

1 **Title: Structure and Density of Silicon Carbide to 1.5 TPa: Implications for Extrasolar**
2 **Planets**

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12 **Abstract**

13 There has been considerable recent interest in the high-pressure behavior of silicon
14 carbide, a potential major constituent of carbon-rich exoplanets. In this work, the atomic-level
15 structure of SiC was determined through *in situ* X-ray diffraction under laser-driven ramp
16 compression up to 1.5 TPa; stresses more than seven times greater than previous static and shock
17 data. Over this stress range we confirm the persistence of the B1-type structure and constrain its
18 equation of state (EOS). Using this data we have determined the first experimentally based mass-
19 radius curves for a hypothetical pure SiC planet. Interior structure models were constructed for
20 planets consisting of a SiC-rich mantle and iron-rich core. Carbide planets are found to be ~10%
21 less dense than corresponding terrestrial planets.

22 **Introduction**

23 A large number of extrasolar planets (>4000) have been discovered in recent years. The
24 chemical compositions of their host stars are diverse, implying that these planets may exhibit a
25 wider range of composition than found in our solar system^{1–3}. Carbon planets are one of the
26 possible types of exoplanets, proposed to form around host stars having a high carbon-to-oxygen

27 ratios^{4,5}. High C/O ratios above 0.8 may alter the condensation sequence in protoplanetary disks
28 and result in the formation of planets with a mantle dominated by C, SiC, and other carbides
29 rather than silicates^{1,5,6}. Although conflicting results have been reported in the literature, recent
30 measurements suggest that a small number of host stars may have high C/O ratios⁷⁻¹¹. In
31 particular, a recent survey⁸ of the chemical composition of a large number of stars found that
32 ~1% had C/O > 0.8 (for comparison, the C/O ratio of the Sun is 0.55¹²). In view of the very large
33 number of planets expected to exist in our galaxy, these results imply there may exist many
34 carbon-rich planets of various sizes. Planetary formation processes could further enhance
35 formation of carbon-rich planets even for host-star C/O ratios as low as 0.65¹³. Moreover, carbon
36 planets are of special interest even if found in only a small percentage of planets as they would
37 represent a highly exotic planetary style that would likely have a very different surface
38 environment, tectonic style, heat flow, and potential for habitability than typical silicate
39 planets¹⁴.

40 Although silicon carbide (SiC) is expected to be a major constituent of this novel
41 planetary type, its properties at the very high pressures of exoplanetary interiors are poorly
42 constrained. Shock-wave experiments have reported a phase transition in SiC with a large
43 volume decrease (~20 %) near 100 GPa¹⁵⁻¹⁷, but the data are restricted to less than 210 GPa¹⁷.
44 Under static compression in a diamond anvil cell (DAC), the phase transformation from B3 to
45 B1 was reported to occur ~60-70 GPa^{18,19}. The equation of state of B1 SiC was determined up to
46 200 GPa^{18,20}. However, the existing studies are in strong disagreement when the reported
47 equations of state are extrapolated to the higher pressures relevant to exoplanets that are several
48 times more massive than Earth.

49 Dynamic ramp compression is a technique to access a low-temperature path, in which
50 materials can be compressed in the solid state to extreme pressures, avoiding the melting that
51 occurs under shock loading²¹. By combining ramp-compression experiments with *in situ* X-ray
52 diffraction (XRD), the atomic-level structure of materials can be determined under extreme
53 conditions reaching into the terapascal regime (1 TPa = 10 million atmospheres)^{21–25}.

54 In this study, the structure and equation of state of SiC were examined using laser-driven
55 ramp compression and *in situ* XRD up to ultrahigh pressures of 1.5 TPa. We present the first
56 direct constraints on the structure and equation of state of SiC under dynamic loading at
57 conditions relevant to the deep interior of extrasolar planets.

58

59 **Results**

60 Ramp-compression experiments were conducted using the Omega-60 and Omega-EP
61 lasers at the Laboratory for Laser Energetics (University of Rochester, NY). The high-intensity
62 Omega lasers were used to both drive a ramp compression wave into the sample on ~6–10
63 nanosecond timescales, and to generate a ~1 or 2-ns duration quasi-monochromatic X-ray
64 source. The X-rays were used to record diffraction peaks from the compressed sample shown in
65 Figs. 1 and 2, and stresses were determined from interferometric wave profile measurements
66 shown in Figs. 3 and 4 (see “Experimental methods” for details). The experimental data are
67 summarized in Supplementary Material Table S1.

68 SiC was compressed to stresses ranging from 80(4) GPa to 1507(64) GPa. Figure 5
69 shows the observed lattice *d*-spacings measured by X-ray diffraction. The low-pressure
70 zincblende-type (B3) phase was observed at 80(4) GPa, the lowest stress achieved in this work.

71 This measured stress is higher than the reported B3-B1 transition stress (~60–70 GPa) from
72 laser-heated diamond cell experiments^{18–20} but below the transition stress observed in shock
73 wave studies (~100 GPa)¹⁵, indicating that kinetics and/or stress heterogeneity play a role in this
74 transformation under dynamic loading. For shots between 140(5) GPa and 371(11) GPa, two or
75 three diffraction peaks were observed that could be indexed as the (111), (002), and (022)
76 reflections of the rocksalt-type B1 phase of SiC (Fig. 5). Thus, the phase transformation to the
77 B1 phase occurs between 80 GPa to 140 GPa under ramp loading. These results are comparable
78 to previous continuum gas-gun experiments¹⁵, and *in situ* XRD results from laser-driven shock
79 experiments, which observed the formation of B1 SiC at ~114 GPa¹⁷. At stresses between
80 433(15) and 621(20) GPa, only one diffraction peak from the sample was observed which could
81 be indexed as the B1 (002) peak, the most intense reflection for this structure. (Fig. 5). The B1
82 (111) peak could not be observed in this stress range due to overlap with diffraction from the
83 pinhole material (Fig. S1). At stresses above 700 GPa, 1-2 diffraction peaks are again observed
84 and these can be indexed to B1 (111) and (002). SiC was compressed to ~9.97(34) to 10.28(20)
85 g/cm³ at peak stress, yielding a 3-fold increase over its ambient density value (Fig. S3 and Table
86 S2). There is no evidence for any further transformations²⁶ or dissociation of SiC (Figs S1 and
87 S2). At stresses above 371 GPa, the increased scatter in densities is likely due to the high X-ray
88 background emitted by the drive plasma for higher stress experiments (Fig. 5).

89 Ramp compression thermodynamically mimics a series of weak shocks, thereby reducing
90 the excess heat arising from increased entropy during compression. However, irreversible work
91 and heating as a result of dissipation and material strength occur during ramp loading, so the
92 loading path is generally considered quasi-isentropic. We have applied to a series of correction to
93 transform from the measured longitudinal stress to the corresponding principal isentrope and

94 isotherm (see “Experimental methods” for details). For our ramp-compressed B1 SiC data, the
95 reduced isentropes and isotherms were fit to a third-order Birch-Murnaghan (BM) equation of
96 state to place constraints on the corresponding bulk modulus, K_0 , and its pressure derivative, K_0' .
97 Considering the uncertainties in the thermodynamic and strength parameters of B1 SiC, however,
98 the differences between reduced isentropes, isotherms and ramp-compressed EOS are small.
99 Here we have assumed a principal isentrope for B1 SiC constructed assuming a yield strength,
100 $Y=20$ GPa, and a Gruneisen parameter, $\gamma_0=1$. The calculated energy and temperature along the
101 Hugoniot, ramp and principal isentropic paths are shown in Figures S6 and S7. The zero-pressure
102 volume, V_0 , of the B1 phase was fixed at 66.3 \AA^3 (Ref.¹⁸). The resultant fit parameters of the
103 reduced isentrope are $K_0 = 272(16)$ GPa and $K_0' = 3.79(11)$. If we fit the data to a Vinet equation
104 of state, we obtain the following fit parameters: $K_0 = 252(21)$ GPa and $K_0' = 4.3(3)$ (Fig. 5 and
105 Table S3).

106 Our results are compared with extrapolations of previous 300-K static equations of state
107 data (black and blue lines) and first-principles calculations at 0 K for B1 SiC (Fig. 5, Table S3).
108 The two previous static equations of state^{18,20} diverge when extrapolated above ~ 600 GPa. Our
109 data are in good agreement with the extrapolation of the results of ref.²⁰ but not ref.¹⁸.
110 Theoretical calculations predict a wide range of stress-density relations for SiC²⁶⁻³³ (Fig. 5, Table
111 S3), but it should be noted that our results are in good agreement with the recent theoretical
112 calculations of ref.²⁶. Thus, the results of this study together with those of Refs.²⁰ and ²⁶ show
113 good overall consistency in the equation of state determined from dynamic experiments, static
114 experiments, and theory at ultrahigh pressures. A summary of experimentally constrained stress,
115 d -spacings, densities and phase assignments for ramp-compressed SiC is shown in Table S2.

116

117 **Discussion**

118 Planetary accretion simulations suggest carbon-rich planets may be dominated by carbon-
119 rich phases in the mantle and Fe-Ni-Si compounds in the core¹. To evaluate the possible interiors
120 of carbon-rich exoplanets, we constructed a simplified internal structure model for an eight-
121 Earth-mass planet in which the mantle consists of SiC or a mixture of C (diamond) and SiC (1:1
122 molar fraction) and the core consists of Fe or an Fe-Si alloy with 15 weight % Si (hereafter Fe-
123 15Si). By analogy with the Earth, the mantle is assumed to be 70% of the planet by mass and the
124 core 30% by mass. The amount and nature of light elements in the core of a carbon planet is
125 uncertain. Here we use ramp-compression data²³ for Fe-Si up to 1300 GPa as a proxy for
126 possible light elements and to minimize the need for extrapolation. The proportion of light
127 element in the core will depend on the partitioning behavior of iron and light elements during
128 planetary formation processes^{34,35}.

129 The interior structure was calculated using the Burnman package³⁶ to solve coupled
130 equations for hydrostatic equilibrium, mass conservation, and the equation of state of each layer
131 given the fractional mass of the core and mantle. We neglect temperature effects which have
132 been shown to have only a small effect on mass-radius relationships^{37,38}. In contrast to previous
133 studies, we included the B3-B1 phase transition at 70 GPa in our model which results in an
134 increase in the planet's radius for a given mass. The calculation used the present results from B1
135 SiC together with previous experimental results for Fe-Si²³, Fe³⁹, and B3 SiC¹⁸. Carbon-rich
136 planets are predicted to have a graphite-rich crust¹⁴, which is not included our model as the effect
137 of the crust on the mass-radius relations is negligible. The surface environment and atmosphere
138 composition of carbon-rich planets may be very different from silicate planets^{14,40}.

139

140 The resulting $8-M_E$ planets have radii of 1.78-1.84 R_E , shown in Fig. 6, where R_E and M_E
141 are the radius and mass of Earth, respectively. The pressure at the core mantle boundary is 1.0-
142 1.2 TPa and the central pressure ranges from 2.7 to 3.4 TPa (Fig. 6). The presence of a light
143 element in the core increases the core size but substantially decreases the central pressure. This is
144 consistent with the previous models for the silicate-dominated planets^{23,38,41,42}. In the mantle, the
145 phase transition from four-coordinated SiC (B3) to six-fold coordination (B1) occurs at a shallow
146 depth ($R/R_P \approx 0.93$, where R_P is a planetary radius) and the B1-phase SiC would comprise more
147 than 77 vol.% of the $8-M_E$ planet's mantle. In contrast, the mantle of an Earth-sized carbon
148 planet would be mainly composed of the low-pressure B3 phase (~90 vol.%, $R/R_P \approx 0.63$)⁴³. The
149 B3-B1 transformation is accompanied by a large density increase of 15-20%^{15,18-20,44}. Such a
150 significant density increase combined with a negative Clapeyron slope for the transition¹⁸⁻²⁰ may
151 inhibit convection across the boundary, favoring a layered thermal convection pattern suggested
152 by previous studies⁴³. The large transition energy²⁷ may also favor layered convection. The
153 reduced mass and heat transfer between the lower and upper mantle may affect the
154 thermochemical evolution of the planet.

155 Incorporation of diamond into the mantle increases in the radius of planets with the same
156 mass up to 2% (Figures 6 and 7). Furthermore, diamond is predicted to transform to the BC8
157 phase above 1 TPa accompanied by ~3% density increase and metallization^{26,45}, although the
158 transformation has not yet been directly experimentally confirmed²¹. In our model the predicted
159 transition occurs for planets greater than about eight Earth masses. The thickness of the BC8-
160 containing layer above the core-mantle boundary would be ~900-1200 km for the largest planet
161 size we considered ($10 M_E$). The presence of a dense, metallic layer above the core may affect
162 heat flow, thermal evolution, and magnetic field generation .

163 SiC is known to be a refractory material with a high melting temperature⁴⁶. Melting is an
164 important process in planetary interiors and strongly influences differentiation and the early
165 history of a planet. Mantle viscosity is also an important factor for the internal structure and
166 dynamics of planets. The high melting temperature and possible high strength of SiC suggest that
167 the interior may be characterized by high viscosity and hence sluggish convection⁴⁷ even for the
168 more highly coordinated B1 phase. It should be noted that there are only limited measurements
169 of the strength of SiC under at extreme conditions¹⁶. Additionally, SiC also has a high thermal
170 conductivity and low thermal expansivity, which may further suppress mantle convection^{18,43,48}.
171 This contrasts with recent studies^{24,49} that have suggested that viscosity decreases at high
172 pressure for large silicate planets due to phase transitions of MgO and the dissociation of
173 MgSiO₃. Recent numerical simulations have reported that the onset of convection in a carbon
174 planet may be highly dependent on the activation volume and the initial temperature¹⁸. More data
175 is needed on the properties of the B1 phase of SiC and SiC + diamond mixtures at extreme
176 conditions to better understand the dynamic behavior of such planets

177 Observations of the mass and radius of exoplanets provide a basis for modeling the
178 possible range of their mineralogies and interior structures. Accurate equations of state to
179 ultrahigh pressures are required to constrain the internal composition of exoplanets. We
180 constructed an experimentally constrained mass-radius curves for a hypothetical pure SiC planet
181 compared with other hypothetical single-phase planets (e.g. H₂O, C, MgSiO₃, Fe-15Si, Fe) for
182 evaluating the possible compositional space of exoplanets. Our results show that a pure SiC
183 planet is expected to be ~10% less than a MgSiO₃ planet. The mass-radius curves for a pure SiC
184 are also compared with those of (yellow region) using the previously reported B1 EOS^{18,20,26-33}
185 shown in Figure 7. These show a wide range of mass-radius curves due to usually long

186 extrapolation of low-pressure data. Our study provides experimental data on the mass-radius
187 relationship for a pure SiC planet to reduce uncertainties from extrapolation of the equation of
188 state. Mass-radius curves are also calculated for a planet composed of a SiC mantle or a C and
189 SiC (1:1 molar fraction) mantle (70% by mass) and an Fe core containing up to 15 weight % Si
190 core (30% by mass). Figure 7 shows the mass-radius relationship for such planets extending up
191 to 10 Earth masses in size. The red and gray regions in the figure represent the range of mass-
192 radius curves depending on the core composition. Our result shows that SiC is less dense than
193 MgSiO_3^{50} and hence a planet with a SiC (red region) or C and SiC (gray region) mantle would be
194 less dense than a corresponding Earth-like planet (30 % Fe + 70 % MgSiO_3) with the same
195 radius⁵⁰. For a given total mass, the radius of a planet decreases with core size (or mass) due to
196 the high density of the metallic core⁴². The mass-radius relation of carbon planets of 1-10 Earth
197 masses in size would overlap those of Earth-like planets (30% core mass) if the core mass of the
198 carbon planet is in the range of 37-50%.

199 Our results provide the first experimentally constrained mass-radius curves for SiC-rich
200 bodies that do not rely on theoretical calculations or long extrapolation of low-pressure data. It
201 should be noted that it is not possible to uniquely identify a planet's interior compositions based
202 on the mass-radius relationship alone. Additional constraints such as detection of atmospheric
203 contents and/or host star composition may help reduce compositional degeneracy⁴. Spectroscopic
204 measurements of exoplanetary atmospheres in future space missions may enable improved
205 characterization of super-Earth planets and resolve the question of the existence and composition
206 of carbon-rich planets.

207 Our results provide experimental evidence for the stability of SiC in the B1 phase over a
208 wide range of exoplanetary conditions that encompass expected mantle conditions of carbon-rich

209 exoplanets. The effect of other potential planetary constituents, if present in sufficient quantity,
210 on phase relations and other properties has been explored only under limited conditions so
211 far^{51,52}. Further work is needed to better understand the mineralogy of carbon-rich rocky planets
212 as a function of pressure, temperature, composition, and oxidation state.

213 In sum, the atomic-level structure of SiC has been determined under dynamic ramp
214 loading to 1.5 TPa using *in situ* X-ray diffraction. Our data confirm that SiC transforms to the B1
215 structure between ~80 and 140 GPa and remains in this phase to at least 1.5 TPa without further
216 phase transformation. The experimental data for the B1 phase are in good agreement with a
217 recent theoretical calculation²⁶. Our work extends the experimental constraints on this material to
218 extreme pressures, more than seven times greater than those achieved with standard static-
219 compression techniques, and more than four times greater than the central pressure of the Earth.
220 SiC was compressed to densities more than three times larger than its ambient value. Our results
221 enable construction of the experimentally constrained mass-radius curves for SiC-rich planets
222 without long extrapolation of low-pressure data. Carbide planets are found to have a lower
223 density than Earth-like planets. Interior structure models of carbide planets have also been
224 constructed, placing constraints on the internal pressure and density distribution, and yielding
225 insights into the possible dynamics and thermal behavior of such bodies.

226

227 **Experimental methods**

228 **Sample preparation and target assembly**

229 Silicon carbide powder (~1-μm mean grain size, 99.8% purity from Alfa Aesar) was used
230 as a starting material. The phase and purity of the sample (B3-type SiC, $F\bar{4}3m$) were confirmed
231 by Raman spectroscopy. Sample disks of ~12–21-μm thickness were produced by compressing

232 the SiC powder in a diamond anvil cell to 5 GPa, thereby reducing the porosity and increasing
233 the uniformity of the sample. The chosen sample thickness range was a balance between
234 maximizing diffraction intensity while avoiding the formation of a shock wave in the sample. A
235 representative compressed sample was examined by scanning electron microscopy. Identifying
236 void space in the resulting image using the optical imaging method of ref.⁵³ leads to an estimated
237 porosity of ~2% (Figure S4).

238 The sample package was mounted on the front of a detector box lined with image plates
239 (Fig. 1A). The target assembly consisted of a SiC layer sandwiched between a [110] diamond
240 ablator and either a [100] LiF window for low-stress experiments (<400 GPa) or a [110]
241 diamond window for higher stresses (>400 GPa) (Fig. 1B). A thin Au layer was added in the
242 latter case to prevent preheating and melting of the sample by X-rays emitted by the drive
243 plasma. Epoxy layers were ~1-μm in thickness. A Ta, W or Pt pinhole (75 to 150-μm thick, 300-
244 μm-diameter hole) was centered over the back of the target package.

245 **X-ray diffraction measurements**

246 Samples were compressed using a laser drive consisting of one or more 351-nm ramp-
247 shaped laser pulses. Either a single 10-ns long pulse (Omega EP) or a composite drive
248 comprising three to six 1-3-ns long laser pulses (Omega-60) was used. The drive laser directly
249 ablates the diamond at the front of the sample package to generate a rapidly expanding plasma
250 that drives a compression wave through the SiC sample. This creates a series of reverberations
251 which generate a uniform stress state within the SiC layer due to impedance mismatch across the
252 boundary between SiC and LiF or diamond window (Fig. 1). The peak drive laser intensities
253 ranged from 7×10^{11} to 6.0×10^{13} W/cm².

254 A quasi-monochromatic He_α X-ray source was produced by laser-irradiating Cu ($\text{He}_\alpha =$
255 8.368 keV) or Ge ($\text{He}_\alpha = 10.249$ keV) foils (Fig. 1A). The foils were positioned 24 mm from the
256 target at an angle of 22.5° (Omega-EP) or 45° (Omega-60) from target normal. At Omega-EP a
257 single 1-2-ns duration square laser pulse was applied with an energy of 1250-1950 J/beam. At
258 Omega-60, 16-18 beams with energies of 400-500 J/beam were applied over a 1-1.6 ns square
259 pulse. The X-ray drive lasers were timed to produce X-rays at the predicted time of peak stress
260 within the SiC layer. The X-rays were transmitted through the sample and collimated by the
261 pinhole. X-ray diffraction data were recorded on five image plates attached to the detector box²².
262 The image plates were filtered by Cu (12.5-μm thick) or Al (50- to 75-μm thick) to reduce
263 satellite emissions such as H_α , He_β and He_γ X-rays, and the bremsstrahlung X-ray background
264 from the drive plasma⁵⁴. Kapton filters (25-μm thick) were placed in front of the Cu or Al filters
265 to block optical light. Diffraction from the pinhole substrate provided ambient reference lines to
266 calibrate the diffraction geometry (see Table S1). The image plates were projected into 2θ - ϕ
267 space (2θ is the scattering angle and ϕ is the azimuthal angle around the incident X-ray
268 direction). In these coordinates, diffraction data are projected as straight lines of constant 2θ
269 (Fig. 2). Interplanar d -spacings were determined from the diffraction angle using Bragg's Law: λ
270 = $2d\sin(\theta)$, where λ is the X-ray wavelength.

271 For each image plate, the observed diffraction features can be assigned to one of the
272 following components of the target assembly: (1) the pinhole substrate material, (2) compressed
273 SiC, (3) Bragg or Laue diffraction from compressed LiF or diamond. We observed between one
274 and four diffraction lines from SiC in these experiments. Diffraction from LiF or diamond could
275 be distinguished based on d spacings as well as characteristic textural features. A representative

276 Debye-Scherrer X-ray diffraction pattern for a typical experiment is shown in Fig. 2. X-ray
277 diffraction patterns and filtered image plates for all shots are shown in Fig. S1.

278 Systematic corrections were applied to correct for the offset of the pinhole substrate from
279 the pinhole center and the offset of the sample and the pinhole, as described in ref.⁵⁵ The
280 uncertainties in the interplanar *d*-spacings include variation in the measured value of two-theta as
281 a function of azimuthal angle (ϕ), uncertainty in fitting the sample peak positions to Gaussian
282 profiles, and the uncertainty in the incident X-ray wavelength (<0.01 Å)⁵⁴. The resultant
283 uncertainties in *d*-spacing (~1%) are given in Supple. Table S2.

284

285 **Stress determination**

286 A line-imaging velocity interferometry system for any reflector (VISAR)⁵⁶ was used to
287 determine the velocity at either the free surface of the rear diamond or at the interface between
288 SiC and the LiF window. The VISAR records the change in the Doppler shift of light reflected
289 off a moving surface (Figs. 3 and 4). Two independent VISAR channels with different velocity
290 sensitivities were used to resolve velocity ambiguities that exceed the frequency response of the
291 system. Example velocity profiles are shown in Figs. 3B and 4B. As the strength of SiC on ramp
292 loading is poorly constrained, we have not made any correction for the difference between axial
293 stress (P_x) and mean pressure (P). The stress within the sample is calculated based on the full-
294 width at half-maximum of the stress histogram over the duration of the 1-2-ns X-ray
295 measurement.

296 For low-stress shots (<400 GPa), the SiC-LiF interface particle velocity, $u_p(t)$, was
297 recorded as a function of time (Fig. 3B, blue and red traces). A correction to the interface
298 velocity was made using the density-dependence of the refractive index of LiF⁵⁷. The stress

299 history within the sample was determined using the radiation hydrodynamics code HYADES⁵⁸.
300 The simulations were used to constrain the stress history within the sample using EOS
301 descriptions of each target material and to provide an initial estimate of the delay time between
302 the laser used for ramp compression and those used to generate the He _{α} emission (Fig. 3). An
303 initial estimate of the stress is determined from a scaling law⁵⁹ that relates laser intensity to
304 diamond ablation stress: $\sigma(\text{GPa}) = 42(\pm 3)[I(\text{TW/cm}^2)]^{0.71(\pm 0.01)}$. The actual stress history within
305 the sample following ramp compression was determined by matching the observed VISAR
306 profile to the hydrodynamic simulations. The SESAME EOS tables⁶⁰ for LiF (#7271) and
307 diamond (#7830) and the Livermore Equation of State (LEOS) table⁶¹ for SiC (#2130) were used
308 in hydrocode simulations with a Steinberg-Guinan strength model⁶² (shear modulus of 243 GPa
309 and yield stress of 20 GPa⁶³).

310 For the high-stress shots (>400 GPa), the diamond window free-surface velocity, $u_{\text{fs}}(t)$,
311 was recorded as a function of time (Fig. 4B, blue and red traces) and used with a backward
312 characteristics analysis to convert to stress using diamond ramp EOS^{64–66}. The stress in SiC
313 during the X-ray probe time was determined by spatially averaging over the sample layer (Fig.
314 4C, D). The stress uncertainty includes the following sources⁵⁵: [1] uncertainty in the equation of
315 state of the window materials (3% of the total stress^{64,67}, systematic); [2] uncertainty in
316 determining the phase of the VISAR fringes (1% of the total stress, random); [3] uncertainty in
317 the stress gradient over the X-ray probe timing. This is estimated as the standard deviation of the
318 stress histogram fitted with a Gaussian function (random); [4] thickness uncertainty for each
319 component of the target assembly (0.05% of the total stress, random); [5] uncertainty in the
320 strength of diamond. Assuming the stress follows along the reversible isentropic path,
321 maintaining diamond strength, leads to a systematic underestimate of the stress (systematic, +50

322 GPa²³). Total stress uncertainties are calculated using sources [1]-[4] in quadrature. The
323 additional +50 GPa uncertainty applies only to the shots with diamond windows [5]. The total
324 stress uncertainties are shown in table S2. A summary of laser power, stress and sample velocity
325 as a function of time for all shots is shown in Fig. S5.

326 **Calculation of principal isentrope and isotherm**

327 In these experiments, the longitudinal stress, P_x , was measured. The longitudinal stress-
328 density, P_x - ρ , relationship was transformed to conditions of the principal isentrope by correcting
329 for the following factors: 1) the deviation between the longitudinal stress and hydrostatic
330 pressure, 2) the thermal pressure of the initial shock (~52 GPa) from a porous starting material
331 (B3 phase), 3) plastic work heating resulting from the strength of SiC and 4) the energy
332 associated with the B3 to B1 phase transition at 100 GPa^{15,17}.

333 1. The longitudinal stress can be described as a sum of a hydrostatic pressure, P_{hyd} and a
334 stress deviator term. Using the von Mises yield criterion, the longitudinal stress is defined as

335
$$P_x = P_{hyd} + \frac{2}{3}Y, \quad (1)$$

336 where Y is the yield strength.

337 2. The pressure-density relationship along the Hugoniot including porosity is given by:

338
$$P_H^* = P_H \frac{1 - \left(\frac{\gamma}{2}\right) \left(\frac{\rho}{\rho_0} - 1\right)}{1 - \left(\frac{\gamma}{2}\right) \left(\frac{\rho^*}{\rho_0} - 1\right)}, \quad (2)$$

339 where ρ is the density, and P_H^* and P_H are the stress states corresponding to the porous Hugoniot
340 and to the crystal Hugoniot, respectively. ρ_0 and ρ_0^* are the initial crystal and the initial bulk
341 density, respectively. γ is the Grüneisen parameter, assumed to be dependent only on density as
342 given by:

343
$$\gamma = \gamma_0 \left(\frac{\rho_0}{\rho} \right)^q. \quad (3)$$

344 where q is the logarithmic density dependence of γ . γ_0 is assumed to be in the range of 0.5-2.0
 345 based on previous studies^{17,18,63}. The laser drive in our experiments produces an initial elastic
 346 shock wave (~81 GPa) in the diamond ablator⁶⁴. This transmits a shock with lower amplitude
 347 (~52 GPa) into the SiC sample due to the impedance mismatch between diamond and SiC. The
 348 thermal pressure between the principal isentrope and the Hugoniot at the conditions of the initial
 349 shock (~ 52 GPa) was determined using the Mie-Grüneisen relation⁶⁸,

350
$$P_H^* - P_S = \gamma \rho (E_H^* - E_S), \quad (4)$$

351 where E_H^* is the energy along the porous Hugoniot, obtained from the Rankine-Hugoniot
 352 relationship:

353
$$E_H^* = \frac{P_H^*}{2} \left(\frac{1}{\rho_0^*} - \frac{1}{\rho} \right). \quad (5)$$

354 The pressure-density relationship of B3 SiC along the principal isentrope up to 52 GPa, P_S , was
 355 determined using the Birch-Murnaghan equation of state:

356
$$P_S = 3K_{0S}f(1 + 2f)^{\frac{5}{2}} \left(1 + \frac{3}{2} [K'_{0S} - 4]f \right), \quad (6)$$

357 where f is the Eulerian strain:

358

359
$$f = \frac{1}{2} \left[\left(\frac{\rho}{\rho_0} \right)^{\frac{2}{3}} - 1 \right], \quad (7)$$

360 and K_{0S} and K'_{0S} are the ambient isentropic bulk modulus and its pressure derivative (at constant
 361 entropy), respectively. The internal energy along the isentrope of the B3 phase up to 52 GPa, E_S ,
 362 can be obtained by integration of the Birch-Murnaghan equation:

363

$$E_S = \int_{1/\rho_0}^{1/\rho} \frac{P_S}{\rho^2} d\rho = \frac{9}{2} \frac{K_{0S}}{\rho_0} (f^2 + [K'_S - 4]f^3). \quad (8)$$

364 3. Material strength under compression results in plastic work heating, a source of
 365 thermal pressure, which causes the hydrostatic pressure to deviate from the isentrope by

366

$$P_{hyd} - P_S = \gamma\rho \int_0^{\epsilon_x} f_{TQ} dW_P, \quad (9)$$

367 where ϵ_x is the natural strain $\ln \frac{\rho}{\rho_0}$, f_{TQ} is the Taylor-Quinney factor, which describes the fraction
 368 of plastic work converted to heat⁶⁹, W_p is the plastic working heating. The Taylor-Quinney factor
 369 represents the relationship between plastic work and heat generation, which is generally assumed
 370 to be $0.7 < f_{TQ} < 0.95$ ⁷⁰. In this work, f_{TQ} is assumed to be 1 indicative of 100% of plastic work
 371 converted to heat. Plastic working heating⁷¹ is defined as

372

$$W_P = \frac{1}{\rho_0} \frac{2}{3} Y \left[d\epsilon_x - \frac{dY}{2G(\rho)} \right], \quad (10)$$

373 where $G(\rho)$ is the shear modulus. In this work, we assume Y is fixed at a constant value ranging
 374 from 5- 20 GPa.

375 4. E_T is the energy associated with the B3 to B1 phase transition²⁷. This transition has a
 376 negative Clapeyron slope^{18,20}, indicating heat is absorbed (endothermic) which decreases the
 377 temperature across the transition.

378 The principal isentrope is calculated from the measured longitudinal stress through 1-4
 379 given by

380

$$P_S = P_x - \frac{2}{3} Y - \gamma\rho(E_H^* - E_S) - \gamma\rho \int_0^{\epsilon_x} \beta dW_P + \gamma\rho E_T. \quad (11)$$

381 The 300-K isotherm is then determined from the principal isentrope using the Mie-
382 Grüneisen relation. The thermodynamic parameters used in the calculation are listed in the
383 Supplemental Material, Table S4.

384 **Stress, energy, and temperature path**

385 While ramp compression reduces the extent of sample heating, irreversible work and
386 heating as a result of dissipation and material strength occur during ramp loading. The quasi-
387 isentropic path followed in a ramp compression experiment lies between the Hugoniot and the
388 isentrope (Figures S6 and S7). In these experiments, the measurements of P_x - ρ provide the
389 energy along the ramp, E_r , using equation (11) and the Mie-Grüneisen relation,

$$390 E_r = E_S + \frac{2Y}{3\gamma\rho} + (E_H^* - E_{H,S}) + \int_0^{\epsilon_x} \beta \frac{1}{\rho_0} \frac{2}{3} Y d\epsilon_x - (E_{Tr}). \quad (12)$$

391 Below the B3-B1 transition pressure transition (<100 GPa), E_r is divided into two
392 regions: 1) the energy along the B3 porous Hugoniot up to 52 GPa, and 2) the energy along the
393 isentrope of the B3 phase from the 52-GPa Hugoniot state to 100 GPa (Figure S6). Above the
394 B3-B1 phase transition, the estimated values of E_r as a function of stress are calculated using
395 Eqn. 12 and are shown as the light-blue band in Fig. S6.

396 Temperature cannot be directly measured in our experiments, but can be estimated from
397 the energy difference between the ramp path and the principal isentrope, calculated using the
398 Debye model given by:

$$399 E_r - E_S = 9nRT \left[\frac{\theta_D(V)}{T} \right]^{-3} \int_0^{\theta_D(V)/T} \frac{x^3}{e^x - 1} dx, \quad (13)$$

400 where $\theta_D(V)$ is the Debye temperature, n is the number of atoms per formula unit, R is the gas
401 constant. The estimated temperature are shown as a blue curve in Figure S7 assuming Y and γ_0
402 are 20 GPa and 1, respectively, corresponding to the E_r shown as a black curve in Figure S6. The

403 estimated temperature ranges from 1180 K to 3440 K corresponding to a stress range from 100
404 GPa to 1600 GPa, which is much lower than the expected melting curve at these conditions
405 shown in Fig. S7.

406

407 **Acknowledgements:** We thank Carol Davis (LLNL) and the staff of the Laboratory for Laser
408 Energetics for expert experimental assistance. The Lawrence Livermore National Laboratory
409 AnalyzeVISAR and AnalyzePXRDIP code were used for data analysis. The authors
410 acknowledge the use of Princeton's Imaging and Analysis Center and Bolton Howes for
411 assistance with porosity estimation.

412 **Funding:** This work was performed under the auspices of the US Department of Energy by
413 Lawrence Livermore National Laboratory under contract number DE-AC52-07NA27344. The
414 research and materials incorporated in this work were partially developed at the National Laser
415 Users' Facility at the University of Rochester's Laboratory for Laser Energetics, with financial
416 support from the U.S Department of Energy under Cooperative Agreement DE-NA0001944 and
417 DE-NA0003611. Additional support was provided by the National Science Foundation EAR-
418 1644614 and the Princeton Center for Complex Materials (PCCM), and a National Science
419 Foundation (NSF)-MRSZECZ program (DRM-2011750).

420 **Autor contributions:**

421 Author contributions: D.K., R.F.S., J.H.E., S.J.T. and T.S.D conceived and designed the
422 experiments. D.K. analyzed the data with support from R.F.S. and F.C. D.K., R.F.S., I.K.O.,
423 M.C.M., M.K.G., J.K.W., J.R.R., and T.S.D. carried out experiments. J.R.R., A.L., F.C., and
424 R.F.S. provided experimental development. D.K. and T.S.D. wrote the manuscript with support

425 from all authors on the manuscript. T.S.D. supervised the project. All authors were involved in
426 discussions related to experimental design and data analysis.

427 **Competing interests:** The authors declare that they have no competing interests.

428 **Data and materials availability:** All data are available in the main text or the supplementary
429 materials. Additional data related to this paper may be requested from the authors.

430

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432

433 **References**

434

- 435 1. Bond, J. C., O'Brien, D. P. & Lauretta, D. S. The compositional diversity of extrasolar terrestrial planets. I. In
436 situ simulations. *Astrophys. J.* **715**, 1050–1070 (2010).
- 437 2. Scora, J., Valencia, D., Morbidelli, A. & Jacobson, S. Chemical diversity of super-Earths as a consequence of
438 formation. *Mon. Not. R. Astron. Soc.* **493**, 4910–4924 (2020).
- 439 3. Jontof-Hutter, D. The compositional diversity of low-mass exoplanets. *Annu. Rev. Earth Planet. Sci.* **47**, 141–
440 171 (2019).
- 441 4. Madhusudhan, N., Lee, K. K. M. & Mousis, O. A possible carbon-rich interior in super-Earth 55 Cancri e.
442 *Astrophys. J.* **759**, L40 (2012).
- 443 5. Kuchner, M. J. & Seager, S. Extrasolar carbon planets. *arXiv:astro-ph/0504214* (2005).
- 444 6. Larimer, J. W. The effect of C/O ratio on the condensation of planetary material. *Geochim. Cosmochim. Acta*
445 **39**, 389–392 (1975).
- 446 7. Stonkutė, E. *et al.* High-resolution spectroscopic study of dwarf Stars in the northern sky: lithium, carbon, and
447 oxygen abundances. *Astrophys. J.* **159**, 90 (2020).
- 448 8. Clark, J. T. *et al.* The GALAH survey: using galactic archaeology to refine our knowledge of *TESS* target stars.
449 *Mon. Not. R. Astron. Soc.* (2021).
- 450 9. Brewer, J. M. & Fischer, D. A. C/O and Mg/Si ratios of stars in the solar neighborhood. *Astrophys. J.* **831**, 20
451 (2016).

452 10. Nissen, P. E. *et al.* High-precision abundances of elements in solar-type stars. Evidence of two distinct
453 sequences in abundance-age relations. *Astron. Astrophys.* **640**, A81–A95 (2020).

454 11. Mashian, N. & Loeb, A. CEMP stars: possible hosts to carbon planets in the early universe. *Mon. Not. R.*
455 *Astron. Soc.* **460**, 2482–2491 (2016).

456 12. Asplund, M., Grevesse, N., Sauval, A. J. & Scott, P. The chemical composition of the Sun. *Annu. Rev. Astron.*
457 *Astrophys.* **47**, 481–522 (2009).

458 13. Moriarty, J., Madhusudhan, N. & Fischer, D. Chemistry in an evolving protoplanetary disk: effects on
459 terrestrial planet composition. *Astrophys. J.* **787**, 81 (2014).

460 14. Sleep, N. H. Planetary interior-atmosphere interaction and habitability. in *Handbook of Exoplanets* (eds. Deeg,
461 H. J. & Belmonte, J. A.) 2937–2958 (Springer, 2018).

462 15. Sekine, T. & Kobayashi, T. Shock compression of 6H polytype SiC to 160 GPa. *Phys. Rev. B* **55**, 8034–8037
463 (1997).

464 16. Vogler, T. J., Reinhart, W. D., Chhabildas, L. C. & Dandekar, D. P. Hugoniot and strength behavior of silicon
465 carbide. *J. Appl. Phys.* **99**, 023512 (2006).

466 17. Tracy, S. J. *et al.* *In situ* observation of a phase transition in silicon carbide under shock compression using
467 pulsed x-ray diffraction. *Phys. Rev. B* **99**, 214106 (2019).

468 18. Miozzi, F. *et al.* Equation of state of SiC at extreme conditions: new insight into the interior of carbon-rich
469 exoplanets. *J. Geophys. Res. Planets* **123**, 2295–2309 (2018).

470 19. Daviau, K. & Lee, K. K. M. Zinc-blende to rocksalt transition in SiC in a laser-heated diamond-anvil cell. *Phys.*
471 *Rev. B* **95**, 134108 (2017).

472 20. Kidokoro, Y., Umemoto, K., Hirose, K. & Ohishi, Y. Phase transition in SiC from zinc-blende to rock-salt
473 structure and implications for carbon-rich extrasolar planets. *Am. Mineral.* **102**, 2230–2234 (2017).

474 21. Lazicki, A. *et al.* Metastability of diamond ramp-compressed to 2 terapascals. *Nature* **589**, 532–535 (2021).

475 22. Rygg, J. R. *et al.* Powder diffraction from solids in the terapascal regime. *Rev. Sci. Instrum.* **83**, 113904 (2012).

476 23. Wicks, J. K. *et al.* Crystal structure and equation of state of Fe-Si alloys at super-Earth core conditions. *Sci.*
477 *Adv.* **4**, eaao5864 (2018).

478 24. Coppari, F. *et al.* Implications of the iron oxide phase transition on the interiors of rocky exoplanets. *Nat.*
479 *Geosci.* **14**, 121–126 (2021).

480 25. Wang, J. *et al.* X-ray diffraction of molybdenum under ramp compression to 1 TPa. *Phys. Rev. B* **94**, (2016).

481 26. Wilson, H. F. & Militzer, B. Interior phase transformations and mass-radius relationships of silicon-carbon

482 planets. *Astrophys. J.* **793**, 34 (2014).

483 27. Karch, K., Bechstedt, F., Pavone, P. & Strauch, D. Pressure-dependent properties of SiC polytypes. *Phys. Rev.*

484 *B* **53**, 13400–13413 (1996).

485 28. Durandurdu, M. Pressure-induced phase transition of SiC. *J. Condens. Matter Phys.* **16**, 4411–4417 (2004).

486 29. Durandurdu, M. *Ab initio* simulations of the structural phase transformation of 2H-SiC at high pressure. *Phys.*

487 *Rev. B* **75**, 235204 (2007).

488 30. Lu, Y.-P., He, D.-W., Zhu, J. & Yang, X.-D. First-principles study of pressure-induced phase transition in

489 silicon carbide. *Physica B: Condens. Matter* **403**, 3543–3546 (2008).

490 31. Lee, W. H. & Yao, X. H. First principle investigation of phase transition and thermodynamic properties of SiC.

491 *Comput. Mater. Sci.* **106**, 76–82 (2015).

492 32. Gorai, S. & Bhattacharya, C. Shock induced phase transition in SiC polytypes. *J. Appl. Phys.* **125**, 185903

493 (2019).

494 33. Gorai, S., Bhattacharya, C. & Kondayya, G. Pressure induced structural phase transition in SiC. *AIP Conf.*

495 *Proc.* **1832**, 030010 (2017).

496 34. Schaefer, L., B. Jacobsen, S., Remo, J. L., Petaev, M. I. & Sasselov, D. D. Metal-silicate partitioning and its

497 role in core formation and composition on super-Earths. *Astrophys. J.* **835**, 234 (2017).

498 35. Fischer, R. A., Cottrell, E., Hauri, E., Lee, K. K. M. & Le Voyer, M. The carbon content of Earth and its core.

499 *Proc. Natl. Acad. Sci. U.S.A.* **117**, 8743–8749 (2020).

500 36. Cottaar, S., Heister, T., Rose, I. & Unterborn, C. BurnMan: a lower mantle mineral physics toolkit. *Geochem.*

501 *Geophys. Geosyst.* **15**, 1164–1179 (2014).

502 37. Seager, S., Kuchner, M., Hier-Majumder, C. A. & Militzer, B. Mass-Radius Relationships for Solid Exoplanets.

503 *Astrophys. J.* **669**, 1279–1297 (2007).

504 38. Hakim, K. *et al.* A new ab initio equation of state of hcp-Fe and its implication on the interior structure and

505 mass-radius relations of rocky super-Earths. *Icarus* **313**, 61–78 (2018).

506 39. Dewaele, A. *et al.* Quasihydrostatic equation of state of iron above 2 Mbar. *Phys. Rev. Lett.* **97**, 215504 (2006).

507 40. Madhusudhan, N., Mousis, O., Johnson, T. V. & Lunine, J. I. Carbon-rich giant planets: atmospheric chemistry,
508 thermal inversion, spectra, and formation conditions. *Astrophys. J.* **743**, 191 (2011).

509 41. Unterborn, C. T., Dismukes, E. E. & Panero, W. R. Scaling the Earth: a sensitivity analysis of terrestrial
510 exoplanetary interior models. *Astrophys. J.* **819**, 32 (2016).

511 42. Dorn, C. *et al.* Can we constrain the interior structure of rocky exoplanets from mass and radius measurements?
512 *A&A* **577**, A83 (2015).

513 43. Nisr, C. *et al.* Thermal expansion of SiC at high pressure-temperature and implications for thermal convection
514 in the deep interiors of carbide exoplanets: thermal expansion of SiC. *J. Geophys. Res. Planets* **122**, 124–133
515 (2017).

516 44. Daviau, K. & Lee, K. High-pressure, high-temperature behavior of silicon carbide: a review. *Crystals* **8**, 217
517 (2018).

518 45. Martinez-Canales, M., Pickard, C. J. & Needs, R. J. Thermodynamically stable phases of carbon at
519 multiterapascal pressures. *Phys. Rev. Lett.* **108**, 045704 (2012).

520 46. Sokolov, P. S., Mukhanov, V. A., Chauveau, T. & Solozhenko, V. L. On melting of silicon carbide under
521 pressure. *J. Superhard Mater.* **34**, 339–341 (2012).

522 47. Deng, J. & Lee, K. K. M. Viscosity jump in the lower mantle inferred from melting curves of ferropericlase.
523 *Nat. Commun.* **29** (2017).

524 48. Li, J., Porter, L. & Yip, S. Atomistic modeling of finite-temperature properties of crystalline β -SiC II. Thermal
525 conductivity and effects of point defects. *J. Nucl. Mater.* **255**, 139–152 (1998).

526 49. Ritterbex, S., Harada, T. & Tsuchiya, T. Vacancies in MgO at ultrahigh pressure: about mantle rheology of
527 super-Earths. *Icarus* **305**, 350–357 (2018).

528 50. Zeng, L., Sasselov, D. D. & Jacobsen, S. B. Mass–radius relation for rocky planets based on PREM. *Astrophys.*
529 *J.* **819**, 127 (2016).

530 51. Allen-Sutter, H. *et al.* Oxidation of the interiors of carbide exoplanets. *Planet. Sci. J.* **1**, 39 (2020).

531 52. Hakim, K., van Westrenen, W. & Dominik, C. Capturing the oxidation of silicon carbide in rocky exoplanetary
532 interiors. *Astron. Astrophys.* **618**, L6 (2018).

533 53. Howes, B., Mehra, A. & Maloof, A. Three-dimensional morphometry of ooids in oolites: a new tool for more
534 accurate and precise paleoenvironmental interpretation. *J. Geophys. Res. Earth Surf.* **126**, (2021).

535 54. Coppari, F. *et al.* Optimized X-ray sources for X-ray diffraction measurements at the Omega laser facility. *Rev.*
536 *Sci. Instrum.* **90**, 125113 (2019).

537 55. Rygg, J. R. *et al.* X-ray diffraction at the National Ignition Facility. *Rev. Sci. Instrum.* **91**, 043902 (2020).

538 56. Celliers, P. M. *et al.* Line-imaging velocimeter for shock diagnostics at the OMEGA laser facility. *Rev. Sci.*
539 *Instrum.* **75**, 4916–4929 (2004).

540 57. Fratanduono, D. E. *et al.* Refractive index of lithium fluoride ramp compressed to 800 GPa. *J. Appl. Phys.* **109**,
541 123521 (2011).

542 58. Larsen, J. T. & Lane, S. M. HYADES-A plasma hydrodynamics code for dense plasma studies. *J. Quant.*
543 *Spectrosc. Radiat. Transf.* **51**, 179–186 (1994).

544 59. Fratanduono, D. E. *et al.* The direct measurement of ablation pressure driven by 351-nm laser radiation. *J. Appl.*
545 *Phys.* **110**, 073110 (2011).

546 60. Lyon, S. P. & Johnson, J. D. Sesame: the Los Alamos National Laboratory equation of state database. (1992).

547 61. More, R. M., Warren, K. H., Young, D. A. & Zimmerman, G. B. A new quotidian equation of state (QEoS) for
548 hot dense matter. *Phys. Fluids* **31**, 3059 (1988).

549 62. Steinberg, D. J. & Lund, C. M. A constitutive model for strain rates from 10^{-4} to 10^6 s $^{-1}$. *J. Appl. Phys.* **65**,
550 1528–1533 (1989).

551 63. Varshney, D., Shriya, S., Varshney, M., Singh, N. & Khenata, R. Elastic and thermodynamical properties of
552 cubic (3C) silicon carbide under high pressure and high temperature. *J. Theor. Appl. Phys.* **9**, 221–249 (2015).

553 64. Bradley, D. K. *et al.* Diamond at 800 GPa. *Phys. Rev. Lett.* **102**, 075503 (2009).

554 65. Rothman, S. D. & Maw, J. Characteristics analysis of isentropic compression experiments (ICE). *J. Phys. IV*
555 *France* **134**, 745–750 (2006).

556 66. Maw, J. R. A characteristics code for analysis of isentropic compression experiments. *AIP Conf. Proc.* **706**,
557 1217–1220 (2004).

558 67. Davis, J.-P., Knudson, M. D., Shulenburger, L. & Crockett, S. D. Mechanical and optical response of [100]
559 lithium fluoride to multi-megabar dynamic pressures. *J. Appl. Phys.* **120**, 165901 (2016).

560 68. Ahrens, T. J. *High-Pressure Shock Compression of Solids*. (Springer, 1993).

561 69. Vogler, T. J. Determining material strength in ramp loading experiments. *AIP Conf. Proc.* **1195**, 1023–1026
562 (2009).

563 70. Taylor, G. I. & Quinney, H. The latent energy remaining in a metal after cold working. *Proc. R. Soc. Lond. A*
564 **143**, 307–326 (1934).

565 71. Fowles, G. R. Shock wave compression of hardened and annealed 2024 aluminum. *J. Appl. Phys.* **32**, 1475–
566 1487 (1961).

567 72. Smith, R. F. *et al.* Equation of state of iron under core conditions of large rocky exoplanets. *Nat. Astron.* **2**,
568 452–458 (2018).

569 73. Smith, R. F. *et al.* Ramp compression of diamond to five terapascals. *Nature* **511**, 330–333 (2014).

570 74. Coppari, F. *et al.* Experimental evidence for a phase transition in magnesium oxide at exoplanet pressures. *Nat.*
571 *Geosci.* **6**, 926–929 (2013).

572 75. Wagner, F. W., Tosi, N., Sohl, F., Rauer, H. & Spohn, T. Rocky super-Earth interiors: structure and internal
573 dynamics of CoRoT-7b and Kepler-10b. *Astron. Astrophys.* **541**, A103 (2012).

574 76. Duffy, T. S. & Smith, R. F. Ultra-high pressure dynamic compression of geological materials. *Front. Earth. Sci.*
575 **7**, 23 (2019).

576

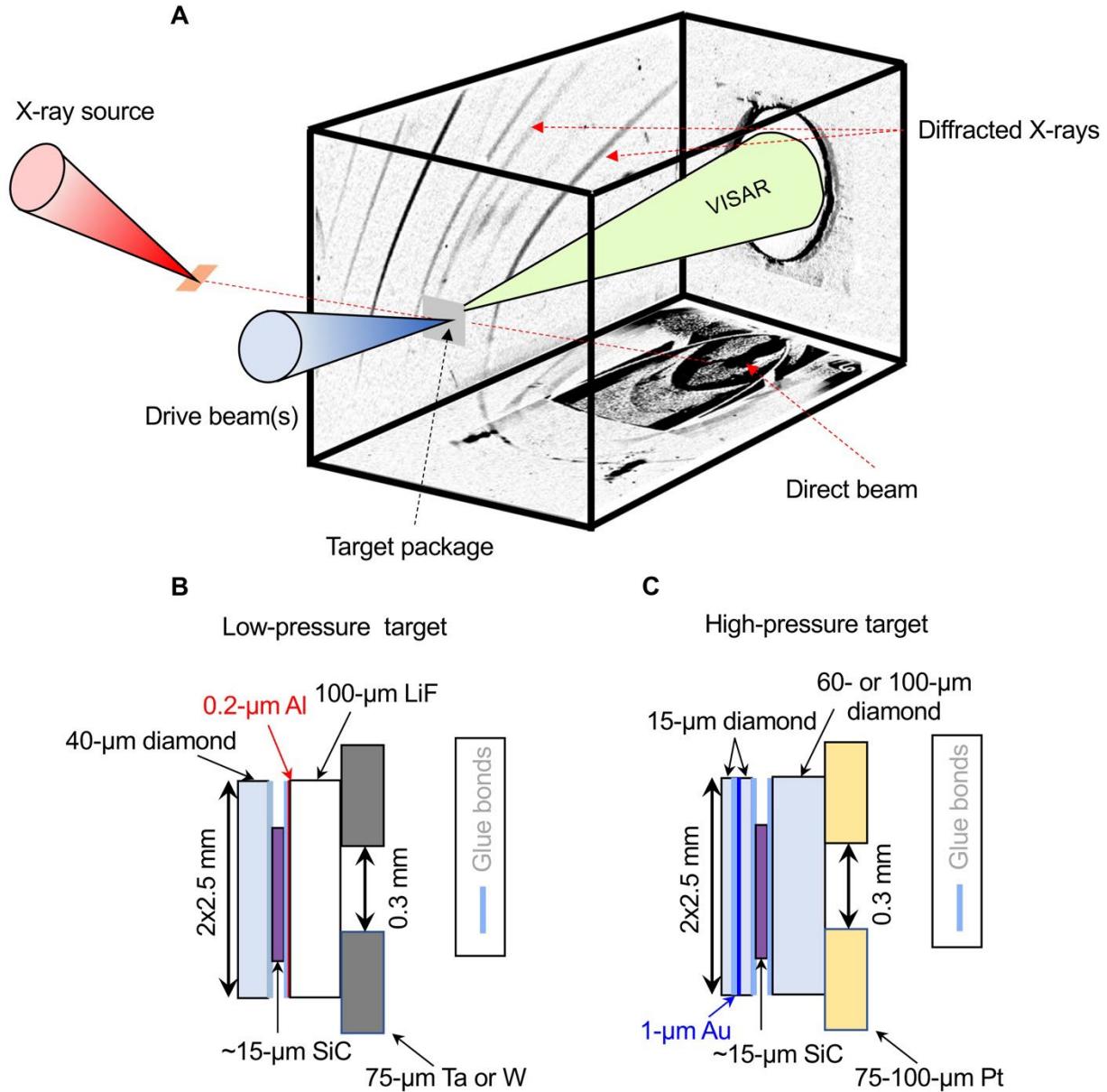
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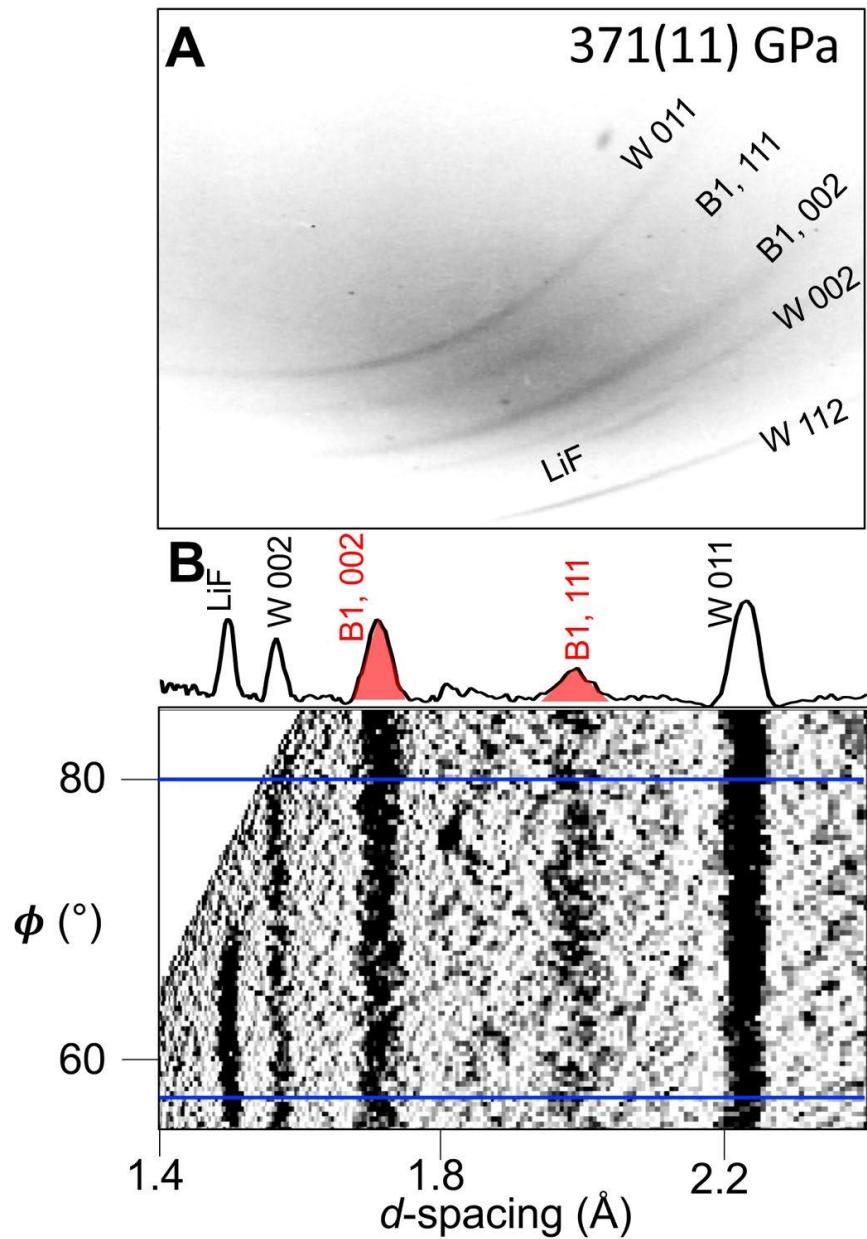


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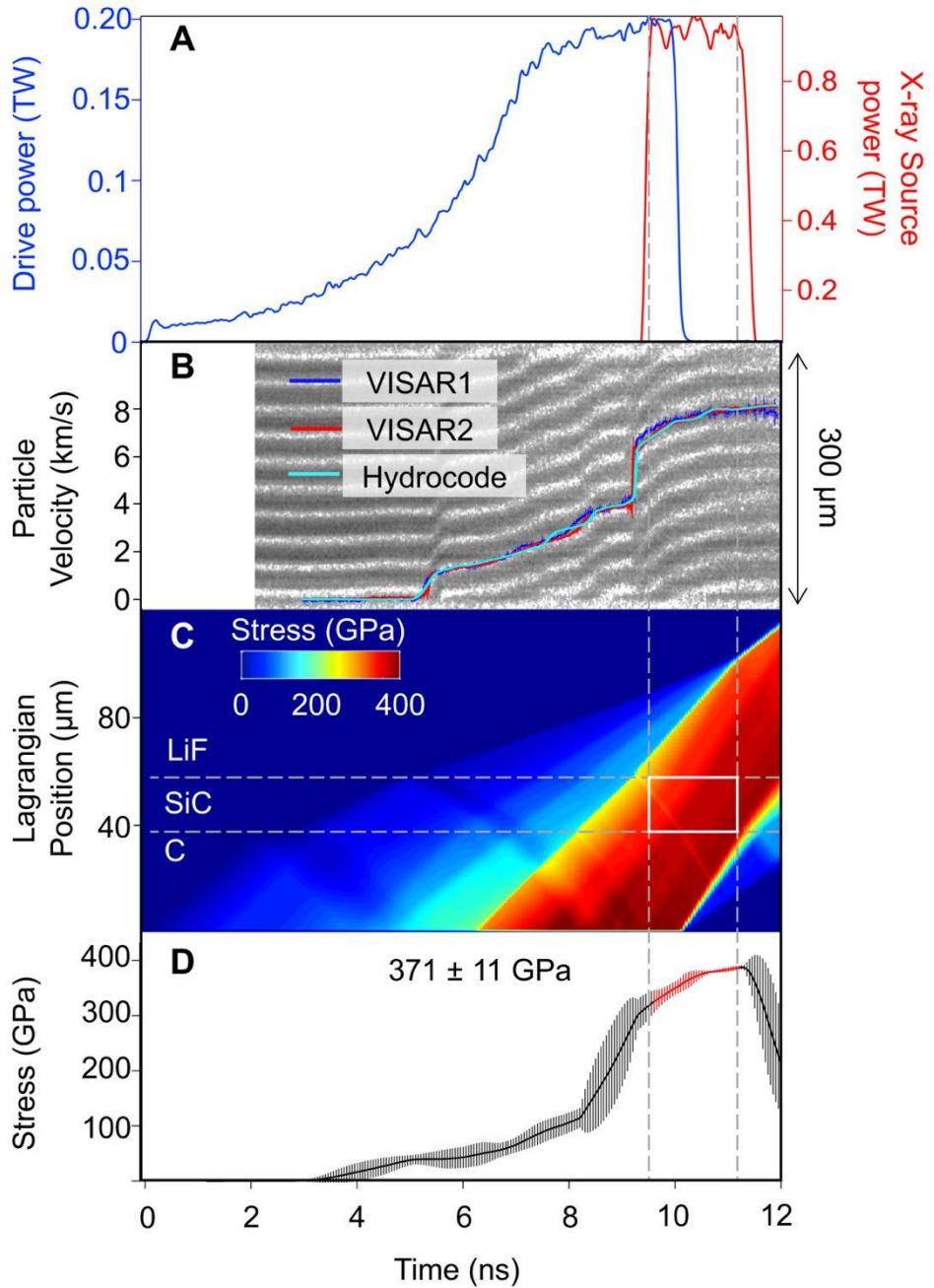
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584 **Figure 1. Experimental set-up and target packages for laser-driven compression.** (A) The
 585 drive laser (blue cone) generates a stress wave through the C/SiC/window target package (gray square). The stress rapidly equilibrates in the SiC foil following reverberation of the compression
 586 wave due to impedance mismatch with the diamond and window interfaces. A quasi-
 587 monochromatic X-ray source was generated by laser irradiation (red cone) of a Cu or Ge foil
 588 (orange square). The diffracted X-rays from the target are recorded on image plates inside of a
 589 detector box. Target assemblies used for low-stress (<400 GPa) and high-stress (>400 GPa)
 590 experiments are shown in (B) and (C), respectively. The front diamond surface is illuminated by
 591 the drive laser using distributed phase plates (800-μm focal spot). The incident X-rays are
 592 collimated using a 300-μm diameter pinhole (W, Ta or Pt).
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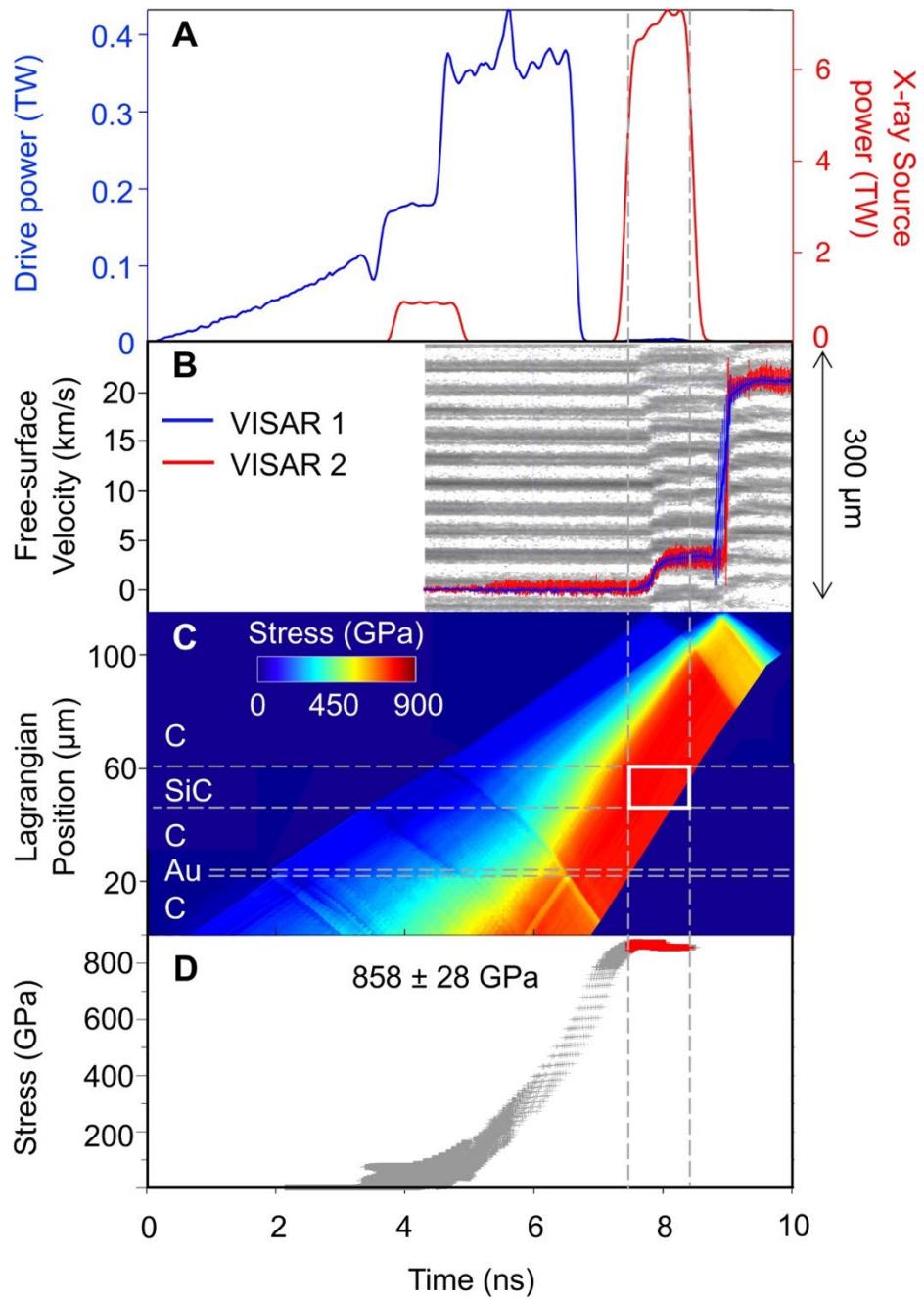
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 596 **Figure 2. X-ray diffraction pattern for a representative experiment on SiC (shot #27430).** (A)
 597 One panel of an unprocessed image plate. (B) Projection of the above image plate into d -spacing-
 598 ϕ space after background subtraction. The one-dimensional X-ray diffraction pattern shown on the
 599 upper axis represents integration in ϕ over the region between the blue lines. The textured
 600 diffraction from LiF can be distinguished from the more extended features from either the W
 601 pinhole or B1 SiC.
 602

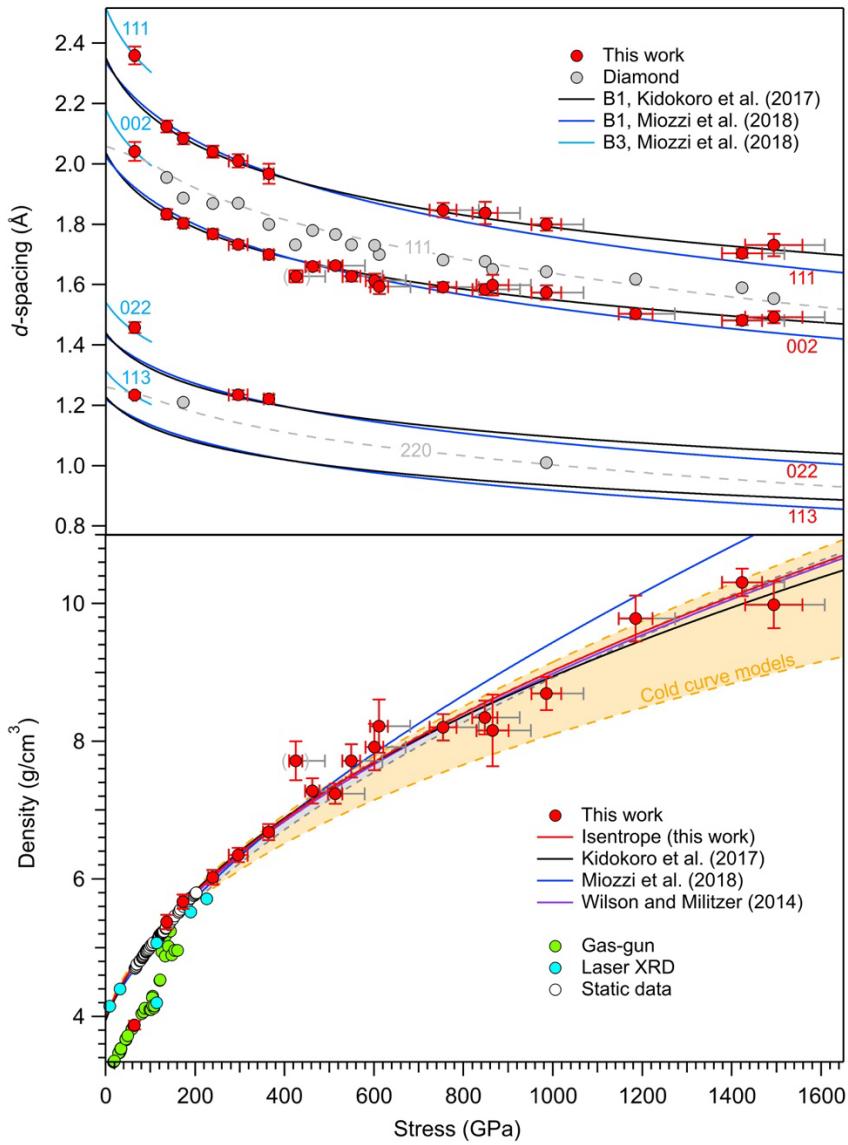


603
604 **Figure 3. Timing and stress determination for SiC ramp-compression experiment (LiF**
605 **window).** (A) Drive laser pulse for shot #27430 (blue trace). X-rays are generated using a 2-ns
606 square pulse (red trace). (B) Raw interferogram from the VISAR records the SiC-LiF interface
607 particle velocity (solid lines) which is reproduced by a hydrocode simulation (light blue curve)
608 to determine (C) stress history in the target package. The white rectangle represents the sample stress
609 condition during the X-ray probe period (bounded by vertical dashed lines). The horizontal dashed
610 lines represent the material layers in Lagrangian coordinates. (D) Calculated stress history of the
611 SiC sample as a function of time. The vertical dashed lines represent stress states over the X-ray
612 probe period.



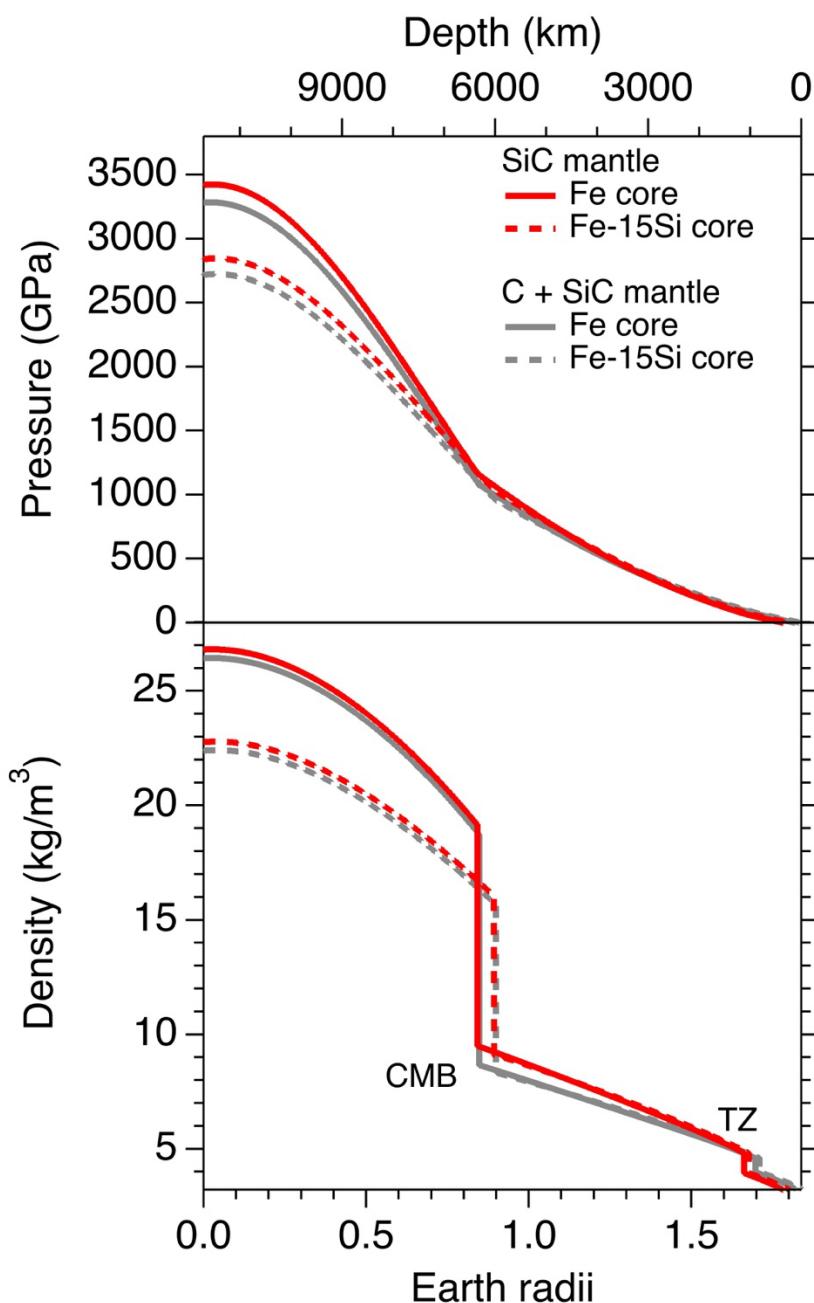
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Figure 4. Timing and stress determination for SiC ramp-compression experiment (diamond window). (A) Drive laser pulse for shot #98955 (blue trace). X-rays are generated using a pair of 1-ns square pulses (red trace). (B) Raw interferogram from VISAR records diamond free-surface velocity that is used to determine the stress history. (C) Calculated map of stress distribution within the target assembly as a function of time determined by the backward characteristics analysis (see Methods). The horizontal dashed lines represent the material layers in Lagrangian coordinates. (D) Calculated stress history of the SiC sample as a function of time. The vertical dashed lines represent stress states over the X-ray probe period.



625 **Figure 5. Measured diffraction peak positions and densities of SiC as a function of stress.**
626 (A) The observed d -spacings of ramp-compressed SiC sample and diamond ablator/window are
627 shown as red and gray circles, respectively. The solid lines are from static equation of state data
628 and their extrapolation as described below. The diamond ramp equation of state⁶⁴ and its
629 extrapolation (>800 GPa) are shown as gray dashed curves. (B) Measured densities of ramp
630 compressed SiC (red circles) and a Birch-Murnaghan (BM) equation of state fit to the reduced
631 isentrope (red curve with assumption of $Y=20$ GPa and $\gamma_0=1$) are compared to the extrapolated
632 EOS of diamond anvil cell experiments (blue²⁰ and black¹⁸ curves) and the range of densities
633 obtained from first principles calculations (orange shaded region^{20,26-33}, see Supplementary Table
634 S3). The gray band shows the pressure-density path along the principal isentrope calculated
635 using the thermodynamic parameters that are listed in the Suppl. Table S4. Our EOS curve is
636 consistent with the theoretical calculations of ref.²⁶ (purple curve), which is only partially visible

637 due to the overlap with our EOS curve. Previous shock compression data (cyan¹⁷ and green
 638 circle^{15,16}) and static diamond anvil cell data (open circles¹⁸) are also plotted. The datum at 433
 639 GPa (gray parentheses) was excluded from fitting to the EOS (See Suppl. Table S2 for details).
 640 Gray error bars represent additional +50 GPa stress uncertainty due to uncertainty in the strength
 641 of diamond.
 642

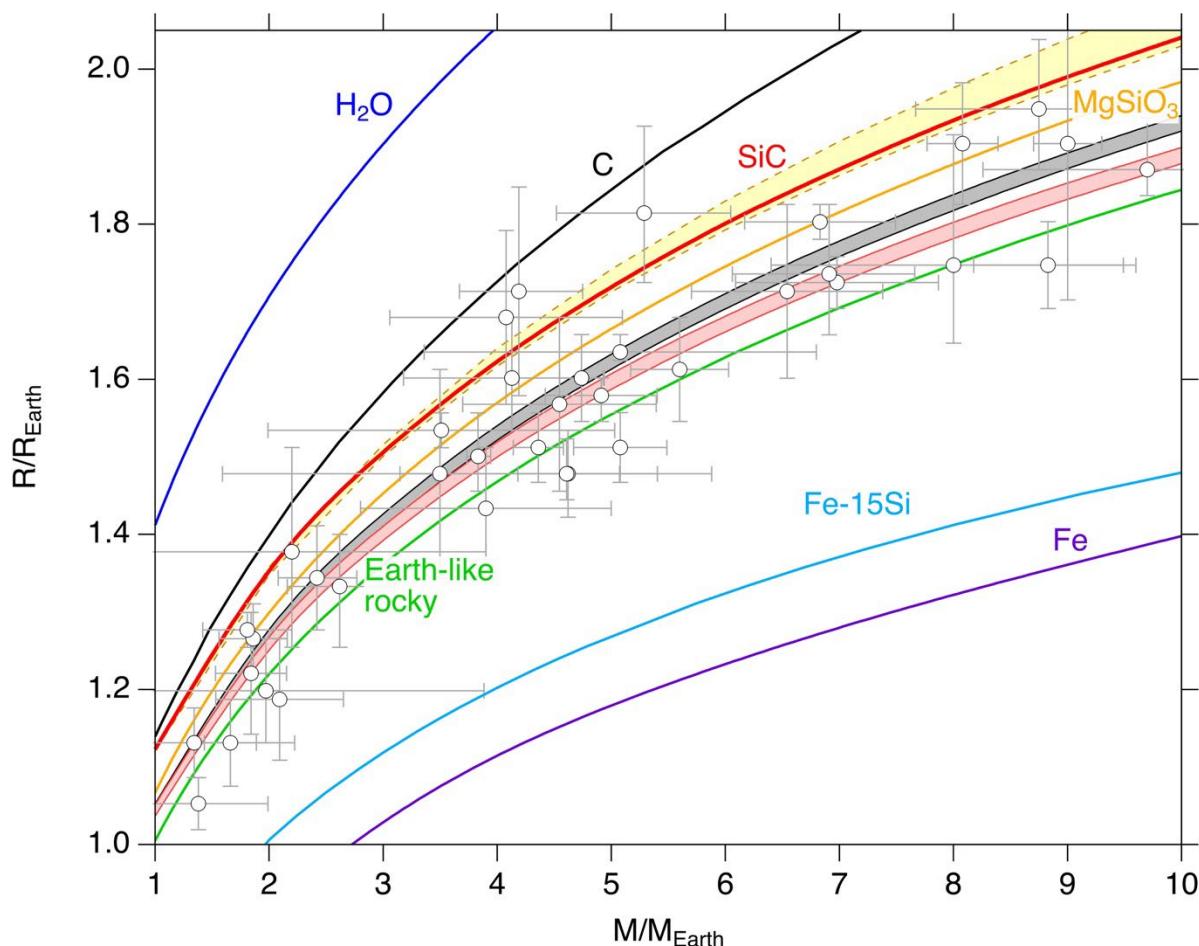


643
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 645 **Figure 6. Interior pressure and density for an eight-Earth-mass carbide planet.** The carbide
 646 planet is modeled with a 30% by mass Fe or Fe-15wt%Si (Fe-15Si) core and a 70% SiC or C and
 647 SiC (1:1 molar fraction) mantle. The addition of Si in the core decreases pressures and densities in

648 the core but increases core size. CMB, core-mantle boundary; TZ, transition zone (due to the B3
 649 to B1 transition in SiC).

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 653 **Figure 7. Calculated mass-radius relationships compared observed exoplanets.** Mass-radius
 654 curves are shown for hypothetical planets composed entirely of SiC (this study), diamond²⁶, H₂
 655 O⁵⁰, MgSiO₃⁵⁰, Fe-15Si²³, and Fe⁷². The mass-radius relationship for an Earth-like rocky planet
 656 is shown in green. The yellow band represents the mass-radius relationship for a SiC planet using
 657 the B1 SiC equation of state from previous studies^{18,20,26-33}. The pink band shows the predicted
 658 mass-radius relationship for a planet consisting of pure SiC mantle and a Fe or Fe-15Si core
 659 (mass fraction of core assumed to be 30%). The gray band shows the mass-radius relationship for
 660 a planet with 50% SiC-50% diamond mantle and a Fe or Fe-15Si core (mass fraction of core
 661 assumed to be 30%). Observed exoplanets are shown as gray circles with error bars
 662 (<http://exoplanetarchive.ipac.caltech.edu>). R_E and M_E are the radius and mass of Earth,
 663 respectively.

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