

# Organosulfate Formation in Proxies for Aged Sea Spray Aerosol: Reactive Uptake of Isoprene Epoxydiols to Acidic Sodium Sulfate

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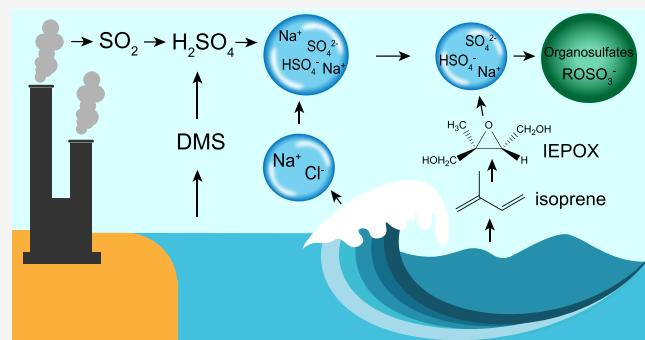
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**ABSTRACT:** Oxidation of isoprene, the biogenic volatile organic compound (BVOC) with the highest emissions globally, is a large source of secondary organic aerosol (SOA) in the atmosphere. Particulate organosulfates formed from acid-driven reactions of the oxidation products isoprene epoxydiol (IEPOX) isomers are important contributors to SOA mass. Most studies have focused on organosulfate formation on ammonium sulfate particles, often at low pH. However, recent work has shown that sea spray aerosol (SSA) in the accumulation mode ( $\sim 100$  nm) is quite acidic (pH  $\sim 2$ ) and undergoes further heterogeneous reactions with  $\text{H}_2\text{SO}_4$  to form  $\text{Na}_2\text{SO}_4$ . Herein, we demonstrate that substantial SOA, including organosulfates, are formed on acidic sodium sulfate particles ( $\text{pH} = 1.4 \pm 0.1$ ) via controlled laboratory experiments.

Comparable organosulfate formation was observed for acidic sodium and ammonium sulfate particles even though acidic particles with sodium versus ammonium as the primary cation formed less SOA volume. Both exhibited core-shell morphology after the reactive uptake of IEPOX; however, organosulfates were identified with Raman microspectroscopy in the core and shell of ammonium sulfate SOA particles, but only in the core for sodium sulfate SOA. Key organosulfates were also identified in ambient samples from the Galápagos Island. Our results suggest that isoprene-derived SOA formed on aged SSA is potentially an important, but underappreciated, source of SOA and organosulfates in marine and coastal regions that could modify SOA budgets.

**KEYWORDS:** marine secondary organic aerosol, heterogeneous chemistry, aerosol acidity, coastal aerosol chemistry, mass spectrometry, Raman microspectroscopy



## INTRODUCTION

Marine aerosol is the largest contributor to global aerosol budgets<sup>1</sup> and impacts Earth's radiative balance through both direct and indirect effects.<sup>2</sup> Understanding marine aerosol is important in both remote and coastal regions, where roughly 50% of the global population lives and where marine aerosol is an important contributor to particulate matter (PM) concentrations,<sup>3</sup> particularly for sub-micrometer particles.<sup>4,5</sup> This makes it critical to understand both primary aerosol emissions, such as sea spray aerosol (SSA),<sup>2</sup> and secondary aerosol that forms from oxidation of gaseous species in coastal and remote marine environments. Marine secondary aerosol includes inorganic sulfate from both natural sources (dimethyl sulfide, DMS, to  $\text{SO}_2$  oxidation) and anthropogenic sources (ships and  $\text{SO}_2$  emissions),<sup>6,7</sup> as well as organic species produced from the reactive uptake or condensation of biogenic volatile organic compound (BVOC)-derived oxidation products.<sup>8</sup>

Marine biota are well-established sources of isoprene,<sup>9–11</sup> with annual global oceanic fluxes of isoprene estimated to

range from 1 to 12 Tg<sup>12</sup> compared to an annual estimated total global flux of  $\sim 600$  Tg.<sup>13</sup> Isoprene-derived secondary organic aerosol (iSOA) is known to be an important contributor to total secondary organic aerosol (SOA) in urban and continental regions,<sup>14–16</sup> accounting for up to 41% of the organic mass fraction in the southeast United States during summer,<sup>17</sup> but its formation in coastal and remote marine regions is far less understood. A major oxidation pathway for isoprene leads to the formation of isomers of isoprene epoxydiols (IEPOX).<sup>18</sup> IEPOX readily undergoes acid-driven reactive uptake onto ammonium sulfate (AS) particles, forming organosulfate compounds,<sup>19</sup> with this

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reaction pathway representing up to ~30% of the particulate organosulfate mass in the eastern United States.<sup>20</sup> Organosulfate compounds have been identified in the atmosphere globally,<sup>14</sup> including in both marine<sup>5,21,22</sup> and continental<sup>23,24</sup> environments. However, most research on iSOA formation has focused on urban and continental environments,<sup>20</sup> while iSOA and organosulfate formation in marine environments remains much less understood.

Studies on the formation of iSOA have primarily focused on the reactive uptake of IEPOX onto ammonium sulfate salts,<sup>19,23,25–28</sup> which are often acidified to replicate urban aerosol pH.<sup>29</sup> In contrast, for marine environments, sodium is the dominant cation,<sup>30</sup> and sodium sulfate (SS) has long been observed to form by the reactive uptake of  $\text{H}_2\text{SO}_4$ <sup>31–34</sup> and displacement of HCl in numerous field studies,<sup>7,35–37</sup> including non-coastal cities (Atlanta)<sup>38</sup> and forested sites<sup>39</sup> hundreds of km inland (eq 1).



Recent measurements by Angle et al.<sup>40</sup> indicate that sub-micrometer marine aerosol particles are emitted with a pH of ~2.  $\text{H}_2\text{SO}_4$  uptake can decrease the pH of SSA particles, as shown in ambient measurements by Keene et al.<sup>41</sup> While Nguyen et al.<sup>42</sup> showed that neutral SS particles do not react with IEPOX to form SOA, to our knowledge, there have not been studies that investigated the SOA production from IEPOX uptake to acidic SS particles and subsequent organosulfate formation. Research on the formation of iSOA from the reactive uptake of IEPOX onto magnesium sulfate seed particles resulted in the enhanced formation of unsaturated oligomeric species,<sup>43</sup> indicating that the seed impacted the chemical composition of the resultant iSOA particles.

In this study, we investigate the potential for proxies of aged SSA ( $\text{Na}_2\text{SO}_4$  acidified with  $\text{H}_2\text{SO}_4$ ) to form iSOA. iSOA was generated on acidic sodium sulfate particles reacted with gaseous *trans*- $\beta$ -IEPOX, a predominant IEPOX isomer.<sup>44</sup> The physical and chemical transformations of the SS aerosol particles were characterized with a wide range of analytical instrumentation, including both online and offline measurements. The analytical techniques consist of a single-particle mass spectrometer (aerosol time-of-flight mass spectrometer, ATOFMS), an aerosol chemical speciation monitor (ACSM), a scanning electrical mobility spectrometer (SEMS), a Raman microspectrometer, and a particle-into-liquid sampler (PILS) whose samples were analyzed by both ion chromatography (IC) and hydrophilic interaction liquid chromatography interfaced with electrospray ionization high resolution quadrupole time-of-flight mass spectrometry (HILIC/ESI-HR-QTOFMS). We demonstrate that acidified SS seed particles are capable of forming significant amounts of iSOA and show that SS seed particles convert similar amounts of inorganic sulfates to organosulfates compared to AS-seeded iSOA particles. Furthermore, we show that SS-seeded iSOA particles result in comparable formation of methyltetrool sulfates (MTSs), which are the single most abundant organosulfates detected in ambient aerosols.<sup>20,45,46</sup> These results indicate that SSA, after conversion of  $\text{NaCl}$  to  $\text{Na}_2\text{SO}_4$ , could potentially be a significant source of iSOA formation in marine environments.

## MATERIALS AND METHODS

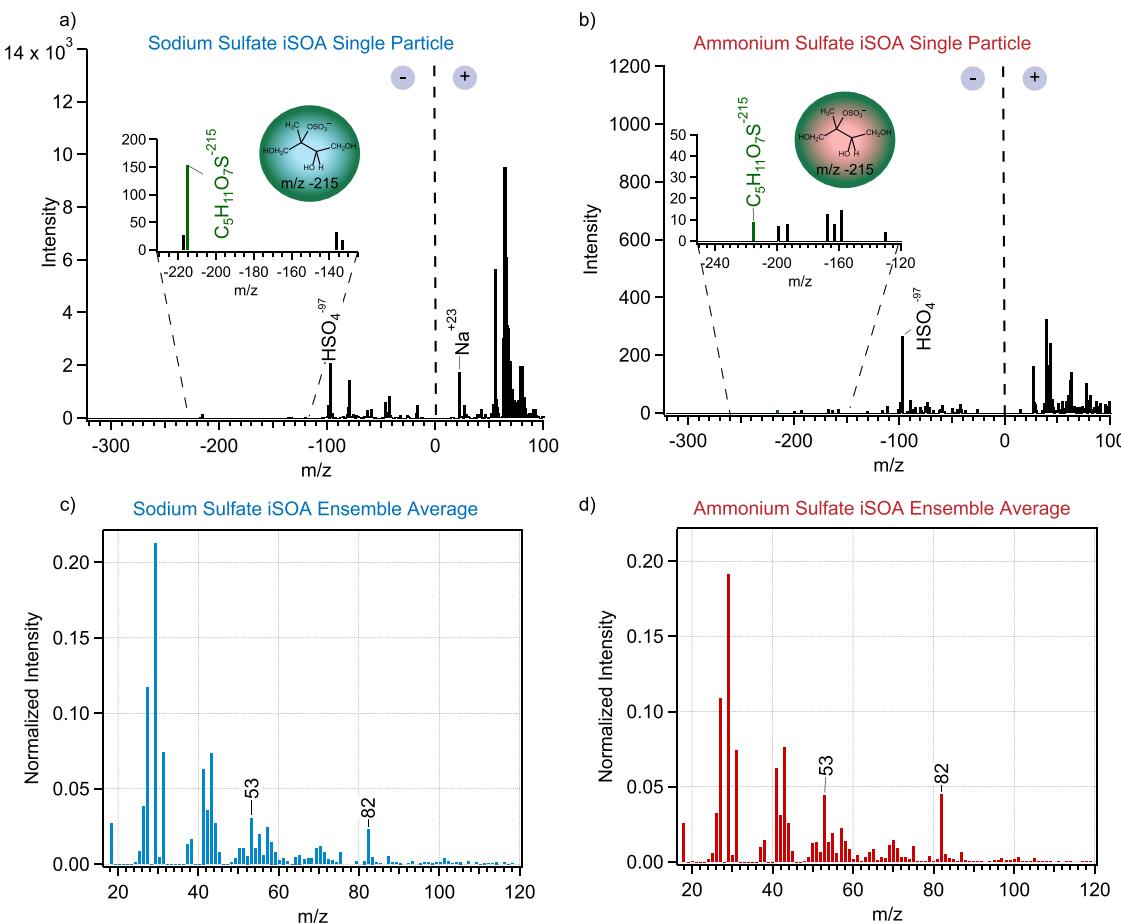
**Aerosol Generation.** SOA was generated from the reactive uptake of *trans*- $\beta$ -IEPOX, which was synthesized via the method described by Zhang et al.<sup>47</sup> The use of the University of North Carolina at Chapel Hill (UNC) 10 m<sup>3</sup> indoor chamber facility to study SOA has been described previously<sup>48</sup> and was operated at 50% relative humidity (RH) under dark conditions, as described by Zhang et al.<sup>27</sup> In short, aerosol particles were generated from solutions of SS (Sigma-Aldrich, ≥99% purity) or AS (Sigma-Aldrich, ≥99% purity) and 18.3 MΩ Milli-Q water (Table S1). The solutions were acidified by adding sulfuric acid (Sigma-Aldrich, ≥98% purity) and measured with a pH probe (Hanna Instruments). Although it has been previously shown that the aerosol pH is lower than its corresponding atomizing solution pH, the aerosol used in the present study approached the lower limit of detection for direct measurement with pH paper.<sup>49</sup> Thus, only the pH of the atomized solutions are reported here (Table S1). To improve run-to-run comparability, the total inorganic sulfate ( $\text{Sulf}_{\text{inorg}}$ ), including both  $\text{SO}_4^{2-}$  and  $\text{HSO}_4^-$ , in the seed aerosol was held constant at 0.12 M, based on prior work.<sup>50</sup> Acidified seed particles were generated from a custom-made atomizer<sup>28</sup> and injected into the chamber. After the seed aerosol concentration had stabilized, gaseous *trans*- $\beta$ -IEPOX was introduced to the chamber through a heated manifold (~60 °C) and injected for 60 min.

**Online Analysis of Aerosol Size and Composition.** A suite of analytical instruments were connected to the environmental chamber to measure real-time SOA formation, including an ATOFMS,<sup>51</sup> an ACSM (Q-ACSM, Aerodyne Research Inc.),<sup>52</sup> and a SEMS system (BGI Inc., Model 2100).

**Aerosol Time-of-Flight Mass Spectrometry.** ATOFMS was used to measure the size and chemical composition of individual particles, and the version of ATOFMS used in this study was previously described by Pratt et al.<sup>51</sup> In short, aerosol particles were introduced into a differentially pumped vacuum chamber and focused into a narrow particle beam by an aerodynamic lens. The particle beam first passes through two continuous wave lasers (405 and 488 nm, respectively) 6 inches apart, and the single particle speed is measured by calculating the time taken to travel between the two lasers. Vacuum aerodynamic diameter ( $d_{\text{va}}$ ) is calculated by generating a calibration curve from the particle speed of polystyrene latex sphere standards of known diameters (90 nm–1.5  $\mu\text{m}$ ). Based on their speed, particles are then desorbed and ionized via a Q-switched Nd:YAG laser (266 nm, fourth harmonic of a 1064 nm source), producing both positive and negative ions that are detected simultaneously via separate time-of-flight mass analyzers, providing size and dual-polarity time-of-flight mass spectra for individual particles.

**Aerosol Chemical Speciation Monitor.** The ACSM measures the time-resolved chemical composition of non-refractory sub-micrometer aerosol. As described in Ng et al.,<sup>53</sup> aerosol particles are focused into a narrow beam via an aerodynamic lens and impacted onto a hot vaporizer. The resulting vapor is ionized via electron ionization (70 eV) and analyzed via a quadrupole mass spectrometer operated in positive ion mode. ACSM calibration and tuning procedures were adopted from our prior studies.<sup>52,54</sup>

**Scanning Electrical Mobility Sizer.** The SEMS consists of a differential mobility analyzer (DMA, BGI model 2002)



**Figure 1.** Single-particle mass spectra collected using the ATOFMS of SS (a) and AS (b) particles after reaction with IEPOX in an atmospheric chamber experiment. Green peak represents the monoisotopic mass of IEPOX-derived methyltetrox sulfates (an abundant organosulfate). Inset represents a zoom-in of the mass spectral region with iSOA products. ACSM organic spectra averaged for 20 min post IEPOX injection with SS (c) and AS (d) seed particles. Peaks labeled represent relevant organosulfate species. Spectra were normalized by the sum of organic species  $m/z$  18–120.

coupled to a mixing condensation particle counter (MCPC, BMI model 1710). The SEMS measures aerosol number, surface area, and volume concentrations within the chamber at a time resolution of  $\sim 1$  min.

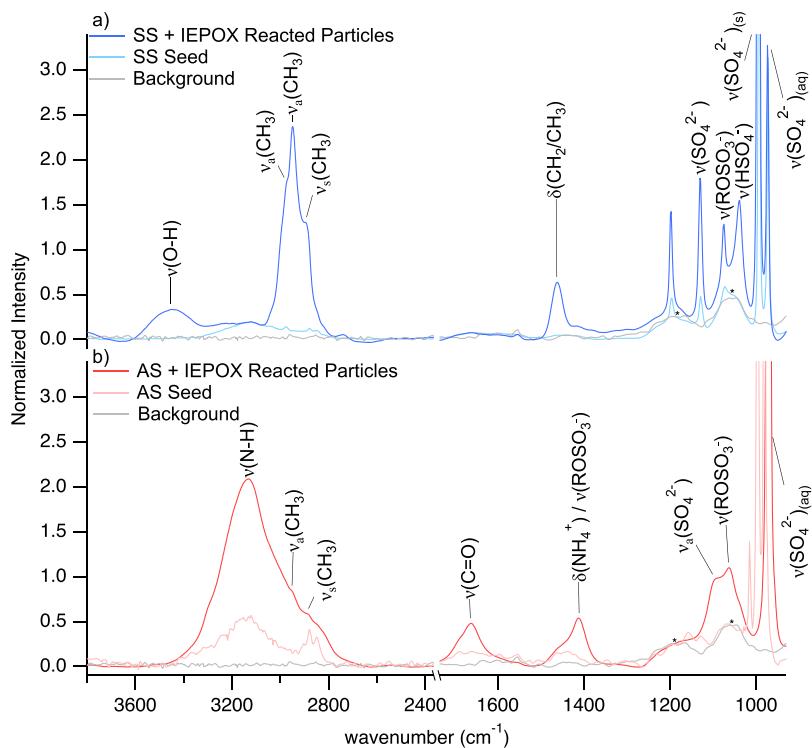
**Offline Analysis. Laboratory Particle Collection.** Particles were impacted onto quartz substrates (Ted Pella Inc.) using a multi-orifice particle sizer (MPS-3, California Measurements Inc.). Particles collected on stages 2 (0.4–2.5  $\mu\text{m}$ ) and 3 ( $<0.4\text{ }\mu\text{m}$ ) were the focus of our analysis as they are closest to the peak of the particle number size distribution from the atmospheric chamber.<sup>55</sup> Additionally, SOA were collected by a PILS system (BMI model 4001) for offline inorganic sulfate quantification via IC, as described previously (Experiments 1 and 2,<sup>56,57</sup> Experiments 3 and 4<sup>58</sup>), and organosulfate quantification via hydrophilic interaction liquid chromatography coupled with electrospray ionization high-resolution quadrupole time of flight mass spectrometry (HILIC/ESI-HR-QTOFMS), as described previously by Cui et al.<sup>59</sup> The PILS vials were stored in the dark at 2 °C immediately after collection and analyzed without further pretreatment except for dilution by acetonitrile for the HILIC/ESI-HR-QTOFMS method described below.

**Field Aerosol Collection and Filter Analysis.** Ambient samples were collected from the Galápagos Science Center on San Cristobal Island in Puerto Baquerizo Moreno, Ecuador

(0.8957°S, 89.6086°W, 15 masl), with a high-volume PM<sub>2.5</sub> air sampler (TE-6070 V-BL, Tisch Environmental) onto prebaked quartz filters. Sampling and extraction conditions were as described previously in Cui et al.<sup>58</sup> The chemical composition of the extracts were analyzed by HILIC/ESI-HR-QTOFMS as described below.

**Raman Microspectroscopy.** Individual particles collected onto quartz substrates during the laboratory studies were analyzed via Raman microspectroscopy with a Horiba LabRAM HR Evolution Raman Spectrometer (Horiba Scientific) coupled to a confocal optical microscope (Olympus, 100  $\times$  0.9 N.A. objective) at ambient temperature and RH. The instrument is equipped with a ND:YAG laser source (50 mW, 532 nm) and a CCD detector. Spectra for each particle were collected in the range of 500–4000  $\text{cm}^{-1}$  for three accumulations of 15 s. An 600 groove/mm diffraction grating with a spectral resolution of  $\sim 1.7\text{ }\text{cm}^{-1}$  was used.

**Ion Chromatography.** IC was used to quantify the inorganic sulfate concentration in the SOA particles from the chamber. For Experiments 1 and 2, a 500  $\mu\text{L}$  aliquot was analyzed by an anion exchange IC (Dionex ICS-2100) equipped with a guard column (Dionex IonPac AG18, 4 mm  $\times$  50 mm) and an anion exchange column (4 mm  $\times$  250 mm, Thermo Fisher) run at a flow rate of 1.0  $\text{L min}^{-1}$ . For Experiments 3 and 4, a 25  $\mu\text{L}$  aliquot of each PILS aqueous



**Figure 2.** Raman spectra of SS (a) and AS (b) seed particles compared to the spectra of seed particles after reaction with IEPOX. For SS + IEPOX, 98 particles were analyzed and averaged; for the well-characterized AS + IEPOX, 82 particles were analyzed and averaged. Intensity for all spectra is normalized to  $792\text{ cm}^{-1}$ , an intense peak associated with the quartz substrate. Asterisks correspond to other peaks associated with quartz substrate.

sample was analyzed by an anion exchange IC (ICS 3000, ThermoFisher). The instrument is equipped with an IonPac AS11-HC guard column (2 mm  $\times$  50 mm ThermoFisher) and an anion-exchange column (2 mm  $\times$  250 mm ThermoFisher) run at a flow rate of 0.4 L  $\text{min}^{-1}$ .

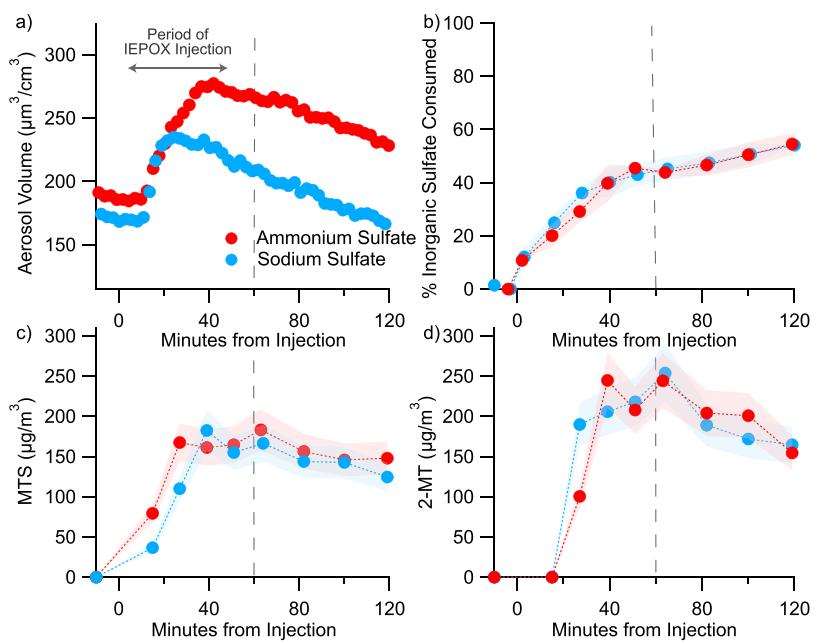
**Hydrophilic Interaction Liquid Chromatography Interfaced with Electrospray Ionization High-Resolution Quadrupole Time-of-Flight Mass Spectrometry.** HILIC/ESI-HR-QTOFMS was used to quantify the concentration of methyltetrol sulfates (MTSs) in the iSOA particles, the predominant IEPOX-derived organosulfate compounds,<sup>45,58,60,61</sup> and 2-methyltetrols (2-MTs) from the laboratory and field. As previously described by Chen et al.,<sup>58</sup> a 50  $\mu\text{L}$  aliquot of each PILS aqueous sample was withdrawn and diluted in 950  $\mu\text{L}$  of acetonitrile (ACN, HPLC-grade, Fisher Scientific) shortly after collection in order to achieve the solvent composition of the organic mobile phase (i.e., 95:5 ACN:H<sub>2</sub>O). Diluted PILS samples were stored in the dark at  $-20\text{ }^\circ\text{C}$  prior to HILIC/ESI-HR-QTOFMS analysis. Analysis was performed following the method of Chen et al.<sup>58</sup> using an Agilent 6520 Series Accurate Mass Q-TOFMS with an ESI source operated in negative mode coupled to an Agilent 6500 Series UPLC equipped with a Waters ACQUITY UPLC BEH Amide column following the published method.<sup>58</sup> The operating conditions of the HILIC/ESI-QTOFMS system (elution gradient program, mass calibration, tuning, voltages, etc.) were previously described by Chen et al.<sup>58</sup>

## RESULTS AND DISCUSSION

**Identification of Organosulfate Formation.** Atmospheric chamber experiments were conducted with both acidified SS ( $\text{pH } 1.4 \pm 0.1$ ) and acidified AS ( $\text{pH } 1.4 \pm 0.1$ )

seed particles and reacted with IEPOX. The chemical composition of the reacted particles was monitored at the single-particle level with ATOFMS and at the bulk aerosol level with an ACSM. IEPOX-derived MTS isomers have been previously observed in ambient particles from field measurements using the ATOFMS as deprotonated molecules ( $\text{C}_5\text{H}_{11}\text{O}_7\text{S}^-$ ) at a mass-to-charge ratio ( $m/z$ ) of 215.<sup>38,62</sup> ATOFMS spectra from organosulfate standards including 3-methyltetrol organosulfates, sodium ethyl sulfate, and hydroxyacetone sulfate were compared (Figure S1). A negative ion peak at  $m/z$  215 was only observed in the spectrum for 3-methyltetrol organosulfates, whereas no peak was observed at  $m/z$  215 in the spectra for the remaining standards, providing further evidence that this peak corresponds to MTS isomers. A negative ion peak at  $m/z$  215 was observed in the ATOFMS spectra of the reacted particles for both AS and SS seed aerosol, indicating that MTS isomers can form from the reaction of IEPOX with sulfate on the acidified SS seed particles. Example ATOFMS spectra of individual particles are shown for acidic SS + IEPOX and acidic AS + IEPOX-reacted particles (Figure 1a,b). We observed particle-to-particle variability in the collected ATOFMS spectra of SS + IEPOX reacted particles, which can be noted in the lower  $m/z$  region of the SS + IEPOX-reacted particles (Figure S2). In addition to MTS isomers observed as deprotonated molecules at  $m/z$  215, the SS-reacted particles contained other isoprene-derived organic compounds previously identified in ambient aerosol: glycolic acid sulfates with deprotonated molecules at  $m/z$  155 and 2-MTs with deprotonated molecules at  $m/z$  135 (Figure S2).<sup>62</sup>

Bulk aerosol average mass spectra from the ACSM demonstrate that organic aerosol forms from the reaction of IEPOX with acidified SS particles (Figure 1c). The average



**Figure 3.** Change in aerosol volume, inorganic sulfate, methyltetrol sulfates (MTS), and 2-methyltetrols (2-MT) for SS (blue) and AS (red) seed particles at  $\text{pH} = 1.4 \pm 0.1$  reacted with IEPOX in an atmospheric chamber. (a) Change in aerosol volume for SS-seeded and AS-seeded iSOA particles. (b) Change in the percentage of inorganic sulfate consumed. Shaded region represents standard deviation for IC analysis. (c) Change in the mass of MTSs. Shaded region represents % relative error for HILIC/ESI-HR-QTOFMS analysis. (d) Change in the mass of 2-MTs. Shaded region represents % relative error for HILIC/ESI-HR-QTOFMS analysis.

organic ACSM mass spectra of SOA produced following 60 min of IEPOX injection were compared for the AS and SS seed particles (Figure 1c,d). The AS and SS experiments produce similar SOA spectra, with both containing mass spectral fragments at  $m/z$  53 and 82,<sup>48</sup> characteristic ACSM mass fragment ions for IEPOX-derived organic compounds.<sup>17,52</sup> The combined ACSM and ATOFMS results support the conclusion that organic aerosol forms from the reaction of IEPOX with acidified SS aerosol via single-particle and bulk analysis.

**Impact of Seed Aerosol Type on SOA Chemical Composition and Morphology.** To further differentiate inorganic and organic sulfate within individual particles, Raman spectra were collected for the SS and AS seed particles and compared to the particles after reaction with IEPOX in the atmospheric chamber (Figure 2 and Table S2). Inorganic sulfate (aqueous) is discernable and located in the seed and reacted particles for both SS and AS particles at 974 and 972  $\text{cm}^{-1}$ , respectively, in the form of the sulfate symmetric stretching mode,  $\nu_s(\text{SO}_4^{2-})$ ,<sup>63</sup> and in the SS-reacted aerosol particles at 1039 in the form of bisulfate  $\nu(\text{HSO}_4^-)$ .<sup>63</sup> For both seed aerosol types, the spectra show the formation of organic compounds observed via the methyl asymmetric stretching mode in the C–H stretching region (2950, 2974, and 2952  $\text{cm}^{-1}$ ),<sup>64</sup> methyl symmetric stretching mode (2897 and 2892  $\text{cm}^{-1}$ ),<sup>64</sup> and additional organic modes in the fingerprint region (1663, 1462, 1410, 1075, and 1063  $\text{cm}^{-1}$ ).<sup>65–69</sup> Peaks at 1075 and 1063  $\text{cm}^{-1}$  in the spectra of the SS- and AS-reacted particles, respectively, correspond to IEPOX-derived organosulfates.<sup>65,70</sup>

The spectra also indicate chemical differences between the SS-reacted aerosol particles and AS-reacted aerosol particles. In the AS-reacted aerosol particles, we observed a peak at 1410  $\text{cm}^{-1}$ , potentially corresponding to the ammonium bend,<sup>69,70</sup> though neutral organosulfate compounds or other organic modes may also contribute.<sup>65,69,70</sup> In the AS-seeded iSOA, we also observed a peak corresponding to carbonyls at 1663

$\text{cm}^{-1}$ .<sup>71</sup> This peak was not observed in the SS particles, suggesting that different products may form in SS-seeded iSOA particles. We also observed a distinct peak in the fingerprint region for the SS particles corresponding to the  $\text{CH}_2/\text{CH}_3$  bend at 1462  $\text{cm}^{-1}$ .<sup>65</sup> Comparing the Raman spectra of the reacted to the seed particles for acidified SS further supports our conclusion that organic aerosol forms from the reaction of IEPOX with SS seed aerosol.

**Impact of Seed Aerosol Type on SOA Mass and Organosulfate Formation.** To investigate how the main cation ( $\text{Na}^+$  vs  $\text{NH}_4^+$ ) in the seed aerosol impacts the formation of organosulfate compounds, SEMS was used to determine aerosol volume formation, IC was used to quantify the percentage of inorganic sulfate consumed, and HILIC/ESI-HR-QTOFMS was used to quantify MTSs and 2-MTs formed in the iSOA during the atmospheric chamber experiments for SS and AS (Figure 3). After 60 min of reaction in the chamber, the volume of SOA formation was greater from the reaction of acidic AS with IEPOX than that from the reaction of acidic SS with IEPOX. Despite the difference in aerosol volume growth, the percentage of inorganic sulfate consumption at 60 min was similar for AS ( $44\% \pm 4$ ) and SS ( $45\% \pm 4$ ) at 60 min, indicating that a comparable amount of organosulfates is formed for both seed types. We also investigated the mass of key iSOA products MTSs and 2-MTs at 60 min. The mass of MTSs formed at 60 min from the reaction of acidic AS with IEPOX ( $180 \pm 30 \mu\text{g}/\text{m}^3$ ) was comparable to the amount formed from the reaction of acidic SS with IEPOX ( $170 \pm 20 \mu\text{g}/\text{m}^3$ ) as well as the mass of 2-MTs formed at 60 min for AS SOA ( $240 \pm 30 \mu\text{g}/\text{m}^3$ ) and SS SOA ( $250 \pm 30 \mu\text{g}/\text{m}^3$ ).

We compared the time-resolved volume, inorganic sulfate consumption, MTS mass, and 2-MT mass for SS and AS seed aerosol types (Figure 3). SEMS results indicate that AS-seeded iSOA particle volume growth is greater than iSOA formed from SS seed aerosol particles (Figure 3a). The volume growth

for the SS and AS experiments were similar until  $\sim$ 20 min after the reaction began, when SS-seeded iSOA particles started to level off and AS-seeded iSOA particles continued to increase. For iSOA from both AS and SS seed aerosols, the amount of inorganic sulfate began decreasing at the beginning of IEPOX injection, indicating that the inorganic sulfate was consumed and converted to organosulfates immediately after reaction with gaseous IEPOX. After 60 min, the injection of IEPOX was stopped and the particles were allowed to react in the chamber for another 60 min. For both the SS- and AS-seeded iSOA particles, the inorganic sulfate continued to be consumed, indicating that further chemical reactions involving sulfate incorporation into organic species were occurring. However, the decrease occurred at a slower rate than during the period of IEPOX injection.

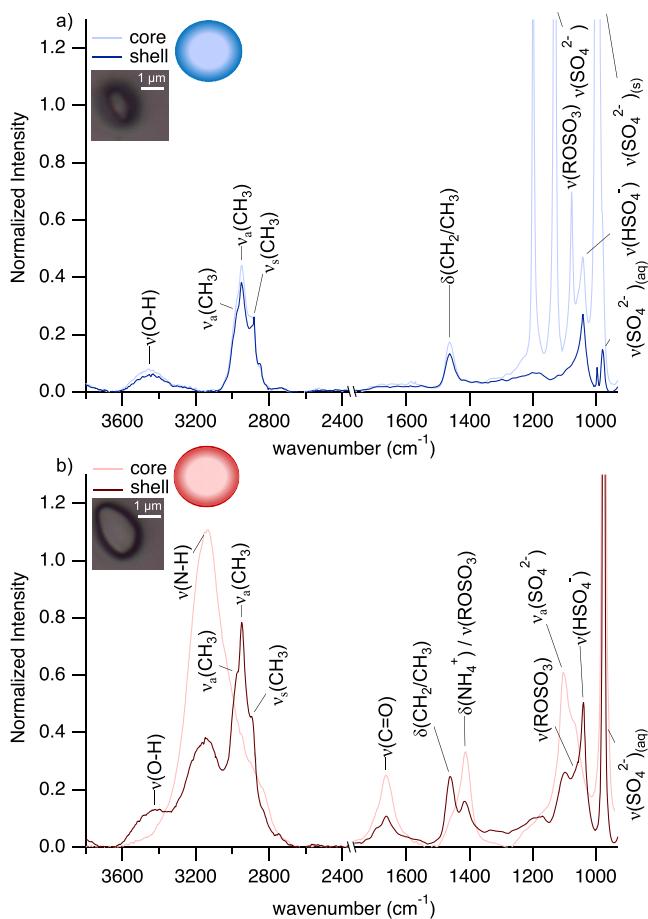
Similar to inorganic sulfate consumption, the formation of MTS and 2-MT products were comparable for the two seed types. The mass of MTSs increased until about 20 min into the reaction, around the same time that the volume formation leveled off in the SS-iSOA and the AS-iSOA continued to increase. After  $\sim$ 30 min of reaction, the MTS mass stopped increasing and remained stable for the experiment for both inorganic seed particle types. The 2-MT mass increased in both the AS and SS iSOA until  $\sim$ 40 min of reaction, after which the mass of 2-MTs in the iSOA slowly decreased in the particles, potentially due to wall loss, for the remainder of the experiment.

The experiments for both particle types were repeated, and a similar trend was observed with increased volume formation in the AS-seeded iSOA compared to the SS-seeded iSOA (Figure S3). The particle size distributions were compared for the SS and AS seed aerosol particles reacted with IEPOX at 0 min (unreacted), 30 min reaction, and 60 min reaction (Figure S4). For both the SS and AS experiments, the particle size modes were similar. Slightly greater MTS and 2-MT masses of formation were observed in the repeats that can be attributed to a greater mass of IEPOX injected during the AS repeat experiment (3.9 mg) compared to the SS (2.9 mg), though other factors, including partitioning and formation of other products, could play a role.

HILIC/ESI-HR-QTOFMS was used to investigate the formation of other known organosulfate products in the iSOA. Similar product formation was observed in the iSOA formed from the two particle types. Future studies will be conducted to compare the product distribution in iSOA-formed IEPOX with acidic SS and acidic AS.

**Particle Morphology.** To investigate the potential for SS seed particles to form a core-shell morphology, potentially resulting in the “self-limiting” effect as described by Zhang et al.,<sup>72</sup> single-particle morphology and chemical composition was investigated using Raman microspectroscopy for SS- and AS-seeded iSOA particles. Both SS- and AS-derived iSOA particles formed core-shell morphologies after reacting with IEPOX (Figure 4 and Table S3), as observed by the optical image of each particle as well as the chemical differences obtained from the spectra taken at the center (core) and edge of the particle (shell). Core and shell spectra for particles (AS,  $n = 18$ ; SS  $n = 25$ ) were compared to assess reproducibility (Figures S5 and S6).

For SS-reacted particles, an organic peak at  $1077\text{ cm}^{-1}$  was observed in the core, corresponding to organosulfates.<sup>70</sup> Inorganic sulfate was observed in the core and the shell in the form of aqueous sulfate at  $997\text{ cm}^{-1}$ ,<sup>64</sup> solid sodium sulfate



**Figure 4.** Representative single-particle Raman spectrum for the core and shell of (a) SS seed particles + IEPOX and (b) AS seed particles + IEPOX. Inset shows 100 $\times$  optical microscope image of aerosol particle corresponding to the spectrum.

at  $996\text{ cm}^{-1}$ ,<sup>64</sup> and bisulfate at  $1042\text{ cm}^{-1}$  as identified by Chen et al.<sup>63</sup> The core and shell also contained peaks at  $1462\text{ cm}^{-1}$  corresponding to bending modes for  $\text{CH}_2/\text{CH}_3$  groups and three peaks at  $2885$ ,  $2951$ , and  $2974\text{ cm}^{-1}$  corresponding to  $\nu_s(\text{CH}_3)$ ,<sup>64</sup>  $\nu_a(\text{CH}_3)$ ,<sup>64</sup> and  $\nu_a(\text{CH}_3)$ ,<sup>65</sup> respectively. Interestingly, the same peak shape with  $\nu_s(\text{CH}_3)$  (2896) and  $\nu_a(\text{CH}_3)$ <sup>64,65</sup> (2950 and 2974) was observed in the shell of the AS particles as well, indicating that the shell of the SS and AS particles likely has a similar organic composition. The shell of the SS particles did not contain a peak representative of organosulfates, indicating that the shell is likely composed of other non-sulfated organic compounds. The core and shell spectra were compared for  $n = 18$  particles (Figure S5), and little particle-to-particle variability was noted. The peaks discussed above were observed in all of the spectra, but there was some variability in the relative intensity of the peaks. These results indicate that, like AS-seeded iSOA particles, both the core and shell of the SS-reacted particles contain organic compounds, which could result in a self-limiting effect in the reactive uptake of IEPOX onto acidic SS particles.<sup>27</sup>

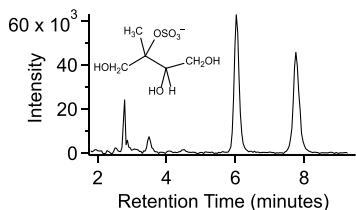
The core and shell for AS-reacted particles were compared as well, and two particle types were observed. The first particle type resembled those previously observed by Olson et al.,<sup>73</sup> with a chemically distinct core and shell. The second particle type was homogeneous with no apparent differences between the core and shell. These particles had a notable absence of a

peak associated with the N–H stretch, indicating that the ammonium may have partitioned out of the particles as ammonia after reaction with IEPOX. Six examples of each particle type are shown (Figure S6), and a representative particle for the core-shell morphology is provided (Figure 4).

In the AS-seeded core-shell particles, an organosulfate mode was identified in the core and shell particles at  $1065\text{ cm}^{-1}$ , agreeing with the results from Olson et al.<sup>73</sup> The  $\nu_a(\text{CH}_3)$  stretching modes at  $2974$  and  $2950\text{ cm}^{-1}$  as well as the  $\nu_s(\text{CH}_3)$  stretching mode at  $2896\text{ cm}^{-1}$  were observed only in the shell of the AS particles, whereas the carbonyl stretch at  $1666\text{ cm}^{-1}$  and the  $\text{CH}_2/\text{CH}_3$  bending modes at  $1461\text{ cm}^{-1}$  were observed in both the core and the shell. These results suggest that the AS-reacted seed particles result in the formation of iSOA particles with different morphologies than those formed from SS seed particles.

The SS- and AS-iSOA particles were reanalyzed 1 month after the experiment was conducted, and it was found that the chemical compositions of both particle types had changed (Figure S7). Ammonium was observed in the SS-iSOA, suggesting that ammonia partitioned into the iSOA particles during the 1 month of storage in the laboratory. Additionally, ammonium was present in both the core and shell of the AS-iSOA, whereas it was only observed in the core of the particles analyzed shortly after the experiment was conducted. These results indicate the potential for both SS-iSOA and AS-iSOA to undergo further reactions after formation in the atmosphere that could alter the particle morphology, as well as chemical composition.

**Organosulfate Identification in an Ambient Marine Environment.** To assess the potential for iSOA and organosulfates to form on SS in an ambient marine-influenced atmosphere, filter samples were collected from the Galápagos Islands of Ecuador. The aerosol composition was analyzed using the HILIC/ESI-HR-QTOFMS method. Methyltetrol sulfates were observed in the extracted ion chromatogram (EIC) for the negative ion at  $m/z 215.0231$  (Figure 5). As



**Figure 5.** HILIC/ESI-HR-QTOFMS extracted ion chromatogram (EIC) for methyltetrol sulfates at  $m/z -215.0231$  in a PM2.5 sample collected from the Galápagos Islands, Ecuador.

Galápagos Islands are a remote environment with primarily marine and, to a lesser extent, terrestrial influences, these organosulfates could be forming from the reaction of IEPOX with smaller SSA, which were recently shown to be very acidic ( $\text{pH} \approx 2$ ).<sup>40</sup> The same number of peaks for the MTS isomers, corresponding to the structural isomers for 2-MTS and 3-MTS as previously shown,<sup>59</sup> and their relative positions (2.7, 3.5, 6.0, and 7.8 min) were observed in the aerosol sample collected from the Galápagos Islands (Figure 5) as were observed in prior HILIC/ESI-HR-QTOFMS data reported by Cui et al.<sup>59</sup> While the formation of iSOA has been primarily investigated in continental environments, the presence of MTSs from a remote environment with marine influences demonstrates that

iSOA from the reaction of IEPOX with aged SSA is an unrecognized and likely important source of MTSs.

**Atmospheric Implications.** SSA is the largest source of aerosol emissions to the atmosphere,<sup>1</sup> and freshly emitted submicron SSA were recently shown to have a pH of  $\sim 2$ .<sup>40,74</sup> SSA particles intrinsically contain SS salts, and the amount of sulfate increases substantially after heterogeneous reactions with  $\text{H}_2\text{SO}_4$ .<sup>35</sup> Thus, it is important to consider the potential of these particles to seed the formation of iSOA. This study demonstrates that acidified SS particles ( $\text{pH } 1.4 \pm 0.1$ ) are capable of seeding the formation of iSOA, resulting in particles with similar amounts of inorganic sulfate incorporated into organosulfates and formation of similar amounts of MTSs when compared to the more-studied iSOA particles formed on AS seed aerosol. As such, SS particles in SSA should be considered when modeling iSOA formation in marine environments in order to accurately predict their formation and impact on the climate. More so, ambient measurements indicate that marine aerosol can undergo long-range atmospheric transport,<sup>3,39,75</sup> and thus, may be a source of iSOA in continental regions in addition to marine environments. The experiments conducted in this study isolated the initial reactive uptake of IEPOX onto particulate sulfate but did not account for additional chemical processes that could occur in ambient marine and coastal environments. Further studies should investigate the potential of SS- and AS-seeded iSOA to undergo photolysis reactions, heterogeneous reactions with other organic gases beyond IEPOX, or further aging via oxidation with OH, which could impact the chemical compositions and physical properties of these particles. Further research is needed on the potential of SSA to seed the formation of organosulfate compounds globally, particularly as a function of aerosol acidity, and their impact on climate.

## ASSOCIATED CONTENT

### Supporting Information

The Supporting Information is available free of charge at <https://pubs.acs.org/doi/10.1021/acsearthspacechem.2c00156>.

(Table S1) Additional information including atmospheric chamber experimental conditions; (Figure S1) single-particle mass spectra for standards: methyltetrol sulfates, sodium ethyl sulfate, and hydroxyacetone sulfate; (Figure S2) four single-particle mass spectra for SS seed particles reacted with IEPOX; (Table S2) experimentally determined Raman modes and tentative assignments for SS and AS seed particles before and after reaction with IEPOX; (Figure S3) change in aerosol volume, inorganic sulfate consumption, MTS, and 2-MT formation for additional experiments; (Figure S4) size distributions for SS- and AS-derived iSOA; (Table S3) experimentally determined Raman modes and tentative assignments for the core (center) and shell (edge) of SS and AS seed particles reacted with IEPOX; (Figure S5) core and shell Raman spectra for SS iSOA; (Figure S6) core and shell Raman spectra for AS iSOA; (Figure S7) core and shell Raman spectra for SS- and AS-derived iSOA after one month of storage (PDF)

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The authors declare no competing financial interest.

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