

Hygroscopicity of Nitrogen Containing Organic Carbon Compounds: *o*-aminophenol and *p*-aminophenol

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Abstract

Nitrogen-containing Organic Carbon (NOC) is a major constituent of atmospheric aerosols and they have received significant attention in the atmospheric science community. While extensive research and advancements have been made regarding their emission sources, concentrations, and their secondary formation in the atmosphere, little is known about their water uptake efficiencies and their subsequent role on climate, air quality, and visibility. In this study, we investigated the water uptake of two sparingly soluble aromatic NOC: *o*-aminophenol (*oAP*) and *p*-aminophenol (*pAP*) under subsaturated and supersaturated environments using Hygroscopicity Tandem Differential Mobility Analyzer (H-TDMA) and a Cloud Condensation Nuclei Counter (CCNC), respectively. Our results show that *oAP* and *pAP* are slightly hygroscopic with comparable hygroscopicities to various studied organic aerosols. The supersaturated single hygroscopicity parameter (k_{CCN}) was measured and reported to be 0.18 ± 0.05 for *oAP* and 0.04 ± 0.02 for *pAP*, indicating that *oAP* is more hygroscopic than *pAP* despite them having the same molecular formulae. The observed disparity in hygroscopicity is attributed to the difference in functional group locations, interactions with gas phase water molecules, and the reported bulk water solubilities of the NOC. Under subsaturated condition, both *oAP* and *pAP* aerosols showed size dependent water uptake. Both species demonstrated growth at smaller dry particle sizes, and shrinkage at larger dry particle sizes. The measured Growth Factor (G_f) range, at RH = 85%, for *oAP* was 1.60-0.74 and for *pAP* was 1.53-0.74 with increasing particle size. The growth and shrinkage dichotomy is attributed to morphological particle differences verified by TEM images of small and large particles. Subsequently, aerosol physicochemical properties must be considered to properly predict the droplet growth of NOC aerosols in the atmosphere.

Introduction

Aerosol particles have major implications on global climate, air quality, and human health. Aerosols can be emitted from a variety of sources (e.g., biomass burning, sea spray, combustion processes, desert dust, and volcanic eruptions)^{1–3} or they can be generated in the atmosphere through various chemical processes (e.g., nucleation, condensation, and oxidation).^{4–6} Once suspended in the atmosphere, aerosols can influence Earth's radiative budget through their direct and indirect interaction with solar radiation.^{2,4,7–9} The direct effect is based on the aerosol's ability to directly scatter and absorb solar radiation. As for the indirect effect, aerosol particles can act either as cloud condensation nuclei (CCN)—taking up water vapor and forming clouds in the process^{4,7,10} or as ice nucleating particles to form ice clouds.¹¹ Hence, aerosols' CCN efficiencies to uptake water (hygroscopicity) can directly affect the amount of cloud's light absorption or scattering that takes place. The latest Intergovernmental Panel on Climate Change (IPCC) estimate of climate radiative forcing from the aerosol/cloud interaction is -0.7 Wm^{-2} (uncertainty range -1.8 to -0.3 Wm^{-2}).¹² The large uncertainty associated with aerosol-cloud interactions highlights the complexity associated with aerosols and their effect on cloud formation. While aerosol particles are composed of organic and inorganic species, the primary source of uncertainty stems from abundant and poorly understood organic aerosols.^{13,14}

Organic aerosols (OA) are ubiquitous in the atmosphere and account for a large fraction of atmospheric aerosols.^{13,15} Within the last decade, the atmospheric chemistry field has made substantial advances in aerosol speciation, chemical compositions of OA, and the chemical processes that OA undergo in the atmosphere. These advances were

highlighted in published reviews that include but not limited to Erven et al.⁶; Hallquist et al.¹⁶; Kanakidou et al.¹³; Kroll and Seinfeld¹⁷; and Mahilang et al.¹⁸ Despite these advances, the uncertainty in the hygroscopic-related properties of atmospheric OA remains large.¹⁵

Nitrogen-containing organic carbon (NOC) are a major class of OA growing in importance due to a decline in regional sulfates and their prominent presence within atmospheric aerosols.^{15,19–23} Wedyan *et al.* (2007)²¹ discovered that 80% of nitrogen over the coastal Gulf of Aqaba was attributed to NOC.²¹ Another study by Cornell *et al.*²⁴ investigated the atmospheric chemical composition in Oahu, Hawaii, and found that approximately 30% of the total nitrogen was stored in NOC form.²⁴ The abundance of NOC in the atmosphere is attributed to the wide range of primary and secondary emission sources.²⁵ Primary sources such as biomass burning and both anthropogenic and biogenic sources, are responsible for approximately 27.4 Tg of NOC annually.¹³ As for secondary sources, the formation of NOC within the atmosphere remains unexplored. Various studies have shown that NOC can be formed in the atmosphere from chemical processes involving secondary organic aerosol (SOA) (e.g., but not limited to Darer et al.²⁶; Liu et al.²⁷; Vidović et al.²⁸; Zhang et al.²⁹). For instance, a wide range of NOC species can be formed from the ozonolysis of α -pinene and *m*-xylene in the presence of ammonia.²⁷ Another study found that the atmospheric presence of NOC is attributed to the reaction of ammonium with oxidized organics (e.g., formate, acetate, pyruvate, malonate).²⁹

To further enhance our knowledge of NOC and their role within atmospheric chemistry and climate impacts, investigating the physiochemical properties of NOC

aerosols is necessary.⁷ One vital physicochemical property is hygroscopicity, which describes the water uptake ability of aerosols under subsaturated and supersaturated conditions.³⁰ While extensive efforts have been allocated to investigate and report the water uptake behavior of atmospherically relevant OA compounds, further research remains needed. This is due to the presence of a wide range of OA in the atmosphere making quantifying their hygroscopic properties a formidable task. To circumvent this challenge, theoretical models have been used to predict the hygroscopicity of OA, which are semi-empirical and often rely on parametrization of water solubility and extrapolation from lab measurements of one class of OA to another.³¹ Such simplistic approach fails to accurately capture the non-ideal behavior that is developed from solubility,³² liquid-liquid phase separation,^{33,34} surface partitioning,³⁵ and morphology.³⁶ Also, these limitations highlight the fundamental importance to investigate the hygroscopicity of different classes of OA through laboratory measurement techniques. As such, various studies have undertaken a systematic approach by investigating the hygroscopicity of OA through the lens of other physicochemical properties such as functional groups, molecular weight, and bulk water solubility.^{37–42} For example, Suda *et al.*⁴² investigated the CCN activity of OA possessing different functional groups and reported a direct influence of functional group on water uptake. Specifically, it was reported that the hygroscopicity of an organic compound with hydroxyl or carboxyl groups is higher than that with nitrate or methylene groups.⁴² Other studies investigated a family of organic compounds, such as carboxylic acids^{37,43} or amino acids.^{37,44} The water uptake of amino acids has been reported due to their important involvement in the nitrogen cycle in the atmosphere (e.g., but not limited to Chan *et al.*⁴⁵; Han *et al.*⁴³; Luo *et al.*⁴⁴; Marsh *et al.*³⁷). Luo *et al.*⁴⁴

investigated the water uptake using a hygroscopicity tandem differential mobility analyzer (H-TDMA) system, under a range of relative humidity (RH), of a series of amino acids, including aspartic acid, glutamine, and serine along with their mixtures with ammonium sulfate. The results of Luo *et al.*⁴⁴ show that bulk water solubility play a major role in understanding the hygroscopicity of these amino acids, and that the presence of ammonium sulfate influence the phase state and hygroscopicity of the mixture.⁴⁴

To contribute to the mechanistic understanding of OA hygroscopicity, we explored two NOC aminophenol isomers: *o*-aminophenol (*oAP*) and *p*-aminophenol (*pAP*).^{46,47} To our knowledge, no previous literature has reported on the hygroscopicity of pure aromatic compounds that contain nitrogen. Previous studies have employed other aromatic compounds to produce secondary organic aerosols (SOA) of mixed chemical composition. For example, the study by *Hilderbrandt Ruiz et al.*⁴⁸ involved measuring the water uptake of SOA produced from the photooxidation of toluene. Another study by *Nakao et al.*⁴⁹ investigated the SOA formation from benzene, toluene, *m*-xylene using an environmental chamber under different conditions. Similarly, various studies have investigated the water uptake of SOA formed from monoterpenes (C₁₀H₁₆) such as α -pinene, β -pinene, limonene, and terpinene (Huff Hartz *et al.*⁵⁰; Prenni *et al.*³⁹; Duplissy *et al.*⁵¹; Frosch *et al.*⁵², to name a few). Additionally, while previous studies have reported the hygroscopicity of SOA from aliphatic amines,⁵³ no previous study has explored aromatic amines, whether in single- or multi-component systems.

Exploring single component composition instead of a mixture of components offer valuable knowledge with respect to how various physicochemical properties such as solubility, O:C ratio, functional groups, chemical composition, and structure impact water

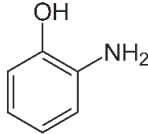
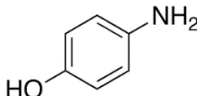
uptake in aerosol particles. This is highlighted by the previously published CCN work that have investigated single component aerosols (e.g., but not limited to Marsh et al.³⁷; Peng et al.⁵⁴; Raymond et al.⁵⁵; Han et al.⁴³; Dawson et al.⁵⁶; Estillore et al.⁵⁷). Therefore, our study aims to enhance our understanding of the physiochemical properties effecting water uptake. The ring structure of *o*AP and *p*AP will provide us with a new perspective on the effect of aromaticity. Additionally, the two functional groups (-NH₂ and -OH) on the aromatic ring in *o*AP and *p*AP add to the growing database of hygroscopicity of OA with different functional groups. Previous studies have invested in understanding the hygroscopic properties of OA with different functional groups (Han et al.⁴³; Petters et al.⁵⁸; Suda et al.⁴², to name a few); however, no previous reports were conducted on compounds containing -NH₂ and -OH together. Finally, *o*AP and *p*AP enable us to explore the effect of solubility on hygroscopicity while maintaining the same O:C ratio, chemical composition, and functional groups.

The objective of this paper is to investigate the water uptake of *o*AP and *p*AP under subsaturated condition using a Hygroscopicity Tandem Differential Mobility Analyzer (H-TDMA) and supersaturated condition using a Cloud Condensation Nuclei Counter (CCNC). Results were analyzed through the lens of solubility and morphological properties and highlight the importance of physicochemical properties on the water uptake of NOC compounds.. Hence, understanding the water uptake of these aminophenol compounds will add a new dimension to the growing knowledge on the impact of functional groups on hygroscopicity.

Experimental and Data Analysis

Chemicals. *o*-aminophenol (99%, CAS: 95-55-6, Sigma-Aldrich) and *p*-aminophenol ($\geq 98\%$, CAS: 123-30-8, Sigma-Aldrich) were used in this study. All chemicals were used as received without further purification. The chemical structures and physical properties of both chemicals are summarized in **Table 1**.

Table 1. The chemical structures and the relevant properties of *o*-aminophenol and *p*-aminophenol

Chemical	Abbv.	Structure	Molecular Weight [g mol ⁻¹]	Density [g mL ⁻¹]	Bulk Solubility, S _{bulk} (in H ₂ O) [g L ⁻¹]
<i>o</i> -aminophenol	<i>oAP</i>		109.13 ^a	1.33 ^a	19.6 ^c
<i>p</i> -aminophenol	<i>pAP</i>		109.13 ^b	1.13 ^b	15.7 ^c

^a National Center for Biotechnology Information (2022). PubChem Compound Summary for CID 5801, 2-Aminophenol. Retrieved April 3, 2022 from <https://pubchem.ncbi.nlm.nih.gov/compound/2-Aminophenol>.

^b National Center for Biotechnology Information (2022). PubChem Compound Summary for CID 403, 4-Aminophenol. Retrieved April 3, 2022 from <https://pubchem.ncbi.nlm.nih.gov/compound/4-Aminophenol>

^c Rumble, J. R.; Lide, D. R.; Bruno, T. J. CRC handbook of chemistry and physics: a ready-reference book of chemical and physical data, 2017

Aerosol Generation. A 0.1 g L⁻¹ solution was prepared for each chemical using ultra-purified water (18 MΩ cm⁻¹, Millipore). The prepared solutions were sonicated to ensure dissolution prior to atomization. The aerosols were then generated using a constant output Collison Nebulizer (Atomizer; TSI 3076). The aerosols generation method has been used and described in previous CCN literature.^{43,55,57,59,60}

The wet aerosols were dried (< 5% RH) using silica gel dryers. The water uptake of the generated dry aerosols was then measured with two different experimental setups; a) H-

TDMA (subsaturated conditions) and b) CCNC (supersaturated conditions) (**Figure 1**). In addition, TEM images were obtained using a sample collection setup shown in **Figure 1**.

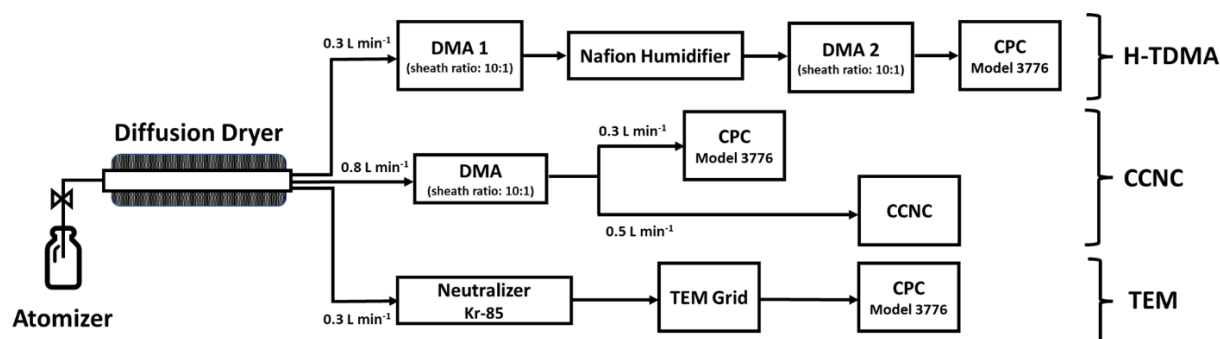


Figure 1. The experimental setup for subsaturated H-TDMA measurements, supersaturated CCNC measurements, and TEM sample collection. Abbreviations are: DMA: Differential Mobility Analyzer, CPC: Condensation Particle Counter, CCNC: Cloud Condensation Nuclei Counter, and TEM: Transmission Electron Microscope.

Hygroscopicity Tandem Differential Mobility Analysis. The water uptake of *oAP* and *pAP* was evaluated under subsaturated conditions using the widely employed H-TDMA. The operation of the H-TDMA has been described in detail in previous literature.^{61,62} An overview of the H-TDMA experimental setup is shown in **Figure 1**. The dry aerosols, generated from a 0.1 g L^{-1} solution, were size selected (30-200 nm) by an electrostatic classifier (DMA 1, TSI 3082; flow rate = 0.3 L min^{-1}). The size-selected particles were then humidified to $\text{RH} = 85\% \pm 5\%$ using a nafion humidification line (PermaPure M.H. series). The size distribution of the wet aerosol particles was then scanned using a second electrostatic classifier (DMA 2, TSI 3085; flowrate 0.3 L min^{-1}), and the concentrations were measured using a Condensation Particle Counter (CPC; TSI 3776; flow rate = 0.3

L min⁻¹). The Growth Factor (G_f) is the ratio between the wet mobility diameter (D_{wet}) and the dry mobility diameter (D_{dry}) (**Eq. 1**). G_f was calculated for each size-selected dry mobility diameter, averaged, and reported with its respective standard deviation.

$$G_f = \frac{D_{wet}}{D_{dry}} \quad (1)$$

The RH of the H-TDMA setup was calibrated with inorganic ammonium sulfate (see *Supplemental Information*; Taylor *et al.*⁶³).

Cloud Condensation Nuclei Analysis. The water uptake of *oAP* and *pAP* was tested under four different supersaturations (0.46, 0.67, 0.86, and 1.08%), that represent warm clouds, using a commercially available Cloud Condensation Nuclei Counter (CCNC; Droplet Measurement Technologies). The operation of the CCNC has been extensively described in previous literature.^{64–66} An overview of the CCNC experimental setup is shown in **Figure 1**. A flow stream (flow rate = 0.8 L min⁻¹) of dried aerosol particles were charged in a Kr-85 bipolar neutralizer to produce an equilibrium distribution of electrically charged particles.⁶⁷ The charged particles were then passed to a DMA (TSI 3080) in scanning mode and a size distribution from 8 to 352 nm was measured every 2.25 minutes. Downstream from the DMA, the aerosol stream was bifurcated with one stream (flow rate = 0.3 L min⁻¹) flowing into a CPC (TSI 3776) where the particle concentration (CN) was measured, while the other stream (flow rate = 0.5 L min⁻¹) flowing into the CCNC where the activated particle concentration (CCN) was measured. The activated fraction is the ratio between CCN and CN and was measured at four different supersaturations, and the critical diameter (D_d), defined as the dry particle diameter size at which 50% of

the particles form droplets, was calculated with Scanning Mobility CCN Analysis (SMCA).⁶⁶ The SMCA applies a charge correction followed by fitting a sigmoid curve through the activated fraction to obtain the D_d .^{66,67} At each supersaturation, 10 critical diameters were obtained for each chemical, and were used to calculate the hygroscopicity. To determine the CCNC's supersaturations, a widely used calibration procedure was conducted with inorganic ammonium sulfate.^{55,56,68–71} Ammonium sulfate has been used in various literature for CCNC calibration for hygroscopicity measurements of organic aerosol.^{55,56,69–71} The calibration was conducted prior to the experiments, and the calibrated supersaturations are reported in the *Supplemental Information*.

Water vapor condensation on aerosol particles and droplet growth can be thermodynamically described by Köhler theory. The theory is based on equilibrium thermodynamics and describes the vapor pressure and saturation, S , as follows:⁷²

$$S = a_w \exp \left(\frac{4\sigma_{s/a} M_w}{RT \rho_w D_{wet}} \right) \quad (2)$$

where a_w represents the water activity coefficient, $\sigma_{s/a}$ is the surface tension of the droplet, R is the universal gas constant, T is the sample temperature, D_{wet} is diameter of the wet droplet, and ρ_w and M_w are the density and molecular weight of water, respectively.⁷²

Based on the Köhler theory, the hygroscopicity of aerosol particles can be expressed in terms of a single hygroscopicity parameter, κ , assuming the solute to be sufficiently dilute (i.e., ideal solubility).³⁰ Hence, supersaturated κ -hygroscopicity (κ_{CCN}) of oAP and pAP can be expressed using the following equation:³⁰

$$\kappa_{CCN} = \frac{4 \left(\frac{4\sigma_s/a M_w}{RT\rho_w} \right)^3}{27D_d^3 \ln^2 s_c} \quad (3)$$

where D_d is the critical activation diameter and s_c is the CCNC instrument supersaturation. The κ_{CCN} derived from the equation above is assumed to be constant as s_c is directly proportional to $D_d^{-1.5}$.³⁰

Transmission Electron Microscopy (TEM). The morphology of nano-sized aerosol particles of *oAP* and *pAP* were measured using a Transmission Electron Microscopy (JEOL 2100 TEM; LaB₆ filament)—a method previously described in literature.^{34,73} An overview of the TEM sample collection is presented in **Figure 1**. In brief, the generated dry aerosol stream (0.3 L min⁻¹) was charged with a neutralizer (Kr-85, TSI 3077A). The charged particles were then deposited for 4 hours onto an electrically grounded lacey carbon-coated copper TEM grid (TED PELLA). The deposited particles were then imaged at an accelerating voltage of 200 kV and a magnification range of 50-150 k. To minimize sample damage, the exposure time was limited to 90 sec.

Results and Discussion

Hygroscopicity under supersaturated conditions

The κ -values for *oAP* and *pAP* are reported and compared to other atmospherically relevant pure organic compounds (obtained from literature; refer to Table 2).. To understand the hygroscopicity of *oAP* and *pAP* aerosols under supersaturated conditions, 10 critical diameters (D_d) were measured at four different supersaturations (0.46, 0.67, 0.86, and 1.08%) and the average κ_{CCN} for each compound was calculated according to **Eq. 3**. Both *oAP* and *pAP* were observed to be slightly hygroscopic, under supersaturated

conditions, with κ -values between 0.04-0.18. The average κ_{CCN} for *oAP* and *pAP* aerosols were 0.18 ± 0.05 and 0.04 ± 0.02 , respectively. These values are consistent with κ -values, obtained from various studied organic compounds, of less than 0.4,^{37,43,74} specifically, κ -values of different carboxylic acids, amino acids, sugars, and cycloalkenes as listed in **Table 2** and **Figure 2a**. Hence, it is evident that the water uptake of both *oAP* and *pAP* falls within the range of hygroscopicity of other atmospherically significant functional groups suggesting that the presence of both functional groups ($-\text{OH}$ and $-\text{NH}_2$) in aminophenols results in an average water uptake behavior similar to that of other functional groups.

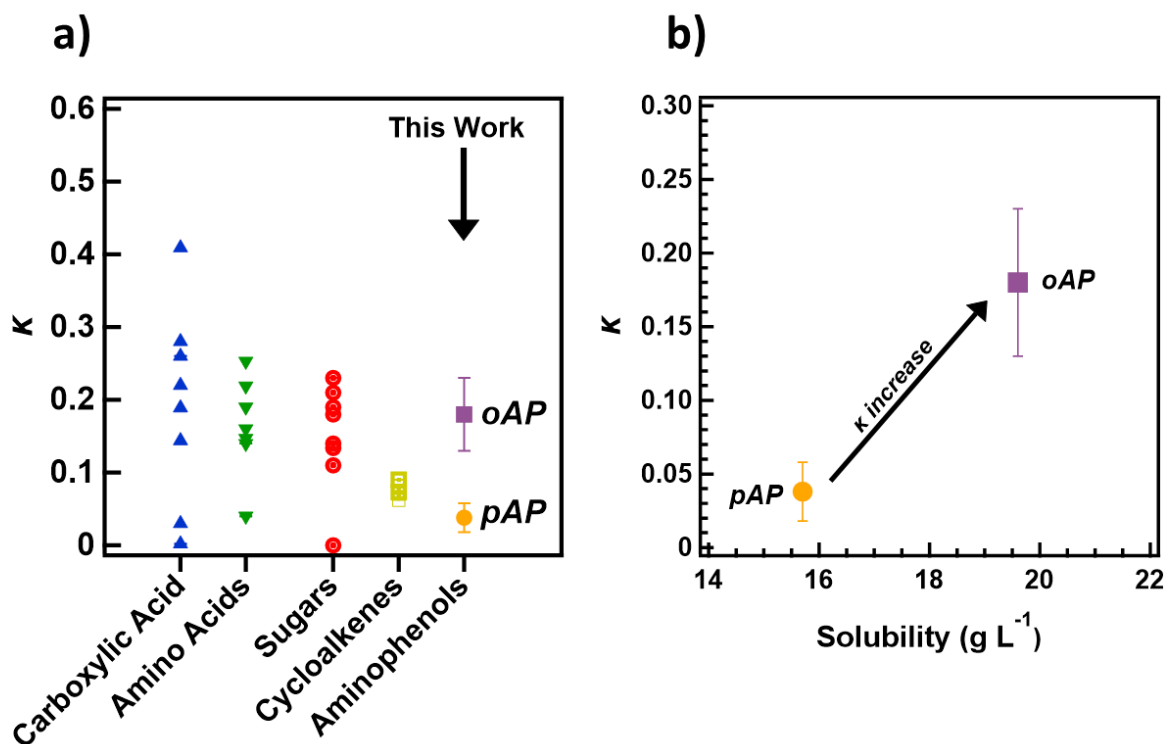


Figure 2. a) Hygroscopicity of organic aerosol reported in literature and compared to the aminophenols studied here, *oAP* and *pAP*. A detailed summary of the chemicals, κ -

values, and their respective references is found in **Table 2**. b) The relationship between hygroscopicity and bulk water solubility for *oAP* and *pAP* at 25 °C. Error bars for *oAP* and *pAP* represent standard deviation of 40 repeated measurements.

Table 2. Summary of literature κ -hygroscopicity of various organic compounds of atmospheric relevance

Organic Groups	Compounds	κ	References
Carboxylic Acid	Malonic Acid	0.26	<i>Han et al.</i> ⁴³
	Succinic Acid	0.002	<i>Peng et al.</i> ⁵⁴
	Pimelic Acid	0.03	<i>Han et al.</i> ⁴³
	Maleic Acid	0.28	<i>Choi and Chan</i> ⁷⁵
	Tartaric Acid	0.22	<i>Han et al.</i> ⁴³
	Citric Acid	0.19	<i>Marsh et al.</i> ³⁷
	Glutaric Acid	0.14	<i>Marsh et al.</i> ³⁷
	Oxalic Acid	0.41	<i>Marsh et al.</i> ³⁷
Amino Acids	Glycine	0.04	<i>Han et al.</i> ⁴³
	L-Aspartic	0.14	<i>Han et al.</i> ⁴³
	L-Glutamine	0.16	<i>Han et al.</i> ⁴³
	Serine	0.19	<i>Han et al.</i> ⁴³
	L-Arginine	0.147	<i>Marsh et al.</i> ³⁷
	L-Lysine	0.219	<i>Marsh et al.</i> ³⁷
	L-Valine	0.253	<i>Marsh et al.</i> ³⁷
Sugars	Fructose	0.23	<i>Han et al.</i> ⁴³
	Sucrose	0.10	<i>Dawson et al.</i> ⁵⁶
	Mannose	0.14	<i>Han et al.</i> ⁴³
	L-Arabitol	0.19	<i>Han et al.</i> ⁴³
	D-Mannitol	-0.01	<i>Han et al.</i> ⁴³
	Galactose	0.134	<i>Marsh et al.</i> ⁴³
	Xylitol	0.18	<i>Bilde et al.</i> ⁷⁶
Cycloalkenes	Cyclopentene	0.09 ^a	<i>Varutbangkul et al.</i> ⁷⁷
	Cyclohexene	0.073 ^a	<i>Varutbangkul et al.</i> ⁷⁷
	Cycloheptene	0.071 ^a	<i>Varutbangkul et al.</i> ⁷⁷
	Cyclooctene	0.062 ^a	<i>Varutbangkul et al.</i> ⁷⁷

^a κ -value was calculated⁷⁸ from G_f reported by *Varutbangkul et al.*⁷⁷

Previous studies by Bilde et al.⁷⁹; Chen et al.⁸⁰; Han et al.⁴³; Suda et al.^{42,81} to name a few reported variation in OA hygroscopicities with organic functionality. For example, Han et al.⁴³ showed a quantitative increase in κ -values based on functional groups in the order of (–COOH or C=O) > (–OH) > (–CH₃ or –NH₂).⁴³ Another study by Suda et al.⁸¹ examined the hygroscopicity of various SOA under subsaturated environments, and reported a relationship between κ -hygroscopicity and polarity in organic compounds, with highly polar SOA being more hygroscopic.⁸¹ The conclusions of these studies suggest that other physicochemical properties contribute, along with the type of functional groups, to the hygroscopicity of organic compounds.

To highlight the significance of other physicochemical properties (specifically solubility), **Figure 2b** shows a significant difference in κ -values between *oAP* and *pAP*. Despite having the same molecular formulae and type of functional groups, *oAP* ($\kappa_{CCN} = 0.18 \pm 0.05$) is more hygroscopic than *pAP* ($\kappa_{CCN} = 0.04 \pm 0.02$). This difference in κ_{CCN} (*i.e.*, hygroscopicity) is attributed to the difference in bulk water solubility between *oAP* and *pAP*. Previous studies have concluded that the bulk water solubility range of 0.1–100 g L^{–1} determines the CCN activity of organic compounds.^{41,82–84} The bulk water solubility of both *oAP* and *pAP* falls in the range of slightly soluble compounds with values of 19.6 g L^{–1} and 15.7 g L^{–1}, respectively. Under subsaturated hygroscopic conditions, compounds with limited solubility are composed of an insoluble core surrounded by a saturated solution.⁴³ Hence, *oAP* having the higher solubility, has a greater molar concentration in the saturated solution, leading to *oAP* being more hygroscopic than *pAP*. This increase in hygroscopicity based on water solubility has been observed in various OA such as, glyoxylic acid and 4-methylphthalic acid in Chan et al.⁴⁰; maleic acid and malonic acid in

Han et al.⁴³; aspartic acid and azelaic acid in Huff-Hartz et al.⁵⁰; and polycatechol and polyguaiacol in Malek et al.⁸⁵, to name a few.

Hygroscopicity under subsaturated conditions

The water uptake of *oAP* and *pAP* under subsaturated conditions was measured and analyzed based on morphological properties. Subsaturated measurements were completed under constant RH ($85 \pm 5\%$), and the measured wet mobility diameters and growth factors are shown in **Figure 3**. For both *oAP* and *pAP* the water uptake was size-dependent highlighting the presence of two distinct regions referred to as Growth and Shrinkage in **Figure 3**. Here, the term 'Growth' refers to G_f values greater than 1, while the term 'Shrinkage' refers to G_f values less than 1. The experimental G_f values for *oAP* and *pAP* inversely correlated with the dry mobility diameter. For *oAP*, the G_f values range between 1.6-0.74 with increasing dry mobility diameter. For dry mobility diameters ≤ 120 nm, *oAP* aerosols were shown to take up water and grow when exposed to humidification. This water uptake behavior is highlighted by the G_f values greater than 1 for dry mobility diameter ≤ 120 nm. *oAP* aerosols shrank for dry mobility diameter ≥ 120 nm where G_f values are less than 1 (**Figure 3a**). Similarly, the G_f values for the *pAP* system follow the same trend as the *oAP* system. The *pAP* G_f values decrease with increasing dry mobility diameter, with G_f values ranging between 1.5-0.74. However, and in contrast to *oAP* aerosols, the *pAP* aerosol shrinks at a smaller dry mobility diameter ≥ 60 nm (**Figure 3b**).

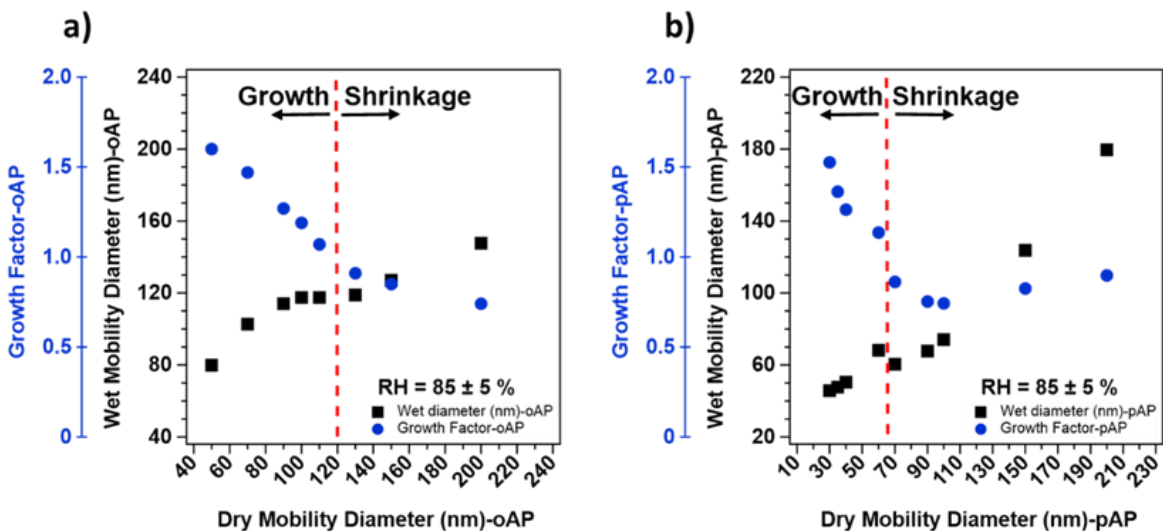


Figure 3. H-TDMA measurements and analysis showing the growth factors versus dry mobility diameters (blue circles) and wet mobility diameters versus dry mobility diameters (black squares) for a) *oAP* and b) *pAP*. Red vertical line indicating the dry mobility diameter where growth to shrinkage transition takes place. Error bars (standard deviation) are included, but small to be visible.

To our knowledge no previous literature has detected a trend similar to our study, where a single component aerosol system undergoes droplet growth and shrinkage across a range of ultrafine dry mobility diameters. Several studies by Cheng *et al.*⁸⁶; Laskina *et al.*⁸⁷; Mahish and Collins⁸⁸ (to name a few) have reported size dependence in the uptake of water of various atmospheric aerosols. However, such observations were made on multicomponent aerosol systems. Laskina *et al.*⁸⁷ reported a difference in the water uptake based on particle sizes obtained from a mixture of ammonium sulfate and different OA. This difference was attributed to changes in the mixing state observed for different sized particles.⁸⁷ A study by Han *et al.*⁴³ reported G_f values less than 1 (200 nm; 90% RH) for azelaic and suberic acid using an H-TDMA setup.⁴³ The shrinking behavior

was attributed to morphological limitations of the aerosol particles.⁴³ Specifically, water adsorption at the surface of the particle could potentially lead to structural rearrangement at the microscale causing particles to collapse and shrink.^{36,89,90}

Figure 4 shows representative TEM images for *oAP* and *pAP* in two particle size ranges. The particle morphology is dependent on particle sizes. At small sizes (≤ 50 nm), the shape of the particles appears spherical for both *oAP* and *pAP* aerosols. These small particles are compact, and hence, when water condenses on their surface, they grow in size. Thus, droplet growth and positive G_f are observed for both *oAP* and *pAP* at small dry mobility diameters.

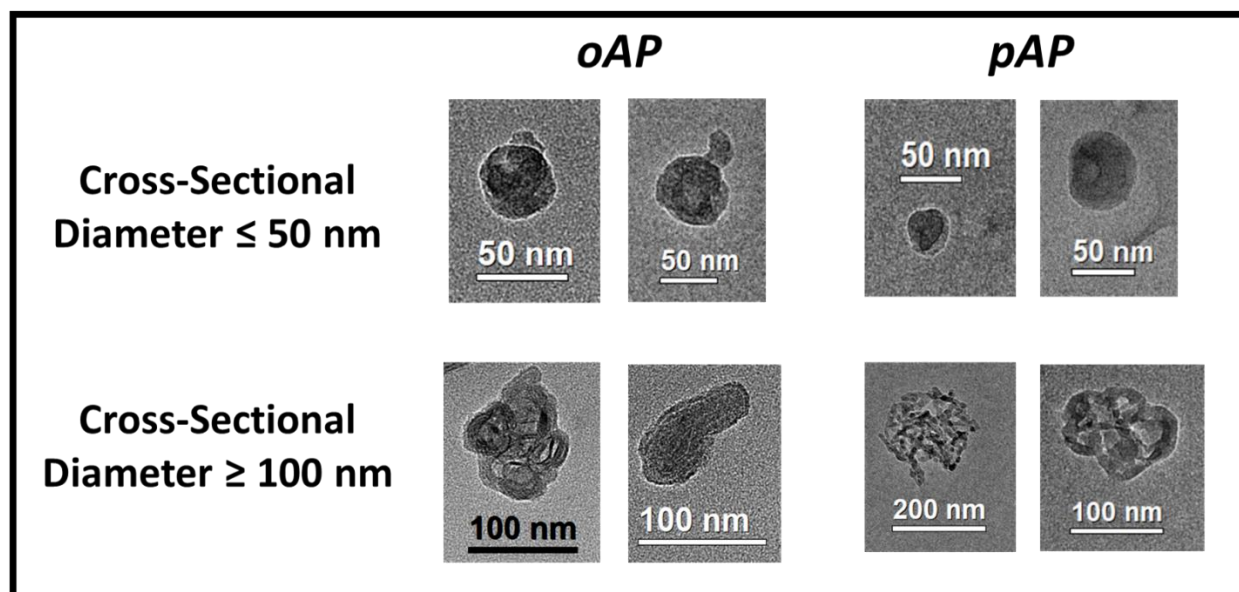


Figure 4. Representative TEM images showing morphological difference between small and large cross-sectional diameter size particles for *oAP* and *pAP*. Sizes within each image represents scale bar. Additional images are provided in **Figures S5 and S6, S7, and S8** in the *Supporting Information*.

Conversely, large size particles (≥ 100 nm) appear to contain void spaces. They can be formed of smaller elongated subunits. Such morphology at large sizes, makes these subunit particles more susceptible to collapse and rearrangement, as previously postulated by *Mikhailov et al.*⁹⁰ TEM images support the observed shrinking in large *oAP* and *pAP* particles, where $G_f = 0.74$, and the collapsed particle is reduced in diameter by ~25%. The size-dependent morphology is consistent with the presence of multiple nucleation sites on the surface of aqueous droplets: as the solvent evaporates, it may leave behind structures with void spaces^{91,92}, which are more prevalent in the *pAP* dry particle size distribution.

The difference between the water uptake of *oAP* particles at larger sizes (< 120 nm) than *pAP* (< 60 nm) can be attributed to chemical bonding. The functional groups ($-OH$ and $-NH_2$) in *oAP* can engage in intramolecular hydrogen bonding due to their *ortho*-positioning on the benzene ring. This reduces the number of hydrogen bonds that they can engage in with other neighboring molecules. Hence, *oAP* is less likely to form structures consisting of sub-units compared to *pAP*. As for *pAP*, due to the functional groups being in *para*-position of the benzene ring, hydrogen bonding with other molecules is more likely to occur. In addition, due to the symmetry of the *para*-structure, *pAP* molecules are prone to having a stronger π - π stacking than *oAP*. This is due to the involvement of the electrons, from $-OH$ and $-NH_2$, in resonance structures causing a high electron density in the benzene ring. Hence, *pAP* particles are more often to consist of sub-units and, as result, more likely to collapse at smaller sizes than *oAP*.

Moreover, the data in **Figure 3** can be used to theoretically estimate the concentration of the aminophenol particles in the droplet at the nanoscale (C_{nano}) from the

value of the wet and dry mobility diameters at $G_f = 1$ assuming spherical particles and using a density of water of 1 g cm^{-3} for comparison with solubility in **Table 1**. **Table 3** lists the parameters used in calculating C_{nano} and the corresponding solute:solvent mass ratio for the particle and bulk systems. This calculation assumes that the mass of *oAP* and *pAP* are fully dissolved in the wet volume. Hence, the C_{nano} calculation can be considered the upper limit of the concentration of a deliquesced aerosol solute.

Table 3: Estimation of particle solubility from hygroscopic growth data in **Figure 3**

o-aminophenol (oAP)							
Wet Diameter	Wet Volume	Dry Diameter	Dry Volume	Mass of oAP	Vol. of H ₂ O	C_{nano}	Solute:solvent mass ratio ^e
$D_{\text{wet}} \text{ (nm)}$	$V_w \text{ (cm}^3\text{)}^a$	$D_{\text{dry}} \text{ (nm)}$	$V_d \text{ (cm}^3\text{)}^a$	$m \text{ (g)}^b$	$V_{\text{H}_2\text{O}} \text{ (cm}^3\text{)}^c$	$\text{(g L}^{-1}\text{)}^d$	
140	1.4×10^{-15}	120	9×10^{-16}	1.2×10^{-15}	5.3×10^{-16}	2264	2.3

p-aminophenol (pAP)							
Wet Diameter	Wet Volume	Dry Diameter	Dry Volume	Mass of pAP	Vol. of H ₂ O	C_{nano}	Solute:solvent mass ratio ^e
$D_{\text{wet}} \text{ (nm)}$	$V_w \text{ (cm}^3\text{)}^a$	$D_{\text{dry}} \text{ (nm)}$	$V_d \text{ (cm}^3\text{)}^a$	$m \text{ (g)}^b$	$V_{\text{H}_2\text{O}} \text{ (cm}^3\text{)}^c$	$\text{(g L}^{-1}\text{)}^d$	
120	9×10^{-16}	65	1.4×10^{-16}	1.6×10^{-16}	7.6×10^{-16}	211	0.21

Notes: ^a volume for spherical particles = $\pi d^3/6$, ^b mass = $V_d \cdot \text{bulk density (g cm}^{-3}\text{)}$. Bulk densities of oAP and pAP are 1.33 and 1.13 g cm^{-3} .⁹³ ^c $V_{\text{H}_2\text{O}} = V_w - V_d$, ^d $S_{\text{nano}} = m \cdot 1000/V_{\text{H}_2\text{O}}$. ^e Mass ratio = $m/(V_{\text{H}_2\text{O}} \cdot 1 \text{ g cm}^{-3})$.

Values of C_{nano} are $\sim 120\times$ and $14\times$ higher than the S_{bulk} for *oAP* and *pAP*, respectively. Also, C_{nano} for *oAP* is $11\times$ higher than that of *pAP* whereas S_{bulk} for *oAP* is $1.2\times$ higher than that of *pAP*. This implies that despite *oAP* and *pAP* having a high bulk water solubility, on the nanoscale they are not fully dissolved. The particles are more likely to be composed of an insoluble core that is consistent with partially soluble species.^{32,43}

These trends highlight the limitation of using bulk water solubility of organic compounds in Köhler based hygroscopic models of nanosized aerosol particles composed of sparingly soluble organic compounds. Consequently, it also emphasizes the importance of accounting for solubility limits when evaluating hygroscopic properties; molar volume alone is not necessarily a driver for droplet growth; especially in this case where structural isomers are examined. The calculated concentration enhancement is due to the increase in surface area with consequences on dissolution rate.^{94,95} Therefore, intermolecular and intramolecular forces need to be taken into account to understand changes to water uptake behavior with size of OA, specifically for isomeric compounds.

Summary and Implications

The hygroscopic properties of two NOC compounds namely, *o*-aminophenol and *p*-aminophenol were investigated under supersaturated and subsaturated conditions. Under supersaturated conditions, the hygroscopicity is reported in terms of the single hygroscopicity parameter (κ). The measured supersaturated hygroscopicity indicates that both *oAP* and *pAP* are slightly hygroscopic with κ -values comparable with other organic compounds of atmospheric-relevance. Under the subsaturated condition of $85 \pm 5\%$ RH, the hygroscopicity studies of these two aminophenols reveal that *oAP* is more hygroscopic than *pAP*, which is attributed to *oAP* having a higher bulk water solubility. The growth factor obtained, under subsaturated conditions, is dependent on particle size for both *oAP* and *pAP*. The range of growth factors calculated were 1.6-0.74 (*oAP*) and 1.5-0.74 (for *pAP*) with increasing dry diameter indicating that the particles undergo both growth and shrink depending on size. To explain this contrast in behavior, TEM images for both compounds were collected at small (≤ 50 nm) and large (≥ 100 nm) cross-

sectional diameters sizes. The difference in morphology between small and large particles aided in explaining the growth and shrinkage behavior. Small size particles were spherical in shape, while large size particles contained void spaces and appeared to consist of small sub-particles. Large particles with void spaces render them more susceptible to collapse during water uptake, hence the shrinking behavior.

To our knowledge, our studies here are the first to quantitatively explore the water uptake of aminophenols and their size-dependent growth factor under subsaturated conditions. Hence, our data add a new dimension to the growing hygroscopicity database of different OA with different functional groups and also highlight that importance of characterizing the physicochemical properties of OA to better explain their hygroscopicity, specifically bulk water solubility limits and morphological changes with water uptake. It should be noted that OA is generally found mixed with other atmospheric components, which could potentially lead to different hygroscopic properties. However, elucidating on the hygroscopic properties of pure OA provides significant insight on the properties that influence water uptake. Accordingly, our study emphasizes that to accurately predict the climate impact of OA, hygroscopicity and water uptake models need to be amended with aerosol physicochemical properties, such as chemical structure, solubility, and morphology. Our current work involves elucidating on the effect of mixing on the hygroscopicity of aromatic NOC.

Author contributions.

HAA conceived the idea with AAA And KAM. KAM designed, collected, and analyzed H-TDMA, and CCN experimental data. DR produced and analyzed TEM images. All authors contributed to the writing and preparation of the manuscript.

Competing interests.

The authors declare that they have no conflict of interest.

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Supporting Information Available

Detailed experimental procedures, and figures and tables showing data analysis. This material is available free of charge on Environmental Science: Processes & Impacts.

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