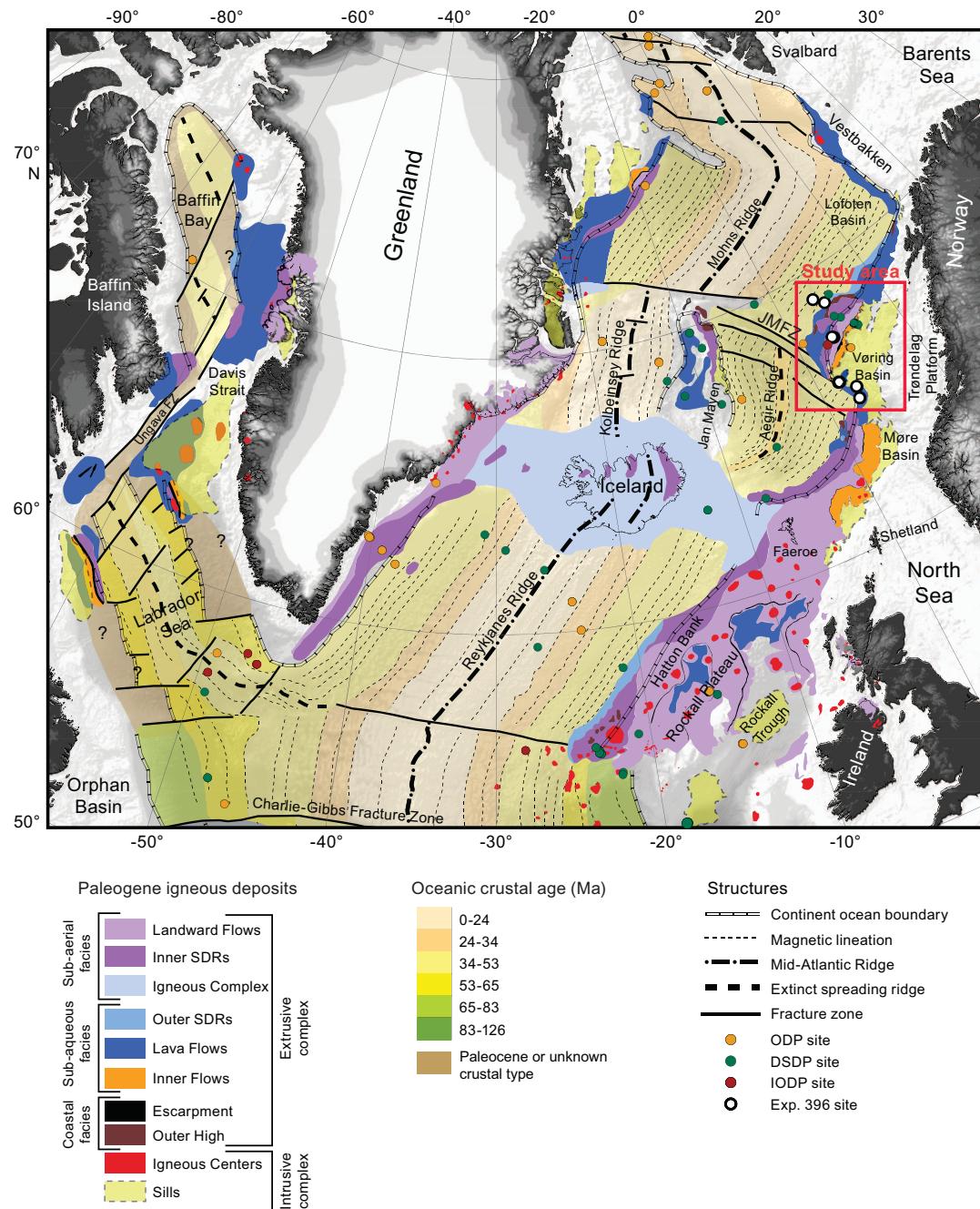
Despite unsurpassed constraints on conjugate crustal structure between the northeast Atlantic Norwegian, Jan Mayen, and Greenland rifted margins, the mechanisms behind rift-related, anomalous excess magmatic productivity are still debated (Lundin and Doré, 2005; Brown and Lesher, 2014; Foulger et al., 2020; Lu and Huismans, 2021). The controversy centers on three competing hypotheses:

- Excess magmatism derived from elevated mantle potential temperatures resulting from mantle plume processes,
- Small-scale convection at the base of the lithosphere enhancing the flux of material through the melt window during rifting and breakup, and
- Mantle source heterogeneity contributing to anomalously high melt production during continental breakup.

Although the mantle plume mechanism requires anomalous high temperatures resulting in high degrees of melting during asthenosphere upwelling, small-scale convection at the base of the lithosphere operates without involving significantly elevated potential temperatures and is inherently connected to the rifting process (Boutilier and Keen, 1999).

Temporal correlations between mass extinctions, global warming, and formation of LIPs have long been recognized but rarely demonstrated (only in the case of the late Permian extinctions) (Vogt, 1972; Eldholm and Thomas, 1993; Wignall, 2001). However, the reasons for the paleoenvironment-

tal crises are highly debated (Bond and Wignall, 2014; Courtillot and Renne, 2003). For example, volcanic eruptions release large volumes of sulfur, halogens, and carbon to the atmosphere (Jones et al., 2016), which may cause environmental disturbances on a variety of timescales. A potential complementary but perhaps asynchronous mechanism is that large volumes of greenhouse gases can be released from metamorphic aureoles around sill intrusions emplaced in sedimentary basins (Svensen et al., 2004). New information on eruption styles, volumes and rates, and sedimentological data in a proximal region to the eruptions is important to document and understand the environmental impact of LIP emplacement.



**Figure F1.** Distribution of Paleogene igneous breakup complexes and oceanic structures in the northeast and northwest Atlantic. Scientific boreholes and proposed drilling sites are shown. Compiled from Abdelmalak et al. (2016a, 2016b, 2017, 2019), Berndt et al. (2001b), Boldreel and Andersen (1994), Davison et al. (2010), Elliott and Parson (2008), Geissler et al. (2016), Reynolds et al. (2017), and Ritchie et al. (1999). JMFZ = Jan Mayen Fracture Zone, FZ = fracture zone, SDR = seaward-dipping reflector.

## 2. Background

Scientific drilling of northeast Atlantic continental margins since the 1970s has been essential for understanding the architecture and implications of igneous deposits associated with continental breakup. In particular, drilling of deep boreholes in the feather edge of the seaward-dipping reflectors (SDRs) offshore mid-Norway and southeast Greenland (Eldholm et al., 1989; Larsen, Saunders, Clift, et al., 1994; Duncan, Larsen, Allan, et al., 1996; Saunders et al., 1998; Larsen et al., 1999) demonstrated that voluminous subaerial volcanic sequences are common along rifted margins. The drilling results further suggested that continental breakup magmatism has had a major impact on the global environment and mass extinctions (Hinz, 1981; Eldholm and Thomas, 1993). Interpretation of industry seismic and borehole data from the Vøring Basin later led to the hypothesis that voluminous intrusion of magma in organic-rich sedimentary basin deposits may have triggered the Paleocene/Eocene Thermal Maximum (PETM) by release of aureole gases through hydrothermal vent complexes (Svensen et al., 2004). This hypothesis will be tested with the results of Expedition 396.

Breakup volcanism along rifted passive margins is highly variable in both time and space. Mantle melting during the formation of mid-oceanic ridges is relatively well understood and mostly a function of spreading rate and mantle potential temperature with melting below the mid-oceanic ridge. It leads to accretion of 6–8 km of magmatic crust at standard mantle potential temperature and full spreading rates larger than 2 cm/y (Bown and White, 1994). On the other hand, factors controlling magmatic activity during continental rifting and breakup are not well known. The variation in the degree of magmatism at rifted margins can, to the first order, be characterized in three contrasting modes of behavior (e.g., Huismans and Beaumont, 2011, 2014; Lu and Huismans, 2021). Mode 1 margins are characterized by a sharp transition from the continent/ocean boundary (COB) to normal thickness (6–8 km) magmatic ocean crust. At Mode 2 margins, magmatic productivity exceeds that expected from decompression melting at normal mantle temperature. Mode 3 margins have little to no magmatism at the COB and a broad transition zone with an exposed mantle at the seafloor preceding formation of mature oceanic crust.

A comprehensive understanding of what controls this range of behaviors and the volume, distribution, and timing of magmatism during continental rifting and breakup is, however, lacking. Excess magmatism at volcanic Mode 2 margins, such as in the northeast Atlantic, has been related to mantle plume and contrasting nonplume mechanisms (McKenzie and Bickle, 1988; Mutter et al., 1988; White and McKenzie, 1989).

Continental breakup may be associated with extensive volcanism over large distances along strike of the rifted margins, as exemplified in the northeast Atlantic (Figure F1). Magmatic products emplaced along these volcanic rifted margins have four major characteristics:

- Wedges of SDRs and associated volcanic seismic facies units interpreted to be massive subaerial and submarine Lava Flows and volcaniclastic sediments are found on both sides of the COB.
- Extensive sill and hydrothermal vent complexes are emplaced in organic-rich sedimentary basins around the incipient breakup axis.
- Thick high-seismic velocity bodies are found in the lower crust along the COB and commonly interpreted to be magmatic underplated material.
- Extensive pre- to synrifting volcanic sequences comprising subaerial lava flows, lava deltas, and volcaniclastic sediments.

Magmatic crustal thickening at these margins often exceeds 20 km, more than three times as thick as normal oceanic crust produced by passive upwelling of normal potential temperature mantle.

It appears that volcanic rifted margins require mantle that is either (1) anomalously hot, (2) actively upwelling at rates higher than the plate half-spreading rate, (3) anomalously fertile, or (4) some combination of these factors.

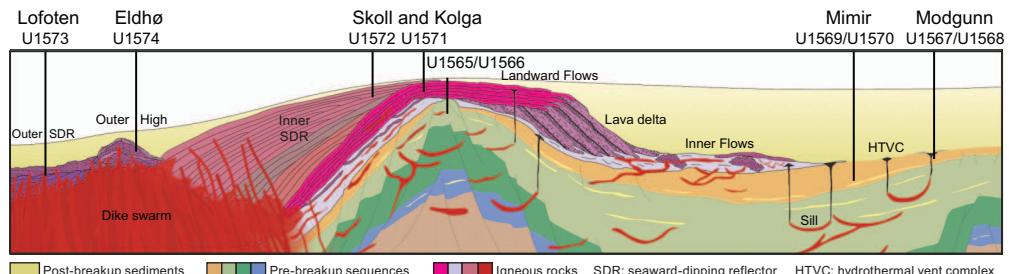
## 2.1. Geological setting

The northeast Atlantic rift system developed as a result of a series of rift episodes succeeding the Caledonian orogeny that ultimately led to continental breakup and passive margin formation in the earliest Eocene (Talwani and Eldholm, 1977; White and McKenzie, 1989; Skogseid et al., 2000; Abdelmalak et al., 2016a; Zastrophnov et al., 2020). The mid-Norwegian Margin is well covered by 2-D and 3-D reflection and refraction seismic surveys, potential field and heat flow data, and bore-hole data that allow a refined structural and stratigraphic framework (Figure F2) (Brekke, 2000; Gernigon et al., 2003, 2020, 2021; Mjelde et al., 2005; Breivik et al., 2006; Theissen-Krah et al., 2017; Zastrophnov et al., 2018, 2020; Polteau et al., 2020).

The mid-Norwegian Margin is segmented by the northwest-trending Jan Mayen Fracture Zone, which separates the Møre and Vøring Margins (Figure F3). The margin segments are characterized by different tectonomagmatic style and sediment distribution (Berndt et al., 2001a; Gernigon et al., 2020). The largest magmatic accumulation is observed in the Vøring segment, and volumes decrease to the south and north (Millett et al., 2020b). In the southern segment, passive margin formation and oceanic spreading was accommodated by the Aegir Ridge between the Møre and Jan Mayen (at the time connected to Greenland) conjugate margins in the Paleocene and Eocene.

Rifting and passive margin formation in the northeast Atlantic was accompanied by strong volcanic activity (White and McKenzie, 1989; Eldholm and Grue, 1994; Larsen and Saunders, 1998; Wright et al., 2012). Evidence for extensive magmatism is provided by SDRs, magmatic intrusions, and high-velocity bodies at the base of the continental crust underlying the COB, which in the distal margin are unequivocally interpreted to be magmatic underplate (Figures F2, F3) (Berndt et al., 2001a; Mjelde et al., 2005; Planke et al., 2005). ODP drilling of the Vøring Margin (Leg 104) and off southeast Greenland (Legs 152 and 163) recovered volcanic rock successions erupted during the initial stages of opening of the northeast Atlantic. Drilled rocks range from prebreakup continental tholeiitic flood basalt to synbreakup picrites to oceanic-type basalts that form the main part of the SDRs (Fitton et al., 2000). Oceanic-type lavas show an increasing degree of melting and contribution from asthenospheric mantle sources with time (Fram et al., 1998; Fitton et al., 1998). Thickness of igneous crust accreted at the southeast Greenland COB increases from about 18 km in the south to about 30 km near the Greenland-Iceland Rise (Holbrook et al., 2001). Similarly, geochemical enrichment of volcanic rocks of the East Greenland Margin (chondrite-normalized [Ce/Y]N and isotopes; Fitton et al., 1998; Tegner et al., 1998; Brown and Lesher, 2014) increases from south to north. In addition to scientific drilling results, evidence from the thick onshore Faroe Islands and East Greenland exposures reveal a temporal evolution that supports fluctuations in mantle potential temperature with time (e.g., Tegner et al., 1998; Millett et al., 2020a). Temporal fluctuations in mantle potential temperature are also invoked from independent geophysical observations, which reveal major uplift and subsidence events during the North Atlantic Igneous Province (NAIP) history that have been invoked as evidence for plume pulsing within the province (e.g. Champion et al., 2008; Hartley et al., 2011; Parnell-Turner et al., 2014).

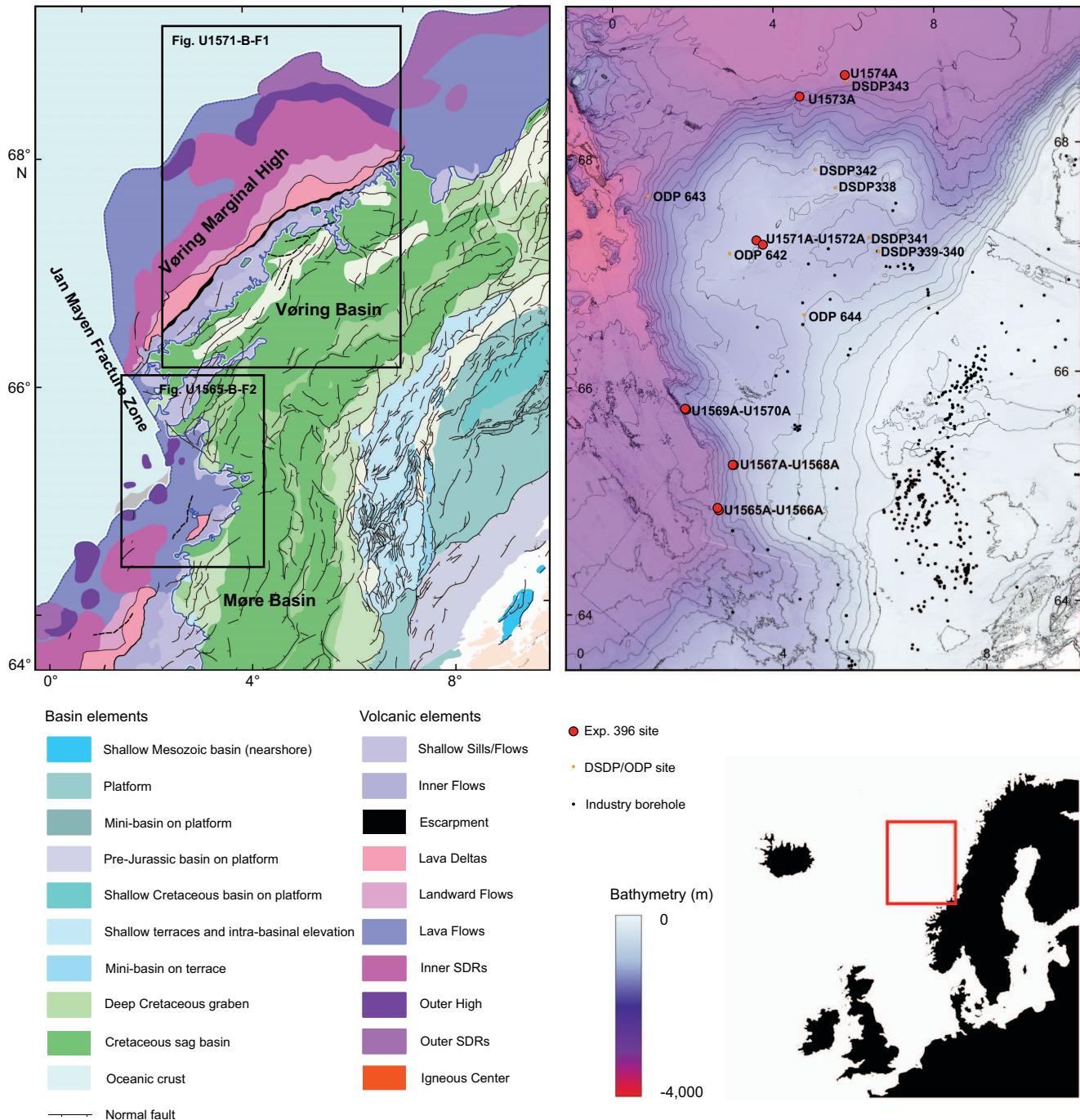
Correlation of crustal thickness and compositional enrichment suggests a combination of changes in source composition, source temperature, and/or melting dynamics (Hole and Millett, 2016; Hole and Natland, 2020, among others). It is not known whether a similar correlation of crustal



**Figure F2.** Schematic transect across a volcanic passive margin indicating the relative position of the volcanic seismic facies units and Expedition 396 site locations.

thicknesses and magma compositions exists along the Norwegian Margin. To establish the relationship between chemistry of the volcanic successions and crustal configuration is a milestone of the proposed investigations. Geochemical data show strong chemical and isotopic similarities between the “upper series” from the Vøring Plateau and southeast Greenland. In contrast, the “lower series” from both areas are fundamentally different from each other in many aspects, pointing to either substantial differences in prebreakup lithosphere composition at the two localities or different styles of mantle-crust interaction (Abdelmalak et al., 2016a).

Periods of elevated magmatism such as the emplacement of the NAIP often coincide with considerable environmental perturbations such as the PETM (56 Ma) and/or long-term climate warming



**Figure F3.** Bathymetric map and structural elements of the study area (after Gernigon et al., 2021). The Vøring transform margin is the landward extension of the Jan Mayen Fracture Zone.

such as the Early Eocene Climatic Optimum (EECO) (~50–53 Ma), suggesting a causal relationship (Bond and Wignall, 2014; Eldholm and Thomas, 1993). The total volume of magma emplaced during the Paleogene is estimated to be  $6–10 \times 10^6 \text{ km}^3$  (Saunders et al., 2007; Horni et al., 2017), with the most voluminous activity roughly coinciding with the Paleocene/Eocene boundary (Storey et al., 2007a), although the full emplacement spans several million years (Wilkinson et al., 2017). Greenhouse gas emissions were likely generated by magmatic degassing (Storey et al., 2007b; Gutjahr et al., 2017) and by explosive discharge of thermogenic gases generated by contact metamorphism (Svensen et al., 2004; Frieling et al., 2016; Aarnes et al., 2010). Therefore, the emplacement of the NAIP is one of the primary contenders for instigating hyperthermal events and long-term warming in the Paleogene, either as a direct forcing or as an instigator of positive climate feedbacks such as methane hydrate dissociation.

### 3. Scientific objectives

The key objective of Expedition 396 is to understand the relationship between rifting, excess magmatism, and paleoclimate and to resolve the relative contribution from plume upwelling, small-scale convection, and mantle heterogeneity and their relation to the formation of volcanic rifted margins in the northeast Atlantic. This requires additional constraints on the following:

- Melting conditions (degree, pressure, and temperature of melting);
- Age distribution of volcanic products, which is essential to constrain magmatic productivity in time and space;
- Variation of pre-, syn-, and postbreakup magmatic activity across the margin;
- Variation of magmatic activity along strike across the major Møre and Vørings Margin segments;
- Eruption rates, environment, and basalt morphologies; and
- The relationship between climate change and timing, volume, and style of magma emplacement.

The sedimentary proxy-based environmental reconstruction likely provides a semiquantitative record of paleoelevations (water depth) and vertical motions as early rifting progresses to seafloor spreading, with the potential additional influence of dynamic support originating from the plume-pulsing hypothesis (e.g., Champion et al., 2008; Parnell-Turner et al., 2014).

Sampling of breakup volcanic successions and proximal sediment cores through IODP drilling, in conjunction with geochemical constraints on melt conditions and integrated quantitative models of melting and mantle convection, is crucial to advance our understanding of breakup processes and resolve competing hypotheses for excess magmatism and paleoenvironmental consequences.

The mid-Norwegian Margin is a unique area in which well-characterized volcanic features, related to volcanic rifted margins worldwide, are readily drillable (Figure F2). Here, igneous rocks and Paleogene sediments are locally buried by minor postbreakup sediments. Large regions along the outer margins have recently been covered by industry 3-D seismic surveys, and this unique database allows the identification of shallow (<200 meters below seafloor [mbsf]) volcanic and Paleogene sedimentary targets.

There are seven main scientific objectives for Expedition 396:

**Objective 1 (primary):** determine the role of the Iceland plume in producing excess magma along the mid-Norwegian segment of the northeast Atlantic volcanic rifted margin during the Paleogene by constraining the conditions of melting (temperature, pressure, mantle sources, and total degree of melting).

**Objective 2 (primary):** determine the cause for along-axis variation in melt production.

In the case of the northeast Atlantic volcanic margins, magmatic productivity changes from the Møre Margin (~12–15 km thick magmatic crust) in the south toward the Vørings Margin (>20 km thick magmatic crust) and the Lofoten Margin (~8 km regular thickness magmatic crust). This pattern suggests a local, structural control because plume models would suggest largest excess

magmatic activity in the southernmost conjugate sections (e.g., Møre–Jan Mayen) closest to the Iceland thermal anomaly.

**Objective 3 (primary):** determine the depositional environment (subaerial vs. submarine) of Inner and Outer Lava Flows (e.g., SDRs) and implications for vertical motions during late synrift, breakup, and early postrift oceanic spreading.

Some of the Lava Flows may not have extruded subaerially (e.g., Planke et al., 2000, 2017). This has important implications for the distribution of buoyancy forces and isostasy during breakup where sections without continental crust would under normal conditions be expected at water depths >2 km (e.g., Kusznir et al., 2004).

**Objective 4 (primary):** determine the timing of magmatic activity and potential relationship(s) through analysis of a large variety of paleoenvironmental proxies proximal to the NAIP.

**Objective 5 (primary):** use the integrated paleoclimatic and paleoenvironmental proxies and geochronological data to assess the relative importance of volcanism and thermogenic release from hydrothermal vent complexes as potential drivers of climate change events.

**Objective 6 (secondary):** study early Eocene hothouse and freshwater incursions into the North Atlantic.

Scientific justification: the early Paleogene (~66–45 Ma) was characterized by warm global greenhouse conditions culminating in the EECO (~53–50 Ma), the warmest sustained climates of the last 65 My (Bijl et al., 2009; Anagnostou et al., 2016; Cramwinckel et al., 2018), inducing an intensified hydrological cycle with strongly increased precipitation at high latitudes (e.g., Pagani et al., 2006; Suan et al., 2017). Paleogene sediments obtained during Integrated Ocean Drilling Program Expedition 302 (Arctic Coring Expedition) show large quantities of free-floating freshwater azolla, which grew and reproduced in the Arctic Ocean by the onset of the middle Eocene (~48 Ma) (Brinkhuis et al., 2006). The drilled sites penetrate the Eocene sediments and will allow testing of the extent of freshwater incursions, constraining paleoceanographic boundary conditions for the excursions, including the evolution of oceanic gateways and their influence on global ocean circulation. The early Eocene sediments also provide a unique opportunity to reconstruct northern mid–high latitude climate, allowing a detailed comparison to Southern Ocean records (e.g., Bijl et al., 2009, 2013; Hollis et al., 2012).

**Objective 7 (secondary):** study carbon capture and storage in basalt provinces.

Industrial scale injection of CO<sub>2</sub> into volcanic reservoirs has been developed in recent years with test sites on Iceland (Matter et al., 2016) and in Washington State (USA) (McGrail et al., 2011). Carbon capture and storage is an integral component of meeting climate change targets of the Intergovernmental Panel on Climate Change (IPCC), and as such, the extensive offshore volcanic sequences of the mid-Norwegian volcanic rifted margin form potential storage targets (Planke et al., 2021). Better petrophysical and geophysical constraints on the volcanic rock sequences collected during Expedition 396 are needed to effectively appraise the reservoir properties and storage potential in this area.

Observations of meteoric water and high dissolved calcium concentration from the bottom of ODP Sites 642 and 643 show that the Vøring Plateau is ideal for studying circulation of freshwater in such large basaltic formations and assessing its potential for CO<sub>2</sub> sequestration. Previous ODP holes did not address the origin of meteoric water and trigger(s) for such large-scale circulation. Dating of borehole water samples with <sup>14</sup>C, <sup>36</sup>Cl, and <sup>234</sup>U/<sup>238</sup>U tracers and systematic analyses of fluid geochemistry (Inagaki et al., 2015) will allow the determination of the source of meteoric waters and provide constraints on their circulation systems, which are crucial for assessing the CO<sub>2</sub> storage potential of breakup basalts. Pore fluids have been sampled along the complete drilling transect to investigate the potential extent of meteoric water flow. The outcome of fluid geochemical analyses will be quantitatively interpreted with hydrological modeling using software such as MODFLOW to investigate the flow path of meteoric water. The sampling for water geochemistry will follow the standard IODP protocol for pore fluid analyses.

## 4. Operations and coring strategy

The mid-Norwegian Margin is the type locality for volcanic rifted margins and is probably the best studied volcanic margin worldwide. The detailed geometry and amount of volcanic products in the form of underplated bodies and intrusive and extrusive volcanic rocks are well constrained through geophysical imaging (e.g., Mjelde et al., 2005; Berndt et al., 2001b; Planke et al., 2017; Abdelmalak et al., 2016a, 2016b, 2017). However, new information on the age, nature, and depositional environment of the volcanic rocks is required to constrain melt production rates and vertical motions.

Answering the fundamental questions outlined in the scientific objectives required extensive sampling of both late synrift/early postrift sediments and magmatic products from the continental into the oceanic domain (Figures F2, F3). Collection of sedimentary records proximal to the volcanic and magmatic activity will allow to use numerous proxies (e.g., geochemical and sedimentological) to distinguish between volcanic and hydrothermal vent complex sources.

The expedition sites were located based on interpretation of very extensive high-quality modern 3-D data and locally 2-D profiles where 3-D data did not exist (Millett et al., 2020b; Gernigon et al., 2021). All proposed sites were located in regions with thin post-Eocene sediments. Site-specific high-resolution 2-D and 3-D P-Cable seismic data were subsequently acquired in 2020 across six of the approved sites (Bünz, 2020).

The expedition boreholes are located along three transects:

- Modgunn transect
  - Targeting the crater of hydrothermal vent complex
  - 5 boreholes (2 sites)
- Mimir transect
  - Targeting Paleogene sediments and the PETM
  - 5 boreholes (2 sites)
- Basement transect (Kolga, Skoll, Eldhø, and Lofoten Basin areas)
  - Targeting both basalt and interbasalt sediments, prebasalt, and postbasalt sequences
  - 11 boreholes (6 sites)

Drilling was highly successful, with clement weather, operations, and core recovery. In total, 21 holes were drilled at 10 different sites (Table T1). The total drilling depth was 3950 m, including more than 600 m of basement rocks. Overall recovery was 57%; however, the recovery of key PETM and basalt stratigraphic intervals was as high as 78% and 61%, respectively. A comprehensive suite of wireline logs were collected in eight boreholes, with a logged interval of 913 m. A summary of the drilling and logging results from the sites are presented in Figure F4 and located along a volcanic margins sketch in Figure F2. The coring results for each of the main recovered stratigraphic units is presented in Table T2, and a summary of the wireline logging intervals is presented in Table T3.

Paleogene sediments were drilled along the 1 and 3 km long Modgunn and Mimir transects (Figures F5, F6). Initially, the primary sites were drilled in both these areas. Initial results from the coring operations, including palynological analyses and physical property interpretations, were integrated with revised interpretation of the high-resolution P-Cable 3-D seismic data to refine the location of additional holes to optimize recovery of the targeted sedimentary strata (e.g., the PETM interval).

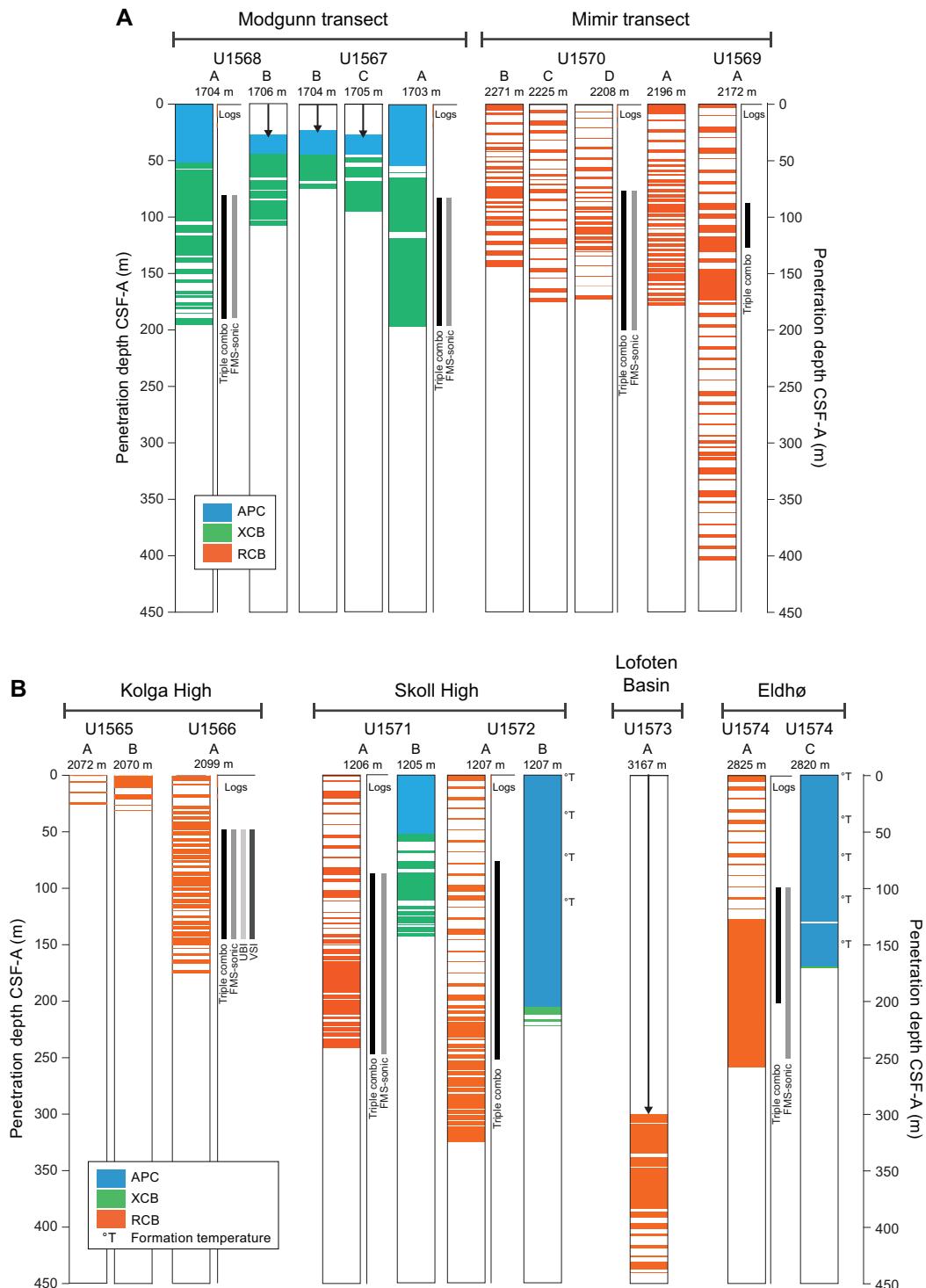
The drilling strategy for the basement transect was somewhat more complicated because several different areas along the margin were targeted (Figure F2). The two Kolga sites targeted subbasalt and initial basalt sequences, followed by the two Skoll sites, which targeted different volcanic seismic facies units and sequences near the feather-edge of the SDRs (Figure F7A). Farther north, the Eldhø (Figure F7B) and Lofoten Basin sites targeted the characteristic Outer High and Outer SDR volcanic seismic facies units, respectively. A summary of the drilling and logging results from the sites are presented in Figure F3.

All basement sites were initially drilled using the rotary core barrel (RCB) coring system for optimal basalt recovery (Figure F4). The main site objectives were typically 100 m of basement drilling or drilling to bit destruction. Half-coring was used extensively in the basement intervals to improve recovery. Wireline logs were acquired in most basement holes; the main objective was to image interbasalt sediments and fracture intervals that are difficult to recover with the RCB system. The advanced piston corer (APC)/extended core barrel (XCB) coring system was used at three basement sites (two at Skoll and Eldhø) to optimize recovery of the Eocene sedimentary rocks. APC coring was guided and constrained by results from the somewhat limited recovery in previously drilled RCB holes at each site. HLAPC coring was again used extensively to ensure high recovery in critical intervals. This approach gave an almost complete recovery of the postbasalt sediments at two sites.

The typical IODP shipboard workflow, including core description, imaging, physical property measurements, and geochemical analyses of rocks and fluids, was followed by the shipboard scientists and IODP staff. Headspace gas was routinely analyzed for safety reasons. Interstitial water (IW) and microbiological samples were also obtained soon after core recovery. Table T4 gives an overview of the comprehensive amount of data acquired in the targeted stratigraphic intervals during the cruise. Figures F8, F9, F10, F11, F12, and F13 highlight results from the different shipboard laboratories (biostratigraphy, core description, geochemistry, wireline logging, and paleomagnetism groups).

**Table T1.** Borehole location, penetration, and core recovery from the 21 holes drilled during Expedition 396. — = not applicable.

Hole	Longitude (WGS84)	Latitude (WGS84)	Water depth (m)	Penetration (m)	Drilled interval (m)	Cored interval (m)	Recovered length (m)	Recovery (%)	Total cores (N)	APC cores (N)	HLAPC cores (N)	XCB cores (N)	RCB cores (N)	Start date (2021)	Start time UTC (h)	End date (2021)	End time UTC (h)	Time on hole (days)
U1565A	2°44.989'E	64°57.761'N	2072	28.0	—	28.0	4.21	15.04	4	0	0	0	4	14 Aug	1600	15 Aug	1515	0.97
U1565B	2°44.845'E	64°57.804'N	2070	31.7	—	31.7	15.45	48.74	5	0	0	0	5	15 Aug	1515	16 Aug	215	0.46
	Site U1565 totals:				0	59.7	19.66	31.89	9	0	0	0	9					1.43
U1566A	2°43.753'E	64°57.887'N	2099	181.7	—	181.7	100.50	55.31	33	0	0	0	33	16 Aug	400	21 Aug	754	5.16
	Site U1566 totals:				0	181.7	100.50	55.31	33	0	0	0	33					5.16
U1567A	3°3.256'E	65°21.851'N	1703	195.9	—	195.9	181.58	92.69	23	6	1	16	0	21 Aug	1205	23 Aug	640	1.77
U1567B	3°3.208'E	65°21.768'N	1704	83.0	25	58.0	55.29	95.33	10	3	0	7	0	23 Aug	640	23 Aug	1850	0.51
U1567C	3°3.219'E	65°21.785'N	1705	106.0	30	76.0	66.48	87.47	13	2	0	11	0	26 Aug	2005	27 Aug	1340	0.73
	Site U1567 totals:				55	329.9	303.35	91.83	46	11	1	34	0					3.01
U1568A	3°3.109'E	65°21.594'N	1704	200.0	—	200.0	156.61	78.31	28	7	0	21	0	23 Aug	1936	25 Aug	2200	2.10
U1568B	3°3.154'E	65°21.663'N	1706	94.6	30	94.6	82.76	87.48	16	2	0	14	0	25 Aug	2201	26 Aug	1920	0.89
	Site U1568 totals:				30	294.6	239.37	82.90	44	9	0	35	0					2.99
U1569A	2°1.608'E	65°49.878'N	2171.1	404.0	—	404.6	144.91	35.82	44	0	0	0	44	27 Aug	1759	30 Aug	2020	3.10
	Site U1569 totals:				0	404.6	144.91	35.82	44	0	0	0	44					3.10
U1570A	1°59.623'E	65°49.890'N	2196.4	200.0	—	200.0	98.59	49.30	36	0	0	0	36	30 Aug	2020	1 Sep	1159	1.65
U1570B	1°57.361'E	65°49.679'N	2270	163.5	—	163.6	71.42	43.66	25	0	0	0	25	1 Sep	1159	3 Sep	100	1.54
U1570C	1°58.707'E	65°49.849'N	2225	200.0	—	200.0	49.59	24.80	23	0	0	0	23	3 Sep	100	4 Sep	030	0.98
U1570D	1°59.174'E	65°49.871'N	2207	200.0	—	200.0	50.49	25.25	28	0	0	0	28	4 Sep	030	5 Sep	2348	1.97
	Site U1570 totals:				0	763.6	270.09	35.75	112	0	0	0	112					6.14
U1571A	3°44.250'E	67°18.402'N	1205.7	247.6	—	247.6	116.55	47.07	39	0	0	0	39	6 Sep	952	10 Sep	730	3.9
U1571B	3°44.277'E	67°18.402'N	1205.3	143.7	—	143.7	112.51	78.30	20	7	0	13	0	10 Sep	730	11 Sep	1351	1.26
	Site U1571 totals:				0	391.3	229.06	62.685	59	7	0	13	39					5.16
U1572A	3°37.162'E	67°19.847'N	1206	330.5	—	330.5	139.53	42.22	46	0	0	0	46	11 Sep	1345	15 Sep	400	3.59
U1572B	3°37.057'E	67°19.918'N	1206	224.3	—	224.3	218.35	97.35	33	16	13	4	0	15 Sep	400	16 Sep	1612	5.10
	Site U1572 totals:				0	554.8	357.88	69.785	79	16	13	4	46					
U1573A	5°47.688'E	68°45.624'N	3167	440.9	300	140.9	100.97	71.66	18	0	0	0	18	17 Sep	153	21 Sep	705	4.22
	Site U1573 totals:				300	140.9	100.97	71.66	18	0	0	0	18					4.22
U1574A	4°38.436'E	68°36.011'N	2825	260.0	—	260.0	88.41	34.00	38	0	0	0	38	24 Sep	1200	30 Sep	520	5.72
U1574B	4°38.463'E	68°36.012'N	2825	9.5	—	9.5	9.72	102.32	1	1	0	0	0	30 Sep	520	30 Sep	1530	0.42
U1574C	4°38.463'E	68°36.001'N	2819	171.5	—	171.5	174.97	102.02	20	19	0	1	0	30 Sep	1530	2 Oct	231	1.46
	Site U1574 totals:				0	441.0	273.10	79.45	59	20	0	1	38					7.60
	Expedition 396 totals:				385	3562.1	2038.89	57.24	503	63	14	87	339					



**Figure F4.** Expedition 396 sites, recovery, and downhole measurements. A. Transects that addressed PETM causes and evolution at Modgunn Arch and Mimir High. B. Transects that addressed igneous basement objectives, including sampling of the subbasalt geology at Kolga High, the Inner SDR units at Skoll High, the Outer High at Eldhø, and the Outer SDR in the Lofoten Basin. FMS = Formation MicroScanner, APC = advanced piston corer, XCB = extended core barrel, RCB = rotary core barrel, UBI = Ultrasonic Borehole Imager, VSI = Versatile Seismic Imager.

## 5. Site summaries

### 5.1. Site U1565

Two short holes were drilled at Site U1565 to determine the nature of the subbasalt basement in the western part of the Kolga High. Both holes recovered altered granitic basement below ~25 m of Quaternary sediments.

**Table T2.** Core recovery for Expedition 396 target intervals.

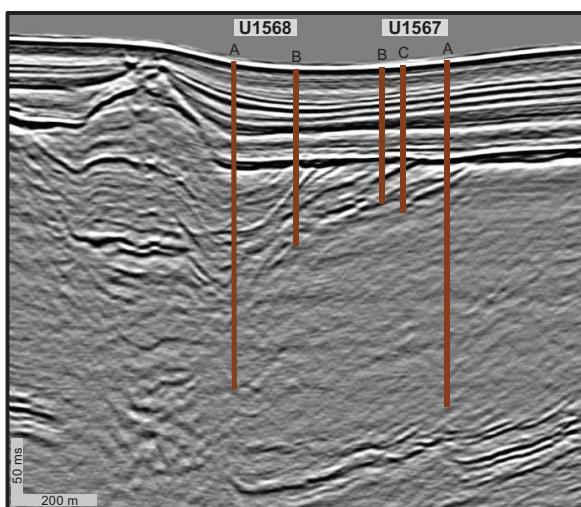
Target interval	Thickness (m)	Recovery (m)	Recovery (%)	Holes
Oligocene–Holocene	969	531	55	19
Eocene	1000	575	57	15
PETM	308	240	78	8
Paleocene	647	311	48	9
Basalt	587	357	61	8
Granite	44	15	34	3
Total cored:	3555	2028	57	21
Total drilled:	3950			

**Table T3.** Summary of downhole measurements, Expedition 396. TD = total depth, DSF = drilling depth below seafloor, WMSF = wireline log matched depth below seafloor, FMS = Formation MicroScanner, UBI = Ultrasonic Borehole Imager, VSI = Versatile Seismic Imager, APC temperature = advanced piston corer temperature tool.

Hole	Water depth (m)	TD (m)	DSF (m)	Pipe depth WMSF (m)	Top basalt WMSF (m)	Top depth	Triple combo (m)			FMS-sonic (m)			UBI (m)				
							Bottom depth	Total thickness	Basalt thickness	Top depth	Bottom depth	Total thickness	Basalt thickness	Top depth	Bottom depth	Total thickness	Basalt thickness
U1566A	2099	182	46	24	46	140	95	94	57	140	83	83	57	140	83	83	
U1567A	1703	196	84		84	195	111		76	197	122						
U1568A	1704	200	82		82	188	106		82	186	104						
U1569A	2171	405	82		84	128	44										
U1570D	2208	200	72		72	196	124		72	195	123						
U1571A	1206	248	87	117	86	247	161	130	86	248	162	131	129	237	109	109	
U1572A	1207	331	87	209	87	250	163	41									
U1574A	2825	260	105	171	104	213	109	42	147	248	101	77					
Total logged interval:							913	307									
Logged sediment interval:							606				695	291				192	192

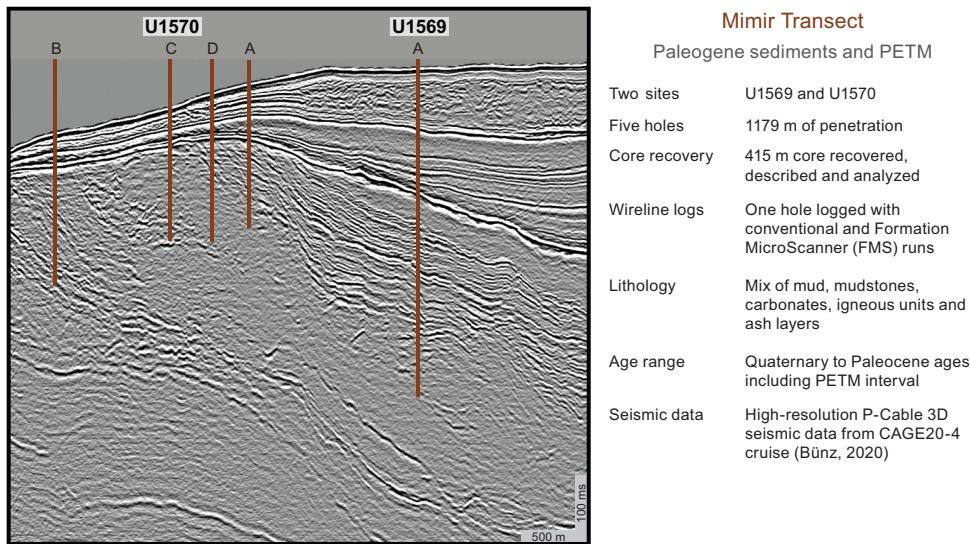
Wireline logs. VSI: 17–137 m (11 stations) in Hole U1566A

APC temperature 2 wells 7 temps

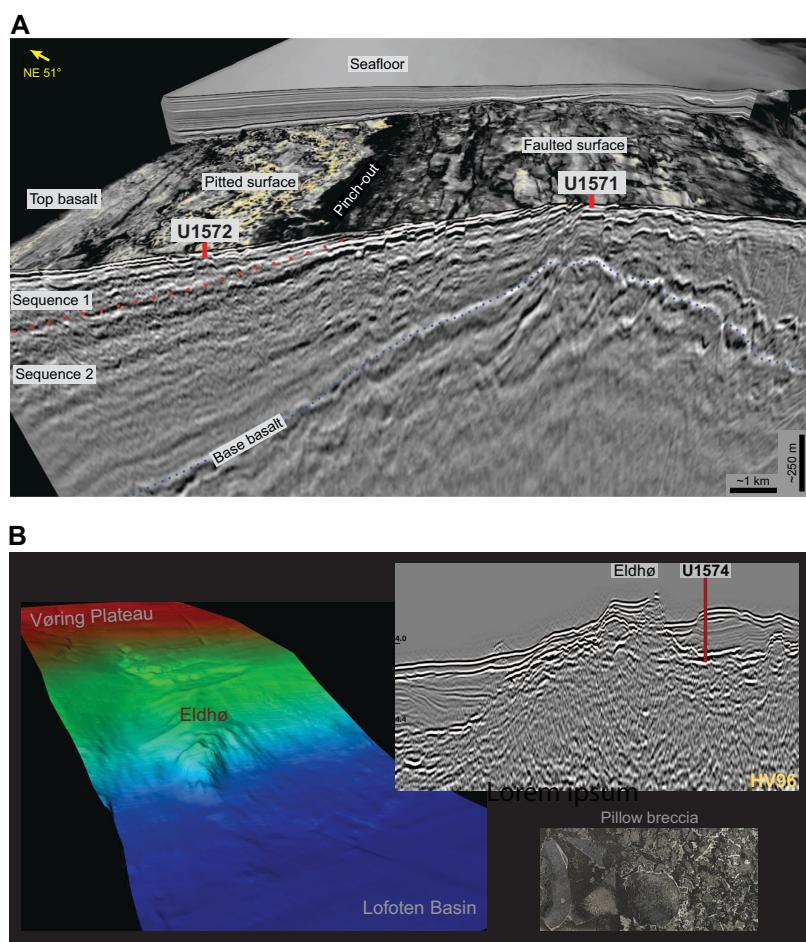


Modgunn Transect	
Upper part of a hydrothermal vent complex	
Two sites	U1567 and U1568
Five holes	85 to 200 m deep (total 710 m)
Core recovery	543 m core recovered, described and analyzed
Wireline logs	Two holes logged; conventional and Formation MicroScanner (FMS) runs
Lithologies	Dominantly mud, mudstones, and ash layers
Age range	Quaternary to middle Paleocene age, including potential PETM interval
Seismic data	High-resolution P-Cable 3D seismic data from CAGE20-4 cruise (Bünz, 2020)

**Figure F5.** High-resolution 3-D seismic data showing the position of Site U1567 and U1568 holes along the Modgunn Transect across the upper part of a hydrothermal vent complex.



**Figure F6.** Arbitrary high-resolution 3-D seismic line along the Mimir Transect on the Vøring transform margin showing the relative location of Site U1569 and U1570 boreholes.



**Figure F7.** Basement transect, Expedition 396. A. 3-D seismic view, Sites U1571 and U1572. Note the distinct difference in top basalt morphology between the two sites: smooth but faulted at Site U1571 and pitted at Site U1572. B. Site U1574 on the southern flank of Eldhø recovered hyaloclastites and pillow basalts below Eocene sediments. Seismic data from the Norwegian Petroleum Directorate (HV96 survey) and 3-D view of bathymetry viewed towards the south are included. Depths: Lofoten Basin = ~3200 m, Vøring Plateau = ~1500 m.

### 5.1.1. Background and objectives

Site U1565 (proposed Site VMVM-20A) is located at 2071 m below sea level (mbsl) on the western flank of the Kolga High, a pronounced structural high in the Jan Mayen Corridor (Figure F14). The site was located using a combination of conventional 3-D and high-resolution seismic data. Comprehensive 3-D seismic mapping shows that the Kolga High is covered by a variable-thickness Paleogene basalt unit overlying a complex and previously untested subbasalt sequence that had been interpreted as potentially sedimentary because of the presence of local intrusions of Paleogene sills and dikes (Millett et al., 2020b; Zastrozhnov et al., 2020). SDR wedges exist on the Møre Marginal High toward the south, and similar, but smaller, seismic reflections lap onto the western part of the Kolga High. A classic lava delta and escarpment defines the eastern limit of the high. The basalt cover is absent in two small ( $5 \times 3$  km sized) areas just west of the summit of the Kolga High (Millett et al., 2020b). Here, the seismic data show a transparent seismic basement that extends down to a strong seismic reflection at about 7.5 s two-way traveltime (TWT). The gap in the basalt cover makes the Kolga High a unique place on the Vøring and Møre Margins where the subbasalt sequence can be sampled with shallow coring.

The primary objective of Site U1565 was to core and sample the subbasalt geology that is largely unconstrained along the mid-Norwegian Margin. Sampling the subbasalt succession would provide important constraints on the rifting history both in time and space as well as constraints on the vertical movements of this margin segment prior to breakup (Gernigon et al., 2015, 2020). We expected to sample the subbasalt rocks after penetrating about 25–30 m of Neogene sediments. The plan was to core for 2 days or to a maximum depth of 200 mbsf to recover representative samples of the basement rocks to characterize the lithologies and alternation. As we encountered altered granitic basement immediately below the soft sediment cover in Hole U1565A, we decided to drill a second hole approximately 140 m to the west to verify the lateral continuity of the basement rocks rather than spending time deepening the initial hole at a very low penetration rate.

### 5.1.2. Operations

Site U1565 consisted of two shallow holes. Hole U1565A was cored using the RCB system to 28.0 mbsf. Coring was stopped because of very slow penetration rates while drilling into granite that was sampled in Cores 396-U1565A-3R and 4R. With the objective to confirm the lateral extent of the granite, Hole U1565B was added to the plan. This hole was drilled 138 m west-northwest of the first hole, in the direction of the next site (U1566). Hole U1565B was cored to 31.7 mbsf and yielded similar core material as Hole U1565A. Operations were ended at 0215 h on 16 August

**Table T4.** Summary of the various Expedition 396 onboard measurements carried out in relation to the target intervals. WRMSL = Whole-Round Multisensor Logger, MAD = moisture and density, SHMSL = Section Half Multisensor Logger, CHNS = carbon-hydrogen-nitrogen-sulfur, ICP-AES = inductively coupled plasma-atomic emission spectrometry, XRD = X-ray diffraction, pXRF = portable X-ray fluorescence spectrometry, PMAG = paleomagnetism.

Shipboard data and number of measurements	Post-Eocene	Eocene	PETM interval	Paleocene	Basalt interval	Granite	Barren	Total
Core recovery (m)	501	432	240	311	356	15		1,854
Recovery (%)	52	43	78	48	61	34		52
P-wave caliper (Gantry)	898	892	226	392	2,225	82		4,715
P-wave (WRMSL)	12,409	13,763	3,990	5,456	85			35,703
Discrete samples: P-wave and MAD	443	185	49	23	93	12		805
Magnetic susceptibility (SHMSL + WRMSL)	38,805	42,722	17,491	22,192	25,379	834		147,423
GRA bulk density (WRMSL)	18,788	20,333	8,230	10,503	11,281	377		69,512
Thermal conductivity	124	89	9	2	54	5		283
Interstitial water (IW)	137	101	28	48	15			329
Carbonates and organics (coulometry and CHNS)	135	110	35	51	19			350
Headspace gas	39	31	20	22	7			119
ICP-AES (hard rock)	5	1			121	6		133
XRD	116	118	72	142	58			506
pXRF	24	41	5		399	17		486
Source rock analyzer (SRA)	13	21	31	25				90
Biostatigraphy	121	96	25	61			67	370
PMAG cubes (JR6 and Kappabridge)	97				83	3		183
PMAG sections (2.5 cm intervals measured)	1,226				281			389,577
Thin sections	14				36	3		53

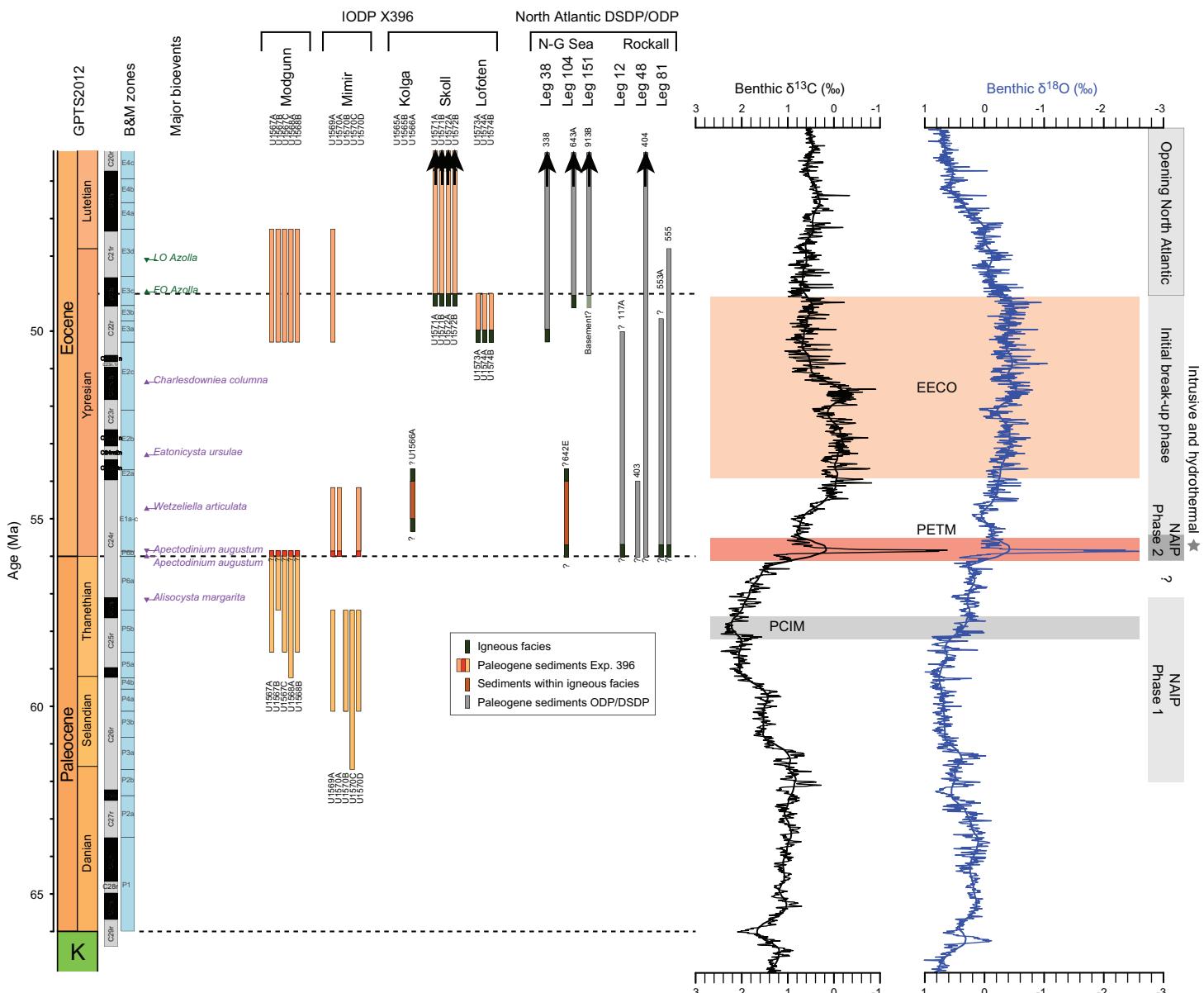
2021, and the drill string was pulled clear of the seafloor to begin the short transit of 900 m to the next site.

Nine cores were recorded for the site. The RCB system was used exclusively. It cored 59.7 m and recovered 19.66 m of core (32.9%). A total of 34.25 h or 1.4 days were spent on Site U1565.

### 5.1.3. Principal results

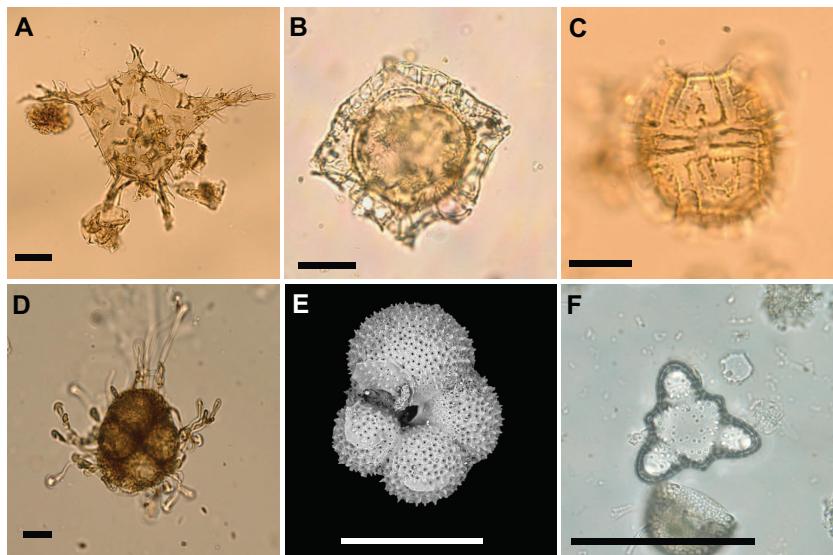
#### 5.1.3.1. Lithostratigraphy

The lithology of Holes U1565A and U1565B is divided into two lithostratigraphic units. Unit I consists of unconsolidated sediment, and Unit II consists of altered granite. The boundary between the two units was determined using the drilling data where the bit depth shows a sharp decrease in the rate of penetration from about 15 m/h to 1–2 m/h at 20.9 mbsf in U1565A and 28.0 mbsf in U1565B. The thickness of granite drilled in Holes U1565A and U1565B is 7.1 m and 3.6 m, respectively. Unit I is 7.1 m thinner in Hole U1565A than in Hole U1565B. High-resolution seismic

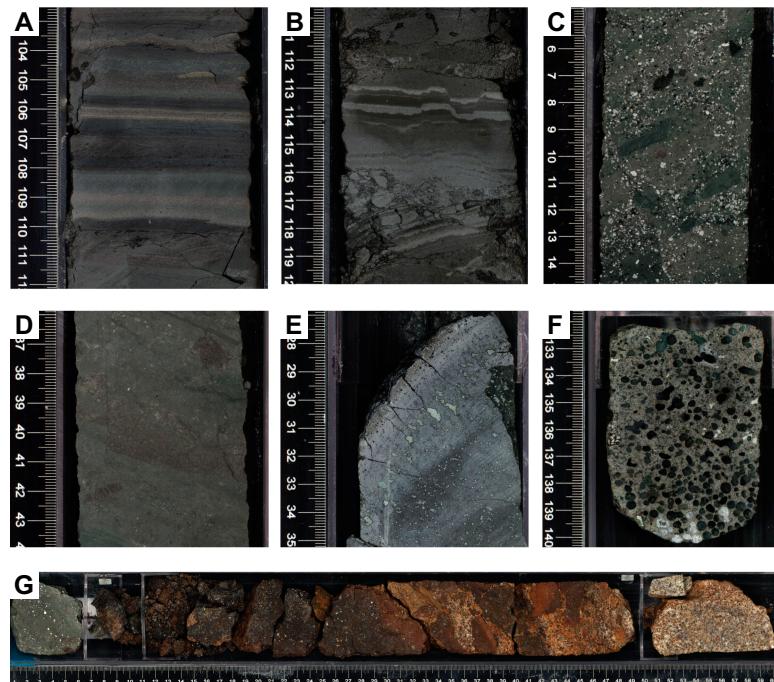


**Figure F8.** Age of cored intervals with respect to target intervals and existing borehole information for the northeast Atlantic margins. GPTS2012 = geomagnetic polarity timescale of Ogg (2012), B&M zones = zonation by Bujak and Mudge (1994), EECO = Early Eocene Climatic Optimum, PETM = Paleocene/Eocene Thermal Maximum, PCIM = Paleocene Carbon Isotope Maximum, NAIP = North Atlantic Igneous Province.

data show that the expected sediment thickness is the same in both holes. The thickness discrepancy in the drilling data is attributed to uncertainties in the determination of the seabed depth for RCB drilling because seabed depth was determined using the increase in torque in Hole U1565A and the mudline recovery in Hole U1565B.

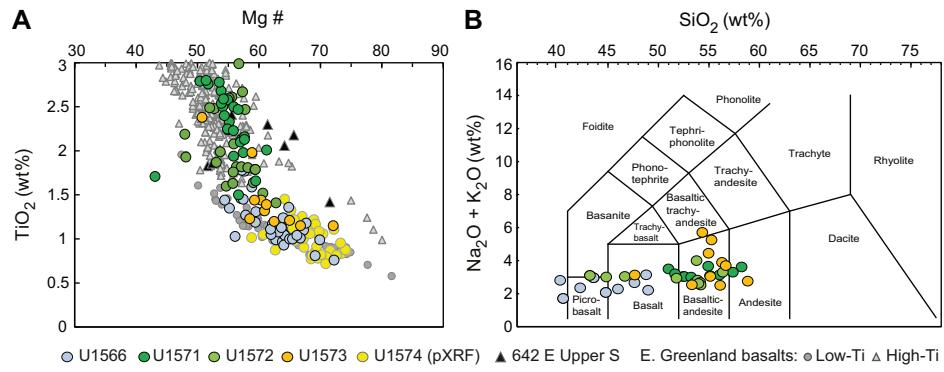


**Figure F9.** Selected microfossils encountered during Expedition 396. A. *Apectodinium augustum*, the dinoflagellate cyst marker taxon for the PETM age. B. Dinoflagellate cyst *Charlesdownia columna* (U1572B). C. Dinoflagellate cyst *Alisocysta margarita* (U1567A). D. Massulae of *Azolla* spp. (U1572A). E. Planktonic foraminifer *Acarinina medizzi* (U1574A). F. Diatom *Sheshukovia flos* (U1567A). Scale bar = ~25  $\mu$ m for A–D and F and ~200  $\mu$ m for E.

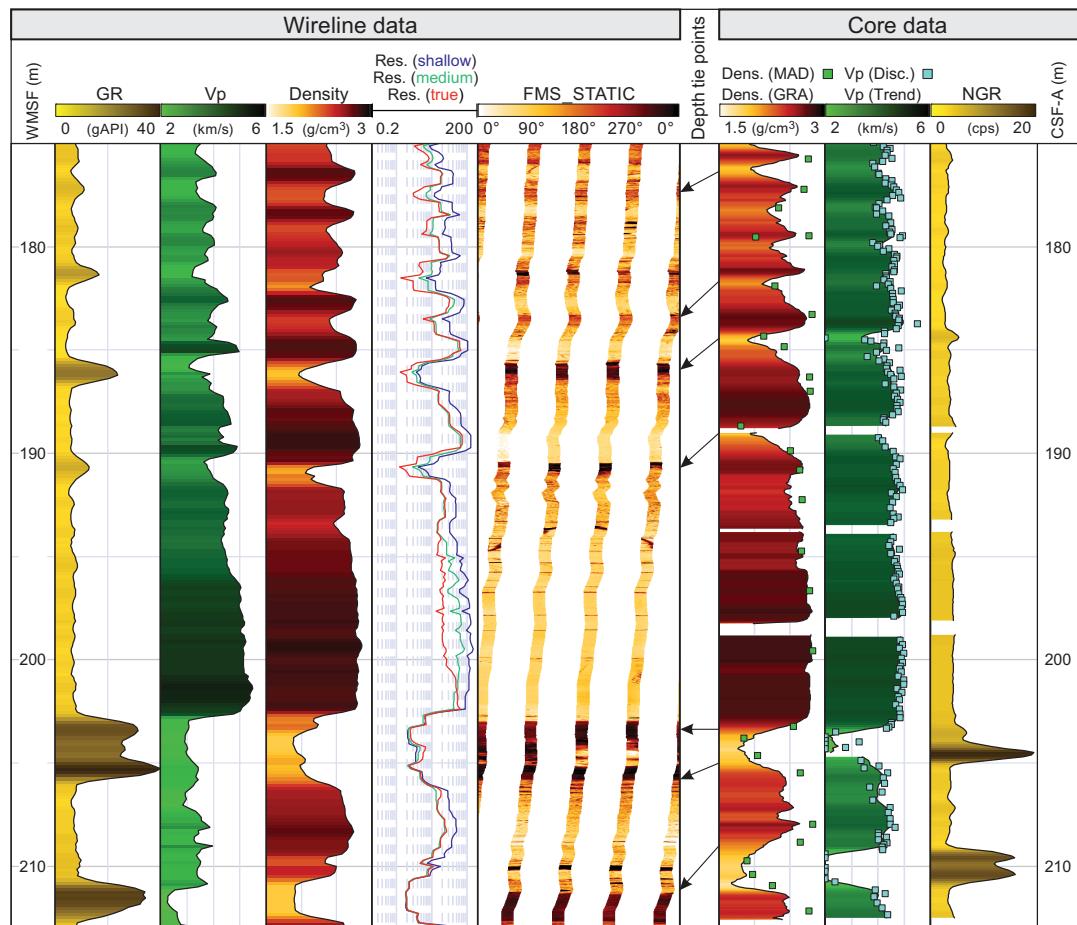


**Figure F10.** Typical lithologies recovered from Site U1566 at the Kolga High. Lithologies for Sites U1573, U1569, U1570, U1574, and U1566 are also shown. A. Dark gray organic-rich claystone with sand and thin parallel lamination (early Eocene; U1573A-7R-1). B. Very dark gray claystone with silt and dark gray to black ash (early Eocene; U1569A-36R-1). C. Dark green tuffaceous siltstone (earliest Eocene/PETM; U1570D-20R-CC). D. Dark greenish gray to very dark gray bioturbated claystone with silt (late Paleocene; U1570C-13R-1). E. Aphyric pillow basalt (early Eocene; U1574A-20R-1). F. Highly vesicular aphyric basalt (early Eocene; U1566A-27R-1). G. Contact between aphyric basalt (early Eocene) and granite with biotite (U1566A-29R-1).

Unit I sediments in Hole U1565A are primarily composed of olive brown to reddish-brown clay, whereas Unit I in Hole U1565B is mainly composed of brown, grayish brown to dark gray clay. Generally, this unit contains varying amounts of silt, sand, calcareous nannofossils and planktonic foraminifers, with diatoms common in the first core section of Hole U1565A. Amphibolite and

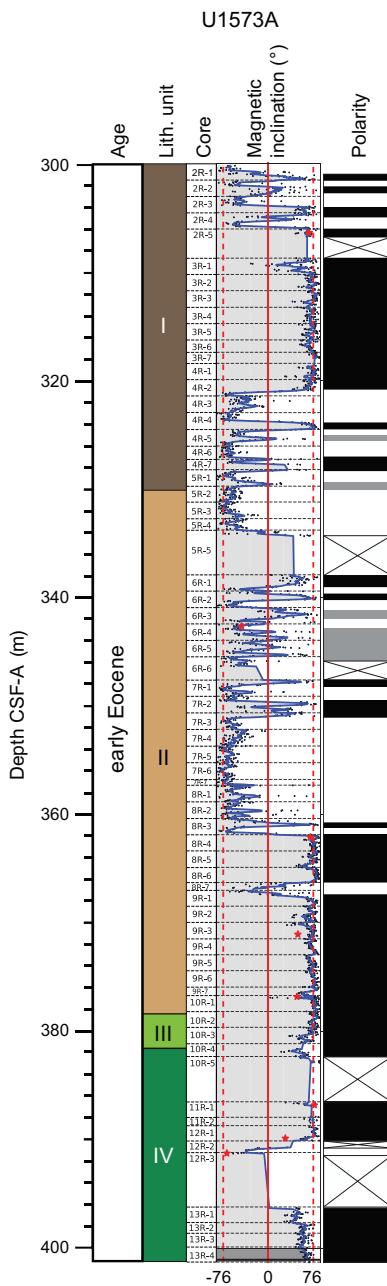


**Figure F11.** Geochemical composition of igneous rocks, Sites U1566 and U1571–U1573 (ICP-AES) and Site U1574 (pXRF). Compositions are compared with basalts recovered in Greenland (Tegner et al., 1998) and the upper series of ODP Hole 642E (Leg 104; Meyer et al., 2009). Note the affinity of Site U1566 and U1574 basalts with low-Ti basalts and the affinity of the Site U1571 and U1572 basalts with high-Ti basalts.



**Figure F12.** Downhole wireline log and discrete core-based physical property data, Hole U1571A. WMSF = wireline log matched depth below seafloor, GR = natural gamma ray logging tool, Res. = resistivity, FMS = Formation MicroScanner, Dens. = density, MAD = moisture and density, GRA = gamma ray attenuation, Disc. = discrete, NGR = natural gamma radiation, cps = counts/s.

granite dropstones also are observed in this core. The unit is overall moderately to very well sorted. Smear slide analyses indicate that Unit I consists of abundant undifferentiated clay-sized minerals, rare to trace foraminifers and nannofossils with larger silt and sand-sized grains comprised of common quartz and undifferentiated feldspar, and rare lithics. Trace amounts of diatoms, zircon, chlorite, and glauconite are also present. Slight to moderate drilling disturbances were observed in parts of the unit. The base of Unit I in Hole U1565B directly above Unit II consists of a mixture of sand, subrounded yellow brown clay clasts, and granules to medium pebbles with a poorly sorted grain size. Smear slide analysis of matrix material from this interval indicates abundant undifferentiated clay-sized minerals with rare nannofossils, rare organic debris, and trace diatoms. Silt and sand grains consist of common undifferentiated feldspar, rare quartz, and rare lithics. Dark gray polymetallic concentric nodules were observed at the base of Unit I in Hole U1565A (interval 396-U1565A-2R-1, 75–79 cm). PXRF data collected from nodules at ~5.80 mbsf (Section 2R-CC) show that their composition consists of low to moderate concentrations of tran-



**Figure F13.** Magnetic declination data for part of Hole U1573A.

sition metals (10.7 wt%  $\text{Fe}_2\text{O}_3$ , 2.6 wt%  $\text{Mn}_2\text{O}_3$ , 210 ppm Ni, 80 ppm Cu, and 94 ppm Zn). The silica content (~35 wt%) is considered relatively high for such nodules and suggests the presence of clay materials.

The transition between Units I and II was not retrieved in Hole U1565A. Based on the low penetration rates and drilling data, we cored 7.1 m into Unit II in Hole U1565A, recovering more than 3 m of granite. In Hole U1565B, the transition between the two units was recovered in Core 4R, and we cored 3.6 m into basement, recovering only 10 pebbles and a 10 cm core of highly altered granite conglomerate. Unit II represents one igneous lithologic unit consisting of moderately to highly altered medium-grained biotite granite with local mafic or felsic enclaves (Figure F10G). Throughout Unit II, the granite rocks remain mineralogically and texturally consistent. Discrete fine-grained enclaves are present in parts of Unit II. Observations of two thin sections confirmed the mineralogy and pervasive alteration observed at macroscopic scale. Alkali feldspars are generally partially or totally replaced by undetermined brownish clay minerals and zeolite. Biotite grains show corona of limonite.

#### 5.1.3.2. Biostratigraphy

Before hitting igneous materials, unconsolidated mixed gray and brown sands, silts, and clays with intervals containing gravel and larger clasts, interpreted as ice-rafted detritus, were recovered from Lithostratigraphic Unit I in Holes U1565A and U1565B. Clay- and silt-rich samples from Sections 396-U1565A-1R-CC through 2-CC and 396-U1565B-1R-CC through 4R-CC were analyzed for microfossil content. All sedimentary samples yield common-to-trace diatoms, in situ and reworked dinoflagellate cysts, reworked sporomorphs, rare ostracods, common-to-trace benthic foraminifers, and common planktonic foraminifers. The age-diagnostic microfossils indicate an overall Quaternary age. In addition, palynomorphs (mainly dinocysts) indicate reworking from Paleogene and (lower) Cretaceous strata, and pollen and spores also indicate reworking from lower Cretaceous rocks. Diatoms and foraminifers are common in the mudline samples and suggest a middle Pleistocene–Holocene age.

#### 5.1.3.3. Paleomagnetism

Lithostratigraphic Unit I sediments show consistent and continuous variations in inclination and intensity despite drilling disturbance. Some of the locally high values of magnetic inclination (>80°) might be due to drilling-induced magnetization. The sediments have an average intensity of  $\sim 3.6 \times 10^{-2}$  A/m, a magnetic susceptibility (MS) (K) of  $\sim 707 \times 10^{-6}$  SI, and a low magnetic anisotropy  $P'$  (~1.001). The nonhorizontal magnetic foliation suggests drilling disturbance. In Holes U1565A and U1565B, the geomagnetic polarity is normal, except at 5.58–5.78 mbsf in Hole U1565A, where a magnetic reversal may occur. The sediments have higher magnetic coercivity than the granites. One discrete sample shows limited demagnetization up to a field of 100 mT indicative of a high coercivity magnetic phase (goethite or hematite) and confirms the normal polarity of this interval.

Granites in Unit II have an average magnetic intensity of  $1.17 \times 10^{-2}$  A/m, likely due to loss of magnetite through alteration, a MS (K) of  $212 \times 10^{-6}$  SI, and a low degree of anisotropy  $P'$  (~1.001). Two discrete samples show a normal polarity and the presence of a low coercivity phase (magnetite, titanomagnetite, or maghemite) consistent with thermal remanent magnetization.

#### 5.1.3.4. Geochemistry

Six interstitial water samples were collected from the subbasalt sequence in Hole U1565B. Slight increases in dissolved Cl, Ca, Na, and Si are documented with depth, whereas Mg and Li decrease with depth. Little to no change is observed for K or Sr. Increasing Ca and decreasing Mg could reflect the diagenetic alteration of volcanoclastic material or seawater-basalt interactions. Downcore increases in alkalinity, Mn,  $\text{NH}_4^+$ , and  $\text{PO}_4^{3-}$  are paired with decreasing pH and may reflect microbially mediated oxidation of organic matter, although little change is observed for  $\text{SO}_4^{2-}$ .

Total carbon (TC), total inorganic carbon (TIC), calcium carbonate ( $\text{CaCO}_3$ ), total organic carbon (TOC), total nitrogen (TN), and sulfur were determined on six discrete sediment samples. Bulk sediment TC and TN were determined by elemental analysis, and TIC was determined by coulometer. TIC content ranges from 1.25 wt% in the upper section of the cored sediments (0.09–2.91 mbsf) to below 0.01 wt% with an average value of 0.69 wt%. Assuming TIC is purely representative

of  $\text{CaCO}_3$ , values range from <0.1 wt% to 10.4 wt%. TC varies between 0.2 and 1.6 wt% (average = 1.2 wt%). TOC is between 0.17 and 0.63 wt% (average = 0.33 wt%) with no clear trend downcore. TOC/TN ratios are between 3 and 25 with an average value of 9, representative of primarily marine organic matter. In all samples, S values are below detection limits.

#### 5.1.3.5. Physical properties

Physical property data acquired at Site U1565 include density, MS, *P*-wave velocity, natural gamma radiation (NGR), color reflectance, X-ray imaging, thermal conductivity, and spectral imaging. Gamma ray attenuation (GRA) bulk density of the sedimentary section varies from 1.4 to 1.79 g/cm<sup>3</sup>. Moisture and density (MAD) measurements collected for each section yielded very similar bulk density values (average = 1.69 g/cm<sup>3</sup>, n. 11), and measured grain density values vary from 2.63 to 2.79 g/cm<sup>3</sup> with an average of 2.72 g/cm<sup>3</sup>. Porosities range from 9% to 67%, reflecting the sediment's uncompact nature. MS measured on whole rounds and split sections are very similar and range from  $\sim 34 \times 10^{-5}$  SI to  $90 \times 10^{-5}$  SI. A MS peak ( $566 \times 10^{-5}$  SI) in Section 396-U1565A-1R-1 correlates to a dropstone found in the core.

Thermal conductivity measured at 40 cm in Section 396-U1565A-2R-1 yielded an average value of 1.42 W/(m·K). NGR values for the sedimentary unit vary from 27.7 to 38.1 counts/s with an average of 33 counts/s. A MAD sample collected at 61 cm in Section 2R-1 yielded a bulk density of 1.78 g/cm<sup>3</sup>, grain density of 2.7 g/cm<sup>3</sup>, and porosity of 55%. *P*-wave velocity measured at 60 cm in Section 2R-1 using the Section Half Measurement Gantry (SHMG) resulted in 1530 m/s in the *x*-direction and 1510 m/s in the *z*-direction.

Cores 396-U1565A-3R and 4R penetrated granitic basement. The average GRA bulk density of the hard rocks in these core sections is 2.275 g/cm<sup>3</sup>, yielding a bulk density of 2.56 cm<sup>3</sup> after a liner correction is applied, which agrees with the three MAD samples collected from these cores. MAD measurements revealed relatively high porosities of  $\sim 10\%$  in the granite cubes. MS values in the hard rocks are significantly lower than those of the overlying sediments (0 to  $22 \times 10^{-5}$  SI).

## 5.2. Site U1566

A complete section of the feather edge of an SDR sequence below a thin Quaternary sediment unit was recovered at Site U1566. The basalt flows and interbasalt sediment units were deposited in a subaerial environment and overlie the weathered granitic basement sampled at Site U1566.

#### 5.2.1. Background and objectives

Site U1566 (proposed Site VMVM-23A) is a companion site to Site U1565. Both sites are located on the western flank of the Kolga High (Figure F14) about 1 km apart at a water depth just above 2000 m. As described for Site U1565, the Kolga High is one of several areas along the rifted margin that are characterized by a thin sediment cover above the top basalt surface and a rapidly westward shoaling of the Base Cretaceous Unconformity (BCU) (Zastrozhnov et al., 2020).

Interpretation of extensive 3-D seismic reflection data shows that the Kolga High is covered by breakup related basalts except for two small windows (5 × 3 km sized) just west of the summit of the Kolga High (Millett et al., 2020b). Characteristic SDRs lap onto the Kolga High along its western margin, thinning toward the basalt windows where the lavas most likely have been eroded, making the Kolga High one of the few sites along the Vøring and Møre Margins where the transition from the lowermost extrusive volcanic rocks to the subbasalt strata is accessible by drilling. A previous attempt to determine the nature of this boundary was carried out during ODP Leg 104 in 1985 on the Vøring Marginal High. Hole 642E showed that the SDRs represent thick piles of subaerially emplaced basalt flows (Eldholm et al., 1989; Planke, 1994), but unfortunately did not get through the entire lava pile. The Norwegian Petroleum Directorate (NPD) drilled the eastern part of the Kolga High in 2014 (Hole 6403/1-U-1) and recovered 38 m of hyaloclastite with a tholeiitic composition (Bakke, 2017) in the upper part of a lava delta, as defined by eastward prograding reflections, but did not penetrate through the entire extrusive volcanic sequence.

The primary objective of Site U1566 was to sample the lowermost part of the SDRs on the Kolga High and, if possible, drill through the entire basalt sequence, but this was not the primary target because the nearby Site U1565 targeted the subbasalt units. By drilling Site U1566 we first aimed

to determine the emplacement environment of the initial volcanism on the Kolga High because it provides important constraints on the vertical movement of the margin prior to the onset of extrusive volcanism and the extension during the final rifting phase prior to breakup. A second objective at Site U1566 was to constrain the timing of the onset of extrusive volcanism in this area and to provide further insights into the hypothesis that breakup and extrusive volcanism on the Møre Margin predated breakup along the Vøring Margin (e.g., Gernigon et al., 2003). A third objective of Site U1566 was to characterize the magma source of the earliest extrusive basalts and determine whether it was primordial, which would imply rapid ascent through the crust, or it evolved during storage at crustal levels with potential incorporation of continental crustal rocks. A fourth objective was to constrain the rates of magma production both from geochemical analyses and potentially by biostratigraphic dating of interlayered sediments. The final objective of Site U1566 was to gather information on the extent of the basement lithologies drilled at Site U1565 and on lava–subsurface interaction if the bottom of the lava pile could be reached. Although Site U1566 was placed only 1 km away from Site U1565, the seismic data were inconclusive about the dip of the basal surface of the lava pile and hence the depth to this interface.

### 5.2.2. Operations

Site U1566 consisted of a single hole cored to 181.7 mbsf and wireline logging. Coring started at 0525 h on 16 August 2021 and ended at 2230 h on 19 August to accommodate the Versatile Seismic Imager (VSI) logging run in daylight hours on 20 August. A total of 33 cores were taken using the RCB system, recovering 100.5 m (55%) of core. Cores 396-U1566A-1R through 3R and 32R–33R were taken with full-length advances. Cores 4R (26.2 mbsf) through 31R (162.3 mbsf) were taken with half-length advances to improve core recovery of the basaltic basement.

All four logging tool strings, the triple combination (triple combo), VSI, Formation MicroScanner (FMS)-sonic tool, and Ultrasonic Borehole Imager (UBI) tool strings, were run to 145 mbsf in the hole. After the wireline logging concluded, the drill string was pulled back to the surface and secured, ending Hole U1566A and Site U1566 at 0230 h on 21 August. The time spent on Hole U1566A was 124 h or 5.2 days.

### 5.2.3. Principal results

#### 5.2.3.1. Lithostratigraphy

The succession recovered at Site U1566 is divided into three lithostratigraphic units. Unit I (0–19 mbsf) is unconsolidated clay with varying amounts of silt and sand as well as foraminifers. Some of the intervals contain clasts and nodules. Unit II (19–148 mbsf) is basalt and interbasaltic sediment. This unit is divided into 14 igneous subunits and 13 sedimentary subunits and has a total cored thickness of 121.8 m. Each igneous subunit is composed of one to several discrete lava flows. The dominant facies is aphyric basalt with rare microphenocrysts of plagioclase characterized by variable vesicularity (Figure F10). The chemostratigraphy of the igneous subunits obtained using pXRF shows significant and abrupt variations of the magma composition downhole. Unit III (148–175.4 mbsf) is weakly to moderately altered medium-grained granite with biotite alternating with decomposed granite and granite derived sandstones (Figure F10).

#### 5.2.3.2. Biostratigraphy

Sedimentary samples analyzed from the first three cores in Hole U1566A yield common to trace diatoms and other biosiliceous debris, trace *in situ* dinoflagellate cysts, sporomorphs, rare ostracods, common to trace benthic foraminifers, and common planktonic foraminifers. Age-diagnostic microfossils indicate an overall Quaternary age for the sedimentary deposits above the uppermost igneous facies.

In addition, a single sample taken from sediments interbedded between igneous rocks in Section 396-U1566A-11R-1 yielded a diverse terrestrial pollen and spore assemblage as well as a few aquatic elements. Age-diagnostic pollen encountered at this level suggest an earliest Eocene age for this sample.

#### 5.2.3.3. Paleomagnetism

Core sections were run through the superconducting rock magnetometer, which measured natural remanent magnetization (NRM) before and after alternating field (AF) demagnetization. Cores

396-U1566A-1R through 3R show normal polarity, followed by reversed polarity from Core 4R to the bottom of the hole, with some data points showing evidence of normal polarity. Lithostratigraphic Unit I has an average initial intensity of magnetization (NRM) of  $4.3 \times 10^{-2}$  A/m. The average NRM of Units II and III is  $\sim 1.4$  A/m, which is two orders of magnitude higher than that of the sediments.

Stepwise demagnetization was conducted on 32 discrete samples from Hole U1566A (thermal demagnetization) with a focus on unaltered basalt flow samples (30 samples). One sediment and one granite unit sample were also analyzed. The basalt sequence records mostly a reverse magnetic polarity.

#### 5.2.3.4. Geochemistry

Alkalinity increases from the mudline (2.3 mM) to the deepest IW sample taken at 18 mbsf (3.1 mM) in Hole U1566A. Measured pH is 7.8 in the mudline and declines to 7.7 at the base of the sediments.  $\text{PO}_4^{3-}$  content ranges from 4.04  $\mu\text{M}$  at 2.95 mbsf to 2.16  $\mu\text{M}$  at 18 mbsf. Ammonium ( $\text{NH}_4^+$ ) is below detection limit in the mudline but increases to 56.6 mM in the deepest sample.

The hard rock samples recovered from Hole U1566A (Unit II) are mostly subalkaline basalt, with one picro-basalt, and three samples from Unit III are dacite to ryolite. Comparison of the data obtained from ICP-AES with pXRF analysis with basalts from the upper series in ODP Hole 642E (Meyer et al., 2009) and flood basalts from East Greenland (Tegner et al., 1998) indicates an overlap of most samples from Site U1566 with the low-Ti basalts of Greenland (Figure F11). Another two samples from Unit III show compositions of a high-K granite (Rickwood, 1989), consistent with the high proportion of alkali feldspar, with a typical signature of an A-type granitoid (Frost et al., 2001).

TC, TIC,  $\text{CaCO}_3$ , TOC, TN, and sulfur were determined on three discrete sediment samples. TIC content ranges from 1.1 wt% at 2.95 mbsf to 0.02 wt% at 18 mbsf.  $\text{CaCO}_3$  values range from 0.2 to 11.3 wt%. TOC ranges from 0.21 to 0.03 wt%. All TOC/TN ratios are below 5, indicating marine organic matter. H and S were below detection limits.

#### 5.2.3.5. Physical properties

Physical properties were measured on all whole-round cores, archive and working-half sections, and discrete samples from Hole U1566A. Measurements included GRA bulk density, MS, NGR, *P*-wave velocity, X-ray imaging, MAD and porosity ( $n = 102$ ) analysis, and thermal conductivity ( $n = 30$ ).

Lithostratigraphic Unit I cores have *P*-wave velocities of 1489 to 2909 m/s and whole-round GRA bulk density from 1.51 to 2.10  $\text{g/cm}^3$ , and whole-round MS has a mean value of  $77.7 \times 10^{-5}$  SI compared to a mean of  $86.9 \times 10^{-5}$  SI from discrete half-round measurements. NGR values for the sediments vary from 19 to 44 counts/s. The bulk density of discrete samples ranges from 1.54 to 1.82  $\text{g/cm}^3$  ( $n = 6$ ), and grain densities range from 2.74 to 2.80  $\text{g/cm}^3$ .

The  $\sim 121$  m thick Unit II succession revealed dominantly very low gamma radiation counts ( $< 10$  counts/s) in the basalt subunits, whereas the interval volcanic sediment subunits show more variation and values up to 62.8 counts/s. Discrete samples have bulk densities ranging from 1.69 to 2.84  $\text{g/cm}^3$ . Discrete measurements of *P*-wave velocity range from 1530 to 5884 m/s and display systematic variations associated with the varying lithologies. In total, 23 successful thermal conductivity measurements were made across a range of lithologies in Unit II with values ranging from 1 to 1.9  $\text{W}/(\text{m}\cdot\text{K})$ . The MS values in the sequence (mean = 482 SI) are significantly higher than those of the overlying soft sediments (mean =  $78 \times 10^{-5}$  SI).

The granitic basement in Unit III has higher gamma radiation readings (mean 53.5 counts/s) compared to the overlying volcanic sequence, and discrete bulk densities range from 1.97 to 2.58  $\text{g/cm}^3$ . SHMG *P*-wave velocity measurements range from 1502 to 4501 m/s, and thermal conductivity measurements have a mean value of 2.1  $\text{W}/(\text{m}\cdot\text{K})$ .

#### 5.2.3.6. Downhole measurements

Wireline data from the four log runs comprise sonic velocities, MS, NGR, bulk density, resistivity measurements, and UBI and FMS image data of the basalt-sediment sequence of Lithostrati-

graphic Unit II. The lava package displays a consistently low background gamma radiation trend, with three clear peaks in the NGR log linked to siliciclastic sediment interlayers. The MS log reveals clear variations, with the basaltic Lava Flows with higher susceptibility values in the basalt flow subunits and almost zero MS of the interbasalt sediments. The porosity-dependent resistivity, bulk density, photoelectric factor, and *P*-wave and *S*-wave velocity logs show systematic variations throughout basalt subunits similar to logs collected in other subaerial basalt flow sequences in the province (e.g. Planke, 1994). For instance, the *P*-wave velocity is as low as 2.5 km/s in the interbasalt sediments subunits and flow tops and increases toward the massive basalt interiors, reaching values up to almost 6 km/s. The imaging logs (FMS-sonic and UBI) similarly display systematic variations in Lava Flow facies, with vesicular low-resistivity flow tops and massive resistive interiors with potential amygdales.

### 5.3. Sites U1567 and U1568

Sites U1567 and U1568 comprise a transect of five holes that sample the upper part of the Modgunn hydrothermal vent complex. The vent complex is truncated by an erosional unconformity of early Eocene age and overlain by 20–50 m of Eocene to recent sediments. It consists of a feeder system and a crater at the paleoseafloor. The crater was rapidly infilled during the PETM, presenting an expanded section that consists of diatom-bearing mudstones and ash layers.

#### 5.3.1. Background and objectives

A key objective of Expedition 396 was to constrain the role of breakup volcanism in the PETM. It has for a long time been argued that LIPs may have a dramatic impact on the Earth's environment (e.g., Eldholm and Thomas, 1993; Svensen et al., 2019) and likely caused mass extinctions (e.g., Ernst and Youbi, 2017). Explosive volcanism influences the climate by the eruption of large volumes of aerosols and ash into the stratosphere, affecting the climate over short time periods of days to years. Long-term climate effects of volcanism are mainly related to emission of sulfur dioxide, which may cause global cooling, and greenhouse gases such as carbon dioxide, which causes global warming.

Very large volumes of greenhouse gases must be emitted in short time periods to trigger global warming. One potential mechanism for the eruption of several thousands of gigatons of carbon-bearing gases is the maturation of sedimentary rocks by igneous intrusions offshore mid-Norway (Svensen et al. 2004). The key aspects of this hypothesis are that (1) magma injected as sill intrusions into organic-rich sedimentary rocks generates large volumes of greenhouse gases, namely methane and carbon dioxide, in metamorphic aureoles around the sills, and (2) pressure buildup due to gas generation causes fracturing of the host rocks, forming hydrothermal vent complexes where the gas and associated solids are transported from the deep sediment basins to the hydrosphere and atmosphere.

Although this process may release the amount of carbon required for the observed climate change, specifically between 2000 and 12000 Gt (Dickens et al., 1995; Dickens, 2011), it is not immediately clear if the bulk of the intrusions were emplaced over the geologically short time span that is required to, for example, explain the observed spike in global warming during the PETM. Also, it is unclear in which form the carbon would have been emitted, namely as CO<sub>2</sub> or CH<sub>4</sub>.

The hydrothermal systems associated with sill intrusions are manifested as pipe-like seismic anomalies that typically root at the tips of saucer-shaped sill intrusions (Planke et al., 2005). These so-called hydrothermal vent complexes are divided into a lower part (the pipe) and an upper part characterized by crater, eye, or dome-shaped seismic features at the paleoseafloor. The pipes are hundreds of meters high, sometimes up to 6 km. The upper part is typically 1–2 km in diameter but may vary in size from some hundreds of meters to >5 km. Many of the crater-shaped depressions are infilled by stratified seismic units and topped by characteristic mounds above the paleoseafloor. More than 1000 of these hydrothermal vent complexes have been mapped in sedimentary basins along the northeast Atlantic margins (Skogseid and Eldholm, 1989; Planke et al., 2005; Manton, 2015; Reynolds et al., 2017; Kjoberg et al., 2017; Omosanya et al., 2018; Roelofse et al., 2021).

Information on the nature of the pipes exists from cuttings from a single well in the Vøring Basin (6607/12-1; Svensen et al., 2003; Frieling et al., 2016) and from field analogs, such as in the Karoo Basin in South Africa (Jamtveit et al., 2004; Svensen et al., 2006) and the Tunguska Basin in Siberia (Svensen et al., 2009, 2018). The nature of the mound structures, however, is largely unconstrained. Their interpretation reaches from mud volcano–like sediment deposition to buildup of massive sulfide deposits that are formed at the hydrothermal vent sites by precipitation from hot vent fluids (Berndt et al., 2016).

Sites U1567 (proposed Site VMVM-31A) and U1568 (proposed Site VMVM-40B) are located on the Modgunn Arch (Figure F14) some 50 km north of the basalt flows sampled at the Kolga High on a dome structure that is interpreted to have formed in the Miocene (Gernigon et al., 2021). The sites were selected because the regional doming uplifted the vent sites to within reach of IODP drilling. Such shallow burial depths of vent complexes are currently only identified in the Modgunn Arch area in the Vøring Basin and in the Danmarkshavn Basin on the northeast Greenland shelf.

The seismic data clearly show the presence of sill intrusions at shallow depth, ~1 km below seafloor, on the Modgunn Arch. The sills are located in a deformed sedimentary unit with weak reflections that terminate against an angular unconformity (Figure F5). A hydrothermal vent complex is interpreted above the upper termination of the sill. The unconformity truncates the dipping reflections in the upper part of the hydrothermal vent complex. At the center of the depression, immediately above the pipe structure, the reflections bend upward, forming a mound. This mound is overlain by draped and onlapping reflections up toward the seafloor.

The aim of Sites U1567 and U1568 was to sample the upper part of the Modgunn hydrothermal vent complex and a nearby reference site along a 500 m long transect, avoiding the dome-shaped central part of the structure, to characterize the geology of the vent complex.

The first drilling objective was to constrain the processes that form the seafloor depressions and the infill. At least four potential scenarios are conceivable for this. First, pipe initiation may be explosive due to interaction of sill intrusions with pre-existing zones of overpressure that may lead to hydrofracturing of the host rock and rapid ascent of expanding gas. This will result in blowouts that can create craters with similar dimensions as the observed eye structure within days (Thatje et al., 1999; Leifer and Judd, 2015) and infill within a few hundred years. Second, they may be caused by long-term seepage of gas and aqueous fluids liquefying the surface sediments such as Lusi mud volcano in Indonesia (Mazzini et al., 2007). Third, they may be caused by repeated mobilization of mud by sills intruding into unconsolidated sediments leading to mud volcanism (e.g., Perez-Garcia et al., 2009) over a time span of thousands of years. Finally, they may result from venting of hydrothermal fluids and precipitation of pyrites and pyrotites over thousands of years (Berndt et al., 2016). Although the first two processes would require a later mobilization of the infill to form the observed domes at the upper termination, for example by gas-controlled seafloor doming (Koch et al., 2016), the domes in the latter two scenarios would be explained by the ejection of vent material.

Each of the four processes would be associated with very different types of material that was ejected during the time that the pipe was formed, both for the solid and the fluid phase. Therefore, understanding the formative processes for the eye structures is intrinsically linked to the second objective: to obtain information on the fluid composition and gas concentration (i.e., CO<sub>2</sub>, CH<sub>4</sub>, or higher hydrocarbons) and the amount of carbon release because this is crucial for understanding the potential role of hydrothermal venting as a driver of PETM climate change. This includes constraining the rates of fluid venting and its potential to reach far up into the water column and/or atmosphere depending on how violent the release was, if submarine venting occurred.

The third objective was to characterize the emplacement environment and age of the hydrothermal vent complexes and how it changed in response to venting and rapid warming during the PETM. This characterization includes the changes in microfossil assemblages before, during, and after the PETM. We also set out to determine whether there is evidence for environmental changes in the immediate vicinity of the proposed source of greenhouse gases.

The fourth objective of Sites U1567 and U1568 was to date the prominent unconformity at the Modgunn Arch and the timing of uplift of the eye structure because this constrains the vertical movement of the margin which is crucial input for melt generation models.

### 5.3.2. Operations

Operations at Site U1567 revolved around three holes. The first hole reached 195.9 mbsf. Three coring systems were used. The APC system was used from the seafloor to 52.9 mbsf. The half-length APC (HLAPC) was deployed for a single core between 52.9 and 54.1 mbsf. After reaching piston coring refusal, the XCB system was deployed for the remainder of the hole. Half-length advances of the XCB system were employed to 69.4 mbsf, and the remainder of the hole was cored using full-length advances. After coring operations ended in Hole U1567A, the hole was prepared for wireline logging and then successfully logged with the triple combo and FMS-sonic tool strings.

Hole U1567B was offset ~165 m west of Hole U1567A in the direction of Site U1568. It was drilled without recovery to 25 mbsf and then cored with the APC system to 49.3 mbsf, where we experienced a partial stroke of the APC system on Core 396-U1567B-4H. We then switched to the XCB system and continued coring with half-length advances to maximize core recovery. Hole U1567B was terminated at 83 mbsf after retrieving Core 11X. After coring concluded, the drill string was pulled back above the seafloor and secured for transit under dynamic positioning (DP) mode to Site U1568. The bit cleared the seafloor at 1850 h on 23 August 2021, ending Hole U1567B.

The 0.25 nmi move to Site U1568 was completed at 1936 h. With the thrusters already down, the move from Site U1567 to Site U1568 consumed only 45 min of vessel operations. Operations at Site U1568 consisted of two holes. Hole U1568A reached 200.0 mbsf. Two coring systems were used in the coring operations. The APC system was used from the seafloor to 51.7 mbsf. The final two cores, 396-U1568A-6H and 7H, were both incomplete strokes, and Core 7H advanced only 0.3 m into the formation. The XCB system was employed to advance the hole to the final depth of 200 mbsf. The hole was then prepared for logging, and we successfully conducted two passes with each the triple combo and FMS-sonic tool strings. With operations in Hole U1568A complete, the vessel was offset along the seismic line by approximately 165 m in the direction of Site U1567.

Hole U1568B was washed down to 30 mbsf. The APC system was deployed and advanced to 49.0 mbsf, where we experienced a partial stroke of the APC system on Core 396-U1568B-3H. The XCB system was deployed, and coring continued with half-length advances to maximize core recovery to 114.8 mbsf. Hole U1568B was terminated at 124.6 mbsf after retrieving Core 17X with a full-length advance. After coring concluded, the drill string was pulled back above the seafloor and secured for the DP transit back to Site U1567, thus ending Hole U1568B at 1900 h on 26 August. After completing operations at Site U1568 and with the APC/XCB drill string still deployed, the vessel moved in DP mode to reoccupy Site U1567. A total of 71.50 h or 3.0 days elapsed while conducting operations at Site U1568.

A total of 43 cores were recovered at Site U1568. The APC system was deployed nine times over a 70.7 m interval and recovered 72.03 m of core (102%). A total of 34 XCB cores were recovered over a 223.9 m interval. The XCB system recovered 75% of the interval cored or 167.34 m of core. The total interval cored at Site U1568 was 294.6 m; 239.37 m of core was recovered for an overall coring efficiency of 81%.

After completing coring operations at Site U1568, we set the end of the pipe at 1666 meters below rig floor (mbrf) and returned to Site U1567 in DP mode to core an additional hole between Holes U1567A and U1567B with the objective of improving the sampling across the Paleocene/Eocene boundary at this site. Once over the new hole coordinates, we spudded Hole U1567C at 2040 h on 26 August. The precision depth recorder seafloor depth was 1706.1 mbsl. We drilled to 30 mbsf without recovery and then APC/XCB cored to 106.0 mbsf (Cores 396-U1567C-1H through 14X). Coring operations were completed at 0820 h on 27 August, and the drill string was recovered back on board and secured for the sea passage to Site U1569. A total of 72.25 h or 3.0 days passed while conducting operations at Site U1567. A total of 46 cores were taken at this site. The total interval cored at Site U1567 was 329.9 m with 303.35 m of core recovered (92%).

### 5.3.3. Principal results

#### 5.3.3.1. Lithostratigraphy

Sites U1567 and U1568 are described together and consist of five holes spread along a 500 m transect of infilled sediments associated with cratering in the Modgunn hydrothermal vent complex. The succession recovered at the two sites is divided into six lithostratigraphic units, all of which are identified in each of the five holes and are preliminarily tied based on unit boundaries. Generally, core recovery was high in all the holes and the depths of most lithostratigraphic unit boundaries are well established. Unit I is gray clay with trace to rare dropstones. Unit II consists of dark brownish gray sediment (clay- to pebble-sized nodules) rich in manganese and iron in the uppermost interval and light yellowish brown sand-rich clay throughout. Unit III is pale yellow to very dark greenish gray clay with silt, including clasts of clay and rare beds of siltstone. Unit IV is dark greenish gray to very dark gray claystone to siltstone with common volcanic ash beds. Unit V is very dark gray to black parallel laminated claystone with common volcanic ash and lacking bioturbation. Unit VI is very dark greenish gray clay- to siltstone with moderate bioturbation.

#### 5.3.3.2. Biostratigraphy

The five holes cored at Sites U1567 and U1568 penetrated strata of predominantly early Paleogene age. A thin layer of Quaternary sediments was present in all holes in roughly equal thickness (~35 m) and was assigned a Quaternary age based on planktonic foraminifer content. No Neogene sediments were identified. Below this, mid-early Eocene (Zone E3) sediments were sampled in Holes U1567A, U1567B, and U1568A. Biostratigraphic and palynological analyses indicated earliest Eocene sediments (Zone E1) were recorded in Hole U1568A, but their presence was not confirmed by biostratigraphic analyses in Hole U1568B and Holes U1567A–U1567C. The PETM succession below is identified based on the presence of the dinocyst marker species *Apectodinium augustum* (P6b) and supported by diatom marker taxa (Figure F9). The extremely expanded PETM deposits sit below the early or earliest Eocene sediments and span between ~20 m in Hole U1567A and a virtually unprecedented 70–80 m in Hole U1568A, given the combined biostratigraphic and lithostratigraphic constraints.

Hole U1567B, and perhaps Hole U1568B, may contain a latest Paleocene sequence (Zone P6a), but this is not seen in any of the other holes, where the PETM succession rests unconformably on early late to middle Paleocene sedimentary successions.

#### 5.3.3.3. Paleomagnetism

The sediments at Sites U1567 and U1568 show low magnetic intensity (less than  $\sim 10^{-3}$  A/m) in ~40% of the recovered sections. In some intervals, both the magnetic intensity and the magnetic inclination angle vary greatly. The consistent distribution of these intervals with depth and across holes suggests that the low intensity is determined by the detrital source rather than representing a drilling artifact. In such intervals, magnetic polarity cannot be reliably determined. The magnetic inclination data after 20 mT AF demagnetization is typically used to determine magnetic polarity and to place the core sections in a magnetostratigraphic age framework. On the other hand, there are also intervals with moderate to high magnetic intensity (greater than  $\sim 10^{-3}$  A/m). These present both magnetic reversals and magnetic excursions, particularly in Holes U1567A (Core 4H), U1567B (Core 2H), and U1568A (Sections 4H-3 through 5H-3). These reversals, having similar inclination angles in normal and reverse polarity, pass the polarity test and therefore are interpreted as recording the geomagnetic field. The uppermost cores of each hole show a dominantly normal polarity and an inclination close to the calculated geomagnetic dipole field inclination (~77°) for these latitudes. These characteristics suggest that the paleomagnetic data of the youngest sediments accurately capture the geomagnetic field and its periodic reversals.

#### 5.3.3.4. Geochemistry

Almost 100 fluid and sediment samples were collected from the five holes along the transect between Holes U1567A and U1568A. For Hole U1568A, the site closest to the center of the Modgunn vent complex, several elemental profiles of IW diverge from the conservative behavior observed in most of the other holes. The first evidence for this can be seen in the transition between Lithostratigraphic Units I and II (36 mbsf), where some elements decrease significantly across the boundary. These include K (from 12 to 11 mM) and Mg (from >50 to <45 mM). Some of

the analyzed elements increase in concentration across this boundary, with the clearest rises observable in the profiles of Ca (from >12 to 13.5 mM), Sr (from 80 to 110 mM), and Li (from 30 to 40 mM). These shifts suggest a complex diagenetic environment, with evidence of marine silicate weathering (e.g., increasing Na and K) but also ash alteration (e.g., rising Ca and decreasing Mg). The elemental data for Unit IV in Hole U1568A corroborate the anomalous values of alkalinity and pH in these strata. Contents of B, Ca, Mg, Si, Sr, and Li are all diminished with respect to overlying and underlying sediments. Mn, Fe, and  $\text{PO}_4^{3-}$  covary in Hole U1568A, displaying clear enrichments in the uppermost section of the hole and much lower values below 50 mbsf. Increased concentrations of these constituents in the upper sections lend support for microbially mediated oxidation of organic matter.

### 5.3.3.5. Physical properties

Physical properties were measured for all cores from the five holes at Sites U1567 and U1568. Lithostratigraphic Unit I consists of up to 35 m thick Quaternary glacial and interglacial deposits with a mean bulk density of 1.81 g/cm<sup>3</sup> and porosities between 47% and 67%. The unit shows a relatively high gamma radioactivity with an across-borehole median value of 40 counts/s. *P*-wave velocities measured by the Whole-Round Multisensor Logger (WRMSL) range from 1480 to 2978 m/s with a median of 1572 m/s, and SHMG *P*-wave point measurements never exceed 1702 m/s, attesting to the soft and poorly compacted nature of the sediments. Unit II is generally less than 2 m thick and consists of Quaternary clay characterized by a rapid decrease in velocity and density with respect to Unit I. The underlying mid-early Eocene Unit III is characterized by a gradual increase in density (mean value = 1.71 g/cm<sup>3</sup>) with a corresponding increase in MS that increases most substantially toward the base of the unit, reaching maximum values in all five holes. The peak MS transition associated with the Unit III/IV boundary generally decreases laterally from Hole U1568A (maximum =  $1772 \times 10^{-5}$  SI; mean =  $451.3 \times 10^{-5}$  SI) to Hole U1567A (maximum =  $740.5 \times 10^{-5}$  SI; mean =  $225.3 \times 10^{-5}$  SI). The highest MS values above this horizon ( $>1700 \times 10^{-5}$  SI as measured by the Section Half Multisensor Logger [SHMSL]) occur in the two holes most proximal to the center of the vent complex (U1568A and U1568B). We suggest that the basal high MS layer (with values greater than  $1700 \times 10^{-5}$  SI) in the two proximal boreholes may pinch out prior to reaching the three more distal Site U1567 boreholes, where values  $\sim 700$ – $800 \times 10^{-5}$  SI are recorded.

The recovered thickness of Unit IV varies from ~6 m in Hole U1567A to more than 65 m in Hole U1568A. The recorded *P*-wave velocity (mean = 1922 m/s; median = 1903 m/s) is the highest amongst all units, associated with a relatively high mean bulk density of 1.74 g/cm<sup>3</sup> and a relatively low MS of  $90 \times 10^{-5}$  SI. The general trend of decreasing MS with depth appears to be linked to a reduction of ash layers with depth in Unit IV. Units V and VI reveal generally low and consistent *P*-wave velocities of around 1600 m/s with accompanying low bulk densities with a mean of 1.43 and 1.42 g/cm<sup>3</sup>, respectively. In summary, the cored sequences revealed physical properties that are consistent with a very weakly compacted mudstone-dominated sequence with fine-scale internal physical property variations linked to ash deposition.

### 5.3.3.6. Downhole measurements

Logging data were collected over a total open hole interval of ~112 m in Hole U1567A and ~105 m in U1568A, yielding high-quality data from both the modified triple combo and FMS-sonic tool strings. Four logging units for Hole U1567A and three for Hole U1568A were defined; they broadly correspond to the lithostratigraphic units. Velocities and densities are low with few anomalies in the reference hole, U1567A, around 1.55 km/s and 1.45 g/cm<sup>3</sup>, respectively. However, a well-defined NGR peak of 65 gAPI is clearly visible. The proximal vent hole, U1568A, shows a large variation of seismic and NGR properties, with a velocity range of 1.5 to 2.7 km/s, densities from 1.55 to 1.9 g/cm<sup>3</sup>, and NGR from 30 to 40 gAPI. Logging Unit 3 in Hole U1568A has an anomalously high photoelectric factor (indicating high average atomic number) at the top, which may indicate the presence of heavy elements such as Ba. FMS image log data from both holes nicely show fine-scale sedimentary structures and stratification.

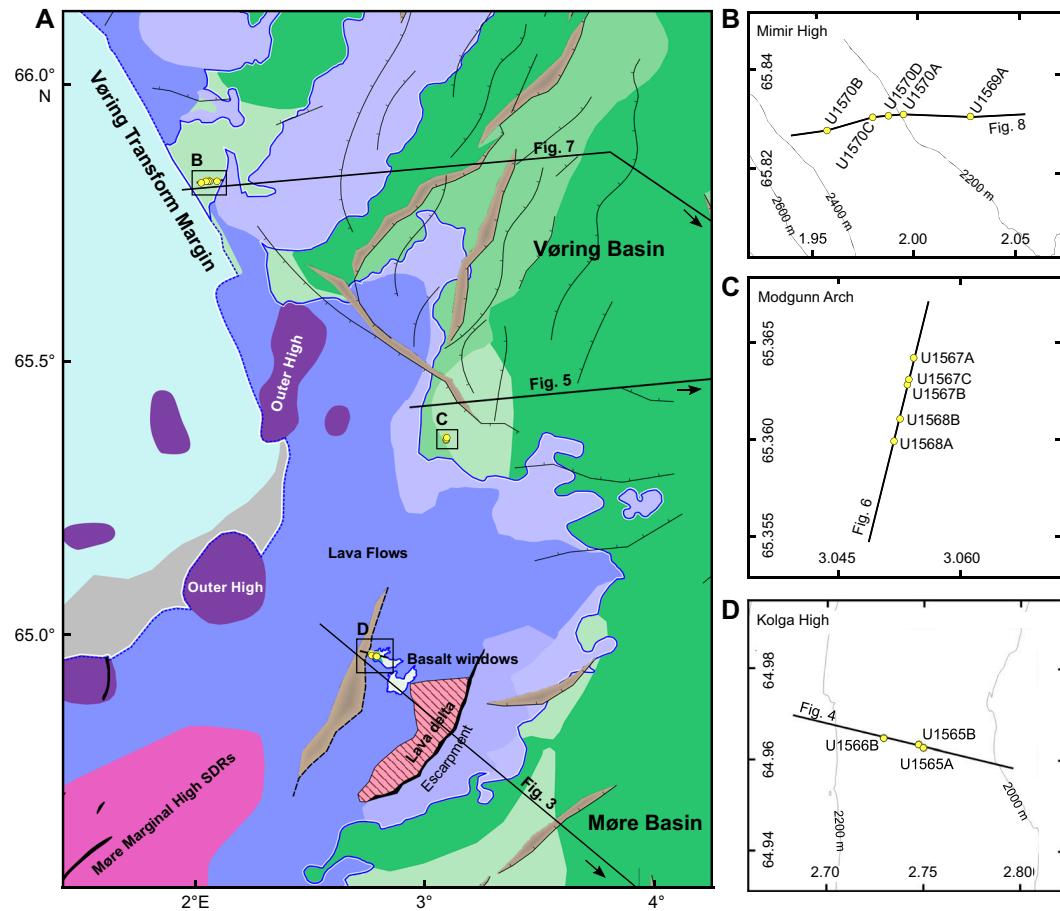
## 5.4. Sites U1569 and U1570

Sites U1569 and U1570 were drilled as a transect of five holes on the Mimir High, which is part of the Vøring transform margin. We recovered an expanded section of Paleocene and Eocene sediments including a thick interval of PETM deposits in three holes. The cores will allow us to constrain the climatic and environmental changes at high resolution and provide constraints on the biogeochemical changes influenced by hydrothermal venting. The expanded sedimentary sections show that the Vøring transform margin was a depocenter during Paleocene rifting and into the early Eocene until the marginal high formed during passage of the Aegir Ridge some 5 My after breakup.

### 5.4.1. Background and objectives

Site U1569 (proposed Site VMVM-55B) and the ribbon of four boreholes that comprise Site U1570 (proposed Site VMVM-56A) are located on the Mimir High at a 2171 to 2270 m water depth. The Mimir High is a transform margin high on the Vøring transform margin (Figure F14). Such highs are characteristic structural elements of transform margins world wide, such as offshore Ghana/Côte d'Ivoire, western Australia, and the Falklands (Basile and Allemand, 2002; Mercier de Lépinay et al., 2016; Baby et al., 2018; Lorenzo and Wessel, 1997). Transform margins are characterized by very narrow COBs and late uplift of the continental crust when the adjacent spreading center passes. Several processes are involved in the formation of the highs. They include lateral heat transport and associated thermal uplift as well as erosion-enhanced uplift (Vågnes, 1997; Basile and Allemand, 2002).

The Vøring transform margin is characterized by diminished breakup-related extrusive volcanism compared to the juxtaposed rifted margin segments of the Vøring and Møre Margins (Berndt et



**Figure F14.** Tectonic and magmatic elements of the Jan Mayen Corridor. Geology from Gernigon et al. (2021); legend in Figure F3.

al., 2001a). Nevertheless, thin extrusive basalts, intrusions, and a high-velocity lower crustal body have been documented for the Vørting transform Margin (Gernigon et al., 2021; Mjelde et al., 2009). The reduced magmatic crustal thickening may be the result of lateral heat transport from the rifted Møre Margin into the continental Vørting Basin to the north lowering the ambient mantle temperature and resulting in reduced melt supply (Berndt et al., 2001b). As a result, the basalt cover on the Vørting transform margin is generally thin and even absent on the Mimir High (Figure F14). However, an erosional truncation at the top of the high does not preclude that some basalt have originally existed and were subsequently eroded.

The absence of a basalt cover makes the Mimir High an exceptional place to study the stratigraphy at intervals coeval to and below the Paleogene basalt. These formations are not well known from drilling; the closest boreholes, 6504/5-15 and 6603/5-16, are more than 100 km away. However, the strata are subcropping along a kilometer-high escarpment on the southwest flank of the Mimir High. This escarpment was sampled during two cruises conducted in 2000 and 2013 using gravity corers, dredges, and an ROV (Polteau et al., 2020). The cruises documented the presence of Upper Cretaceous to Lower Eocene successions that include outcropping doleritic sills with columnar jointing. Although the recovered late Campanian and early Maastrichtian sediments were deposited in a coastal to shallow marine setting, the Paleocene samples indicate deep marine conditions with abrupt shoaling and formation of the Mimir High in the early Eocene. Toward the Vørting Basin, seismic data clearly show several large sill intrusions that affected this part of the margin and indicate that the strata generally dips from the Mimir High to the northeast (Figure F6).

The first objective of Sites U1569 and U1570 was to constrain the paleoenvironmental conditions during the Paleocene–Eocene transition. In particular, we hoped to obtain high-resolution information about the timing, duration, and nature of the environmental changes due to rapid warming and subsequent cooling during the earliest Eocene.

The second objective of Sites U1569 and U1570 was to tie the borehole stratigraphy with reflection seismic data to constrain the vertical movements of this margin segment as an input for geodynamic melt production models. In particular, the timing of uplift and any further constraints on the amount of crustal stretching are important.

The third objective of Sites U1569 and U1570 was to characterize the Paleogene sediments and their potential for releasing hydrocarbons when intruded by sills. Here, the key parameter is the amount of TOC in the sediments that may be transformed to greenhouse gases by contact metamorphism.

In the case that any igneous rocks were sampled, a further objective was to constrain the melting and emplacement conditions. This includes the composition of the magma and in particular the temperatures at which melting took place.

#### 5.4.2. Operations

The original plan for Site U1569 called for dual coring of key stratigraphic intervals, including a hole to 300 mbsf to be cored using the APC/XCB system and a second hole to 800 mbsf to be cored using the RCB system. In the case we could not properly sample the complete stratigraphic target at Site U1569, four offset holes of 200 m each were proposed along a transect of outcropping strata extending west of Site U1569, where deep stratigraphic targets were present at shallower subseafloor depths, as an alternative drilling strategy.

Hole U1569A was cored to the shortened depth of 404.6 mbsf after coring through the PETM, our scientific target, and penetrating 50 m into the late Paleocene. Estimating that the base of the Paleocene could be much deeper and to improve core recovery of the Paleocene/Eocene boundary, we ended coring in Hole U1569A in favor of sampling the older strata at much shallower depths at proposed alternate ribbon Site U1570. Coring the PETM strata at a shallower location also allowed us to sample less mature PETM sediments with better microfossil preservation.

After coring ended in Hole U1569A, we conditioned the hole and released the RCB bit at the bottom of Hole U1569A in preparation for wireline logging. The hole was then displaced with 125 bbl of 10.5 ppg mud, and the end of the pipe was set at 84.0 mbsf. The triple combo tool string was

rigged up and deployed to log Hole U1569A. The initial logging run failed to pass a bridge at 127.4 mbsf. With very little open hole to log, the tools were pulled back to the surface and logging operations were terminated. After logging concluded, the drill string was pulled back to the surface. The remaining top half of the mechanical bit release (MBR) cleared the rotary table at 2220 h, ending Hole U1569A and Site U1569. A total of 74.25 h or 3.1 days were recorded while on Site U1569.

A total of 44 RCB cores were retrieved from the site, five of which were half-length advances. The RCB system cored 404.6 m and recovered 144.91 m of core (36%). The coring plan at Site U1570 was the same for all four holes: to core to 200 mbsf and sample the Paleocene/Eocene boundary interval at considerably shallower depths than at Site U1569. Three of the four holes were completed to 200 mbsf, and the fourth one was terminated at 163.6 mbsf after we determined that the science objective for the hole had been met. All four holes were cored with the RCB coring system with a combination of half-length (4.6 m) and full-length (9.6 m) advances. The half-length advances were used to maximize core recovery in zones of interest at the expense of the additional wireline time. Coring in the final hole, U1570D, was ended after reaching the Environmental Protection and Safety Panel approved depth of 200 mbsf. After conditioning Hole U1570D for wireline logging, the RCB bit was released at the bottom of the hole and the end of pipe was set at 76.0 mbsf. The triple combo and FMS-sonic tool strings were successfully deployed to the full depth of the hole. After logging concluded, the drill string was pulled back to the surface. The bottom-hole assembly (BHA) was racked back in the derrick, and the outer core barrel was disassembled and inspected. The rig floor was secured at 2348 h, and the bridge was notified that operations were complete at Site U1570. A total of 112 RCB cores were taken at the site. The RCB system cored 763.6 m and recovered 270.09 m of core (35.4%). A total of 147.50 h or 6.2 days were recorded while on Site U1570.

### 5.4.3. Principal results

#### 5.4.3.1. Lithostratigraphy

Sites U1569 and U1570 consist of five holes on the northeastern flank of the Mimir High, a transform margin high. The holes sampled northeast dipping stratigraphy along a 5 km transect where Paleogene-age seismic reflections truncate below thin Quaternary sediments at ribbon Site U1570. The succession recovered at the two sites is divided into ten lithostratigraphic units. As expected from the seismic interpretation of the high-resolution 3-D seismic data, none of the five holes include all ten units. Unit boundaries are identified based on lithologic change and supported by physical properties and biostratigraphy. Where recovery is high, boundaries between units are well established. Where core recovery is low, unit boundaries are placed at the top of the underlying unit.

- Unit I is grayish brown and brown unconsolidated mud.
- Unit II consists of dark greenish gray consolidated mud which is locally slightly bioturbated.
- Unit III is grayish brown sand rich clay with nodules.
- Unit IV is very dark gray mudstone.
- Unit V is divided into two subunits. Subunit Va is very dark gray mudstone with sparse parallel lamination and slight localized bioturbation along with rare limestone intervals and common ash beds. Subunit Vb is very dark gray mudstone with sparse parallel lamination and slight localized bioturbation along with common ash beds and diagenetic pyrite.
- Unit VI is hypohyaline garnet dacite with cordierite, pyrite, and graphite.
- Unit VII is dark gray ash-rich mudstone with parallel lamination and common pyrite.
- Unit VIII is dark greenish gray to dark gray bioturbated mudstone.
- Unit IX is light gray limestone.
- Unit X is very dark greenish gray bioturbated mudstone with common ash along with some beds of limestone and volcaniclastic conglomerate.

#### 5.4.3.2. Biostratigraphy

Quaternary sediments from Holes U1569A and U1570A–U1570D (~10–80 m thick) are characterized by glacially influenced hemipelagic mud, which unconformably overlies lower to middle Miocene strata followed by uppermost Oligocene and lower Eocene strata. The lower Eocene deposits are thick (>100 m), and much of the lower Eocene is apparently missing. The Eocene

deposits include a thick PETM succession that is at least 10–40 m thick. The sequences from U1569A and U1570D, spanning the Paleocene/Eocene boundary, do not show any record of the presence of uppermost Paleocene sediments. Dinocyst chronobiostratigraphy suggests that at least ~1 My of the latest Paleocene is missing. The oldest sediments recovered are assigned to the middle Paleocene (Selandian; 61.6–59.2 Ma), recorded in samples from Holes U1569A and U1570C.

#### 5.4.3.3. Paleomagnetism

For the post-PETM strata (Lithostratigraphic Unit I, Subunits IIa and IIb, Units III and IV, Subunits Va and Vb, and Units V and VI), the average intensity of magnetization varies from  $2.26 \times 10^{-2}$  to  $3.22 \times 10^{-2}$  A/m. The median coercivity varies from 1.06 to 2.17, indicating the magnetic mineralogy is dominated by diagenetic hematite or goethite. There are six conspicuous episodes of magnetic reversals marked by magnetic inclinations  $< 0^\circ$ . The bulk MS of one discrete sample (396-U1569A-23R, 28–30 m) is over  $1 \times 10^{-3}$  SI, indicating a small ferromagnetic contribution, possibly from magnetite. This is also the only discrete sample presenting subhorizontal magnetic fabrics. The rest of the samples are paramagnetic, with a bulk MS below  $1 \times 10^{-3}$  SI, and have oblique magnetic fabrics, which may be owed to drilling-related disturbances. A low corrected degree of anisotropy ( $P_j < 1.028$ ) indicates the absence of micaceous minerals.

The average intensity of magnetization of PETM strata (Unit VIII) is lower than the post-PETM strata ( $0.06 \times 10^{-2}$  to  $0.13 \times 10^{-2}$  A/m). However, the median coercivity is higher (1.32–3.04), indicating higher hematite or goethite proportions. At least two short-lived episodes of normal polarity truncated the dominantly reverse magnetic polarity. All four discrete samples present oblique magnetic fabrics, low bulk MS ( $K_m \leq 0.4 \times 10^{-3}$  SI) and low corrected degree of anisotropy ( $P_j \leq 1.018$ ). Pre-PETM strata (Units VIII–X) have the highest coercivity (2.77–4.41) compared to the post-PETM and PETM units. High coercivity may be due to higher proportions of hematite and goethite. Their intensity of magnetization ( $0.06 \times 10^{-2}$  to  $0.59 \times 10^{-2}$  A/m) is lower than post-PETM strata but similar to PETM strata. All discrete samples have low bulk MS ( $K_m \leq 0.5 \times 10^{-3}$ ) and a low corrected degree of magnetic anisotropy ( $P_j \leq 1.077$ ). Only two discrete samples (396-U1570B-21R, 56–58 cm, and 25R, 92–94 cm) present subhorizontal magnetic fabrics. The rest have oblique fabrics, possibly due to drilling-related disturbances.

#### 5.4.3.4. Geochemistry

Geochemical data from the two Mimir High sites show substantial variability both downcore and between holes. In Hole U1569A, nonconservative behavior is observed for a range of measured IW variables, with periods of high pH and low alkalinity similar to those observed at Sites U1567 and U1568. This is also similar to the profiles observed in Hole U1570C and suggests lithologic changes control the IW composition. In Holes U1570A and U1570B, however, elemental distributions appear to be primarily conservative, and they suggest diffusion and advection are the primary controls on IW composition, potentially linked to the extensive fracturing of these rocks. In all holes, carbonate content is high in the upper units before dropping to lower values farther downcore. Occasional peaks in carbonate and organic carbon content suggest infrequent authigenic mineral precipitation. Lithostratigraphic Unit IV in Hole U1570D was investigated via ICP-AES, appears to be dacitic, and may represent shallow emplacement into the sediment package. XRD analysis indicates the primary sedimentary components are quartz, calcite, and a range of clays. In some samples, it was possible to identify clinoptilite as the primary clay mineral, a form of zeolite linked to ash weathering. The direct impact of volcanic ash deposition is present in the form of a range of pyroxene minerals.

#### 5.4.3.5. Physical properties

A total of 151 samples across the five holes cored at Sites U1569 and U1570 were measured for MAD, and 1344 *P*-wave velocity measurements were made on discrete samples to constrain and supplement physical property measurements conducted on the WRMSL and SHMSL. The results showed distinct petrophysical properties for the 10 lithostratigraphic units that were described at the sites. Similarly to Sites U1567 and U1568, peaks in MS are related to the presence of ash beds, such as in Lithostratigraphic Subunit Va. However, the magnetic signature of ash beds dissipates in Subunit Vb, and this characteristic is interpreted as evidence of increased alteration of the ashes. The dacitic Unit VI, which was sampled in Holes U1570A and U1570D, has noticeably high seis-

mic velocity, relatively high density, and low MS. However, the highest  $P$ -wave velocity of all five holes is recorded in Unit IX (mean = 4172.6 m/s), which represents the thin limestone bed found in Hole U1570B. This unit also has high density (GRA bulk density mean = 2.31 g/cm<sup>3</sup>) and the lowest porosity (MAD porosity = 19%) values.

#### 5.4.3.6. Downhole measurements

Wireline logging was conducted in Holes U1569A and U1570D, which extended to 404.6 m drilling depth below seafloor (DSF) and 200 m DSF, respectively. Two wireline logging tool strings, the triple combo and the FMS-sonic, were planned for Hole U1569A; however, an obstruction encountered at 127.4 m wireline log depth below seafloor (WSF) during the first deployment of the triple combo prevented us from completing the logging plan. In contrast, two logging runs were performed successfully in Hole U1570D with the triple combo and the FMS-sonic tool strings over a total open hole logged interval of ~117 m.

Six logging units were defined in Hole U1570D, constrained by a combination of the lithostratigraphy, physical properties, and wireline data. In the main open hole logged interval of Hole U1570D, the wireline properties reveal several transitions that correspond closely to the lithostratigraphic units.

Logging Unit 1 corresponds to an interval of the hole logged behind the pipe. It is best defined by slightly higher gamma radiation readings (~10 gAPI) from the seafloor to ~22 m wireline log matched depth below seafloor (WMSF) compared to the gamma radiation of the underlying Logging Unit 2 (~4 gAPI). This change corresponds well to the measured whole-round NGR values for the poorly recovered sediments in Lithostratigraphic Units I and II. The open hole part of Logging Unit 2 reveals gamma radiation of ~45–50 gAPI, low bulk density of ~1.6–1.7 g/cm<sup>3</sup>, and low consistent  $V_p$  of around ~1.65 km/s. Logging Unit 3 exhibits relatively high  $V_p$  (~1.8–2.3 km/s) and resistivity compared to its surrounding claystone-dominated sequence and corresponds closely to Lithostratigraphic Unit VI. Logging Unit 4 is characterized by relatively high gamma radiation, fluctuating between ~45 and 85 gAPI. Logging Unit 5 comprises a thin ~3 m thick interval with elevated density (~1.95 g/cm<sup>3</sup>) and  $V_p$  (2 km/s) and extreme photoelectric effect (PEF) enrichment up to ~33 b/e<sup>-</sup> coupled with a gamma radiation reduction. The remainder of the logged section is assigned to Logging Unit 6 and shows minor fluctuations in gamma radiation, density, and PEF.

### 5.5. Sites U1571 and U1572

Sites U1571 and U1572 sampled the upper termination of the Inner SDRs on the central part of the Vøring Margin. At both sites, two holes were drilled: the first using the RCB to sample the basaltic basement and the second one using the APC to sample the overlying sediment package. The reason for drilling two sites only 5 km apart was the starkly different seismic character of the top basalt surface that we interpreted to be the boundary between different emplacement environments, with the upper sequence being emplaced in a wet or an aquatic environment. The sediments are of late early Eocene age, which could be consistent with late emplacement of the last flood basalts on the Vøring Marginal High. There is, however, a distinct possibility that the overlying sediments were not deposited immediately after the lava flows.

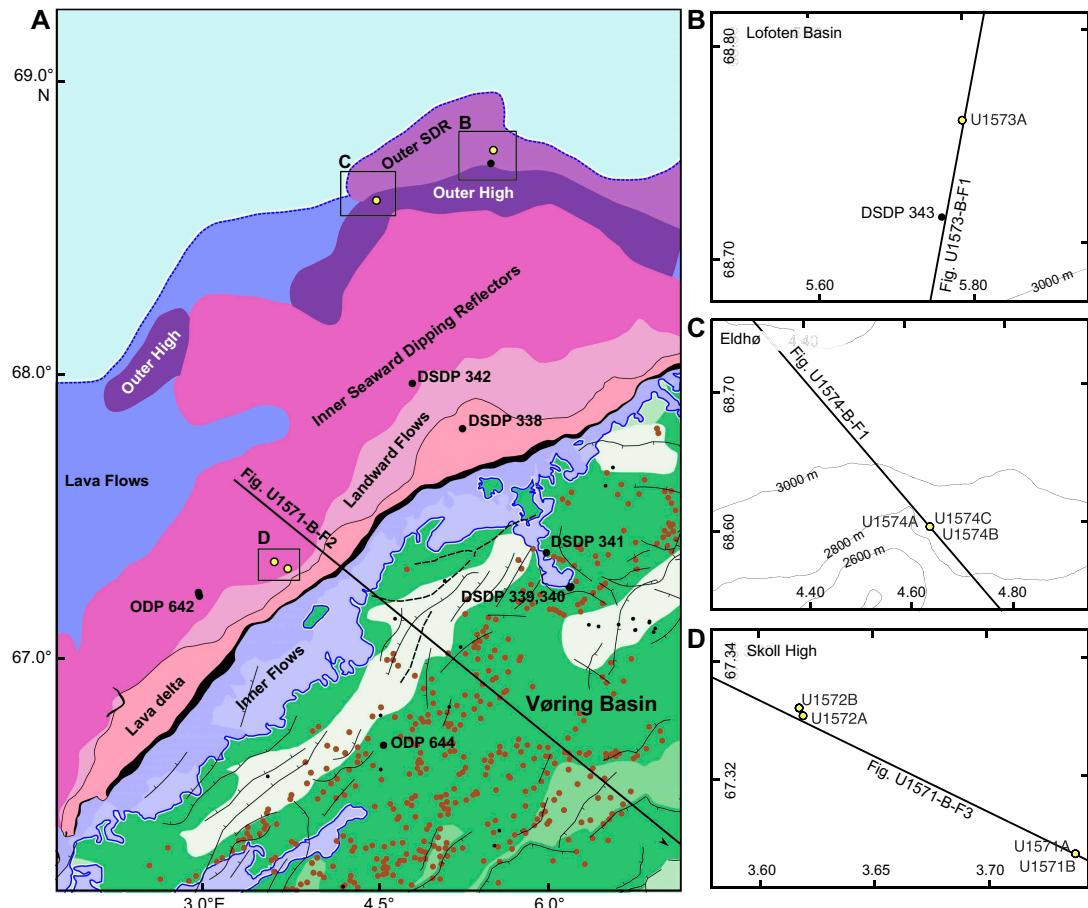
#### 5.5.1. Background and objectives

Sites U1571 (proposed Site VMVM-61A) and U1572 (proposed Site VMVM-07A) are located on the Skoll High on the Vøring Marginal High, about 15 km west of the Vøring Escarpment (Figure F15). The western part of the Vøring Marginal High is characterized by SDR sequences, first noted by Hinz and Weber (1976), that have early been interpreted as up to 6 km thick volcanic complexes near the COB (Skogseid and Eldholm, 1987, 1989). The first sampling of basalts on the Vøring Marginal High was accomplished in 1974 during DSDP Leg 38 (Talwani, Udintsev, et al., 1976). Holes 338 and 342 were drilled into basement, and 11 and 12 m of basalt were recovered, respectively, in 6 cores. Subsequently, a 1277 m deep borehole was drilled into the upper termination of the SDRs during ODP Leg 104 during the summer of 1985 (Eldholm et al., 1987). Hole 642E yielded more than 900 m of basalt and penetrated at least 135 lava flows and 3 dikes. The flows were interlayered with 59 volcaniclastic sedimentary units that mainly represent weathered basalt flow tops and one ignimbrite. Based on their geochemical and petrological composition, the volca-

nic succession was divided into an upper 800 m thick tholeiite series and a lower series that is at least 100 m thick and consists of dominantly basaltic andesite flows (Eldholm et al., 1987; Abdelmalak et al., 2016).

The extent and thickness of the breakup-related basaltic complex has subsequently been mapped along the mid-Norwegian continental margin using the concepts of seismic volcanostratigraphy and igneous seismic geomorphology (Berndt et al., 2001b; Planke et al., 2000, 2017; Millett et al., 2020b). The seismic reflection data reveal substantial differences in the volcanic emplacement environments across the Vøring Marginal High in time and space. Planke et al. (2017) documented distinct differences in the physical volcanology across the marginal high based on geomorphometric analysis of the top basalt surface. The eastern parts of the margin are characterized by a lava flow field that suggests that lava was flowing from eruption centers in the west toward the paleo-coastline in the east where they form a lava delta (Abdelmalak et al., 2016b). Incised valleys on the top basalt surface further suggest erosion by rivers flowing from west to east. The top basalt surface farther west is characterized by a rough morphology with numerous semicircular anomalies that have been interpreted as rootless cones due to lava flowing across a wet substrate leading to phreatomagmatic eruptions. This pitted surface belongs to a seismic sequence that onlaps onto the eastern flow field and must be younger than the subaerial basalts drilled in Hole 642E (Figure F7A). Strikingly, the two domains also exhibit very different structural deformation. Although there are only few normal faults with little displacement that offset the sequence with the pitted surface, the eastern lava flows toward the escarpment are offset by numerous north–south trending normal faults with offsets of tens of meters.

Site U1571 is located on the eastern, faulted lava flow field, whereas Site U1572 is located on the pitted surface farther west. The morphological difference between the two sites and the onlap



**Figure F15.** Tectonic and magmatic elements of the Vøring Basin. Legend in Figure F3. Brown dots indicate hydrothermal vent complexes.

relationship suggest that the lavas have been emplaced diachronously, possibly in different environments. The first objective of these sites was to obtain information on the emplacement environment based on the facies of the recovered basalt and sediments. This information will constrain the vertical movements of the margin, for example, whether both facies were indeed emplaced in a subaerial environment or whether there is a difference between the relatively well constrained subaerial flows in the east and the possibly shallow marine flows in the west. The second objective was to establish the time of emplacement of the two sequences to constrain the temporal evolution of the breakup volcanism. The third objective was closely linked to the first two: by analyzing the emplacement environment, we hoped to obtain information on the environmental and by proxy the climatic changes during and shortly after the main phase of volcanism. This will primarily be based on biostratigraphic and geochemical analysis of the sedimentary succession immediately above the top of the basalt and possibly the sedimentary layers between basalt flows. The fourth objective was to constrain the conditions in the melt region, specifically the composition of the mantle before melting and the ambient temperature from the geochemistry and petrology of the encountered rocks. Reaching this objective will help to parameterize geodynamic models that will distinguish between the relative importance of the drivers of excess volcanism.

Furthermore, Sites U1571 and U1572 will address two auxiliary objectives laid out in the *Scientific Prospectus*. First, we hope to provide information on the suitability of the breakup basalt sequences and interbedded sediments for permanent geological storage of CO<sub>2</sub>. This will require obtaining information on their porosity and permeability as well as their geochemical composition, in particular the abundance of olivine, which may react with CO<sub>2</sub>. Second, we hope to obtain information on the earliest incursion of deep water into the North Atlantic and the extent of major freshwater incursions into the young ocean (e.g., during the Azolla event) (Brinkhuis et al., 2006).

Finally, and akin to the other sites, we will use the drilling results to test and calibrate the volcanological predictions of seismic volcanostratigraphy and igneous seismic geomorphology for the Landward Flows and SDRs and smaller-scale volcanic elements identified in the seismic reflection data.

### 5.5.2. Operations

Site U1571 consisted of two holes. The first hole was cored using the RCB system from the seafloor to 247.6 mbsf and then logged with the triple combo, FMS-sonic, and UBI logging tool strings. Although the original plan called for a single hole to 240 mbsf, the poor recovery through a critical stratigraphic interval suggested an additional hole should be cored at the site. The APC and XCB systems were selected for the second hole. Hole U1571B was cored with the APC system to refusal at 57.2 mbsf. The XCB coring system was deployed and cored with good results through 143.7 mbsf. Polycrystalline diamond compact (PDC) XCB cutting shoes were used to core through the hard layers between 125 and 143.7 mbsf. Coring in Hole U1571B was terminated at 1005 h on 11 September 2021. The drill string was pulled back to the surface after clearing the seafloor at 1110 h. On clearing the seafloor, the vessel began moving in DP mode to the next site, which was located 3.1 nmi away. The APC/XCB coring bit cleared the rotary table at 1351 h, ending Hole U1571B and Site U1571. A total of 124 h or 5.2 days were recorded while on Site U1571.

A total of 59 cores were taken at the site. The RCB coring system was used for Hole U1571A. The RCB system cored 247.6 m and recovered 116.55 m of core (47.1%). Hole U1571B used both the APC system and the XCB system. Seven APC cores were taken over a 57.2 m interval. Recovery was 58.73 m (102.7%). The XCB system was used for 13 cores over a 86.5 m interval and recovered 62.2% of the cored interval.

Site U1572 also consisted of two holes. The first hole was cored using the RCB system from the seafloor to 330.5 mbsf and then logged with the triple combo tool string. Further logging was canceled because of degrading hole conditions. The original plan called for a single hole to 320 mbsf. Poor recovery through a critical zone suggested an additional hole should be cored at the site. The APC/XCB coring system was selected for the second hole at the site. Hole U1572B was cored with the APC and HLAPC systems to APC refusal at 209.6 mbsf. The XCB system was deployed, and the basement contact interval was cored to 224.3 mbsf. PDC XCB cutting shoes were used to core through the hard layers between 209.6 and 224.3 mbsf. Coring in Hole U1572B was terminated at

1045 h on 16 September. The drill string was pulled back to the surface after clearing the seafloor at 1235 h. On clearing the rotary table, the APC/XCB BHA was secured at 1612 h, ending Hole U1572B. The bridge was notified that the drill floor activities were complete. The thrusters and hydrophones were pulled back inside the hull of the vessel, and the vessel began the sea passage to Site U1573. A total of 122.5 h or 5.1 days were recorded while on Site U1572.

A total of 79 cores were recorded for the site. The RCB system was used for Hole U1572A. The RCB system cored 330.5 m and recovered 139.53 m of core (42.2%). Hole U1572B used the APC, HLAPC, and XCB systems. There were 16 APC cores recorded over a 149.3 m interval. Recovery was 154.45 m, giving an efficiency of 103.5%. The HLAPC system was used for 13 cores over a 60.3 m interval and recovered 99.1% of the cored interval. The XCB system was used for 4 cores over a 14.7 m interval and recovered 28.3% of the cored interval.

### 5.5.3. Principal results

#### 5.5.3.1. Lithostratigraphy

The succession recovered from the four holes cored at Sites U1571 and U1572 is divided into eight lithostratigraphic units. Units I–VI are sedimentary, Unit VII is igneous, and Unit VIII is characterized by primarily igneous rock interbedded with sediment and volcanics.

Units were determined based on macroscopic observations and microscopic analysis and are supported by physical properties and biostratigraphy. Hole U1572A is the only hole that includes all eight units. In Holes U1572B and U1571B, APC drilling was utilized, and recovery of sediments is more than 100% in many sediment cores. Where core recovery is low across unit boundaries or fall-in drilling is observed, boundaries were placed at the top of the underlying unit. Unit I is grayish brown and brown unconsolidated mud. Unit II is greenish gray consolidated mud and nanno-fossil ooze. Unit III is greenish gray mud, very dark gray diatomite, and diatom ooze. Unit IV consists of greenish gray radiolarian ooze interbedded with gray to black ash. Unit V is very dark gray mudstone with parallel lamination and dark gray ash. Unit VI is biosiliceous ooze with mudstone and common ash. Unit VII is basaltic andesite and basalt Lava Flows. Unit VIII consists of basalt and interbasalt sediments and is divided into two distinct subunits. Subunit VIIIa is essentially composed of hyaloclastite alternating with volcanic mudstone. Subunit VIIIb is an alternation of basaltic lava flows and minor mudstone. The presence of Unit VII as the top basement unit at Site U1572 constitutes the main difference between the two sites and likely is the reason for the difference in seismic character.

#### 5.5.3.2. Biostratigraphy

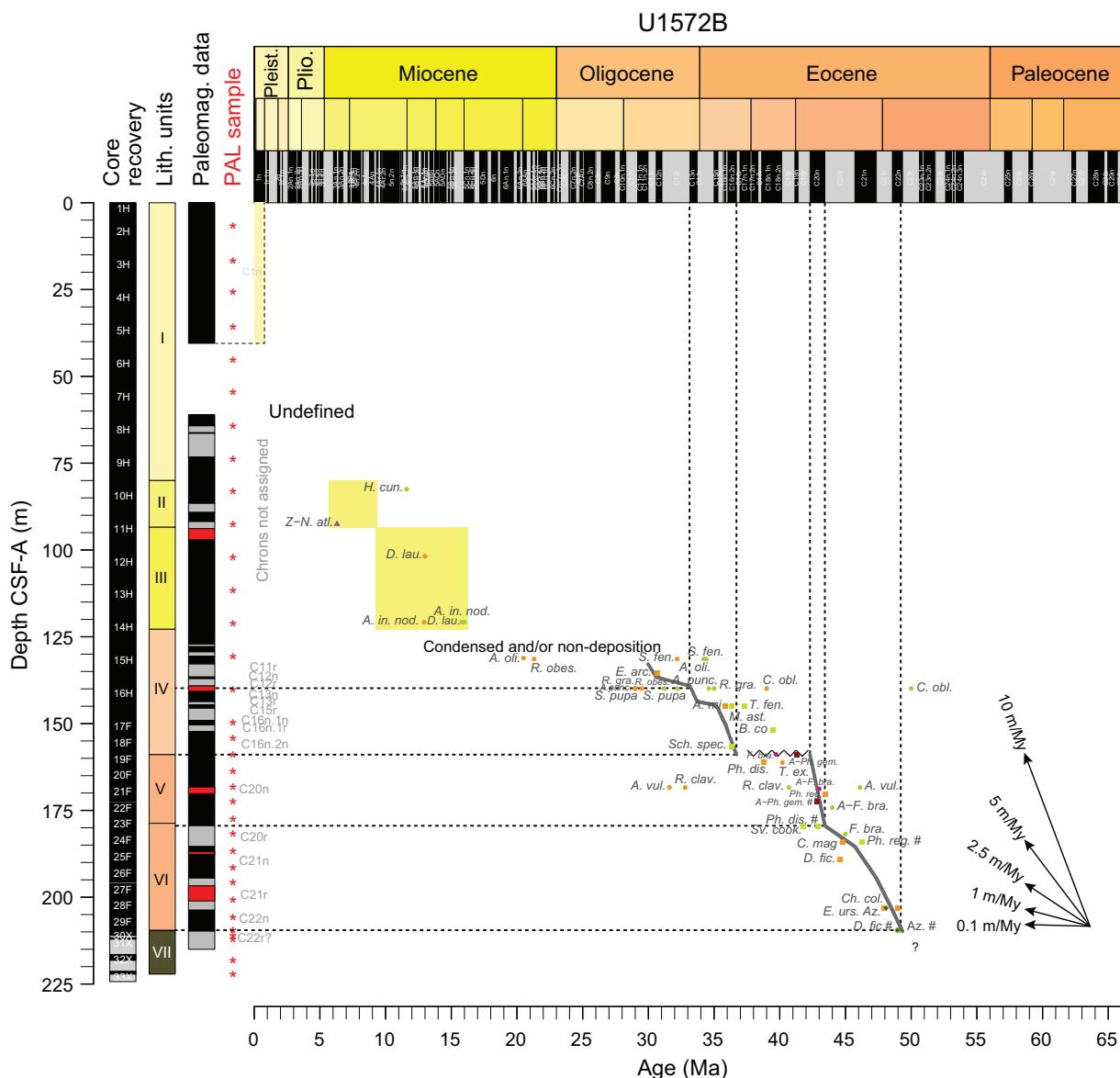
The Quaternary sediments recovered from the holes at Sites U1571 and U1572 (max = ~80 m) are (similar to most other holes drilled during Expedition 396) characterized by glacially influenced hemipelagic muds, which typically unconformably overlie Miocene strata. Siliceous and organic-walled microfossils were found in various abundances below the Quaternary deposits. Also, indications are that the successions at Sites U1571 and U1572 are very similar; lower to late Middle Miocene strata overlie upper lower to upper Eocene strata, with some lower Oligocene sediments, especially in Hole U1572B.

In Holes U1571B, U1572A, and U1572B, sediments recovered from directly above the igneous facies yield *Azolla* spp. and associated dinocyst marker species. The well-documented Nordic Sea *Azolla* phase has been dated as occurring between ~49 and 48 Ma (mid C22n to mid C21r) at the end of the early Eocene (Ypresian) (e.g., Brinkhuis et al., 2006; Barke et al., 2011). This constrains the earliest sediment accumulation to the very end of, or directly postdating, the EECO as defined by Westerhold et al. (2018). The middle to late Eocene appears to be marked by a progressive deepening or increasing distance to shore. Because of the general absence of microfossils, no biostratigraphic age constraints are available from in between the igneous facies.

The microfossil-rich sedimentary succession in Hole U1572B was utilized to construct an integrated age-depth model, and the correlation to the shipboard paleomagnetic reversal interpretation and lithostratigraphic units is shown in Figure F16, Table T5, and Figure F17.

### 5.5.3.3. Paleomagnetism

The low recovery in Holes U1571A and U1572A prevents the construction of a meaningful magnetostratigraphic record; however, better magnetostratigraphic results were obtained from Holes U1571B and U1572B. Lithostratigraphic Unit I sediments show mainly normal polarities in the uppermost 40 m, although the increase of the  $I_{20mT}/NRM$  ratio with depth suggests a possible climatic change. The upper and lowermost parts of Unit II show normal polarity with a reverse polarity in the lower part. Unit III has the lowest magnetic intensities ( $INRM < 10^{-4}$  A/m) for these sites. The dominantly normal polarities might be caused by drilling-induced magnetization. Unit IV shows evenly normal and reverse polarities. Unit V has exclusively normal polarities that are validated by discrete samples. Unit VI shows a mixture of normal and reverse polarities. The basaltic sequences show strong magnetization intensities with dominantly reverse polarities, all validated by discrete samples. The magnetic carrier in these basalts has a low coercivity and may potentially carry a viscous remanent magnetization that could be used for core reorientation. The anisotropy of magnetic susceptibility (AMS) of sediments shows that correction of magnetic inclination for inclination shallowing will not be necessary. The basalts carry a strong AMS resulting from viscous flow.



**Figure F16.** Age-depth estimates, Hole U1572B. Interpretation utilizes paleomagnetic reversal data, lithostratigraphic analyses, trends in physical property data, and bioevents. Abbreviations of bioevents can be found in Table T5. PAL = micropaleontological samples.

### 5.5.3.4. Geochemistry

The recovered basalt sequence from the Skoll High shows chemical and spatial variations between Sites U1571 and U1572. For example,  $\text{TiO}_2$  concentrations indicate the presence of occasional high-Ti basalts, usually associated with LIPs and overlap with high-Ti basalts from East Greenland and flows found during previous drilling in Hole 642E during ODP Leg 104 (Figure F11). The compositions of the uppermost basalts at the two sites appear to be representative of two different flow units. This supports the lithostratigraphic interpretation, which indicates an additional unit (VII) at the top of the basalts at Site U1572. IW recovered from intrabasalt sediments shows unusually high pH and low alkalinity for saline water. Basalt alteration and associated clay authigenesis at Sites U1571 and U1572 likely result in enriched IW Ca contents (released during clay formation) and depleted Mg and K as these elements are incorporated into the clays. Elements in IW samples from the basalt–sediment transition to 50–100 m above it show largely conservative

**Table T5.** Calcareous, biosiliceous and organic microfossil biostratigraphic tie-points used in the age-depth model for Hole U1572B and abbreviations of taxa as used in Figure F16.

Bio event	Top depth CSF-A (m)	Abbreviation	Age (Ma)	Reference	Type
FO <i>H. cuniformis</i>	82.50	H.cun.	<11.6	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
Zone <i>Neogloboquadrina atlantica</i> (sin.)	92.53	Z.N.atl.	6.3	Spiegler and Jansen, 1989	FORAM
LO <i>D. lauta</i>	101.80	D.lau.	13.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>A. ingens nodus</i>	120.80	A.in.nod.	15.8	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>D. lauta</i>	120.80	D.lau.	16.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>A. ingens nodus</i>	120.80	A.in.nod.	13.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>A. oligocenica</i>	131.10	A.olii.	20.5	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>A. oligocenica</i>	131.40	A.olii.	34.4	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>S. fennerae</i>	131.40	S.fen.	34.2	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>R. obesa</i>	131.40	R.obes.	21.3	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>S. fennerae</i>	131.40	S.fen.	32.2	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>E. arcuata</i>	135.42	E.arc.	30.7	Köthe et al., 2008	DINO
FO <i>A. punctifera</i>	139.90	A.punc.	35.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>C. oblongus</i>	139.90	C.obl.	50.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>R. granda</i>	139.90	R.gra.	34.6	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>R. obesa</i>	139.90	R.obes.	31.2	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>S. pupa</i>	139.90	S.pupa	32.2	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>A. punctifera</i>	139.90	A.punc.	29.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>C. oblongus</i>	139.90	C.obl.	39.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>R. granda</i>	139.90	R.gra.	29.5	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>S. pupa</i>	139.90	S.pupa	29.0	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>M. asterium</i>	144.95	M.ast.	36.34	Heilmann-Clausen and van Simaeys, 2005	DINO
FO <i>T. fenestrata</i>	144.95	T.fen.	37.29	Average from Bijl et al., 2021	DINO
LO <i>A. michoudii</i>	144.95	A.mi	35.85	Eldrett et al., 2004	DINO
FO <i>B. compta</i>	151.94	B.co	39.45	Williams et al., 1993	DINO
FO <i>Sch. spec.</i>	156.52	Sch.spec.	36.37	Average from Bijl et al., 2021	DINO
Acme <i>Ph. germinatum</i>	158.85	A-Ph.gem.	41.27	Eldrett et al., 2004	DINO
LO <i>F. brachiata</i>	158.85	F.bra.	39.7	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>Ph. Distinctum</i>	161.22	Ph.dis.	38.8	Average from Eldrett et al., 2004	DINO
LO <i>T. excavata</i>	161.22	T.ex.	40.2	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>A. vulgaris</i>	168.40	A.vul.	46.1	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>R. claviger</i>	168.40	R.clav.	40.7	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>A. vulgaris</i>	168.40	A.vul.	31.6	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>R. claviger</i>	168.40	R.clav.	32.8	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>F. brachiata acme</i>	168.80	A-F.bra.	43	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>Ph. Regalis</i>	170.35	Ph.reg.	43.44	Average from Eldrett et al., 2004	DINO
Acme <i>Ph. germinatum</i>	172.36	A-Ph.gem.#	42.87	Eldrett et al., 2004	DINO
FO <i>F. brachiata acme</i>	174.20	A-F.bra.	44	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>Ph. Distinctum</i>	179.56	Ph.dis.#	42.9	Average from Eldrett et al., 2004	DINO
FO <i>Sv. cooksoniae</i>	179.56	Sv.cook.	41.8	Average from Eldrett et al., 2004	DINO
FO <i>F. brachiata</i>	181.80	F.bra.	<45	Eldrett and Harding (2009)	DIAT
LO <i>T. hebergei</i>	181.80	T.hei.	<53.9	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
LO <i>T. regina</i>	181.80	T.reg.	<54.5	J.A. Barron, 2021 database (Lazarus et al., 2014 in Janus)	DIAT
FO <i>Ph. Regalis</i>	184.18	Ph.reg.#	46.27	Average from Eldrett et al., 2004	DINO
LO <i>C. magna</i>	184.18	C.mag	44.77	Average from Eldrett et al., 2004	DINO
LO <i>D. ficusoides</i>	189.02	D.fic.	44.58	Average from Eldrett et al., 2004	DINO
LO <i>Azolla</i>	203.13	Az.	48.1	Eldrett et al., 2004	AZOL
LO <i>Ch. Columna</i>	203.13	Ch.col.	48.96	Average from Eldrett et al., 2004	DINO
LO <i>E. ursulae</i>	203.13	E.urs.	47.96	Eldrett et al., 2004	DINO
FO <i>Azolla</i>	209.49	Az.#	48.96	Eldrett et al., 2004	AZOL
FO <i>D. ficusoides</i>	209.49	D.fic.#	48.96	Eldrett et al., 2004	DINO

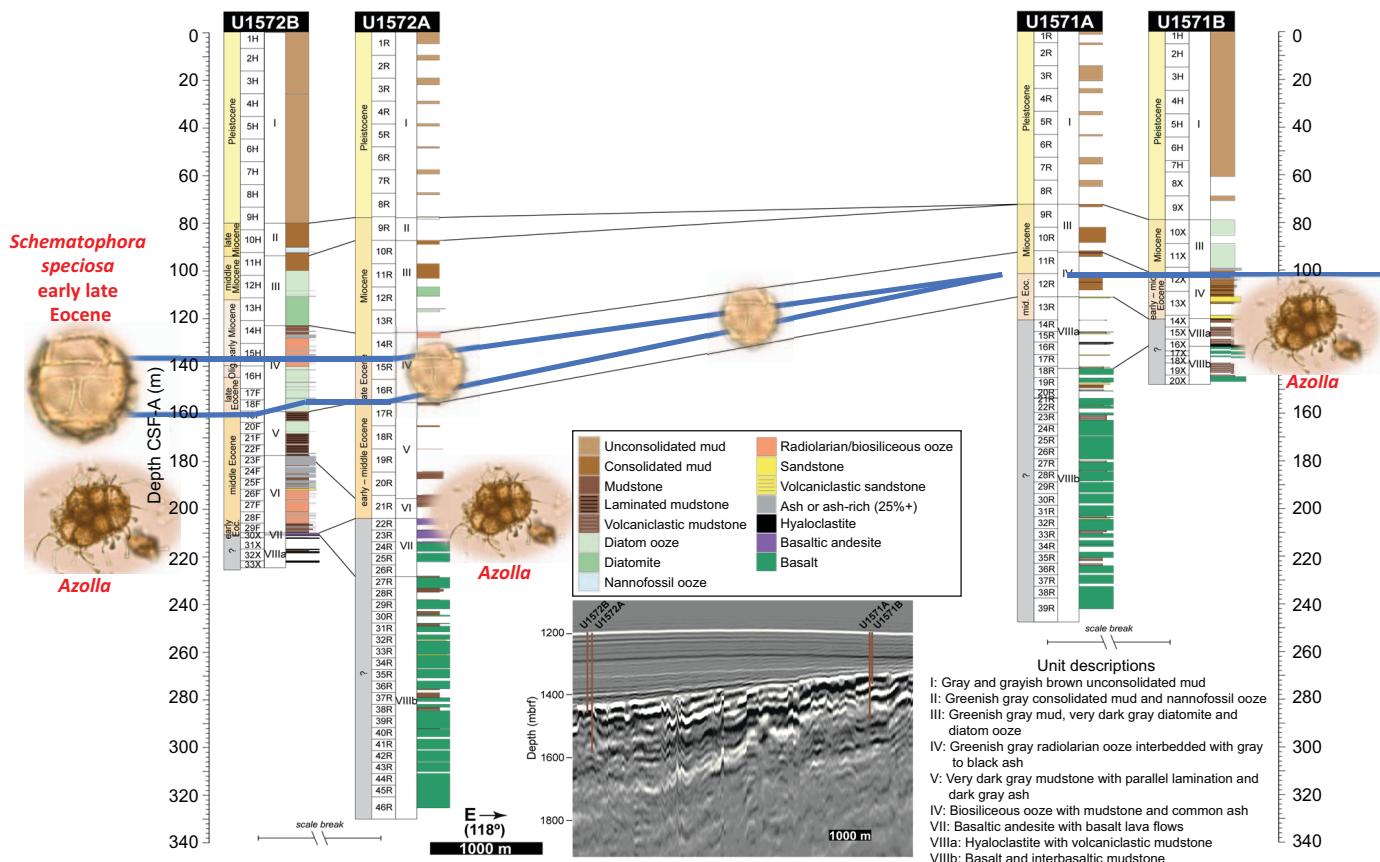
behavior downcore. Dissolved silica concentrations are sporadically enriched, and these intervals correspond to ash- and/or diatom-rich strata (e.g., Unit III). This is indicative of marine silicate weathering, and the high interstitial silica preserves some original features of volcanic glass, diatoms, and organic matter.

### 5.5.3.5. Physical properties

Shallow Quaternary Lithostratigraphic Unit 1 at Sites U1571 and U1572 comprises relatively high GRA bulk densities (~1.74 g/cm<sup>3</sup>) and NGR (42 counts/s) that transition abruptly downward into Units II and III, which generally comprise lower GRA bulk density (~1.4–1.5 g/cm<sup>3</sup>) and NGR (~23–25 counts/s) with an accompanying reduction in MS. The underlying Miocene to Eocene strata comprising Units IV–VI are best preserved in Hole U1572B where Units IV and VI reveal elevated NGR (mean = >30 counts/s) compared with the intervening Unit V (mean = 21 counts/s); MS is also elevated in these intervals, which is likely linked to increased ash content in the Eocene.

The basaltic andesite-dominated Unit VII is characterized by high GRA bulk densities (2.53 g/cm<sup>3</sup>) and *P*-wave velocities (4120 m/s) and low NGR (11 counts/s) in line with the data recorded for the basalt interiors found in Subunit VIIb but at slightly higher NGR associated with the more evolved compositions. Point magnetic susceptibility (MSP) and MS (592 and 401 × 10<sup>-5</sup> SI) are a factor of 10 higher than for the overlying sedimentary units but about half that observed for the underlying basalt-dominated Unit VIII.

A clear distinction is seen between the hyaloclastite Subunit VIIa and lava dominated Subunit VIIb. The former has relatively low  $P$ -wave velocities (~1800 m/s) and GRA bulk densities (~1.55 g/cm<sup>3</sup>) more in line with the low velocity, low bulk density interbasaltic sediments. Basaltic Lava Flow units from Unit VIII reveal clear asymmetrical profiles of  $P$ -wave velocity and GRA bulk density ranging from ~ 5000 m/s and 2.9 g/cm<sup>3</sup>, respectively, in the massive flow interiors to as low



**Figure F17.** Lithostratigraphic and chronobiostratigraphic correlation panel, Sites U1571 and U1572. Note the absence of the upper Eocene section at Site U1571. ? = unknown age.

as  $\sim$ 2000 m/s and  $1.5 \text{ g/cm}^3$ , respectively, for porous and variably altered flow tops and bases. Calculated grain densities average  $3 \text{ g/cm}^3$  for the basaltic sequences and exceed  $3.1 \text{ g/cm}^3$  in some cases, attesting to the basic undifferentiated nature of the basaltic magmas that fed the sequence. Abundant primary pore spaces in the form of unfilled vesicles are identified throughout the volcanic sequence, leading to calculated porosities commonly  $>30\%$ , whereas in certain cases these pores are filled with secondary minerals such as calcite and zeolite.

#### 5.5.3.6. Downhole measurements

Wireline logging operations at the Skoll High delivered a comprehensive data set covering approximately 147 m (about 80 m basaltic basement) in Hole U1571A and approximately 157 m (about 45 m basaltic basement) in Hole U1572A. Gamma radiation logging data aided in constraining the nature and distribution of more evolved igneous compositions in an uppermost interval of mixed volcanic and sedimentary interlayers. A key objective of the Skoll High sites was to constrain the nature of different seismic facies associated with the top basalt reflector between the two sites, in particular the “pitted surface” seismic facies associated with Hole U1572A. The transition from sediment to volcanic dominated sequences in Hole U1571A comprises a more protracted sequence of interlayered hard volcanic units and softer sediments. In contrast, the transition from the sediments to the volcanic sequence is sharper in Hole U1572A, with an initial abrupt increase in density, PEF, and MS followed by a gradual increase in density over the following  $\sim 10$  m culminating in a sharp increase in resistivity at  $\sim 220$  m. Core observations suggest abundant lava-water interaction for Lithostratigraphic Unit VIII, suggesting massive subaqueous flows and volcaniclastic units. This is consistent with the wireline data.

Gamma radiation also aided the identification and thickness quantification of several interlava sedimentary units that will form key future targets for palynology and isotopic age dating efforts to appraise the timing and environmental conditions of the eruption sequence. FMS image log data revealed high resolution imaging of intricate volcanic intrafacies features such as ubiquitous unfilled vesicles, mineralized amygdalites, pipe vesicles, and pillow lava features indicative of aqueous eruptions during the volcanic pile evolution.

### 5.6. Site U1573

Site U1573 consisted of one RCB hole with the aim to characterize the basement nature of the Outer SDRs. Hole U1573A resulted in coring about 90 m of late early Eocene sediments that lie on top of basaltic basement. The 54 m of drilled basement consists of massive sheet flows that were emplaced in a shallow marine setting and are interlayered by sediments. The upper 300 m of the shallow sediments at the site were washed because an approaching storm did not leave enough time to core the entire hole.

#### 5.6.1. Background and objectives

Site U1573 (proposed Site VMVM-09A) is located in the southern part of the Lofoten Basin just north of the Vøring Plateau at a 3167 m water depth (Figure F15). This part of the Vøring Margin has rarely been explored since the 1960s and 1970s. Site U1573 is located on Magnetic Anomaly 23 near the border between the Vøring and Lofoten Margins. The only available basement samples prior to Expedition 396 in this area are from DSDP Leg 38 Site 343 (Talwani, M., Udintsev, G., et al., 1976), during which basement from 253 to 284 mbsf was drilled and about 5 m of highly altered basalts that were interpreted to be part of a sill intrusion were recovered.

Compared to the Vøring Margin, the Lofoten Margin is characterized by a narrower rift system, less extensive breakup-related magmatic crustal thickening, and a rapid increase of water depth from the shelf into the Lofoten Basin. The Lofoten Margin consists of three segments that are characterized by different fault dip directions and separated by crustal transfer systems (Tsikalas et al., 2001). The southernmost transfer system, the Bivrost Lineament, separates the Lofoten Margin from the Vøring Margin. Wide-angle seismic data suggest about 12 km thick oceanic crust in the southern part of the Lofoten Basin, much thinner than the about 20 km thick crust of the Vøring Plateau (Breivik et al., 2009). This difference in crustal thickness should have resulted in a very different vertical movement history of the Lofoten Margin than that of the Vøring Margin. A previously calculated early Eocene bathymetric gradient from the Vøring to the Lofoten Margin of

a 500–1000 m water depth on the Lofoten Margin during breakup (Tsikalas et al., 2008) would imply submarine lava emplacement on the Lofoten Margin as opposed to the subaerial volcanic emplacement environment on the Vøring Margin (Eldholm et al., 1989). This bathymetric gradient has been attributed to rapid initial margin subsidence along a deep-rooted low-angle detachment fault (Tsikalas et al., 2008). Reflection seismic data reveal a horizontal basement reflection underlain by SDRs with lower seismic amplitude and shorter lateral extent than the main set of SDRs on the Vøring Margin. These Outer SDRs are a common feature of volcanic rifted margins and have been interpreted as submarine Lava Flows (Berndt et al., 2001b).

The stark difference between the observed subaerial Lava Flows at ODP Site 642 and the submarine conditions predicted by gravity modeling require a distinctly different subsidence history for the Vøring and Lofoten Margins. Thus, the first objective of drilling Site U1573 was to obtain information on the emplacement environment based on the facies of the recovered basalt and sediments. As such, reaching this first objective directly tests the predictions of the concept of seismic volcanostratigraphy. The second objective was to establish the timing of volcanism on the Lofoten Margin, in particular whether there is evidence of a time gap between the emplacement of the SDRs of the Vøring Margin on Skoll High (Sites U1571 and U1572) and the Outer SDRs. The third objective of Site U1573 was to constrain the conditions in the melt region during the presumably late phase of breakup volcanism. This includes not only any constraints on the composition of the mantle before melting and the ambient temperature and pressure from the geochemistry and petrology of the encountered rocks but also any signs of evolution of the melts in crustal magma chambers during breakup volcanism.

### 5.6.2. Operations

Site U1573 consisted of a single hole drilled first without coring using the RCB system with a center bit from the seafloor to 300.0 mbsf and then cored to a final depth of 440.9 mbsf, including 54.4 m into basement. The average penetration rate in the basement section was ~1.82 m/h. There were frequent jammed liners and core catchers which contributed to relatively low core recovery rates within basement (average = 43%). Coring in Hole U1573A was terminated at 1830 h on 20 September 2021 because of severe weather and heavy sea conditions that were expected over the operational area for the following day. The last core recovered was Core 396-U1573A-19R at 1925 h on 20 September. The drill string was pulled back on board at 2200 h. The BHA was broken down, and the drill collars were secured in the drill collar rack. The rig floor was secured at 0705 h on 21 September, ending Hole U1573A and Site U1573. A total of 101.0 h or 4.2 days were spent on site. A total of 18 cores were recovered using the RCB system over 140.9 m recovering 101 m (72%).

### 5.6.3. Principal results

#### 5.6.3.1. Lithostratigraphy

Site U1573 consists of a single hole that was drilled to 440.9 mbsf, with the recovered core interval spanning 300 to 440.9 mbsf. The 140.9 m thick succession consists of both sediments and basalts that are divided into four lithostratigraphic units. The criteria for the division of the units are based on significant lithologic changes, which are supported by physical properties variability. Units I–III are sedimentary, and Unit IV consists of basalt and interbasaltic sediments.

- Unit I consists of dark greenish gray to very dark gray claystone with rare weak parallel lamination or fining-upward sequences. Slight to heavy bioturbation is present throughout the unit and, diagenetic pyrite is observed locally.
- Unit II consists of dark greenish gray to very dark gray claystone with moderate bioturbation alternating with very dark gray organic-rich claystone with thin parallel lamination. Rare ash beds or ash-rich beds are present locally.
- Unit III consists of very dark gray to very dark bluish gray volcaniclastic claystone mudstone with sand.
- Unit IV consists mainly of gray to dark gray aphyric aphanitic to phaneritic basalt and interbasaltic claystone to sandstone.

### 5.6.3.2. Biostratigraphy

All sedimentary core catcher samples and selected samples from intercalated sedimentary beds in the igneous facies from Hole U1573A were processed for calcareous, siliceous, and organic microfossils. Calcareous and siliceous biogenic material was largely absent, with the exception of rare calcareous nannofossil occurrences in the lower several meters above basalt. However, most samples were productive for palynology, including some samples from between the igneous facies, and most yielded age diagnostic dinocysts, which indicate a late early Eocene age for the entire interval.

### 5.6.3.3. Paleomagnetism

Lithostratigraphic Unit I is characterized by an average intensity of magnetization of 0.05 A/m and a median coercivity of 2.73. This coercivity may be owed to sedimentary hematite or goethite. The unique discrete sample measured is not ferromagnetic, devoid of magnetite, because the average bulk MS is  $0.447 \times 10^{-3}$  SI.

Unit II has an 0.05 A/m average intensity of magnetization. The median coercivity (3.04), MS ( $K < 1.03 \times 10^{-3}$  SI), and corrected degree of anisotropy ( $P_j < 1.105$ ) are higher than in Unit I. High K is owed to small contributions from magnetite. High coercivity and  $P_j$  may be due to a higher goethite/hematite ratio and/or a larger iron sulfide (e.g., pyrrhotite) population in the ash. Multiple minerals carrying remanence are evident from the three types of remanent magnetization identified in four discrete samples. The low and medium coercivity component is erased by AFs of 20 and 70 mT, respectively. The high coercivity component is stable throughout the demagnetization process, specifically in the 120 mT field.

Unit III has an average magnetization intensity of 0.07 A/m, which is slightly higher than Units I and II. In contrast, the medium coercivity (2.53) is lower than in Units I and II. Lower median coercivity can be related to domination of volcanic material, which usually is richer in iron sulfide (e.g., pyrrhotite) but poorer in goethite and hematite. Note that goethite and hematite have higher coercivity than iron sulfides. No discrete samples were measured.

Unit IV has high average magnetization intensity (2.6 A/m) and high median coercivity (6.02). Such high coercivities may be owed to a larger goethite/hematite ratio. The unit is ferromagnetic with the highest bulk MS of all the units, indicating the dominance of magnetite. However, the corrected degree of anisotropy is the lowest among all the units. This indicates that the magnetic fabrics are carried dominantly by magnetite and the populations of pyrrhotite and hematite are small. Three magnetization components are conspicuous from the demagnetization behavior of the discrete samples.

Unit I has four short periods of magnetic reversals between Sections 396-U1573A-2R-1 (300 mbsf) and 2R-5 (306.71 m). This is followed by a long episode of magnetic reversal in Sections 4R-2 (321.05 mbsf) through 5R-5 (333.91 mbsf) in Unit II. After this, the magnetic polarity flips several times at a relatively higher frequency (within a couple meters) until Section 7R-3 (350.98 mbsf). It then remains reversed through Section 8R-3 (361.72 mbsf). After that, the polarity stays dominantly normal in the rest of the hole through Section 18R-2 (436.74 mbsf). Preliminary data from paleomagnetic analyses suggest a predominance of normal polarity at the top of the cored succession, which we tentatively assign to magnetochron C22n (48.57–49.34 Ma).

### 5.6.3.4. Geochemistry

Site U1573 in the Lofoten Basin provided a unique opportunity to sample an Outer SDR sequence on a magma-rich passive continental margin. The sampled lavas are noteworthy in that the higher silica and low-Ti compositions are chemically distinct from both typical mid-ocean-ridge basalts and LIP magmatism. The sedimentary package recovered above the lavas is limited to 100 m. As seen at previous sites, high organic matter production and/or preservation is found in volcaniclastic sediments, suggesting a causal relationship. The IW samples show a diagenetic environment controlled by basalt weathering, with processes consuming alkalinity in the overlying sedimentary strata. There is also strong evidence of authigenic carbonate formation above the sediment/basalt contact, linked to the liberation of calcium during the weathering of primary silicates.

### 5.6.3.5. Physical properties

Sedimentary Units I and II are characterized by downward increasing density (from ~1.6 to ~2 g/cm<sup>3</sup>) consistent with increasing compaction. NGR values fluctuate between ~30 and 60 counts/s and display a general increase with depth over Units I and II. Increased scatter in MS is observed over ash-rich intervals within sedimentary sections. The hard rock basement sequence comprises generally very low NGR values, typically under 10 counts/s, consistent with the basalt-dominated compositions and has generally high density (mean GRA bulk density = 2.28 g/cm<sup>3</sup>; mean MAD grain density = 2.83 g/cm<sup>3</sup>) and velocities (mean = 3660 m/s). The presence and nature of interlava sediments with lower velocity (~2000 m/s) compared to the higher velocity basaltic flows (~4500 m/s) give first insights into the geological causes of the Outer SDR seismic reflectivity.

## 5.7. Site U1574

Three holes were drilled at Site U1574 targeting the basement and the overlying sedimentary sequence of an Outer High, Eldhø, located on the northern flank of the Vøring Plateau. Hole U1574A was drilled with the RCB system and sampled 94 m of pillow basalts, hyaloclastites, and a ~30 m thick interbasalt sedimentary interval. Both triple combo and FMS-sonic wireline log data were acquired in the basement. Only one APC core was collected in Hole U1574B, whereas a 170 m thick sequence of sediments was almost completely recovered using APC coring in Hole U1574C. The last core of the expedition was an XCB core recovering 1.5 m of spectacular pillow basalts, sediments, and hyaloclastites.

### 5.7.1. Background and objectives

Site U1574 (proposed Site VMVM-80A) probes the top of an Outer High called Eldhø (Figure F15). It is located on the northern flank of the Vøring Plateau at a 2825 m water depth. Outer Highs are conspicuous volcanic seismic facies units that are located seaward of the Inner SDR facies units (Hopper et al., 1992; Planke et al., 2000; Franke et al., 2019). Although most volcanic seismic facies units are relatively continuous along a rifted margin, Outer Highs have a limited lateral extent of typically 10–20 km (Berndt et al., 2001a). They form elongate structures that rise above the top of the adjacent Inner and Outer SDRs. Seismically, Outer Highs frequently display a high-amplitude top reflection of normal seismic polarity, chaotic internal reflection patterns, and high-amplitude reflections that extend from the Outer High into adjacent sedimentary sequences that overlie the SDRs.

Because Outer Highs are characteristic features of volcanic rifted margins located between Inner and Outer SDRs, their presence likely signifies specific geological processes and provides clues to the evolution of volcanic passive margins. The precruise interpretations of Outer Highs mainly relied on seismic volcanostratigraphic interpretations, field analogs, and geological reasoning. The facies unit was interpreted as dominantly hyaloclastite deposits formed when the eruption fissures were submerged during the initial phase of seafloor spreading. In this situation, eruptions would happen at sea level or in shallow marine conditions and lava would get into direct contact with seawater leading to phreatomagmatic eruptions. The formation of Surtsey between 1964 and 1968 may serve as a recent analog (Kokelaar, 1983).

The structure targeted at Site U1574 is a typical example of an Outer High (Figure F7B). We call this structure Eldhø, meaning fire mountain in Scandinavian languages. The name is also inspired by the achievements of Professor Olav Eldholm, who was a pioneer in drilling and defining the concepts of volcanic margins and LIPs (e.g., Eldholm et al., 1989). Eldhø rises more than 1500 m above the adjacent Inner SDRs in the southeast and about 650 m above the Outer SDRs in the northwest with numerous high-amplitude seismic reflections that transgress into the adjacent sedimentary basins south of the high. Its lateral extent is about 17 km in the dip (northwest–southeast) direction and about 30 km along strike (southwest–northeast). A sedimentary basin, the Eldhø Basin, which has kilometer-thick postbasalt sediment deposits, is located landward of Eldhø.

The first objective of Site U1574 was to obtain information on the emplacement environment based on observations of the drilled volcanic facies. If Eldhø was indeed emplaced close to sea level, this will validate these aspects of seismic volcanostratigraphy, and their presence may then

be used as a marker for vertical movement. The second objective was to constrain the conditions in the melt region, specifically the composition of the mantle before melting and the ambient temperature and pressure conditions in the melt region during the late phase of breakup volcanism based on the geochemical composition and petrology of the encountered rocks. Reaching the first two objectives will allow to parameterize geodynamic models for excessive breakup volcanism.

Site U1574 was also drilled to provide information on the suitability of the breakup basalt sequences and interbedded sediments for permanent geological storage of CO<sub>2</sub>. If the Outer Highs indeed consist mainly of hyaloclastite deposits as postulated, they should be characterized by high porosity and permeability, which may make them a suitable target CO<sub>2</sub> storage. Finally, we hoped to obtain information on the earliest incursion of deep water into the North Atlantic and the extent of freshwater incursions into the young ocean during the early to middle Eocene (Brinkhuis et al., 2006) by coring the overlying sedimentary strata.

### 5.7.2. Operations

Site U1574 consisted of three holes. The first hole was cored using the RCB coring system from the seafloor to 260.0 mbsf and then logged with the triple combo and FMS-sonic logging tool strings. The original plan for Site U1574 was to core a single hole to 310 mbsf, but the objectives were achieved by 260 mbsf. After completing Hole U1574A, the Co-Chief Scientists decided to use the remaining expedition time to piston core the sediment section above the basement and to attempt to improve recovery of the sediment/basement interface. The APC/XCB coring system was selected for the second hole at the site. After assembling and lowering the APC/XCB system to the seafloor, the first APC core misfired and penetrated the seafloor at least twice. Hole U1574B was terminated after the poor mudline core at 9.5 mbsf. The vessel was offset 20 m south, and Hole U1574C was piston cored to 169.2 mbsf. We conducted four borehole temperature measurements with the advanced piston corer temperature (APCT-3) tool on Cores 396-U1574C-4H (38.0 mbsf), 7H (66.1 mbsf), 10H (94.6 mbsf), and 13H (118.7 mbsf). The Icefield MI-5 core orientation tool was deployed for the first 10 cores. The XCB coring system was deployed after the last APC core and immediately cored basement. A PDC XCB cutting shoe was used to core a hard layer at 169.2 mbsf. Coring in Hole U1574C was terminated at 1630 h on 1 October 2021 at a final depth of 171.5 mbsf. The last core was on deck at 1655 h. All coring tools were put away. The drill string was pulled back to the surface after clearing the seafloor at 1920 h. The drill collars were laid out to the drill collar racks and the outer core barrel was disassembled and secured. The drill floor was secured for transit at 0231 h on 2 October, ending Hole U1574C and Site U1574. A total of 182.5 h or 7.6 days were recorded while on Site U1574.

A total of 59 cores were recorded for the site. The RCB system was used exclusively for Hole U1574A. It cored 260.0 m and recovered 88.41 m of core (34%). The APC system was used in both Holes U1574B and U1574C. There were 20 APC cores recovered over a 178.7 interval. The recovery for the 20 cores was 183.27 m (103%). The XCB system was used for 1 core over a 2.3 m interval and recovered 64% of the cored interval.

### 5.7.3. Principal results

#### 5.7.3.1. Lithostratigraphy

Site U1574 is located on the lower slope of the Vøringsfjellet Plateau, on the southern flank of Eldhø. This Outer High was likely formed by subaqueous volcanism during the initial phase of seafloor spreading. The succession of recovered material consists of both sediments and basalts and is divided into five lithostratigraphic units. The division of the units is based on changes in lithology, which coincide with changes in physical property trends. Units I–IV are sedimentary, and Unit V consists of basalt, hyaloclastite, and minor interbasaltic sediments.

Unit I consists of brown and brownish gray unconsolidated mud with intervals of rare pebbles, likely related to ice rafted debris, and rare to common foraminifers.

Unit II consists of pale yellow and grayish green consolidated mud, in which the first Eocene-aged dinocyst zonations are encountered.

Unit III is very dark gray mudstone with rare beds of sandstone and ash and some parallel lamination and bioturbation, showing an increase in organic content downhole.

Unit IV is very dark gray to very dark grayish brown organic rich mudstone with thin parallel lamination. Compared to overlying units, Unit IV is richer in foraminifers and nannofossils, with abundances increasing downhole in the unit.

Unit V mainly consists of aphyric- to plagioclase-phyric pillow basalt with localized hyaloclastite and rare interbedded mudstone. Unit V is divided into two subunits (Va and Vb) based on lithologic and geochemical variations. The two subunits are separated by a large interval of no recovery to very poor recovery of sediments, and both are overlain by hyaloclastites. Subunit Va contains both massive and pillow basalts and shows significant variation of composition with depth. Subunit Vb contains primarily pillow basalt and presents remarkably uniform composition.

#### 5.7.3.2. Biostratigraphy

Site U1574 holes comprise up to ~170 m of sedimentary strata above the igneous basement. Except for Cores 396-U1574A-1R through 3R, all sedimentary core catcher samples from Hole U1574A were processed for palynology. These samples, including two taken from between basalt flows (Samples 32R-CC and 33R-CC), are assigned a late early Eocene age based on marker dinocyst taxa and broader assemblage characteristics. The overlying Quaternary strata yield calcareous and rare biosiliceous fossil remains, but these are largely absent from most of the expanded late early Eocene strata. A significant exception is the lower ~20 m of sediment on top of the igneous basement (Samples 17R-CC and 18R-CC), which contains calcareous remains, including calcareous nannofossils, planktonic and benthic foraminifers, and micromolluscs, and includes fine preservation of aragonitic pteropods. Benthic foraminifers are mostly mineral filled, and diatoms are replaced by or coated with pyrite microframboids. These samples are tentatively attributed to planktonic foraminifer Zone E7a of Wade et al. (2011), in line with inferences from dinocyst biostratigraphy.

#### 5.7.3.3. Paleomagnetism

Lithostratigraphic Units I–III are characterized by low average intensity of magnetization (0.02, 0.006, and 0.02 A/m, respectively), whereas their median coercivity is 1.69, 1.31, and 0.87, respectively. The lower intensity of magnetization and coercivity may be owed to a smaller population of magnetic minerals. No discrete samples were measured for these three units. Unit IV has a slightly higher average intensity of magnetization (0.006 A/m) than Units I–II. The bulk MS is  $1.06 \times 10^{-3}$  to  $.014 \times 10^{-3}$  SI, indicating a ferro + paramagnetic AMS without magnetite. Median coercivity (2.61) is much higher than in Units I–III. The corrected degree of anisotropy in the range of ~1.1 and median coercivity of 2.61 indicate the absence of micaceous minerals and a higher population of goethite and hematite, respectively. Only one of the two discrete samples presents stable magnetization directions (Sample 396-U1574A-17R-1, 111–113 cm). The low coercivity remanent magnetization component is erased by a field of 10 mT. The high coercivity component is removed at an applied field of 100 mT.

The average intensity of magnetization (4.38 A/m) in Unit V is several orders of magnitude higher than in the other units. The median coercivity (2.61) is the same as in Unit IV. The bulk MS is  $\sim 15 \times 10^{-3}$  SI, and the corrected degree of anisotropy is ~1.1. Two magnetization components are evident. The low coercivity component is erased by application of a field of 10 mT. The high coercivity component is removed by a field of 40–60 mT. The magnetic signals are apparently carried by magnetite.

Magnetic polarity is normal through Section 396-U1574A-4R-2 (31.05 mbsf). The polarity then remains dominantly reversed to the bottom of the hole (Section 38R-3; 258.15 mbsf). This long, dominantly reverse polarity episode is truncated by short-lived positive polarity events.

#### 5.7.3.4. Geochemistry

Site U1574 consists of two holes, one that targeted the basement (U1574A) and one that was aimed at obtaining high recovery of the overlying sediments (U1574C). The hard rock sampled from Hole U1574A consists of low-Ti basalts. Significant variability of some geochemical parame-

ters ( $\text{TiO}_2$ , Mg#, Ni, and Zr) downcore suggests recovery of multiple flow units. Interbasalt sediments contain what appears to be dolomite.

The sediments consist of an expanded section from the early late Eocene period. IW compositions indicate a diagenetic environment dominated by the weathering of primary silicate phases and the formation of authigenic clays and carbonate. Calcium and strontium decrease downcore due to their release in the formation of smectite and other clay minerals. The same formation reactions lead to the removal of Mg and K from the IW, resulting in decreasing Mg and K content downcore. Enrichment of other elements (e.g., Si and B) in the upper 40 m of the core suggest a lithology-controlled elemental composition of the IW.

#### 5.7.3.5. Physical properties

The uppermost Lithostratigraphic Unit I has high bulk densities (mean =  $1.82 \text{ g/cm}^3$ ) and NGR values (average = 38 counts/s) and relatively low *P*-wave velocities ( $\sim 1575 \text{ m/s}$ ) and magnetic susceptibilities ( $\sim 45 \times 10^{-5} \text{ SI}$ ). Eight MAD samples collected throughout this unit have an average bulk density of  $1.82 \text{ g/cm}^3$ , grain density of  $2.76 \text{ g/cm}^3$ , and porosity of 54%. Six thermal conductivity measurements range from 1.1 to 1.5  $\text{W}/(\text{m}\cdot\text{K})$  and have an average of  $1.3 \text{ W}/(\text{m}\cdot\text{K})$ .

In Unit II, *P*-wave velocities, GRA bulk densities, magnetic susceptibilities, and NGR average  $\sim 1550 \text{ m/s}$ ,  $1.43 \text{ g/cm}^3$ ,  $\sim 21 \times 10^{-5} \text{ SI}$ , and 19 counts/s, respectively. Analysis of two MAD samples revealed very high porosities of 83%, a grain density of  $2.70 \text{ g/cm}^3$ , and a bulk density of  $1.31 \text{ g/cm}^3$ , consistent with poorly compacted clay-rich sediments.

A total of 25 discrete MAD sample analyses indicate a mean bulk density of  $1.59 \text{ g/cm}^3$  for the  $\sim 90 \text{ m}$  thick Eocene Unit III, slightly lower than the mean GRA bulk density of  $1.62 \text{ g/cm}^3$ . Porosities range from 37% to 82%. The baseline values of the measured physical properties generally increase with depth throughout Unit III. Whole-round *P*-wave velocity values average  $1560 \text{ m/s}$ ,  $\sim 250 \text{ m/s}$  slower than velocities measured through 30 discrete *P*-wave caliper measurements using the SHMG, and have a significant spread with a standard deviation of  $120.5 \text{ m/s}$ . MS increases compared to the overlying Unit II (average =  $105 \times 10^{-5} \text{ SI}$ ) and features increased scattering linked to the occurrence of ash layers.

Reduced MS scatter (mean =  $\sim 90 \times 10^{-5} \text{ SI}$ ) is observed for the underlying Unit IV. Mean GRA ( $\sim 1.83 \text{ g/cm}^3$ ) and MAD ( $n = 8$ ,  $\sim 1.73 \text{ g/cm}^3$ ) bulk densities correspond well to an associated average porosity of 59%. Whole-round *P*-wave velocities with a mean of  $1565 \text{ m/s}$  are within 20 m/s of the mean recorded using *P*-wave caliper measurements and similar to average *P*-wave velocities found for Unit III. NGR values increase slightly compared to the overlying unit and have a constant value around 51 counts/s.

The massive basalt and pillow basalts from Unit V have significantly higher *P*-wave velocities, bulk densities ( $2.4\text{--}2.6 \text{ g/cm}^3$ ), and magnetic susceptibilities ( $1000\text{--}2000 \times 10^{-5} \text{ SI}$ ) than those found for the overlying sedimentary successions. Similar to interbasaltic sediments, the NGR counts, *P*-wave velocities, and bulk densities measured for the hyaloclastite intervals feature a similar but inverse response to those measured for the pillow and massive basalts because *P*-wave velocities were on average  $1000\text{--}1500 \text{ m/s}$  lower than the  $4000\text{--}5000 \text{ m/s}$  typically associated with the pillow lavas and massive intervals. Mean *P*-wave velocities of the massive basalt intervals typically featured on the lower end of this range. The unit has relatively low porosities (15%–20%), although significantly higher porosities are recorded for the interbasaltic layers.

#### 5.7.3.6. Downhole logging

The priority target for Hole U1574A was to characterize the basement nature of Eldhø. Logging operations collected important data for intervals with poor core recovery in challenging logging conditions where hole instability resulted in only partial data collection for the bottom  $\sim 50 \text{ m}$  of the open hole section. The basaltic composition pillow lavas are characterized by high densities ( $\sim 2.6 \text{ g/cm}^3$ ) and fluctuating velocities ranging from  $\sim 3$  to  $5.2 \text{ km/s}$  with associated low gamma radiation values ( $\sim 9 \text{ gAPI}$ ). FMS imaging of the formation reveals pillow structures and the presence of Lava Flow units, which correspond to highly vesicular basalt in the recovered core. Wireline data also show the presence of three thin elevated gamma radiation sediment layers along with a much thicker interval of high gamma radiation ( $\sim 70 \text{ gAPI}$ ) stratigraphy associated with an inter-

val of ~30 m with almost no core recovery. This layer is imaged with the FMS, which records a resistive fractured nature with traces of potential pillow-like features and possible sediment layering toward the base.

## 6. Preliminary scientific assessment

During Expedition 396, we collected data from 21 boreholes in five different geological settings along the volcanic rifted margin off mid-Norway. The boreholes probe almost all elements of a typical volcanic rifted margin and the associated sedimentary archive. Recovery ranges from the subbasalt stratigraphy through hydrothermal vent complex deposits, thick flood basalts deposits in SDR sequences to postbreakup phreatomagmatic volcanoes, and basalts representing thickened oceanic crust formed during early seafloor spreading. The cored sediments include a comprehensive record of the upper Paleocene and lower Eocene strata, including expanded sections of the PETM. Thick lower Eocene sediments recovered above the top basalt horizon document the environment and subsidence history of the margin during the initial stages of northeast Atlantic Ocean formation. Along with an enormous amount of high-quality 2-D and 3-D seismic data and legacy boreholes from DSDP, ODP, and the NPD, Expedition 396 results will allow us to assess and reevaluate the development of volcanic rifted margins and the implications of LIPs on paleoclimate and mass extinctions.

The first main objective of the expedition was to constrain the tectonomagmatic processes that lead to excess volcanism during the formation of a new ocean. Magmatism is controlled by the composition and rheological properties of the upper mantle, the initial pressure and temperature conditions, and the tectonically controlled unloading of the mantle through space and time. These boundary conditions will ultimately control the response of the system, including induced convection, passive, and active upwelling—processes that move mantle material into the melt window. We drilled 587 m of breakup-related basalts in eight different holes that cover a wide range of compositions. Cores from these holes will allow us to constrain the pressure and temperature conditions and the composition of the upper mantle during various phases of breakup volcanism through geochemical analyses, whereas the age of the magmatism will be constrained by geochronological analyses. The igneous rocks comprise the main Inner SDR units, the Outer High, and the Outer SDR unit as well as an enigmatic dacitic lava flow along the Vøring transform margin. These Expedition 396 data will be complemented by boreholes into a lava delta on the Møre Margin by NPD and samples of the earliest lava flows on the Vøring Margin that were recovered in ODP Hole 642E. High-precision radiometric dating, biostratigraphy of intrabasalt sediments, and paleomagnetic studies will allow us to constrain the age and rates of the magmatism in unprecedented detail. Furthermore, the samples allow a direct assessment of the paleowater depth, specifically, subaerial vs. marine emplacement of the extrusive basalts. Along with paleoenvironmental studies of the immediate sedimentary succession above the basalts by biostratigraphic and sedimentological analyses, this will constrain the vertical movements of the margin and by proxy (and in conjunction with the geophysically constrained magmatic crustal thickening) the rifting history of the margin. This is a main boundary condition for numerical simulation of the tectonomagmatic breakup processes that will allow us to reach the first objective of the expedition.

The second main objective of the expedition was to study the environmental consequences of breakup volcanism through time. Breakup volcanism in the North Atlantic was accompanied by a phase of rapid global warming, the so-called PETM. With the results of Expedition 396, we will test the hypothesis that the observed global warming was induced by magmatic intrusions into the surrounding sedimentary basins. It has been proposed that the intrusions led to contact metamorphism in the host rock and expulsion of carbon-bearing hydrothermal fluids through pipe-like vent complexes into the young ocean basin and from there onward into the atmosphere.

During Expedition 396, we successfully recovered 240 m of sediments of PETM age from eight boreholes within vent structures and their immediate vicinity in the southern Vøring Basin (Figure F8). These cores will allow to test the hypothesis by biostratigraphic, sedimentological, and geochemical analyses along with integration of wireline logs and high-resolution 3-D seismic data. A total of 1955 m of Paleocene and Eocene age sedimentary rocks were drilled during Expedition

396, with a recovery of 1125 m (58%) (Table T6). These cores represent a comprehensive archive of the environmental changes that took place from the onset of the EECO through the PETM and into the middle and late Eocene. The proxy data derived from the postcruise analyses will provide insights into the various drivers and responses in the earth system during these times of a global hothouse climate and subsequent cooling.

The comprehensive Expedition 396 database will also be used to study the opening of the north Atlantic seaways, the onset of ocean circulation patterns, and the occurrence of freshwater incursions into the young ocean basins. We have encountered the remains of the freshwater fern *Azolla* spp. in several of the expedition boreholes, and further analyses will provide a crucial documentation between its massive occurrence in the Arctic Ocean, Norwegian/Greenland Sea, and North Sea areas and sites in the Labrador Sea on the east coast of North America.

Finally, the recovered 587 m of various basaltic facies units will be invaluable to achieve the other secondary objective, namely, the study of the suitability of the breakup basalt successions for the permanent storage of CO<sub>2</sub>. Some of the recovered basalts are very vesicular providing large reaction surfaces, making them potential reservoir rocks for permanent carbon storage. Detailed analyses and integration with geophysical data will allow us to extrapolate these findings in three dimensions and form the basis for a robust assessment of the storage potential.

In spite of the ambitious work program set out in the expedition's *Scientific Prospectus*, it was possible to significantly surpass our operational and scientific objectives. The expedition results provide a much broader than anticipated scientific database and should allow us to reach all the expedition's objectives during postcruise work.

**Table T6.** Assessment, Expedition 396. TD = total depth.

Location	Proposed drilling	Target	Max depth (mbsf)	Drilled holes	Main result	Logs	Total (m)	Recovery (m)	Target (m)	Comments	
Kolga High basalts	20A 1 hole	Subbasalt	200 (30)	U1565 2 holes	Granitic basement		60	20	11	Target reached. Initial basalt and interbasalt sediments cored and logged in Hole U1566A. Subbasalt granitic rocks cored at both sites.	
	23A 1 hole	Initial basalt and subbasalt	200 (25)	U1566 1 hole	Basalt flows with sediments, granite below	4 runs	182	100	158		
N Modgunn vent complex	31A 1 hole	Sediments across PE transition	200 (55)	U1567 3 holes	Paleocene, Paleocene–Eocene transition, PETM, vent infill	2 runs	384	303	282	Target reached. Hole U1567A is a reference hole across the Paleocene/Eocene boundary. The other four holes provide a detailed sampling of the vent complex in-fill, including the PETM. Abundant ash layers recovered, and two holes were logged.	
	40B 2 holes	Vent complex, Paleogene	200 (55)	U1568 2 holes	Vent infill, PETM, late Paleocene	2 runs	325	239	252		
Mimir High Paleogene	55B 2 holes	early Eocene to late Paleocene	800 (255)	U1569 1 hole	Expanded early Eocene, PETM, late Paleocene	Bridge	405	145	268	Target reached. Site U1569 sampled the early Eocene, the PETM, and the late Paleocene before TD at 405 mbsf. Ribbon sites provided complimentary sampling of the early Eocene, the PETM, and mid-late Paleocene, with very pristine microfossils. Abundant ash layers. A thin igneous unit was cored in two holes.	
	56A (alt) ribbon	Paleocene to earliest Eocene	200	U1570 4 holes	Expanded early Eocene, PETM, Paleocene; dacite unit	2 runs	764	270	603		
Skoll High basalts	61A 1 hole	Lava field— <i>Landward Flows</i>	240 (125)	U1571 2 holes	Subaerial basalt flows with sediments	3 runs	391	229	164	Target reached. Hole U1571A cored and logged a sequence of subaerial tabular lava flows and interbasalt red sediments. A very different basalt sequence was sampled in the upper 50 m of the basalt basement in U1572A, likely deposited in an aquatic environment; interlava black sediments.	
	7A 1 hole	Pitted basalt flows/SDR	320 (220)	U1572 2 holes	Mixed basalts with various sediments	1 run collapse	555	359	142		
Lofoten Basin Eldhø basalts	9A 2 holes	Basalts of Outer SDR	550 (450)	U1573 1 hole	Sheet flows with interbasalt sediments		441	141	59	Target reached. Hole U1573A cored a sequence of basalt flows with interbedded marine sediments, one of which could be data. A very different basalt sequence was sampled at U1574, including hyaloclastites, pillow basalts, and massive flows. Dominantly Eocene and some Quaternary sediment cores above the basaltic basement. Note: Hole U1574B (10 m) failed.	
	80A 1 hole	Volcanoclastic Outer High	310 (210)	U1574 3 holes	Hyaloclastite, pillow basalts, massive flows with Eocene sediments	2 runs	442	273	96		
Totals	12 holes		3720 (1935)	21 holes			17 runs	3949	2079	2035	Note: In total, 385 m was washed.

## 7. Outreach

The onshore participants of Expedition 396 included an outreach officer from the United States who communicated the goals of the expedition to students and the public across the globe through live broadcasts and social media with the support of shipboard scientists and marine technicians.

Using the direct satellite link, computer devices, and Zoom videoconferencing software, expedition scientists and the outreach officer connected with people all over the world. A typical session included an introduction to the current location and expedition goals followed by a tour of the laboratories. Scientists were stationed at each laboratory to give an overview of the science being conducted and to answer questions. We finished each live event with a 10–15 min question and answer session.

During Expedition 396, we completed 15 broadcasts and connected with 794 participants in 7 different countries. We also conducted 2 live events on YouTube, which had 677 views by the end of the expedition.

The outreach officer managed the *JOIDES Resolution* Expedition 396 website (<https://joidesresolution.org/expedition/mid-norwegian-continental-margin-magma>) and completed seven blog entries on various topics ranging from the expedition scientific objectives to expedition-related science topics and sailing in the time of COVID-19. The expedition website and blog posts totaled over 1000 views by the end of the expedition.

The outreach officer and the onboard imaging specialist also added at least one post every day to the *JOIDES Resolution* Facebook, Twitter, and Instagram social media accounts, totaling 258 posts by the end of the expedition. These posts were seen a total of 374,000 times and resulted in an increase of 51 new Instagram followers, 445 new Twitter followers, and 118 new Facebook page fans over the 2 month expedition. That is a total increase of 614 newly engaged individuals.

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