1	Why does rapid contraction of the radius of maximum wind precede
2	rapid intensification in tropical cyclones?
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#### Abstract

24 The radius of maximum wind (RMW) has been found to contract rapidly well preceding rapid 25 intensification in tropical cyclones (TCs) in recent literature but the understanding of the involved dynamics is incomplete. In this study, this phenomenon is revisited based on ensemble 26 27 axisymmetric numerical simulations. Consistent with previous studies, because the absolute 28 angular momentum (AAM) is not conserved following the RMW, the phenomenon can not be 29 understood based on the AAM-based dynamics. Both budgets of tangential wind and the rate of 30 change in the RMW are shown to provide dynamical insights into the simulated relationship between the rapid intensification and rapid RMW contraction. During the rapid RMW contraction 31 32 stage, due to the weak TC intensity and large RMW, the moderate negative radial gradient of radial 33 vorticity flux and small curvature of the radial distribution of tangential wind near the RMW favor 34 rapid RMW contraction but weak diabatic heating far inside the RMW leads to weak low-level 35 inflow and small radial absolute vorticity flux near the RMW and thus a relatively small 36 intensification rate. As RMW contraction continues and TC intensity increases, diabatic heating 37 inside the RMW and radial inflow near the RMW increase, leading to a substantial increase in 38 radial absolute vorticity flux near the RMW and thus the rapid TC intensification. However, the 39 RMW contraction rate decreases rapidly due to the rapid increase in the curvature of the radial 40 distribution of tangential wind near the RMW as the TC intensifies rapidly and RMW decreases.

### 41 **1. Introduction**

42 Understanding the relationship between changes in intensity and the radius of maximum wind 43 (RMW) in tropical cyclones (TCs) is important for understanding the dynamics of TC 44 intensification because RMW contraction is often related to intensification (Shapiro and 45 Willoughby 1982; Schubert and Hack 1982). Based on the balanced vortex dynamics, Shapiro and Willoughby (1982) showed that tangential wind tendency in response to diabatic heating in the 46 47 eyewall is the greatest inside the RMW. This means that TC intensification and RMW contraction 48 would occur simultaneously, as mentioned in many earlier observational studies (e.g., Willoughby et al. 1982; Willoughby 1990) and numerical studies (e.g., Judt and Chen 2013). However, more 49 50 recent observational and numerical studies (Stern et al. 2015; Oin et al. 2016; Li et al. 2019) found 51 that the RMW contraction often stops well before the end of intensification, which is in contrast 52 with that implied by the result of Shapiro and Willoughby (1982). Stern et al. (2015) attributed the 53 cessation of RMW contraction well before the end of intensification to the increased curvature of 54 the radial distribution of tangential wind at the RMW as TC rapidly intensifies.

55 Another striking feature regarding the relationship between TC intensification and RMW 56 contraction was recently documented by Wu and Ruan (2021), who showed that statistically rapid 57 RMW contraction often precedes rapid intensification (RI) in observations. This phenomenon also 58 appeared in many previous observational and numerical studies (e.g., Corbosiero et al. 2005, see 59 their Fig. 4; Shimada and Horinouchi 2018, see their Fig. 7; Chen et al. 2011, see their Fig. 11; 60 Stern et al. 2015, see their Fig. 4), but its dynamics did not receive any special attention in those 61 studies. Results from statistical analyses for the North Atlantic TCs during 2000-2017 in Wu and 62 Ruan (2021) revealed that rapid RMW contraction does not necessarily cause RI and generally 63 only after the RMW contracts to a certain small size could a TC undergo RI. They found that the 64 TC intensification rate (IR) and the RMW contraction rate in the early intensification stage depend 65 on the amount and radial location of convective heating inside the RMW and the size of the RMW. This dependence can be explained based on the tangential wind budget equation and balanced 66 67 vortex dynamics (Schubert and Hack 1982). Namely, more convective heating inside the RMW 68 drives stronger inflow to transport higher absolute angular momentum (AAM) inward to spin up 69 the inner core in the lower troposphere; and the smaller RMW for a given intensity usually implies 70 higher inner-core inertial stability and thus higher efficiency of eyewall heating in spinning up the 71 inner core (Schubert and Hack 1982). However, Wu and Ruan (2021) did not provide any detailed 72 insight into the dynamical processes that cause the rapid contraction preceding RI and they did not 73 discuss what determines the amount and radial location of diabatic heating in the eyewall relative 74 to the RMW.

In addition, many previous studies have qualitatively explained the relationship between TC intensification and RMW contraction based on the AAM-based dynamics (e.g., Holland and Merrill 1984; Pu et al. 2009; Chen et al. 2011, Kilroy et al. 2016). By this view, the RMW is assumed to be approximately a material surface, namely, the AAM following the RMW is nearly conserved during the RMW contraction and TC intensification. The AAM-based dynamics can be explained more transparently based on the AAM ( $M_m$ ) passing through the RMW ( $r_m$ ) as given below

82

$$M_m = r_m V_m + \frac{1}{2} f r_m^2,$$
 (1)

83 where *f* is the Coriolis parameter,  $M_m$  and  $V_m$  are the AAM and the maximum tangential wind at 84  $r_m$ , respectively. Differentiating (1) with respect to time ( $\tau$ ) following the RMW, one can obtain

85 
$$\frac{dV_m}{d\tau} = \left(\frac{V_m}{r_m} + f\right) \left(-\frac{dr_m}{d\tau}\right) + \frac{\dot{M}_m}{r_m}.$$
 (2)

86 where  $\dot{M}_m$  denotes the rate of change of  $M_m$  following the RMW. Note that both  $dV_m/d\tau$  and

87  $dr_m/d\tau$  also follow the RMW rather than the air parcel. Equation (2) indicates that the RMW contraction  $(-dr_m/d\tau > 0)$  is associated with intensification as long as the AAM does not 88 89 decrease too much following the RMW. Equation (2) also indicates that the contribution of RMW contraction to TC intensification depends on  $V_m/r_m + f$ , which is equivalent to the effect of 90 inertial stability in the balanced vortex dynamics. If  $\dot{M}_m$  is assumed to be small enough and 91 92 omitted, Eq. (2) implies that the rapid TC intensification lags the rapid RMW contraction because 93 the IR depends on not only the contraction rate but also the size of RMW and TC intensity. Namely, 94 for the same RMW contraction rate, the weak intensity and large RMW are unfavorable for RI, or 95 alternatively, RI may occur when the TC is strong enough with a relatively small RMW even the 96 RMW contraction rate becomes small. This seems to provide a reasonable explanation for the 97 observed relationship between TC intensification and RMW contraction as discussed in Wu and 98 Ruan (2021) as well. However, Stern et al. (2015) showed that the AAM could change up to 50% 99 following the RMW in their numerical simulated TCs, demonstrating that the RMW does not follow an AAM surface. Namely,  $\dot{M}_m$  in Eq. (2) could not be ignored during TC intensification. 100 101 Stern et al. (2015) found that the RMW first moves inward across AAM surfaces from higher to lower values of AAM during the initial rapid RMW contraction period, giving rise to negative  $\dot{M}_m$ 102 103 but once the RMW contraction ceases, AAM surfaces move inward across the RMW, corresponding to positive  $\dot{M}_m$ , and thus the TC continues to intensify more rapidly. This finding 104 105 indicates that it is inappropriate to understand the relationship between TC intensification and 106 RMW contraction solely based on the above AAM-based argument with the assumption of AAM being conserved along the RMW. Note that, here, it does not mean that any physical conservation 107 108 law has been "violated", because the RMW itself is not a material surface but simply the location 109 of a maximum in a scalar field as also suggested by Stern et al. (2015). Note that the negligible 110  $\dot{M}_m$  along the RMW was also assumed in the theoretical intensification model of Emanuel (2012).

111 Therefore, it is necessary to provide a detailed evaluation of Eq. (2) during TC intensification.

112 In this study, the relationship between TC intensification and RMW contraction is revisited 113 based on comprehensive diagnostics of ensemble axisymmetric numerical simulations of TC 114 intensification. The main objectives are threefold: to provide an evaluation of the AAM-based 115 dynamics using the simulation results, to examine how the amount and radial location of diabatic 116 heating in the evewall relative to the RMW evolve during the TC intensification, and to provide a 117 more complete dynamical understanding of the relationship between TC intensification and RMW 118 contraction based on the budgets for the RMW contraction and the tangential wind following the 119 RMW. The rest of paper is organized as follow. Section 2 describes the model setup and 120 experimental design. Section 3 provides an overview of the relationship between the RMW 121 contraction and TC intensification in the ensemble simulations. The AAM-based dynamics of Eq. 122 (2) is evaluated based on ensemble simulations in Section 4. Section 5 explains why the simulated 123 rapid TC intensification lags the rapid RMW contraction based on the budgets of tangential wind 124 and the rate of change in the RMW. Main conclusions are summarized in the last section.

# 125 **2. Model setup and experimental design**

The model used in this study is the axisymmetric version of the nonhydrostatic Cloud Model (CM1), version 19.8 (Bryan and Fritsch 2002), with the detailed model setup described in Li et al. (2020b). The model domain extends radially up to 3100 km with the grid spacing of 1 km within 100-km radius and linearly stretched to 12 km at the lateral boundary. The model atmosphere has 59 vertical levels with stretched grids below 5.5 km and uniform grid spacing of 500 m above. An *f* plane with the Coriolis parameter of  $5 \times 10^{-5}$  s<sup>-1</sup> is assumed. The cloud microphysics parameterization is the double-moment scheme of Thompson et al. (2008), and no cumulus

convective parameterization is used. The Newtonian cooling capped at 2 K dav<sup>-1</sup> is used to mimic 133 134 the radiative cooling (Rotunno and Emanuel 1987), and dissipative heating is not considered in this 135 study. The horizontal and asymptotic vertical mixing lengths are set at 700 m and 70 m, respectively. The surface enthalpy exchange coefficient is constant of  $1.2 \times 10^{-3}$ , and the surface 136 137 drag coefficient depends on surface wind speed as given in Donelan et al. (2004). The sea surface 138 temperature (SST)-dependent sounding experiments in Li et al. (2020b) are used, in which the 139 atmospheric sounding for each SST bin is sorted from reanalysis data to ensure the consistency 140 between SST and initial atmospheric sounding as in nature. There are four SST ensemble 141 experiments with their corresponding atmospheric soundings sorted from the western North Pacific 142 (WNP; labeled as SST28 WNP-SST31 WNP), and four experiments with their corresponding 143 atmospheric soundings sorted from the North Atlantic (NA; labeled as SST27 NA-SST30 NA). 144 For more details on the sounding construction, the readers are referred to Li et al. (2020b). Note 145 that some of the simulation results were also used in Wang et al. (2021) to validate an energetically 146 based TC intensification dynamical system model.

147 Each ensemble experiment has 21 simulations. The radial profile of the initial tangential wind 148 is given following Wood and White (2011). In the standard run, the initial maximum tangential wind speed is 15 m s<sup>-1</sup> with the initial RMW of 80 km at the surface and decreases linearly with 149 150 height to zero at about 18 km. In the other 20 runs, the initial maximum tangential wind speed and the initial RMW are consecutively perturbed by the increments of  $\pm 0.1$  m s<sup>-1</sup> (for 10 runs) and 151 152  $\pm 0.4$  km (for 10 runs), respectively. The simulation time for each run is 240 hours with the model 153 output saved at every 10 min for budget analyses. To avoid the impacts of small-scale processes 154 on TC IR or contraction rate, we only focus on the ensemble mean results as in Li et al. (2020b) 155 and Wang et al. (2021). Note that for each experiment, although the timing of contraction and 156 intensification varies among those ensemble runs (cf. Fig. 4 in Wang et al. 2021), our preliminary

157 results have indicated that the ensemble mean results discussed here are insensitive to the 158 perturbation increments within reasonable ranges.

# 159 **3.** An overview of the ensemble simulation results

160 The radius-time Hovmöller plots of the ensemble-mean AAM and tangential wind at 2-km 161 height for each experiment are shown in Fig. 1. The 2-km height is used as in Stern et al. (2015) 162 because we focus on the low-level flow but try to avoid the effects of unbalanced boundary-layer 163 flow and vertical mixing on the AAM following the RMW. It can be seen from Fig. 1 that for each 164 experiment, there is an initial adjustment period, during which a weak initial vortex core weakens 165 and a new inner vortex occurs with a rapid, discontinuous contraction of the RMW. To avoid the 166 impact of the adjustment period on the analysis of the primary intensification period, we only focus 167 on the primary intensification phase, namely the period after the initial discontinuous RMW 168 contraction and the new inner vortex attains an intensity of 12 m s<sup>-1</sup> (results are insensitive to this 169 threshold) as marked by the first horizontal purple line in Fig. 1 until the end of the simulated TC 170 intensification (as the first time at which the ensemble mean intensity change rate attains zero) as 171 marked by the second horizontal purple line in Fig. 1. During the earlier stage of this primary 172 intensification period, the RMW first contracts slowly or barely changes with a relatively slow IR 173 (compared to the peak rate of intensification). Subsequently, the RMW contracts rapidly and then 174 the RMW contraction rate slows down and approaches zero finally. The IR still increases during 175 the late stage of RMW contraction, as implied by the increasing density of the contours of tangential 176 wind speed. This means that the change rates of the simulated TC intensity and RMW are not 177 simultaneous, and the rapid contraction precedes the RI.

A closer comparison between the TC IR and change rate of RMW at 2-km height in each
experiment is shown in Figs. 2a,b. Note that the small-scale perturbations of tangential wind less

180 than 6 h in time and less than 8 km in radial direction are filtered out similar to Stern et al. (2015) 181 and Li et al. (2019). The "tendencies" of the intensity and RMW refer to the rates of their changes 182 in the subsequent 6 h forward difference after filtering rather than 24 h used in the observational 183 study of Wu and Ruan (2021). As expected, the peak of RMW contraction rate is prior to the peak 184 of IR in all experiments, indicating that the rapid contraction precedes the RI. Note that the "rapid" 185 periods for both the IR and RMW contraction rates refer to the periods around their corresponding 186 peaks and are not defined based on any specific threshold. Figures 2c,d show the TC IR and the 187 rate of change of RMW based on the filtered tangential wind averaged the layer between the surface 188 and 500 m height. Note that as done above, the initial adjustment period and the period after the 189 end of intensification are not shown. Overall, the results are similar to those at 2-km height shown 190 in Figs. 2a,b. Namely, the peak of contraction rate occurs prior to the peak of IR by about 10 h to 191 more than 40 h, consistent with the observations documented in Wu and Ruan (2021). Furthermore, 192 with the decrease in the time to attain the RI as SST increases, the time between the peak of RMW 193 contraction rate and the peak of IR also decreases. This may be partly because the inner-core 194 diabatic heating increases more rapidly with higher SST (Črnivec et al. 2016), and thus earlier RI 195 following the RMW rapid contraction. In addition, the RMW contraction tends to cease (as the 196 contraction rate attains zero) well before the end of intensification in all experiments no matter 197 above or within the boundary layer, consistent with previous observational and numerical studies 198 (Qin et al. 2016; Stern et al. 2015; Li et al. 2019).

We can see from Fig. 1 that the AAM following the RMW is not conserved, which is in contrast with the assumption in some previous studies (e.g., Holland and Merrill 1984; Emanuel 201 2012). As we can see more clearly from Figs. 3a,b, in the very beginning of the primary intensification stage as the RMW contracts slowly or barely changes, the AAM following the RMW increases by roughly 30–50%. However, subsequently the AAM following the RMW 204 decreases by roughly 20–40% with the rapid RMW contraction. Then, as the contraction rate slows 205 down, the AAM following the RMW increases again by about 20-60%. These are generally 206 consistent with those found by Stern et al. (2015). As a result, although the AAM surface at the 207 current RMW contracts in the following 6 h during the primary intensification stage (Fig. 1; Figs. 208 3c,d) as the RMW contracts (Figs. 2a,b), they contract at different rates (Figs. 3e,f). Note that 209 assuming the RMW propagates continuously, the AAM surface at the current RMW must contract 210 if there is intensification. The contraction rate of the AAM surface is much smaller than that of the 211 RMW during the rapid contraction phase (Figs. 3e,f), and the peak of the AAM surface contraction 212 rate generally lags the peak of the RMW contraction rate (Figs. 3c,d), suggesting that the term  $\dot{M}_m/r_m$  in Eq. (2) is nonnegligible following the RMW. Therefore, the relationship between the 213 214 TC intensification and RMW contraction could not be understood simply based on the conservation 215 of AAM following the RMW as hypothesized in some previous studies (e.g., Holland and Merrill 216 1984; Pu et al. 2009; Chen et al. 2011; Kilroy et al. 2016; Emanuel 2012). This will be further 217 quantitatively evaluated based on the results at the 2-km height in the next section.

### 218

### 4. An evaluation of the AAM-based dynamics

219 As mentioned above, based on Eq. (2), if the effect of  $M_m$  is omitted, the slow IR during the rapid contraction phase could be considered as a result of small  $V_m/r_m + f$ , and the RI during the 220 slow contraction period could be considered as a result of the increased  $V_m/r_m + f$ . As shown in 221 Figs. 4a,b,  $V_m/r_m + f$  is indeed small during the rapid contraction phase because of the large 222 223 RMW and weak TC intensity in all experiments, and then  $V_m/r_m + f$  increases with the decrease 224 of RMW. To evaluate the AAM-based dynamics, the simulated IR, the diagnosed IR from Eq. (2), and the contributions by the RMW contraction  $[(V_m/r_m + f)(-dr_m/d\tau)]$  and  $\dot{M}_m$   $(\dot{M}_m/r_m)$  are 225 shown in Figs. 4c–h. Note that the simulated  $dr_m/d\tau$  and  $\dot{M}_m$  at each time  $\tau_0$  are calculated as 226

the change rate of  $r_m$  and  $M_m$  between  $\tau_0$  and  $\tau_0 + 6h$ , and the  $V_m$  and  $r_m$  on the rhs of Eq. 227 228 (2) are their averages between  $\tau_0$  and  $\tau_0 + 6h$  based on the model outputs at 10-min intervals. 229 As expected, the evolution and magnitude of the diagnosed IR are comparable well with that of the 230 simulated IR in all experiments (Figs. 4c,d). However, although  $V_m/r_m + f$  is small during the 231 rapid contraction phase (Figs. 4a,b), the rapid contraction still implies a large IR (Figs. 4e,f). The 232 RMW contraction-implied IR decreases with the decreasing RMW contraction rate (Figs. 4e,f) although  $V_m/r_m + f$  shows considerable increase due to the TC intensification (Figs. 4a,b). This 233 234 means that the evolution of the diagnosed IR associated with the RMW contraction during the early intensification stage is generally similar to the evolution of the RMW contraction rate. However, 235 236 this is in contrast with the evolution of the simulated IR, suggesting that the term associated with 237 the RMW contraction in Eq. (2) could not determine the simulated IR alone. This is because the diagnosed IR due to  $\dot{M}_m/r_m$  is negative during the rapid contraction phase in all experiments, 238 showing a minimum at the peak of the RMW contraction rate (Figs. 4g,h). The negative  $\dot{M}_m/r_m$ 239 240 during the rapid contraction phase is consistent with the decrease of AAM following the RMW as 241 mentioned earlier (Fig. 1; Figs. 3a,b). As the contraction rate slows down, the diagnosed IR due to 242  $\dot{M}_m/r_m$  increases gradually and becomes positive, consistent with the increase of the simulated IR (Figs. 4c,d). The diagnosed IR due to  $\dot{M}_m/r_m$  shows a great contribution to the peak of the 243 244 simulated IR (Figs. 4g,h).

The above analysis indicates that the simulated slow IR during the rapid contraction phase is mainly associated with the decrease of AAM following the RMW, and the simulated RI during the slow contraction phase is largely associated with the increase of AAM following the RMW. This means that  $\dot{M}_m$  is not negligible following the RMW and is key to the lag of the RI behind the rapid RMW contraction in the AAM-based dynamics. Therefore, it is inappropriate to explain the relationship between RMW contraction and TC intensification by omitting the effect of  $\dot{M}_m$  as assumed in many previous studies (e.g., Holland and Merrill 1984; Pu et al. 2009; Chen et al. 2011, Emanuel 2012; Kilroy et al. 2016). Note that  $\dot{M}_m$  itself is a function of TC IR and RMW contraction rate by definition. Therefore, the AAM-based dynamics provides little insights into the physical/dynamical understanding on the relationship between TC intensification and RMW contraction.

# 256 5. New insights from tangential wind and RMW budgets

To help understand the relationship between the simulated TC intensification and RMW contraction during the primary intensification stage after the initial adjustment, we conducted the budgets for both tangential wind and the rate of change in the RMW as done in Li et al. (2019), with the results discussed in this section. Note that the present study focuses on the early intensification stage based on the ensemble-mean budget results, rather than the late intensification stage based on the single simulation results in Li et al. (2019). The tangential wind tendency equation in the axisymmetric, cylindrical coordinates can be written as (Li et al. 2019, 2020a)

264 
$$\frac{\partial V_m}{\partial \tau} = -u_m \xi_{a,m} - w_m \frac{\partial V_m}{\partial z} + F_{h,m} + F_{v,m}, \tag{3}$$

265 where the subscript "m" denotes the variables at the RMW, u and w denote radial wind and vertical velocity,  $\xi_a = \frac{\partial v}{\partial r} + \frac{v}{r} + f$  denotes absolute vertical vorticity. Note that at the RMW, the 266 radial gradient of tangential wind is zero by definition and thus  $\xi_a$  equals to  $V_m/r_m + f$ , which 267 268 also appears in Eq. (2). The terms on the rhs of Eq. (3) denote the radial (horizontal) flux of absolute 269 vorticity (ADV H), vertical advection (ADV V), radial diffusion due to subgrid-scale horizontal 270 mixing (DIFF H), and vertical diffusion due to turbulent mixing including surface friction 271 (DIFF V). As in Li et al. (2019, 2020a), all instantaneous terms in Eq. (3) are directly from the model output and thus ideally no computational error in the budget is introduced. 272

The diagnostic equation for the RMW contraction rate, based on the kinematic/geometric definition of the RMW, can be given below following Stern et al. (2015),

275 
$$\frac{dRMW}{d\tau} = -\frac{(\partial/\partial r)(\partial V_m/\partial \tau)}{\partial^2 V_m/\partial r^2},$$
 (4)

where the numerator and denominator on the rhs denote the radial gradient of tangential wind tendency and the curvature of the radial distribution of tangential wind at the RMW, respectively. Because of the negative definite nature of the curvature at the RMW, the sign of the RMW tendency depends on the radial gradient of tangential wind tendency. Substituting the tangential wind tendency equation (3) into Eq. (4), we get (Li et al. 2019)

281 
$$\frac{dRMW}{d\tau} = \frac{(\partial/\partial r)(u_m\xi_{a,m})}{\partial^2 V_m/\partial r^2} + \frac{(\partial/\partial r)(w_m\frac{\partial V_m}{\partial z})}{\partial^2 V_m/\partial r^2} - \frac{(\partial/\partial r)(F_{h,m})}{\partial^2 V_m/\partial r^2} - \frac{(\partial/\partial r)(F_{v,m})}{\partial^2 V_m/\partial r^2}.$$
 (5)

The terms on the rhs of Eq. (5) denote contributions of the radial gradient of tangential wind tendency to the RMW tendency due to radial flux of absolute vorticity (ADV\_H\_S15), vertical advection (ADV\_V\_S15), radial diffusion (DIFF\_H\_S15), and vertical diffusion including surface friction (DIFF\_V\_S15).

286 Equations (3) and (5) can be used to understand the relationship between TC intensification 287 and RMW contraction. Considering that the eyewall above the boundary layer is spun up by the 288 vertical advection of AAM from the boundary layer (e.g., Zhang et al. 2001; Heng and Wang 2016; 289 Peng et al. 2018; Li et al. 2019, 2020a) and the RMW within the boundary layer also influences 290 the RMW above the boundary layer (Stern et al. 2014, 2015), we focus on the RMW tendency and 291 tangential wind tendency below 500-m height in the boundary layer as in Li et al. (2019). Similar 292 to Stern et al. (2015) and Li et al. (2019), the small-scale features in the vertically-averaged 293 tangential wind as well as those tangential wind tendency terms in the rhs of Eq. (3) with time 294 scales less than 6 h and radial scales less than 8 km are filtered out. Similarly, the simulated TC IR and  $dRMW/d\tau$  at each time  $\tau_0$  are calculated as the change rate of TC intensity and RMW 295

between  $\tau_0$  and  $\tau_0 + 6h$ , and all budget terms for both tangential wind and RMW tendencies are their averages between  $\tau_0$  and  $\tau_0 + 6h$  using the model outputs at 10-min intervals.

298 Figures 5 and 6 show the budget results using Eqs. (3) and (5), respectively. We can see that 299 both the simulated TC IR and RMW contraction rate in all experiments are well captured by the 300 budgets based on Eq. (3) and Eq. (5) (Figs. 5a,b; Figs. 6a,b), implying that the model outputs at 301 10-min intervals can be used to examine the time evolution of TC intensity and RMW contraction. 302 As shown in previous studies (e.g., Zhang et al. 2001; Heng and Wang 2016; Li et al. 2019, 2020a), 303 the radial vorticity flux dominates the inner-core spinup in the boundary layer (Figs. 5c,d), and all 304 the vertical advection and diffusion terms including surface friction contribute to the inner-core 305 spindown in the boundary layer (Figs. 5e-j). This means that the TC IR largely depends on the radial vorticity flux, which is determined by the secondary circulation  $(u_m)$ , TC intensity  $(V_m)$ , and 306 307 the size of RMW (Eq. 3). Note that after the peak of IR, the simulated IR decreases (Figs. 5a,b) 308 although the radial vorticity flux still rapidly increases (Figs. 5c,d). This is because the surface 309 friction, which is approximately proportional to the square of the maximum tangential wind, also 310 increases rapidly with the rapid increase of TC intensity (not shown). As a result, the associated 311 negative contribution of vertical diffusion and vertical advection in the boundary layer to the IR 312 also rapidly increases (Figs. 5e,f,i,j; Li et al. 2020a).

From Figs. 6c–j, we can see that the radial vorticity flux forcing dominates the RMW contraction rate (Figs. 6c,d) and the vertical advection forcing and the diffusion forcings (including surface friction) tend to reduce RMW contraction (Figs. 6e–j), consistent with the budget analyses of Li et al. (2019). During the late stage of TC intensification (after the peak of TC IR), the vertical advection forcing (Figs. 6e,f) decreases, the horizontal diffusion forcing (Figs. 6g,h) increases with the increasing TC intensity and contributes largely to the cessation of RMW contraction, and the vertical diffusion forcing including surface friction (Figs. 6i,j) also plays an important role in 320 prohibiting RMW contraction. Note that although the direct effect of the vertical diffusion 321 including surface friction on RMW contraction discussed here is negative, its net effect may be 322 positive because surface friction would enhance the low-level radial vorticity flux and its negative 323 radial gradient (Stern et al. 2015; Li et al. 2019, 2020a). Another striking feature of the budgets is 324 that although the contribution of vertical diffusion including surface friction to the IR is much 325 larger than that of horizontal diffusion (Figs. 5g-j), the contribution of vertical diffusion forcing to 326 the RMW contraction rate is smaller than that of horizontal diffusion forcing (Figs. 6g-j) during 327 the late stage of TC intensification. This is because the contours of the horizontal diffusion forcing 328 are more parallel to the RMW than that of the vertical diffusion forcing (not shown, cf. Li et al. 329 2019), and thus the radial gradient of the horizontal diffusion is larger than that of the vertical 330 diffusion as indicated in Li et al. (2019).

331 To address why the simulated rapid contraction precedes the RI, we mainly focus on the early 332 intensification stage, the period after the initial adjustment and prior to the peak of the simulated 333 IR, which includes the rapid contraction phase (cf. Figs. 5a,b; Figs. 6a,b). During the early 334 intensification stage, the surface friction is relatively small because of the weak surface wind speed, 335 and the simulated IR is dominated by the radial vorticity flux (Figs. 5c,d), which increases with the 336 increasing TC intensity and the decreasing RMW. However, as the TC evolves, the increase in 337 intensity and decrease in RMW would increase the curvature of the radial distribution of tangential 338 wind and thus is unfavorable for the RMW contraction as implied by Eqs. (4) or (5) (Stern et al. 339 2015; Li et al. 2019) although the radial vorticity flux forcing, which dominates the RMW 340 contraction, increases as the TC intensifies during the early intensification stage (Figs. 6c,d). 341 During the rapid RMW contraction phase with slow TC intensification, the TC intensity is 342 relatively weak and the RMW is relatively large (Figs. 7a,b), giving rise to a small curvature (Figs. 343 7e,f). As a result, even though the negative gradient of radial vorticity flux and thus the negative 344 gradient of net tangential wind tendency near the RMW is moderate (Figs. 6c.d: Figs. 7e.f), it can 345 induce a large RMW contraction rate (Figs. 6a,b). However, the small radial vorticity flux near the 346 RMW (Figs. 5c,d) is unfavorable for the RI as implied by Eq. (3) (Figs. 5a,b). With the 347 intensification of the TC and the continuous decrease of the RMW, the curvature of the radial 348 distribution of tangential wind near the RMW increases continuously (Figs. 7e,f), leading to the slowdown of the RMW contraction (Figs. 6a,b). The small RMW and strong intensity during the 349 350 slow contraction phase (Figs. 7c,d) are more favorable for RI than the relatively large RMW and 351 weak intensity during the rapid RMW contraction phase (Figs. 7a,b).

352 In addition to the TC intensity and size of RMW, the radial vorticity flux also depends on the 353 low-level inflow as inferred from Eq. (3), which is partly controlled by diabatic heating in the 354 eyewall. It is our interest to further compare the TC structure and diabatic heating during different 355 phases. Figure 8 shows the composite TC structures and diabatic heating rates at times when the 356 RMW contraction rate (Fig. 8a) and the IR (Fig. 8b) reach their respective peaks among all 357 experiments. Note that the composites are conducted here because the TC structures are similar 358 among all experiments for both times (Figs. 7a-d). As expected, diabatic heating in the inner core 359 and thus the low-level inflow are much weaker during the rapid RMW contraction phase than 360 during the RI phase. Therefore, based on Eq. (3), during the rapid RMW contraction phase (Fig. 361 8a), the weak low-level inflow associated with weak diabatic heating inside the RMW, together 362 with the relatively large RMW and weak TC intensity (Figs. 7a,b), results in weak low-level radial 363 vorticity flux near the RMW and thus small IR. Note that here the weak diabatic heating is itself 364 largely a reflection of the weak intensity. As the TC intensifies, the enhanced diabatic heating and 365 thus low-level inflow, and the increasing TC intensity and decreasing RMW favor the high IR 366 during the RI phase (Fig. 8b).

367

Furthermore, we can see from Figs. 8a and 8b that the maximum heating rate is located further

368 inside the RMW during the rapid RMW contraction phase than during the RI period. This can be 369 understood based on the boundary layer dynamics discussed in Kepert (2017), who revealed that 370 the radial location of the evewall updraft is a function of the RMW. The evewall updraft in a larger 371 RMW TC would be located more inside the RMW than in a smaller RMW TC (see also Li and 372 Wang 2021). Consistently, the radial inflow also penetrates more inside the RMW in a large RMW 373 TC than in a small RMW TC. Because the relatively high absolute vorticity inside the RMW, the 374 more inwardly penetrated inflow would lead to relatively large tangential wind tendency also more 375 inside the RMW, as implied by the balanced vortex dynamics (Shapiro and Willoughby 1982; 376 Schubert and Hack 1982). This favors the contraction of the RMW but is less favorable for TC 377 intensification because the tangential wind tendency near the RMW is relatively small. This point 378 is illustrated by comparing the radial distributions of tangential wind and its tendency in the 379 boundary layer at the time of the peak in RMW contraction rate and the time of the peak in IR 380 shown in Figs. 8c and 8d. During the rapid RMW contraction phase, larger tangential wind 381 tendencies are well inside the RMW, which contribute relatively small tangential wind increasing 382 tendency near the RMW (Fig 8c) but contribute largely to the RMW contraction rate because of 383 the small curvature of the radial distribution of tangential wind. However, with the decreasing 384 RMW following the rapid contraction phase, the diabatic heating rate is larger and becomes closer 385 to the RMW (Fig. 8b), this would induce much larger tangential wind tendencies near the RMW 386 (Fig. 8d), contributing largely to the TC IR but little to the RMW contraction rate due to the large 387 curvature of the radial distribution of tangential wind. These results further indicate that in addition 388 to the dependence of the radial vorticity flux on TC intensity and the size of RMW (or the inertial stability in the inner core of the TC), the dependence of the strength and radial location of diabatic 389 390 heating in the eyewall relative to the RMW on the size of the RMW and TC intensity is also a key 391 to the simulated lag of RI behind the rapid RMW contraction. This latter dependence results

392 primarily from the unbalanced boundary layer dynamics as discussed above (Kepert 2017) while 393 the former dependence results mainly from the balanced vortex dynamics. Therefore, both balanced 394 and unbalanced dynamics work cooperatively to determine the relationship between the simulated 395 rapid RMW contraction and TC intensification.

# **6.** Conclusions and discussion

397 In this study, the relationship between the TC IR and RMW contraction rate is revisited 398 through ensemble numerical experiments using an axisymmetric cloud model. The focus is on the 399 asynchrony between the rapid contraction and RI, namely, the rapid contraction precedes the RI, 400 which has been reported in previous observational and numerical studies (e.g., Corbosiero et al. 401 2005; Chen et al. 2011; Stern et al. 2015; Wu and Ruan 2021) but the understanding of the involved 402 dynamics is incomplete. Most previous studies explained the relationship between RMW 403 contraction and TC intensification qualitatively based on the assumption of AAM-conservation 404 following the RMW (e.g., Holland and Merrill 1984; Pu et al. 2009; Chen et al. 2011, Kilrov et al. 405 2016). Under this assumption, the IR would depend primarily on the RMW contraction rate. This 406 hypothesis has been invalidated in a recent study based on high-resolution numerical simulations 407 by Stern et al. (2015).

408 Our results show that the AAM following the RMW is not conserved and can vary by ~30– 409 50% no matter during the rapid contraction phase or during the RI phase, consistent with the results 410 of Stern et al. (2015). As a result, the IR of the simulated TC could not be diagnosed by the RMW 411 contraction rate only because the RMW is not a material surface. During the rapid contraction 412 phase, the RMW shifts rapidly inward across the AAM surfaces, namely the AAM following the 413 RMW decreases with relatively small IR of the simulated TC. However, during the slow RMW 414 contraction phase, the AAM surfaces moves inward passing through the RMW, corresponding to 415 the high IR. Therefore, the AAM conservation following the RMW as assumed in previous TC 416 intensification theory (e.g., Emanuel 2012) is not valid as also pointed out by Stern et al. (2015). 417 The AAM-based diagnostics provides little physical/dynamical insights into the understanding of 418 the relationship between TC IR and RMW contraction rate because the rate of change in AAM 419 following the RMW needs to be explained.

420 To understand why the rapid RMW contraction precedes the RI in the simulated TCs, we 421 performed both the tangential wind budget and the budget for the rate of change in RMW in the 422 simulated TCs. Results show that during the rapid RMW contraction phase, the slow intensification 423 is primarily due to the small radial vorticity flux near the RMW, while the rapid RMW contraction 424 is largely due to the moderate negative radial gradient of the radial vorticity flux and the small 425 curvature of the radial distribution of tangential wind near the RMW. The relatively small radial 426 vorticity flux near the RMW results primarily from the weak TC intensity and large size of RMW 427 (and thus weak inner-core inertial stability) and weak low-level inflow. The moderate radial 428 gradient of the radial vorticity flux results primarily from the more inwardly displaced eyewall 429 updraft (and thus convection) relative to the RMW due to more inwardly penetrated low-level 430 inflow controlled by the boundary layer dynamics as proposed by Kepert (2017). With the decrease 431 of RMW due to the rapid contraction and the increase of TC intensity, the inner-core inertial 432 stability increases and the low-level inflow becomes less penetrative relative to the RMW, resulting 433 in the eyewall updraft less inwardly displaced. As a result, the radial vorticity flux increases with 434 the maximum slightly inside and close to the RMW, giving rise to high intensification rate and thus 435 RI of the simulated TCs. Although during the RI phase, the negative radial gradient of the radial 436 vorticity flux near the RMW is also large, the RMW contraction is greatly suppressed by the rapidly 437 increasing curvature of the radial distribution of tangential wind.

438

Note that this study only discusses one scenario regarding the relationship between the rapid

439 RMW contraction and TC RI, i.e., with the rapid RMW contraction preceding RI. Wu and Ruan 440 (2021) showed that there are also considerable portion of TCs experiencing rapid RMW contraction 441 and RI simultaneously in observations. They found that those TCs are characterized by the 442 combination of large RMW and sufficient convective heating close to the storm center (cf. their 443 Fig. 12c). As shown in our idealized simulations, the large RMW always occurs during the earlier 444 stage of the intensification period, however, during which the convective heating is often weak or 445 not "sufficient". We thus hypothesize that in idealized conditions, it is easier for the rapid RMW 446 contraction to precede RI than in nature, but this needs to be further examined in a future study. In 447 addition, results from this study do not mean that there must be an RI (often quantitatively defined 448 as the near-surface maximum wind speed increase more than 15 m s<sup>-1</sup> per day; Kaplan and DeMaria 449 2003) following rapid RMW contraction, as mentioned earlier and in Wu and Ruan (2021). This is 450 because in nature there are many deleterious environmental effects to TC intensification, such as 451 vertical wind shear and oceanic cold core eddies, which are not included in our idealized 452 simulations.

453 Note also that as our main conclusions in this study are based on the axisymmetric model 454 experiments, it is unclear whether and how much the asymmetric eddy processes may contribute 455 to the relationship between the RMW contraction and TC intensification rate. Nevertheless, since 456 most of our results are consistent with previous three-dimensional model simulations (e.g., Stern 457 et al. 2015; Li et al. 2019) and observations (Wu and Ruan 2021), our main conclusions could be also applicable to three-dimensional simulations. In addition, this study only provides some 458 459 dynamical/physical insights into the relationship between the rapid RMW contraction and TC RI qualitatively. It will be a good topic for a future study to quantify the relationship, e.g., the lag time 460 between the rapid RMW contraction and TC RI, so that the results can provide guidance to real-461 462 time TC intensity forecasts.

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#### 468 **References**

- Bryan, G. H., and J. M. Fritsch, 2002: A benchmark simulation for moist nonhydrostatic numerical
  model. *Mon. Wea. Rev.*, 130, 2917–2928, https://doi.org/10.1175/1520-0493(2002)130,2917:
  ABSFMN.2.0.CO;2.
- Chen, H., D. Zhang, J. Carton, and R. Atlas, 2011: On the rapid intensification of Hurricane Wilma
  (2005). Part I: Model prediction and structural changes, *Wea. Forecasting*, 26, 885901, https://doi.org/10.1175/WAF-D-11-00001.1.
- 475 Corbosiero, K. L., J. Molinari, and M. L. Black, 2005: The structure and evolution of Hurricane
  476 Elena (1985). Part I: Symmetric intensification, *Mon. Wea. Rev.*, 133, 2905-2921,
  477 https://doi.org/10.1175/MWR3010.1.
- 478 Črnivec, N., R. K. Smith, and G. Kilroy, 2016: Dependence of tropical cyclone intensification rate
- 479 on sea-surface temperature. *Quart. J. Roy. Meteor. Soc.*, 142, 1618–1627, https://
  480 doi.org/10.1002/qj.2752.
- Donelan, M. A., B. K. Haus, N. Reul, W. J. Plant, M. Stiassnie, H. C. Graber, O. B. Brown, and E.
  S. Saltzman, 2004: On the limiting aerodynamic roughness of the ocean in very strong winds. *Geophys. Res. Lett.*, **31**, L18306, https://doi.org/10.1029/ 2004GL019460.
- Emanuel, K. A., 2012: Self-stratification of tropical cyclone outflow. Part II: Implications for storm
  intensification. J. Atmos. Sci., 69, 988–996, https://doi.org/10.1175/JAS-D-11-0177.1.
- Heng, J., and Y. Wang, 2016: Nonlinear response of a tropical cyclone vortex to prescribed eyewall
  heating with and without surface friction in TCM4: Implications for tropical cyclone
  intensification. J. Atmos. Sci., 73, 1315–1333, https://doi.org/10.1175/JAS-D-15-0164.1.
- Holland, G. J., and R. T. Merrill, 1984: On the dynamics of tropical cyclone structural changes. *Quart. J. Roy. Meteor. Soc.*, 110, 723–745, https://doi.org/10.1002/qj.49711046510.
- Judt, F., and S. S. Chen, 2013: Reply to "Comments on 'Convectively generated potential vorticity
  in rainbands and formation of the secondary eyewall in Hurricane Rita of 2005." *J. Atmos. Sci.*, 70, 989–992, https://doi.org/10.1175/JAS-D-12-0151.1.
- Kaplan, J., and M. DeMaria, 2003: Large-scale characteristics of rapidly intensifying tropical
  cyclones in the North Atlantic basin. *Wea. Forecasting*, 18, 1093–1108,
  https://doi.org/10.1175/1520-0434(2003)018,1093: LCORIT.2.0.CO;2.
- Kepert, J. D., 2017: Time and space scales in the tropical cyclone boundary layer, and the location
  of the eyewall updraft. *J. Atmos. Sci.*, 74, 3305–3323, https://doi.org/10.1175/JAS-D-17-

499 0077.1.

- Kilroy, G., R. K. Smith, and M. T. Montgomery, 2016: Why do model tropical cyclones grow
  progressively in size and decay in intensity after reaching maturity? *J. Atmos. Sci.*, 73, 487–
  503, https://doi.org/10.1175/JAS-D-15-0157.1.
- Li, T.-H., and Y. Wang, 2021: The role of boundary layer dynamics in tropical cyclone
  intensification. Part II: Sensitivity to initial vortex structure. *J. Meteor. Soc. Japan*, 99, 537–
  554, https://doi.org/10.2151/jmsj.2021-028.
- Li, Y., Y. Wang, and Y. Lin, 2019: Revisiting the dynamics of eyewall contraction of tropical
  cyclones. J. Atmos. Sci., 76, 3229–3245, https://doi.org/10.1175/JAS-D-19-0076.1.
- Li, Y., Y. Wang, and Y. Lin, 2020a: How much does the upward advection of the supergradient
  component of boundary layer wind contribute to tropical cyclone intensification and maximum intensity? J. Atmos. Sci., 77, 2649–2664, https://doi.org/ 10.1175/JAS-D-19-0350.1.
- 511 Li, Y., Y. Wang, Y. Lin, and R. Fei, 2020b: Dependence of superintensity of tropical cyclones on
- 512 SST in axisymmetric numerical simulations. *Mon. Wea. Rev.*, 148, 4767-4781,
   513 https://doi.org/10.1175/MWR-D-20-0141.1.
- Peng, K., R. Rotunno, and G. H. Bryan, 2018: Evaluation of a time- dependent model for the
  intensification of tropical cyclones. *J. Atmos. Sci.*, 75, 2125–2138,
  https://doi.org/10.1175/JAS-D- 17-0382.1.
- Pu, Z., X. Li, and E. J. Zipser, 2009: Diagnosis of the initial and forecast errors in the numerical
  simulation of the rapid intensification of Hurricane Emily (2005). *Wea. Forecasting*, 24,
  1236–1251, https://doi.org/10.1175/2009WAF2222195.1.
- Qin, N., D.-L. Zhang, and Y. Li, 2016: A statistical analysis of steady eyewall sizes associated with
   rapidly intensifying hurricanes. *Wea. Forecasting*, **31**, 737–742, https://doi.org/10.1175/
   WAF-D-16-0016.1.
- Rotunno, R., and K. A. Emanuel, 1987: An air–sea interaction theory for tropical cyclones. Part II:
  Evolutionary study using a nonhydrostatic axisymmetric numerical model. *J. Atmos. Sci.*, 44,
  542–561, https://doi.org/10.1175/1520-0469(1987)044,0542: AAITFT.2.0.CO;2.
- 526 Schubert, W. H., and J. J. Hack, 1982: Inertial stability and tropical cyclone development. J. Atmos.
- 527 *Sci.*, **39**, 1687–1697, https:// doi.org/10.1175/1520-0469(1982)039,1687:ISATCD.2.0.CO;2.
- 528 Shapiro, L. J., and H. E. Willoughby, 1982: The response of balanced hurricanes to local sources
- 529 of heat and momentum. J. Atmos. Sci., **39**, 378–394, https://doi.org/10.1175/1520-0469(1982)

- 530 039,0378:TROBHT.2.0.CO;2.
- Shimada, U., and T. Horinouchi, 2018: Reintensification and eyewall formation in strong shear: A
  case study of Typhoon Noul (2015). *Mon. Wea. Rev.*, 146, 2799-2817, https://doi.org/10.1175/MWR-D-18-0035.1.
- Stern, D. P., J. R. Brisbois, and D. S. Nolan, 2014: An expanded dataset of hurricane eyewall sizes
  and slopes. *J. Atmos. Sci.*, 71, 2747–2762, https://doi.org/10.1175/JAS-D-13-0302.1.
- Stern, D. P., J. L. Vigh, D. S. Nolan, and F. Zhang, 2015: Revisiting the relationship between
  eyewall contraction and intensification. *J. Atmos. Sci.*, **72**, 1283–1306, https://doi.org/
  10.1175/JAS-D-14-0261.1.
- Thompson, G., P. R. Field, R. M. Rasmussen, and W. D. Hall, 2008: Explicit forecasts of winter
  precipitation using an improved bulk microphysics scheme. Part II: Implementation of a new
  snow parameterization. *Mon. Wea. Rev.*, 136, 5095–5115,
  https://doi.org/10.1175/2008MWR2387.1.
- 543 Willoughby, H. E., 1990: Temporal changes of the primary circulation in tropical cyclones. *J.*544 *Atmos. Sci.*, 47, 242–264, https://doi.org/10.1175/1520-0469(1990)047,0242:TCOTPC.
  545 2.0.CO;2.
- Willoughby, H. E., J. A. Clos, and M. G. Shoreibah, 1982: Concentric eye walls, secondary wind
  maxima, and the evolution of the hurricane vortex. *J. Atmos. Sci.*, **39**, 395–411,
  https://doi.org/10.1175/1520-0469(1982)039,0395:CEWSWM.2.0.CO;2.
- Wang, Y., Y. Li, J. Xu, Z.-M. Tan, and Y. Lin, 2021: The intensity-dependence of tropical cyclone
  intensification rate in a simplified energetically based dynamical system model. *J. Atmos. Sci.*, **78**, (in press), https://doi.org/10.1175/JAS-D-20-0393.1.
- Wood, V. T., and L. W. White, 2011: A new parametric model of vortex tangential-wind profiles:
  Development, testing, and verification. *J. Atmos. Sci.*, 68, 990–1006, https://doi.org/
  10.1175/2011JAS3588.1.
- Wu, Q., and Z. Ruan, 2021: Rapid contraction of the radius of maximum tangential wind and rapid
  intensification of a tropical cyclone. *J. Geophys. Res. Atmos.*, **126**, e2020JD033681,
  https://doi.org/10.1029/2020JD033681.
- 558 Zhang, D., Y. Liu, and M. K. Yau, 2001: A multiscale numerical study of Hurricane Andrew (1992).
- 559 Part IV: Unbalanced flows. *Mon. Wea. Rev.*, **129**, 92–107, https://doi.org/10.1175/ 1520560 0493(2001)129,0092:AMNSOH.2.0.CO;2.

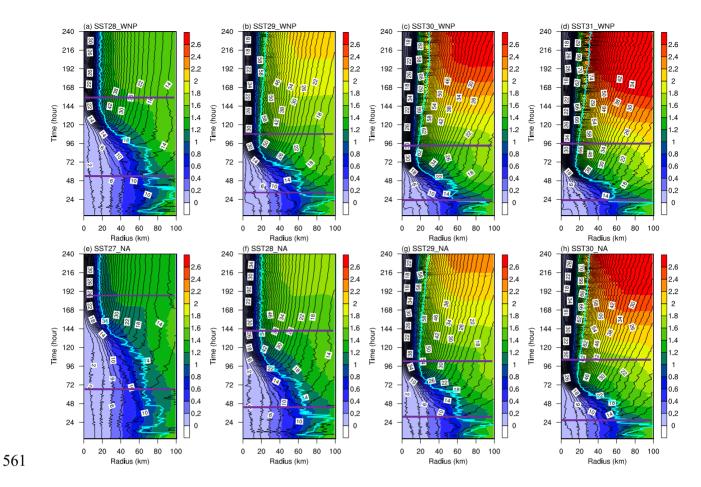


Figure 1. Radius-time Hovmöller plots of hourly ensemble-mean absolute angular momentum
(shades; 10<sup>6</sup> m<sup>2</sup> s<sup>-1</sup>) and tangential wind (contours; at an interval of 2 m s<sup>-1</sup>) at 2-km height,
overlaid by the radial position of RMW (cyan curve), and the two purple horizontal lines mark
the intensification period after the initial adjustment in each experiment.

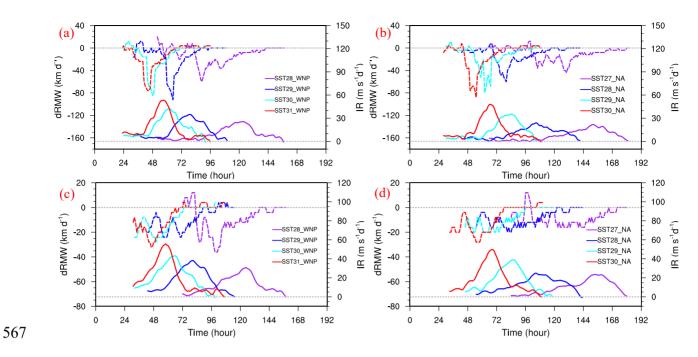
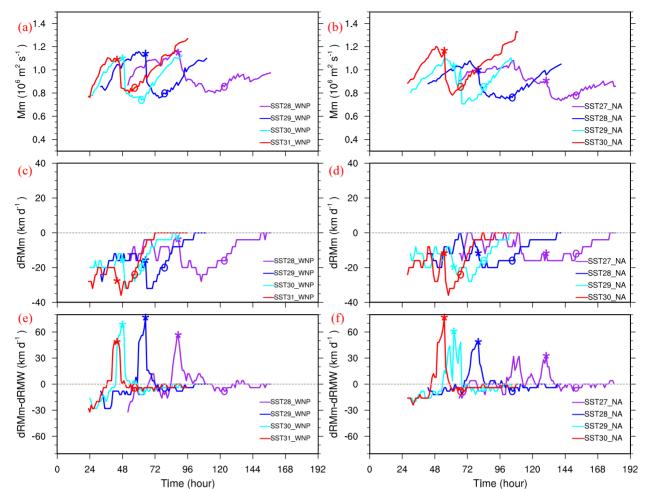


Figure 2. Time series of tendency of RMW (dRMW; dashed curves) and tendency of maximum
tangential wind speed (IR; solid curves) (a,b) at 2-km height and (c,d) averaged below 500 m.
Only results during the primary intensification period after the initial adjustment are shown.



572 Time (hour)
573 Figure 3. (a,b) Time series of absolute angular momentum following the RMW at 2-km height.
574 (b,c) Time series of change rate in the radius of absolute angular momentum surface from the
575 current RMW to the radius at the subsequent 6 h for the same absolute angular momentum
576 surface (dRMm). (e,f) Time series of the difference between dRMm and dRMW. Only results
577 during the primary intensification period after the initial adjustment are shown. The asterisk and
578 circle on each curve mark the time of the peak RMW contraction rate and the time of the peak
579 simulated IR, respectively.

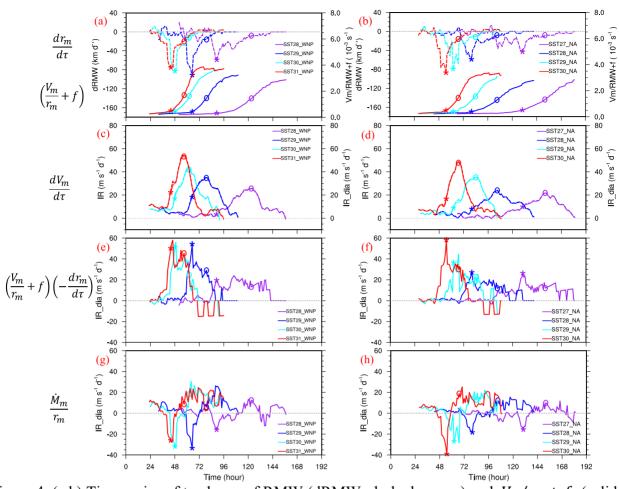
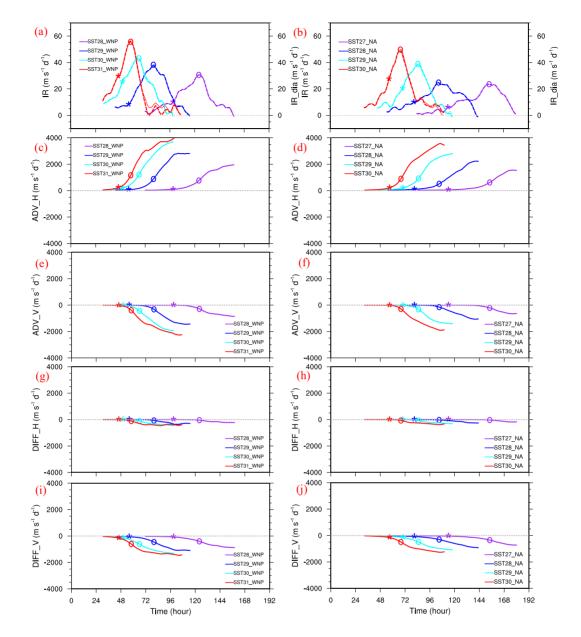
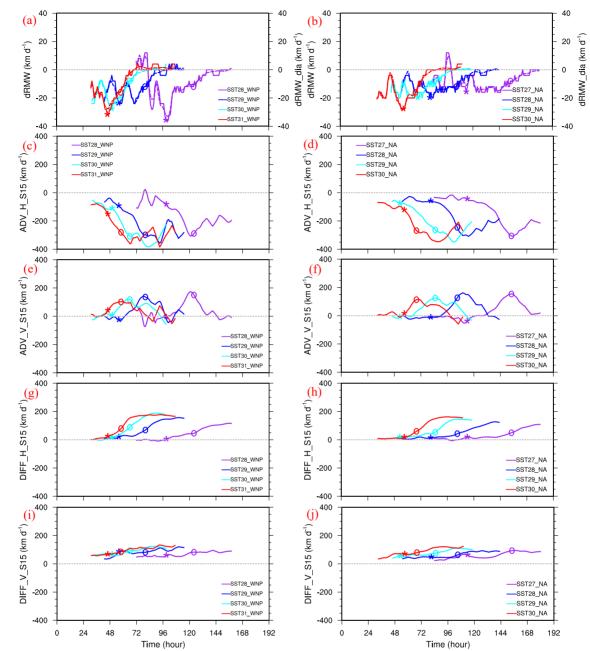


Figure 4. (a,b) Time series of tendency of RMW (dRMW; dashed curves) and  $V_m/r_m + f$  (solid curves). (c,d) Time series of the simulated IR (solid curves) and diagnosed IR based on Eq. (2) (dashed curves). Note that the solid curves are mostly under the dashed curves. (e,f) Times series of  $(V_m/r_m + f)(-dr_m/d\tau)$ . (g,h) Times series of  $\dot{M}_m/r_m$ . All results are from 2-km height, and only results during the primary intensification period after the initial adjustment are shown. The asterisk and circle on each curve mark the time of the peak RMW contraction rate and the time of the peak simulated IR, respectively.



589

Figure 5. (a,b) Time series of the simulated IR (solid curves) and diagnosed IR based on Eq. (3) (dashed curves) averaged below 500 m. Note that the solid curves are mostly under the dashed curves. (c–j) Time series of the diagnosed IR averaged below 500 m due to (c,d) radial vorticity flux, (e,f) vertical advection, (g,h) radial diffusion, and (i,j) vertical diffusion including surface friction. Only results during the primary intensification period after the initial adjustment are shown. The asterisk and circle on each curve mark the time of the peak RMW contraction rate and the time of the peak simulated IR, respectively.



597 Figure 6. (a,b) Time series of the simulated tendency of RMW (solid curves) and diagnosed 598 599 tendency of RMW based on Eq. (5) (dashed curves) averaged below 500 m. (c-j) Time series 600 of the contributions of the radial gradient of tangential wind tendency to the RMW tendency 601 averaged below 500 m due to (c,d) radial vorticity flux, (e,f) vertical advection, (g,h) radial 602 diffusion, and (i,j) vertical diffusion including surface friction. Only results during the primary 603 intensification period after the initial adjustment are shown. The asterisk and circle on each 604 curve mark the time of the peak RMW contraction rate and the time of the peak simulated IR, 605 respectively.

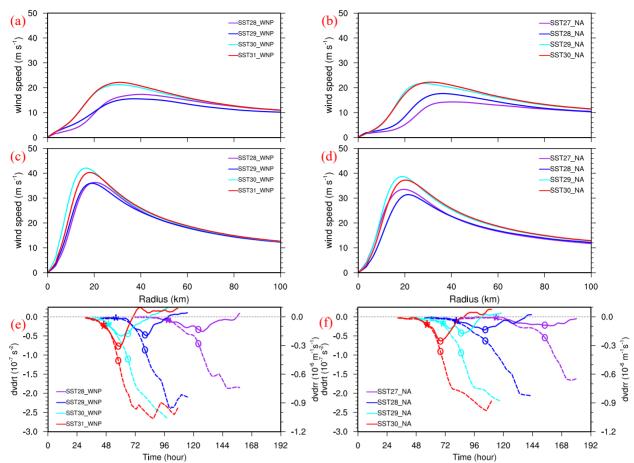
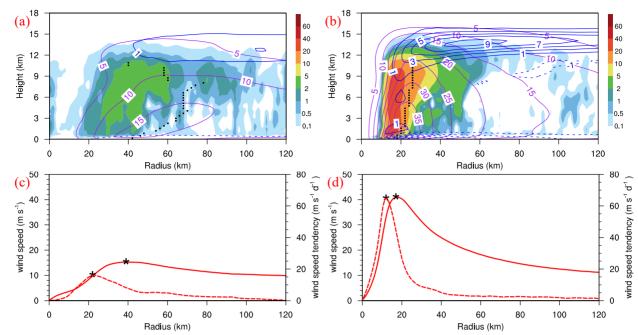


Figure 7. (a,b) Radial profile of tangential wind speed averaged below 500 m at the time of the peak RMW contraction rate. (c,d) As in (a,b), but at the time of the peak simulated IR. (e,f) Time series of the radial gradient of tangential wind tendency (dvdrt ;solid curves) and the curvature of the radial distribution of tangential wind at the RMW (dvdrr; dashed curves). The asterisk and circle on each curve mark the time of the peak RMW contraction rate and the time of the peak simulated IR, respectively.



<sup>614</sup> Figure 8. (a,b) Radial-vertical cross sections of the composite diabatic heating rate (shading; K h<sup>-</sup> <sup>615</sup> <sup>1</sup>), tangential wind speed (purple contours; at an interval of 5 m s<sup>-1</sup>), and radial wind speed with <sup>617</sup> negative value dashed (blue contours; at an interval of 2 m s<sup>-1</sup>) at the time of (a) the peak RMW <sup>618</sup> contraction rate and (b) the peak simulated IR. The dotted black line shows the location of the <sup>619</sup> RMW. (c,d) Radial profile of tangential wind speed (solid) and tendency of tangential wind speed <sup>620</sup> in the subsequent 6 h (dashed) averaged below 500 m at the time of (c) the peak RMW contraction <sup>621</sup> rate and (d) the peak simulated IR. The asterisk marks the radius of the corresponding peak value.