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# Hurricane-Induced Oceanic Carbon Changes in the Upper Ocean

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**Abstract:** Changes in marine carbon cycling due to hurricanes with different intensity and translation speeds have not been systematically investigated. This study uses an idealized coupled physical-biogeochemical model and a suite of model sensitivity analyses to better quantify the relationship between hurricane characteristics and marine property changes, including variations in air-sea carbon flux and partial pressure of carbon dioxide in water ( $p\text{CO}_2\text{w}$ ). We find that strong (category 4–5), mid-speed (5–8 m/s) storms cause the most carbon flux from the atmosphere to the ocean, and that the relationship between air-sea carbon flux and hurricane properties is non-linear. Climate models that do not consider synoptic-scale, storm-induced physical-biogeochemical coupling may underestimate regional carbon sinks.

**Keywords:** hurricane; carbon cycling; hurricane intensity; hurricane translation speed; idealized model



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## 1. Introduction

Recent studies suggest that hurricane activity has increased in the North Atlantic since the 1970s [1–3]. The hyperactive 2020 Atlantic hurricane season, with a record high number (30) of storms with highly varying translation speeds and intensities, and the highly active 2021 Atlantic hurricane season, with twenty-one named storms, has highlighted the need to understand how hurricanes perturb the ocean's base state under various conditions. It is known that hurricanes can lead to higher nutrient loading, increased algal biomass, and changes in oceanic carbon concentrations, which may subsequently affect regional and basin-scale carbon flux [4–6]. Improved understanding of multifaceted marine property variations and regional atmosphere-ocean  $\text{CO}_2$  exchange in response to hurricanes is therefore critical for better quantification of the ocean's role in the global carbon cycle and for predicting future climate change.

The partial pressure of carbon dioxide in water ( $p\text{CO}_2\text{w}$ ) is a function of ocean temperature, salinity, alkalinity, and dissolved inorganic carbon (DIC) concentrations, which in turn are functions of phytoplankton primary production and gas exchange [4,7,8]. Hurricanes affect  $p\text{CO}_2\text{w}$  physically, by upwelling and entraining cold, carbon-rich water, and biologically, by upwelling and entraining nutrients that stimulate phytoplankton blooms [9–11]. The resulting high phytoplankton concentrations decrease  $p\text{CO}_2\text{w}$  by decreasing the amount of DIC in the water. Studies have shown that the extent of these physical and biological responses depend on the hurricane's intensity and translation speed, as well as the initial conditions of the upper water column [9,12–17]. Increased intensity and decreased translation speed are correlated with lower sea surface temperature (SST), more nutrient injection to surface waters, and larger phytoplankton blooms [4,9,12,18–22]. Due to the time it takes for the water column to stabilize enough to support new growth and for phytoplankton to assimilate nutrients into new growth, phytoplankton response to a hurricane's passage is time-lagged by a few days [12,23]. Hurricanes can reduce  $p\text{CO}_2\text{w}$  to the point where the ocean can switch from a source of

carbon into a sink [24,25]. Several studies have shown a net decrease in  $p\text{CO}_2\text{w}$  after a hurricane's passing [10,24,26], though this is dependent on many factors, including the upper ocean condition, the hurricane's translation speed, and its intensity. Other studies have shown an increase in  $p\text{CO}_2\text{w}$  after a hurricane's passage, and a net efflux of  $\text{CO}_2$  from the ocean to the atmosphere [8,26]. The many variables involved in predicting and understanding each individual hurricane-ocean interaction make an overall estimation and prediction of future hurricanes' impact challenging.

While numerous factors affect hurricane impact on  $p\text{CO}_2\text{w}$ , we hope to piece apart the most important factors. This effort will begin with an idealized modeling study, using biological and physical profiles from the Gulf of Mexico, that will allow us to examine the effect of translation speed and hurricane intensity in a controlled setting. This is a building block upon which we will later add increasing complexity to better understand how hurricanes affect ocean  $p\text{CO}_2\text{w}$  in a wider context. The Gulf of Mexico was chosen because all necessary in situ depth profiles for biological and physical variables were readily available.

This study aims to examine drivers of  $p\text{CO}_2\text{w}$  change due to variations in hurricane intensity and translation speed via a coupled physical-biogeochemical idealized model. We will address three specific questions: (1) What is the role of hurricane-induced phytoplankton blooms in reducing, neutralizing, or reversing carbon efflux from the ocean? (2) Is there a spatial asymmetry in  $p\text{CO}_2\text{w}$  variation like that of ocean temperature and phytoplankton bloom in response to hurricane passage [9,23,27]? (3) Which combinations of hurricane translation speeds and intensities lead to the largest  $p\text{CO}_2\text{w}$  change? Addressing these questions using field data is very challenging because of the scarcity of observations and their limited temporospatial coverage. The advantage of using an idealized model for this investigation lies in the fact that many variables can be controlled, so the influence of the variables of interest (hurricane intensity and translation speed) can be examined separately. This systematic investigation can offer insights that would be difficult to gather with other methods. It is our hope that the idealized modeling and process analyses reported herein, while difficult to be validated directly against in situ observations, can provide valuable information to aid the design of future ocean carbon field campaigns and realistic modeling programs.

## 2. Materials and Methods

The modeling system used in this study consists of the Regional Ocean Modeling System (ROMS) [28] coupled to a biological model considering both nutrient and carbon cycling [29,30]. The Fennel model has been successfully applied and validated for the Gulf of Mexico [31]. The idealized model domain is a horizontally uniform and vertically stratified ocean basin (3600 km long by 2800 km wide), with 61 vertical levels and grid spacing of 8 km (Figure 1). The Coriolis force has been approximated using an f-plane to simulate 15° N. The boundary conditions are closed on the western boundary and use the gradient condition on all others. The bathymetry is set constant at 2000 m. The Generic Length Scale (GLS) vertical mixing scheme with the Kantha and Clayson stability function was applied [32–34]. More details of this coupled modeling system configuration are given in [23].

Our investigation consisted of a control model experiment and a suite of model sensitivity experiments. As was done in [23], in all cases, an idealized hurricane moved due westward with wind speeds calculated from the axisymmetric Holland wind model [35]. The model started with calm winds at the location  $x = 3000$  km,  $y = 1500$  km, and ramped up to full strength in 24 h (Figure 1). The wind speed was converted to wind stress using the formula

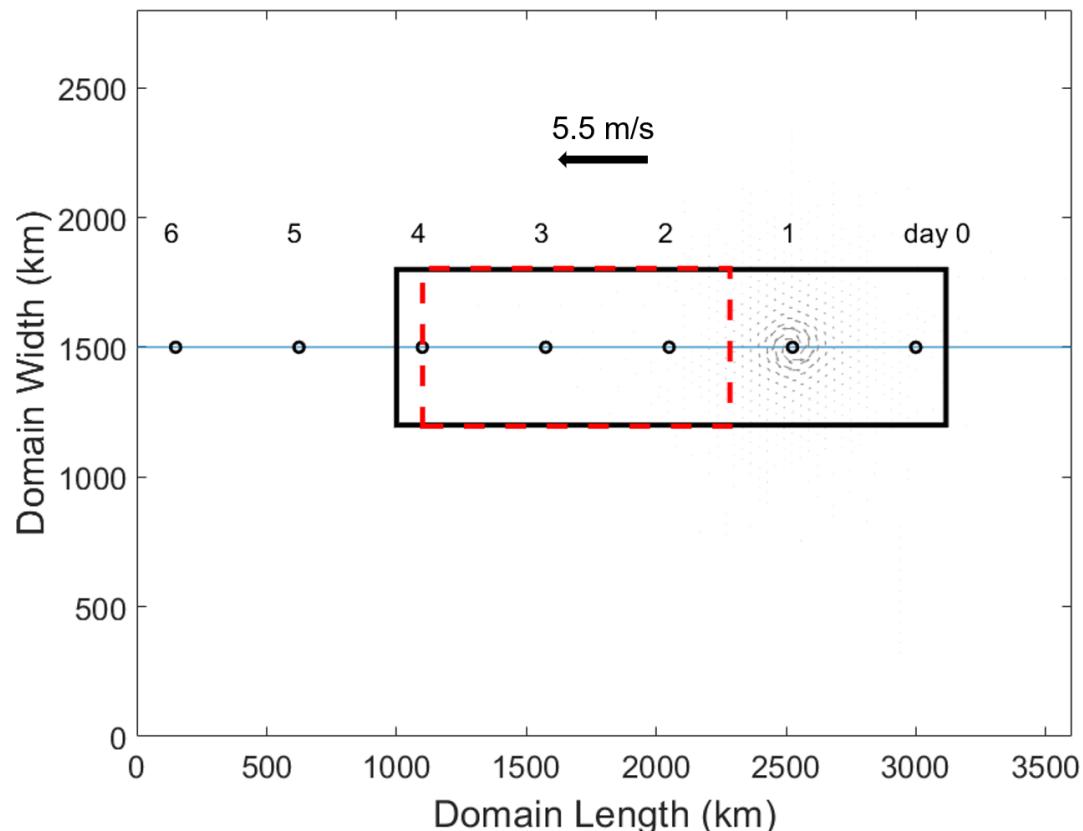
$$\text{Stress} = \rho C_d W^2 \quad (1)$$

where  $\rho$  is air density,  $C_d$  is the drag coefficient, and  $W$  is the wind speed in  $\text{m s}^{-1}$ . For the control run, the hurricane had a Category 5 (per Saffir-Simpson scale, <https://www.nhc.noaa.gov/aboutsshws.php>, accessed on 10 December 2021) intensity and a translation

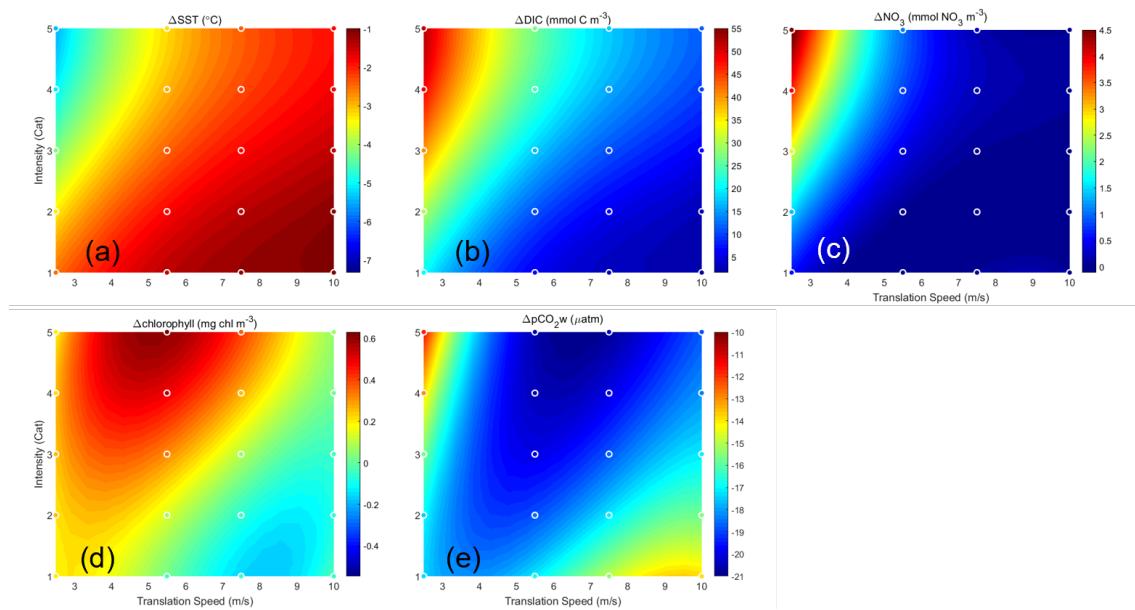
speed of  $5.5 \text{ m s}^{-1}$ . With this speed, the storm system could make landfall (reaching the western ocean boundary) in about 6.4 days. The hurricane wind forcing was one way, from the atmosphere to the ocean; ocean feedback to the hurricane was omitted in this study, but will be discussed in a future correspondence. The simulation was run for 10 days, with output every two hours. Additional sensitivity runs were designed to examine ocean response to each combination of hurricane intensity (Categories 1 through 5) and translation speed ( $2.5 \text{ m s}^{-1}$ ,  $5.5 \text{ m s}^{-1}$ ,  $7.5 \text{ m s}^{-1}$ , and  $10 \text{ m s}^{-1}$ ). The minimum central pressure and maximum wind speeds are shown in Table 1.

**Table 1.** Minimum central pressure (mB) and maximum wind speed ( $\text{m s}^{-1}$ ) for each hurricane category. Minimum central pressure is user-defined in the Holland wind model, while wind speed is calculated by the model.

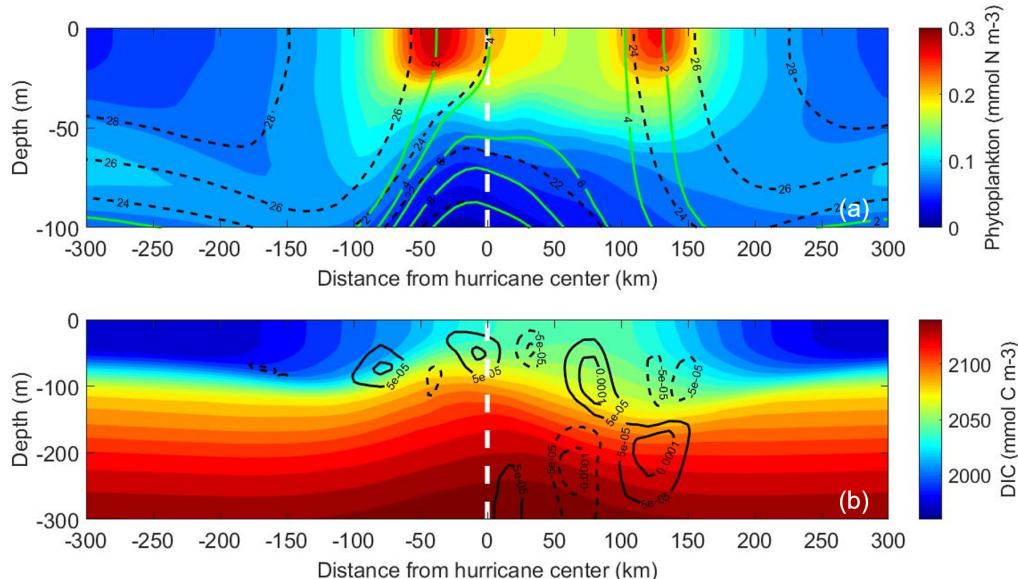
Hurricane Category	Minimum Central Pressure (mB)	Maximum Wind Speed ( $\text{m s}^{-1}$ )
Cat 1	975	37.5
Cat 2	965	44.6
Cat 3	950	54.5
Cat 4	930	66.8
Cat 5	920	72.8



**Figure 1.** The ROMS model domain, 3600 km by 2800 km. Black circles denote the location of the control run hurricane's center at the beginning of the day labeled above it. The black rectangle shows the averaged area used in Figure 2, while the dashed red rectangle shows the zonally averaged area used in Figure 3. Grey vectors indicate the hurricane wind field, and the hurricane direction with the control run's translation speed is shown.



**Figure 2.** Comparison of the change in surface properties between all combinations of translation speed and intensity. Values were computed by averaging spatially over a subregion defined by 1000 km to 3116 km, and 1200 km to 1800 km (a common area affected by all translation speeds). The difference between the last time step and the first time step was calculated. Dots indicate the calculated values. Overlaid is a third-order polynomial fitting (color shading) of the data points, extrapolated to translation speeds 2.5 through 10 m s<sup>-1</sup> for Categories 1–5. Six properties are shown: (a)  $\Delta SST$  (°C), (b)  $\Delta DIC$  (mmol m<sup>-3</sup>), (c)  $\Delta NO_3$  (mmol m<sup>-3</sup>), (d)  $\Delta chlorophyll$  (mg m<sup>-3</sup>), and (e)  $\Delta pCO_2w$  (μatm).



**Figure 3.** Vertical profiles of the control simulation. (a) Time-averaged phytoplankton (colors), temperature (black dashed contours), and nitrate (green contours). (b) Time-averaged DIC (colors) and vertical velocity (m s<sup>-1</sup>, black contours; solid lines indicate upward motion, and dashed lines indicate downward motion). All variables have been zonally averaged and time-averaged over three inertial periods. The view is westward in the direction of the translating hurricane, and the white dashed line indicates the hurricane center.

The biological model of this coupled system included key interactions for both the nitrogen and carbon cycles [29–31]. The Fennel biological model computes twelve biological

tracers: nitrate ( $\text{NO}_3$ ), ammonium ( $\text{NH}_4$ ), chlorophyll, phytoplankton, zooplankton, large and small detritus concentration (for both carbon and nitrogen), total inorganic carbon (TIC), alkalinity, and oxygen [29,31]. Chlorophyll concentrations in the model are related to phytoplankton concentrations; they increase based on temperature, photosynthetically available radiation, and concentrations of  $\text{NO}_3$  and  $\text{NH}_4$ , and decrease as a result of phytoplankton mortality and grazing by zooplankton populations [29]. The physical, biological, and nutrient profiles used for model initial conditions were based on hydrographic observations from Stations 1 and 12 of the Gulf of Mexico and East Coast Carbon Cruise 2 (GOMECC-2, <https://www.aoml.noaa.gov/ocd/gcc/GOMECC2/>, accessed on 10 December 2021). The initial atmospheric  $\text{pCO}_2$  ( $\text{pCO}_{2a}$ ) was set at 370  $\mu\text{atm}$  and was held constant throughout the 10-day simulation. TIC and alkalinity were chosen so that initial  $\text{pCO}_2w$  was 370  $\mu\text{atm}$  as well, making initial  $\text{CO}_2$  air-sea flux zero. This way, the changes in  $\text{pCO}_2w$  and  $\Delta\text{pCO}_2$  due to the hurricane could be clearly examined. As in McGee and He (2018), a simple day-night cycle was applied for the short-wave radiation ( $E_s$ ) in all simulations, given by

$$E_s = -533 \cos \frac{2(\pi)t}{24}, \text{ if } E_s > 0, \quad (2)$$

$$E_s = 0, \text{ if } E_s < 0 \quad (3)$$

where  $t$  is time in hours [23,36]. This day-night cycle does not include the effects of cloud cover, which would reduce the amount of short wave radiation available for phytoplankton growth.

### 3. Results

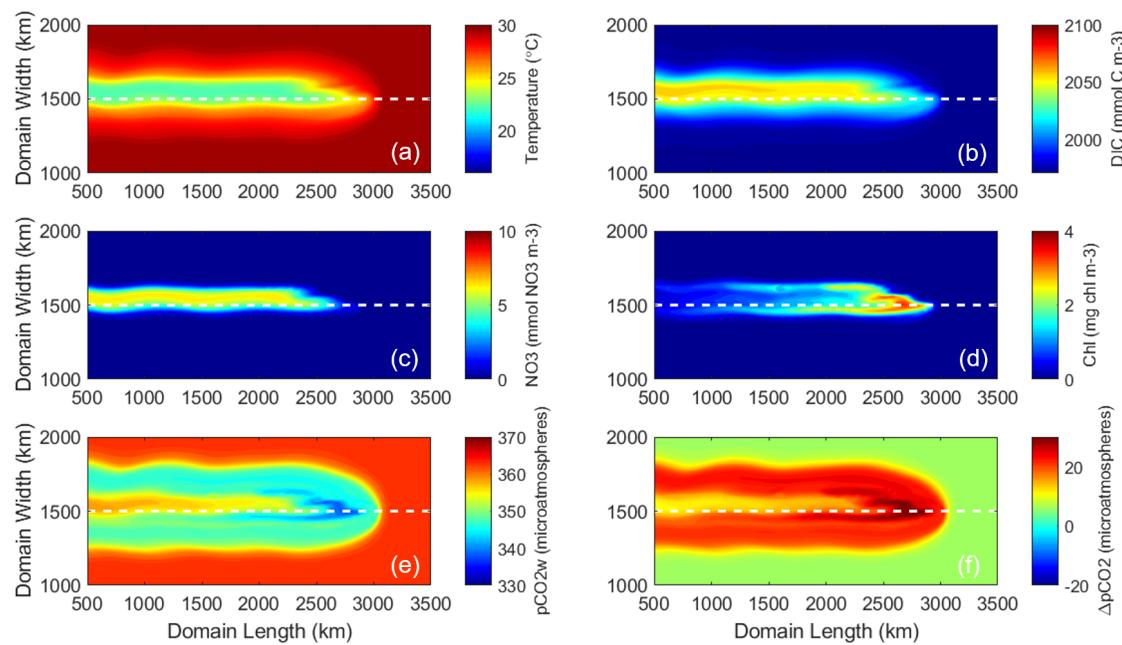
#### 3.1. Control Run

The control run examined the ocean's response to the idealized hurricane at the surface and subsurface. The tight coupling of the marine physical and biogeochemical systems was evident in the model solutions. Model results show an asymmetric (right-hand biased) upwelling of cold water and nutrients DIC and  $\text{NO}_3$  (Figure 4a–c). The greatest upwelling occurred on the right side of the storm's track. Upwelled nutrients stimulated new phytoplankton growth, as indicated by chlorophyll (chl) concentration, which by Day 8 was an order of magnitude larger than the initial concentration (Figure 4d). The chl concentration is highest around the 2800 km length mark where the full-strength storm first enters the study domain. This is because phytoplankton in this area have more time to react to the storm passage and assimilate nutrients that are upwelled than those further to the left/west.

Compared to other marine properties, the response of  $\text{pCO}_2$  occurred over a larger spatial area (Figure 4e). The initial ocean  $\text{pCO}_2w$  of 370  $\mu\text{atm}$  decreased to a minimum of 331.2  $\mu\text{atm}$  in response to this hurricane. The area of largest  $\text{pCO}_2w$  decrease (between 2500 and 3000 km domain length marks) corresponded to the region with cool temperatures (Figure 4a), low DIC (Figure 4b), and the highest concentration of phytoplankton growth (thus carbon uptake, Figure 4d). West of this area (between 500 and 2000 km domain length marks), a larger injection of DIC from the deep ocean slowed the decrease of  $\text{pCO}_2w$  due to the physically (temperature) and biologically induced carbon uptake, resulting in an asymmetric (right-hand biased) band of  $\text{pCO}_2w$  with elevated values compared to surrounding waters. Averaged over the common study area (the black box in Figure 1), the change in  $\text{pCO}_2w$  was  $-20.6 \mu\text{atm}$ ; SST decreased by  $3.2^\circ\text{C}$ ; DIC increased by  $27.1 \text{ mmol C m}^{-3}$ ; and chl increased by  $0.63 \text{ mg chl m}^{-3}$ . Because  $\text{pCO}_{2a}$  was held constant at 370  $\mu\text{atm}$  and  $\Delta\text{pCO}_2$  (defined as  $\text{pCO}_{2a}$ - $\text{pCO}_2w$ ) was positive over this area (Figure 4f), the ocean acted as a carbon sink in response to this hurricane.

The three-dimensionality of ocean responses can be examined in a subsurface transect view. The subsurface conditions for phytoplankton, temperature,  $\text{NO}_3$ , DIC, and ocean vertical velocity (Figure 3) were averaged across 1184 km (red box in Figure 1) and over three inertial periods (one inertial period at  $15^\circ$  N is 46.36 h), the same spatial and temporal averages that were done in [23,36]. These spatial and temporal averages reveal the overall

trends and allow for direct comparison with the previous two papers. In the subsurface, greater upwelling of colder water,  $\text{NO}_3$ , phytoplankton concentration, and DIC all occurred on the right side of the hurricane's track (Figure 3a,b). This asymmetry is due to two factors. First, resonance between inertial oscillations in the ocean and the clockwise turning of the hurricane's wind stress vector with time lead to increased upwelling and entrainment on the right side of the track [9,27,37]. Second, recirculation cells, indicated by alternating upward and downward motions (Figure 3b), destroy potential vorticity at the surface and restratify the water column; the resulting decreased turbulence leads to better conditions for phytoplankton growth [23,36]. These recirculation cells are larger and deeper on the right side of the track than the left.

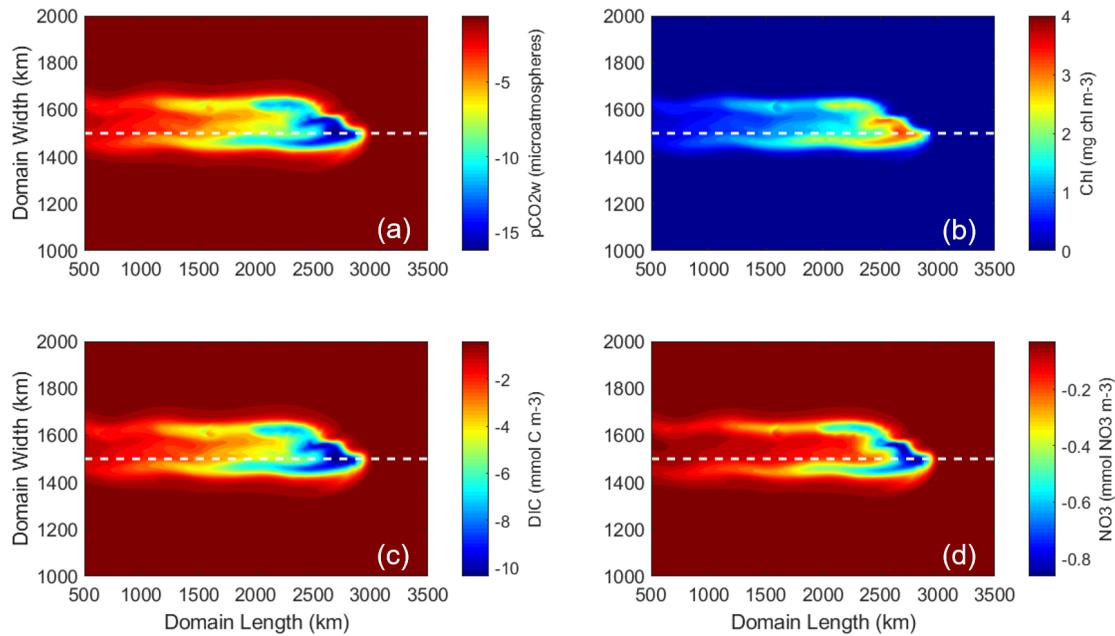


**Figure 4.** Surface plots of six variables in the control simulation at Day 8, 6 h. (a) SST ( $^{\circ}\text{C}$ ). (b) DIC ( $\text{mmol m}^{-3}$ ). (c)  $\text{NO}_3$  ( $\text{mmol m}^{-3}$ ). (d) Chlorophyll ( $\text{mg m}^{-3}$ ). (e)  $\text{pCO}_{2w}$  ( $\mu\text{atm}$ ). (f)  $\Delta\text{pCO}_2$  ( $\mu\text{atm}$ ), defined as  $\text{pCO}_{2a}-\text{pCO}_{2w}$ .

### 3.2. Sensitivity Experiments

#### 3.2.1. Effects of Biology

The first sensitivity test (S1) allowed no biological growth. The (Control-S1) difference maps (Figure 5) sampled on Day 8, at 6 h therefore highlight the effect of biological growth in the Control. The chl concentration in the control run was up to  $3 \text{ mg m}^{-3}$  higher than in S1 (Figure 5b). In the area with the most new growth, nearly  $0.8 \text{ mmol m}^{-3}$  more  $\text{NO}_3$  as well as approximately  $10 \text{ mmol m}^{-3}$  more carbon (Figure 5c,d) were taken up by the phytoplankton. As a result, in the area with new growth,  $\text{pCO}_{2w}$  in the Control was as much as  $16 \mu\text{atm}$  lower than that in S1 (Figure 5a). While  $\Delta\text{pCO}_2$  remained positive in S1 (not shown), without the effect of biology, the ocean in S1 was a less effective carbon sink than in the Control.



**Figure 5.** Surface plots of the control simulation results minus S1 (simulation without biological growth) results at day 8, 6 h. (a)  $p\text{CO}_2\text{w}$  ( $\mu\text{atm}$ ). (b) Chlorophyll ( $\text{mg m}^{-3}$ ). (c) DIC ( $\text{mmol C m}^{-3}$ ). (d)  $\text{NO}_3$  ( $\text{mmol m}^{-3}$ ).

### 3.2.2. Effects of Translation Speed and Intensity

Twenty model sensitivity experiments (each simulation ran for the same 10 days as the Control did) were performed to elucidate variations in SST, surface DIC,  $\text{NO}_3$ , chl, and  $\Delta p\text{CO}_2\text{w}$  in response to different combinations of hurricane intensities (Categories 1 through 5) and translation speeds ( $2.5 \text{ m s}^{-1}$ ,  $5.5 \text{ m s}^{-1}$ ,  $7.5 \text{ m s}^{-1}$ , and  $10 \text{ m s}^{-1}$ ). To obtain a consistent analysis for each marine property, their mean changes ( $\Delta$ ) were quantified by spatially averaging their differences before and after hurricane impacts (sampled at the first and last time step of each model experiment) over a subset of the model domain defined by 1000 km to 3116 km domain length marks, and 1200 km to 1800 km domain width marks (black box in Figure 1). This subdomain is the common region affected by hurricanes in all sensitivity simulations. For each variable, the ensemble of 20 mean change values ( $\Delta$ ) were fit into a function of hurricane intensity and translation speed using third-order (in both x and y) polynomials to reveal the intrinsic relationship between the simulated marine property changes and their corresponding hurricane intensity and translation speed. These functional fittings are very successful, with  $R^2$  values all above 0.95.

Such an analysis reveals three interesting patterns describing the resulting marine property-hurricane relationships. Pattern 1 is a quasi-linear relationship, as seen in  $\Delta\text{SST}$  (Figure 2a),  $\Delta\text{DIC}$  (Figure 2b), and  $\Delta\text{NO}_3$  (Figure 2c). Here the largest SST cooling (up to  $-5^\circ\text{C}$ ) and the maximum increases in surface DIC (up to  $50 \text{ mmol C m}^{-3}$ ) and surface  $\text{NO}_3$  (up to  $4.5 \text{ mmol NO}_3 \text{ m}^{-3}$ ) occurred when the ocean was impacted by the Category 5 hurricane moving at the slowest speed under study ( $2.5 \text{ m s}^{-1}$ ). As hurricane intensity decreased and translation speed increased, the magnitudes of  $\Delta\text{SST}$ ,  $\Delta\text{DIC}$ , and  $\Delta\text{NO}_3$  all rapidly decreased. This pattern occurs due to physical processes and is driven mainly by hurricane-induced upwelling. The weaker storms and the strong, but fast-moving storms are not as capable as the strong and slow-moving storms in transporting cold, nutrient-rich deep ocean water to the surface.

Pattern 2 has a dipole structure, as seen in the surface chl response (Figure 2d). Interestingly, the largest chl increase ( $0.6 \text{ mg chl m}^{-3}$ ) occurred when the ocean was impacted by the Category 5 hurricane moving at  $5.5 \text{ m s}^{-1}$  and it did not co-locate with the maximum  $\text{NO}_3$  increase. This is because the maximum phytoplankton growth rate depends on both

nutrient concentration ( $\text{NO}_3$ ) and water temperature [31,38]. The nutrient and temperature conditions produced by the Category 5 storm moving at  $5.5 \text{ m s}^{-1}$  constitute a “sweet spot” for rapid phytoplankton growth.  $\Delta\text{chl}$  decreases as storms become weaker and/or move faster, and can become negative. The most negative  $\Delta\text{chl}$  (up to  $-0.2 \text{ mg chl m}^{-3}$ ) occurred when the ocean was impacted by a fast moving ( $8\text{--}9 \text{ m s}^{-1}$ ) Category 1 hurricane. There SST cooling was weak and fewer nutrients were available. More phytoplankton were lost to mortality and zooplankton grazing than produced, resulting in a surface chl deficit after the storm passage. Pattern 2 is therefore controlled primarily by biological reactions.

Pattern 3, as seen in  $\Delta\text{pCO}_2\text{w}$  (Figure 2e), has the most complex tripole structure. While  $\Delta\text{pCO}_2\text{w}$  is negative for all cases, showing that partial pressure of ocean  $\text{CO}_2$  was reduced after passages of all hurricanes under study, there are two local maxima (representing the smallest  $\text{pCO}_2\text{w}$  changes) and one local minimum (representing the largest  $\text{pCO}_2\text{w}$  changes). The first local maximum occurred in the fast moving ( $9\text{--}10 \text{ m s}^{-1}$ ) Category 1 hurricane, where SST cooling was weakest. Higher SST makes  $\text{CO}_2$  less soluble in seawater, thus increasing  $\text{pCO}_2\text{w}$ . The second local maximum occurred in the slowest moving ( $2.5 \text{ m s}^{-1}$ ) Category 5 hurricane. Here the large amount of DIC (which increases  $\text{pCO}_2\text{w}$ ) made available by strong hurricane-induced upwelling outweighed the enhanced solubility (which decreases  $\text{pCO}_2\text{w}$ ) associated with upwelling cooling. The local minimum occurred during the Category 4–5 hurricane moving at a speed of  $5\text{--}8 \text{ m s}^{-1}$ . Under such storm conditions,  $\text{pCO}_2\text{w}$  can become  $21 \mu\text{atm}$  lower than its pre-storm condition, making the ocean the greatest carbon sink among all storm cases under consideration. Overall, the complex ocean  $\text{pCO}_2\text{w}$  response to hurricanes results from a delicate interplay among chemical process (temperature dependent solubility), physical transport (upwelling of DIC-rich deep water), and biological reactions (phytoplankton uptake of  $\text{pCO}_2\text{w}$ ).

#### 4. Discussion

Our analyses show that right-side asymmetry in the track of hurricanes in the North Atlantic open ocean, as seen in SST cooling and phytoplankton blooms, is also a feature of ocean  $\text{pCO}_2$ . The comparisons between the control run and S1 illustrate the importance of new phytoplankton growth in reducing oceanic  $\text{pCO}_2\text{w}$ . The maximum change in  $\text{pCO}_2\text{w}$  between the simulation with phytoplankton growth (Control) and the simulation without (S1) was  $16 \mu\text{atm}$  (Figure 5a). Since the maximum  $\text{pCO}_2\text{w}$  decrease in the control run was  $33 \mu\text{atm}$  (Figure 4e), new growth accounted for nearly half (approximately 48%) of the control run’s maximum decrease in  $\text{pCO}_2\text{w}$ .

The three (quasi-linear, dipole, and tripole) patterns identified from model experiments among hurricane characteristics (intensity and speed) and hurricane-induced marine property distributions are intriguing findings that warrant further examinations by future, carefully designed observing programs. Overall, the  $\text{pCO}_2\text{w}$  responses to hurricanes result from a delicate interplay among chemical process (temperature dependent solubility), physical transport (upwelling of DIC-rich deep water), and biological reactions (phytoplankton uptake of  $\text{pCO}_2\text{w}$ ).

There are a number of limitations and simplifications in this study, which will be outlined here. In this study, the vertical ocean profile, hurricane translation speed, and hurricane intensity (after the initial ramp-up) are held constant, in contrast to the highly dynamic condition of the open ocean. The simplified conditions in this study are a deliberate choice made to isolate the effects of hurricane intensity and translation speed, and reveal physical processes that are of further interest. Because of the difficulties in gathering data under hurricane conditions and the idealized nature of this model, there are limited opportunities to provide model validation. Instead, we have relied upon model parameterizations used successfully in previous studies to represent ocean conditions, and based our analysis primarily on the differences in variable responses ( $\Delta\text{variable}$ ) instead of on the absolute values themselves. Another limitation of this modeling study is that the ocean variables do not affect the atmospheric variables in the Holland wind model, and so the ocean and atmosphere in this model are not fully coupled. Changes to  $\text{pCO}_2\text{a}$

due to a storm's influence needs to be accounted for as well in future research. Coupling an atmosphere-ocean model will also allow investigations into the effect of light levels on phytoplankton growth, as well as possible effects of hurricane precipitation. Likewise, this study used vertical ocean profiles from the Gulf of Mexico. As ocean vertical profiles change, the response to cyclones will also change. Further studies using different configurations of thermocline and nutricline depth common to the North Atlantic are needed to determine the effect of vertical ocean profiles on hurricane-induced changes in oceanic carbon concentrations.

Over the past 60 years, global tropical cyclone translation speeds have decreased by 10% [3,39]. Cyclone translation speeds have slowed in every basin except the northern Indian Ocean, with the strongest slowing trends found over the western North Pacific Ocean and around Australia (16% and 14%, respectively) [39]. During this time, global temperatures have risen by 0.5 °C [39]. This adds evidence to the idea that tropical cyclone behavior will change as the climate warms, with overall increases in intensity and precipitation rates and slower translation speeds [39–41]. The average annual translation speeds in the North Atlantic from 1950 to 2010 ranged between 20 and 25 km h<sup>−1</sup> (5.5–7 m s<sup>−1</sup>). Based on our findings, those storms could have caused a significant pCO<sub>2</sub>w change. There is evidence to suggest that extreme translation events are becoming more common, with extreme slow-translation events (over the past forty years) increasing in the southern portion of the North Atlantic Ocean [42]. As translation speeds decrease over time, the role of phytoplankton blooms in facilitating air-sea carbon exchange may increase, potentially becoming as important as physical factors such as hurricane-induced SST decrease. However, if hurricane translation speeds decrease to an average of 2.5 m s<sup>−1</sup>, the larger amount of DIC upwelled by such hurricane conditions would make the ocean a less effective carbon sink.

The cumulative effect of synoptic processes such as hurricanes on air-sea carbon cycling can be substantial. Without considering the coupling between physical and biological processes at sufficient temporospatial scales, climate models can underestimate the effect of hurricanes on pCO<sub>2</sub>w and underestimate the regional ocean's role as a carbon sink. Current climate models do not resolve synoptic, storm-induced processes, and future improvements may need to resolve synoptic, mesoscale, and even submesoscale physical-biogeochemical coupling in order to accurately capture air-sea carbon carbon exchange before, during, and after hurricanes. With this study, we provide a systematic, idealized study of the effect of varying hurricane translation speed and intensity on SST, phytoplankton, NO<sub>3</sub>, DIC, and pCO<sub>2</sub>w in the Gulf of Mexico. This information will aid the design of needed ocean carbon field campaigns and realistic modeling programs that will be able to fully understand and quantify the effect of hurricanes on oceanic carbon concentrations.

**Author Contributions:** Conceptualization, R.H.; methodology, R.H.; validation, L.M. and R.H.; formal analysis, L.M.; investigation, L.M.; resources, R.H.; data curation, L.M.; writing—original draft preparation, L.M.; writing—review and editing, R.H.; visualization, L.M.; supervision, R.H.; project administration, R.H.; funding acquisition, R.H. and L.M. All authors have read and agreed to the published version of the manuscript.

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**Informed Consent Statement:** Not applicable.

**Data Availability Statement:** A publicly available dataset and community model were used in this study. The Regional Ocean Modeling System model code can be accessed at <https://www.myroms.org> (accessed on 10 December 2021) with the permission of The ROMS/TOMS Group. The Gulf of Mexico and East Coast Carbon Cruise 2 data can be publicly accessed at <https://www.aoml.noaa.gov>.

gov/ocd/gcc/GOMECC2/ (accessed on 10 December 2021). Input files to reproduce the results of this study are openly available in Zenodo at <https://doi.org/10.5281/zenodo.5295831> (accessed on 10 December 2021).

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## Abbreviations

The following abbreviations are used in this manuscript:

MDPI	Multidisciplinary Digital Publishing Institute
DOAJ	Directory of open access journals
pCO <sub>2</sub> w	Partial pressure of carbon dioxide in water
pCO <sub>2</sub>	Partial pressure of carbon dioxide
pCO <sub>2</sub> a	Partial pressure of carbon dioxide in air
DIC	Dissolved inorganic carbon
SST	Sea surface temperature
ROMS	Regional Ocean Modeling System
chl	Chlorophyll
NO <sub>3</sub>	Nitrate
ΔpCO <sub>2</sub>	pCO <sub>2</sub> a-pCO <sub>2</sub> w

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