

1 Low Temperature Raman Imaging of Component 2 Distribution in Micron-Size Droplets

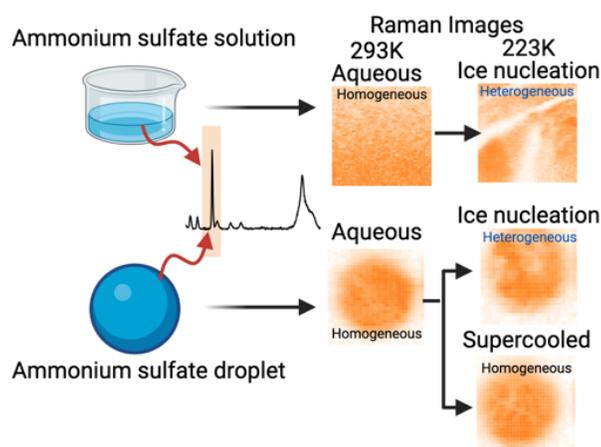
3 *Qishen Huang*^{1,2}, Peter J. Vikesland^{1,2}*

4 ¹Department of Civil and Environmental Engineering, Virginia Tech, Blacksburg, Virginia 24061,
5 United States

6 ²Virginia Tech Institute of Critical Technology and Applied Science (ICTAS) Sustainable
7 Nanotechnology Center (VTSuN), Blacksburg, Virginia 24061, United States

8 Keywords: Raman imaging, ice nucleation, ammonium sulfate, micron-size droplets, Surface-
9 enhanced Raman spectroscopy (SERS), cloud formation, atmospheric droplets.

10 TOC Graphic



11

12

13

14 **Abstract**

15 Atmospheric droplets exhibit heterogeneity due to their distinct mixing states and phases. Ice
16 nucleation is a common and stochastic atmospheric process that potentially enhances differences
17 between individual droplets. In this study, we investigated the distribution of chemical moieties
18 within micron-size droplets by recording spatial maps of sulfate, ice crystals, and gold
19 nanoparticles (AuNPs) via Raman imaging at 293K and 223K. We observed a spatially even
20 distribution of sulfate in ammonium sulfate (AS) solutions and droplets at 293K, as well as in
21 supercooled AS droplets at 223K. Spatially enriched sulfate, expelled from ice crystals, appeared
22 in frozen droplets and bulk solution at 223K. Interestingly, a fraction of the frozen droplets
23 exhibited spatially enriched sulfate distributions, while others froze evenly. A higher percentage
24 of more evenly distributed droplets were found for higher initial AS concentrations. We
25 differentiated the droplets as supercooled or frozen according to the brightness of the collected
26 optical images. Using Raman images of the AS droplets, we determined that >93% of the
27 supercooled droplets were evenly distributed, while >90% of the frozen droplets were spatially
28 enriched due to ice nucleation. We suggest that localized ice nucleation, within supercooled
29 droplets, and glass formation, within completely frozen droplets are the major factors contributing
30 to the remainder of the droplets (~10%). The different component distributions reflect the
31 stochastic nature of ice nucleation. We also investigated the distribution of functionalized pH
32 sensing AuNPs within frozen AS droplets and observed an independent distribution of AuNPs that
33 differed from either the AS distribution or the ice nucleation sites. The relative peak intensity of
34 the Raman spectrum of AuNPs changed at 223K compared to its room temperature spectrum,
35 which suggests that the formation of ice crystals in the vicinity of AuNPs at 223K altered the
36 spectral behavior of the functionalized nanoparticles.

37 Introduction

38 Particles and droplets dictate cloud radiative properties by scattering and absorbing solar and
39 terrestrial radiation;¹ by serving as cloud condensation nuclei;^{2,3} and through their involvement in
40 atmospheric chemical processes.^{4,5} Atmospheric ice particles, which can serve as cloud
41 condensation nuclei and atmospheric reaction sites, form either through homogeneous nucleation
42 of supercooled droplets or by heterogeneous nucleation from ice nuclei (e.g., immersion freezing,
43 contact freezing, deposition nucleation).^{1,6,7} Due to the relative rarity of airborne ice nuclei,
44 supercooled droplets are atmospherically ubiquitous and are found from the surface layer to the
45 upper troposphere.^{8,9} Accordingly, homogeneous nucleation is important for the formation of
46 mixed-phase clouds as it occurs in supercooled aqueous droplets when the ice saturation ($S_i =$
47 $\frac{P_{H_2O}}{P_{ice}}$) is high (where P_{H_2O} is the equilibrium partial pressure of water at the freezing temperature
48 and P_{ice} is the vapor pressure of ice at the freezing temperature).¹⁰⁻¹³

49 Individual aerosols and droplets can be subjected to different atmospheric processes and thus
50 exhibit unique properties that reflect their prior histories. Atmospheric aerosols and droplets with
51 the same composition can have differential properties depending on differences in mixing state.¹⁴
52 The phases within individual aerosols and droplets may be solid or liquid, or may be a mixture of
53 both as a result of their composition, the ambient conditions, and the stochastic nature of nucleation
54 and crystallization.¹⁵⁻¹⁹ Accordingly, even for consistent environmental conditions and droplet
55 compositions, droplets can stabilize in various states and phases. For example, nucleation and
56 crystallization can be hindered in some droplets thus resulting in formation of an amorphous solid
57 or 'glassy' state instead of ice crystal formation.²⁰

58 Ammonium sulfate (AS) particles and droplets are important constituents of upper tropospheric
59 clouds.²¹ Accordingly, understanding the supercooling and nucleation behavior of AS particles and
60 droplets is important. Previous studies have investigated ice formation temperature and nucleation
61 kinetics;^{10,22} however, the exact distribution of AS in droplets upon freezing requires study since
62 the AS internal distribution can affect droplet properties, such as pH, and in turn alter droplet-
63 atmosphere interactions.^{23,24}

64 Vibrational spectroscopies, such as infrared (IR)^{21,25,26} and Raman,²⁷⁻²⁹ have recently been shown
65 to be useful tools for the investigation of ice nucleation and phase transitions in cloud droplets.
66 Raman spectroscopy specifically, provides abundant chemical detail and higher spatial resolution
67 compared to IR.³⁰ Using single-particle Raman spectroscopy, researchers have characterized
68 mixing-states, phase transition, and ice nucleation of atmospherically relevant particles and
69 aerosols.³¹⁻³⁵ The studies rely mainly on Raman spectral variations (e.g., peak location, peak
70 intensity) which provide chemical information about the system. Raman images, compared to
71 Raman spectra, contain additional information that reflect the spatial distribution of the collected
72 spectra. Schill and Tolbert applied Raman line-mapping to investigate heterogeneous ice
73 nucleation on simulated sea-spray aerosol.³⁶ Such line-maps can be further expanded to 2-D and
74 3-D Raman images to obtain detailed spatial distribution and chemical variation of the interested
75 species. Functionalized gold nanoparticles (AuNPs) enabled surface enhanced Raman
76 spectroscopy can provide steady and intense Raman signals with high enhancement factor;
77 moreover, the spectra of functionalized AuNPs are responsive to chemical changes (e.g., pH) in
78 their vicinity.^{23,37,38} Particles in the atmosphere are important as they can serve as cloud
79 condensation and ice nuclei, and recent studies have shown the formation of gold nanocrystals in
80 supersaturated solutions.⁷ Accordingly, the addition of AuNPs to micron-size droplets may provide

81 information with respect to the crystallization of droplet components. In this study, we explored
82 the feasibility of using confocal Raman and surface-enhanced Raman spectroscopic (SERS)
83 imaging to characterize the heterogeneity of individual micron-scale droplets, to probe droplet
84 component distribution, and to examine the correlation between the distribution patterns of
85 chemical constituents, ice crystals, and AuNPs in both supercooled and phase-transitioned droplets.

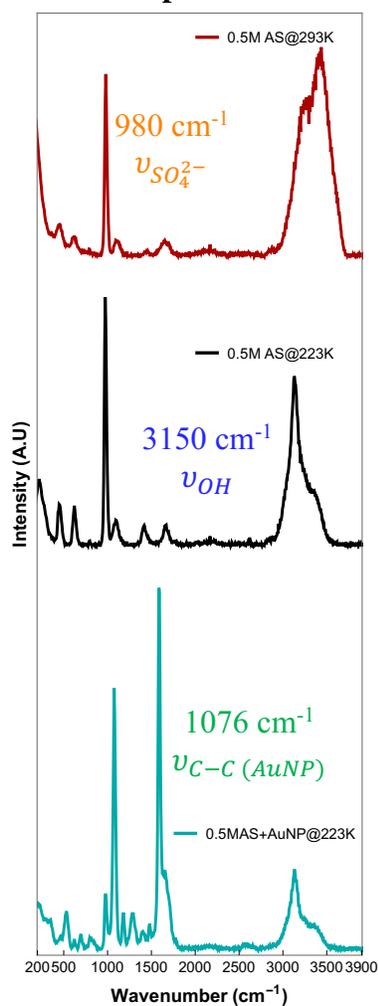
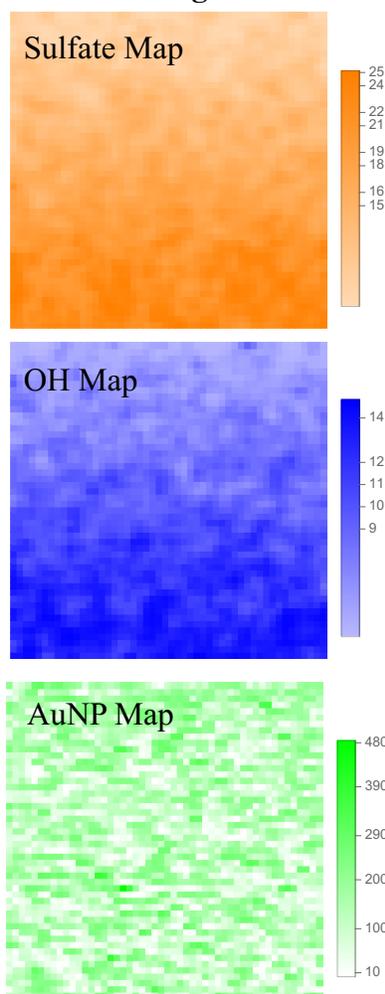
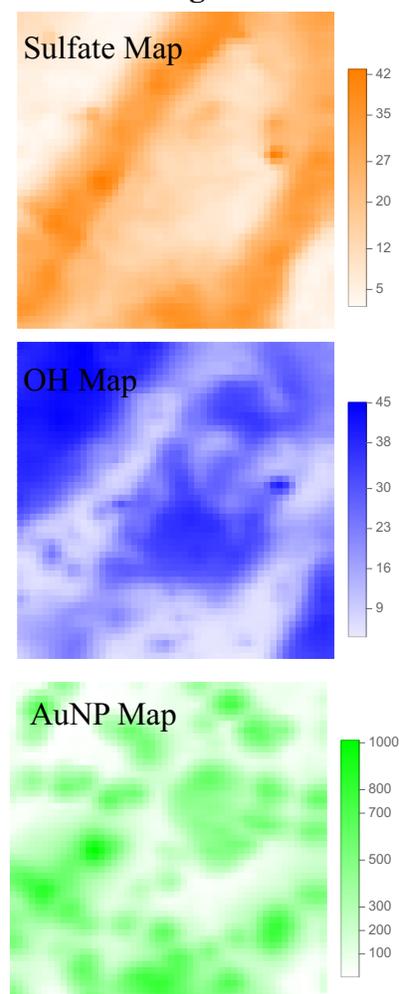
86 **Experimental Methods**

87 **Droplet generation and collection.** Micron-size droplets were generated from prepared solutions
88 (0.1-3M ammonium sulfate (AS); 1M ammonium nitrate (AN); and 1M phosphate buffer (PB))
89 using a commercial atomizer (TSI 3076, TSI Inc.). Suspensions and droplets containing pre-
90 synthesized pH nanoprobes - 4-mercaptobenzoic acid (4-MBA) functionalized AuNPs³⁷ - were
91 studied separately. The atomizer was operated in a custom chamber at 95% RH.³⁸ Mixtures
92 containing 2 mL of probe suspension and 3 mL of experimental solution were aerosolized and then
93 collected for 2 minutes on a superhydrophobic filter fixed ≈ 0.5 cm away from the atomizer outlet.
94 The superhydrophobic filter was prepared by coating 100 μ L of AEROSIL-acetone suspension (4
95 g/L) onto a PVDF filter (Millipore Sigma) via air-drying. The contact angle of droplets collected
96 on this substrate was $\approx 147^\circ$ (Figure S1). The filter was transferred into a Linkam THM600
97 thermostage for temperature control and Raman imaging.

98 **Freezing solution and droplets.** The freezing of solutions and droplets was characterized using a
99 Linkam THM600 thermostage. The accuracy of the temperature control of the thermostage was
100 evaluated using a thermogun (Figure S2). The rate of temperature decrease provided by the
101 thermostage can vary from 0.1 to 100K/min with 173K the lowest temperature possible. A
102 temperature change rate of 10K/min was elected for this study.

103 **Raman imaging and image processing.** Raman imaging was conducted at both 293K (20°C, 95%
104 RH) and 223K (-50°C) to observe both solutions and droplets. All frozen droplets were
105 equilibrated at 223K for 20 min prior to Raman imaging, as the observed optical images showed
106 negligible change with either longer freezing times or lower temperatures (down to 193K, Figure
107 S3). Raman images were collected using a confocal Raman microscope (WITec Alpha 500R) and
108 a 50× objective lens and a 633 nm laser. The laser spot has a lateral size of 0.68 μm , a vertical size
109 of 3.2 μm , and an excitation volume of 1.5 μm^3 . The spatial resolution of each Raman image was
110 1 $\mu\text{m}^2/\text{pixel}$ with each pixel containing the averaged Raman signal from the excitation volume.
111 The laser was initially focused at the droplet centroid and the integration time was 1 s for all
112 measurements. All Raman images were denoised and smoothed using a Gaussian filter. This image
113 processing method was selected out of three common methods (median filter, Gaussian filter, and
114 fast and flexible denoising convolutional neural network) based upon having the lowest mean
115 square error (MSE) and highest peak signal-to-noise-ratio (PSNR) (Figure S4).

116 **Results and Discussion**

a. Raman spectra**b. Raman images @ 293K****c. Raman images @ 223K**

117

118 *Figure 1. a) Raman spectra of 0.5M ammonium sulfate (AS) at 293K and 223K, and Raman*
 119 *spectrum of 0.5M AS + 4MBA functionalized AuNP at 223K. The peaks used to generate the*
 120 *Raman images in b) and c) are indicated. b) Raman images of sulfate (980 cm⁻¹), O-H (3230 cm⁻¹),*
 121 *and AuNP (1076 cm⁻¹ from 4-MBA) collected at 293K. Raman images are created based upon*
 122 *the peak intensities indicated in panel a). Each image contains 50×50 pixels that cover a 50×50*
 123 *μm² area. c) Raman images of sulfate (980 cm⁻¹), O-H (ice crystal, 3150 cm⁻¹), and AuNP (1076*
 124 *cm⁻¹) at 223K.*

125 Bulk ammonium sulfate solutions.

126 We collected Raman images of bulk AS in both liquid and frozen form at 293K and 223K to
 127 investigate the spatial variation of chemical moieties before and after freezing. The Raman signal
 128 intensity is proportional to the number density of molecules for conditions with controlled laser

129 intensity and integration time;³⁹ accordingly, a Raman image of a particular vibrational mode
130 reflects the spatial distribution of a given ion or molecule within the imaged area.

131 We initially collected Raman spectra to determine how the AS spectrum changes following ice
132 nucleation. As shown in Figure 1a, the AS solution spectrum at 293K has two major peaks: one
133 peak at 980 cm^{-1} , corresponding to the sulfate symmetric stretch,^{40,41} and another broad band at
134 3050-3300 cm^{-1} , which reflects the water molecule O-H vibration. The width of a Raman band,
135 represented by the full width at half maximum (FWHM), can reflect molecular crystallinity;^{42,43}
136 therefore, the measured changes in the Raman spectrum of AS system at 223K reflect ice
137 crystallization upon freezing. A sharp band at $\approx 3150 \text{ cm}^{-1}$, with a lower FWHM compared to the
138 broad O-H vibration band in aqueous solution, appeared upon freezing of AS solution. This sharp
139 band is consistent with the previously reported vibration of O-H in crystallized ice, and the shape
140 of the band (from 3050 cm^{-1} to 3500 cm^{-1}) reflects hexagonal ice crystals that formed as aqueous
141 water freezes.^{27,44-48} Using a deuterated AS/D₂O solution, we observed a similar, but shifted, peak,
142 further suggesting the origin of the peak is the O-H/O-D vibration (Figure S5). Importantly, the
143 FWHM of the 980 cm^{-1} peak changed negligibly between 293K and 223K (Figure S6) and thus
144 sulfate crystallization cannot be discriminated by monitoring the FWHM at 980 cm^{-1} . Yeung et al.
145 observed a peak shift of the sulfate band following sulfate crystallization after efflorescence;
146 however, no well-resolved shift of the sulfate peak location was observed in this study.

147 Changes in the Raman maps following freezing events in the AS/H₂O system provide additional
148 information about crystallization. Figure 1b&1c display Raman maps of sulfate and the O-H
149 vibration intensity at 293K and 223K respectively. These images were generated based upon the
150 intensities of the characteristic peaks of sulfate (980 cm^{-1}) and O-H (3150 cm^{-1}) and show the
151 distribution of sulfate and water molecule within the imaged area. Both the sulfate and O-H maps

152 collected at 293K show the homogeneous spatial distribution of sulfate and water in the aqueous
153 phase. When the temperature was decreased to 223K at a rate of 10 K/min, below the apparent
154 freezing point observed in this study (~253K), the sulfate and O-H peak intensities exhibit obvious
155 spatial variations as reflected in the Raman maps (Figure 1c). In particular, the low temperature
156 Raman maps exhibit spatially enriched sulfate and O-H distributions.

157 In some images, the sulfate and O-H enriched areas are completely complementary, while in other
158 images, their distribution overlaps (Figure 1b&1c, Figure S7). The difference in the overlap
159 between the sulfate and the O-H distributions indicates different freezing processes in the AS
160 system. To quantify the difference in the overlap between the sulfate and O-H distributions, we
161 defined the overlap ratio:

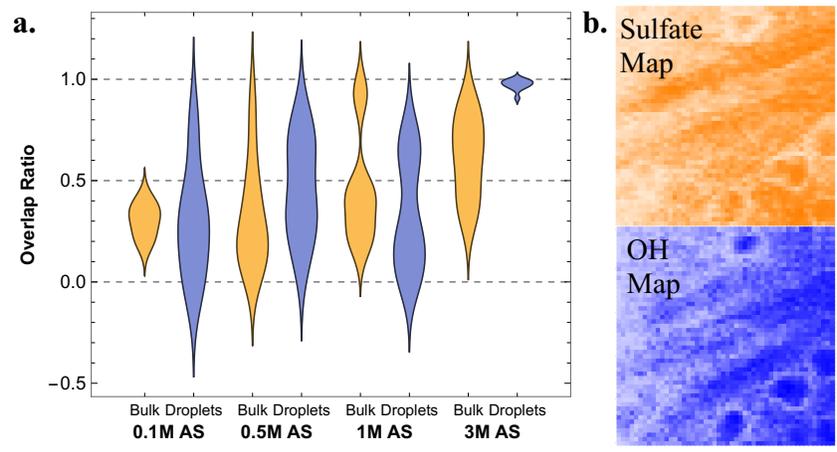
$$\text{Overlap ratio (OR)} = \frac{\text{overlapped pixels with above - average O - H and sulfate signal}}{\text{pixels with above - average O - H signal}}$$

162

163 OR reflects the overlapping regions of ice crystal and sulfate within a Raman image, and
164 numerically represents the ratio of pixels with O-H and sulfate signals that are greater than the
165 mean intensity divided by the pixels with O-H signals that are greater than the mean O-H intensity.

166 A detailed explanation of OR can be found in the Supporting Information. Low OR suggests sulfate
167 expulsion from the ice crystal structure. Previous studies have shown that sulfate and other salts
168 are excluded from ice crystals due to the mismatch between the crystal structure of AS and ice.
169 Ions in frozen electrolyte system, in the absence of supercooling, tend to enrich at the interfaces
170 between individual ice domains.⁴⁹⁻⁵¹ Sulfate expulsion from ice would suggest that the majority of
171 ammonium ions have the same spatial distribution, even though the two ions can be imbalanced
172 within interfacial areas when ice crystals form.^{51,52} The orange patterns in Figure 2a show the OR

173 distribution of various concentration AS bulk solutions. OR values are represented using a
 174 probability distribution since they can vary for a given AS concentration due to the stochastic
 175 nature of ice nucleation; however, in general, higher initial concentrations of AS have greater
 176 probability of higher OR values. This result is represented by the shift in the width of the
 177 probability distribution: the majority of OR values are lower than 0.5 at 0.1M&0.5M AS whereas
 178 1M&3M AS showed a higher probability of OR values above 0.5. The enhanced probability of
 179 high OR values reflect coexisting regions of sulfate and ice. Figure 2b provides example Raman
 180 images that show virtually identical distributions of sulfate and ice crystals. The increase in AS
 181 concentration enhanced the sulfate signal intensity uniformly in the sample. The greater probability
 182 of high OR values in high AS concentration bulk systems reflects a reduction in sulfate expulsion
 183 from newly formed ice crystals following nucleation and the inhibition of AS crystallization due
 184 to the greater viscosity of higher AS concentrations.^{17,49,53,54}



185
 186 *Figure 2. a) The distribution of overlap ratio (OR) calculated for Raman images of ammonium*
 187 *sulfate (AS) solution and droplets at 223K. The patterns denote the distribution of the OR*
 188 *represented by 8-10 Raman maps collected for corresponding AS bulk solutions and 10-15 Raman*
 189 *maps for droplets, wider parts represent higher probability of the corresponding OR. b) Raman*
 190 *images of 3M AS bulk system at 223K with identical distribution of sulfate and ice crystal that are*
 191 *denoted by high OR values.*

192 **Ammonium sulfate droplets.**

193 The observed changes in the spatial distribution measured by Raman spectroscopy following
194 freezing of AS demonstrate our general approach. We applied a similar imaging procedure to 10-
195 40 μm AS droplets to investigate ice crystallization within droplets and to probe the differences
196 between individual droplets. The findings of our previous study show that the atomization process
197 preserves droplet composition compared to original bulk solutions and that the droplets remain
198 stable under 293K and 95% RH.^{23,38} For Raman imaging during droplet freezing, the low
199 temperature limits the diffusion of water molecules thus leading to elongated equilibration
200 times.^{55,56} The change in the droplet size was negligible according to the optical images at 223K
201 (Figure S3). Therefore, we assume that AS concentrations in droplets are identical to the original
202 bulk solutions for Raman imaging at both 293K and 223K in this study. The droplets studied herein
203 were $\approx 13 \pm 3.9 \mu\text{m}$ on average. This average approximates the average diameter of droplet
204 clouds and cirrus clouds of $\approx 10 \mu\text{m}$.⁵⁷ The droplet temperature was decreased at a rate of 10 K/min
205 to minimize mass transport from supercooled droplets to ice particles due to vapor pressure
206 differences and to suppress heterogeneous nucleation.²⁰ Ice particles possess a lower equilibrium
207 vapor pressure relative to supercooled water droplets due to their greater stability. This vapor
208 pressure difference can induce ice crystal growth.^{7,58}

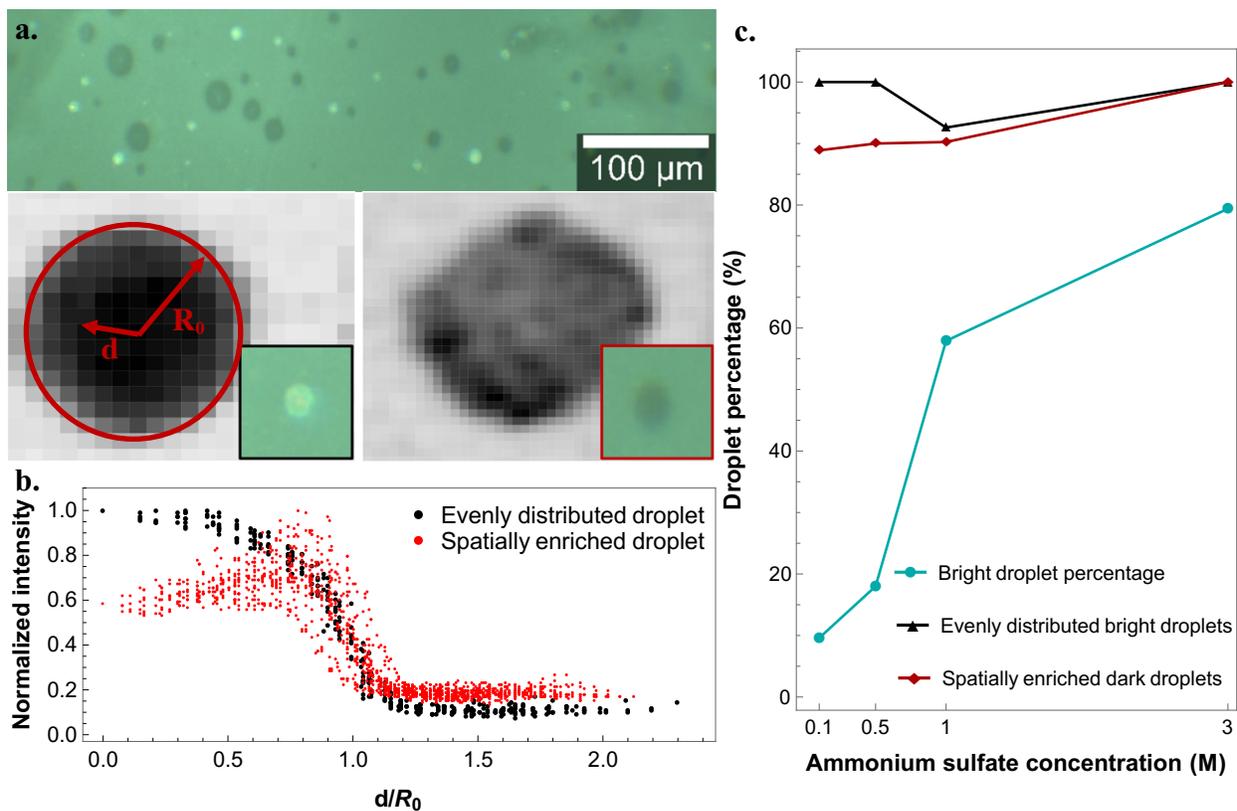
209 According to the collected optical images, about 10-80% of the AS droplets exhibited a visible
210 change in their appearance from bright to dark at 223K. This change was dependent on the initial
211 AS concentration (Figure 3a&3c). This temperature is comparable with literature-reported values
212 for the freezing of AS droplets.^{25,26,47,59} The occurrence of different brightness levels at 223K
213 suggests that the droplets are in different states or phases under the same temperature condition.
214 To characterize the difference between the bright and dark droplets, we collected Raman images
215 of aqueous AS droplets at room temperature as well as bright and dark AS droplets at 223K. The

216 averaged Raman spectra from various AS droplets at 223K showed shifts in O-H peak location
217 and variations in O-H peak shapes that indicate the different states of droplets after freezing (Figure
218 S6). Moreover, Raman images of aqueous AS droplets at 293K consistently exhibited evenly
219 distributed sulfate and O-H (Figure S8). In contrast, low temperature Raman maps of AS droplets
220 showed two distinct patterns: evenly distributed or spatially enriched sulfate (Figure 3a).

221 To better visualize and determine the chemical moiety distribution within droplets, we transformed
222 the Cartesian coordinates of a droplet image into distance from the droplet center (similar to polar
223 coordinates) and Raman peak intensity to present our collected Raman images (Figure 3b). This
224 method was used to simultaneously compare multiple collected maps and objectively differentiate
225 evenly distributed and spatially enriched analyte distributions (Figure S9). The x-axis of the
226 transformed graph represents the distance (d) of each pixel from the pre-defined droplet center,
227 while the y-axis represents the relative intensity at a given distance. We used a normalized distance
228 (d/R_0 ; where R_0 is the equivalent radius of a droplet) to denote the droplet edge ($d/R_0 \approx 1$) and to
229 generate comparable length scales when overlaying transformed images because the chemical
230 species distribution was independent of droplet size (Figure S10). In a typical droplet Raman map,
231 pixels outside the droplet have low intensity, whereas pixels inside the droplet have higher signal
232 intensities. We applied a threshold to set the droplet area by counting the pixels inside droplets.
233 The as produced intensity-distance profiles succinctly categorize the homogeneity in the droplet
234 signal intensity and its components. The signal intensity of an evenly distributed droplet, such as
235 an aqueous AS droplet, should be flat when the normalized distance (d/R_0) is less than one and
236 then decrease sharply near the droplet edge ($d/R_0 \approx 1$), thus resulting in a clearly defined reverse-S
237 shape as illustrated in Figure 3b. In contrast, spatially enriched droplets have less defined patterns,

238 with the highest signal appearing at the most concentrated area in the droplet; however, similar to
239 homogeneous droplets, the signal intensity decreases at the droplet edge.

240 Across all AS droplets, >93% of the bright droplets had evenly distributed Raman images while
241 >90% of the dark droplets showed spatially enriched Raman images (Figure 3c). According to the
242 observed change in the Raman image for the bulk AS system from evenly distributed at 293K to
243 spatially enriched at 223K, which reflect ice nucleation and crystallization during the freezing, the
244 same change in Raman images for AS droplets should also correlate with the ice nucleation and
245 crystallization in AS droplets.



246
247 *Figure 3. a) Example optical images (upper panel) and example Raman images of bright and dark*
248 *droplets; b) the intensity-relative distance profile of droplets generated from the Raman images*
249 *for evenly distributed droplet and spatially enriched droplet. c) The bright droplet percentage*
250 *obtained from optical images (cyan line, more than 100 droplets were counted for each*
251 *concentration) at 223K and the percentile of evenly distributed bright droplets (black line) and*

252 *spatially enriched dark droplets (red line) according to the intensity-relative distance profile*
253 *transformed from the collected Raman images of >20 droplets for each concentration at 223K.*

254 The droplet brightness change reflects the phase transition in the AS droplets. Our confocal Raman
255 imaging system uses a reflected light microscope (incident light is transmitted through the
256 objective); therefore, less absorption (more transmitted light), and greater forward-scattered or
257 backscattered light will cause a brighter droplet image. Previous studies on ice and water droplets
258 with similar sizes to those here (e.g., 10-40 μm), have shown that the infrared light absorption
259 efficiency of ice droplets is $>5\times$ greater than that of water droplets.⁶⁰ The same study also showed
260 a comparable extinction efficiency for ice and water droplets; however, the specific direction of
261 scattered light needs to be taken into account. The side scattering of ice clouds is stronger than that
262 of water cloud droplets in the range of 60-130°, whereas water cloud droplets exhibit greater
263 forward (10 - 60°) and backscatter ($> 160^\circ$).⁶¹ We calculated greater backscattering of water
264 droplets at higher angle via Mie-scattering calculations at 532 nm using previously reported
265 refractive indices for water and ice (Figure S11).⁶² The lower absorption efficiency and greater
266 forward-scattered and backscattered light collectively suggest that the brighter droplets observed
267 in this study are in a supercooled aqueous state.

268 The strong correlation between the evenly distributed Raman images and the bright optical droplet
269 images, as well as the correlation between the dark droplets and the spatially enriched Raman
270 images (Figure 3c), suggest that bright evenly distributed droplets are supercooled droplets while
271 the dark spatially enriched droplets are frozen droplets that experienced ice crystallization and the
272 expulsion of AS upon phase transition.¹⁷ We note that we also observed $<7\%$ of spatially enriched
273 bright droplets, and a low percentage ($<10\%$) of evenly distributed dark droplets. We interpret the
274 spatially enriched bright droplets to reflect the coexistence of crystallized ice and supercooled

275 viscous water within a droplet, as reported previously.^{29,44} The evenly distributed dark droplets
276 suggest that a phase transition occurred within the AS droplets without crystallization. The evenly
277 distributed Raman images reflect the amorphous solid state of these dark AS droplets. Prior studies
278 have shown that atmospheric droplets can form glasses, an amorphous state that behaves
279 mechanically like a solid, when droplet viscosity is high or when organic matter is present.^{17,63} As
280 the temperature decreases, the viscosity of the supercooled droplets increases and that may hinder
281 nucleation and crystallization; therefore, we suggest the evenly distributed dark droplets are
282 droplets that underwent glass formation.^{17,64}

283 Following the classification of the two major droplet states at 223K, we investigated how the AS
284 concentration affected supercooling and crystallization in droplets. We evaluated the OR of
285 collected Raman images for both supercooled and crystallized AS droplets for AS concentrations
286 ranging from 0.1-3 M. The distribution of AS and water molecules in supercooled droplets should
287 be evenly distributed as observed for the aqueous phase, resulting in overlapping sulfate and O-H
288 Raman images. Micron-size AS droplets can separate into ice crystals and freeze-concentrated AS
289 domains below the eutectic temperature (< 254.5 K)⁴⁹ that result in compensating Raman images
290 of O-H and sulfate. We observed a greater OR for higher AS concentrations which suggests that
291 more droplets were supercooled under high AS concentrations.¹⁰ Herein, we assumed homogenous
292 ice nucleation as the dominating process for AS droplets. Because homogeneous nucleation occurs
293 more readily when the ice saturation ($S_i = \frac{P_{H_2O}}{P_{ice}}$) is high, droplets with high AS exhibit a decrease
294 in S_i and are more likely to be supercooled. We considered heterogenous nucleation as negligible
295 for the following reasons. First, the estimated relative humidity (RH, equivalent to ice saturation
296 S_i) of AS aerosols at equilibrium at 223K using the E-AIM model is 61.7%.⁶⁵ At this moderate
297 RH, the temperature change rate we used (10K/min) should suppress mass transfer that may initiate

298 heterogeneous nucleation.²⁰ Second, the chamber environment within our thermostage contained
299 pure air and sample droplets with no source of dust or particles that potentially serve as nuclei for
300 heterogeneous nucleation, which can suppress the probability of heterogeneous nucleation. Third,
301 our substrate is superhydrophobic which minimizes the effect of the surface on heterogeneous
302 nucleation.⁴⁸ Surface roughness can promote heterogeneous nucleation;^{66,67} therefore, we further
303 sought to evaluate the probability of surface-induced heterogeneous nucleation by testing two
304 other hydrophobic filters (PTFE filter and silane-coated aluminum). We observed similar
305 phenomenon for the three substrates (Figure S12).

306 The bright droplet percentage denoted by the cyan line (indicated by circles) in Figure 3c shows
307 an increasing percentage of bright droplets with an increase in the AS concentration that reflects
308 the higher probability of droplet supercooling at high AS concentration as the bright droplet
309 percentage is strongly correlated with the percentage of supercooled droplets. This result agrees
310 with the droplet results in OR in Figure 2a. This bright droplet percentage is equivalent to the
311 frozen fraction commonly defined in ice nucleation studies.^{50,68} We also determined the bright
312 droplet percentage ($\frac{\text{Number of bright droplets}}{\text{Number of total droplets}} \times 100\%$) of 1M ammonium nitrate (AN) and 1M
313 phosphate buffer (PB, solution pH = 7.4) droplets (Figure S13). The droplet number percentage of
314 1M AN was comparable to 1M AS. AS and AN solution have similar viscosity^{53,69}, and E-AIM
315 calculation gave similar equilibrium RH value (<1% difference) for AS and AN droplets which
316 denote similar water activity in droplets. Given that homogeneous nucleation dominates in this
317 study, we expect ice nucleation activity in AS and AN droplets to be similar, since the process is
318 primarily dictated by water activity and is affected by viscosity.^{17,70} The phosphate droplets
319 exhibited a significantly higher ratio of bright droplets at the same temperature thus implying a
320 higher probability of droplet supercooling and lower ice nucleation activity in PB droplets. This

321 difference between PB droplets and ammonium droplets suggests the effect arises from viscosity
322 differences,^{71,72} and different ice nucleation processes in the multicomponent PB.

323 The ratio of bright droplets and dark droplets can be correlated with the Boltzmann distribution.
324 According to the Boltzmann distribution, this ratio can be expressed in terms of energy:

$$325 \quad \frac{N_{state1}}{N_{state2}} = e^{-\frac{\Delta\epsilon}{kT}}$$

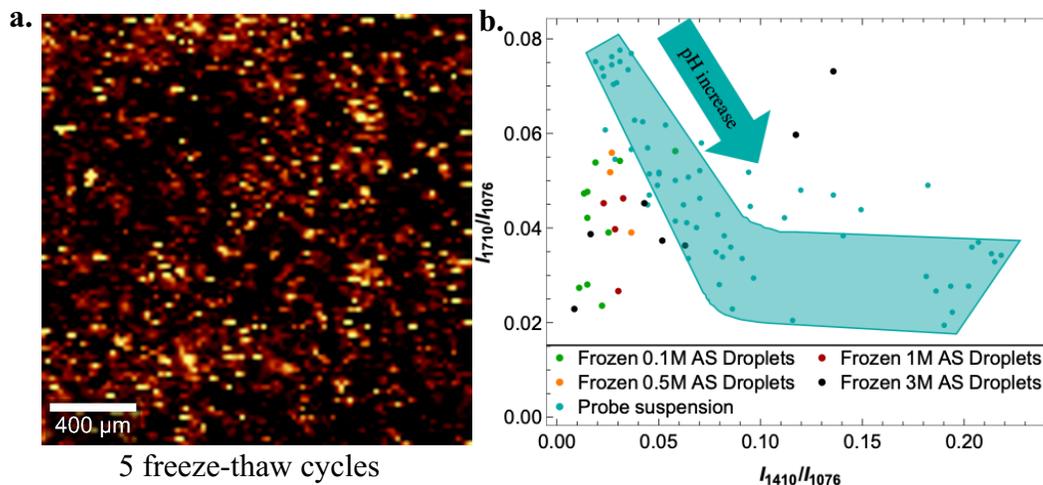
326 where $\Delta\epsilon$ is the thermodynamic energy difference between two states (e.g., supercooled droplet
327 vs. crystallized droplets, or crystallite vs. glass formation). We estimated the apparent $\Delta\epsilon$ between
328 supercooling and crystallization for 0.1, 0.5, 1, and 3M AS to be ~ 2.2 kT, 1.5 kT, -0.3 kT, and -
329 1.3 kT respectively. These values refer to the difference in thermodynamic free energy. The droplet
330 number percentage of AS, AN, and PB droplets changed only after the system reached 223K
331 (Figure S13). Similarly, the change in the brightness of the AS droplets occurred after the system
332 reached a certain temperature (223K, Figure S3). These two phenomena reflect that kinetic
333 activation is required to initiate nucleation and crystallization.¹² The values achieved from the
334 Boltzmann distribution were about 10 \times lower than the kinetic activation energy for bulk
335 homogeneous ice nucleation reported previously.¹² Since freezing is a kinetically controlled
336 nucleation process, and previous studies have shown that larger droplets tend to crystallize at a
337 higher rate⁵⁴, the smaller thermodynamic energy can be attributed to: 1) the dominant kinetic
338 parameters during ice nucleation; 2) the lower energy requirements for micron-size droplet
339 freezing compared to bulk solution. We note that the Boltzmann analysis provides only a
340 qualitative assessment of the energy given the finite number of droplets. Moreover, we categorize
341 the evenly distributed dark droplets as those that experienced glass formation; accordingly, with

342 additional droplet data that correlates with both evenly distributed and spatially enriched dark
343 droplets, we will be able to estimate the free energy difference between glass formation and
344 crystallization via droplet counting and the Boltzmann analysis.

345 **Gold nanoparticle in ammonium sulfate system.**

346 Using Raman imaging, we have shown that the distribution of sulfate and ice in droplets upon
347 freezing varies between individual droplets for AS concentrations $< 3\text{M}$. In parallel, we observed
348 the distribution of the 4-mercaptobenzoic acid (4-MBA) functionalized AuNPs (pH nanoprobe)
349 within the frozen AS system. This study was conducted to investigate the correlation between the
350 distribution of the pH nanoprobe and the ice crystal growth. Raman images of the pH nanoprobe
351 were collected according to its peak intensity at 1076 cm^{-1} which originates from the 4-MBA
352 benzene breathing mode.^{23,37} Similar to the AS solution, the pH nanoprobe were evenly distributed
353 at room temperature, but spatially enriched after being frozen in both bulk and droplets which
354 reflects the previously reported nanoparticle embedded in ice (Figure 1, Figure S7).⁷³⁻⁷⁵ We also
355 conducted multiple freeze-thaw cycles on the pH nanoprobe suspension, and demonstrated the
356 stabilization effect of polymer coating on the pH nanoprobe (Figure 4a). The collected images
357 show the independent distribution of AuNPs relative to the distribution of either AS or ice crystal
358 under the conditions of this study (Figure 1c&S7, AuNPs are enriched within both AS rich
359 domains and ice crystal domains). The aggregation of AuNPs without polymer coating should
360 have identical distribution with salt due to the exclusion by the growing ice crystal.⁷⁵ The
361 independent distribution of the pH nanoprobe observed in this study can be attributed to the
362 colloidal stabilization effect of the PEG polymer coating and the hydroxyl end group of PEG. We
363 observed no apparent change in the droplet number ratio of the droplets that contained AuNPs
364 ($\sim 22\%$ in 1M AS droplets), which suggests negligible immersion freezing and little effect on the

365 kinetics of ice nucleation in the presence of relatively low concentrations of pH nanoprobes ($\sim 10^9$
366 NPs/L).



367
368 *Figure 4. a) The distribution of pH nanoprobe in the pH nanoprobe suspension after 5 freeze-thaw*
369 *cycles. b) The correlation between two peak ratios I_{1710}/I_{1076} and I_{1410}/I_{1076} at different solution pH*
370 *(from 2-12). The cyan region denotes the peak ratios of pH nanoprobe in suspensions of different*
371 *pH at 293K. Both peak ratios can be applied to quantify the suspension pH in aqueous phase. The*
372 *correlation of two peak ratios under 223K completely deviated from that observed in aqueous*
373 *suspension (cyan region).*

374 In our previous studies, we applied 4-MBA functionalized AuNPs with SERS imaging to achieve
375 pH quantification and imaging of micron-size droplets.^{23,38} The weak correlation between
376 aggregated AuNP and ice crystals in the AS system provided us with the potential to characterize
377 pH in different domains within the AS system, and to investigate the pH difference between
378 supercooled and crystallized droplets. Investigating pH changes that occur following droplet
379 freezing may provide important information on ion enrichment and potentially the imbalance of
380 ions caused by ice crystallization.⁵¹ We evaluated the spectral behavior of the spatially enriched
381 pH nanoprobes. In aqueous solutions, the Raman spectrum of the pH nanoprobe correlates well
382 with the surrounding pH due to the protonation and deprotonation of the carboxyl group of 4-
383 MBA. At low pH conditions, the higher abundance of protonated 4-MBA results in higher intensity
384 of 1710 cm^{-1} peak (I_{1710}) and reduced intensity of the 1410 cm^{-1} peak (I_{1410}). To minimize signal

385 variability the pH sensitive peaks are normalized to the pH intensity peak at 1076 cm^{-1} (I_{1076}) which
386 is originated from benzene ring-breathing mode.^{23,37,38} pH nanoprobe suspensions with different
387 pH values (4 - 11) were prepared using 1M HCl and 1M NaOH after synthesis, and characterized
388 by Raman spectroscopy after pH adjustment. The correlation between the two peak ratios (I_{1410}/I_{1076}
389 and I_{1710}/I_{1076}) in aqueous solutions are depicted by the cyan region in Figure 4b. At 223K, the
390 correlation between the two peak ratios of the pH nanoprobe upon freezing (represented by dots
391 in non-cyan colors) deviated from that obtained in the aqueous phase. Moreover, the peak ratios
392 changed greatly within droplets with the same AS concentration (Figure 4b). The large variance
393 in spectral behavior suggests that the aggregation of AuNPs alters the Raman spectra at 223K. A
394 previous study has shown that changes in the orientation and binding modes of a molecule on the
395 gold nanoparticle surface can alter its spectrum.⁷⁶ Therefore, the spectral changes we observed can
396 be attributed to the change in molecular orientation in the aggregation of AuNPs and the change
397 in protonation of the carboxylic group of the pH nanoprobe in the ice crystals at 223K. The spectral
398 changes also indicate a state different from the aqueous state at 293K, which can be related to the
399 aggregation and assembly of pH nanoprobe at 223K. To achieve pH characterization in different
400 domains of ice, future study is needed to figure out how the functional group (4-MBA) was
401 affected by the assembly and aggregation of pH nanoprobe during ice crystal growth.

402 **Conclusions**

403 In this study we demonstrate that low temperature (223K) Raman and SERS imaging are reliable
404 tools to investigate molecule and ion distribution which can provide micro-meter scale resolution.
405 Using low temperature Raman imaging, we observed the exclusion of sulfate from ice crystal
406 during the freezing of AS solution. Both AS bulk system and AS droplets exhibited various
407 freezing processes indicated by the OR value evaluated from Raman images. The droplets at 223K

408 showed different states that reflects the stochastic nature of ice nucleation. The population of
409 droplets in different states that are characterized by Raman imaging can provide qualitative
410 estimation of the thermodynamic energy difference between the states. We studied the distribution
411 of pH nanoprobe separately using SERS imaging, which has the same experimental setup with
412 Raman imaging. The distribution of pH nanoprobe are independent from either icy crystal or
413 sulfate distribution, which provide the potential for the pH nanoprobe on charactering pH in
414 different ice domains. However, the freezing processes altered the spectral behavior of pH
415 nanoprobe; therefore, future studies on factors affect pH nanoprobe SERS spectra is required. Our
416 findings in this study show the potential of low-temperature Raman and SERS imaging on micron-
417 level characterization that can provide fundamental physical chemistry insights.

418 AUTHOR INFORMATION

419 **Corresponding Author**

420 *Qishen Huang – Virginia Tech, Blacksburg, Virginia, 24060, United States; The author is
421 currently at The Pennsylvania State University, University Park, Pennsylvania, 16802, United
422 States; Email: qqh5081@psu.edu; Phone: (434)-326-8290

423 **Funding Sources**

424 This research was supported by the National Science Foundation (CBET-1705653).

425 ASSOCIATED CONTENT

426 **Supporting Information.** Additional information on experimental details and supportive graphs
427 involved in the discussion and method section are included in the supporting information.

428 ACKNOWLEDGMENT

429 We would like to thank Dr. Weinan Leng for his assistance in Linkam thermostage operation and
430 Raman experiments, and the Virginia Tech Institute of Technology and Applied Science (ICTAS)
431 and center for Sustainable Nanotechnology for providing the instruments.

432 **References**

- 433 (1) Signorell, R.; Bertram, A.; Knopf, D. A.; Lopez, M. D. Homogeneous Ice Freezing
434 Temperatures and Ice Nucleation Rates of Aqueous Ammonium Sulfate and Aqueous
435 Levoglucosan Particles for Relevant Atmospheric Conditions. *Phys. Chem. Chem. Phys.*
436 **2009**, *11* (36), 7759. <https://doi.org/10.1039/b916865f>.
- 437 (2) Rosenfeld, D. Suppression of Rain and Snow by Urban and Industrial Air Pollution.
438 *Science* (80-.). **2000**, *287* (5459), 1793–1796.
439 <https://doi.org/10.1126/science.287.5459.1793>.
- 440 (3) Baker, M. B.; Peter, T. Small-Scale Cloud Processes and Climate. *Nature* **2008**, *451*
441 (7176), 299–300. <https://doi.org/10.1038/nature06594>.
- 442 (4) Krämer, M.; Schiller, C.; Voigt, C.; Schlager, H.; Popp, P. J. A Climatological View of
443 HNO₃ Partitioning in Cirrus Clouds. *Q. J. R. Meteorol. Soc.* **2008**, *134* (633 B), 905–912.
444 <https://doi.org/10.1002/qj.253>.
- 445 (5) Jensen, E.; Pfister, L. Implications of Persistent Ice Supersaturation in Cold Cirrus for
446 Stratospheric Water Vapor. *Geophys. Res. Lett.* **2005**, *32* (1), 1–4.
447 <https://doi.org/10.1029/2004GL021125>.
- 448 (6) Pruppacher, H. R.; Klett, J. D. *Microphysics of Clouds and Precipitation*, 2nd ed.;

- 449 Atmospheric and Oceanographic Sciences Library; Springer Netherlands: Dordrecht,
450 2010; Vol. 18. <https://doi.org/10.1007/978-0-306-48100-0>.
- 451 (7) Murray, B. J.; O'sullivan, D.; Atkinson, J. D.; Webb, M. E. Ice Nucleation by Particles
452 Immersed in Supercooled Cloud Droplets. *Chem. Soc. Rev.* **2012**, *41* (19), 6519–6554.
453 <https://doi.org/10.1039/c2cs35200a>.
- 454 (8) Tabazadeh, A.; Djikaev, Y. S.; Reiss, H. *Surface Crystallization of Supercooled Water in*
455 *Clouds*; 2002; Vol. 99. <https://doi.org/10.1073/pnas.252640699>.
- 456 (9) Rosenfeld, D.; Woodley, W. L. Deep Convective Clouds with Sustained Supercooled
457 Liquid Water down to - 37.5 °C. *Nature* **2000**, *405* (6785), 440–442.
458 <https://doi.org/10.1038/35013030>.
- 459 (10) Bertram, A. K.; Koop, T.; Molina, L. T.; Molina, M. J. Ice Formation in (NH₄)₂SO₄-H₂O
460 Particles. *J. Phys. Chem. A* **2000**, *104* (3), 584–588. <https://doi.org/10.1021/jp9931197>.
- 461 (11) Espinosa, J. R.; Vega, C.; Sanz, E. Homogeneous Ice Nucleation Rate in Water Droplets.
462 *J. Phys. Chem. C* **2018**, *122* (40), 22892–22896. <https://doi.org/10.1021/acs.jpcc.8b04788>.
- 463 (12) Ickes, L.; Welti, A.; Hoose, C.; Lohmann, U. Classical Nucleation Theory of
464 Homogeneous Freezing of Water: Thermodynamic and Kinetic Parameters. *Phys. Chem.*
465 *Chem. Phys.* **2015**, *17* (8), 5514–5537. <https://doi.org/10.1039/c4cp04184d>.
- 466 (13) Koop, T.; Luo, B.; Tsias, A.; Peter, T. *Water Activity as the Determinant for*
467 *Homogeneous Ice Nucleation in Aqueous Solutions*; 2000; Vol. 406.
468 <https://doi.org/10.1038/35020537>.

- 469 (14) Ault, A. P.; Axson, J. L. Atmospheric Aerosol Chemistry: Spectroscopic and Microscopic
470 Advances. *Anal. Chem.* **2017**, *89* (1), 430–452.
471 <https://doi.org/10.1021/acs.analchem.6b04670>.
- 472 (15) Martin, S. T. Phase Transitions of Aqueous Atmospheric Particles. *Chem. Rev.* **2000**, *100*
473 (9), 3403–3453. <https://doi.org/10.1021/cr990034t>.
- 474 (16) Marcolli, C.; Luo, B.; Peter, T. Mixing of the Organic Aerosol Fractions: Liquids as the
475 Thermodynamically Stable Phases. *J. Phys. Chem. A* **2004**, *108* (12), 2216–2224.
476 <https://doi.org/10.1021/jp036080l>.
- 477 (17) Zobrist, B.; Marcolli, C.; Pedernera, D. A.; Koop, T. *Do Atmospheric Aerosols Form*
478 *Glasses?*; 2008; Vol. 8. <https://doi.org/10.5194/acp-8-5221-2008>.
- 479 (18) Schlenker, J. C.; Martin, S. T. Crystallization Pathways of Sulfate-Nitrate-Ammonium
480 Aerosol Particles. *J. Phys. Chem. A* **2005**, *109* (44), 9980–9985.
481 <https://doi.org/10.1021/jp052973x>.
- 482 (19) Freedman, M. A. Phase Separation in Organic Aerosol. *Chem. Soc. Rev.* **2017**, *46* (24),
483 7694–7705. <https://doi.org/10.1039/c6cs00783j>.
- 484 (20) Koop, T.; Ng, H. P.; Molina, L. T.; Molina, M. J. A New Optical Technique to Study
485 Aerosol Phase Transitions: The Nucleation of Ice from H₂SO₄ Aerosols. *J. Phys. Chem.*
486 *A* **1998**, *102* (45), 8924–8931. <https://doi.org/10.1021/jp9828078>.
- 487 (21) Chelf, J. H.; Martin, S. T. Homogeneous Ice Nucleation in Aqueous Ammonium Sulfate
488 Aerosol Particles. *J. Geophys. Res. Atmos.* **2001**, *106* (1), 1215–1226.

489 <https://doi.org/10.1029/2000jd900477>.

490 (22) Rigg, Y. J.; Alpert, P. A.; Knopf, D. A. Immersion Freezing of Water and Aqueous
491 Ammonium Sulfate Droplets Initiated by Humic-like Substances as a Function of Water
492 Activity. *Atmos. Chem. Phys.* **2013**, *13* (13), 6603–6622. [https://doi.org/10.5194/acp-13-](https://doi.org/10.5194/acp-13-6603-2013)
493 6603-2013.

494 (23) Huang, Q.; Wei, H.; Marr, L. C.; Vikesland, P. J. Direct Quantification of the Effect of
495 Ammonium on Aerosol Droplet PH. *Environ. Sci. Technol.* **2021**, *55* (1), 778–787.
496 <https://doi.org/10.1021/acs.est.0c07394>.

497 (24) Pye, H.; Nenes, A.; Alexander, B.; Ault, A.; Barth, M.; Clegg, S.; Collett Jr., J.; Fahey, K.;
498 Hennigan, C.; Herrmann, H.; Kanakidou, M.; Kelly, J.; Ku, I.-T.; McNeill, V. F.; Riemer,
499 N.; Schaefer, T.; Shi, G.; Tilgner, A.; Walker, J.; Wang, T.; Weber, R.; Xing, J.; Zaveri,
500 R.; Zuend, A. *The Acidity of Atmospheric Particles and Clouds*; 2019.
501 <https://doi.org/10.5194/acp-2019-889>.

502 (25) Cziczo, D. J.; Abbatt, J. P. D. Deliquescence, Efflorescence, and Supercooling of
503 Ammonium Sulfate Aerosols at Low Temperature: Implications for Cirrus Cloud
504 Formation and Aerosol Phase in the Atmosphere. *J. Geophys. Res. Atmos.* **1999**, *104*
505 (D11), 13781–13790. <https://doi.org/10.1029/1999JD900112>.

506 (26) Hung, H. M.; Malinowski, A.; Martin, S. T. Ice Nucleation Kinetics of Aerosols
507 Containing Aqueous and Solid Ammonium Sulfate Particles. *J. Phys. Chem. A* **2002**, *106*
508 (2), 293–306. <https://doi.org/10.1021/jp012064h>.

509 (27) Celli, M.; Ulivi, L.; Del Rosso, L. Raman Investigation of the Ice Ic-Ice Ih

- 510 Transformation. *J. Phys. Chem. C* **2020**, *124* (31), 17135–17140.
511 <https://doi.org/10.1021/acs.jpcc.0c05136>.
- 512 (28) Ben Mabrouk, K.; Kauffmann, T. H.; Aroui, H.; Fontana, M. D. Raman Study of Cation
513 Effect on Sulfate Vibration Modes in Solid State and in Aqueous Solutions. *J. Raman
514 Spectrosc.* **2013**, *44* (11), 1603–1608. <https://doi.org/10.1002/jrs.4374>.
- 515 (29) Jenniskens, P.; Banham, S. F.; Blake, D. F.; McCoustra, M. R. S. Liquid Water in the
516 Domain of Cubic Crystalline Ice Ic. *J. Chem. Phys.* **1997**, *107* (4), 1232–1241.
517 <https://doi.org/10.1063/1.474468>.
- 518 (30) Riemer, N.; Ault, A. P.; West, M.; Craig, R. L.; Curtis, J. H. Aerosol Mixing State:
519 Measurements, Modeling, and Impacts. *Rev. Geophys.* **2019**, *57* (2), 187–249.
520 <https://doi.org/10.1029/2018RG000615>.
- 521 (31) Yeung, M. C.; Chan, C. K. Water Content and Phase Transitions in Particles of Inorganic
522 and Organic Species and Their Mixtures Using Micro-Raman Spectroscopy. *Aerosol Sci.
523 Technol.* **2010**, *44* (4), 269–280. <https://doi.org/10.1080/02786820903583786>.
- 524 (32) Mael, L. E.; Busse, H.; Grassian, V. H. Measurements of Immersion Freezing and
525 Heterogeneous Chemistry of Atmospherically Relevant Single Particles with Micro-
526 Raman Spectroscopy. *Anal. Chem.* **2019**, *91* (17), 11138–11145.
527 <https://doi.org/10.1021/acs.analchem.9b01819>.
- 528 (33) Guo, X.; Xiao, H. S.; Wang, F.; Zhang, Y. H. Micro-Raman and FTIR Spectroscopic
529 Observation on the Phase Transitions of MnSO₄ Droplets and Ionic Interactions between
530 Mn²⁺ and SO₄²⁻. *J. Phys. Chem. A* **2010**, *114* (23), 6480–6486.

- 531 <https://doi.org/10.1021/jp9104147>.
- 532 (34) Craig, R. L.; Bondy, A. L.; Ault, A. P. Computer-Controlled Raman Microspectroscopy
533 (CC-Raman): A Method for the Rapid Characterization of Individual Atmospheric
534 Aerosol Particles. *Aerosol Sci. Technol.* **2017**, *51* (9), 1099–1112.
535 <https://doi.org/10.1080/02786826.2017.1337268>.
- 536 (35) Collier, K. N.; Brooks, S. D. Role of Organic Hydrocarbons in Atmospheric Ice
537 Formation via Contact Freezing. *J. Phys. Chem. A* **2016**, *120* (51), 10169–10180.
538 <https://doi.org/10.1021/acs.jpca.6b11890>.
- 539 (36) Schill, G. P.; Tolbert, M. A. Heterogeneous Ice Nucleation on Simulated Sea-Spray
540 Aerosol Using Raman Microscopy. *J. Phys. Chem. C* **2014**, *118* (50), 29234–29241.
541 <https://doi.org/10.1021/jp505379j>.
- 542 (37) Wei, H.; Willner, M. R.; Marr, L. C.; Vikesland, P. J. Highly Stable SERS PH
543 Nanoprobes Produced by Co-Solvent Controlled AuNP Aggregation. *Analyst* **2016**, *141*
544 (17), 5159–5169. <https://doi.org/10.1039/c6an00650g>.
- 545 (38) Wei, H.; Vejerano, E. P.; Leng, W.; Huang, Q.; Willner, M. R.; Marr, L. C.; Vikesland, P.
546 J. Aerosol Microdroplets Exhibit a Stable PH Gradient. *Proc. Natl. Acad. Sci.* **2018**, *115*
547 (28), 7272–7277. <https://doi.org/10.1073/pnas.1720488115>.
- 548 (39) Peter Larkin. *Infrared and Raman Spectroscopy, 2nd Edition*, 2nd ed.; Elsevier, 2017.
- 549 (40) Venkateswarlu, P.; Bist, H. D.; Jain, Y. S. Laser Excited Raman Spectrum of Ammonium
550 Sulfate Single Crystal. *J. Raman Spectrosc.* **1975**, *3* (2–3), 143–151.

- 551 <https://doi.org/10.1002/jrs.1250030205>.
- 552 (41) Zhang, Y. H.; Chan, C. K. Understanding the Hygroscopic Properties of Supersaturated
553 Droplets of Metal and Ammonium Sulfate Solutions Using Raman Spectroscopy. *J. Phys.*
554 *Chem. A* **2002**, *106* (2), 285–292. <https://doi.org/10.1021/jp012694j>.
- 555 (42) Agarwal, U. P.; Ralph, S. A.; Reiner, R. S.; Baez, C. Probing Crystallinity of Never-Dried
556 Wood Cellulose with Raman Spectroscopy. *Cellulose* **2016**, *23* (1), 125–144.
557 <https://doi.org/10.1007/s10570-015-0788-7>.
- 558 (43) Graf, D.; Molitor, F.; Ensslin, K.; Stampfer, C.; Jungen, A.; Hierold, C.; Wirtz, L. Raman
559 Imaging of Graphene. *Solid State Commun.* **2007**, *143* (1–2), 44–46.
560 <https://doi.org/10.1016/j.ssc.2007.01.050>.
- 561 (44) Mallamace, F.; Broccio, M.; Corsaro, C.; Faraone, A.; Majolino, D.; Venuti, V.; Liu, L.;
562 Mou, C. Y.; Chen, S. H. *Evidence of the Existence of the Low-Density Liquid Phase in*
563 *Supercooled, Confined Water*; 2007; Vol. 104. <https://doi.org/10.1073/pnas.0607138104>.
- 564 (45) Sun, Q.; Zheng, H. Raman OH Stretching Vibration of Ice Ih. *Prog. Nat. Sci.* **2009**, *19*
565 (11), 1651–1654. <https://doi.org/10.1016/j.pnsc.2009.06.010>.
- 566 (46) Sbroscia, M.; Sodo, A.; Bruni, F.; Corridoni, T.; Ricci, M. A. OH Stretching Dynamics in
567 Hydroxide Aqueous Solutions. *J. Phys. Chem. B* **2018**, *122* (14), 4077–4082.
568 <https://doi.org/10.1021/acs.jpcc.8b01094>.
- 569 (47) Ishizaka, S.; Wada, T.; Kitamura, N. In Situ Observations of Freezing Processes of Single
570 Micrometer-Sized Aqueous Ammonium Sulfate Droplets in Air. *Chem. Phys. Lett.* **2011**,

571 506 (1–3), 117–121. <https://doi.org/10.1016/j.cplett.2011.02.058>.

572 (48) Koop, T.; Ng, H. P.; Molina, L. T.; Molina, M. J. A New Optical Technique to Study
573 Aerosol Phase Transitions: The Nucleation of Ice from H₂SO₄ Aerosols. *J. Phys. Chem.*
574 *A* **1998**, *102* (45), 8924–8931. <https://doi.org/10.1021/jp9828078>.

575 (49) Bogdan, A. Double Freezing of (NH₄)₂SO₄/H₂O Droplets below the Eutectic Point and
576 the Crystallization of (NH₄)₂SO₄ to the Ferroelectric Phase. *J. Phys. Chem. A* **2010**, *114*
577 (37), 10135–10139. <https://doi.org/10.1021/jp105699s>.

578 (50) Chong, E.; King, M.; Marak, K. E.; Freedman, M. A. The Effect of Crystallinity and
579 Crystal Structure on the Immersion Freezing of Alumina. *J. Phys. Chem. A* **2019**, *123*
580 (12), 2447–2456. <https://doi.org/10.1021/acs.jpca.8b12258>.

581 (51) Watanabe, H.; Otsuka, T.; Harada, M.; Okada, T. Imbalance between Anion and Cation
582 Distribution at Ice Interface with Liquid Phase in Frozen Electrolyte as Evaluated by
583 Fluorometric Measurements of PH. *J. Phys. Chem. C* **2014**, *118* (29), 15723–15731.
584 <https://doi.org/10.1021/jp5031653>.

585 (52) Gerardo Wolfgang, G. Ion Distribution and Phase Boundary Potentials during the
586 Freezing of Very Dilute Ionic Solutions at Uniform Rates. *J. Colloid Interface Sci.* **1967**,
587 *25* (2), 270–279. [https://doi.org/10.1016/0021-9797\(67\)90030-6](https://doi.org/10.1016/0021-9797(67)90030-6).

588 (53) Nikam, P. S.; Aber, J. S.; Kharat, S. J. Viscosities of Ammonium Sulfate, Potassium
589 Sulfate, and Aluminum Sulfate in Water and Water + N,N-Dimethylformamide Mixtures
590 at Different Temperatures. *J. Chem. Eng. Data* **2008**, *53* (10), 2469–2472.
591 <https://doi.org/10.1021/je800330d>.

- 592 (54) Murray, B. J.; Bertram, A. K. Inhibition of Solute Crystallisation in Aqueous H⁺-NH₄⁺-
593 SO₄²⁻-H₂O Droplets. *Phys. Chem. Chem. Phys.* **2008**, *10* (22), 3287–3301.
594 <https://doi.org/10.1039/b802216j>.
- 595 (55) Lu, J. W.; Rickards, A. M. J.; Walker, J. S.; Knox, K. J.; Miles, R. E. H.; Reid, J. P.;
596 Signorell, R. Timescales of Water Transport in Viscous Aerosol: Measurements on Sub-
597 Micron Particles and Dependence on Conditioning History. *Phys. Chem. Chem. Phys.*
598 **2014**, *16* (21), 9819–9830. <https://doi.org/10.1039/c3cp54233e>.
- 599 (56) O’Meara, S.; Topping, D. O.; McFiggans, G. The Rate of Equilibration of Viscous
600 Aerosol Particles. *Atmos. Chem. Phys.* **2016**, *16* (8), 5299–5313.
601 <https://doi.org/10.5194/acp-16-5299-2016>.
- 602 (57) Sassen, K.; Zhao, H. Lidar Multiple Scattering in Water Droplet Clouds: Toward an
603 Improved Treatment. *Opt. Rev.* **1995**, *2* (5), 394–400. [https://doi.org/10.1007/s10043-995-](https://doi.org/10.1007/s10043-995-0394-2)
604 [0394-2](https://doi.org/10.1007/s10043-995-0394-2).
- 605 (58) Murphy, D. M.; Koop, T. Review of the Vapour Pressures of Ice and Supercooled Water
606 for Atmospheric Applications. *Q. J. R. Meteorol. Soc.* **2005**, *131* (608), 1539–1565.
607 <https://doi.org/10.1256/qj.04.94>.
- 608 (59) Larson, B. H.; Swanson, B. D. Experimental Investigation of the Homogeneous Freezing
609 of Aqueous Ammonium Sulfate Droplets. *J. Phys. Chem. A* **2006**, *110* (5), 1907–1916.
610 <https://doi.org/10.1021/jp0542701>.
- 611 (60) Mitchell, D. L. Parameterization of the Mie Extinction and Absorption Coefficients for
612 Water Clouds. *J. Atmos. Sci.* **2000**, *57* (9), 1311–1326. <https://doi.org/10.1175/1520->

- 613 0469(2000)057<1311:POTMEA>2.0.CO;2.
- 614 (61) Sassen, K.; Liou, K.-N. Scattering of Polarized Laser Light by Water Droplet, Mixed-
615 Phase and Ice Crystal Clouds. Part II: Angular Depolarizing and Multiple-Scattering
616 Behavior. *J. Atmos. Sci.* **1979**, *36* (5), 852–861. [https://doi.org/10.1175/1520-](https://doi.org/10.1175/1520-0469(1979)036<0852:sopllb>2.0.co;2)
617 [0469\(1979\)036<0852:sopllb>2.0.co;2](https://doi.org/10.1175/1520-0469(1979)036<0852:sopllb>2.0.co;2).
- 618 (62) Adriani, A.; Deshler, T.; Di Donfrancesco, G.; Gobbi, G. P. Polar Stratospheric Clouds
619 and Volcanic Aerosol during Spring 1992 over McMurdo Station, Antarctica: Lidar and
620 Particle Counter Comparisons. *J. Geophys. Res.* **1995**, *100* (D12).
621 <https://doi.org/10.1029/95jd02029>.
- 622 (63) Debenedetti, P. G.; Stillinger, F. H. *Supercooled Liquids and the Glass Transition*; 2001;
623 Vol. 410. <https://doi.org/10.1038/35065704>.
- 624 (64) Kissinger, J. A.; Buttke, L. G.; Vinokur, A. I.; Guzei, I. A.; Beyer, K. D. Solubilities and
625 Glass Formation in Aqueous Solutions of the Sodium Salts of Malonic Acid with and
626 Without Ammonium Sulfate. *J. Phys. Chem. A* **2016**, *120* (21), 3827–3834.
627 <https://doi.org/10.1021/acs.jpca.6b02656>.
- 628 (65) Wexler, A. S.; Clegg, S. L. Atmospheric Aerosol Models for Systems Including the Ions
629 H⁺, NH₄⁺, Na⁺, SO₄²⁻, NO₃⁻, Cl⁻, Br⁻, and H₂O. *J. Geophys. Res. Atmos.* **2002**, *107*
630 (14). <https://doi.org/10.1029/2001JD000451>.
- 631 (66) Campbell, J. M.; Meldrum, F. C.; Christenson, H. K. Is Ice Nucleation from Supercooled
632 Water Insensitive to Surface Roughness? *J. Phys. Chem. C* **2015**, *119* (2), 1164–1169.
633 <https://doi.org/10.1021/jp5113729>.

- 634 (67) Bi, Y.; Cao, B.; Li, T. Enhanced Heterogeneous Ice Nucleation by Special Surface
635 Geometry. *Nat. Commun.* **2017**, *8*. <https://doi.org/10.1038/ncomms15372>.
- 636 (68) Vali, G.; DeMott, P. J.; Möhler, O.; Whale, T. F. Technical Note: A Proposal for Ice
637 Nucleation Terminology. *Atmos. Chem. Phys.* **2015**, *15* (18), 10263–10270.
638 <https://doi.org/10.5194/acp-15-10263-2015>.
- 639 (69) Wahab, A.; Mahiuddin, S. Electrical Conductivity, Speeds of Sound, and Viscosity of
640 Aqueous Ammonium Nitrate Solutions. *Can. J. Chem.* **2001**, *79* (8), 1207–1212.
641 <https://doi.org/10.1139/cjc-79-8-1207>.
- 642 (70) Koop, T.; Luo, B.; Tsias, A.; Peter, T. Water Activity as the Determinant for
643 Homogeneous Ice Nucleation in Aqueous Solutions. *Nature* **2000**, *406* (6796), 611–614.
644 <https://doi.org/10.1038/35020537>.
- 645 (71) Sadeghi, R.; Jamehbozorg, B. Volumetric and Viscosity Studies of Interactions between
646 Sodium Phosphate Salts and Poly(Propylene Glycol) 400 in Aqueous Solutions at
647 Different Temperatures. *Fluid Phase Equilib.* **2009**, *284* (2), 86–98.
648 <https://doi.org/10.1016/j.fluid.2009.06.016>.
- 649 (72) Nikam, P. S.; Aber, J. S.; Kharat, S. J. Viscosities of Ammonium Sulfate, Potassium
650 Sulfate, and Aluminum Sulfate in Water and Water + N,N-Dimethylformamide Mixtures
651 at Different Temperatures. *J. Chem. Eng. Data* **2008**, *53* (10), 2469–2472.
652 <https://doi.org/10.1021/je800330d>.
- 653 (73) Deville, S. *Freezing Colloids: Observations, Principles, Control, and Use*; Engineering
654 Materials and Processes; Springer International Publishing: Cham, 2017.

655 <https://doi.org/10.1007/978-3-319-50515-2>.

656 (74) Richardson, H. H.; Hickman, Z. N.; Govorov, A. O.; Thomas, A. C.; Zhang, W.;
657 Kordesch, M. E. Thermo-optical Properties of Gold Nanoparticles Embedded in Ice:
658 Characterization of Heat Generation and Melting. *Nano Lett.* **2006**, *6* (4), 783–788.
659 <https://doi.org/10.1021/nl0601051>.

660 (75) Liu, B.; Liu, J. Freezing Directed Construction of Bio/Nano Interfaces: Reagentless
661 Conjugation, Denser Spherical Nucleic Acids, and Better Nanoflakes. *J. Am. Chem. Soc.*
662 **2017**, *139* (28), 9471–9474. <https://doi.org/10.1021/jacs.7b04885>.

663 (76) Al-Johani, H.; Abou-Hamad, E.; Jedidi, A.; Widdifield, C. M.; Viger-Gravel, J.; Sangaru,
664 S. S.; Gajan, D.; Anjum, D. H.; Ould-Chikh, S.; Hedhili, M. N.; Gurinov, A.; Kelly, M. J.;
665 Eter, M. El; Cavallo, L.; Emsley, L.; Basset, J. M. The Structure and Binding Mode of
666 Citrate in the Stabilization of Gold Nanoparticles. *Nat. Chem.* **2017**, *9* (9), 890–895.
667 <https://doi.org/10.1038/nchem.2752>.

668

669