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Sewer subsidies from overflows and pipe leaks dominate urban stream solute loads in all storm events

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Stream channel burial drastically alters watershed flowpaths by routing surface waters underground and increasing the potential for interactions between stream water and urban infrastructure such as storm and sanitary sewers. While numerous studies have investigated storm event solute loads from urban watersheds, the influences of stream channel burial and sewer overflows are often overlooked. This study uses grab samples and natural abundance stable isotope tracers to quantify the event dynamics of solute concentration-discharge relationships as well as cumulative loads in a buried urban stream. Our results demonstrate that different solutes, as well as different sources of the same solute (atmospheric NO_3^- and sewer-derived NO_3^- differentiated by the $\Delta^{17}\text{O}$ tracer), are delivered via separate watershed flowpaths and thus have different timings within the event and contrasting relationships to flow. This inter-event variability reveals dynamics that result from temporal and spatial heterogeneity in infiltration, exfiltration, and pipe overflows. These results can help guide system-wide infrastructure maintenance as cities seek to meet challenges in sustaining and improving water quality as infrastructural systems age.

KEYWORDS

urban areas, buried streams, sewage subsidies, nitrate isotopes, hysteresis, stormwater management

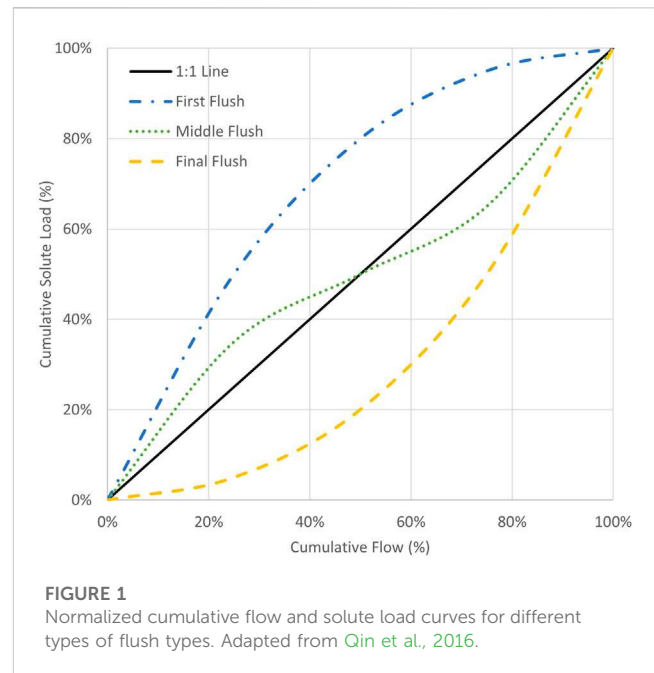
1 Introduction

Storms mobilize solutes from across watersheds and connect storage zones to the stream that are disconnected in dry weather conditions. This is especially true in urban watersheds where large areas of impervious surfaces are connected to the stream channel via stormwater pipe networks and lead to high and flashy event flows (Leopold, 1968; Graf, 1975). Additionally, urban watershed flowpaths are complicated by stream burial, where surface stream channels are deliberately relocated to subsurface pipes (Elmore and Kaushal, 2008; Forgrave et al., 2022), often in the storm drainage network (Broadhead et al., 2013). Burial is common in many older US cities (Elmore and Kaushal, 2008; Roy et al., 2015; Napieralski and Welsh, 2016; Weitzell et al., 2016; Hopkins and Bain, 2018; Forgrave et al., 2022) and is concentrated in spatial extent, resulting in urban areas where complete burial has created “stream deserts” or areas of land with no surface streams (Napieralski et al., 2015; Napieralski and Carvalhaes, 2016). This process of stream burial has three significant impacts on the connectivity between watershed solute sources and the stream network: 1) direct connections to the storm drainage network increase loads of surface solutes (Kaushal and Belt, 2012; Hobbie et al., 2017; Fork et al., 2018), 2) relocating streams underground increases the

likelihood of connectivity to subsurface solutes such as sewage leaks (Broadhead et al., 2013; 2015; Divers et al., 2014; Hopkins and Bain, 2018; Delesantro et al., 2022), and 3) burial disconnects streams from their surface floodplains and wetlands—locations that can slow down flow and promote nutrient uptake, settling, and denitrification (Groffman and Crawford, 2003; Klocker et al., 2009; Pennino et al., 2014; Newcomer Johnson et al., 2016). These compounding disturbances to watershed flowpaths have significant biogeochemical consequences (Beaulieu et al., 2014; 2015; Divers et al., 2014; Pennino et al., 2014), particularly during storm events through activation of flowpaths and hydrologic connections to surface and subsurface solute reservoirs.

Two common methods to quantify solute event dynamics are concentration-discharge hysteresis and load-volume curves. Solute concentration-discharge (c - Q) relationships often exhibit hysteresis during the storm hydrograph where concentrations at a particular discharge differ between the rising and falling limb of the hydrograph (Evans and Davies, 1998). These hysteretic patterns can clarify solute sources and active catchment flow paths for any given storm event (Carey et al., 2014; Bowes et al., 2015; Lloyd et al., 2016a; Lloyd et al., 2016b; Duncan et al., 2017; Baker and Showers, 2019). Solute concentrations that increase with rising streamflow are categorized as “transport-limited” indicating that discharge determines mobilization of readily available solute pools, also referred to as “flushing” from near-stream sources (Bowes et al., 2015). Conversely, when solute concentrations decrease with increasing discharge, these solutes are considered “source-limited” due to finite supply that is exhausted and diluted with higher streamflow (Bowes et al., 2015). When solute concentrations remain constant, despite changes in flow, this condition is called chemostasis (Godsey et al., 2009; Thompson et al., 2011).

Another method for quantifying the timing of solute delivery is through the creation of a dimensionless plot—called an LV curve—of cumulative pollutant load *versus* cumulative runoff volume for each event (Figure 1) (Bertrand-Krajewski et al., 1998). “First flush” is the phenomenon when the majority of the total load (60%–80%) of a given pollutant is contained in the first 20%–30% of runoff volume (Saget et al., 1996; Deletic, 1998; Lee et al., 2002; Barco et al., 2008; McCarthy, 2009; Bach et al., 2010) which would plot above the 1:1 line on an LV curve (Bertrand-Krajewski et al., 1998). Some studies note that first flush is particular to certain contaminants (Kim and Furumai, 2016), while other studies note that first flush occurs only in small sewer catchments (Lee et al., 2002; Zhang et al., 2012) or varies on an event-by-event basis (Saget et al., 1996). The first flush phenomenon is more likely to occur in catchments with more impervious surfaces (Lee et al., 2002) and is highly dependent on the time of concentration (i.e., the length of time needed for water to flow from the most remote point in a watershed to the watershed outlet (Kang et al., 2008)). Contaminant mobilization can also follow a middle-flush pattern with continued contaminant mobilization throughout the extent of the hydrograph rise and fall (Surbeck et al., 2006; Qin et al., 2016). In this middle flush scenario, increased effective contributing area of runoff throughout a storm corresponds to an increasing number of large but diffuse catchment reservoirs of accumulated contaminants (Peter et al., 2020). Final-flush behavior is when the largest proportion of pollutant load is delivered in the latter portion of the storm event (Lee and Bang, 2000; Flint and Davis, 2007).



In urban systems, solute event dynamics are complicated because of the variety of potential contributing sources as well as the complex flow paths that result from stream burial, incision, and storm drain inputs (Kaushal & Belt, 2012). Urban solute sources include runoff from upland impervious surfaces carried through storm sewers, infiltration into pervious upland areas and subsequent transport through soil flow paths, flushing of riparian soils, leaking water and sewer infrastructure, as well as sanitary and combined sewer overflows. Variability in travel times of solutes from different sources in different parts of the watershed through a variety of transport mechanisms increases the complexity of understanding stream solute concentrations (Scholefield et al., 2005).

An additional confounding factor in urban stream solute load analysis is sewer overflows which contribute significant pollution and bypass natural catchment flowpaths that could mitigate water quality impacts. Overflows come from both combined and sanitary sewers. Combined sewers are designed so the sewage and stormwater travel in the same pipes. Under dry weather conditions, the pipes only convey sewage, and all flow is directed to the treatment plant; however, during wet weather, the additional stormwater causes the total water flow in the pipe to exceed a fixed threshold and discharges a mix of raw sewage and stormwater directly into surface water. In separated sewers, sanitary and stormwater travel in different pipes and should not interact, but there can be enough infiltration and inflow of freshwater, whether from groundwater or rainfall, to cause surcharge and discharge of the sanitary system to surface water (Kracht et al., 2007; Karpf and Krebs, 2011; Bhaskar and Welty, 2012; Pangle et al., 2022).

Both sanitary sewage and stormwater contribute nitrate (NO_3^-) to urban streams and can be differentiated using natural abundance stable isotope tracers (Kendall et al., 2007). In particular, $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of NO_3^- can indicate both nitrogen sources and some aspects of terrestrial nitrogen cycling (i.e., denitrification). In comparison, mass-independent $\Delta^{17}\text{O}$ of NO_3^- is not affected by denitrification and can thus be used to precisely quantify

contributions of unprocessed atmospheric NO_3^- to surface and groundwater (Michalski et al., 2003). All terrestrial systems have a constant ratio between $\delta^{17}\text{O}$ and $\delta^{18}\text{O}$ due to the mass dependence of kinetic and equilibrium fractionations. Because all terrestrial nitrate sources follow this mass-dependent relationship ($\delta^{17}\text{O} = 0.52 \times \delta^{18}\text{O}$), they have a characteristic slope called the “terrestrial fraction line”. Deviations from this relationship are known as “mass-independent fractionations” and are conservative tracers indicative of the proportion of unaltered atmospheric NO_3^- . This mass-independent relationship is imparted to nitrate molecules during interactions in the atmosphere with ozone that lead to excess ^{17}O (Michalski et al., 2003). The anomaly stemming from interactions with ozone between $\delta^{17}\text{O}$ and $\delta^{18}\text{O}$ is known as $\Delta^{17}\text{O}$, where $\Delta^{17}\text{O} = \delta^{17}\text{O} - (\delta^{18}\text{O} \times 0.52)$. Positive $\Delta^{17}\text{O}$ values are unambiguous tracers of atmospheric deposition in terrestrial and aquatic systems (Michalski et al., 2003; Riha et al., 2014; Rose et al., 2015). $\Delta^{17}\text{O}$ presents the opportunity to use a two-end-member mixing model to partition atmospheric NO_3^- (i.e., stormwater runoff) from human and terrestrial sources (i.e., sewage). In contrast, quantification of atmospheric nitrate using the $^{18}\text{O}/^{16}\text{O}$ abundances of nitrate, while easier to measure, are highly uncertain due to large ranges in the isotopic composition of ^{18}O in many forms of reactive nitrogen compounds in the atmosphere (Elliott et al., 2019). In urban streams, the potential terrestrial sources are human sewage, fertilizer, and animal waste (e.g., domestic dogs). Sewage is a common nitrate source to urban streams due to pipe leaks and overflows (Kaushal et al., 2011; Divers et al., 2014). While fertilizer can be a significant contribution to total NO_3^- load in arid urban watersheds (Hale et al., 2014), fertilizer is often hard to detect in watersheds with more frequent precipitation because it does not have much opportunity to build up on the landscape in between storm events. (Elliott and Brush, 2006; Anisfeld et al., 2007; Divers et al., 2014). Pet waste can also be a NO_3^- source to urban and suburban watersheds (Hobbie et al., 2017), but is generally indistinguishable from sewage using $\delta^{15}\text{N}$. Because of prior work in this watershed, both isotopic and inverse modeling, we know that the mass-dependent source is dominated by sewage (Divers et al., 2013; 2014).

While numerous studies have investigated storm event solute loads from urban watersheds (Thompson et al., 2011; Carey et al., 2014; Duncan et al., 2017; Blaszcak et al., 2019), the influences of stream channel burial and sewer overflows are often overlooked, despite the fact that these infrastructure challenges have drastic effects on the routing of stormwater flow, the timing of event responses, and the hydrologic connections between solute sources and the stream (Forgrave et al., 2022). This study uses grab samples from different storm event sizes to highlight unique solute delivery patterns in a buried urban stream. Specifically, we investigate 1) how do solute concentrations and loads change throughout storm events? and 2) how can triple NO_3^- isotope analyses be used to distinguish timing and load percentages of atmospheric and sewage-derived sources? We expect to see a difference between a first flush of runoff chemistry such as atmospheric NO_3^- and road-related solutes such as chloride from road salt, and then potentially a second flush of sewage-related solutes from overflows after pipes exceed capacity. This research will reveal the storm event-driven solute dynamics of buried urban streams and highlight the important impacts to watershed connectivity and biogeochemical loads. To reduce

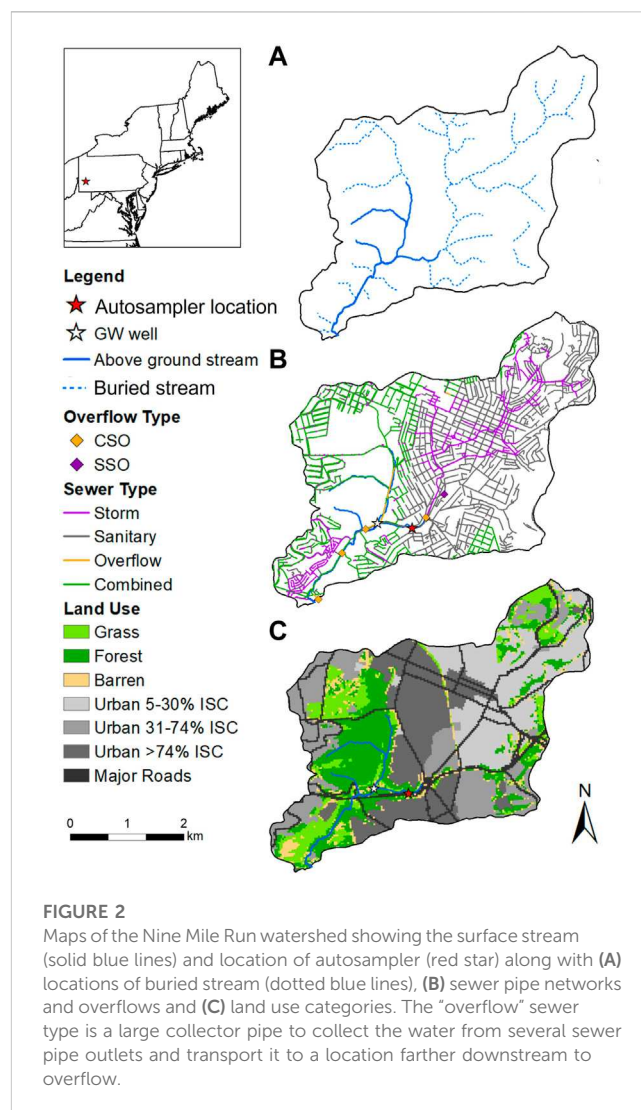


FIGURE 2

Maps of the Nine Mile Run watershed showing the surface stream (solid blue lines) and location of autosampler (red star) along with (A) locations of buried stream (dotted blue lines), (B) sewer pipe networks and overflows and (C) land use categories. The “overflow” sewer type is a large collector pipe to collect the water from several sewer pipe outlets and transport it to a location farther downstream to overflow.

urban surface water contamination, these solute dynamics must be an overarching consideration in urban stream restoration activities, stormwater management, and nutrient load planning.

2 Materials and methods

2.1 Study location

Nine Mile Run drains 15.7 km² of Eastern Pittsburgh, PA and adjacent communities. The watershed is predominantly urban land use (66%); the remainder is 18% forested, 10% grassy field, and 5% barren land (PAMAP Program Land Cover for Pennsylvania, 2005) (Figure 2B). Half of the watershed is served by combined sewers and half by separated storm and sanitary sewer networks (Figure 2A). A majority of the headwaters of Nine Mile Run were buried in the 1920s (Historic Pittsburgh Maps, 2021) during a time of rapid residential development (Hopkins et al., 2013). In combined sewer systems, buried stream flow was often deliberately routed into the sewer network (Hopkins and Bain, 2018). Today, an estimated 78% of all channels in the Nine Mile Run watershed

are buried (Rivers 2nd Nature, 2006), with only the downstream portion in a restored surface channel (Figure 2). This restoration project was completed in 2006 and cost over \$7.7 million (2007 dollars) to restore the stream channel with the goal of promoting ecological recovery (Bain et al., 2014). However, the restoration effort was restricted to the stream valley itself without focusing on the upstream watershed causes of stream impairment and thus did not mitigate large storm flows (Forgrave et al., 2022).

In addition to flooding issues, Nine Mile Run is strongly influenced by sewage inputs at both low flows and storm events. Previous research using inverse modeling and dual NO_3^- isotope analysis on biweekly grab samples from Nine Mile Run indicates that up to 94% of the NO_3^- is sewage-derived at baseflow conditions, and 66% of during storm events (Divers et al., 2014). The pollution is a result of leakage from a sewage system developed in the early 1900s made of concrete, brick, and vitrified clay pipes installed within the stream channel (Hopkins and Bain, 2018). In addition to leaks, Nine Mile Run receives inputs from four combined sewer overflows (CSOs) and one sanitary sewer overflow (SSO) (Figure 2A), however, there are an additional 38 maintenance holes for access to the main trunk sewer that runs alongside the stream. Several of these maintenance holes routinely overflow during storm events, and in some instances, the maintenance hole cover is completely blown off during a high flow storm event. The City of Pittsburgh is in the top three cities in the U.S. for highest number of CSOs (EPA, 2004), and the overflows can be triggered by precipitation events as low as a tenth of an inch (ALCOSAN, 2012). While sanitary sewers are, in theory, not supposed to be connected to the storm drainage system, the authors have observed and the local watershed organization has acknowledged SSOs overflowing during storms, indicating significant stormwater flow in the sewer pipes and flow that exceeds capacity, resulting in overflow to the stream (Rivers 2nd Nature, 2004; UpstreamPGH, 2020).

2.2 Sample collection and laboratory analysis

In a plunge pool 300 m downstream of where the stream first emerges from underground pipes, we installed an ISCO autosampler (Teledyne Isco, Lincoln, NE) secured in a metal barrel and activated through a cell phone modem. The autosampler was programmed to collect grab samples at 15-minute intervals during storms, and then switched to every 30 min after the storm to collect the hydrograph recession. Water samples were collected from the autosampler within six hours of the end of collection and brought to the lab for filtering and partitioning.

One subset of each sample was vacuum-filtered through 0.45 μm filters and frozen for analysis of nitrogen species, anion concentrations, and NO_3^- isotopes. Nitrate (NO_3^-), nitrite (NO_2^-), ammonium (NH_4^+), and orthophosphate (PO_4^{3-}) concentrations were measured on a Lachat Quickchem 8,500 Flow Injection Analysis (Hach). Fluoride (F^-), chloride (Cl^-), and sulfate (SO_4^{2-}) ion concentrations were measured on a Dionex ICS2000 Ion Chromatograph. Nitrate isotopes ($\delta^{15}\text{N}-\text{NO}_3^-$, $\delta^{18}\text{O}-\text{NO}_3^-$) were measured on an Isoprime Continuous Flow Mass Spectrometer following the denitrifier method (Sigman et al., 2001). Samples were corrected using international reference standards

USGS-32, USGS-34, USGS-35, and IAEA-N3; these standards were also used to correct for linearity and instrument drift. Samples with more than 2% nitrite of the total oxidized nitrogen ($\text{NO}_2^- + \text{NO}_3^-$) had the nitrite removed through sulfamic acid reduction (Granger and Sigman, 2009) prior to isotopic analysis. $\Delta^{17}\text{O}-\text{NO}_3^-$ values were analyzed similarly to $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values but with the additional step of the N_2O gas heated to 800°C in a gold tube where it decomposed into N_2 and O_2 gases (Kaiser et al., 2006) prior to analysis by the IRMS. The remainder of each sample was left unfiltered, frozen, and then later analyzed for total nitrogen and total phosphorus on the Lachat combined with a persulfate pre-digestion (Hosomi and Sudo, 1986).

Water level was collected with a HOBO pressure transducer (Onset Computer Corporation, Bourne, MA) in a stilling located less than 100 m upstream from the autosampler location, between where the stream emerges and the plunge pool. The stage-discharge relationship has already been established and was shown to be consistent with the USGS gage (03085049) that existed in this watershed from 2006 to 2009 (Kolmogorov-Smirnov test, $p < 0.0001$).

2.3 Solute load calculations and LV curves

The load is the total mass in kilograms of a given solute that is mobilized in a given storm event. Loads of each solute for each storm were calculated by multiplying the concentration (in mg/L) by the flow (in m^3/s) for each sample, then multiplying by the total number of seconds between each sampling interval. Cumulative mass load percent for a given solute was compared to cumulative flow volume percent at each sample interval to create an LV curve (Bertrand-Krajewski et al., 1998). For this study, the numeric thresholds used to define first flush were 80% of the total solute load delivered in the first 30% of the storm event, and 80% of the total load not delivered until 80% of the total event flow to be “final flush” and everything in between as “middle flush”. Due to fast-rising limbs observed there are large jumps in cumulative discharge, increasing by nearly 40% in a single 15-minute time interval.

For some storm events, the autosampler did not capture the entirety of the rise and recession due to large storm size, imprecise weather predictions, or technical difficulties in remotely triggering autosampler initiation. In these cases, the calculated event loads only account for the proportion of the total event that was successfully collected via the autosampler. We recognize that missing a portion of the storm in the data could potentially bias the LV analysis and have noted on figures when part of the storm event is missing. However, with the number of storms and the fact that different portions are missing in the different events collected, the data still allows for generalized findings on the average solute load response across a range of event sizes.

2.4 Nitrate load partitioning with $\Delta^{17}\text{O}$

For each storm event collected with the autosampler, we use a two-end-member mixing model to determine the fraction of the total nitrogen load contributed by atmospheric sources and sewage (Michalski et al., 2004; Hale et al., 2014). Although prior studies have conducted source apportionment in Nine Mile Run using dual

nitrate isotopes (Divers et al., 2014) and inverse modeling (Divers et al., 2013), here we use a mass-independent anomaly of the least abundant oxygen isotope (^{17}O) in the nitrate molecule following Eq. 1 below.

$$\Delta^{17}\text{O} = \delta^{17}\text{O} - (0.52 \times \delta^{18}\text{O}) \quad (1)$$

Because $\Delta^{17}\text{O}$ is not fractionated by biological processes, it is a conservative tracer of atmospheric nitrate and can be used for more precise source apportionment relative to the $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of nitrate. This approach is particularly powerful in urban systems like Nine Mile Run that are dominated by two sources: untreated sewage and atmospheric nitrate from stormwater runoff (Divers et al., 2013; 2014). The endmembers used for load apportionment were $\Delta^{17}\text{O} = 21\text{‰}$ for atmospheric nitrogen and $\Delta^{17}\text{O} = 0\text{‰}$ for sewage. The atmospheric value is the average from locally-collected, event-integrated precipitation from 43 events over a full year [$\Delta^{17}\text{O} = +20.9 \pm 3.8\text{‰}$] (Supplementary Figure S1). To account for the extreme within-event variability reported for $\Delta^{17}\text{O}$ values of precipitation nitrate (Rose et al., 2019), we designed a large-volume precipitation collector to sample throughout the entire precipitation event and measure volume-integrated $\Delta^{17}\text{O}$ values. The mean $\Delta^{17}\text{O}$ value and range from all precipitation samples collected falls within the range of published values (+20 to +30‰, Michalski et al., 2003). The variability in this end member for atmospheric $\Delta^{17}\text{O}$ is due to seasonal differences in photochemical oxidation pathways (Michalski et al., 2003). $\Delta^{17}\text{O}$ values in precipitation sample collected were seasonally variable, however, there is also considerable event-to event variation within a given season, so we chose to represent atmospheric nitrate inputs with the average $\Delta^{17}\text{O}$ value for the precipitation end member across all sampled storms to keep a consistent method across all events collected. Based on the standard deviation of observed $\Delta^{17}\text{O}$ values ($\pm 3.8\text{‰}$) across storm events, the maximum error introduced by using an average $\Delta^{17}\text{O}$ value as the atmospheric end member is less than 5% of the load attributed to the atmospheric sources.

The atmospheric fraction (f_{atm}) of the nitrate load for each collected water sample was determined with Eq. 3 and the fraction from sewer (f_{sewer}) from Eq. 4. The total NO_3^- load was calculated as the NO_3^- concentration multiplied by discharge and these two fractions to calculate total load (in mass of $\text{NO}_3\text{-N}$) from each source Eqs 5, 6.

$$\text{sample} = f_{\text{atm}} + f_{\text{sewer}} \quad (2)$$

$$f_{\text{atm}} = \frac{\Delta^{17}\text{O}_{\text{sample}}}{20.9\text{‰}} \quad (3)$$

$$f_{\text{sewer}} = 1 - f_{\text{atm}} \quad (4)$$

$$N\text{Load}_{\text{atm}} = [\text{NO}_3^-]_{\text{sample}} \times Q \times f_{\text{atm}} \quad (5)$$

$$N\text{Load}_{\text{sewer}} = [\text{NO}_3^-]_{\text{sample}} \times Q \times f_{\text{sewer}} \quad (6)$$

3 Results

3.1 Timing of solute changes

We collected a total of five storm events through 2019 and 2020, for four of these events (Storms 1, 3, 4 and 5), samples were collected at 15-minute time intervals for the whole event and sample collection for Storm 2 started at 15-minute and then switched to 30-min to spread out the samples and capture the complete recession (Table 1). Three storms (Storms 2, 3, and 5) had multiple discharge peaks throughout the event corresponding to multiple peaks in precipitation, while the other two summer storms (Storms 1 and 4) had a single discharge rise and recession. Below we briefly describe the hydrological and chemical characteristics of each of the five storms where these differences are compared across storms in Figure 3.

Storm 1 discharge peaked at 5.25 m³/s, rising to this peak discharge in 30 min, then had a secondary peak in flow an hour later after another pulse of precipitation. Only nine grab samples were collected during Storm 1 because the autosampler started collecting before the storm event occurred, thus missing most of the recession. As discharge increased, NO_3^- , Cl^- , F^- , and SO_4^{2-} concentrations rapidly decreased, while PO_4^{3-} concentrations increased (Figure 3). NH_4^+ initially increased at the beginning of the storm, subsequently decreased, and then increased again. $\delta^{15}\text{N}\text{-NO}_3^-$ values decreased and $\delta^{18}\text{O}\text{-NO}_3^-$ values increased along with the discharge (Figure 3).

Storm 2 was an extended rainfall event over several hours that produced several consecutive discharge peaks where the largest peak was 11.6 m³/s. After the first collection of 24 grab samples, we

TABLE 1 List of grab sample storm collections. Precipitation amounts listed are the amount directly before the peak discharge captured by the autosampler rather than the entire event from start of precipitation to finish to account for storms where it rains throughout the entirety of the event causing subsequent discharge peaks after sampling concluded.

Storm	Date	Times	Number of samples	Sampling interval (minutes)	Peak discharge (m ³ /sec)	Event precipitation (mm)	Notes
1	June 10, 2019	15:00 -19:15	18	15	5.25	7.9	Missing second half of recession
2	October 6-7, 2019	22:30-5:30, 9:15-18:15	43	15, then 30	11.66	44.1	4-hour gap in the middle of collection
3	January 11-12, 2020	20:15-2:00	24	15	1.22	3.01	
4	July 22, 2020	6:11-11:56	24	15	21.50	10.6	
5	August 28, 2020	17:42-21:57	18	15	4.06	4.4	Follows a larger event (discharge peak of 41.9 m ³ /s at 13:25)

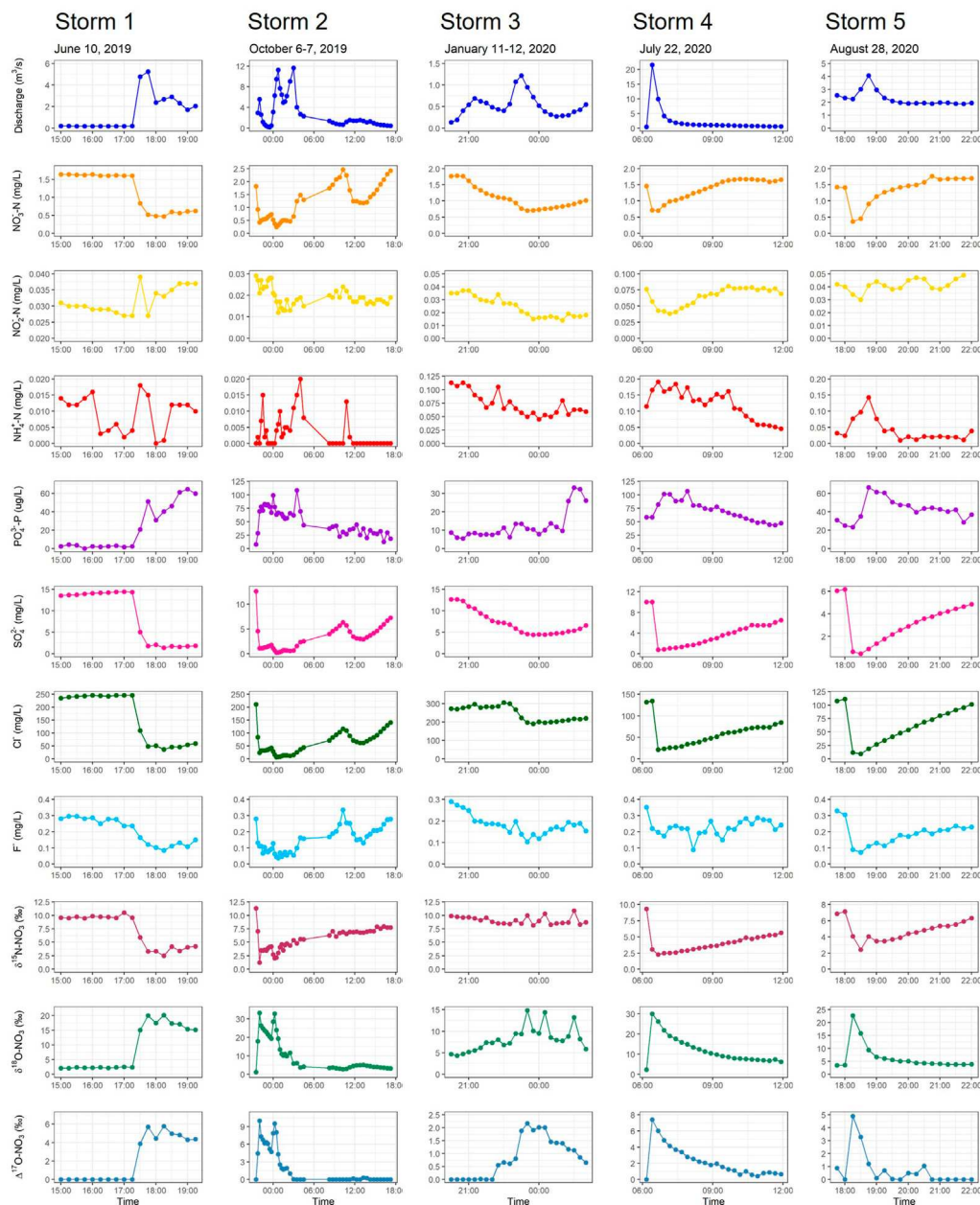


FIGURE 3

Time series plots for Nine Mile Run discharge and grab sample chemistry for five storm events. Each column is a different storm event while each row is a different chemical parameter. Note different y-axis scale ranges for different storm events.

restarted the autosampler and continued sample collection for the entire recession. However, there is a 4-hour gap during this restart process (5:15 to 9:30am). With the initial increase in flow, NO_3^- , F^- , and Cl^- concentrations decreased sharply and then decreased again with subsequent discharge peaks followed by an increase in concentration during the recession. In the case of NO_3^- and F^- , recession concentrations were higher than the pre-storm concentrations (Figure 3). PO_4^{3-} concentration increased with the initial runoff inputs and then progressively decreased throughout the rest of the sample collection. $\delta^{15}\text{N}-\text{NO}_3^-$ values decreased dramatically with the initial discharge peak and then steadily

increased throughout the rest of the flow event. $\delta^{18}\text{O}-\text{NO}_3^-$ values increased with the first two discharge peaks (2019-10-06 22:00 and 2019-10-07 0:15), then decreased back to baseflow values for the rest of the sample collection period.

Storm 3 was a small winter rain event (3.0 mm of precipitation) where the stream rose to two discharge peaks of 0.75 and 1.25 m^3/s . NO_3^- concentration slowly decreased in contrast to the abrupt changes seen in the summer and fall storms (Figure 3). NH_4^+ and F^- also decreased in concentration but not as smoothly as NO_3^- . PO_4^{3-} concentration slowly increased throughout most of the storm, but then increased suddenly at the end of the event. Cl^-

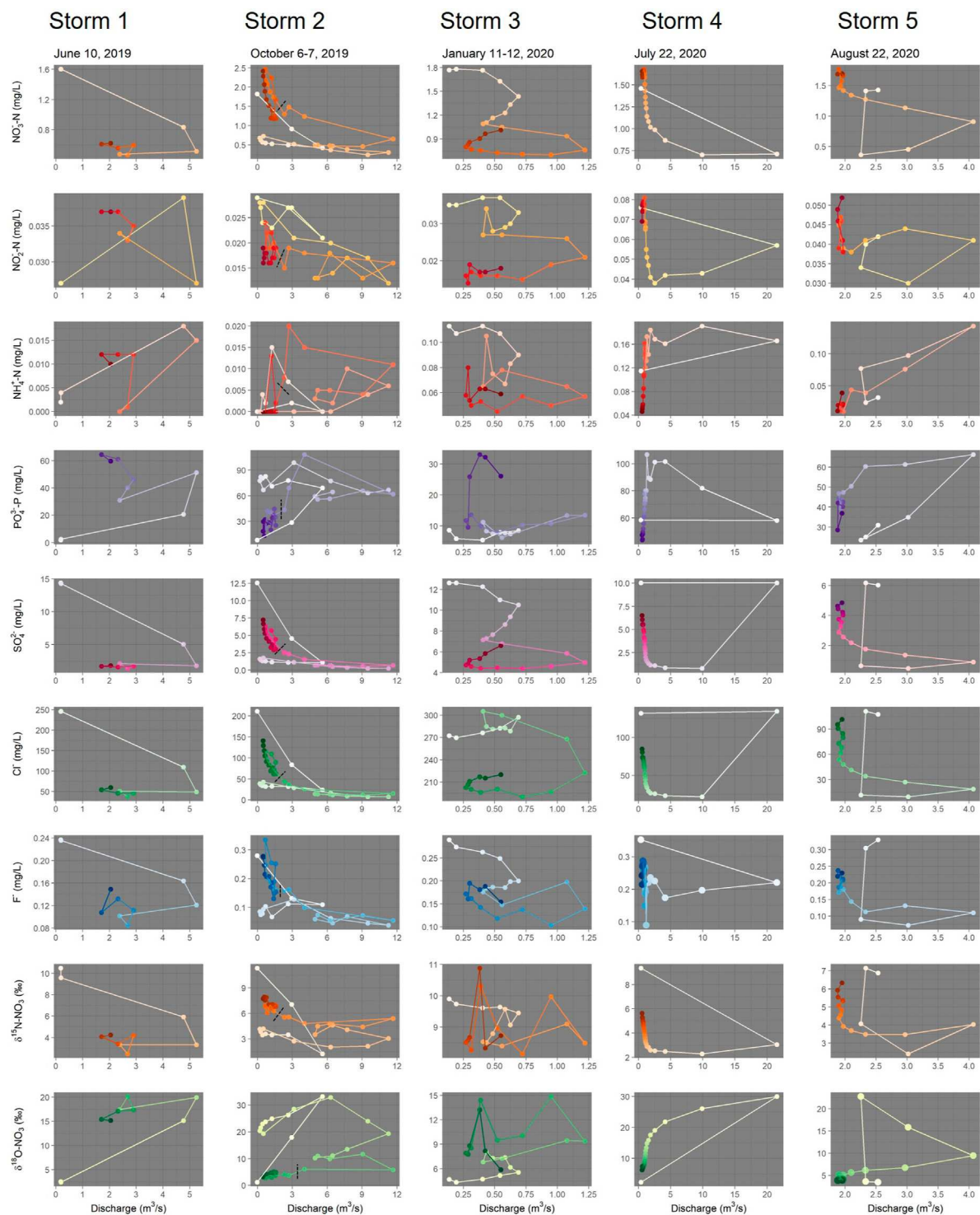


FIGURE 4

Hysteresis for solutes collected by the autosampler for each storm event. Each column is a different storm event while each row is a different chemical parameter; the color goes light to dark over time with white at the beginning of the event and the darkest color at the end. Note different axis scales for different storm events. Black dotted lines on Storm 2 indicate the 4-hour gap in collection.

concentration increased slightly with the first discharge peak and then decreased with the second peak. $\delta^{15}\text{N}$ values initially decreased slightly but then increased several times in the latter half of the event.

$\delta^{18}\text{O}-\text{NO}_3^-$ values followed a similar pattern, but the initial change was a gradual increase in $\delta^{18}\text{O}-\text{NO}_3^-$ values followed by three more increases in $\delta^{18}\text{O}-\text{NO}_3^-$ values later in the event (Figure 3).

TABLE 2 Solute loads for all storm events collected with the autosampler.

Storm	NO ₃ ⁻ -N (kg)	NO ₂ ⁻ -N (kg)	NH ₄ ⁺ -N (kg)	PO ₄ ³⁻ -P (kg)	F ⁻ (kg)	Cl ⁻ (kg)	SO ₄ ²⁻ (kg)
1	13.4	0.8	0.2	1.0	2.8	1,392.0	56.8
2	158.0	3.0	0.7	9.4	18.9	6,394.3	326.1
3	11.6	0.3	0.8	0.1	1.9	2,740.9	76.1
4	49.2	2.8	8.0	3.7	10.9	3,828.5	263.7
5	47.9	1.5	1.8	1.6	6.6	2,135.3	110.5

TABLE 3 Solute loads for all storm events normalized by event duration.

Storm	NO ₃ ⁻ -N (kg/hr)	NO ₂ ⁻ -N (kg/hr)	NH ₄ ⁺ -N (kg/hr)	PO ₄ ³⁻ -P (kg/hr)	F ⁻ (kg/hr)	Cl ⁻ (kg/hr)	SO ₄ ²⁻ (kg/hr)
1	6.7	0.4	0.1	0.5	1.4	696.0	28.4
2	8.0	0.2	0.04	0.5	1.0	323.8	16.5
3	1.9	0.0	0.1	0.0	0.3	456.8	12.7
4	8.2	0.5	1.3	0.6	1.8	638.1	44.0
5	10.6	0.3	0.4	0.4	1.5	474.5	24.6

Storm 4 was an intense summer storm that rose to a peak discharge of 21.5 m³/s in only 15 min in response to 0.42 mm of precipitation (Figure 3). NO₃⁻ and Cl⁻ concentrations decreased rapidly as streamflow increased. F⁻ concentrations decreased by 0.1 mg/L in response to the initial input of stormwater but then stayed fairly constant for the remainder of the event. PO₄³⁻ and NH₄⁺ concentrations both sharply increased with increasing discharge and then decreased gradually throughout the rest of the event. δ¹⁵N-NO₃⁻ values decreased sharply and δ¹⁸O-NO₃⁻ values increased sharply along with the sudden discharge increase and then both values slowly returned to pre-storm conditions as the storm peak flow receded.

Storm 5 was another summer storm where the autosampler collected the secondary storm peak (4.1 m³/s) after a much larger initial discharge rise of 41.9 m³/s (Figure 3). Total precipitation for the whole event was 52.0 mm while the precipitation directly before ISCO collection was only 4.4 mm. Concentrations of NO₃⁻, Cl⁻, SO₄²⁻, and F⁻ decreased then slowly increased throughout the recession. PO₄³⁻ and NH₄⁺ concentrations increased with streamflow and decreased afterward. δ¹⁵N-NO₃⁻ values decreased and δ¹⁸O-NO₃⁻ values increased along with discharge, but δ¹⁸O-NO₃⁻ values returned to pre-storm conditions before δ¹⁵N-NO₃⁻ values.

Most solutes show clockwise hysteresis patterns in Storms 1, 3, and 4 (Figure 4) suggesting sources that are quickly connected to the stream via fast flow pathways such as pipes, with the exceptions of PO₄³⁻ which is counter-clockwise in these events and NH₄⁺ which is counterclockwise in Storm 4. All solutes in Storm 2 have a complex hysteresis with multiple loops in different directions corresponding to the several discharge peaks throughout the storm event. For Storm 5, most solutes show a “figure-8” pattern with an initial counterclockwise solute dilution, followed by concentration increase during discharge recession. This is likely because Storm 5 is a smaller storm directly after a much larger storm that was missed by the

autosampler so the initial start of the plotted hysteresis loop is affected by the recession of the previous storm.

3.2 Event solute loads

Event solute loads varied widely across the five storm events for each solute, but the largest event-to-event variation was in PO₄³⁻, where loads were 88x larger in Storm 2 than in Storm 3. The highest overall load of NO₃⁻, PO₄³⁻, F⁻ and Cl⁻ was exported from the longest storm (Storm 2), while the largest mass of NH₄⁺, NO₂⁻, and SO₄²⁻ was in the storm with the highest peak discharge (Storm 4) (Table 2; Table 3).

In this study, the timing of solute loads varied by event and parameter (Figure 5). Two notable solutes are Cl⁻ and SO₄ in Storm 4 where 69% and 75% of the mass load, respectively, are delivered in the first 15 min of the storm event which also coincides with 40% of total event runoff, but this is still categorized as middle flush. Conversely, the NO₂⁻ and NO₃⁻ loads are delivered more towards the end of Storm 4 where 80% of the total mass load for NO₂⁻ and NO₃⁻ is delivered at 86% and 89%, respectively, of the total event runoff. In other events with less rapid increases in discharge compared to Storm 4, SO₄²⁻ and Cl⁻ still plot above the 1:1 line indicating solute delivery that slightly precedes the comparable percentage of total water volume. In Storms 1 and 3, all solutes except PO₄³⁻ have cumulative load increase faster than cumulative discharge. In Storms 2 and 5, which have multiple discharge peaks, solute load delivery changes mid-storm between the multiple discharge peaks. In Storm 2, there is a first flush of SO₄²⁻ and Cl⁻ and then they lag flow for the rest of the event, while NO₃⁻ is equal to flow for the first 10% of the event and then lags after stormwater volume for the rest of the event (Figure 5). PO₄³⁻ cumulative mass lags cumulative flow in all storm events except for the second half of Storm 5.

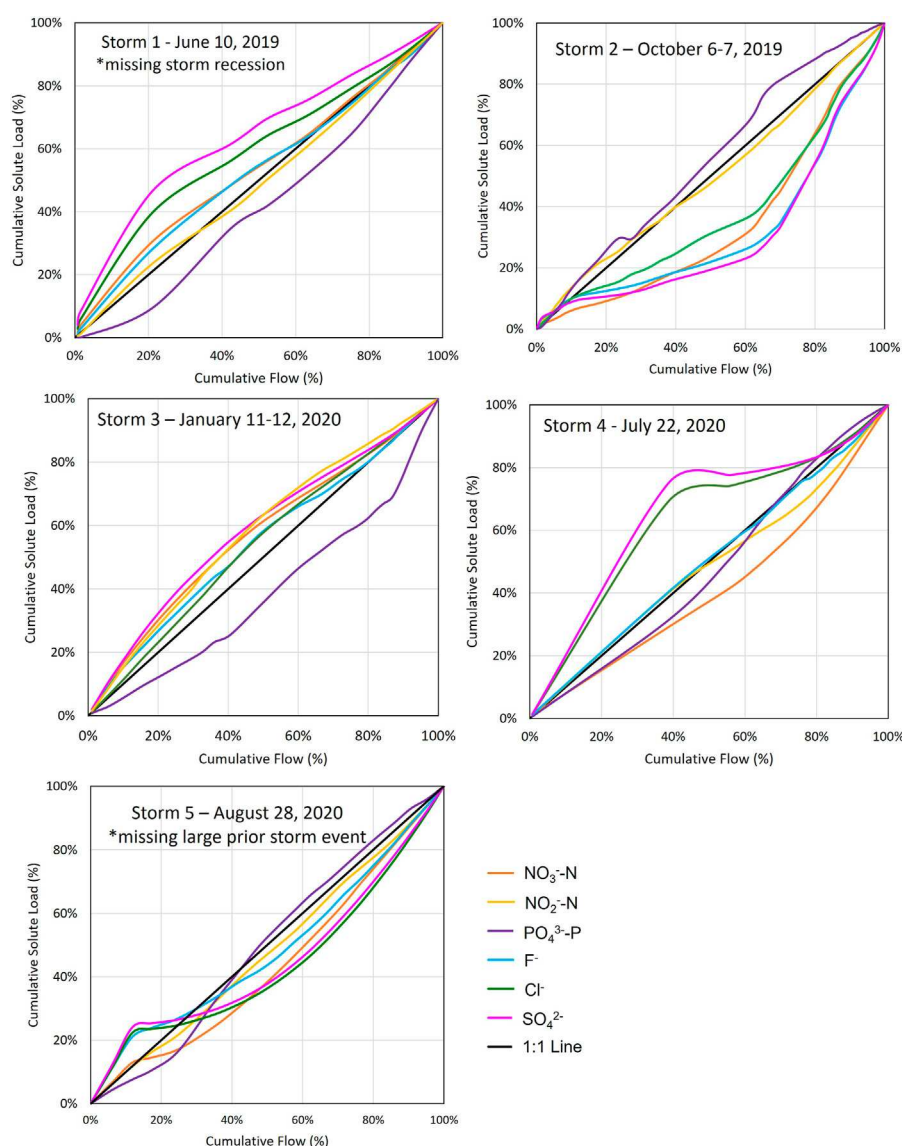


FIGURE 5

LV curves for each storm event collected by the autosampler. The 1:1 line indicating uniform solute delivery is in black. The intercept of these plots represents beginning of autosampler collection, not necessarily beginning of storm flow.

3.3 Nitrate source partitioning

For all grab samples, $\delta^{15}\text{N}-\text{NO}_3^-$ and $\delta^{18}\text{O}-\text{NO}_3^-$ values in the pre-event sampling (after autosampler initiation but before discharge rise) are within the range of $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of sewage-derived NO_3^- , while $\Delta^{17}\text{O}-\text{NO}_3^-$ are zero. In the beginning of each event, $\delta^{15}\text{N}-\text{NO}_3^-$ values decreased while $\delta^{18}\text{O}-\text{NO}_3^-$ and $\Delta^{17}\text{O}-\text{NO}_3^-$ values increased as impervious surface runoff delivered atmospherically deposited NO_3^- , and then returned to a predominantly sewage source for the rest of the event (Figure 6). Although the NO_3^- load from the atmosphere ranges from 2.3% to 23.6% of total NO_3^- load across the five storm events as indicated by the $\Delta^{17}\text{O}-\text{NO}_3^-$ partitioning, (Table 4; Figure 7), the majority of the NO_3^- load exported from all storms is sewage (76%–98%). The lowest atmospheric NO_3^- contributions occurred during the

longest-duration storms characterized by several discharge peaks (Storms 2, 3, and 5). This is because the majority of the atmospheric NO_3^- is delivered from the first peak from rain washing atmospherically deposited dry N deposition off the landscape, particularly from directly connected impervious surfaces, while successive peaks have increasing percentages of sewer-derived NO_3^- . These three events with lower atmospheric deposition contribution are also the events where the autosampler missed the first peak of the storm, making the calculated mass from atmospheric deposition conservative.

More NO_3^- from atmospheric deposition is exported during storms that have a higher peak discharge ($R^2 = 0.94$); however, peak discharge is only weakly correlated ($R^2 = 0.24$) with the relative percentage of atmospheric contribution compared to sewage. One caveat to note is that the relative percentages are biased by the

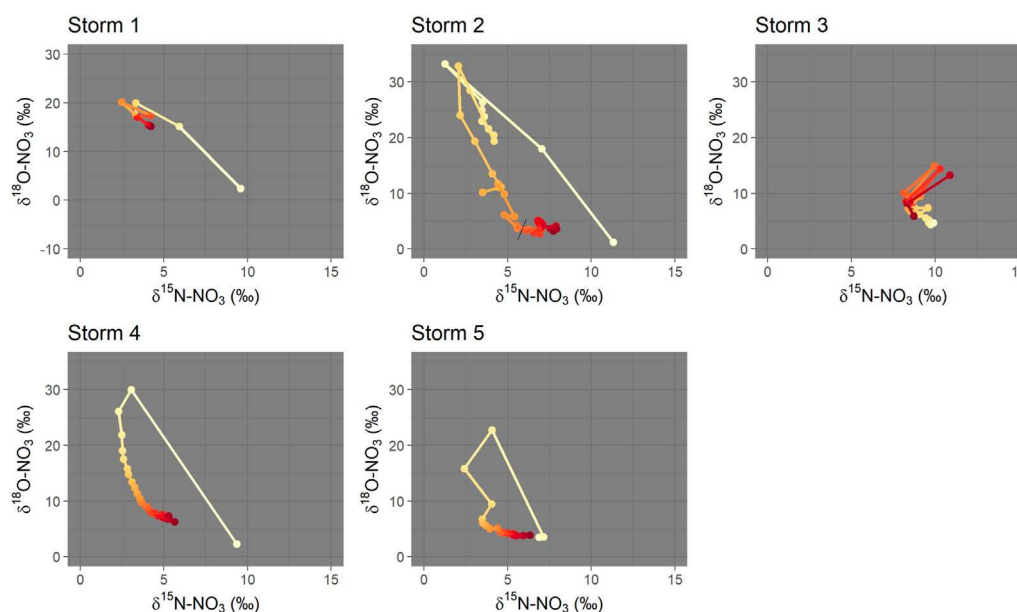


FIGURE 6

Dual plot of NO_3^- isotope values for each event collected by the autosampler. Time is represented in color from light yellow at the beginning of the storm event to dark red at the end. The black dotted line in Storm 2 indicates the 4-hour gap in sample collection.

TABLE 4 Load calculations of atmospherically deposited and sewer-derived NO_3^- from grab sample data collected by the autosampler.

Storm number	Duration (hours)	Total Mass $\text{NO}_3\text{-N}$ exported (kg)	Mass atm $\text{NO}_3\text{-N}$ (kg)	Mass sewer $\text{NO}_3\text{-N}$ (kg)	% atm $\text{NO}_3\text{-N}$	% sewer $\text{NO}_3\text{-N}$
1	2.00	13.4	3.0	10.0	22.6	77.4
2	19.75	158.0	7.0	151.0	4.4	95.6
3	6.00	11.6	0.4	11.1	3.6	96.4
4	6.00	49.2	9.8	39.4	20.0	80.0
5	4.50	47.9	1.1	46.8	2.2	97.8

portion of the storm event collected (inconsistent due to issues described previously, as well as inaccuracies in weather prediction). For example, storms when more of the recession was collected, or the storm event occurred immediately after a previous storm, have a higher percentage of sewer-derived NO_3^- and less atmospheric contribution. Due to the small number of storms collected with the autosampler, attempts to break these storms down further run into issues of statistics of small numbers and preclude quantitative evaluation of these relationships.

4 Discussion

4.1 Concentration responses to storm events

Nine Mile Run is distinct among urban streams documented in the literature for the high percentage of buried channels (Forgrave et al., 2022). This efficient drainage leads to fast runoff responses to

rainfall, and rapid dilution of baseflow chemistry. NO_3^- , NO_2^- , Cl^- , F^- , SO_4^{2-} are diluted in every event, while PO_4^{3-} always increases in concentration, and NH_4^+ responses vary among events (Figure 3). The differences in magnitude and direction of change among the various solutes demonstrate that different solute pools within the watershed are activated at different times throughout a storm event.

This dilution pattern is more apparent in the c-Q hysteresis loops (Figure 4) where the majority of solutes rotate clockwise during hysteresis, i.e., higher background concentrations are diluted by surface runoff. In other urban studies, clockwise hysteresis loops have been attributed to pollutants rapidly entering the stream from surface runoff or “new water” while counter-clockwise hysteresis implies slower groundwater inputs or “old water” (Baker and Showers, 2019; Duncan et al., 2017; Lloyd et al., 2016). The clockwise hysteresis observed in this study is consistent with fast solute inputs, but likely from both surface and subsurface flowpaths as the distinction between them is blurred by buried surface waters receiving contribution from storm drainage and considerable infiltration from shallow soil flowpaths.

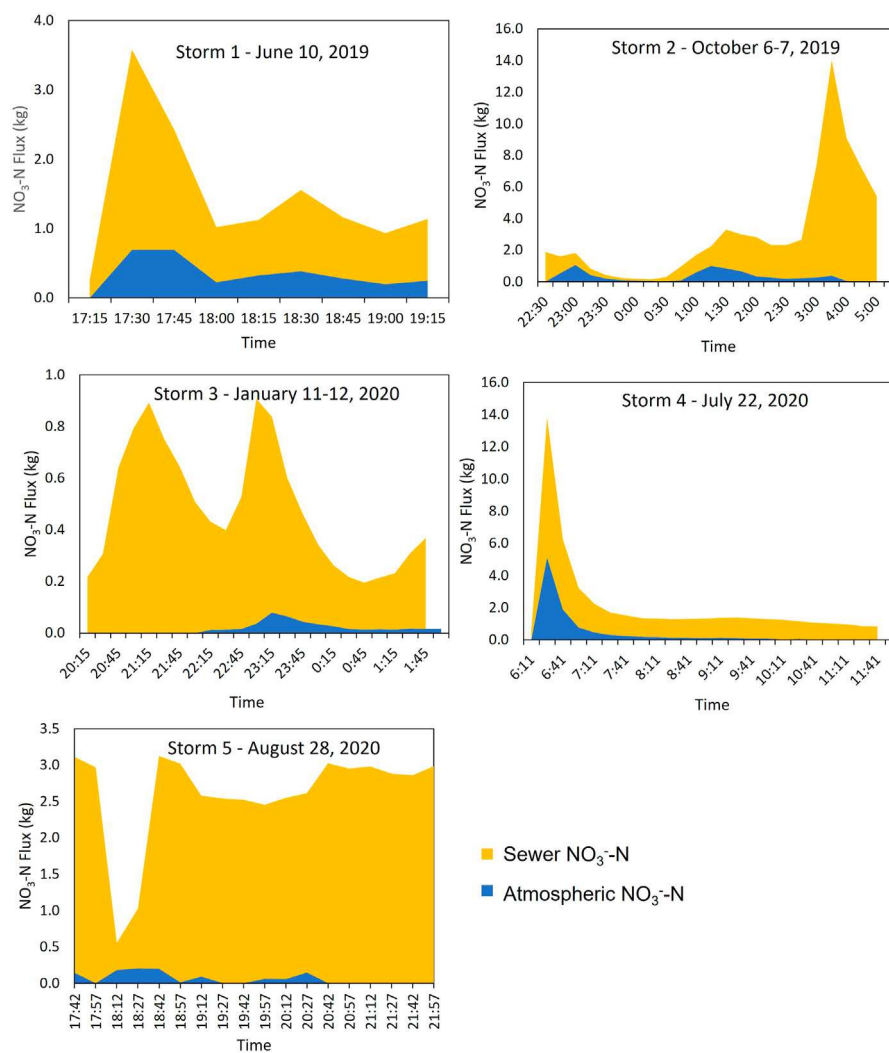


FIGURE 7

Time series plots of NO_3^- loads separated into sewer and atmospheric sources for storms collected with the autosampler. Note different y-axes ranges among storm events.

Exceptions to this overall pattern are PO_4^{3-} and sometimes NH_4^+ which show counter-clockwise flushing behavior, suggesting they are delivered from slower subsurface flowpaths that become hydrologically connected to the stream during the event (Scanlon et al., 2001; Butturini and Sabater, 2002; Duncan et al., 2017). Variable NH_4^+ event responses result from its low concentration and are harder to attribute to a particular source. Higher PO_4^{3-} concentrations can be evidence of reducing conditions as PO_4^{3-} is released during break down of Fe and Al oxides in reducing conditions (Patrick and Khalid, 1974). These increases in PO_4^{3-} correspond with higher $\delta^{15}\text{N}-\text{NO}_3^-$ values (Figure 4), which is indicative of sewage, so these reducing zones could be near the pipes and subsidized by sewer leaks, creating patches of variably saturated, hot spots of biological activity, which are then activated during a storm event (McClain et al., 2003; Bernhardt et al., 2017).

An additional complicating factor to the observed PO_4^{3-} concentration dynamics is additions of orthophosphate to the

municipal drinking water supply starting in 2019. The Pittsburgh Water and Sewer Authority started introducing 1700 $\mu\text{g/L}$ PO_4^{3-} in five locations through the water supply network to build biofilm on the interior walls of the pipes and prevent corrosion with a goal of concentrations of 500 $\mu\text{g/L}$ at the tap (Balangoda et al., 2023; submitted; Spencer-Williams et al., 2022). Due to the need to pressurize drinking water distribution systems as well as the advanced age of the pipes, this network leaks an estimated 40%–50% of its total volume, resulting in at least 28 million gallons per day of phosphate-enriched water subsidizing the groundwater and surface water every day and contributing to storm event solute loads (Balangoda et al., 2023; submitted). While we cannot definitively distinguish whether phosphate observed in storm events originates as leaks from sewer pipes, drinking water pipes, or natural sources, based on our observations in this and previous studies (Divers et al., 2014; Delesantro et al., 2022), we surmise that infrastructure leakage is the dominant source of both phosphate and ammonium.

TABLE 5 Solute loads categorized into different flushing behavior.

Storm	First flush (80% of load at < 30% of flow)	Middle flush (80% load at 30%-80 of flow)	Final flush (80% load at >80% of flow)
1	—	NH ₄ ⁺ , NO ₂ ⁻ , PO ₄ ³⁻ , F, Cl ⁻ , SO ₄ ²⁻ , Sewer NO ₃ ⁻ , Atmospheric NO ₃ ⁻	—
2	—	Atmospheric NO ₃ ⁻	NO ₂ ⁻ , PO ₄ ³⁻ , F, Cl ⁻ , SO ₄ ²⁻ -NH ₄ ⁺ , Sewer NO ₃ ⁻
3	—	NH ₄ ⁺ NO ₂ ⁻ , F, Cl ⁻ , SO ₄ ²⁻ , Sewer NO ₃ ⁻	PO ₄ ³⁻ , Atmospheric NO ₃ ⁻
4	—	Cl ⁻ , SO ₄ ²⁻ , Atmospheric NO ₃ ⁻ NH ₄ ⁺ , PO ₄ ³⁻	NO ₂ ⁻ , F, Sewer NO ₃ ⁻
5	—	Atmospheric NO ₃ ⁻ , NH ₄ ⁺	NO ₂ ⁻ , PO ₄ ³⁻ , F, Cl ⁻ , SO ₄ ²⁻ , Sewer NO ₃ ⁻

4.2 Storm event solute loads

Even as concentrations decrease for most parameters during storms, the solute loads increase because of the large volume of stormwater flowing through the stream. However, the LV curves from grab sample collections reveal that not all solute loads increase at the same time. Additionally, different sources of the same solute can respond at different times, and not all storm events have the same responses (Table 5). Sewer NO₃⁻, NO₂⁻, PO₄³⁻, and F⁻ show final flush behavior in more events, where Cl⁻, SO₄²⁻, NH₄⁺, and atmospheric NO₃⁻ are more commonly middle flush, though no solute is completely consistent across all events (Figure 5).

It is also important to note that the amount of the total rise and recession captured for any given autosampler event can introduce considerable bias to these thresholds.

No solutes exhibited first flush behavior. The closest species/storm pair was atmospheric NO₃⁻ in Storm 2 where it just barely missed threshold for first flush, reaching 80% of total mass load at 36% of the total flow. The lack of first flush is an interesting finding because it contrasts with the conceptual model of urban streams as dominated by surface runoff from impervious surfaces (Shuster et al., 2005; Schueler et al., 2009). Studies that have shown a significant first flush in urban areas have mostly been in small stormwater catchments that do not have any interaction with the subsurface (Hathaway et al., 2012; Yang and Toor, 2016). In Nine Mile Run, this rapid first addition of stormwater does not have large amounts of solutes when compared to the loads of the rest of the event. Additionally, not all the landscape area is connected to the storm drainage system (Figure 2). Other urban stream studies have made a distinction between impervious area and directly connected impervious area (Roy and Shuster, 2009; Baruch et al., 2018), a difference that is particularly relevant for buried streams where there are fewer connections between the landscape surface and the stream. Buried streams only receive surface runoff from the storm drainage network, not direct inputs from the landscape itself. Areas of land with no storm drainage (purple lines in Figure 2) will not drain to the buried stream, so a smaller portion of the total impervious surface is actually connected impervious area that could contribute to a first flush phenomenon.

The middle-flush scenario, when load increases linearly with discharge throughout the event, was the most common pattern among the different solutes and events in this study as well as others (Qin et al., 2016; Peter et al., 2020). One likely theory for the dominance of middle flush responses in urban watersheds is that pervasive contaminants exist in the urban subsurface in nearly

unlimited quantities compared to the runoff volumes (Peter et al., 2020). As some solute pools become depleted, the increase in effective contributing area can activate additional solute reservoirs throughout the watershed (Qin et al., 2016; Knapp et al., 2022).

Storms with a longer portion of the total collected event as recession (Storms 2 and 4) tend to have a higher percentage of sewer-derived NO₃⁻ as well as more solutes exhibiting final flush behavior. In this case, a longer recession results in storm events where more of the water samples are collected are on the falling limb of the hydrograph. It is possible that these patterns of final flush and increased sewer NO₃⁻ are in all events, but this behavior was missed due to truncated recession from limited grab sampling capacity. The impact of the increased solute load is seen when comparing Storm 3 and Storm 4 which have the same overall duration of collection time, but for Storm 3, the flow fluctuates through multiple storm peaks, while for Storm 4, there is one single peak and then the remainder of the event is recession. These storms with longer recessions also have higher loads of PO₄³⁻, Cl⁻, and F⁻, supporting relatively higher wastewater contributions as they are all additives to drinking water, and in the case of PO₄³⁻ also present in both human waste and detergents, and thus present in wastewater.

Deconvolving these patterns involves both spatial and temporal dimensions due to the similarity in chemistry among the various solute pools as well as spatial connectivity among flowpaths. For example, PO₄³⁻ can come from organic matter decomposition in soils, flushing of reduced soils, drinking water leaks, or sewer pipe leaks, and these various pools contribute different amounts throughout a storm event. While atmospheric NO₃⁻ and sewer-derived NO₃⁻ can be distinguished with nitrate isotopes, both soil-mediated sewer leaks and direct inputs from overflows are characterized by higher δ¹⁵N-NO₃⁻ values. Additionally, the various flowpaths are spatially intertwined such as in the case of combined sewer pipes intentionally designated to carry both stormwater and sanitary sewage, but also through unintentional connections such as exfiltration and infiltration (Bhaskar and Welty, 2012; Pangle et al., 2022). This pipe exfiltration is even more significant during storm events when sewer pipes are at capacity.

4.3 Sewage-derived nitrate is pervasive even in the absence of overflows

Part of the timing differences in NO₃⁻ load delivery is due to location and the relative contribution from multiple sources among

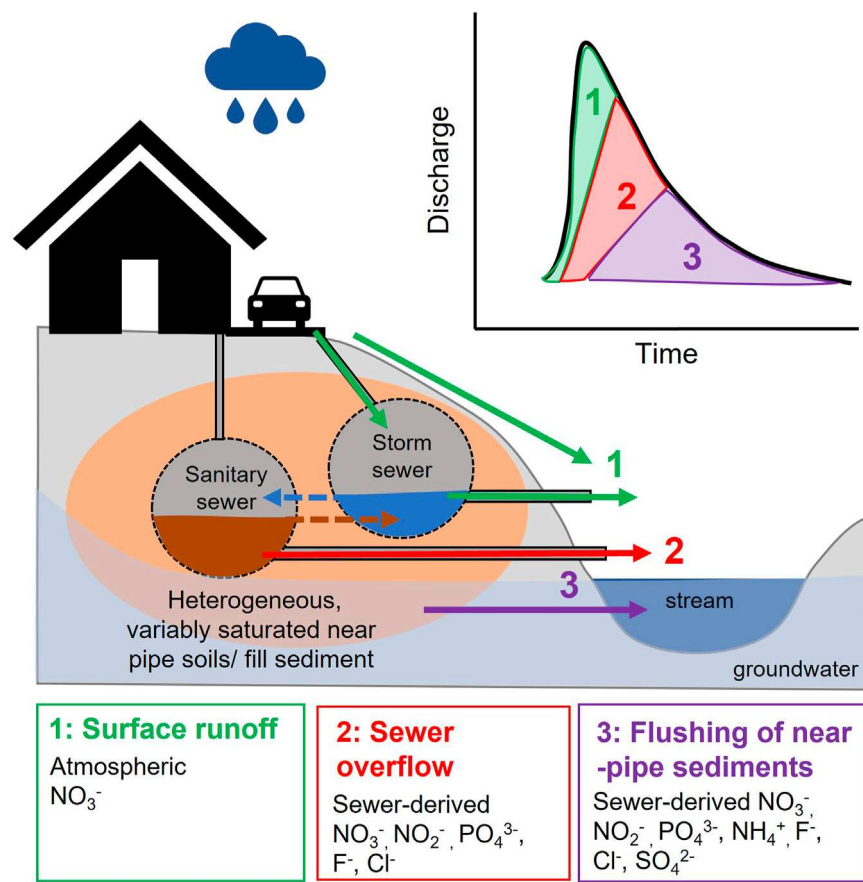


FIGURE 8

Conceptual diagram that shows the main solute pools in urban watershed and the sequential and overlapping timing of the flowpaths connecting these solute pools to the stream throughout a storm event. While the sewer pipes are shown at similar depths, this is not meant to portray actual pipe locations or scale.

events. Dual NO_3^- isotope data indicate a mix of sewer and atmospheric sources throughout each storm event (Figure 5), and atmospheric nitrate loads based on $\Delta^{17}\text{O}$ measurements suggest that sewer NO_3^- contributions are pervasive throughout all storm events and are at least 75% of the total NO_3^- load in any storm event.

Storms following a relatively dry period (Storms 1 and 4) have more atmospheric deposition sourced NO_3^- in the storm load because during the dry period, reactive nitrogen particulates and aerosols accumulate on landscape surfaces before washing off into the stream during the storm event (Riha et al., 2014) (Figure 3; Table 3). Storms with a long recession (Storms 2 and 4) or those immediately following another event (Storms 3 and 5) have a higher percentage of sewage-derived NO_3^- because once the atmospheric deposition source has been depleted, sewage-derived NO_3^- —which is an effectively unlimited source—dominates (Figure 3; Table 3). Additionally, storm events with a longer duration of precipitation, take more time to return to baseflow conditions and thus have a longer period of time in saturated soil conditions whereby wastewater is more directly and quickly conveyed to the stream.

It is surprising that there were no thresholds in event size or intensity that could be connected to the distinctive chemistry of an active combined sewer overflow. This contradicts our hypothesis,

where we expected to see a distinction between runoff chemistry and sewage overflows. While it is possible that more distinct thresholds of chemical change would be more apparent if more events were sampled, the solute event loads we observed in this study show flushing of various subsurface solute reservoirs with considerable sewer influence in all sizes of storms in all seasons.

As other studies have noted, the connections between sewer systems and surface waters vary in both space and time (DeSilva et al., 2005; Delesantro et al., 2022). Breaks and fractures in the sewer network are dependent on pipe age, pipe material, topography, intrusion by tree roots, and other disturbance factors. The height of the water table relative to the sewer pipes is also very important for determining how much infiltration or exfiltration occurs from these sewer pipe fractures (Divers et al., 2013; Duncan et al., 2017). This connectivity changes over time due to seasonal changes in water table elevation as well as changes in the pipe such as pipe fractures becoming blocked with debris and no longer leak, or debris can be washed away by heavy storms and increase leakage (DeSilva et al., 2005). These factors combine to create the event-specific solute behavior observed in this study where the majority of nitrate is sewer-derived but not necessarily from direct contributions from a sewer overflow.

5 Conclusions

Solute loads in buried urban streams respond rapidly and variably to storm events. In this study, by considering both solute concentrations and load patterns, we clarify the timing and delivery mechanisms associated with multiple solutes (Figure 8). Grab samples from storm events of different sizes and seasons reveal three different solute flowpaths that vary in relative importance through the course of a storm event: 1) elevated background concentrations of NO_3^- , NO_2^- , Cl^- , F^- , SO_4^{2-} are diluted by runoff during storm events, and increased PO_4^{3-} concentrations and $\delta^{15}\text{N}$ - NO_3^- values during storms potentially arise from (2) direct inputs from sewer overflows as well as 3) flushing of biochemically active reducing soil zones near sewer pipes. Nitrate isotopes further reveal that NO_3^- responses are driven by inputs of accumulated atmospheric deposition in the early parts of the storms, whereas wastewater-derived NO_3^- is present throughout the storm events and especially dominant during storm recession. Solute load responses also reflect the dominance of hydrologic control on solute loads as seen by all solutes showing middle flush or final flush delivery patterns. It is important to note that while we present these different flowpaths as distinct, as shown in Figure 8, they are not completely sequential, as flushing of near-pipe reducing zones and direct sewer overflow inputs tend to occur at the same time in storm events. These two subsurface flowpaths are coupled due to infiltration of shallow subsurface event water into sewer pipes triggering overflow events as well exfiltration of sewer solutes that will drain through the shallow subsurface during storm recession or remain in the near-pipe sediments to be flushed in subsequent storm events.

The results presented here expand our understanding of the storm event solute transport in a largely buried urban stream network, through the multitude of solutes and isotopes measured and analysis of cumulative loads. Despite their prevalence in urban areas (Napieralski and Carvalhaes, 2016), buried streams are understudied due to the difficulty of accessing and sampling. The unique situation of Nine Mile Run where 98% of the original stream channels are buried allows for transferrable insights into how buried streams interact with urban infrastructure and the effects of this hydrologic connectivity on downstream loads. While non-buried urban streams also have significant inputs from sewer infrastructure, the deliberate location of stream underground makes these connections more efficient and likely to occur in a wider range of event conditions. Previous research on solute transport in buried stream channels has emphasized that burial decreases nutrient uptake, thus increasing downstream loads (Beaulieu et al., 2014; Pennino et al., 2014; Beaulieu et al., 2015). This research builds on this finding to demonstrate that, in addition to decreased uptake, stream burial also increases the hydrologic connections to additional solute sources from runoff and leaking sewer infrastructure, resulting in substantial impacts on downstream ecosystems.

The biogeochemical dynamics revealed in this study highlight the importance of subsurface solute pathways to urban surface water quality, especially in watersheds with aging pipe infrastructure creating opportunities for sewage leakage and biogeochemical transformations (Delesantro et al., 2022). In the complicated, multiple-stakeholder, stormwater management environment, these fundamental sources of impairment cannot be ignored. If ongoing infrastructure repair focuses solely on combined sewer overflow abatement, opportunities to address parallel and potentially more

persistent water quality impairments from leaking sewers are missed. Given the dynamics documented here, more sophisticated prioritization of repair targets and increased recognition of subsurface interactions between infrastructure and streams to improve urban water quality and mitigate water quality impairments downstream of urban systems (Chung et al., 2023, in press).

Data availability statement

The raw data supporting the conclusion of this article will be made available by the authors, without undue reservation.

Author contributions

RF collected all samples, conducted the laboratory analyses, and did the majority of the data analysis and writing. DB and EE assisted in data analysis as well as writing and editing.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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Supplementary material

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fenvs.2023.1117809/full#supplementary-material>

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