



# Circum-Arctic Map of the Yedoma Permafrost Domain

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Ice-rich permafrost in the circum-Arctic and sub-Arctic (hereafter pan-Arctic), such as late Pleistocene Yedoma, are especially prone to degradation due to climate change or human activity. When Yedoma deposits thaw, large amounts of frozen organic matter and biogeochemically relevant elements return into current biogeochemical cycles. This mobilization of elements has local and global implications: increased thaw in thermokarst or thermal erosion settings enhances greenhouse gas fluxes from permafrost regions. In addition, this ice-rich ground is of special concern for infrastructure stability as the terrain surface settles along with thawing. Finally, understanding the distribution of the Yedoma domain area provides a window into the Pleistocene past and allows reconstruction of Ice Age environmental conditions and past mammoth-steppe landscapes. Therefore, a detailed assessment of the current pan-Arctic Yedoma coverage is of importance to estimate its potential contribution to permafrost-climate feedbacks, assess infrastructure vulnerabilities, and understand past environmental and permafrost dynamics. Building on previous mapping efforts, the objective of this paper is to compile the first digital pan-Arctic Yedoma map and spatial database of Yedoma coverage. Therefore, we 1) synthesized, analyzed, and digitized geological and stratigraphical maps allowing identification of Yedoma occurrence at all available scales, and 2) compiled field data and expert knowledge for creating Yedoma map confidence classes. We used GIS-techniques to vectorize maps and harmonize site information based on expert knowledge. We included a range of attributes for Yedoma areas based on lithological and stratigraphic information from the

source maps and assigned three different confidence levels of the presence of Yedoma (confirmed, likely, or uncertain). Using a spatial buffer of 20 km around mapped Yedoma occurrences, we derived an extent of the Yedoma domain. Our result is a vector-based map of the current pan-Arctic Yedoma domain that covers approximately 2,587,000 km<sup>2</sup>, whereas Yedoma deposits are found within 480,000 km<sup>2</sup> of this region. We estimate that 35% of the total Yedoma area today is located in the tundra zone, and 65% in the taiga zone. With this Yedoma mapping, we outlined the substantial spatial extent of late Pleistocene Yedoma deposits and created a unique pan-Arctic dataset including confidence estimates.

**Keywords:** late Pleistocene, syngenetic permafrost, Beringia, Siberia, Alaska, Yukon

## 1 INTRODUCTION

Vast portions of the Arctic and sub-Arctic regions (hereafter pan-Arctic), predominantly in Siberia, the Far East of Russia, Alaska and Yukon, are covered by ice-rich fine-grained permafrost deposits that contain large late Pleistocene syngenetic ice wedges (Sher et al., 1971; Kanevskiy et al., 2011; Schirrmeister et al., 2013). Accompanied by wedge-ice growth in polygonal tundra-steppe landscapes, the sedimentation process was driven by cold continental climatic and environmental conditions in unglaciated regions during the late Pleistocene, inducing the accumulation of the unique Yedoma deposits up to 50 m thickness. Because of the fast incorporation of material into syngenetic permafrost and ice wedge growth during its formation, Yedoma deposits include a high volume of ground ice (including pore, segregated, and wedge ice) and preserved organic matter consisting of floral and faunal fossil remains that are perennially frozen and distinct from the recent organic matter cycles and have remained relict for tens of thousands to hundreds of thousands of years (Schirrmeister et al., 2011; Strauss et al., 2017). The first descriptions of Yedoma deposits date back to the early 19th century, when Arctic explorers described these deposits erroneously as buried glacial ice remnants based on the obvious high amount of “pure” ice that was observed in coastal exposures (von Kotzebue, 1821).

The permafrost region and the Yedoma domain in particular have been characterized as a very large and potentially vulnerable organic carbon pool (Schuur et al., 2015). Recently, the Yedoma domain has even been discussed as one of the potential climate tipping elements (Lenton et al., 2019). Scientific interest in Yedoma is rising as, besides the vulnerability of its frozen organic matter pool to degradation, melt of the high excess ground ice upon Yedoma thaw will cause substantial ground volume loss. Resulting surface subsidence will pose a serious threat to any infrastructure built on permafrost (de Grandpré et al., 2012; Hjort et al., 2018; Streletskiy et al., 2019; Chen et al., 2021; Schneider von Deimling et al., 2021). More broadly, Yedoma thaw implies substantial consequences for landscape reorganization by surface subsidence (Günther et al., 2015; Antonova et al., 2018), thermal erosion (Kanevskiy et al., 2016; Fuchs et al., 2020; Shur et al., 2021b; Morgenstern et al., 2021), thermokarst formation (Jones et al., 2011; Nitze et al., 2017; Ulrich et al., 2017; Veremeeva et al., 2021), and relief

inversion as well as land loss by coastal erosion (Günther et al., 2013; Farquharson et al., 2018). Such changes of the periglacial landscape inventory certainly affect surface and sub-surface drainage patterns across scales, trigger ecosystem responses, and thus alter floral and faunal species composition and distribution (Pastick et al., 2019; Jones et al., 2020).

Previous studies estimated that Yedoma deposits cover about one million km<sup>2</sup> of a large region in central and eastern Siberia (Romanovskii, 1993; Zimov et al., 2006), especially Yakutia (Grigoriev and Kunitsky, 2000; Konishchev, 2011), although these estimates are generally based on maps with scales smaller than 1:10,000,000. During the late Pleistocene, the extent of Yedoma and Yedoma-like deposits may have been up to three million km<sup>2</sup> (Zimov et al., 2006; Walter et al., 2007; Strauss et al., 2017). Detailed compilations of current Yedoma distribution were started by using Quaternary geological maps at scales of 1:500,000 to 1:1,000,000 for Northeast Siberia (Grosse et al., 2013). Map analysis and field studies showed that significant areas of Yedoma deposits were reworked by thermokarst and thermal erosion due to the climate warming at the end of the Pleistocene and beginning of the Holocene with extensive thermokarst lake formation, expansion, and drainage (Kaplina, 2009; Morgenstern et al., 2013; Lenz et al., 2016). Estimations of Yedoma and drained lake basin coverage in the Yana-Indigirka and Kolyma Lowlands show that the modern area of drained lake basins could reach up to 80–100% in low-lying relief regions with Yedoma remnants covering now less than 20% while in elevated regions 40–55% of Yedoma uplands remained (Veremeeva, 2016; Veremeeva et al., 2021). Building on previous efforts, the objective of our present study is to compile the first detailed circum-Arctic digital map of the Yedoma domain coverage using high and medium resolution geological and stratigraphical maps from the regions of former Beringia, the unglaciated landmass consisting of today's Siberian (Russia), Alaskan (United States) and Yukon (Canada) territories.

## 1.1 Terminology

### 1.1.1 Origin of the Term Yedoma

There is no clear agreement on the origin of the term “Yedoma” and its association with ice-rich fine-grained sediments. One version is that “Yedoma” was originally used by local people in Yakutia in a geomorphological sense to describe hills composed of Ice Complex deposits, which are “eaten” (Russian

root yed from yest' = to eat, to reduce, to erode) by thermokarst that is forming typical landforms of alternating hills with lake depressions. However, there is no absolute confidence regarding this meaning. The term Yedoma was also used in the north-European part of Russia, where its originality from "samoyed" (Finnish language "erämaa") is hypothesized with the meaning of remote forests or pasture and something "far away" or "wilderness" (Chudinov, 1910).

### 1.1.2 Previous Scientific Use of the Term Yedoma

As summarized by Kanevskiy et al. (2011) and Schirrmeister et al. (2013), the terminology used in scientific publications for the ice-rich, syngenetically frozen deposits of late Pleistocene age varies across Beringia. These deposits are often referred to as "Ice Complex", "Yedoma", or "Edoma" in Siberia, while in North America other terms such as "muck" or "upland silt" are partially considered as equivalent terms (Péwé, 1975; Shur et al., 2004; Froese et al., 2009). The term "muck", originating from gold mining operations, seems to have the broader meaning of an overburden material including ice-rich deposits of Holocene age (Péwé, 1975), but with remnants dating back to at least 700 ka in Yukon (Froese et al., 2008). Solov'ev (1959) defined "Ice Complex" as frozen deposits of various age, composition, genesis, and thickness, with syngenetic ice wedges. Ice Complexes formed in the Pleistocene and the oldest directly dated Ice Complexes are the mid-Pleistocene Yukagir Suite of Bol'shoi Lyakhovsky Island (New Siberian Islands) with an age of about 2,00,000 years (Tumskoy, 2012; Wetterich et al., 2019). The lower Ice Complex of the Batagay Megalump was even dated to an age of about 650 000 years (Murton et al., 2021). Usually, the Yedoma Ice Complex formed in late Pleistocene during the period 60,000–12,500 years (Sher et al., 1971; Sher et al., 1987; Schirrmeister et al., 2002a; Kaplina, 2009; Strauss et al., 2017). It is considered a dominant relief-forming deposit in many regions of western Beringia. The term "Yedoma Suite", describing the late Pleistocene Ice Complex, was first introduced into the regional stratigraphic scheme of northeastern Siberia (Sher et al., 1987). As there are many usages and viewpoints of "Yedoma" in literature, it is difficult to agree on one overall definition. Especially, uncertainties using this term in mining and scientific traditions complicate a circum-Arctic Yedoma study. Nevertheless, it is agreed that ice constitutes the largest portion of Yedoma deposits with 60–95 volume percent (vol %) including pore, segregated, and wedge ice (Kanevskiy et al., 2011; Schirrmeister et al., 2013; Strauss et al., 2013; Ulrich et al., 2014; Kanevskiy et al., 2016; Strauss et al., 2017; Wetterich et al., 2019; Schirrmeister et al., 2020; Wetterich et al., 2020; Kanevskiy et al., Forthcoming 2021). Further, clastic and organic Yedoma components (Schirrmeister et al., 2011; Schirrmeister et al., 2013) suggest aeolian, fluvial, or slope sedimentation with simultaneous freezing (i.e. syngenetic permafrost formation).

### 1.1.3 Yedoma Definition

Our summary definition considers that Yedoma deposits are perennially frozen (permafrost), fine-grained, organic-bearing, and ice-rich sediments of late Pleistocene age, that they can reach thickness of up to tens of meters and are framed by large

syngenetic ice wedges, and that they are preserved in remnant Yedoma uplands. With this definition, we cover the Yedoma Suite and other types of the late Pleistocene Ice Complexes and late Pleistocene muck, being located in West, Central and East Siberia, in Chukotka and the Russian Far East as well as in Alaska and Yukon areas. Even though syngenetic Pleistocene permafrost sediments mostly have large ice wedges, this is not true for all cases (e.g. Fraser and Burn, 1997; Froese et al., 2009).

### 1.1.4 Yedoma Domain

Analogue to the permafrost region, where "just" a part of the sediments are perennial frozen (e.g. >90% coverage in the continuous permafrost region below the active layer), the "Yedoma domain" term refers to the region where Yedoma deposits are expected to occur but may also include drained lake basins, river valleys etc. Thus, this region is an outline of the maximum extent of Yedoma deposits during the late Pleistocene.

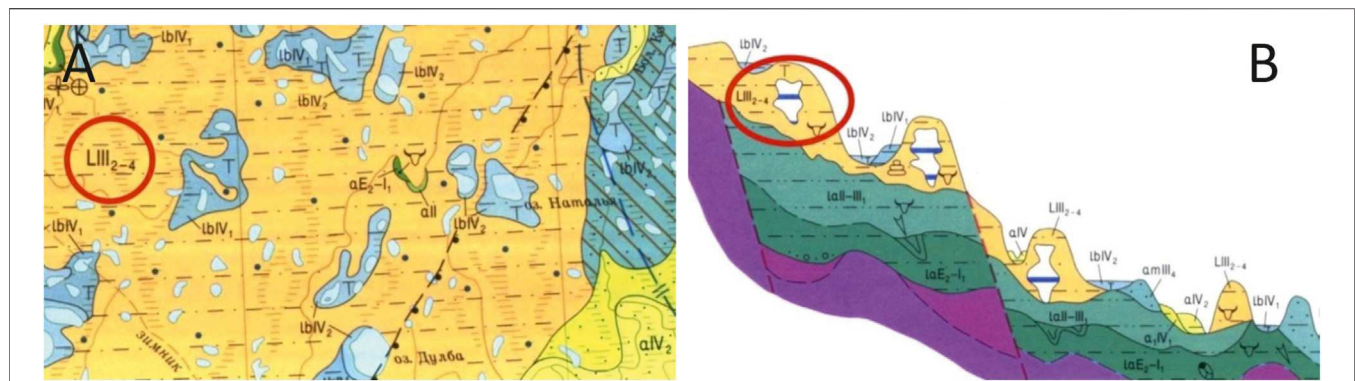
## 2 MATERIALS AND METHODS

We continued the work of Grosse et al. (2013) to extend their digital database of Yedoma distribution in East and Central Siberia by updating and merging different sources such as maps, scientific descriptions and synthesizing scientific literature.

### 2.1 Digitization of Maps

Our study focused on extracting geologic units from existing maps that are labeled as Yedoma or that we interpreted as Yedoma, based on cryolithology, ground ice conditions, geochronology, geomorphologic, and spatial associations (Figure 1). For compiling this map, we used maps of the previous Yedoma coverage estimates, included the digitized areas from Grosse et al. (2013), and extracted areas of potential Yedoma distribution from additional surface geological and Quaternary geological maps:

- 1:500,000 Geological maps of Yakutia (**Supplementary Table S2**)
- 1:1,000,000 Quaternary formation maps of Russia (**Supplementary Table S3**)
- 1:2,500,000 Quaternary map of the territory of Russian Federation (**Supplementary Table S4**)
- Permafrost Characteristics of Alaska
- Late Quaternary Aeolian Deposits of Northern North America: Age and Extent, Geological Survey of Canada (Wolfe et al., 2009), used in version 1 of the map available at Strauss et al. (2016)
- Yukon Digital Surficial Geology Compilation as digital 1:50,000 (50k), 1:100,000–125,000 (100k–125k) and 1:250,000 (250k) datasets (replaced the "Late Quaternary Aeolian Deposits of Northern North America" in this dataset, Yukon Geological Survey, 2014)
- Dataset "The Lena River Delta - inventory of lakes and geomorphological terraces" by Morgenstern, et al. (2011)
- Digital Database and Maps of Quaternary Deposits in East and Central Siberia by Bryant et al. (2017).



**FIGURE 1** | Example subsets of the maps with Yedoma units. This example from Grosse et al. (2013) was taken from map R-(55)-57. **(A)** Map view: Orange-colored unit LIII<sub>2-4</sub> (label circled in red) is described in the legend as loess-like deposits (L) of late Pleistocene stratigraphy (III), specifically covering the periods Zyryan, Kargin, and Sartan (2–4). The lithologic signature (dash-dot-dash) illustrates that this unit consists of mostly silty to fine-sandy deposits, and the dark blue dots indicate very high ground ice contents. Abundance of lakes (light blue) and lacustrine deposits (medium and dark blue, lbIV<sub>1</sub> and lbIV<sub>2</sub>) within the unit indicate degradation of at least the upper part of the Yedoma unit by Holocene thermokarst. **(B)** Cross-profile view: the same unit LIII<sub>2-4</sub> (red circle) is shown to blanket terrace slopes. Bison skull symbol indicates the presence of fossil bones of the mammoth fauna. White features with blue stripes (e.g. in the red circle) indicate the presence of large syngenetic ice wedges typical for Yedoma.

- Quaternary deposits map of Yana-Indigirka and Kolyma lowlands tundra zone, R-55-57, based on Landsat imagery by Veremeeva (2021).
- Sedimentological, biogeochemical and geochronological data of Yedoma and thermokarst deposits in West-Alaska by Jongejans et al. (2018a) and Jongejans et al. (2018b).

**Supplementary Figure S5** shows the approximate map frames of the used sources. The digitization was done using GIS techniques (ArcGIS) and semiautomated vectorization of raster images (Adobe Photoshop and Illustrator). Using digitized geologic maps does not violate the spatial resolution of the specific data. Even though the digital form of the map removes the constraints imposed by the scale of a paper map, the detail and accuracy represents various scale-specific generalization levels of each map. Similarly, where this database is used in combination with other data of higher resolution, the resolution of the combined output will be limited by the lower resolution of these data.

The original geological Quaternary maps of Russia were issued by the Department of Natural Resources of the Russian Federation or its predecessor the Department of Geology of the Soviet Union and have their foundation in decades of geological field and remote sensing work and mapping at scales 1:50,000 to 1:500,000 by Russian geologists and cartographers in the respective regions.

We eventually merged the Yedoma spatial units extracted from across these multiple maps into a single data layer for the pan-Arctic domain. We further synthesized data on the thickness of Yedoma for our database using deposit thickness information provided on the maps and from published datasets from drill cores and exposures.

### 2.1.1 Geological Maps of Yakutia, Scale 1:500,000

We used nine geological maps of Yakutia of scale 1:500,000 to refine the spatial distribution of Yedoma deposits. The data were digitized and edited at a scale of 1:500,000. The Yedoma

identification criteria are shown in the supplement (**Supplementary Table S2**).

### 2.1.2 Russian Quaternary Formation Maps, Scale 1:1,000,000 Map

We used 23 Quaternary formation maps of scale 1:1,000,000 to refine the spatial distribution of Yedoma deposits. The data was edited at a scale of 1:1,000,000. On these maps, we also extracted point features, which contain information about boreholes in potential Yedoma deposits. Yedoma selection criteria are listed in **Supplementary Table S3**.

### 2.1.3 Quaternary Map of the Territory of Russian Federation, Scale 1:2,500,000

The 1:2,500,000 scaled map published in 2010 by the Ministry of Natural Resources and Ecology of the Russian Federation, Federal Agency of Mineral Resources was used to extract the areas declared as “Ice loess (Yedoma) deposits” (**Supplementary Table S4**).

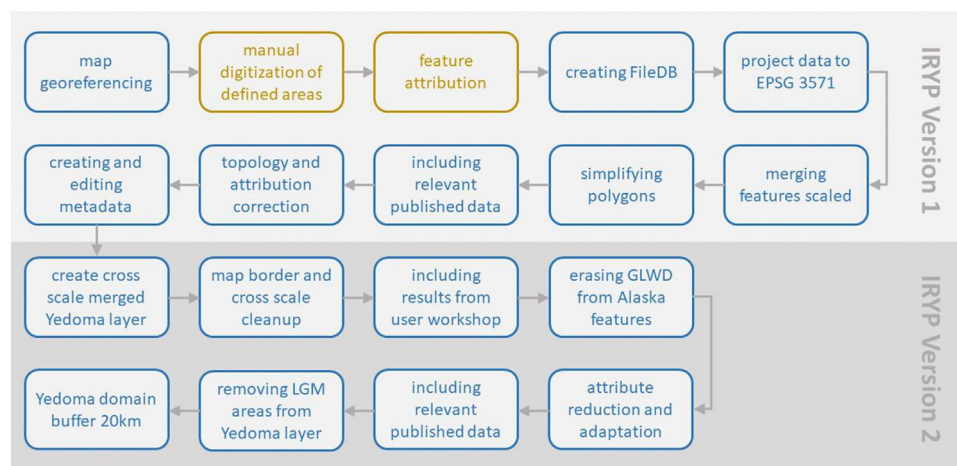
### 2.1.4 Alaskan Permafrost Map, Scale 1:7,200,000

For Yedoma deposits in Alaska, we used the “Permafrost Characteristics of Alaska” map, scale 1:7,200,000 published by Jorgenson et al. (2008). Here, we extracted the areas declared as “loess” and “ice rich” (**Supplementary Table S5, S6**). To reduce the area of this highly generalized Yedoma polygon features and to improve the Yedoma domain estimate, we used the “Global Lakes and Wetlands Database” (GLWD) (Lehner and Döll, 2004) and clipped the Yedoma domain polygon with GLWD lake polygons. Moreover, we used Yedoma distribution maps for Alaska derived from extensive geomorphologic mapping in the field and literature synthesis by Kanevskiy et al. (2011; 2016; Forthcoming 2021) for a visual plausibility check.

### 2.1.5 Yukon Loess Map

For the western Yukon loess, we used the map “Loess of Northern North America” published by Wolfe et al. (2009) in the dataset





**FIGURE 2 |** Workflow of the major steps of the Ice-Rich Yedoma Permafrost (IRYP) mapping from the scanned Russian maps or Alaskan and Canadian datasets to data interpretation, data harmonization to compiling the overall map. Software applied in these steps included ArcGIS (blue) and Avenza MAPublisher (yellow).

version in Strauss et al. (2016). In this version, we refined it by using the Yukon Digital Surficial Geology Compilation (Yukon Geological Survey, 2014). We selected the Yukon 1:50,000, 1:100,000–125,000 and 1:250,000 maps. The selection criteria are shown in the supplement (**Supplementary Table S8**). We clipped this with late Pleistocene glacial areas to avoid overestimation.

### 2.1.6 Inclusion of Other Datasets

There are a few datasets with more detailed mapping. For example, Veremeeva (2021) used Landsat images that allow the area of the Yedoma uplands and alases (partially to fully drained depressions formed by thermokarst, i.e. ground-ice melt induced surface subsidence and lake formation) to be clearly identified. We included this detailed dataset in our circum-Arctic version. Moreover, we used the difference between our first estimate for this region compared to Veremeeva (2021) as an indicator for and potential overestimation of Yedoma coverage, also caused by ongoing degradation (thermokarst and other processes).

For the Baldwin Peninsula, one of the first (1826) sites where Yedoma was encountered scientifically (at Cape Blossom (Beechey, 1831, see; Shur et al., Forthcoming 2021a), we integrated the Yedoma areas defined in the land cover classification map published by Jongejans et al. (2018) (**Supplementary Table S1**). We used data from Morgenstern et al. (2011) for refining the Lena Delta and Bryant et al. (2017) as well as Grosse et al. (2013) for East and Central Siberia.

## 2.2 Dataset Merging and Harmonization

To harmonize the different datasets and to avoid merging artifacts, we applied map edge cleaning while merging data from different database layers. For the digitalization and spatial integration, we used Adobe Photoshop CS6 (Version: 13.0 x64), Adobe Illustrator CS6 (Version 16.0.3 x64), Avenza MAPublisher 9.5.4 (Illustrator Plug-In) and ESRI ArcGIS 10.2.2 for Desktop (Advanced License). The applied working steps were

done as described in **Figure 2** and the **Supplementary Methods** in detail.

To improve working efficiency with many georeferenced maps in a map view, overlapping map parts, legends, and further information were clipped. For georeferencing, we used map corner coordinates as control points. For older maps, more (up to 50) control points had to be set at grid intersections. Then, to simplify further data processing, the geological maps were reprojected using WGS84 as datum to a common projection. In order to define the relevant map contents, we used different sources from different continents for our study. Even if having the same publisher, which is the case for the Russian geological maps, the maps were compiled with diverse geological foci and during different times. For every map, we defined the relevant categories, and for the Russian maps, the legend content had to be translated to English. After this, we used Avenza MAPublisher (Adobe Illustrator) for vectorization. From the Russian geological maps in 1:500,000 scale, we digitized 818 (multi-)polygon features. For the 1:1,000,000 maps, about 2,138 (multi-) polygon features and 75 point features of deposit thickness were digitized, and from geological maps of scale 1:2,500,000, we deduced about 771 (multi-)polygon features. We connected the digitized polygons to the specific relevant attributes. We provided information for stratigraphic units that were assigned to Yedoma in **Supplementary Tables S5–S8**, as well as in Grosse et al. (2013 table D1).

Subsequently, the Yedoma database was built by merging (projection: EPSG: 3571 WGS 1984 North Pole LAEA Bering Sea) the different data extracts. We defined topology rules to identify and solve errors like gaps between bordering areas and areas overlapping within a layer. Some content errors have been resolved by checking the dataset manually. For the Alaskan dataset, we cut out lakes and used the “Global Lakes and Wetlands Database” to clip lake polygons. We removed overlaps of the generated Yedoma polygons with areas of LGM glaciation. This removed a large portion of the areas generated from the

Yukon Digital Surficial Geology Compilation, and a few small areas in Western Siberia.

## 2.3 Field Sites and Photos

For Yedoma ground truthing, we synthesized available field sites from scientific literature data on Yedoma characteristics and occurrences. Our database includes more than 360 boreholes and exposures. These points were summarized by Schirrmeister et al. (2011), Kanevskiy et al. (2011) and updated with additional points by Strauss et al. (2013) and Strauss et al. (2017). For Chukotka we use a field guide compiled by Kotov and Brazhnik (1991).

As an additional visualization opportunity, we initiated the Yedoma photo database as a feature collection, having a point as spatial representation. (**Supplementary Figure S2**). This point feature class (collection/compilation of points in a geospatial layer) includes the spatial location of photos showing Yedoma, which were taken by several members of the International Permafrost Association (IPA) Yedoma Action Group. The picture's credits and license information are given in the attribute table.

To get a general impression of Yedoma thickness, which is needed for volume estimation, we included the data synthesized by Strauss et al. (2013; 2017). We compiled field and literature data for available Yedoma depth estimates and updated this with newly available or newly recognized depth estimates (see **Supplementary Table S9**).

## 2.4 Inclusion of Expert Knowledge

During the International Conference on Permafrost (ICOP) held in Potsdam in 2016, a 2-day Yedoma workshop took place (**Supplementary Figure S3**). Preliminary maps were discussed and checked by a group of experts. For this reason, the IRYP map was split into four regions: Western and Central Siberia (supervised by I. Streletskaia), East Siberia/Yakutia (supervised by V. Kunitsky and C. Siegert), Far East Siberia and Chukotka (supervised by D. Shmelev and G. Kraev) and Alaska and Canada (supervised by Y. Shur and D. Froese). We used first-hand field knowledge from various Yedoma sites and scientific literature in order to decide if digitized areas should be attributed to Yedoma.

## 2.5 Confidence Estimation and Area Estimate Harmonization

Despite studies on Yedoma mystery and controversy for more than 150 years (Shur et al., *Forthcoming 2021a*), we had to agree on a technical way to compile Yedoma deposits on the broad scale, including local peculiarities. To account for this level of controversy and uncertainty, we implemented a Yedoma confidence classification system based on source details and additional information such as ancillary data on lithology and stratigraphy, which can be attributed to Yedoma. This data is often available in the original source maps and helped us to define our confidence level for our Yedoma classification of a specific area or region. We included a range of attributes for Yedoma area polygons based on lithological and stratigraphical information from the original source maps as well as a confidence level for our

classification of an area as Yedoma (three levels: confirmed, likely, or uncertain, corresponding to high, medium, and low confidence). We based the confidence level on the amount of independent sources providing information about the sites. In case that just the stratigraphic context of the maps matched the Yedoma description, we classified the Yedoma as “uncertain” (**Table 1**). When one map explicitly defined a certain area as Yedoma, and this classified polygon intersected with polygons of other maps, or if lithologic and stratigraphic context were plausible for Yedoma, we used the classification “likely”. We reached the level “confirmed” when field data was available, or it was mentioned on the map explicitly as Yedoma, or the context of lithology and close field data fit, or experts agreed that this is Yedoma.

As the Alaskan Yedoma domain is mapped in less detail when compared to the Siberian (see **Figure 3**) or the Canadian datasets, and thus including much more areas and other areas, we applied a correction factor of 0.5 for the Alaskan area estimation to avoid overestimation of the Yedoma area (**Supplementary Table S10**). This factor is a first guess based on findings from Veremeeva (2021) reducing our mapped areas by 50%, even though it is in Siberia with better spatial resolution (**Supplementary Table S10**).

## 2.6 Estimating the Yedoma Domain Outline

For getting the outline of the maximum Yedoma domain, and to be able to compare it to older regional estimates (**Supplementary Figure S4**), we decided to add a buffer of 20 km to each digitized Yedoma polygon and merge these areas. This buffer distance was chosen as the average to compensate for the lower resolution of older maps as well as a reasonable distance to account for the fact that natural borders are never as sharp as illustrated in the maps. The buffer was clipped with the coastline and the resulting area is shown in **Figure 4**.

# 3 RESULTS

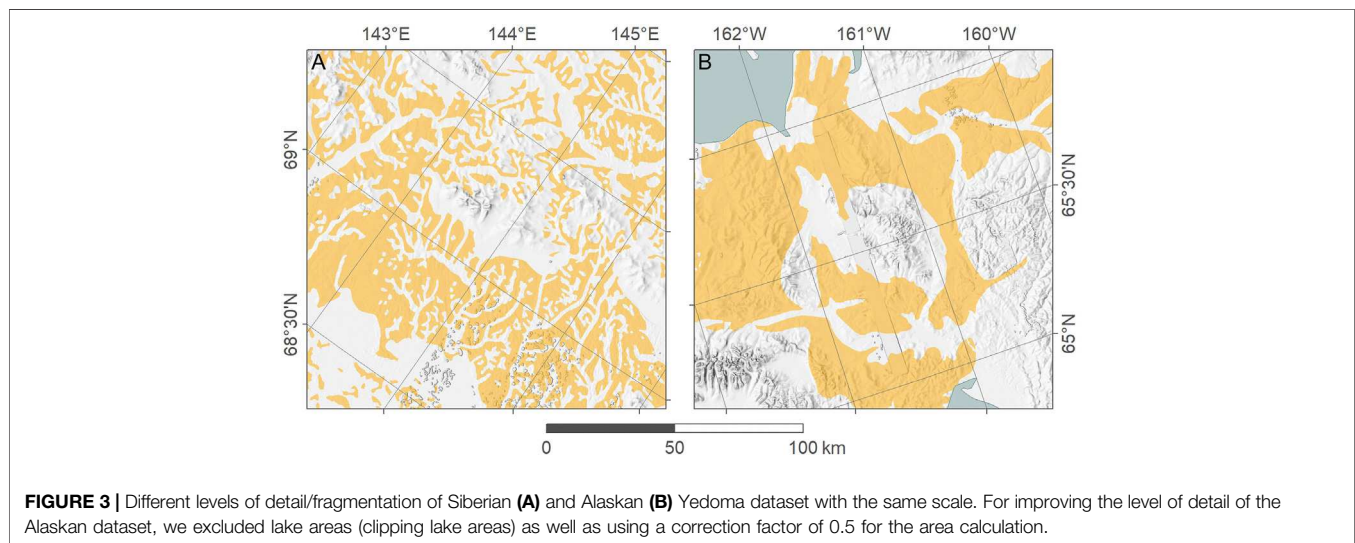
## 3.1 Coverage and Regional Classification

We estimated the Northern Hemisphere Yedoma domain to cover approximately 2,587,000 km<sup>2</sup> located between the Yamal Peninsula in Russia all the way to the Yukon, Canada (**Figure 4**). In this domain, we estimate that 480,000 km<sup>2</sup> to be underlain by Yedoma. 35% of the total Yedoma area today is located in the tundra zone, and 65% in the taiga and boreal ecozone (tree line from Walker et al., 2005).

Our database includes more than 360 boreholes and exposures and more than 13,600 digitized Yedoma areas. For the area calculations please note the above-mentioned Alaskan correction factor of 0.5. The globally northernmost Yedoma occurrence was found on 138.9°E, 76.19°N (Kotelný Island, confidence class: confirmed), the easternmost in 136.0°W, 62.88°N (close to the village of Pelly Crossing, Canada, confidence class: uncertain), the southernmost in Central Yakutia 127.3°, 61.0° (close to the village Sinsk, Yakutia, confidence class: confirmed), and the westernmost 83.0°E, 70.3°N (Yenisei River, confidence class: uncertain). If only including the confirmed areas, the latitudinal extent stretches from 76.19°N to 61.0°N and longitudinal extent from 141°W to 83.28°E.

**TABLE 1 |** Confidence classification scheme. In the column “confidence ID”, the first number is connected to the confidence classification’ and the second number to the confidence source.

Confidence classification	conf_source (the source that constitutes the confidence)	Confidence ID (classification and source)
confirmed (1)	1_field data	11
confirmed (1)	2_explicit classification in map	12
confirmed (1)	3_context of lithology and field data	13
confirmed (1)	4_expert_knowledge	14
likely (2)	1_intersection with explicit classification of 2.5M map	21
likely (2)	2_lithologic and stratigraphic context	22
uncertain (3)	1_stratigraphic context	31



Separated from east to west (a bit misleading as we are crossing the international date line), 19% of the Yedoma area is found in North America (18.5% Alaska, 0.1% in Canada) and 81% in northern Asia.

### 3.2 Confidence

We found 83% (399,060 km<sup>2</sup>) of our digitized Yedoma areas to be confirmed, 5% (21,873 km<sup>2</sup>) to be likely and 12% (58,808 km<sup>2</sup>) to be uncertain. Looking at the Yedoma domain (Figure 4) and its confidence estimates (Figure 5), Yakutia proves to be the heartland of Yedoma.

### 3.3 Available Products

All products from the present Yedoma mapping attempt are based on the Ice-Rich Yedoma Permafrost (IRYP) database. A preliminary version was published by Strauss et al. (2016) and with this study, we enhanced and finalized the database. All reported results and products of this study are based on this second version (IRYPv2).

The main product of this study is the map (Figure 4 and map version in the supplement) and its confidence estimate (Figure 5 and map version in the supplement).

Moreover, we worked closely together with the geospatial visualization team of the Alfred Wegener Institute (Germany;

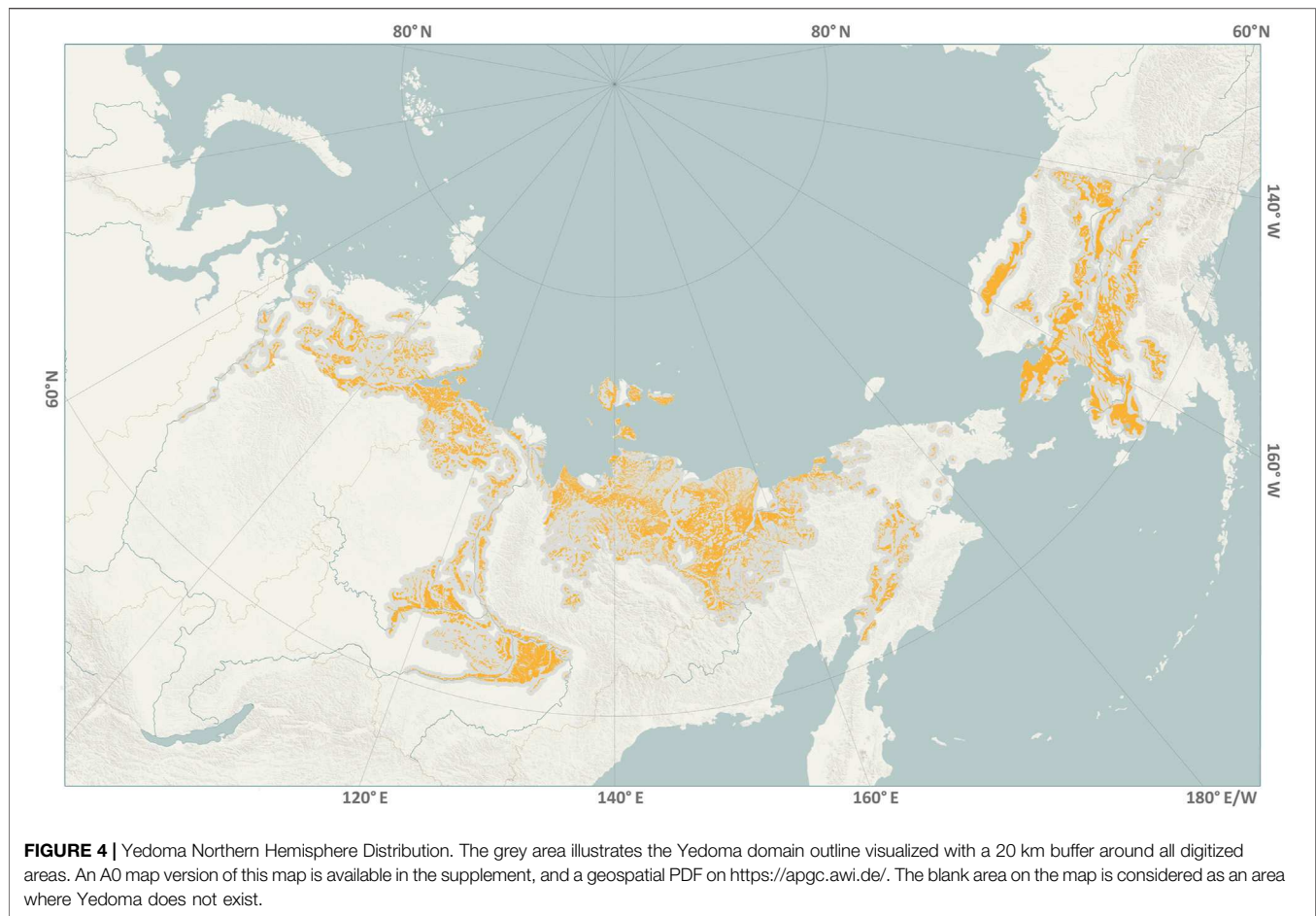
maps@awi.de) and the Arctic Permafrost Geospatial Centre (APGC; <https://apgc.awi.de/>) to make a WebGIS available allowing for interaction with the spatial data sets (Supplementary Figure S1). We further published a visual tool called the Yedoma picture database (Supplementary Figure S2) that is hosted and visualized by the APGC and PANGAEA, whose pictures were used for the confidence assessment.

## 4 DISCUSSION

### 4.1 Previous Large Scale Mapping of Yedoma Distribution

Various maps of Yedoma distribution exist in the literature. However, most of them are small-scale (1:10,000,000 or smaller) maps that tend to under or overestimate the spatial extent of Yedoma. The Yedoma map most widely referred to is that of Romanovskii (1993) (Supplementary Table S1 and Supplementary Figure S4), but the first Yedoma map, to our knowledge, was published by Tormidiaro (1980) (Supplementary Table S1 and Supplementary Figure S4). In another adaptation of the small-scale map of Yedoma distribution, Konishchev (2011) separated the Yedoma deposits into three categories of spatial





distribution, ranging from widely distributed to fragmented to sporadic (**Supplementary Table S1**).

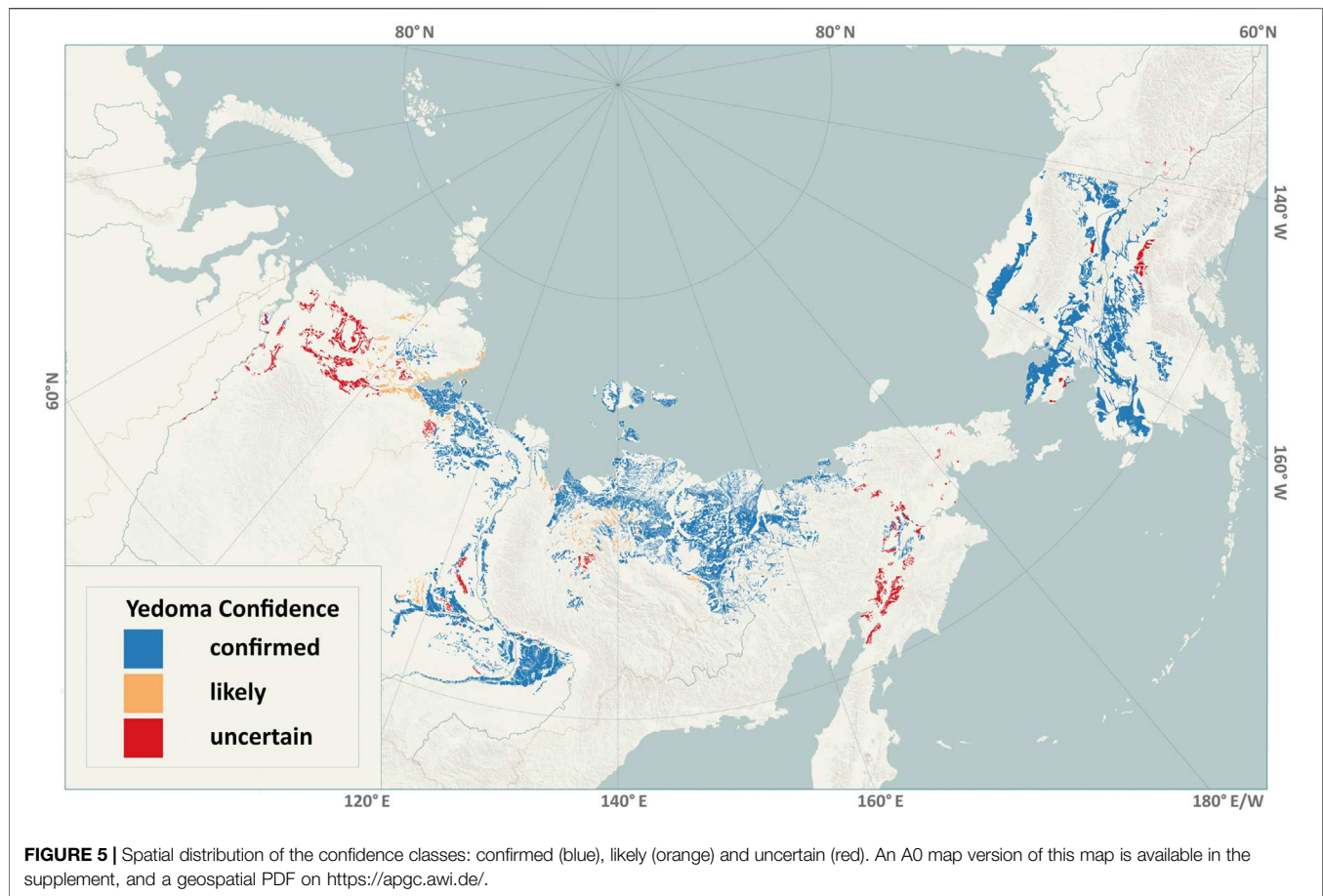
The area covered by Yedoma in Romanivskii's (not including spatially not quantified "Yedoma in river valleys") map is 1,141,390 km<sup>2</sup>. This map has been subsequently adapted by Siegert and Romanovskii (1996) and was furthermore published by Schirrmeister et al. (2002b), showing Yedoma extent in central and northeastern Siberian lowlands and indicating additional Yedoma occurrences in formerly unglaciated valleys. Besides Romanovskii (1993), we digitized the maps by Tormidiaro (1980), Grigoriev and Kunitsky (2000) and Konishchev (2011) (**Supplementary Figure S4** and **Supplementary Table S11**).

There are two other citable estimates for the extent of the Yedoma region in Siberia: 1,000,000 km<sup>2</sup> assumed by Zimov et al. (2006), and 1,186,000 km<sup>2</sup> estimated by Strauss et al. (2013). In our study, we estimated an extent of 1,957,885 km<sup>2</sup> for the Siberian Yedoma domain, which is by about 49, 42 and 39% more than found in Zimov et al. (2006), Romanovskii (1993), and Strauss et al. (2013), respectively. The main reason for this is that we included a broader east-to-west extent, as well as that Romanovskii's (1993) not quantifiable signatures on Yedoma distribution in river valleys (**Supplementary Table S1**).

## 4.2 Previous Regional Mapping of Yedoma Distribution

Several maps of regional Yedoma extent exist, for example, for the Yana-Indigirka and Kolyma lowlands (Tomirdiario, 1980), or Yakutia (Grigoriev and Kunitsky, 2000) (**Supplementary Table S1**). The areas of all these maps are not comparable to our new numbers, as field studies and high-resolution satellite imagery clearly indicate that the Yedoma deposit cover in the Yedoma domain is not uniform, but can range from complete to heavily fragmented remnants (e.g. **Figure 3**). For example, Grosse et al. (2006) indicated that only about 22% of the Yedoma surface area in the Lena Anabar lowland near Cape Mamontov Klyk was preserved in its original state. Within the Yana-Indigirka lowland, the Yedoma area on Buor-Khaya Peninsula is 15% (Günther et al., 2013) and on Shirokostan Peninsula 42% (Tarasenko et al., 2013). In the State Geological Map from the year 2000, 31% is covered by alases and 27% by Yedoma in the Indigirka-Kolyma region (Shmelev et al., 2017). Following Lomachenkov (1966), 65–75% of the Yana-Indigirka region is covered by alases and 25% by Yedoma. For the Kolyma Lowland tundra zone, Veremeeva and Glushkova (2016) found a coverage of 72% for alases and 16% for Yedoma. Strauss et al. (2013) generalized the Yedoma domain for Northern Yakutia and found 56% covered by alases and 30% by Yedoma deposits.





This fragmentation in Yedoma extent varies by location and map source. In most cases it is captured in high-resolution scales of 1:1,000,000 and better. Examples for a much higher resolution mapping are Veremeeva (2021), who published an update for the Yana-Indigirka and Kolyma Lowlands (**Supplementary Table S1**) as well as Jongejans et al. (2018b) for the Baldwin Peninsula, and Morgenstern et al. (2011) for Lena Delta Yedoma.

### 4.3 Applicability

The spatial dataset of Yedoma distribution will have broad applications. A range of studies citing the first version of the IRYP dataset (Strauss et al., 2016) suggest that an accurate spatial Yedoma extent dataset is useful in the context of modelling, reconstructing paleo-environmental dynamics and past ecosystems such as the mammoth-steppe-tundra, or assessments of ground ice distribution and future thermokarst vulnerability (e.g. Wild et al., 2019; Nitzbon et al., 2020; Douglas et al., 2021).

Moreover, the map is a crucial improvement needed to refine the present-day Yedoma permafrost organic matter inventory as included in Hugelius et al. (2013) or Strauss et al. (2017).

### 4.4 Dataset Limitations

Our data set is a significant improvement to previous mapping attempts of the Yedoma region. However, when mapping such a

large area in uninhabited terrain, several potential sources of errors need to be discussed such as the fragmentation of Yedoma, the resolution of the final product, the definition of Yedoma itself.

#### 4.4.1 Map Resolution

The fragmentation of the Yedoma extent can be captured in high-resolution scales of 1:1,000,000 and higher. In the State Geological Map (2000) of the Indigirka-Kolyma region (R-55-57), 31% is occupied by alases and 27% by Yedoma, while the more detailed estimate on the same study region using Landsat images reveal an alas area of 50% and a Yedoma area of 13% (Shmelev et al., 2017; Veremeeva, 2021). Therefore, the coverage of Yedoma could be about two times overestimated on maps of 1,000,000 scale due to the high level of initial Yedoma relief dissection by thermokarst and thermal erosion, which could be mapped on a more detailed scale using remote sensing data. Taking our estimate (based on a map of the region which was also quantified by Veremeeva (2021)), our database included 50% more Yedoma area compared to the identical region like in Veremeeva (2021) (**Supplementary Table S10**). For our map, we included Veremeeva's (2021) update for this specific region, but having this in mind, the rest of our circum-Arctic Yedoma mapping approach is potentially overestimating the Yedoma distribution. Moreover, since the year of the map's data acquisition, the degradation of Yedoma deposits likely continued. Especially

when considering that the Yedoma domain belongs to one of the hotspots of warming, with a temperature increase of more than  $\sim 4^{\circ}\text{C}$  (NASA Visualization Studio, 2021) since von Bunge's (1883) first steps to the today's accepted interpretation of how the ground ice in Yedoma was formed.

#### 4.4.2 Yedoma Existence Under Overlying Sediments

Another important point of discussion was mentioned by Grosse et al. (2013). They pointed to a need to carefully evaluate the Yedoma coverage from the fact that a two-dimensional surface source map, as true for nearly every classical map, cannot properly represent the total spatial extent of geologic horizons that potentially overlap in three dimensions. The Yedoma units represented in the maps (and thus the ones in our database) are likely those that crop out close to the land surface. Although thermokarst depressions and lakes indicate that the entire underlying Yedoma has thawed out in many cases, shallow lakes may not have fully degraded Yedoma below yet (Shmelev et al., 2017; Walter Anthony et al., 2021). Similarly, younger aeolian or alluvial deposits may cover late Pleistocene Yedoma surfaces. Therefore, it is still unclear whether there are any Yedoma remains beneath the lacustrine and alluvial sediments. Such geological situations would make the identified extent of Yedoma in our database a conservative estimate at the lower end.

#### 4.4.3 General Map Accuracy

As also summarized by Grosse et al. (2013), the general accuracy of the Yedoma distribution shown in this database is tied to several potential general error sources: 1) uncertainties in the original geologic and cryostratigraphic mapping, 2) uncertainties in converting and scaling the original field and remote sensing data to a paper map product, 3) uncertainties in the technical conversion from a paper map to a digital geodatabase, and 4) uncertainties in interpretation of Quaternary geological units across multiple map tiles as a Yedoma unit.

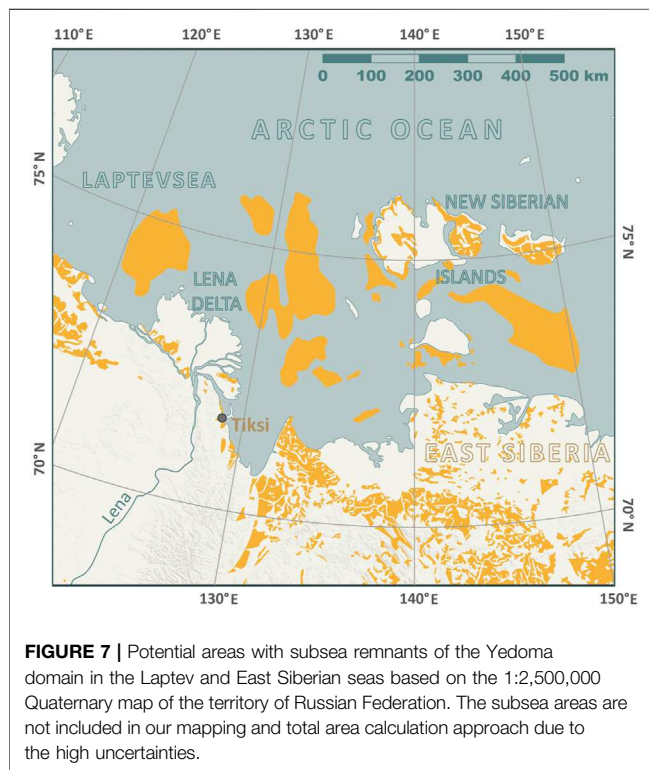
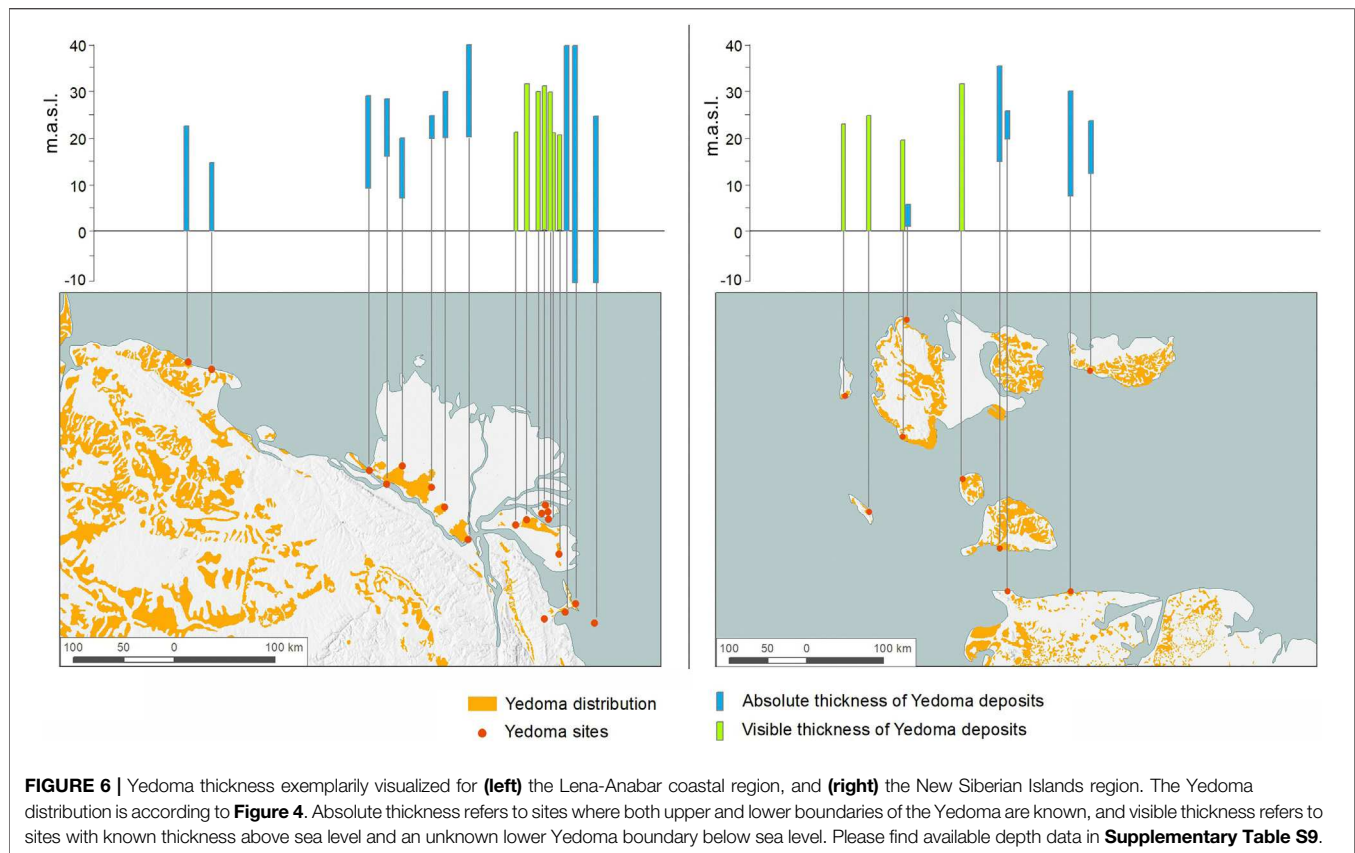
Our study cannot influence the uncertainties mentioned in (1) and (2) and assumes that quality control of the original geologic, stratigraphic, and cartographic work was rigorous. For (3), geometric uncertainties are mostly included, such as the accuracy of georeferencing the maps, and the line thickness of geologic boundaries and their digitization accuracy. To estimate transparently and to minimize errors in (4), we relied on existing scientific literature, our field data and field knowledge of Yedoma and the involvement of expert assessment. Therefore, we introduced the confidence classes, making use of the fact that most of the maps provide evidence that a certain unit is ice-rich, syngenetic, late Pleistocene Yedoma by cryostratigraphic signatures. In a few cases, no signature for high ground-ice content or syngenetic ice wedges was included. However, it was possible to trace such units by comparing them with neighboring polygons in directly adjacent map tiles, connecting them to field or literature data, or intersecting this with other maps and inferring that they are the same units if spatially connected.

#### 4.4.4 Yedoma and Yedoma-like Deposit Definitions

Another important topic to address here are the scientific differences and viewpoints on Yedoma, Ice Complex and muck definitions. This is quite clear for the Yedoma core region, but gets more difficult to the Yedoma domain borders, meaning the easternmost and westernmost sites (Figure 5). The key features of Canadian mucks (eastern edge of Beringia) are that they are primarily found in the extensive-discontinuous permafrost zone. In contrast to the East Siberian Yedoma core area, they are also strongly controlled by aspect (north-facing sites and narrow valleys with relict ice) and they also occur in high relief landscapes. Following our definition, the Holocene part like organics (sologenic) are not included here (Fraser and Burn, 1997; Froese et al., 2009). On the other side, the Western Taymyr to Yamal Peninsula coast is the westernmost area of distribution of Yedoma (Sopochynaya Karga, Marre-Sale, Dikson, Krestyanka River). In contrast to the Yakutian Yedoma, the Western Taymyr Yedoma is less thick (up to 12 m), and includes smaller syngenetic ice wedges (Streletskaia et al., 2007). Nevertheless, the grain-size fractions and cryogenic structure of the Yedoma deposits of Western Taymyr are similar to sections in Yakutia and Alaska (Tomirdiario and Chernen'kiy, 1987). Here, we see these differences as regional expressions of Yedoma, marking differences in source areas, sedimentation rates, moisture availability and climate during aggradation of these fine-grained ice-rich sediments across Beringia. Secondly, we feel there is more to gain by considering these as regional variants of processes that were active across the region during the Late Pleistocene.

#### 4.4.5 Yedoma Thickness

One important step forward in our dataset and map is for the first time the inclusion of Yedoma thickness, which is needed for volume estimation. However, published data on Yedoma thickness are rather scarce. Synthesized by Strauss et al. (2013; 2017), we compiled field and literature data for available Yedoma depth estimate and updated this with newly available or newly recognized depth estimates (see **Supplementary Table S9**). Moreover, as visualized in the example in Figure 6, it is difficult to quantify the depth of Yedoma deposits if the base is not reached by the cliff or drilling. Local heterogeneity and depths ranging from 5 to 50 m (mean 17 m, **Supplementary Table S9**) make it impossible at this current stage to visualize the thickness spatially explicit. There are first approaches to use the depth of thermokarst basins for Yedoma depth estimation such as the study by Veremeeva and Glushkova (2016). For the Kolyma Lowland tundra region, they showed that average depths of Yedoma-alas relief dissection by thermokarst and thermal erosion are 20–25 m between highest elevation of Yedoma uplands and alas levels. This is consistent with Yedoma thickness data of studied exposures and boreholes. Average heights of Yedoma upland and alas relief for low, moderate and high Yedoma fraction regions in Kolyma Lowland tundra correspond to different elevation levels but have similar relative elevation differences between average Yedoma and alas heights across all relief types, which are about 10 m (Veremeeva et al., 2021).



#### 4.4.6 Subsea Yedoma Deposits

Another unsolved question is the quantification of potentially preserved Yedoma under today's sea coverage. The area of this potential subsea Yedoma is highly uncertain: there are only several maps estimating potential areas in the Laptev Sea (1:1,000,000 Quaternary geological map S-52 and S-53,54, and the 1:2,500,000 Quaternary map of the territory of Russian Federation), visualized in Astakhov et al. (2021) and the western part of the East Siberian Sea as indicated in Walter et al. (2007). Although we did not include the potential subsea regions in our map or any submarine Yedoma remnant area estimate in our calculations, we here show an example of potential regions with submarine Yedoma remnants (**Figure 7**). The outline of the areas shown here enclose 77,620 km<sup>2</sup>.

## 5 OUTLOOK

Our Yedoma mapping includes a large and diverse amount of available geospatial data from which the Yedoma extent could be extracted. Additional datasets and maps are likely to become available in the future. In particular, Soviet and Russian geological mapping efforts over many decades have been included in geological maps of 1:200,000 scale, however, many of these are not yet published or available for the Northeast Siberian Yedoma



regions. A substantial treasure of data on Yedoma distribution and thickness is likely buried in largely inaccessible data from prospecting drilling campaigns focused on engineering, infrastructure planning, exploration, or hydrogeological studies in Siberia as well as in North America. Making such datasets accessible and digitally available would provide a tremendous boost to understanding the distribution of ice-rich syngenetic permafrost at much higher levels of spatial detail. Additional scientific studies on Yedoma exposures and drill cores will further enhance knowledge on Yedoma distribution and thickness and can be integrated in future updates of this new pan-Arctic map and database. A highly promising approach will be the application of geomorphic mapping of Yedoma coverage based on terrain texture and thermokarst and thermal erosion features associated with Yedoma deposits using high-resolution satellite images and digital elevation models. Such remote sensing datasets also allow enhanced quantification of the spatial heterogeneity of remnant Yedoma uplands well below the resolution of currently used mapping scales in this study.

## 6 CONCLUSION

The Yedoma geospatial database and map strongly reduces uncertainties in Yedoma distribution. We quantified the Yedoma domain for a land area of 2,587,000 km<sup>2</sup> in Central and East Siberia, Alaska and Yukon with areas of digitized Yedoma deposits of 480,000 km<sup>2</sup>. Today's Yedoma areas are highly fragmented, which mainly resulted from permafrost degradation processes during the Lateglacial and Holocene such as thermokarst, thermal erosion, and fluvial and coastal erosion. This first complete, digital circum-Arctic Yedoma map is an important step forward to understand the past and present spatial heterogeneity of Yedoma deposits, which is of utmost importance for assessing vulnerabilities and risks in a rapidly warming Arctic.

## DATA AVAILABILITY STATEMENT

The datasets presented in this study can be found in online repositories. The names of the repository/repositories and accession number(s) can be found in the article/**Supplementary Material**.

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## AUTHOR CONTRIBUTIONS

JS and GG designed this study. SL, GG and USGS volunteers conducted all digitizing, data archiving and visualizations. JS, GG, and LS drafted the first versions of the manuscript. All authors contributed data to this study and contributed to manuscript finalization.

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## SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/feart.2021.758360/full#supplementary-material>

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