

NO at low concentration can enhance the formation of highly oxygenated biogenic molecules in the atmosphere

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The interaction between nitrogen monoxide (NO) and organic peroxy radicals (RO₂) greatly impacts the formation of highly oxygenated organic molecules (HOM), the key precursors of secondary organic aerosols. It has been thought that HOM production can be significantly suppressed by NO even at low concentrations. Here, we perform dedicated experiments focusing on HOM formation from monoterpenes at low NO concentrations (0 – 82 pptv). We demonstrate that such low NO can enhance HOM production by modulating the RO₂ loss and favoring the formation of alkoxy radicals that can continue to autoxidize through isomerization. These insights suggest that HOM yields from typical boreal forest emissions can vary between 2.5%–6.5%, and HOM formation will not be completely inhibited even at high NO concentrations. Our findings challenge the notion that NO monotonically reduces HOM yields by extending the knowledge of RO₂–NO interactions to the low-NO regime. This represents a major advance towards an accurate assessment of HOM budgets, especially in low-NO environments, which prevails in the pre-industrial atmosphere, pristine areas, and the upper boundary layer.

Atmospheric aerosols are crucial to Earth's radiative forcing and climate, yet their influence remains poorly quantified¹. Globally, organic aerosols contribute significantly to the total aerosol mass^{2–5}. Oxidation pathways of various volatile organic compounds (VOCs) involves autoxidation of peroxy (RO₂) radicals⁶. The products, known as highly oxygenated organic molecules (HOM)⁷, have recently been observed in various environments in the atmosphere. Owing to their high oxidation state and low volatility, HOM are a major source of secondary organic aerosol (SOA)⁸ and contribute significantly to new particle formation (NPF)^{9–11}.

NO_x concentrations are substantially elevated in the present-day atmosphere due to anthropogenic emissions. NO_x profoundly influences HOM formation. First, it can regulate the atmospheric oxidation capacity and consequently affect the oxidation of VOCs^{12–15}. Second, NO_x can greatly influence the extent of RO₂ radical autoxidation¹⁶ and thus HOM composition by directly reacting with RO₂ radicals, leading to enhanced formation of organic nitrates and suppressed formation of HOM dimers^{6,17–25}. Up to now, the second

effect remains poorly quantified, presenting an important obstacle towards a complete understanding of HOM budget and its impacts on aerosol formation in the atmosphere. Currently, most laboratory experiments have been conducted either at effectively zero NO_x (with very high VOC: NO_x), or at relatively high NO_x concentrations. It is widely accepted that high NO_x suppresses RO₂ autoxidation and HOM formation¹⁶. Yet, little is known about HOM formation at low but non-zero NO_x concentrations (i.e., NO ranges from 5–100 pptv and NO₂ ranges from 0.1–5 ppbv), characteristic of the pre-industrial atmosphere, the pristine environments, the upper boundary layer, and likely the future atmosphere if NO_x concentration continuously declines. Because the interactions between RO₂ and NO_x are so complex (involving the different roles of NO, NO₂, and NO₃), the overall influence on HOM formation can be highly nonlinear. Therefore, simply extrapolating the results derived from high-concentration experiments to low-concentration conditions, or even interpolating between high-NO_x and effectively zero NO_x, may lead to substantial biases^{26,27}.

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In this work, we perform dedicated experiments at the CERN CLOUD (Cosmics Leaving Outdoor Droplets) chamber with molecular-level observations and simulations focusing on the HOM formation at low NO_x concentrations²⁸. We demonstrate that the role of NO_x in HOM formation is more nuanced than simple suppression, and low NO concentrations can cause a previously overlooked increase in HOM yields.

Results and discussion

An overview of experimental observations

In the experiments, we kept the chamber conditions stable at a temperature of 278 K and 38% relative humidity. A 2:1 mixture of α -pinene and Δ -3-carene comprised the VOC precursors, representative of the

typical monoterpene profile above the canopy in a boreal forest in southern Finland (the SMEAR II station, the reference site for this study)^{29,30}. In different experiments, the steady-state mixing ratio of total monoterpenes was approximately 300, 600, and 1200 pptv. We controlled NO_x in three distinct sequences: experiments with only NO_2 to represent the nighttime condition (Fig. 1a); experiments with variable NO_x concentrations with a constant NO to NO_2 ratio (~ 0.007) to represent the morning (Fig. 1b, c); and experiments at constant NO_x with variable NO to NO_2 ratios to represent the evolution from night to noon (Fig. 1d). These experiments are referred to as “pure NO_2 ”, “constant NO/NO_2 ”, and “variable NO/NO_2 ” hereafter in this work. The highest NO_2 and NO mixing ratios in these experiments were 4.6 ppbv and 82 pptv, respectively. A vital feature of these experiments is the

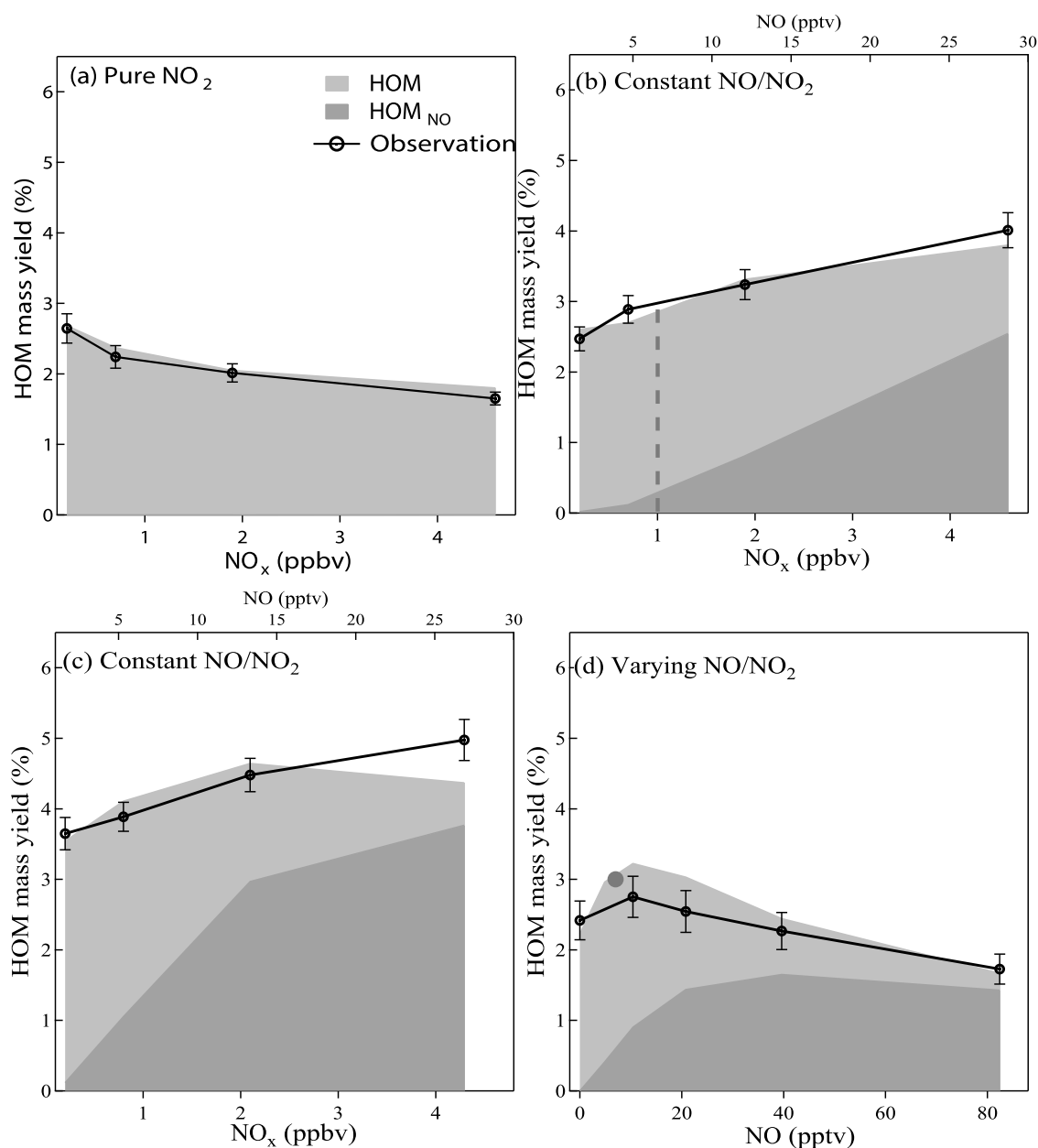


Fig. 1 | Comparison of measured and modelled mass yields of highly oxygenated organic molecules (HOM). In (a) a pure NO_2 experiment run with 1200 pptv monoterpene, (b) a constant NO/NO_2 experiment with 1200 pptv monoterpene, (c) a constant NO/NO_2 experiment with 300 pptv monoterpene and (d) a varying NO/NO_2 experiment with around 1 ppbv NO_2 and 1200 pptv monoterpene. The brown solid circle in (d) denotes the HOM mass yield obtained from the constant NO/NO_2

experiment with around 1 ppbv NO_2 and 7 pptv NO (the brown dash line in (b)). Model simulated HOM represents molecules formed without NO participation and are denoted in green; HOM_{NO} represents molecules formed with NO 's involvement and are denoted in blue. The propagated error of the HOM yield varies within 6–8% among different experiments with the calculation method provided in the Methods.

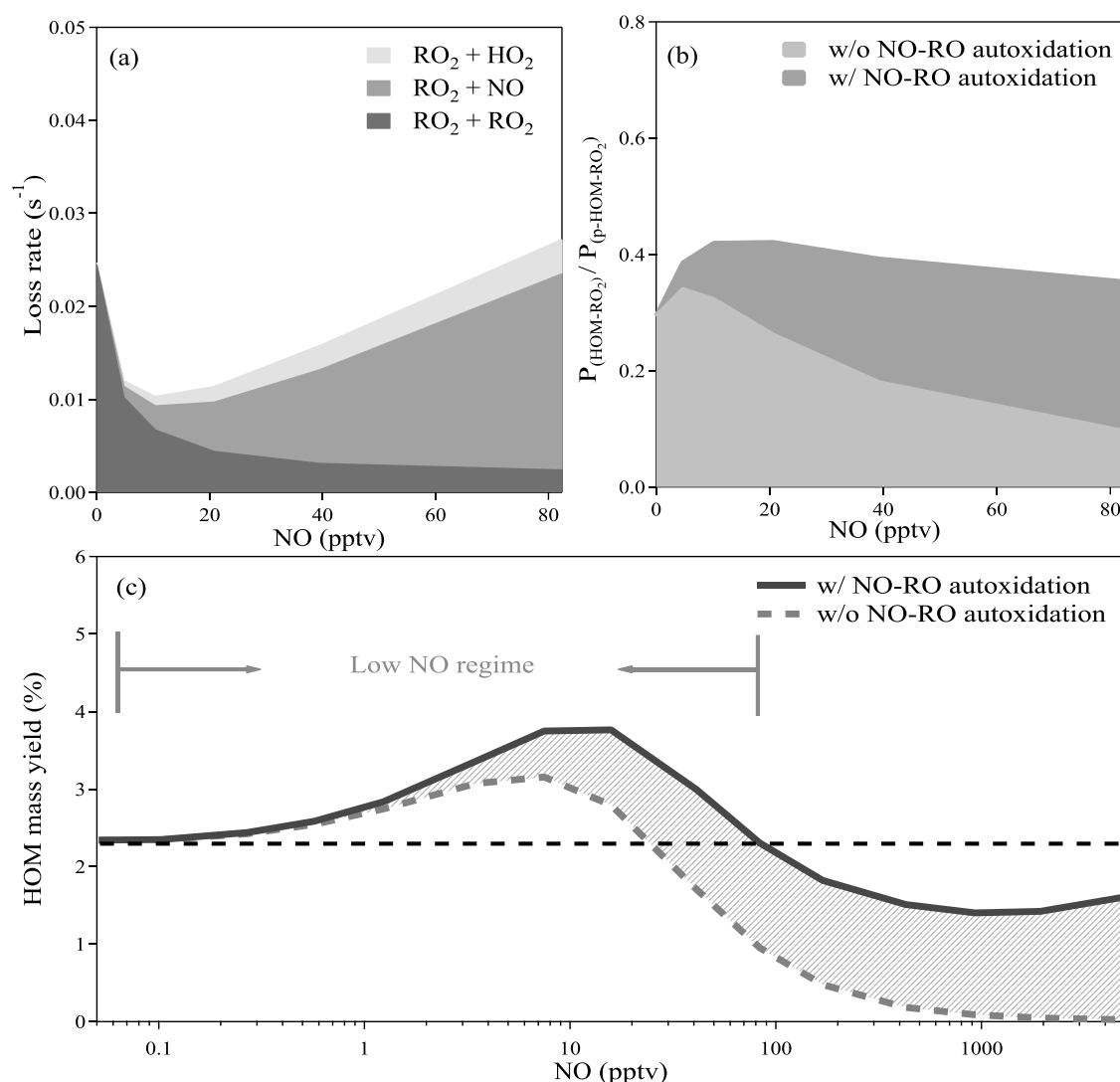


Fig. 2 | The mechanism through which NO can enhance the formation of highly oxygenated organic molecules (HOM) from monoterpene oxidation at 278 K. (a) Loss rates of p-HOM-RO₂ (RO₂ that can undergo autoxidation with $n_O < 7$) as a function of NO concentration in the varying NO/NO₂ experiment. Reactions of RO₂ + HO₂, RO₂ + NO, and RO₂ + RO₂ are marked by pink, blue, and red, respectively. (b) The fraction of p-HOM-RO₂ that can undergo autoxidation to

HOM-RO₂ with (the sum the blue and green fillings) and without (green fillings) NO-induced RO autoxidation in the varying NO/NO₂ experiment. (c) Comparison between HOM yields with (solid blue line) and without NO (dashed green line) RO autoxidation. In the simulation of Fig. 2c, the concentrations of monoterpene, O₃, and NO₂ are constrained at 1200 pptv, 40 ppbv and 1 ppbv.

precise control and monitoring of NO and monoterpenes at low concentrations, which to the best of our knowledge, has never been achieved in previous studies. This allows us to successfully reproduce the chemical environment of the boreal forest atmosphere, and to investigate the HOM production in detail in the low-NO regime. A list of experiments illustrated in Fig. 1 is provided in Supplementary Table 1, and detailed information on the experimental design and instrumentation is provided in Methods.

A nitrate-based chemical ionization mass spectrometer (CI-API-TOF)^{6,31} was deployed to measure gas-phase HOM; these were classified as molecules without (CHO) and with nitrogen atoms (CHON), the latter likely being organic nitrates. We show the variations of some fingerprint CHON- and CHO- HOM for different monoterpene and NO_x values in Supplementary Fig. 1. The corresponding dependence of HOM concentrations and mass yields on NO_x is illustrated in Fig. 1 and Supplementary Fig. 2a–d. The response of HOM production to NO_x levels shows marked differences for different NO_x sequences, indicating that NO_x influences HOM formation in a complicated manner. For example, in the pure

NO₂ sequence, the total HOM yield decreases monotonically as the NO₂ concentration rises (Fig. 1a). However, in a nearly identical sequence with constant NO/NO₂, differing only by a small amount of NO (maximum 28 pptv), HOM yields increase monotonically, from about 2.5% to 4.0% as the NO_x concentration rises (Fig. 1b). This enhancement is more pronounced at low monoterpene concentrations (Fig. 1c) than at high monoterpene concentrations when the NO_x concentration is identical. These results clearly show that the roles of NO and NO₂ in HOM formation are different – while NO₂ suppresses HOM formation monotonically³², NO at low concentrations can enhance HOM formation. The influence of NO on HOM yields is seen more clearly in the varying NO/NO₂ sequence; When NO increases from 0 to 82 pptv at an approximately constant NO₂ concentration near 1 ppbv, it first enhances HOM production, but as it further increases the enhancement gradually diminishes and turns into suppression (Fig. 1d). Altogether, these observations clearly show that NO has a non-linear effect on HOM yields, challenging the notion that NO monotonically suppresses HOM production by inhibiting the autoxidation of RO₂^{15,16}.

A mechanistic understanding of the roles of NO_x in HOM formation

We use the Aerosol Dynamics gas- and particle-phase chemistry model for laboratory CHAMber studies (ADCHAM) to further obtain a mechanistic understanding of the influence of NO on HOM formation¹¹. This model simulates HOM formation on a molecular level. We extend the peroxy radical autoxidation mechanism (PRAM) module^{11,33} in ADCHAM by incorporating NO₃ chemistry and refining the representation of alkoxy radical (RO) isomerization, which better describes HOM formation through the interactions between monoterpenes and NO_x. As shown in Fig. 1 and Supplementary Fig. 2, the model not only reproduces the HOM distribution at various experimental conditions, but also captures the variation of HOM concentrations for the different NO_x sequences we employed in our experiments, showing that the model setup is reasonable and results are robust.

With the help of ADCHAM, we track each reaction channel and divide HOM into HOM_{NO} and HOM_{non-NO}. HOM_{NO} denotes HOM formation directly involving NO in at least one reaction, forming either an organic nitrate or an RO radical. HOM_{non-NO} denotes HOM formation without any reactions involving NO. We can see that the HOM_{NO} increase significantly in three NO-involved sequences (Fig. 1b–d). The fraction of HOM_{NO} is approximately 85 % at low monoterpene concentrations (Fig. 1c) or high NO (Fig. 1d). This high fraction of HOM_{NO} in the model also is consistent with our PMF analysis (see Methods and Supplementary Fig. 3). This also agrees with our previous PMF analysis of ambient data at our reference site, where the dominant daytime factors are NO-involved (accounting for ~90% of the total HOM concentration) with around 80 pptv NO (Supplementary Fig. 4)¹⁷, suggesting NO is involved in the formation of the majority of HOM at quite low concentrations.

Next, we use the varying NO/NO₂ sequence to investigate the non-linear effects of NO on HOM yields. We find two main causes. First, NO suppresses RO₂ production and slows the second-order RO₂ cross-reactions dramatically. When these cross-reactions are the dominant RO₂ sink, they can terminate RO₂ chemistry before isomerization reactions initiate autoxidation. Second, NO enhances RO formation, which also undergoes autoxidation and leads to second-generation RO₂. Both these effects enhance HOM formation when RO₂ + NO is not the major RO₂ sink but suppress HOM formation when RO₂ + NO becomes the dominant RO₂ sink. This drives the nonlinear HOM production with a peak at low but non-zero NO.

An important detail is that only some RO₂ can undergo rapid autoxidation. According to the latest understanding of bicyclic monoterpene oxidation, only a fraction of first-generation RO₂ can undergo cyclobutyl ring breaking, a necessary step for further autoxidation and formation of HOM-RO₂³⁴. This suggests a fixed yield of RO₂ that can undergo autoxidation. The non-linear effect of NO on HOM yield is essentially caused by influencing the further oxidation of these potential HOM-RO₂ (p-HOM-RO₂, defined as RO₂ that can undergo autoxidation with $n_O < 7$) to HOM-RO₂ (defined as RO₂ that is autoxidized to $n_O \geq 7$, Fig. 2b), e.g., from C₁₀H₁₅O₄ to C₁₀H₁₅O_{≥7} in the ozonolysis of monoterpene. In Fig. 2a, we show the loss rate of the p-HOM-RO₂ at different NO levels. NO has a non-linear influence on the loss rate of the p-HOM-RO₂ and thus the possibility that p-HOM-RO₂ becomes HOM-RO₂. When NO is absent, p-HOM-RO₂ loss is primarily due to the cross-reaction with non-HOM-RO₂ (defined as RO₂ that cannot autoxidize). Low but non-zero NO (e.g., 10 pptv) drastically reduces the total loss rate of p-HOM-RO₂ by efficiently reducing the concentration of non-HOM-RO₂. Although NO also consumes p-HOM-RO₂, at low NO this effect is more than compensated by the decrease in non-HOM-RO₂, which dominates the loss of p-HOM-RO₂ at this NO level. As the NO concentration further increases, it progressively becomes the dominant sink of the p-HOM-RO₂ (Fig. 2a) and starts to inhibit HOM production. This intervention in the RO₂ cross-reactions

can also partially explain the suppression effect of NO₂ on HOM formation, in addition to the termination of chain propagation of acyl RO₂³¹. Elevated NO₂ produces considerable non-HOM-RO₂ via NO₃ oxidation, which reduces the probability of p-HOM-RO₂ producing HOM.

Aside from modulating the HOM-RO₂ loss rate, NO also affects the formation of HOM-RO₂ directly by converting RO₂ to RO. As studied under high-NO conditions³⁵, a fraction of RO can isomerize and form new RO₂ via intra-molecular hydrogen abstraction followed by instantaneous O₂ addition. In the context of HOM formation, we refer to RO isomerization as the “RO autoxidation channel” (Supplementary Fig. 5), as it is similar to the definition of RO₂ autoxidation. It should be noted that the RO autoxidation channel has been shown to play a non-negligible role in HOM formation following α-pinene ozonolysis under NO_x-free conditions, as RO₂ cross-reactions also lead to RO formation³⁶. But the presence of NO leads to additional RO formation, making this channel more important. In the case of alkane photo-oxidation, which can have very low yields of highly oxygenated products, the addition of NO was recently shown to greatly enhance the concentration of the most oxygenated products³⁷. This was inferred to be due to several steps of RO isomerization taking place, causing the efficient formation of highly oxygenated species even at 10 ppbv NO.

We cannot directly measure the short-lived RO radicals, so we use distinct HOM products as markers to examine the importance of this channel. In the ozonolysis of monoterpenes, HOM-RO₂ produced via direct autoxidation contain an odd number of hydrogen atoms (n_H) and an even number of oxygen atoms (n_O), e.g., C₁₀H₁₅O_{8,10}. The corresponding closed-shell HOM from either unimolecular termination or bimolecular reactions with RO₂ and HO₂ must be C₁₀H_{14,16}O_{7,9} and C₁₀H₁₆O_{8,10} (Supplementary Fig. 5), which are the major HOM products in many experiments without NO_x^{6,36,38}. In contrast, RO isomerization will produce an RO₂ with an odd n_O , and the consequent closed-shell species are e.g. C₁₀H₁₄O_{8,10} (Supplementary Fig. 5). In Supplementary Fig. 6, we show the variation of C₁₀H₁₄O_{8,10} at different NO levels. As NO rises, they increase by ca. 30 and 100 %, respectively, illustrating the enhancement of the RO autoxidation channel (Fig. 2b). In Fig. 2c, we show the importance of RO autoxidation by comparing the results with and without this reaction channel. The ADCHAM model can only reproduce the experimental data when the RO reaction channel is implemented. Without the RO reaction channel, the predicted HOM yield is significantly lower than the experimental observation, and the difference diverges at higher NO concentrations. This channel forms ca. 25% of HOM even at only 10 pptv NO, and it is responsible for almost all the HOM formation at above 100 pptv NO. Because of this, the HOM yield remains at a considerable level at high NO concentration, although still lower than the yield under NO-free conditions. It is worth noting that RO autoxidation also explains the large fraction of CHO-HOM formed with the direct involvement of NO (Supplementary Fig. 2).

We summarize the dependences of HOM yields on NO and NO₂ in Fig. 3. In general, NO₂ suppresses HOM production monotonically, and NO has a non-linear effect (Fig. 3a). It was previously thought that HOM suppression by NO₂ was mostly due to the direct reaction of RO₂ + NO₂ outcompeting RO₂ autoxidation³². However, we find that NO₂ can suppress HOM production indirectly by increasing the loss rate of HOM-RO₂ - it promotes NO₃-initiated monoterpene oxidation, which forms a considerable amount of non-HOM-RO₂, enhancing the second-order RO₂ cross-reactions. When NO is also present, it counteracts the suppression effects of NO₂: on the one hand, it can significantly consume NO₃; on the other hand, it reduces the total loss rate of HOM-RO₂ (Fig. 2a). Our results suggest that the highest HOM yield at different NO₂ levels is similar when NO is present, although the optimal NO concentration varies slightly. This makes the enhancement of HOM yield by NO more significant at high NO₂ (up to 140% with NO₂ at 10 ppbv, Fig. 3a). The enhancement of HOM yield at low NO is

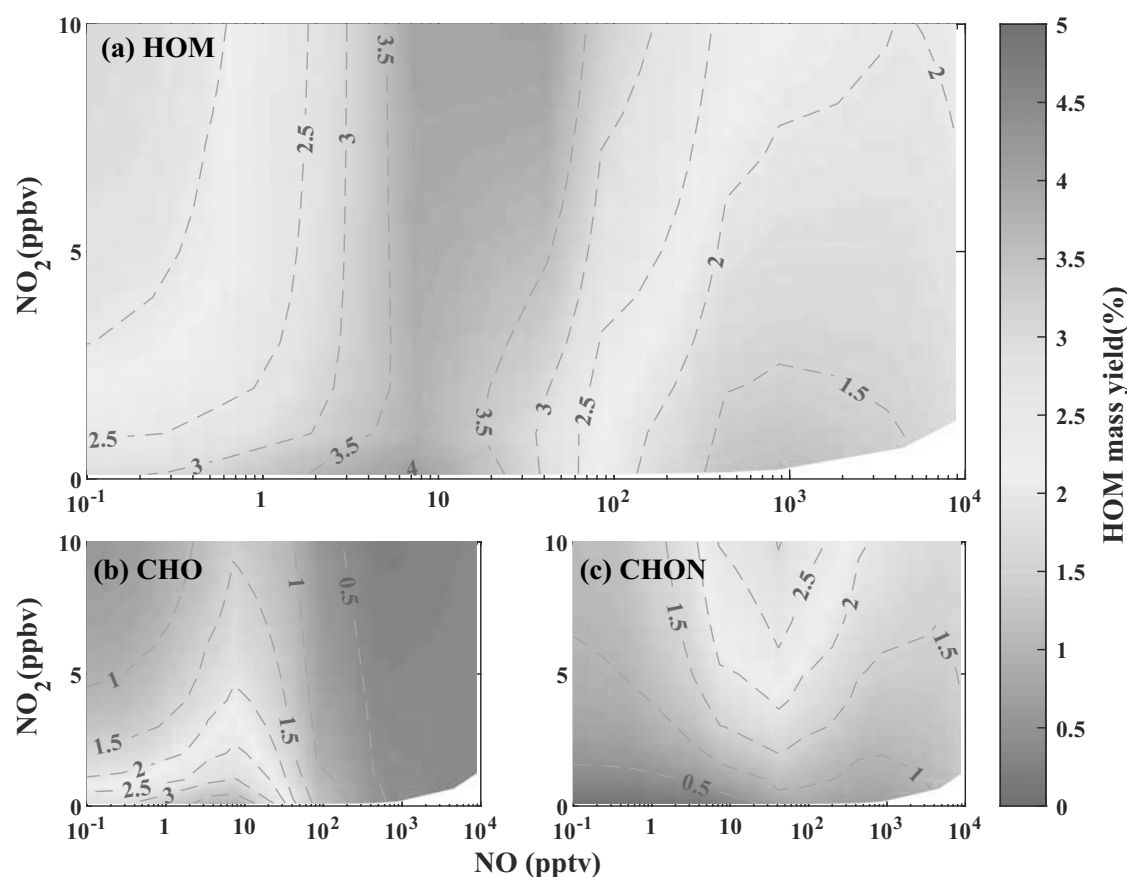


Fig. 3 | Model simulated non-linear dependence of the yields of highly oxygenated organic molecules (HOM) on NO_2 and NO at 278 K. (a) total HOM, (b) CHO-HOM, and (c) CHON-HOM with 1200 pptv monoterpene. O_3 concentrations

are constrained at 40 ppbv at each experiment. The contour lines indicate the corresponding HOM yields.

pronounced for CHO HOM at low NO_2 concentrations (e.g., <5 ppbv) (Fig. 3b) and for CHON HOM at high NO_2 concentrations (e.g., >5 ppbv) (Fig. 3c).

It is worth noting that the quantitative relationships outlined here may be sensitive to conditions in human-influenced environments, where a greater abundance of non-HOM- RO_2 originating from small molecule VOCs may lead to a stronger loss of HOM- RO_2 . In this case the maximum HOM yield is expected to occur at a higher concentration of NO .

Atmospheric observational evidence and implications

We demonstrate that HOM formation is governed by a complex competition between multiple reactions including RO_2 autoxidation, RO isomerization, and various uni-molecular and bi-molecular termination reactions. Therefore, HOM yields vary depending on NO , NO_2 , and monoterpene concentrations and are expected to vary significantly in different environments rather than remain constant. As shown in Fig. 4, we reproduce the HOM yields in a boreal forest in southern Finland by implementing these insights. The details of the simulation inputs and assumptions are provided in Methods. The calculated HOM yield as a function of measured NO_x and monoterpene concentrations varies by almost a factor of three, from 2.5% to 6.5%. This large variation can cause a notable change in both SOA formation and particle growth.

Although this study investigates HOM formation at low and moderate NO levels, the highlighted importance of the RO autoxidation channel also advances our understanding of HOM formation at high NO concentrations, common for monoterpene oxidation in the

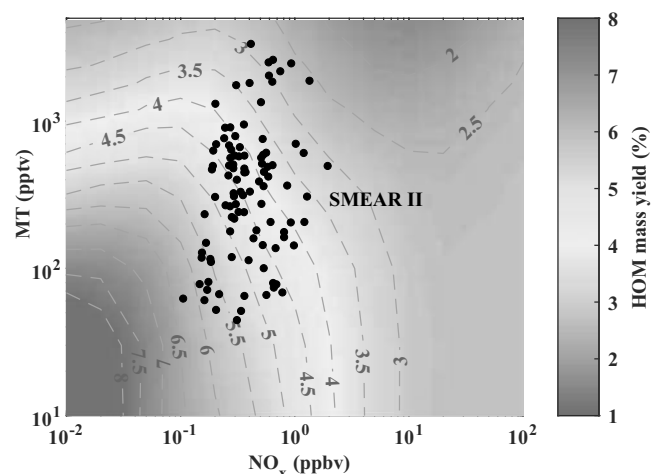


Fig. 4 | Contour plot of the yields of highly oxygenated organic molecules (HOM) from monoterpene oxidation vs monoterpene and NO_x concentration. The observation data of monoterpene and NO_x from the SMEAR II boreal forest station⁵⁸ are used to predict the HOM yield (black points). A series of simulations were conducted with monoterpene concentration varying from 100 to 1800 pptv, and NO_x concentration from 0.01 to 100 ppbv. With the assumption NO is solely formed from NO_2 photolysis, NO/NO_2 varied to follow the observed diurnal cycle at SMEAR II station from almost zero at night to 14.3% at 10 am. We take the daytime average HOM yield as the HOM yield at fixed NO_x and monoterpene concentration.

polluted urban boundary layer. It has been argued that HOM formation can be completely suppressed when the $\text{RO}_2 + \text{NO}$ reaction overwhelms RO_2 autoxidation¹⁵. However, our measurements in polluted East China show that the HOM yield of monoterpene oxidation remains at approximately 1–2% on average (Supplementary Fig. 7), instead of approaching zero as anticipated. This discrepancy can be explained by RO autoxidation, as we demonstrated in Fig. 2c. Besides monoterpene oxidation, the RO autoxidation channel should be important in the oxidation of other VOCs, such as alkanes^{8,37}. This is consistent with the significant anthropogenic HOM production observed in polluted megacities with high NO concentrations⁸. We conclude that high NO cannot completely inhibit HOM formation even in severely polluted urban environments.

In summary, combining experimental data, ambient observations, and molecule-level model simulations, we demonstrate that NO has a non-linear effect on HOM formation. Though NO_x maxima are known for phenomena such as O_3 production and HO_x concentrations, the maxima occur at different points. For HOM production, the maximum is at very low NO, meaning that environments can be “low NO_x ” for O_3 yet “high NO_x ” for HOM. We propose a new division of NO regimes including “Zero NO_x ”, “Low NO”, and “High NO” (Supplementary Table 2) in HOM-related studies. In this new division, previous studies almost completely overlooked the low-NO regime, in which HOM formation is enhanced due to a reduced sink of HOM-RO_2 as well as enhanced RO autoxidation. Conceptually, this non-linear effect is similar to the radical termination via production of organonitrates (at intermediate NO_x)³⁹ and even the catalytic ozone depletion in the stratosphere (with NO_x suppressing HO_x and ClO_x catalysis at low concentrations but ultimately enhancing catalytic ozone loss at sufficiently high NO_x)⁴⁰. All these reveal the notably different outcomes between low-concentration and high-concentration reactions in atmospheric photochemistry. In fact, from a global perspective, high-NO environments are the exceptions, existing only near the ground surface of regions with strong human activity, whereas low-NO environments likely prevail in the vast majority of Earth’s boundary layer. As shown in Supplementary Fig. 8, the low-NO regime ($\text{NO} < 30$ pptv) prevailed at our boreal-forest reference site, accounting for more than half of the observation period from 2002 to 2018. In a megacity of east China, although only 6% of the observation data was lower than 30 pptv, the frequency of low-NO conditions has increased significantly in recent years due to NO_x emission controls. Therefore, this refined understanding of HOM formation in the low-NO regime is crucial for an accurate evaluation of HOM and aerosol budgets, as well as climate change.

Methods

The CLOUD chamber

The CLOUD chamber is a well-controlled stainless-steel cylinder with a volume of ca. 26.1 m³, located at CERN, Geneva, Switzerland^{28,41}. Plenty of efforts were made to keep the chamber ultra-clean, including using an electro-polished inner surface, rinsing with 373 K ultrapure water and flushing the chamber with humidified synthetic air containing several ppmv of ozone thereafter before each campaign. Synthetic air used throughout the campaign was produced by mixing ultra-clean cryogenic liquid nitrogen and oxygen. These efforts achieved an as low as possible contamination level that the background VOCs concentration is below ppbv level and the condensable vapors are mostly below the detection limits, and allowed investigating the chemical processes with atmospherically relevant concentration of reaction precursors⁴², e.g. pptv level monoterpene and NO_x . It is worth noting that with the ultra-clean electro-polished stainless-steel surface, the influence of HO_2 produced from the chamber wall which was a long-standing trouble for Teflon chambers is believed to be ignorable during the campaign.

Several light system covering different regions of the UV and visible spectrum was equipped in the chamber to simulate the atmospheric photochemistry, including a Krypton-Fluoride excimer UV-laser (3 W, $\lambda = 248$ nm), two UV LEDs (2×16.5 W, $\lambda = 370\text{--}390$ nm) and four Hamamatsu Xenon arc lamps (4×200 W, $\lambda = 250\text{--}580$ nm)^{9,43}. For varying NO/NO_2 experiment run during CLOUD11, an additional UV-sabre (400 W UVS3, centered on 385 nm) system was installed to further photolyze NO_2 into NO. The NO/NO_2 ratio can be adjusted by changing the UV-sabre light intensity. The NO_2 photolysis frequency, $j(\text{NO}_2)$, was measured by NO_2 actinometry and varying the UV-sabre intensity.

Experimental design

Our objective of this study is to investigate the role of NO_x in the oxidation of monoterpene and HOM formation in the atmospherically relevant concentrations (low NO_x environment). The chamber conditions were kept stable at the temperature of 278 K and relative humidity of 38%. Ozone was injecting at a constant rate to keep the concentration around 40 ppbv. The monoterpene used in this study is a mix of α -pinene and Δ -3-carene with a volume mixing ratio of 2:1, which represents the condition at our reference site of SMEAR II station in southern Finland^{29,30}. The steady-state concentration of monoterpene mixture was set to 300 pptv, 600 pptv and 1200 pptv. In this study, with the target to understand the role of NO_x in detail, we conducted three different experiment runs, including pure NO_2 run, constant NO/NO_2 ratio runs and, varying NO/NO_2 run. Pure NO_2 run and constant NO/NO_2 ratio runs were conducted during November of 2015 (CLOUD10 campaign). The varying NO/NO_2 experiment was performed in November of 2016 (CLOUD11 campaign). NO_2 was injected into the chamber under dark condition during pure NO_2 run to maintain a NO free environment. NO was injected into the chamber for constant NO/NO_2 runs, and mostly oxidized to NO_2 by reacting with O_3 . For each experiment run, monoterpene was injected into the chamber to get its steady-state concentration before injecting NO_x . The steady-state concentrations of NO_x were set to 0, 0.7, 1.9 and 4.6 ppbv for the pure NO_2 run and constant NO/NO_2 ratio runs, representing very clean to slightly polluted environments, and at approximately 1 ppbv for the varying NO/NO_2 experiment to represent the mean value at the SMEAR II station. Throughout all the NO/NO_2 experimental runs, the UV LEDs light system was kept on to photolyze some NO_2 to NO. The ratio of NO/NO_2 was kept at $\sim 0.7\%$ in constant NO/NO_2 experimental runs. Noting that due to the titration of NO in the NO injection run (constant NO/NO_2 experiment), the steady state concentration of O_3 decreased slightly from 40 ppbv at 0 ppbv NO_x to around 36 ppbv at 4.6 ppbv NO_x . During the varying NO/NO_2 experimental run, NO concentration was adjusted from 0 pptv to 82 pptv by changing the intensity of UV-sabre.

Measurements

Measurement of gas-phase HOM and OH radicals. Concentrations of highly oxygenated molecules (HOM) were measured with a nitrate-ion based chemical ionization atmospheric pressure interface time-of-flight mass spectrometer (CI-API-TOF)⁴⁴. The nitrate ions are produced by exposing nitric acid (HNO_3)-containing sheath flow to soft x-ray radiation. The sulfuric acid and targeted HOM are charged in the drift tube before entering to the API and analyzing in the TOF chamber. To quantify HOM, the system was calibrated and corrected using the methods similar to our previous work^{9,10}. Since the loss rate of sulfuric acid is well characterized, the concentration and calibration coefficient can be obtained. This calibration coefficient was also used to quantify HOM with the consideration of weight-dependent ion transmission⁴⁵ and sampling loss in the inlet. We processed the raw data using the MATLAB tofTools package (version 603)⁴⁴. High-resolution analysis was used to identify the peaks with different elemental formulae.

OH concentrations in the chamber were determined by the following equation.

$$[\text{OH}] = \frac{k_{\text{loss}}[\text{SA}]}{k_{\text{SO}_2 + \text{OH}}[\text{OH}] \times [\text{SO}_2]} \quad (1)$$

Where k_{loss} refers to the loss of sulfuric acid to the aerosol particles (condensation sink) and to the wall (wall loss); $k_{\text{SO}_2 + \text{OH}}$ refers to the reaction rate constant between OH and SO_2 at 278 K. OH concentrations were also simulated by the ADCHEM model. The simulated concentration was identical to the estimated concentration via Eq. (1). NO_3 concentrations were calculated by the change in monoterpene concentration after adjusting NO_x concentration at each experimental steps.

Measurement of monoterpenes and trace gases. A newly developed proton transfer reaction time-of-flight mass spectrometer, named PTR3, was deployed to measure the concentrations of monoterpenes and other VOCs⁴⁶. Since α -pinene and Δ -3-carene fragment differently in the instrument, they can be calibrated individually. Sulfur dioxides (SO_2) and ozone (O_3) were measured using Thermo Scientific gas monitors (model 42i for SO_2 and model 49i for O_3). In view of the low concentration of nitrogen oxide (NO), an advanced NO monitor (ECO PHYSICS, model: CLD 780 TR) was used to measure NO accurately. The detection limit of this NO monitor is 3 pptv for 1-min integration time. Nitrogen dioxide (NO_2) was measured using a cavity-attenuated phase-shift nitrogen dioxide monitor (CAPS NO_2 , Aerodyne Research Inc.) close to the injection of the chamber and by Cavity-Enhanced Differential Optical absorption spectroscopy (CE-DOAS) at the top of the chamber. NO and NO_2 were injected at the bottom of the chamber close to a mixing fan and quickly dispersed into the chamber. NO_2 measurements at the top and bottom of the chamber suggest the NO_2 concentrations to be well mixed, with concentrations at the top and bottom typically not differing by more than 25%. All the gases were sampled from the middle of the chamber. For the NO_2 injection runs, including the pure NO_2 experiment run and the varying NO/ NO_2 experiment run, all the trace gases were homogeneous distributed in the chamber. For the constant NO/ NO_2 experiment runs, the NO mixing ratio close to its injection port was about 5 times higher than that in the middle of the chamber. However, due to the fast reaction of NO and O_3 , the space with highly concentrated NO should be much smaller than the overall chamber volume.

Ambient measurement campaigns of HOM. Ambient HOM measurement data from two field campaigns were used to verify the role NO_x in HOM formation in the real atmosphere. One campaign was conducted at a boreal forest site (SMEAR II station) in southern Finland between 15 and 24 May 2013 to represent the low- NO_x environment¹¹, and the other was at an urban site in east China (SORPES station) during 2 August and 6 September 2019 to represent the high- NO_x environment²². Detailed description of the campaigns can be found in Roldin et al., 2019¹¹ and Liu et al., 2021²². In short, a CI-API-HTOF at SMEAR II and a CI-API-LTOF at SORPES were used to measure HOM and sulfuric acid. VOCs precursors were measured by a proton transfer reaction mass spectrometer (PTR-MS, Ionicon). NO_x and other trace gases (e.g., O_3) were measured by TEI gas analyzers.

Positive matrix factorization (PMF)

PMF is a widely used receptor model for source apportionment analysis⁴⁷. In this study, we first performed the high-resolution peak fitting for the varying NO/ NO_2 experiment run, during which an LTOF based nitrate CIMS was deployed with the resolution up to $\sim 10,000$ Th/Th that allowed us to identify the peaks more accurate. Then, we prepared the data matrix and error matrix according to the methods described by Yan et al.¹⁷ for the PMF model input. An IGOR based analyzing interface SoFi (solution finder, version 6.3) and ME-2 was used to perform the high-resolution PMF analysis⁴⁸. A detailed description of

PMF used for a nitrate-CIMS based dataset can be found in our previous study¹⁷. 5 factors resolved from the PMF analysis provide the optimal solutions. We shortly introduce the 5 factors as follows.

Δ -3-carene ozonolysis factor (D3C factor): 400 pptv Δ -3-carene was injected into the chamber as the first run-step, during which only carene and ozone reacted in the chamber. Therefore, PMF resolved the Δ -3-carene ozonolysis factor as the first factor, the spectrum of which is a bit different from the well-known α -pinene ozonolysis factor (Supplementary Fig. 3)⁶. The pre-dominant molecule is $\text{C}_{10}\text{H}_{14}\text{O}_9$, followed by two RO_2 radicals of $\text{C}_{10}\text{H}_{15}\text{O}_8$ and $\text{C}_{10}\text{H}_{15}\text{O}_{10}$. Dimer formation at this stage is not significant, due to the lower concentration of precursors which that limit the RO_2 cross-reactions.

α -pinene and Δ -3-carene ozonolysis factor (AP + D3C factor): 800 pptv α -pinene was added after the Δ -3-carene step reaching the steady state to obtain the targeted mixture of monoterpenes. PMF identified this system as a pure ozonolysis factor, the spectrum of which is very similar to Nighttime Factor 1 resolved from field observation at SMEAR II station (Fig. S4)¹⁷. Compared to the D3C factor, the intensity of $\text{C}_{10}\text{H}_{14}\text{O}_7$ increased dramatically and was to be one of the main fingerprint molecules of the AP + D3C factor. In contrast, RO_2 radicals (e.g., $\text{C}_{10}\text{H}_{15}\text{O}_8$ and $\text{C}_{10}\text{H}_{15}\text{O}_{10}$) decreased significantly, probably due to the enhanced RO_2 cross-reactions which consumed RO_2 radicals but promoted the dimer formations, e.g., $\text{C}_{19}\text{H}_{28}\text{O}_{11}$ and $\text{C}_{20}\text{H}_{32}\text{O}_{11}$. In addition, the intensity of OH-initiated HOM molecules, e.g., $\text{C}_{10}\text{H}_{16}\text{O}_7$ and $\text{C}_{10}\text{H}_{16}\text{O}_9$, enhanced significantly. It could be that OH production from monoterpene ozonolysis was elevated, or the HOM yield from OH oxidation of α -pinene was higher than OH oxidation of Δ -3-carene.

NO_3 oxidation factor (NO_3 factor): Around 1 ppb NO_2 was injected into the chamber after the above two steps. NO_2 reacted with O_3 to form NO_3 radical, which oxidized monoterpenes to be an additional HOM source. Therefore, PMF resolved an NO_3 factor aside from the pure ozonolysis factor at this stage. The spectrum of the NO_3 factor was similar to that of the Nighttime Factor 2 resolved from field observation at SMEAR II station (Fig. S4)¹⁷. CHON monomers of $\text{C}_{10}\text{H}_{15}\text{O}_{8,10}\text{N}$ and dimers of $\text{C}_{20}\text{H}_{31}\text{O}_{11,13}\text{N}$ were the major fingerprint molecules. A significant formation of $\text{C}_{10}\text{H}_{16}\text{O}_{10}\text{N}_2$ was observed at this stage (in the NO_3 factor), during which only NO_3 and NO_2 presented as the N-containing precursor gases. This is different from previous understanding that CHON₂ can be only formed from a CHON radical produced from NO_3 oxidation and terminated by NO. One possible explanation is that $\text{C}_{10}\text{H}_{16}\text{O}_{10}\text{N}_2$ in this factor is acyl nitrate formed via the reaction of an acylperoxy radical with NO_2 .

NO involved NO_3 oxidation factor (NO_3 + NO factor): After producing NO by photolyzing NO_2 in the chamber, a new factor was resolved and defined as NO_3 + NO factor (Supplementary Fig. 3). This factor is a transition factor from the NO_3 factor to a NO heavily participating factor (NO factor). A bunch of CHON molecules, e.g., $\text{C}_{10}\text{H}_{15}\text{O}_7\text{N}$ and $\text{C}_{10}\text{H}_{15}\text{O}_9\text{N}$, were formed from the reaction between RO_2 radical and NO. Additional CHON₂ molecules other than $\text{C}_{10}\text{H}_{16}\text{O}_{10}\text{N}_2$, e.g., $\text{C}_{10}\text{H}_{16}\text{O}_9\text{N}_2$, were formed from the NO_3 -initiated CHON radical reacting with NO. Dimer formation became very weak. Fragmented molecules (e.g., $\text{C}_5\text{H}_6\text{O}_7$), as well as C_{15} HOM molecules (e.g., $\text{C}_{15}\text{H}_{17}\text{O}_{9,10}$), started to appear in this factor because of the NO, which can react with RO_2 to form RO to enhance the possibility of fragmentation.

NO heavily participating factor (NO factor): When further stepping up the NO concentration, PMF resolved another factor, in which NO participated the HOM formation heavily (Supplementary Fig. 3). In this factor, a considerable amount of fragmented molecules (e.g., $\text{C}_5\text{H}_6\text{O}_5$, $\text{C}_5\text{H}_6\text{O}_7$, $\text{C}_7\text{H}_{11}\text{O}_8$, and $\text{C}_8\text{H}_{13}\text{O}_9\text{N}$) were observed. HOM dimers almost disappeared, suggesting the reaction with NO was becoming the dominating fate of RO_2 instead of its cross-reaction.

In summary, optimal PMF solutions provide 4 types of factors for all the four experimental runs, including two ozonolysis factors (D3C

factor and MT factor), an NO₃ oxidation factor (NO₃ factor), an NO involved NO₃ oxidation factor (NO₃ + NO factor) and a NO heavily-participating factor (NO factor). Factors appear successively from the D3C factor to the NO factor accompanying with stepping up of NO₂ or NO. The spectrum of HOM changes as follows: it is dominated by typical CHO HOM in the ozonolysis factors, but switches to CHON monomers (e.g., C₁₀H₁₅O_{8,10}N) and dimers (e.g., C₂₀H₃₁O_{11,13}N) in NO₃ factor. In the NO₃ + NO factor, CHON (e.g., C₁₀H₁₅O₇₋₁₀N) dominates and CHON₂ (e.g., C₁₀H₁₆O_{9,10}N₂) appears extensively. Finally, a significant amount of fragmented molecules (e.g., C₅H₆O₅, C₅H₆O₇, C₇H₁₁O₈) appears in the NO factor. HOM dimers decreases gradually and almost disappear in the NO factor. This is in accordance with our previous understanding that CHON radicals and CHON dimers can be only formed from the NO₃ involved reaction chain; most CHON₂ are only formed from a CHON radical produced from NO₃ oxidation and terminated by NO, except for some acyl nitrate formed via the reaction of an acylperoxy radical with NO₂ (e.g., C₁₀H₁₆O₁₀N₂).

ADCHEM model

We use the Aerosol Dynamics gas- and particle-phase chemistry model for laboratory CHAMber studies (ADCHAM)⁴⁹ and the trajectory model for Aerosol Dynamics gas- and particle-phase CHEMistry and radiative transfer (ADCHEM)⁵⁰. They share a detailed gas-phase kinetic code that combines peroxy radical autooxidation mechanism (PRAM) and Master Chemical Mechanism version 3.3.1 (MCMv3.3.1) using the Kinetic Pre-Processor (KPP)⁵¹. As described in Roldin et al., 2019⁵¹, PRAM explicitly simulates the formation of peroxy radicals (RO₂) and their oxidation products. We provide the yield, branch ratio, and autooxidation rate constants of main chemical reactions in Supplementary Table 3. Given that the mechanism of fragmentation chemistry is not well understood, ADCHAM tends to underestimate the fragments, the production of which is accompanied by an increase in the NO/MT ratio. Therefore, the model slightly underestimates the HOM yield in the highest NO_x step of the constant NO/NO₂ experiment with 300 pptv monoterpene (Fig. 1c).

In this work, we added the following mechanisms into the model to improve the simulation.

- (1) NO₃ chemistry in HOM formation, which included RO₂ formed when monoterpenes are oxidized by NO₃ and RO formed via RO₂-NO₃ reactions.
- (2) Terminal reaction by bimolecular reactions between NO₂ and certain RO₂ (e.g., C₁₀H₁₅O₅).
- (3) Various isomers of RO₂ were considered, which greatly reduced the uncertainty of simulating responses that are highly dependent on RO₂ structures.
- (4) Refine the representation of RO isomerization

Since HOM have very low saturation vapor pressures, we assume that they are lost losses irreversibly onto the chamber wall, and calculate the wall loss from the following equation⁵²:

$$k_{\text{wall}} = C_{\text{wall}} \sqrt{D} \quad (2)$$

where C_{wall} is an empirical parameter, which is $0.0075 \text{ cm}^{-1} \text{ s}^{-0.5}$ derived from dedicated sulfuric acid decay experiments at 5 °C. The diffusion coefficients D_i for each HOM_{*i*} are approximated with the equation $D_i (\text{cm}^2 \text{ s}^{-1}) = 0.31 \cdot M_i^{-1/3}$, where M_i is the mass of molecule with the unit of g mol^{-1} .

HOM yield calculation

HOM yield calculation in CLOUD chamber and SMEARII station: we defined the yield of HOM as the fraction of reacted monoterpenes that produced HOM. Due to the complex formation of HOM through multiple oxidations, the production rate of HOM was replaced by

the HOM loss rate under steady state:

$$\text{MT} - \text{HOM yield} = \frac{k_{\text{loss}}[\text{HOM}]}{k_{\text{MT} + \text{oxidants}}[\text{oxidants}] \times [\text{MT}]} \quad (3)$$

where oxidants consist of O₃, OH, and NO₃ in this study and k_{loss} is the total loss rate of HOM to the chamber walls in CLOUD experiments, to the ground or other surfaces (Dry deposition) and aerosol particles (condensation) in field observations. We propagate the error for each chamber experiment with four different NO_x levels. Within each selected time windows, absolute errors of [HOM], [SO₂], [sulfuric acid], CS, [O₃], [α-pinene], [Δ-3-carene] are derived from the measurements as the 3-sigma standard deviation. Since NO₃ concentrations are calculated based on the consumed monoterpene concentration, they share the same relative error with [α-pinene] and [Δ-3-carene]. For the loss of HOM, we assume zero error for k_{wall} and k_{dilution} . The propagated error of the HOM yield varies slightly within 6–8% among different experiments (Fig. 1).

When calculating condensation for HOM measured at SMEAR II, pure liquid saturation vapor pressures(p_0) were calculated based on saturation concentrations (C_{sat}), which can be parameterized by the numbers of carbon(n_C), oxygen(n_O), and nitrogen(n_N) atoms⁵³:

$$\log_{10}(C_{\text{sat}}) = (n_O - n_C)b_C - (n_O - 3n_N)b_O - 2 \frac{(n_O - 3n_N)n_C}{(n_C + n_O - 3n_N)}b_{CO} - n_Nb_N \quad (4)$$

where $n_O = 25$, $b_C = 0.475$, $b_O = 0.2$, $b_{CO} = 0.9$, and $b_N = 2.5$, respectively.

As volatilities of organic molecules detected span over a wide range, we grouped the hundreds of organic molecules detected in the chamber into different bins within a volatility basis set (VBS)⁵⁴. Then we can calculate the HOM mass yield via formula (3), in which monoterpene concentrations were measured, HOM was also measured and grouped into a series of bins, k_{loss} , OH and NO₃ radical were simulated by the model. To minimize the potential uncertainty, we selected HOM species with $C^*(300 \text{ K})$ below $3 \times 10^{-4} \mu\text{g m}^{-3}$ as those would condense onto aerosol particles irreversibly.

HOM yield calculation at SORPES station: In polluted east China, where VOCs species distribution and photochemistry are far more complex than in the boreal environment, it is challenging to simulate HOM concentrations. Especially during nighttime, when the NO concentration is very high and consume most known oxidants, e.g., NO₃ radical and O₃. Therefore, we only calculated the daytime HOM yield through the following Eq. (5):

$$\text{MT} - \text{HOM yield} = \frac{[\text{ELVOCs}]_{\text{MT}} \cdot \text{CS}}{k_{\text{MT} + \text{OH}}[\text{MT}] \cdot [\text{OH}] + k_{\text{MT} + \text{O}_3}[\text{MT}] \cdot [\text{O}_3]} \quad (5)$$

Here, $[\text{ELVOCs}]_{\text{MT}}$ is the concentration of monoterpene-derived ELVOCs, [MT] is the concentration of monoterpenes. Monoterpene derived HOM are selected as HOM molecules with a carbon number of 10 and a double bond equivalent (DBE) number between 2 and 4⁸. Since the atmospheric oxidation capacity in polluted east China is usually strong, HOM are possibly produced through multi-generational oxidation. We estimated the contribution of multi-generational by the following method, and excluded it from the yield calculation. First, dinitrates (HOM molecules with two nitrate groups) are regarded as the multi-generational products, since one oxidation step can only add one nitrate group to the product molecule; Second, HOM molecules with a DBE of 2 are produced either from OH-initiated oxidation of monoterpenes or from the multi-generational oxidation process of monoterpenes. The calculated contribution of multi-generational oxidation to monoterpene HOM ranges from 15.8% to 33.5%. Here, we only selected non-nitrates and mononitrates with a DBE of 3–4 to calculate the yield to

exclude the effects of multi-generational oxidation (Supplementary Fig. 7).

Molecules with $C^*(300\text{ K})$ of below $3 \times 10^{-4} \mu\text{g m}^{-3}$ were further selected as ELVOCs those can condense onto aerosol particles irreversibly. Since the oxidation products with higher volatility are not included due to their long lifetimes cause the steady-state assumption not to hold, the HOM yield calculated here, which is considerably under-estimated, can be considered as the low limit. The OH concentration was calculated by applying the Eq. (1), based on the assumption that gaseous sulfuric acid is mostly produced from the oxidation of SO_2 by OH and primarily lost by condensing onto particles. $k_{\text{OH}+\text{SO}_2}$ is a termolecular reaction constant for the rate-limiting step of the formation pathway of H_2SO_4 at the temperature of the sampling time in the atmosphere⁵⁵, and k_{loss} is the loss rate of H_2SO_4 by condensation to aerosol surface (CS, condensation sink). Although reactions of SO_2 with products from the ozonolysis of alkenes generate a moderate amount of nighttime sulfuric acid, with little effect on daytime sulfuric acid⁵⁶. This is also one reason that we calculated the yield of MT-HOM only during the daytime in this study. The error of OH does not change the relative distribution of RO_2 from different precursors.

In the atmosphere without direct anthropogenic emissions, NO concentrations, mainly from the photolysis of NO_2 , show a clear diurnal cycle from almost zero during the night to its maximum before noontime. For example, NO concentration and NO to NO_2 ratio revealed its peak value at around 10:00–11:00 am at our reference station of SMEAR II¹⁷. Correspondingly, the HOM yield will vary nonlinearly and diurnally. Therefore, the HOM yield is determined by the ratio of NO/NO_2 at a specific time of the day. We thus take the daytime average HOM yield as the HOM yield at fixed NO_x and monoterpene concentration. We conducted a series of simulations with monoterpene concentration varying from 100 to 1800 pptv, and NO_x concentration from 0.01 to 100 ppb. With the assumption NO is solely formed from NO_2 photolysis, the NO/NO_2 ratio varied to follow the observed diurnal cycle at SMEAR II station from almost zero at night to 14.3% at 10 am. The average daily dependence of HOM on monoterpene and NO_x concentrations can be estimated on the basis of integration with this function. We then take the daily average HOM yield as the HOM yield at fixed NO_x and monoterpene concentration, and illustrated in Fig. 4.

Data availability

The observation data that support the main findings of this study are available at figshare (<https://doi.org/10.6084/m9.figshare.22724648.v1>)⁵⁷.

Code availability

Data processing techniques are available on request from the corresponding author.

References

- Boucher, O. et al. Clouds and Aerosols. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change [Stocker, T. F., D. Qin, G.-K. Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P. M. Midgley (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA. (2013).
- Jimenez, J. L. et al. Evolution of organic aerosols in the atmosphere. *Science* **326**, 1525–1529 (2009).
- Hallquist, M. et al. The formation, properties and impact of secondary organic aerosol: current and emerging issues. *Atmos. Chem. Phys.* **9**, 5155–5236 (2009).
- Zhang, Q. et al. Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes. *Geophys. Res. Lett.* **34**, <https://doi.org/10.1029/2007GL029979> (2007).
- Ng, N. L. et al. Organic aerosol components observed in Northern Hemispheric datasets from Aerosol Mass Spectrometry. *Atmos. Chem. Phys.* **10**, 4625–4641 (2010).
- Ehn, M. et al. A large source of low-volatility secondary organic aerosol. *Nature* **506**, 476–479 (2014).
- Bianchi, F. et al. Highly Oxygenated Organic Molecules (HOM) from gas-phase autoxidation involving peroxy radicals: a key contributor to atmospheric aerosol. *Chem. Rev.* **119**, 3472–3509 (2019).
- Nie, W. et al. Secondary organic aerosol formed by condensing anthropogenic vapours over China's megacities. *Nat. Geosci.* **15**, 255–261 (2022).
- Kirkby, J. et al. Ion-induced nucleation of pure biogenic particles. *Nature* **533**, 521–526 (2016).
- Tröstl, J. et al. The role of low-volatility organic compounds in initial particle growth in the atmosphere. *Nature* **533**, 527–531 (2016).
- Roldin, P. et al. The role of highly oxygenated organic molecules in the Boreal aerosol-cloud-climate system. *Nat. Commun.* **10**, 4370 (2019).
- Seinfeld, J. H. & Pandis, S. N. *Atmospheric chemistry and physics: from air pollution to climate change*. (John Wiley & Sons, 2016).
- Shrivastava, M. et al. Urban pollution greatly enhances formation of natural aerosols over the Amazon rainforest. *Nat. Commun.* **10**, 1046 (2019).
- Tan, Z. et al. Daytime atmospheric oxidation capacity in four Chinese megacities during the photochemically polluted season: a case study based on box model simulation. *Atmos. Chem. Phys.* **19**, 3493–3513 (2019).
- Pye, H. O. T. et al. Anthropogenic enhancements to production of highly oxygenated molecules from autoxidation. *Proc. Natl Acad. Sci.* **116**, 6641–6646 (2019).
- Praske, E. et al. Atmospheric autoxidation is increasingly important in urban and suburban North America. *Proc. Natl Acad. Sci.* **115**, 64–69 (2018).
- Yan, C. et al. Source characterization of highly oxidized multifunctional compounds in a boreal forest environment using positive matrix factorization. *Atmos. Chem. Phys.* **16**, 12715–12731 (2016).
- Yan, C. et al. Size-dependent influence of NO_x on the growth rates of organic aerosol particles. *Sci. Adv.* **6**, eaay4945 (2020).
- Lee, B. H. et al. Highly functionalized organic nitrates in the southeast United States: contribution to secondary organic aerosol and reactive nitrogen budgets. *Proc. Natl Acad. Sci.* **113**, 1516–1521 (2016).
- Massoli, P. et al. Ambient measurements of highly oxidized gas-phase molecules during the southern oxidant and aerosol study (SOAS) 2013. *ACS Earth Space Chem.* **2**, 653–672 (2018).
- Xu, Z. N. et al. Multifunctional products of isoprene oxidation in polluted atmosphere and their contribution to SOA. *Geophys. Res. Lett.* **48**, e2020GL089276 (2021).
- Liu, Y. et al. Formation of condensable organic vapors from anthropogenic and biogenic volatile organic compounds (VOCs) is strongly perturbed by NO_x in eastern China. *Atmos. Chem. Phys.* **21**, 14789–14814 (2021).
- Zhao, Y., Thornton, J. A. & Pye, H. O. T. Quantitative constraints on autoxidation and dimer formation from direct probing of monoterpene-derived peroxy radical chemistry. *Proc. Natl Acad. Sci.* **115**, 12142–12147 (2018).
- Lehtipalo, K. et al. Multicomponent new particle formation from sulfuric acid, ammonia, and biogenic vapors. *Sci. Adv.* **4**, eaau5363 (2018).
- Schervish, M. & Donahue, N. M. Peroxy radical chemistry and the volatility basis set. *Atmos. Chem. Phys.* **20**, 1183–1199 (2020).

26. Zhang, X. et al. Influence of vapor wall loss in laboratory chambers on yields of secondary organic aerosol. *Proc. Natl Acad. Sci.* **111**, 5802–5807 (2014).
27. Shilling, J. E. et al. Loading-dependent elemental composition of α -pinene SOA particles. *Atmos. Chem. Phys.* **9**, 771–782 (2009).
28. Kirkby, J. et al. Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation. *Nature* **476**, 429–433 (2011).
29. Rinne, J., Hakola, H., Laurila, T. & Rannik, Ü. Canopy scale monoterpene emissions of *Pinus sylvestris* dominated forests. *Atmos. Environ.* **34**, 1099–1107 (2000).
30. Hakola, H., Hellén, H., Hemmilä, M., Rinne, J. & Kulmala, M. In situ measurements of volatile organic compounds in a boreal forest. *Atmos. Chem. Phys.* **12**, 11665–11678 (2012).
31. Jokinen, T. et al. Atmospheric sulphuric acid and neutral cluster measurements using CI-API-TOF. *Atmos. Chem. Phys.* **12**, 4117–4125 (2012).
32. Rissanen, M. P. NO₂ suppression of autoxidation–inhibition of gas-phase highly oxidized dimer product formation. *ACS Earth Space Chem.* **2**, 1211–1219 (2018).
33. Roldin, P. In *Supplement to: Roldin, Pontus; Ehn, Mikael; Kurtén, Theo; Olenius, Tinja; Rissanen, Matti P; Sarnela, Nina; Elm, Jonas; Rantala, Pekka; Hao, Liqing; Hyttinen, Noora; Heikkinen, Liine; Worsnop, Douglas R; Pichelstorfer, Lukas; Xavier, Carlton; Clusius, Petri; Öström, Emilie; Petäjä, Tuukka; Kulmala, Markku; Vehkamäki, Hanna; Virtanen, Annele; Riipinen, Ilona; Boy, Michael* (2019): The role of highly oxygenated organic molecules in the Boreal aerosol-cloud-climate system. *Nat. Commun.*, **10**, <https://doi.org/10.1038/s41467-019-12338-8> (PANGAEA, 2019).
34. Iyer, S. et al. Molecular mechanism for rapid autoxidation in α -pinene ozonolysis. *Nat. Commun.* **12**, 878 (2021).
35. Matsunaga, A. & Ziemann, P. J. Branching ratios and rate constants for decomposition and isomerization of β -hydroxyalkoxy radicals formed from OH radical-initiated reactions of C₆–C₁₃ 2-methyl-1-alkenes in the presence of NO_x. *J. Phys. Chem. A* **123**, 7839–7846 (2019).
36. Molteni, U. et al. Formation of highly oxygenated organic molecules from α -Pinene Ozonolysis: chemical characteristics, mechanism, and kinetic model development. *ACS Earth Space Chem.* **3**, 873–883 (2019).
37. Wang, Z. et al. Efficient alkane oxidation under combustion engine and atmospheric conditions. *Commun. Chem.* **4**, 18 (2021).
38. Ehn, M. et al. Gas phase formation of extremely oxidized pinene reaction products in chamber and ambient air. *Atmos. Chem. Phys.* **12**, 5113–5127 (2012).
39. Perring, A. E., Pusede, S. E. & Cohen, R. C. An observational perspective on the atmospheric impacts of alkyl and multifunctional nitrates on ozone and secondary organic aerosol. *Chem. Rev.* **113**, 5848–5870 (2013).
40. Wennberg, P. O. et al. Removal of stratospheric O₃ by radicals: in situ measurements of OH, HO₂, NO, NO₂, ClO, and BrO. *Science* **266**, 398–404 (1994).
41. Duplissy, J. et al. Effect of ions on sulfuric acid–water binary particle formation: 2. Experimental data and comparison with QC-normalized classical nucleation theory. *J. Geophys. Res.: Atmospheres*. **121**, 1752–1775 (2016).
42. Schnitzhofer, R. et al. Characterisation of organic contaminants in the CLOUD chamber at CERN. *Atmos. Meas. Tech.* **7**, 2159–2168 (2014).
43. Kupc, A. et al. A fibre-optic UV system for H₂SO₄ production in aerosol chambers causing minimal thermal effects. *J. Aerosol Sci.* **42**, 532–543 (2011).
44. Junninen, H. et al. A high-resolution mass spectrometer to measure atmospheric ion composition. *Atmos. Meas. Tech.* **3**, 1039–1053 (2010).
45. Heinritzi, M. et al. Characterization of the mass-dependent transmission efficiency of a CIMS. *Atmos. Meas. Tech.* **9**, 1449–1460 (2016).
46. Breitenlechner, M. et al. PTR3: an instrument for studying the life-cycle of reactive organic carbon in the atmosphere. *Anal. Chem.* **89**, 5824–5831 (2017).
47. Paatero, P. & Tapper, U. Positive matrix factorization: a non-negative factor model with optimal utilization of error estimates of data values. *Environmetrics* **5**, 111–126 (1994).
48. Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U. & Prévôt, A. S. H. SoFi, an IGOR-based interface for the efficient use of the generalized multilinear engine (ME-2) for the source apportionment: ME-2 application to aerosol mass spectrometer data. *Atmos. Meas. Tech.* **6**, 3649–3661 (2013).
49. Roldin, P. et al. Modelling non-equilibrium secondary organic aerosol formation and evaporation with the aerosol dynamics, gas- and particle-phase chemistry kinetic multilayer model ADCHEM. *Atmos. Chem. Phys.* **14**, 7953–7993 (2014).
50. Roldin, P. et al. Development and evaluation of the aerosol dynamics and gas phase chemistry model ADCHEM. *Atmos. Chem. Phys.* **11**, 5867–5896 (2011).
51. Damian, V., Sandu, A., Damian, M., Potra, F. & Carmichael, G. R. The kinetic preprocessor KPP - a software environment for solving chemical kinetics. *Comput. Chem. Eng.* **26**, 1567–1579 (2002).
52. Simon, M. et al. Molecular understanding of new-particle formation from α -pinene between –50 and +25 °C. *Atmos. Chem. Phys.* **20**, 9183–9207 (2020).
53. Mohr, C. et al. Molecular identification of organic vapors driving atmospheric nanoparticle growth. *Nat. Commun.* **10**, 7 (2019).
54. Donahue, N. M., Epstein, S. A., Pandis, S. N. & Robinson, A. L. A two-dimensional volatility basis set: 1. organic-aerosol mixing thermodynamics. *Atmos. Chem. Phys.* **11**, 3303–3318 (2011).
55. Finlayson-Pitts, B. J. & Pitts, J. N. *Chemistry of the Upper and Lower Atmosphere* (eds B. J. Finlayson-Pitts & J. N. Pitts) 294–348 (Academic Press, 2000).
56. Yang, L. et al. Towards building a physical proxy for gas-phase sulfuric acid concentration based on its budget analysis in polluted Yangtze River Delta, east China. *Environ. Sci. Technol.* **55**, 6665–6676 (2021).
57. Nie, W. (2023): NO at low concentration can enhance the formation of highly oxygenated biogenic molecules in the atmosphere. figshare. Dataset. <https://doi.org/10.6084/m9.figshare.22724648.v1>
58. Huang, W. et al. Measurement report: molecular composition and volatility of gaseous organic compounds in a boreal forest – from volatile organic compounds to highly oxygenated organic molecules. *Atmos. Chem. Phys.* **21**, 8961–8977 (2021).

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Author contributions

W.N., C.Y., A.D., M.E. and N.M.D. designed the study; W.N., C.Y., L.Y., P.R. and Y.L. analyzed the data; W.N., C.Y., L.Y., A.D., M.E. and N.M.D. wrote the manuscript; A.V., D.S., H.F., A.A., F.B., J.C., L.D., D.D., J.Du., A.H., X-C.H., V.H., T.J., C.K., K.L., L.N., R.M., V.M., B.M., A.M-O., L.L.Q., S.S., M.S., C.T., A.T., R.V., A.C.W., R.W., M. W., P.Y., H.L., W.H., X.Q., S.L., T.L. and X.C. collected the data and contributed to the analysis; U.M., T.P., J.Do., U.B., I.E.H., J.K., D.W. and M.K. contributed to data interpretation and editing of manuscript.

Competing interests

The authors declare no competing interests.

Additional information

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