Characteristic scales during the onset of radiatively driven convection: linear analysis and simulations

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This paper considers the initial stage of Radiatively Driven Convection (RDC), when the perturbations from a quiescent but time-dependent background state are small. Radiation intensity is assumed to decay exponentially away from the surface, and we consider parameter regimes in which the depth of the water is greater than the decay scale of e of the radiation intensity. Both time-independent and time-periodic radiation are considered. In both cases, the background temperature profile of the water column is time-dependent. A linear analysis of the system is performed based on these time-dependent profiles. We find that the perturbations grow in time according to $\exp[(\sigma(t)t)]$, where $\sigma(t)$ is a time-dependent growth rate. An appropriately defined Reynolds number is the primary dimensionless number characterising the system, determining the wavelength, vertical structure, and growth rate of the perturbations. Simulations using a Boussinesq model (the Stratified Ocean Model with Adaptive Refinement) confirm the linear analysis.

1. Introduction

This paper focusses on radiatively driven convection (RDC), which occurs when heat is applied to a fluid by absorption of radiation penetrating a finite distance from a boundary. To achieve convection, the resulting heating must result in an unstable buoyancy distribution developing in the fluid. In fluids where the buoyancy increases with temperature, the radiation must be applied from below. This occurs, for example, in the interior of stars, where the radiation from the inner core drives convection in the outer layer (Spiegel 1971). Conversely, in fluids where buoyancy decreases with temperature, such as fresh water below the temperature of maximum density, RDC requires that radiation be applied from above. An example of the latter in a geophysical setting occurs in temperate lakes during spring when the water column is below the critical temperature and solar radiation heats the surface layer (see e.g. Bouffard et al. 2019; Cannon et al. 2019; Austin et al. 2022). An important difference between RDC in the interior of stars and RDC in lakes is that in the former, the horizontally (or ensemble) averaged temperature profile can be assumed to be in a statistically steady state, that is, the amount of heat received from the inner core is eventually transferred to the surface and lost to space, whereas in the case of RDC in lakes heat continuously accumulates in the system. Austin et al. (2022) shows that some heat loss occurs during nighttime, but it is small compared to the net daytime heat input. Thus, the averaged temperature never achieves steady state. Additionally, and just as important, in lakes the radiation intensity is itself time-dependent, following a diurnal cycle.

RDC as it applies to temperate lakes has been the subject of several recent observational studies that focus on vertical velocity (Bogdanov et al. 2019; Bouffard et al. 2019; Cannon et al. 2019) and the scale of convection cells (Forrest et al. 2008; Yang et al. 2017; Bogdanov et al. 2019; Austin 2019; Austin et al. 2022).

These studies show that in RDC systems that are driven by a cyclical radiation which spends a significant amount in the "off" state each cycle follows a consistent pattern:

(i)Onset: the beginning of each cycle starts from a relatively quiescent state.

(ii)Linear phase: warming of the water column develops a top-heavy buoyancy distribution on which perturbations grow. In this stage, the effect of perturbations on the averaged buoyancy field is negligible. The latter is still driven solely by the absorbed radiation.

(iii)Non-linear phase: the amplitude of perturbations saturates due to nonlinear interactions.

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(iv)Recovery phase: As the intensity of the radiation wanes, turbulent fluctuations decrease in intensity, and eventually the system relaxes to a mostly quiescent state with little or no residual stratification.

Early studies (Mironov and Terzhevik 2000; Mironov et al. 2001) suggest that if the depth is horizontally uniform, when turbulence develops, the vertical divergence of the total heat flux (the sum of turbulent and radiative heat fluxes) becomes constant with depth, that is, the rate of heating becomes uniform, or, which is the same, the stratification profile becomes frozen in time (even as the fluid heats up). This suggests that the stratification during the nonlinear phase is determined by the length of the linear phase, since the stratification ceases to grow once turbulence sets it. Since advection operates on the averaged vertical temperature gradient, whose temperature contrast at the end of the linear phase is proportional to the time lapsed since the onset of radiation, the latter also gives an estimate for the temperature fluctuations, at least until the waning solar radiation alters the balance and turbulence starts eroding the temperature gradient. From this point of view, the linear phase sets the characteristics of the turbulent phase. This provides the motivation for the present study.

Motivated by observations of RDC in Lake Superior, Christopher et al. (2023) recently studied the onset of RDC by applying a time-periodic heat flux to the surface of a fluid. Applying Floquet theory, they calculated the critical Rayleigh number and normalised wavenumber as a function of normalised frequency of thermal forcing (their figures 2 and 3), and showed that the critical Rayleigh number captures stability properties in two-dimensional numerical simulations (their figures 8 and 9).

Although appropriate to consider the stability of the background RDC state, the analysis in Christopher et al. cannot be used to study the evolution of perturbations at geophysical scales. The maximum normalised frequency considered in Christopher et al. (2023) is 100, while, for example, the normalised frequency in geophysical settings such as Lake Superior is of $O(10^7)$. Moreover, their analysis does not provide the growth rate and vertical structure of the most unstable mode, nor are any characteristics of the system described when the Rayleigh number exceeds the critical value.

Christopher et al.'s analysis applies to cases where the time scale of evolution of the perturbations is comparable to or larger than the period of the forcing, but at more geophysically relevant scales the perturbations grow on a time scale much shorter than the forcing period.

Yet, the onset is still followed by a period in time during which the perturbations are still small, so a linearised treatment is still appropriate. With this in mind, here we develop a theory that applies to systems in which appropriately defined Reynolds and Péclet numbers are large, and the forcing can be time-dependent. This is not intended to provide a stability analysis of RDC in the traditional sense. The aim of the latter is to determine over which range of values of the relevant parameters the system develops instabilities. In our analysis, we consider the linearised regime in the limit of large Reynolds numbers, where we expect the system to be unstable, and we concentrate on two questions:

1. What are the wavelength, vertical structure, and growth rate of the growing perturbations during the initial linear growth stage for RDC driven at geophysical scales?
2. How do these features relate to environmental parameters such as radiation intensity and penetration depth?

We consider the problem from the point of view of an initial-value problem. At $t = 0$ radiation is applied to the surface of a motionless, unstratified fluid with a realistic e-folding decay scale, where the radiation can be either time-independent or can have a diurnal cycle, and we follow the growth of a perturbation from $t = 0$ driven by the time-evolving background state.

We find that perturbations do not follow the typical exponential growth $\exp(\alpha t)$, which may be expected for instabilities that grow on an otherwise constant-in-time background, (α being the constant in-time growth rate). Instead, the perturbations grow as $\exp[(\sigma(t)t)]$ and $\sigma(t) \propto t^{n/2}$ with $n = 1$ for time-independent radiation and $n = 2$ for time-periodic radiation (the latter for times shorter than the period).

This paper is organised as follows. In Section 2 we linearise the equations of motions by considering a time-varying basic state buoyancy profile which is heated by radiative forcing, whose evolution is considered in Section 3; in Section 4 we estimate scalings for velocity and buoyancy by balancing the dominant terms in perturbation equations and introduce the relevant nondimensional parameters; in Section 5 we explore the behaviour of the perturbations under linearised dynamics in the limit of large Reynolds number; in Section 6 we use Direction Numerical Simulations (DNS) to confirm the prediction obtained from the linearised equations; finally we provide a summary and conclusions. Several Appendices discuss technical points.

2. RDC in the linearised regime

We consider a fluid with a linear equation of state for the density $\rho = \rho_0(1 - \alpha T)$, where T is the temperature. We assume a negative thermal expansion coeficient α. The fluid is subject to a radiative forcing applied to the surface. The applied heat flux $S_0 e^{z/Z_0} F(t)$ decays exponentially away from the surface and can be modulated in time (figure 1). The problem under the Boussinesq approximation can be formulated as follows

$$\frac{D\tilde{u}}{Dt} = -\frac{1}{\rho_0}\nabla\tilde{p} + \tilde{b}e_3 + \nu\nabla^2\tilde{u}, \qquad \nabla\cdot\tilde{u} = 0, \tag{2.1}$$

$$\frac{D\tilde{b}}{Dt} = -\frac{B}{Z_0}F(t)e^{\frac{z}{Z_0}} + \kappa\nabla^2\tilde{b}. \tag{2.2}$$

Here $\tilde{u} = (\tilde{u}, \tilde{v}, \tilde{w})$ is the velocity, $\tilde{p}$ is the pressure deviation from the hydrostatic profile, $\tilde{b} = g(\rho_0 - \rho)/\rho_0 = \alpha g T$ is the buoyancy, $B = (-\alpha g S_0/\rho_0 C_p)$ is the buoyancy flux due to the radiative heat flux S_0, ρ_0 is the reference density, C_p is the heat capacity, g is the gravitational acceleration, Z_0 is the e-folding decay scale of the radiation flux, ν is the molecular viscosity, κ is the molecular heat diffusivity, and e_3 is the unit vector pointing upward. The surface through which radiation is applied is at $z = 0$ and the domain extends below to $z = -H$. All thermodynamic quantities are evaluated at the reference temperature. Mutatis mutandis, the same configuration applies to radiation applied to the bottom of a fluid with a positive expansion coeficient, with the understanding that in this case $Z_0 < 0$.

For time-independent radiation,

$$F(t) = 1, \tag{2.3}$$

while for diurnal solar radiation

$$F(t) = \sin(\Omega t) \quad \text{if} \quad t < \tau_{DTL}, \quad \text{else} \quad 0, \tag{2.4}$$

where $\Omega = \pi/\tau_{DTL}$ and τ_{DTL} is the daytime length.

We decompose the motion into a basic state and perturbations

$$\tilde{u} = 0 + u(x,t), \qquad \tilde{b} = \bar{b}(z,t) + b(x,t), \qquad \tilde{p} = \bar{p} + p(x,t). \tag{2.5}$$

Here $x = (x, y, z)$ and $\bar{b}(z, t = 0) = 0$. The basic state satisfies

$$0 = -\frac{1}{\rho_0}\nabla\bar{p} + \bar{b}e_3, \tag{2.6}$$

$$\frac{\partial\bar{b}}{\partial t} = -\frac{B}{Z_0}F(t)e^{\frac{z}{Z_0}} + \kappa\frac{\partial^2\bar{b}}{\partial z^2}. \tag{2.7}$$

From here, we follow the same approach used to study the stability of Rayleigh-Bénard convection (Chandrasekhar 1961). Subtracting the basic state momentum equation (2.6) from (2.1) and then neglecting squares of perturbations, the linearized perturbation momentum equation reads

$$\frac{\partial u}{\partial t} = -\frac{1}{\rho_0}\nabla p + be_3 + \nu\nabla^2 u, \tag{2.8}$$

From (2.8) and the incompressibility condition we derive a single equation for the Laplacian of the vertical velocity

$$\frac{\partial\nabla^2 w}{\partial t} = \nabla_h^2 b + \nu\nabla^2\nabla^2 w. \tag{2.9}$$

Here, $\nabla_h^2 = \partial^2/\partial x^2 + \partial^2/\partial y^2$ is the horizontal Laplacian operator. Subtracting the basic state buoyancy equation (2.7) from (2.2) and then neglecting squares of perturbations, the linearized perturbation buoyancy equation reads

$$\frac{\partial b}{\partial t} + w\frac{\partial\bar{b}(z,t)}{\partial z} = \kappa\nabla^2 b. \tag{2.10}$$

Substituting $\nabla_h^2 b$ in (2.9) into $\nabla_h[(2.10)]$, these two equations can be combined as

$$\frac{\partial^2\nabla^2 w}{\partial t^2} + (\nabla_h^2 w)\frac{\partial\bar{b}}{\partial z} - (\nu + \kappa)\frac{\partial}{\partial t}\nabla^2\nabla^2 w + \nu\kappa\nabla^2\nabla^2\nabla^2 w = 0. \tag{2.11}$$

(2.7), (2.9) and (2.10) will be used to find scalings for the linear system. (2.11) will be used to find the growth rate and the spatial structure of the perturbations.

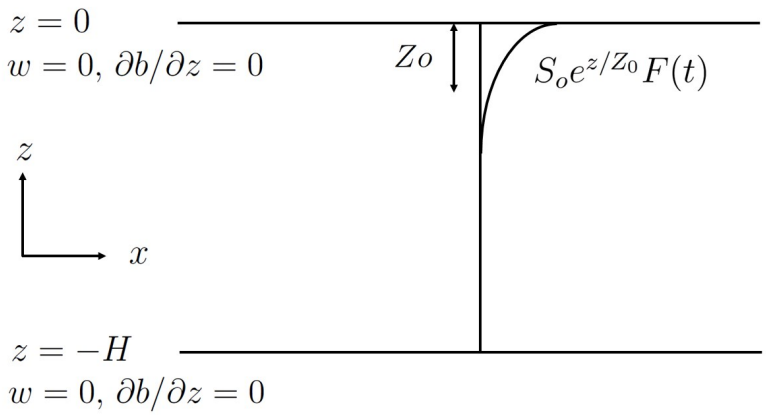

Figure 1: Schematic of RDC.

3. Evolution of the background profile

With an appropriate choice of a vertical velocity scale W_0, a buoyancy scale b_0, and a time scale t_0 (to be defined in the next section), we define dimensionless vertical velocity, buoyancy, time, and coordinates

$$\hat{w} = w/W_0, \quad \hat{b} = b/b_0, \quad \hat{t} = t/t_0, \quad \hat{x} = x/Z_0, \quad \hat{y} = y/Z_0, \quad \hat{z} = z/Z_0. \tag{3.1}$$

We also define a Reynolds number $Re \equiv W_0 Z_0/\nu$ and a Péclet number $Pe \equiv W_0 Z_0/\kappa$. From the buoyancy scale, we can derive a temperature scale $T_0 = b_0/(-\alpha g)$. In the following we will dispense from decorating non-dimensional variables, and all variables except those in section 4 are non-dimensional. The background buoyancy profile satisfies

$$\frac{\partial \bar{b}}{\partial t} = -F(t)e^z + \frac{1}{Pe}\frac{\partial^2 \bar{b}}{\partial z^2}, \tag{3.2}$$

where the Péclet number Pe is the ratio of the perturbation time scale (to be defined more precisely later) to the diffusive time scale Z_0^2/κ. This equation needs to be solved subject to boundary and initial conditions. For the latter, we simply choose $\bar{b}(z, 0) = 0$. At the bottom ($z = -H$) of the water column, the natural choice is a no-flux condition. At the surface, we assume that the latent, sensible and long-wavelength radiative heat flux are small compared to the incoming short-wave heat flux, and thus we approximate the surface boundary condition with a no-flux condition as well. This approximation is suggested by the observations of Austin et al. (2022) who reports that the total increase in the heat content of the water column as a function of time can be, to a great degree of accuracy, predicted by integrating the equation for heat over the water column with no-flux conditions at both boundaries. As we shall see, for large values of Pe, the evolution of the perturbations is primarily controlled by the evolution of the stratification in the bulk of the water column driven by the absorbed radiation. With these boundary conditions, (3.2) can be solved by writing the solution as a standard trigonometric series. The vertical gradient of the general solution is thus

$$\frac{\partial \bar{b}}{\partial z}(z, t) = \sum_{m=1}^{\infty} a_m R(\lambda_m^2/Pe, t) \sin(\lambda_m z), \tag{3.3}$$

where the wave number $\lambda_m = m\pi/H$, and the coeficients

$$a_m = \frac{2\lambda_m}{H}\frac{1 - (-1)^m e^{-H}}{1 + \lambda_m^2}. \tag{3.4}$$

The function $R(s, t)$ describes the relaxation of the solution to the stationary state (which is not steady for diurnal radiation)

$$R(s, t) = \begin{cases} 1 - e^{-\alpha t} & \text{if } F_s(t) = 1, \text{ (Steady radiation)} \\ e^{-\alpha t}) & \text{if } F(t) = \sin(\Omega t), \text{ (Diurnal radiation).} \end{cases} \frac{s^2 + \Omega^2}{(s \sin(\Omega t) - \Omega(\cos(\Omega t) -} \tag{3.5}$$

In the case of steady radiation, $R(s, t) = R(st, 1)t$. For diurnal radiation this is not in general true. However, here we are interested in perturbations that grow on a time scale much shorter than the

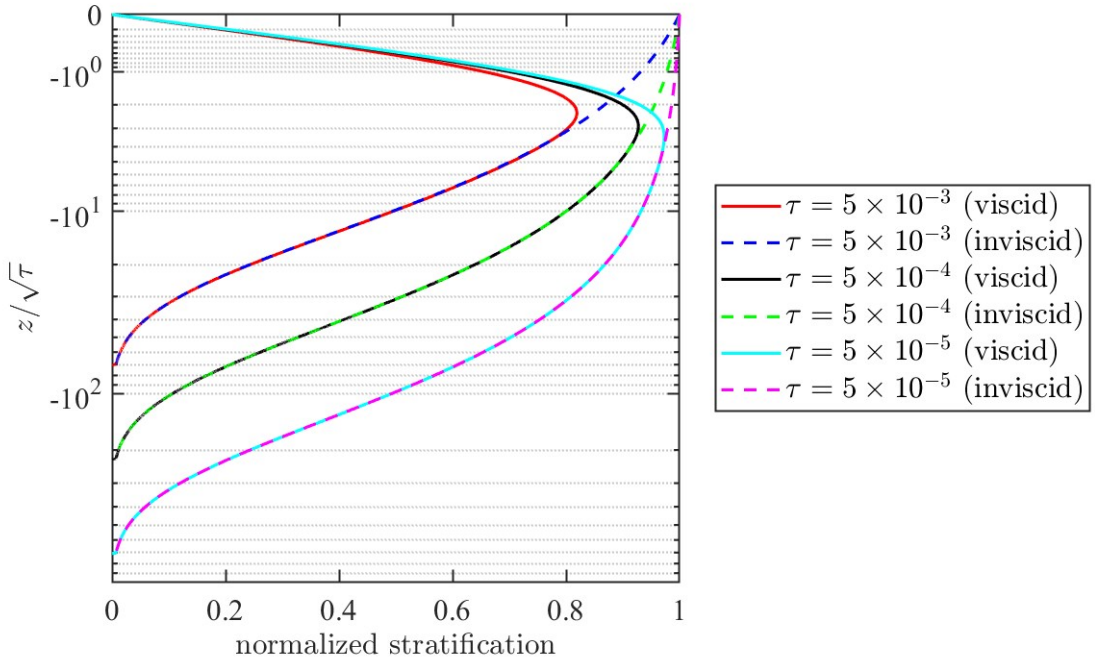

Figure 2: Profiles of $\partial\bar{b}/\partial z$ normalized with fast time (solid lines) and e^z (dashed lines) for different values of τ plotted against $z/\sqrt{\tau}$. The viscous profiles depart from the inviscid solution starting at a depth which deepens as $\sqrt{\tau}$.

diurnal period, i.e. $\Omega t \ll 1$. In this case we can approximate

$$R(s,t) \sim \Omega \frac{st - (1 - e^{-st})}{s^2} + O((\Omega t)^2), \tag{3.6}$$

and therefore $R(s,t) \sim R(st,1)t^2 + O((\Omega t)^2)$. Thus we can write a general form for the background shear

$$\frac{\partial \bar{b}}{\partial z}(z,t) = \left[\sum_{p=1}^{\infty} a_p R(\lambda_p^2 \tau, 1) \sin(\lambda_p z) \right] t^n \equiv S_n(z,\tau) t^n, \tag{3.7}$$

with $\tau = t/Pe$, $n = 1, 2$ for steady and unsteady radiation respectively, and $S_n(z,\tau)$ is the term in square brackets in (3.7) with the appropriate choice for R, the relaxation function. Thus, for large values of the Péclet number, there are two time scales that control the evolution of the background stratification profile: the "fast" time t over which the profile evolves in a self-similar manner, and the "slow" time τ over which the overall shape of the profile changes as the diffusive boundary layer grows at the surface. In particular, the inviscid solution

$$\bar{b}(z,t) = -e^z \int_0^t F(t')dt'. \tag{3.8}$$

is recovered in the limit $\tau \to 0$.

For finite, but small values of τ the inviscid solution approximates well the actual solution except for the surface boundary layer whose thickness grows as $\sqrt{\tau}$ (figure 2). Of course, for this to work, the penetration depth (which in our units is 1) must be much larger than the thickness of the surface boundary layer during the time over which the analysis is carried. In practise, this limits the analysis to times shorter than Z^2/κ. Thus, our analysis cannot be applied to Rayleigh-Benard convection driven by a time-dependent surface heat flux (SHF convection), because that would require taking the $Z_0 \to 0$ limit. Physically, in time-dependent SHF the driving signal is carried into the fluid by the developing boundary layer itself. Whereas in RDC we have a non-trivial inviscid solution which is modified over a slow time by diffusion effects, in SHF convection the inviscid solution is trivial, and the background system evolves under the slow time alone.

4. Scaling and normalisation

In this section, we temporarily revert to dimensional quantities. There are two length scales in RDC: the depth of water H and the radiation penetration scale Z_0. As near-surface water warms gradually and just starts to sink, the depth of water during the initial stages of RDC should not play a role provided $H/Z_0 \gg 1$. Thus Z_0 is the natural length scale during the onset of this process .

The basic state is the time-dependent solution to (2.7) given by (3.7). Clearly, both $\bar{b}$ and $\partial\bar{b}/\partial z$ change continuously over time. Thus, perturbations grow against a background state which is itself changing.

We derive scales by balancing the dominant terms in (2.9) and (2.10). Since growing perturbations are forced by the basic state, which is time dependent in RDC, time-derivative terms must be retained. In the vertical momentum equation (2.9), we assume the local vertical acceleration and buoyancy balance

$$\frac{\partial^2 w}{\partial t^2} \sim b. \tag{4.1}$$

In the buoyancy equation (2.10), we assume a balance between the local rate of increase in buoyancy and vertical advection of buoyancy

$$\frac{\partial b}{\partial t} \sim w \frac{\partial \bar{b}(z,t)}{\partial z}. \tag{4.2}$$

Three equations sufice to solve for the vertical velocity scale W_0, the buoyancy scale b_0, and the time scale t_0.

4.1. Time independent radiation

The inviscid solution for the background buoyancy provides a scale for the buoyancy

$$b_0 = \frac{B}{Z_0} t_0, \tag{4.3}$$

which in conjunction with (4.1) and (4.2) allows us to determine the other scales

$$W_0 = (BZ_0)^{1/3}, \qquad b_0 = \frac{B^{2/3}}{Z_0^{1/3}}, \qquad t_0 = \frac{Z_0^{2/3}}{B^{1/3}}. \tag{4.4}$$

Substituting (3.1) and (4.4) into (2.11) we have

$$\frac{\partial^2 w}{\partial t^2} - \frac{1}{Re}\frac{1}{Pe}\frac{\partial^2 \theta}{\partial t} + \frac{1}{RePe}\nabla^2\nabla^2 w = \nabla_H^2 w S_1(z,\tau)t - \nabla_H^2 w. \tag{4.5}$$

The physical interpretation of the characteristic scales (4.4) is that as the RDC develops, the perturbations grow to $O(W_0)$ and $O(b_0)$ over a time $O(t_0)$, after which the system becomes nonlinear.

Also, note that $W_0 = (BZ_0)^{1/3}$ in (4.4) has the same form as the Deardorff (1970) scaling $W^* = (B^* H)^{1/3}$, which characterises the vertical velocity in SHF convection (B^* is the buoyancy flux applied to the surface and H is the depth of the water). However, the fact that the two scalings are of the same form should be viewed as a coincidence, because the two problems differ fundamentally. The Deardorff scaling characterises vertical velocity in SHF convection in the nonlinear steady state stage, while (4.4) are scalings for the linear, time-dependent stage of RDC in which the radiation profile penetrates into fluid with an e-folding decay scale Z_0.

4.2. Diurnal solar radiation

In this case, the inviscid solution indicates that

$$b_0 = \frac{B}{Z_0}\Omega t_0^2. \tag{4.6}$$

Substituting W_0, b_0, and t_0 into (4.1), (4.2), (4.6), we have

$$\frac{W_0}{t_0} = b_0, \qquad \frac{1}{t_0} = \frac{W_0}{Z_0}, \qquad b_0 = \frac{B}{Z_0}\Omega t^2, \tag{4.7}$$

which together yield

$$W_0 = (B\Omega Z_0^2)^{1/4}, \qquad b_0 = (B\Omega)^{1/2}, \qquad t_0 = \frac{Z_0^{2/4}}{B\Omega}, \tag{4.8}$$

which are to be interpreted as their counterpart in the steady radiation case. Substituting (3.1) and (4.8)

into (2.11) and working under the assumption that $\Omega t_0 \gg 1$ so that the background solution retains the self similar profile in the fast time t, we have

$$\frac{\partial}{\partial t_2}\nabla^2 w - \nabla_H^2 w S_2(z,\tau)t^2 - \frac{1}{Re}\nabla^4 w + \frac{1}{Pe}\frac{\partial}{\partial t}\nabla^2 w + \frac{1}{RePe}\nabla^6 w = 0. \tag{4.9}$$

5. Evolution of perturbations

The equations for the evolution of the perturbation under steady and periodic radiation conditions can be condensed in a single equation which in non dimensional form reads

$$\left(\frac{\partial}{\partial t} + \frac{1}{Pe}\frac{\partial}{\partial \tau} - \frac{1}{Re}\nabla^2\right)\left(\frac{\partial}{\partial t} + \frac{1}{Pe}\frac{\partial}{\partial \tau} - \frac{1}{Pe}\nabla^2\right)\nabla^2 w = S_n(z,\tau)t^n \nabla_H^2 w, \tag{5.1}$$

with $n = 1$ describing the steady radiation case and $n = 2$ the diurnal cycle. This equation describes the evolution of the perturbed vertical velocity $w(x, y, z, t, \tau)$.

In the geophysical settings of interest, the Reynolds and Péclet numbers are large, though not infinite, and therefore it is of consequence to consider whether (5.1) can be further simplified. To this extent, it would be tempting to discard altogether the terms proportional to Re^{-1} and Pe^{-1}. However, this would not be appropriate, since, on physical grounds, we expect that viscosity and diffusivity at suficiently small scales cannot be ignored. However, it is reasonable to expect that no high-frequency oscillations should be expected in τ (otherwise it would be a fast time), and thus we can neglect the term $Pe^{-1}\partial/\partial\tau$. Equation (5.1) is invariant under rotations around the vertical axis. Thus, we consider perturbations confined to a two-dimensional vertical plane. Since equation (5.1) is not homogeneous in fast time, the typical solution form $e^{\gamma t}$ with γ being a constant growth rate cannot be applied. Therefore, we seek solutions in the form of

$$w(x,t) = e^{iKx}\psi(z,t,\tau), \tag{5.2}$$

and we expand $\psi(z,t,\tau) = \sum_m f_m^K(t)\varphi_m^K(z,\tau)$. The $\varphi_m^K(z,\tau)$ functions are the eigenvectors of the Sturm-Liouville problem

$$\frac{d^2\varphi_m}{dz^2} - K^2\varphi_m = -\frac{1}{D_m}K^2 n! S_n(z,\tau)\varphi_m, \qquad \varphi_m(0,\tau) = \varphi_m(-H/z_0,\tau) = 0, \tag{5.3}$$

where $D_m(K)$ is the corresponding eigenvalue, and normalized such that

$$\int_{-H/z_0}^{0} K^2 n! S_n(z,\tau)\varphi_m^K(z,\tau)\varphi_l^K(z,\tau)dz = \delta_{ml} \tag{5.4}$$

with $n = 1$ is for steady state radiation and $n = 2$ for diurnal radiation. As we shall see, the eigenvalues D_m's control the growth of the perturbations in the inviscid limit.

A second simplification that we seek is to replace $n!S_n(z,\tau)$ with its limiting value $n!S_n(z,0) = e^z$ in (5.1) and the attendant Sturm-Liouville problem. This assumption is justified by considering that, by the time the linear stage of perturbation growth comes to an end, τ is still very small. Indeed, one way to interpret the Péclet number is to see it as the ratio of t_0 (the time scale of growth of perturbations) to the diffusive time scale Z^2/κ. Thus at large Péclet numbers the perturbations ought to experience a background state whose only mode of change is self-similar. In Appendix A we verify that the effect of the neglected surface boundary layer on the spectral properties of the Sturm-Liouville problem is small, especially on the eigenvalues D_m's, which control how perturbations grow in time.

For a given K, there are countable eigenvalues $D_1(K) > D_2(K) > \ldots$, and the corresponding eigenvectors form an orthonormal basis. $D_1(K)$ as a function of the wavelength $\lambda = 2\pi/K$ is shown in Figure 3a, with a few representative eigenvectors shown in figure 3b. Note how, as K increases, the region over which the eigenvector is non-negligible decreases. In fact, with a simple rescaling of the problem, it can be shown that the extent of the non-zero region decreases as $1/\sqrt{K}$. Although it is not possible to give an analytic expression for the D_m's, we have $D_m(K) = O(K^2)$ for $K \ll 1$, whereas $D_m(K) = O(K^0)$ for $K \gg 1$. In particular, the fact that the eigenvalues saturate at large wave numbers will have significant consequences. For a given horizontal wavenumber K, substituting (5.2) into (5.1) yields a system of coupled ODEs for the amplitude functions $f_m^K(t)$. In Appendix B we show that when the Reynolds number is large (with fixed Prandtl number), to leading order the system decouples, so that each amplitude function $f_m^K(t)$ evolves independently from the other. For a given $f_m^K(t)$, the inviscid growth rate is controlled by the eigenvalue $D_m(K)$. Because we are interested in the characteristics of the fastest growing perturbations, we focus on the case $m = 1$. From now on, $D(K) = D_1(K)$ and

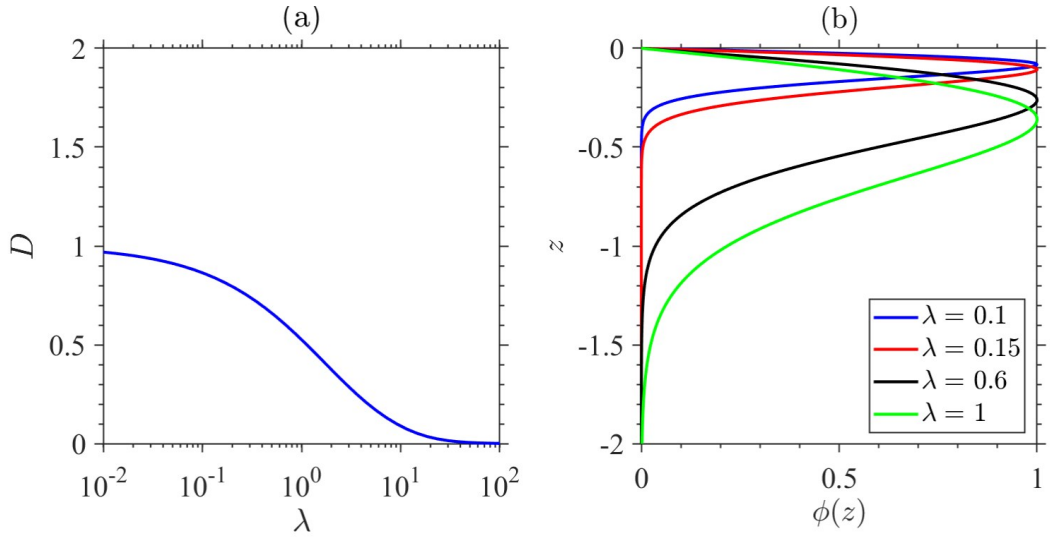

Figure 3: Solution of the Sturm-Liouville problem (5.3). (a) eigenvalue D as a function of λ. (b) eigenvector $\varphi(z)$ for the maximal D for a given λ. Solutions are obtained by solving (5.3) using a domain $-5 \leq z \leq 0$, boundary conditions $\varphi(-5) = \varphi(0) = 0$, and 1000 uniform grids. The eigenvalue D controls the growth rate in (5.8)-(5.9), (5.14), and (5.17), while the eigenvector $\varphi(z)$ represents the vertical structure of perturbations.

$f^K(t) = f_1^K(t)$. By introducing the renormalized wavenumber $K = K/\sqrt{DRe}$ to order $O(Re^{-1/2})$ the amplitude $f^K(t)$ satisfies the following equation

$$\frac{d^2 f^K}{dt^2} + K^2\left(1 + Pr\frac{1}{dt}\right)\frac{df^K}{dt} Pr^{-1}K^4 - D(K)\frac{t^n}{n!} f^K = 0, \tag{5.5}$$

where $Pr = Pe/Re$ is the Prandtl number of the fluid.

In the previous equation, we can discern two limits. In the limit $K \to \infty$ with Re constant, (5.5) tends to a simple differential equation with constant coeficients, whose characteristic polynomial has roots

$$\sigma_{1,2}^{\infty}(K) = -K^2\{1, Pr^{-1}\} \tag{5.6}$$

both of which are real and negative and thus exponentially damped. This is to be expected on physical grounds, as for suficiently small wavelengths diffusive effects will smooth out and eventually dampen fluctuations. Conversely, when $K \to 0$, with K constant, the equation reduces to the modified Airy's equation

$$\frac{d^2 f^0}{dt^2} - D(K)\frac{t^n}{n!} f^0 = 0, \tag{5.7}$$

whose solution grows asymptotically as

$$f^0(t) \sim \frac{1}{D(K)^{1/12} t^{1/4}} e^{(\sigma_t t)^{3/2}}, \qquad \sigma_t = \frac{2}{3} D(K)^{1/3}, \qquad \text{as} \quad t \gg 1, \tag{5.8}$$

when $n = 1$ and as

$$f^0(t) \sim \frac{1}{D(K)^{3/8} t^{1/2}} e^{(\sigma_d t)^2}, \qquad \sigma_d = \frac{D(K)^{1/4}}{2^{3/4}}, \qquad \text{as} \quad t \gg 1, \tag{5.9}$$

when $n = 2$. In practice, these asymptotic formulas apply already for $t = O(1)$, as seen by comparing asymptotic solutions with the numerical integration of the corresponding equations (Figure 4). Note that σ_d and σ_t tend to saturate as $K \to \infty$, whereas the viscous damping timescale becomes increasingly shorter. On physical grounds, we can thus expect that at a certain wavenumber, dependent on the Reynolds number, a cross-over occurs whereby viscous damping dominates and so very little energy should be found above such wavenumber. To verify our intuition, we consider the general solution of (5.5). It can be written in terms of special functions.

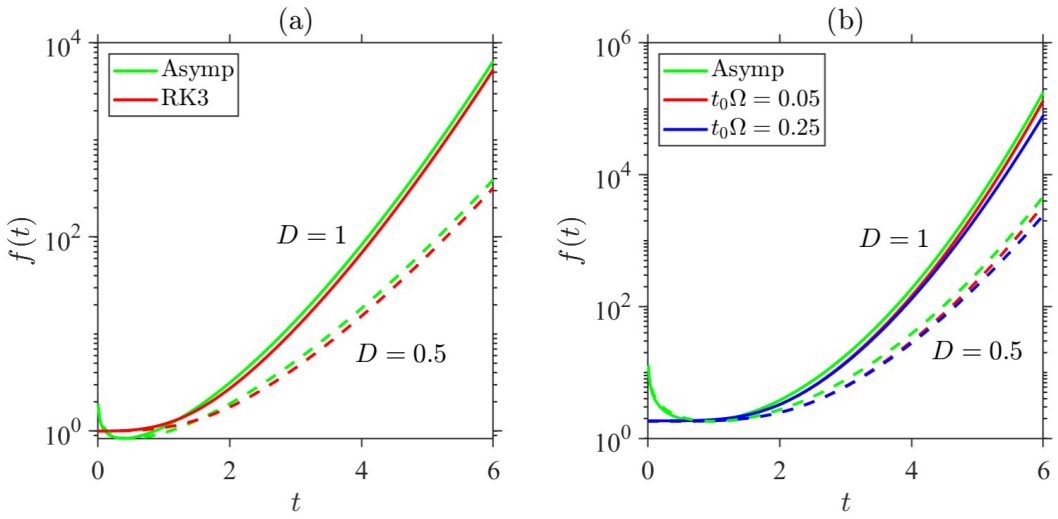

Figure 4: Comparison between asymptotic solution and numerical solution using a third-order Runge-Kutta method with dt = 0.01, and initial conditions f(0) = 1 and f′(0) = 0. The comparison is to show asymptotic solution captures the growth of numerical solution as t ⪆ 1. Solid curves are for D = 1; dashed for D = 0.5. (a) Asymptotic solution (5.8) and numerical solution of the modified Airy equation (5.7) with n = 1. (b) Asymptotic solution (5.9) and numerical solution of (5.7) with n = 2.

5.1. Time independent radiation

We consider first the case when n = 1. In this case we have

$$f^K(t) = c_1(K)A(K,t) + c_2(K)B(K,t). \tag{5.10}$$

Here $c_1(K)$ and $c_2(K)$ are integration constants and the functions A and B can be expressed in terms of Airy's Ai and Bi functions

$$A(K,t) = e^{\overline{\sigma}(K)t}\mathrm{Ai}(\zeta(K,t)), \quad B(K,t) = e^{\overline{\sigma}(K)t}\mathrm{Bi}(\zeta(K,t)), \quad \zeta(K,t) = \frac{(1-Pr^{-1})^2K^4 + 4D(K)t}{4D(K)^{2/3}}, \tag{5.11}$$

and $\overline{\sigma}(K)$ is the arithmetic mean of the $\sigma_i^\infty(K)$'s. The integration coeficients are given by

$$c_1(K) = \frac{(\overline{\sigma}(K)-1)\mathrm{Bi}(\zeta(K,0))f'(0) + \mathrm{Bi}'(\zeta(K,0))f(0)}{\Delta(K)}, \tag{5.12}$$

and

$$c_2(K) = -\frac{(\sigma(K)-1)\mathrm{Ai}(\zeta(K,0))f'(0) - \mathrm{Ai}'(\zeta(K,0))f(0)}{\Delta(K)}, \tag{5.13}$$

with $\Delta(K) = \mathrm{Ai}(\zeta(K,0)\mathrm{Bi}'(\zeta(K,0)) - \mathrm{Ai}'(\zeta(K,0)\mathrm{Bi}(\zeta(K,0))$.

For t > 1 the solution can be asymptotically expressed as

$$\begin{aligned} f^K(t) &\sim e^{\Sigma_t t}\left[K^4\left(\frac{Pr-1}{Pr}\right)^2 + 4Dt\right]^{-1/4}, \\ \Sigma_t &= -\frac{K^2}{2}\frac{1+Pr}{Pr} + \frac{1}{12Dt}\left[K^4\left(\frac{Pr-1}{Pr}\right)^2 + 4Dt\right]^{3/2} - \frac{1}{12Dt}\left[K^4\left(\frac{Pr-1}{Pr}\right)^2\right]^{3/2}. \end{aligned} \tag{5.14}$$

The solution when molecular viscosity and diffusivity are equal (Pr = 1) is particularly illuminating, since in this case the solution is simply the inviscid solution multiplied by the viscous damping term.

$$f^K(t) \sim \frac{1}{D(K)^{1/12}t^{1/4}}e^{-K^2t+(\sigma_t t)^{3/2}}, \qquad \sigma_t = \frac{2}{3}\sqrt{2/3}\,D(K)^{1/3}. \tag{5.15}$$

Neglecting the (weak) algebraic dependence on time of the prefactor, we see that evolution of a perturbation with a horizontal wavenumber K is controlled by the time- and wave-number-dependent

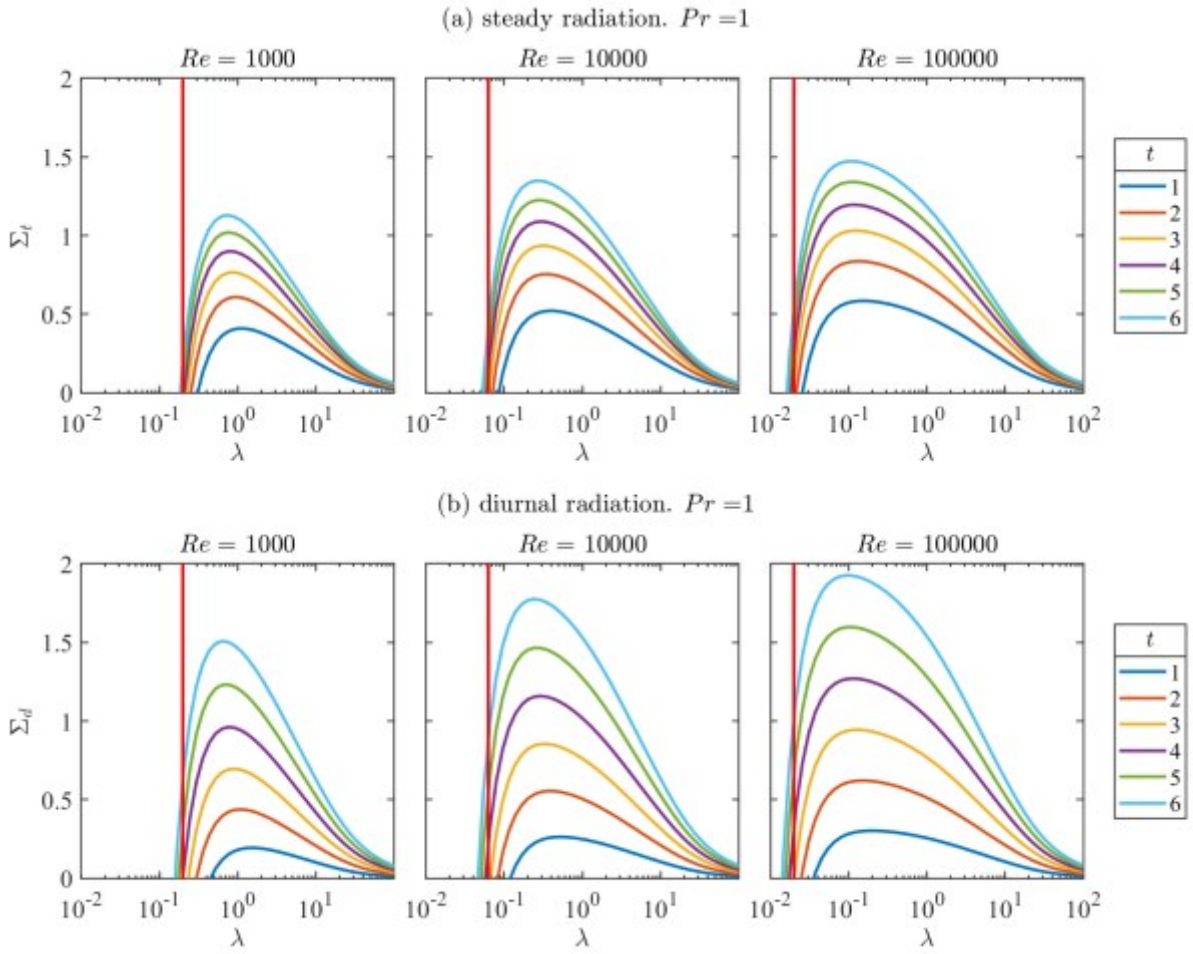

Figure 5: Asymptotic growth rate Σ vs wavelength at different times. The vertical red lines indicated the estimated high-frequency cutoff $K = 1$, i.e. $\lambda \sqrt{Re} \approx 2\pi$. (a) Steady radiation; (b) Diurnal radiation.

growth rate

$$\Sigma_t = \frac{1}{Re} - D(K) + \frac{K^2}{3} D(K) t Re^2. \tag{5.16}$$

Figure 5a shows the growth rate for a range of times as a function of the wavelength. Note that Σ_t is at most O(1), which vindicates the choice of t_0 as the relevant time scale for the process.

For a given Re, there is a Reynolds number dependent t_{min} such that for $t < t_{min}$ the time-dependent growth rate $\Sigma_t(K, t) < 0$. This can be easily seen considering that as a function of K the viscous damping is a convex function which tends to a value O(1/Re) for $K \to 0$, whereas the inviscid growth is a concave function which approaches zero as K t for small values of K (figure 6)

To find how t_{min} depends on Re, we consider the following ansatz $t_{min} = \beta Re^{-2}$. When substituted into (5.16), the Reynolds number is factored out, leaving an equation for β that can be easily solved numerically. We obtain $\beta \approx 491$ and a Reynolds-independent marginal wavelength $\lambda_{mar} \approx 10.7$. Past t_{min} the range of wavelengths that experiences growth widens. The upper limit of the range increases as Re $t - t_{min}$. The lower limit of the range decreases as $Re^{-1/2}(t - t_{min})^{-1/4}$. The peak of the growth rate rapidly shifts to smaller wavelengths (figure 5a). The existence of a minimum time that needs to elapse before instabilities can grow implies that instabilities will appear only after the background stratification has grown suficiently. Since the non-dimensional time-dependent background stratification is $N^2 = -te^z$, the minimum near-surface background stratification necessary to sustain perturbations is $-t_{min}$.

However, we recall that the superexponential growth given by (5.16) is not expected until $t \approx 1$. Therefore, for large values of the Reynolds number, by the time the solution enters the superexponential phase, a range of wavelengths is already poised to grow. Moreover, for large values of Re, the growth rate for most wavelengths experiencing growth is only weakly dependent on the Reynolds number. Only

wavelengths that are close to either side of the interval see a significant departure from purely inviscid growth. Thus, when comparing our results with numerical simulations, we use the inviscid limit.

The asymptotic expression is more complicated when $Pr = 1$, but as long as Pr remains finite, there is no qualitative change in the behavior of the solution and little quantitative change. In fact, as we shall see, the predicted growth rate and peak wavelength at $Pr = 1$ match the values observed in the numerical simulation at $Pr = 10$.

The above analysis suggests that instabilities will always appear, provided that enough time is allowed for the stratification to reach its critical value, regardless of the magnitude of the Reynolds number. However, as we remarked above, equation (5.5) was derived in the large Re limit. Superexponential growth requires t_{min} to be $\gg 1$ or larger. For values of Re greater than 20, by the time the superexponential phase begins, there is already a range of wavelengths with positive growth rates. Whether values of Re smaller than 20 can still be considered "large" for the purpose of the theory remains to be seen. It is certainly reasonable to assume that no matter how small the Reynolds number is, a suficiently large accumulation of negative buoyancy at the surface should be able to overcome the stabilising effect of viscosity and diffusion, but this is a question that our theory cannot answer with certainty. Moreover, our theory works in the small τ limit. Under highly diffusive conditions, the growth of the surface boundary layer cannot be neglected.

5.2. Diurnal solar radiation

For the diurnal radiation case ($n = 2$), the general solution to (5.5) can be expressed in terms of a combination of Hermite polynomials and hypergeometric functions, which play the role of Airy's functions for the $n = 1$ case. In this case, it is the hypergeometric function that dominates the behaviour for large values of t. For large t the asymptotic solution is

$$f^K(t) \propto c_2(K) e^{\Sigma_d t} \left(\sqrt{2Dt} + \left(\frac{K^4 \left(\frac{Pr-1}{Pr}\right)^2}{(32D)^{1/2}} \right)^{-1/4} \right), \qquad \Sigma_d = -\frac{K^2}{2}\left(\frac{1+Pr}{Pr}\right) + \left(\frac{K^4}{4}\left(\frac{Pr-1}{Pr}\right)^2 + 2Dt^2\right)^{1/2}, \tag{5.17}$$

where in this case $c_2(K)$ is the coeficient of integration associated to the Hypergeometric function. The growth rate Σ_d is dependent on time and wave number, which, when $Pr = 1$ can be written as (Figure 5 b)

$$\Sigma_d = -\frac{K^2}{D(K)Re} + \frac{\sqrt{Dt}}{2\sqrt{2}}. \tag{5.18}$$

Relative to the steady radiation case, the growth is more gradual at the onset, and steeper at later times, but the pattern is otherwise very similar.

In this case as well the background stratification must grow beyond a Re-dependent critical threshold for instabilities to grow. The analysis is very similar to the one done for the steady radiation case, with the exception that now $t_{min} \propto 42 Re^{-1}$ has the same value for the marginal stability wavelength. The upper limit of the unstable range increases as $Re\, t$, whereas the lower range decreases as $(Re\, t)^{-1/2}$ (figure 5b).

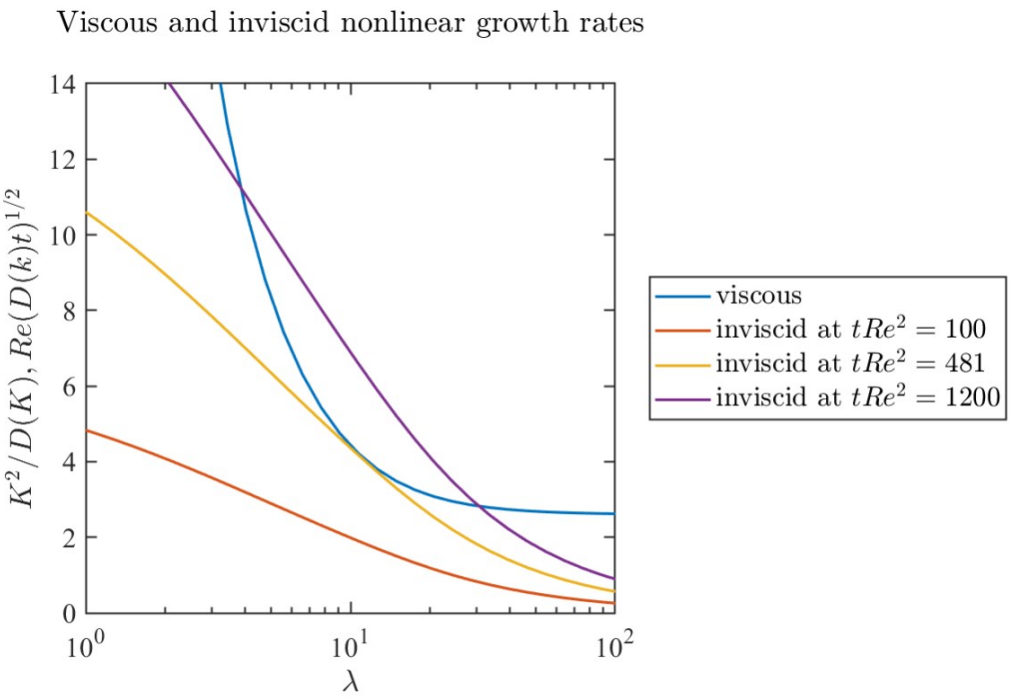

Figure 6: Viscous damping and inviscid growth rate as a function of wavelength under steady radiation conditions. The viscous damping is time independent, whereas the inviscid growth rate accelerates with time. Up to $t \approx 481 Re^{-2}$ viscous damping dominates. Past this time, a widening range of wavenumbers experiences net growth.

6. Comparison with numerical simulations

To validate the theory developed in the preceding section, we use the Stratified Ocean Model with Adaptive Refinement (SOMAR) to compare the prediction of our theory with the DNS simulations of RDC. SOMAR solves the Navier-Stokes equations under the Boussinesq approximation (Santilli and Scotti 2015; Chalamalla et al. 2017) using an operator splitting technique. The finite-volume discretization is second-order in space, while a third-order Runge-Kutta method is adopted for time marching. The Poisson equation is solved with an eficient multigrid solver. SOMAR adjusts the time step based on a CFL condition to maintain stability under advection. Viscous terms are treated implicitly.

We solve equations (2.1) and (2.2) in a rectangular domain extending from the surface to the bottom. Equations are discretised with uniform grids in each direction. The initial condition is a motionless, unstratified fluid plus perturbations of two types. The first type is initialised with single-mode perturbations with the most unstable wavelength predicted by (5.14) and (5.17) and the corresponding vertical structure shown in Figure 3b, with the peak normalised by W_0 being 0.001. The second type is initialised with random temperature perturbations (normalised by T_0) uniformly distributed in the range -0.01 to 0.01 over the entire domain. The boundary conditions are periodic in the horizontal direction. At the surface and at the bottom of the domain we use free-slip conditions and zero buoyancy flux. The use of free-slip conditions at the bottom removes the need to resolve the viscous sublayer. As convective activity is mostly confined with Z_0, the effect is negligible at large Reynolds numbers. The range of random temperature perturbations (-0.01 to 0.01) is chosen so that nonlinear terms is at least two orders of magnitudes smaller than linear terms. The effect of the magnitude of random perturbations is examined in Appendix D.

The configuration of all cases simulated and compared with the theory is described in Table 1. Cases TLR and THR consider time-independent radiation. The other two cases, DLR and DHR consider periodic radiation conditions. The non-dimensional parameters in the time periodic cases correspond to typical values found in Lake Onego (DLR) and Lake Michigan (DHR).

Together, these 4 cases cover a wide range of Re from 200 to 65000 with $Pr = 10$, which is typical for freshwater below the critical temperature. From our theory, we expect that over such a range of Reynolds numbers, the critical wavelength and vertical structure should vary appreciably. To capture these characteristics in each case, the choice of grid spacing is made with the following criteria in mind:

Case	Section	Re	Pe	$t_0\Omega$	δ/S	D	λ_p	S	N_h	N_v
TLR	5.1.1	932	9323	–	0.009	0.64	0.6	1.2	166	122
THR	5.1.2	65444	654444	–	0.003	0.89	0.1	0.4	164	164
DLR	5.2.1	196	1959	0.09	0.013	0.51	1	1.7	161	173
DHR	5.2.2	20908	209080	0.26	0.004	0.83	0.15	0.5	163	205

Table 1: Configurations for the numerical simulations considered in sections 5.1 and 5.2. For each simulation, we list the Reynolds number $Re = W_0Z_0/\nu$, the Péclet number $Pe = W_0Z_0/\kappa$, the ratio δ/S of the thickness of the viscous surface layer at $\tau = 1/Pe$ to the thickness of the most unstable mode S, time scale ratio $t_0\Omega$, the eigenvalue of the most unstable mode D, the wavelength of the most unstable mode λ_p, the vertical extent (measured from the surface) of the most unstable mode S, the number of grid points N_h that resolve one horizontal wavelength , and the number of grid points N_v that resolve S in vertical. All lengths are measured in units of Z_0. In all the cases considered, the non-dimensional depth H of the domain is 5.

First, the vertical domain is 5 times Z_0, and horizontal domain is sized to contain 100 wavelengths with the highest growth rate λ_p. Second, at least 161 horizontal grid points resolve the wavelength with the highest growth rate λ_p and at least 40 grid points resolve the cutoff wavelength λ_{cut}. In the vertical we use at least 122 grid points to resolve the sharp vertical variation S near the surface (see λ_p, S, N_h and N_v in table 1). Third, and finally, in addition to the CFL condition necessary to ensure the numerical stability of the scheme, the time step is further subject to the constraint that it should not exceed $t_0/40$. These three rules ensure that the numerical setup does not bias the simulations. Sensitivity to grid resolution tests are reported in Appendix C.

To distinguish between the linear stage (where our theory is expected to hold) and the non-linear stage we compare $\partial b/\partial t$ and $w(\partial b/\partial z)$ in the buoyancy equation and $\partial w/\partial t$ and $w(\partial w/\partial z)$ in the z-momentum equation. We compute the rms of each term for comparison. The average is taken over the vertical range in which the eigenfunction of the most unstable wavelength varies (S in table 1 and figure 3b). The theoretical solution is normalised so that at $t = 2$ it coincides with the numerical solution.

6.1. Time independent radiation

Case TLR initialized with single-mode perturbations is shown in figure 7a. The leftmost column displays the time evolution of the horizontally averaged buoyancy. The red curve shows the value assumed by the theory, while the blue dashed line is calculated from the SOMAR output. The two profiles are still identical at $t = 6$, while significant difference exists at $t = 7$, signalling the end of the linear stage. Even towards the end of the linear stage, the energy is still concentrated at the wavelength of the initial condition (last column). Buoyancy and vertical velocity are shown in the central columns. The perturbations evolve as a series of downward localized jets.

The growth of the perturbations in the simulation agrees well with the one predicted by the theory until the magnitude of the nonlinear terms catches up with the magnitude of the linear terms around $t = 6.7$ (figure 7b).

We estimate the time dependent growth rate from the simulation as

$$\Sigma_t \approx \frac{2}{3}\frac{d}{dt}(\ln \tilde{w}), \tag{6.1}$$

which compares extremely well with the theoretically predicted value (figure 7c) from $t = 3$ to $t = 6$, after which the former declines because the system approaches the nonlinear stage. The discrepancy at earlier times may simply indicates that (6.1) is sensitive to small change in w, since $\partial w/\partial t$ and $df^K(t)/dt$ agree well (figure 7b). Note that over the interval $t = 2$ and $t = 6$ the magnitude of the perturbation grows by 3 orders of magnitudes.

Between $0 < t < 1$, $\partial w/\partial t$ drops slightly (figure 7b). Presumably, during this time, viscosity in (2.9) and diffusivity in (2.10) play a role in the evolution of perturbations, because $\partial \bar{b}/\partial z$ in (2.10), which grows

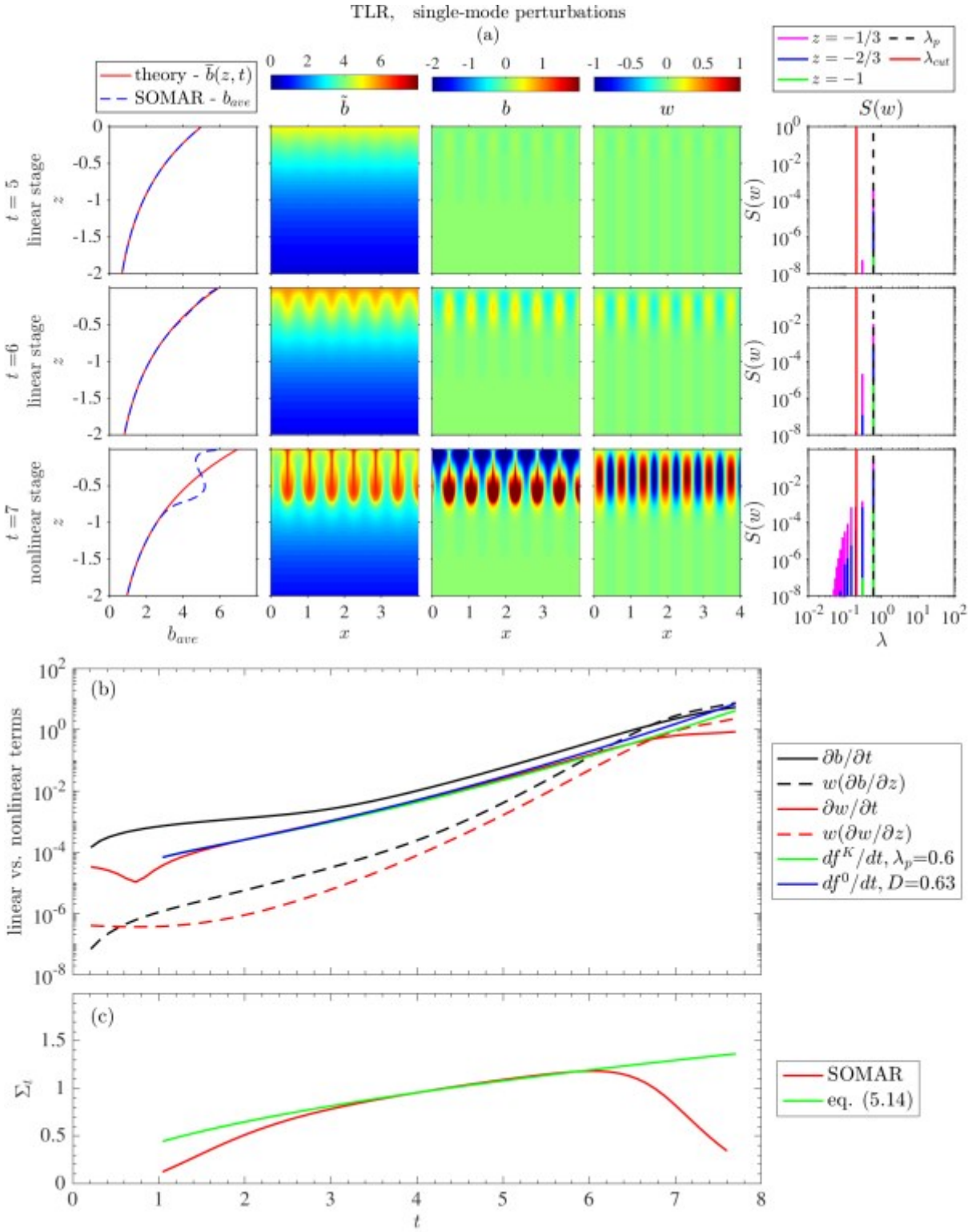

Figure 7: Evolution of RDC under time-independent radiation profile, with Re = 932 and single-mode perturbations. (a) From rows 1-3, time advances. Column 1, comparison between theoretical basic state buoyancy $\bar{b}(z,t)$ and horizontally averaged buoyancy profile b_{ave}. Column 2, side view of total buoyancy b. Column 3, side view of perturbation buoyancy b. Column 4, side view of vertical velocity, which is also perturbation vertical velocity w. Column 5, vertical velocity spectrum at $z = -1, -2/3$, and $-1/3$. λ is the wavelength of the perturbations. λ_p represents the theoretical most unstable wavelength. λ_{cut} the theoretical cutoff wavelength. Note the appearance of a spectral line at half the forcing wavelength due to the quadratic term in the equation of motion at $t = 5$. Only when the flow becomes fully nonlinear we observe energy at wavelengths smaller than λ_{cut}. (b) rms of linear and nonlinear terms in the legend and $df^K(t)/dt$ and $df^0(t)/dt$ as a function of time. λ_p is calculated from (5.14), and D is computed via (5.3). (c) growth rate as a function of time. The Red curve is computed from the SOMAR output. The green curve is the theoretical value (5.14).

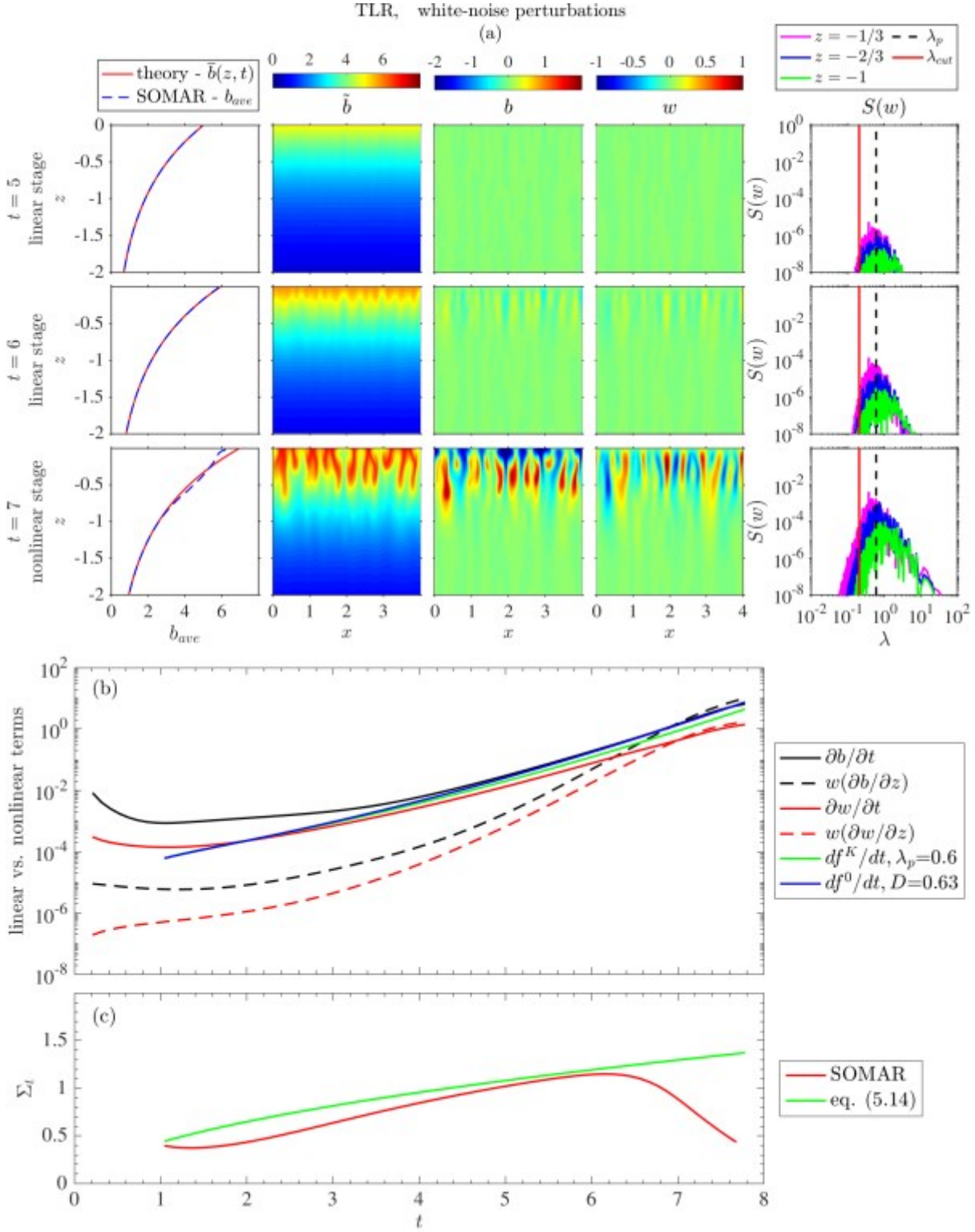

Figure 8: Same as figure 7 but with white-noise perturbations.

from 0, is not strong enough. This does not contradict the balance (4.1) and (4.2) and the subsequent analysis in sections 4 and 5. (4.1) and (4.2) target the evolution of perturbations when $\partial b/\partial z$ grows continuously, and the analysis which follows applies for $t > 1$.

Simulations of case TLR with white-noise perturbations are shown in figure 8. The major difference in this case is that the spectrum contains energy over the range of wavelengths that experience positive growth, peaked on the wavelength of maximal growth ($\lambda_p = 0.6$ at $Re = 932$) predicted by the theory, which also predicts well the vertical envelope of the perturbations, with a peak at $z = -0.25$ (figure 3b).

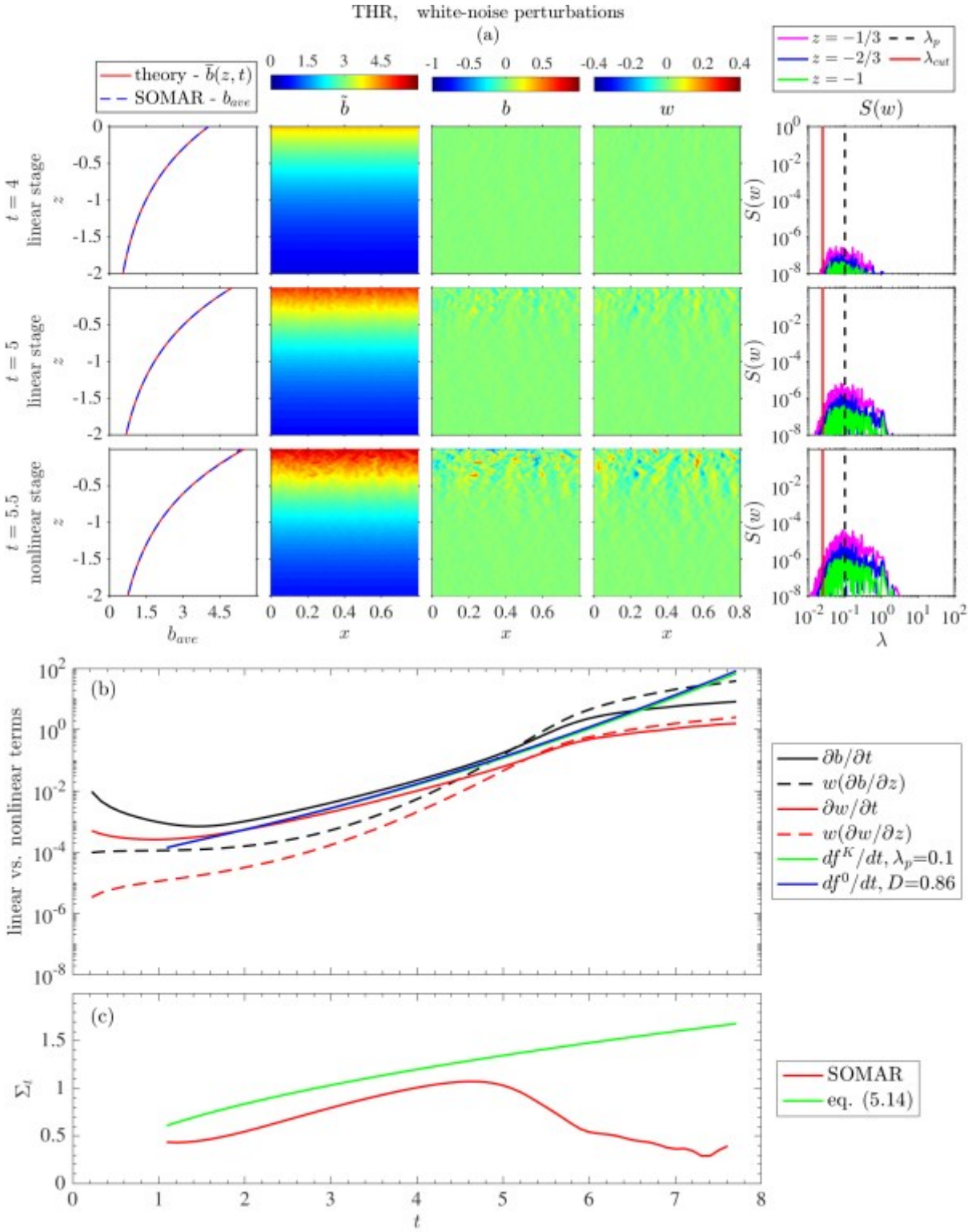

Figure 9: Same as figure 7, but with Re = 65444, white-noise perturbations and smaller colorbar range for columns 1-4.

Towards the end of the linear stage, there is a shift of the spectrum to larger wavelengths, likely due to the nonlinear merger of plumes. Of course, in the nonlinear stage three dimensional effects become dominant, which are not captured by our simulations. The spectra at different depths essentially overlap, indicating strong coherence across the vertical dimension, consistent with a single mode being energized.

Overall, the theory predicts well how the perturbation grows in time, though the measured time-dependent growth rate Σ_t is somewhat smaller than the one calculated from the theory, though the

latter is the growth rate expected from the most unstable mode, whereas in the simulation we have a combination of modes across a range of horizontal wavenumbers that, while following the same growth pattern, have different values of the eigenvalue $D(K)$. Thus it is to be expected that the measured Σ_t be lower than the theoretical value based on the most unstable wavelength.

Case THR considers a much larger value of the Reynolds number ($Re = 65444$). In this case as well the averaged numerical buoyancy profile agrees well with the profile used for the theoretical analysis up to $t = 5.5$, after which the numerical solution becomes dominated by nonlinearity (figure 9a, column 1). The spectra at different depths do not exactly overlap, as was the case at lower values of the Reynolds number, though still peaking near the expected wavelength of maximal growth ($\lambda_p = 0.1$). At large values of the Reynolds number, the growth rate of higher vertical modal orders is less sensitive to the modal number, and therefore we expect that the flow initialized with random perturbations will exhibit a mixture of contribution from different modes. Overall, the theory captures well the growth of the perturbations (figure 3b), though the presence of a continuum of energized wavelengths is reflected in the overall time-dependent rate of growth, which, while following the expected increase in time between $t = 2$ and $t = 4.5$ is lower than the growth rate based on the most unstable wavelength and mode (figure 9c).

6.2. Diurnal solar radiation

Simulations forced by diurnal solar radiation with single-mode perturbations agree well with theory, as was for case TLR in the preceding section. Therefore, in this section we focus on simulations with white-noise perturbations.

Case DLR uses parameters typical of springtime conditions in a shallow temperate lake (Lake Onego, Bouffard et al. (2019)). Compared to the other cases considered, diffusion effects are stronger, which results in a relatively low value for the Reynolds number ($Re = 196$). The agreement between the buoyancy profile computed from the simulation and the one used in the theory is good up to $t = 6$, after which nonlinearity dominates (first column in figure 10a; also figure 10b indicates that the nonlinearity begins at $t = 6.6$). The spectra at $t = 5$ peak near $\lambda \approx 1.2$, close to the theoretical prediction $\lambda_p = 1$. Beginning at $t = 6$ we observe a shift of the energy to longer wavelength. This may be due to the flow becoming more nonlinear. At $\lambda_p = 1$ the corresponding mode-1 structure function (figure 3b), extends to $z = -1.7$, with a peak at $z = -0.4$, which captures well the vertical envelope of the fluctuations. The growth of the rms fluctuations is well described by the theory (figure 10b). Over $2 \lesssim t \lesssim 6.6$, the amplitude of the fluctuations increases by four orders of magnitude. Setting $\tilde{w} = \exp(\sigma t^2)$ we compute the growth rate $\Sigma_d = \sigma t$ as

$$\Sigma_d \approx \frac{1}{2}\frac{d}{dt}(\ln \tilde{w}). \tag{6.2}$$

The computed growth rate is in good agreement up to $t = 6.2$ (figure 10c), after which the computed growth rate declines as the linear and nonlinear terms become comparable (figure 10b). As it was for the steady radiation cases, the calculated growth rate when the flow is initialized with a spectrally broad-banded initial condition is somewhat lower than the theoretically predicted value. However, the trend in time is very similar, and the same considerations that we presented in the steady radiation case apply here as well. Overall, the agreement between theory and simulations is good.

Case DHR is representative of a deeper lake, close to critical temperature on a cloudy day. Such conditions are often observed in Lake Michigan and Lake Superior (Cannon et al. 2019; Austin 2019). With the role of viscosity and diffusivity diminished, the superexponential growth rate is larger, and therefore the duration of the linear phase is shorter. Prior to $t = 5$, the theoretical basic state buoyancy (4.6) and SOMAR output coincide, after which nonlinearity becomes apparent ($t = 5.5$) (figure 11a, column 1). The spectra peak around the theoretical wavelength of maximal growth ($\lambda_p = 0.15$), which is almost an order of magnitude shorter than in the previous case. The vertical envelope of the fluctuations is confined in the upper 10% of the water column during the linear phase, as expected from the mode-1 function associated to λ_p. The computed time-dependent growth rate follows the theoretical profile well, though it is smaller than the value expected for the the wavelength of maximal growth. Again, this is to be expected, since the range of active wavenumbers is wide.

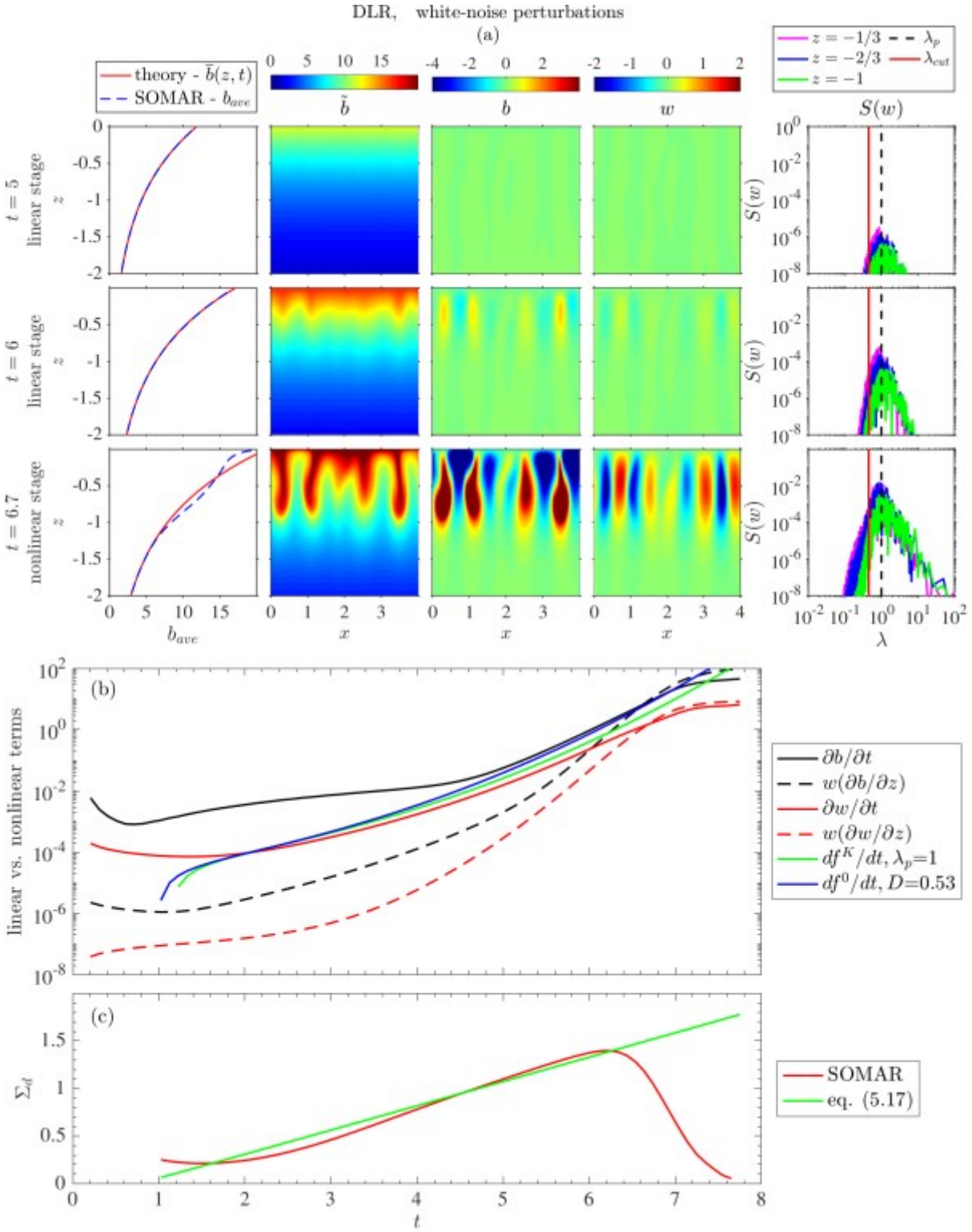

Figure 10: RDC under time-dependent radiation at Re = 196 initialized with white-noise. Symbols and colors as in figure 7. Note that the colorbar range for columns 1-4 is greater.

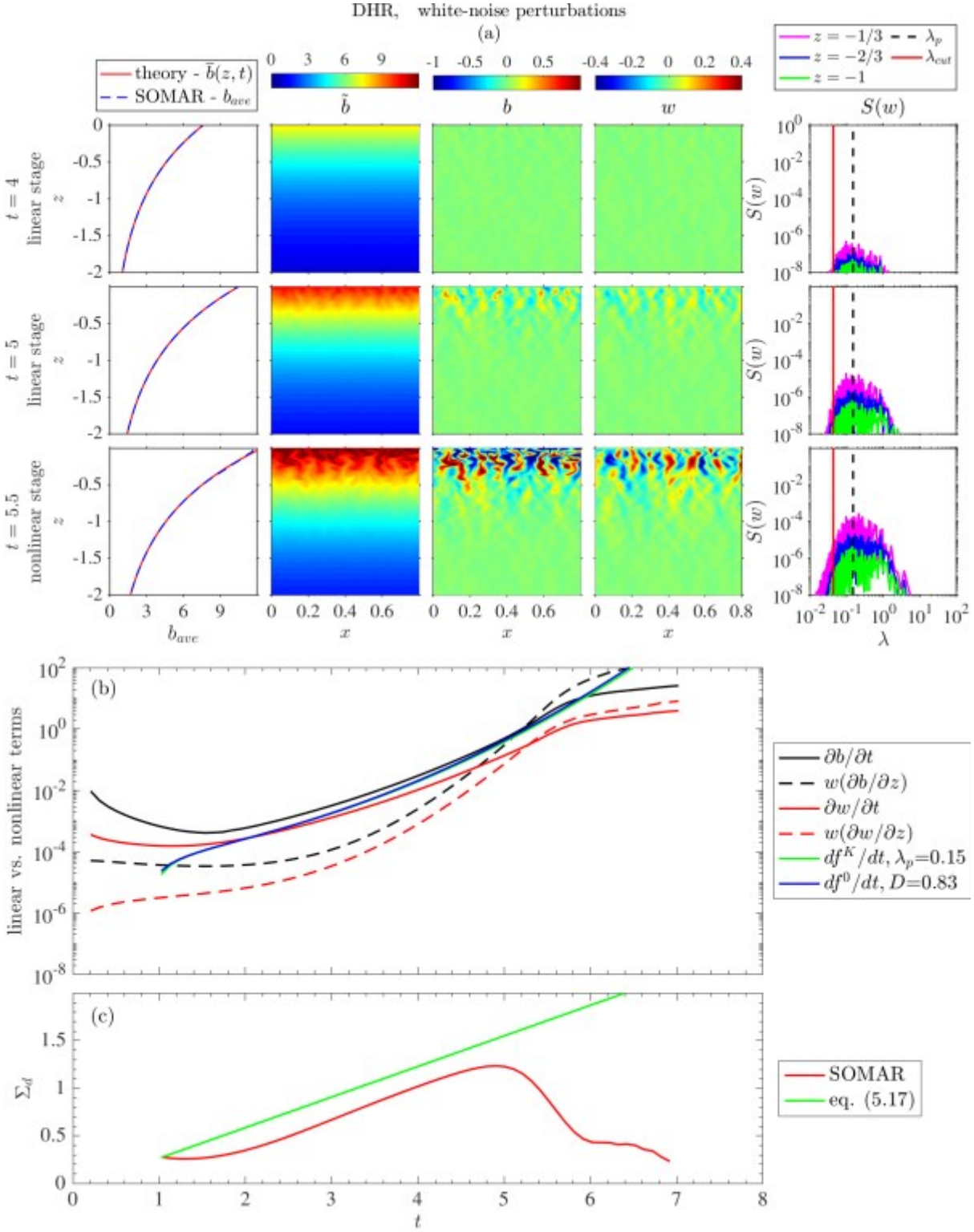

Figure 11: R D C under time-dependent radiation profile at R e = 20908 with white-noise initial condition. Symbols and colors as in figure 7. Note that the colorbar range for columns 1-4 is smaller.

7. Summary and conclusions

We have considered the characteristics of the perturbations at the onset of Radiatively Driven Convection. In this stage, the perturbations are small enough that it is permissible to linearise the equations of motions around a quiescent background state characterised by a time-varying buoyancy profile. We considered both steady radiation, which has been considered in recent laboratory experiments (Bouillaut et al. 2019) and has applications to stellar interior (Kippenhahn et al. 1990), as well as time dependent cases, more representative of Radiatively Driven Convection in temperate lakes during spring time and which may have application to atmospheric convection (see, e.g. Figure 4 in Deardorff (1974)). As opposed to the more traditional Rayleigh-Bénard convection, where the background state moves heat at fixed rate from bottom to top boundary, while remaining constant in time, in Radiatively Driven Convection the background state evolves in time, as heat is continuously added by radiation, and only after the onset of convection heat can be effectively redistributed in the interior. By considering the most likely balance in the dynamical equations, we build velocity and time scales which together with viscosity allow the definition of a Reynolds number. We develop a theory that is valid for large values of the Reynolds number that predicts the wavelength, vertical structure, and growth rate of the perturbations during the initial linear stage. Due to the time-dependent nature of the background state, the growth of perturbations not significantly affected by viscous damping has the form $\exp[(\sigma t)^n]$ where the exponent is $n = 3/2$ for time-independent radiation and $n = 2$ for diurnal radiation. We have confirmed the analysis by comparison with highly resolved DNS.

The theory gives a way to estimate the duration of the linear phase. Indeed, the latter terminates when the amplitude of the nonlinear terms becomes comparable to the amplitude of the linear terms. In Appendix D we show that the duration of the linear phase over a range of intensities of the residual flow at dawn is found to be between $3t_0$ and $6t_0$, where the linear time scale t_0 is given by the last of (4.8). For Lake Superior, assuming a water temperature of $T = 3.5\ ^oC$, and a e-folding scale $Z_0 = 10$, we obtain $t_0 \approx 80$ min on a cloudy day when the radiative intensity is $S_o \approx 200\ W/m^2$. On a sunny day, with the radiative intensity $S_o \approx 800\ W/m^2$, the linear time scale is reduced to $t_0 \approx 60$ min.

If we accept that once turbulence sets in the vertical gradient in temperature stabilises, the latter is then set by the length of the linear phase. Ceteris paribus, a weaker initial circulation left over from the previous cycle will result in a longer linear phase and thus a stronger stratification at the onset of the turbulent phase, that is, a stronger temperature contrast upon which advection acts. This should result in larger rms temperature fluctuations (relative to the background) during the day. With lakes as our geophysical example of RDC under periodic radiation forcing, one way to test this prediction from an observational point of view would be to correlate the strength of the fluctuations in temperature to the strength of the residual circulation at first daylight.

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Appendix A

In this appendix we consider the Sturm-Liouville problem 5.3 subject to a more general profile for the background stratification which includes the upper boundary layer. We compare three profiles: The inviscid profile considered earlier, and two profiles in which the buoyancy gradient is (a) equal to zero within the region $-\sqrt{\tau} \le z \le 0$; or (b) decreases linearly to zero within the same region. We discretise the Sturm-Liouville problem with standard second-order differences. For the latter two cases, we resolve the boundary layer with at least 20 points. Figure 12 compares the largest eigenvalue of the problem as a function of wavelength: when $\tau = 5 \times 10^{-3}$, the eigenvalues computed with diffusive boundary layers of type (a) or (b) are virtually indistinguishable. Compared to the inviscid profile, saturation at small wavelengths occurs more slowly, but for the same wavelength the difference is never greater than 10%. The difference is much smaller for $\tau = 5 \times 10^{-4}$. The corresponding eigenfunctions show little difference between the two boundary layer cases. Relative to the eigenfunctions obtained from the inviscid profile, we observe a slight shift downward of the peak in the eigenfunction when the boundary layer is included. However, the difference is small. Since the focus of this paper is on the small τ regime, this justifies the use of the inviscid profile ($\tau = 0$) in the calculations presented in the main paper.

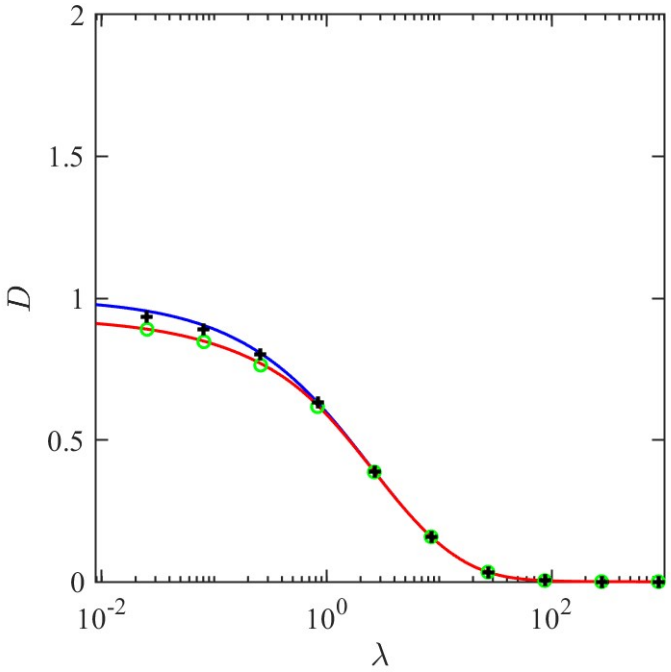

Figure 12: D as a function of λ for three different buoyancy profiles: inviscid profile (blue); profile with zero stratification in the boundary layer at $\tau = 5 \times 10^{-3}$ (red) and $\tau = 5 \times 10^{-4}$ (crosses); profile with stratification linearly approaching zero with the boundary layer at $\tau = 5 \times 10^{-3}$(circles).

Appendix B

In this appendix we present a more formal derivation of eq. (5.5). Let us consider a given K and let $\mathcal{D}^2 \equiv (-K^2 + d^2/dz^2)$ and $\mathcal{D}^2_H \equiv -K^2$. We substitute (5.2) into (5.1) and project over the eigenfunctions obtained solving the Sturm-Liouville problem (5.3) to obtain a set of coupled ODEs for the functions $f_m(t)$. In particular, we need to consider integrals over the vertical domain $\gamma = [-H/Z_0, 0]$ that can be written as

$$I^p_{mn} = \int_\gamma \mathcal{D}^{2p}\varphi_m(z)\ \varphi_n(z)dz, \tag{7.1}$$

with $p = 1, 2, 3$. When $p = 1$, the integral reduces to

$$I^1_{mn} = -\frac{\delta_{mn}}{D_m}, \tag{7.2}$$

where δ_{mn} is the Kronecker's delta. When $p = 2$, we can use integration by part and the properties of the eigenfunctions to obtain

$$I^2_{mn} = \int_\gamma \mathcal{D}^2\varphi_m(z)\mathcal{D}^2\varphi_n(z)dz = \frac{K^2}{D_m D_n}\delta_{mn} + \epsilon^1_{mn}, \tag{7.3}$$

with $\epsilon^p_{mn} \equiv \int_\gamma (e^{pz} - 1)\varphi_m(z)\varphi_n(z)d\mu$, where $d\mu = K^2 e^z dz$ is the weighted measure of the interval. Finally, again using integration by part

$$I^3_{mn} = \frac{K^4}{D_m D_n} - \frac{\delta_{mn} + \epsilon^2_{mn}}{D_m} + \frac{\delta_{mn} + \epsilon^1_{mn}}{K^2} + \frac{\int_\gamma e^z \frac{d\varphi_m}{dz}\varphi_n d\mu}{K^2}. \tag{7.4}$$

The functions $f_m(t)$ then must obey the following system of coupled ODEs:

$$\frac{d^2 f_m}{dt^2} + \sum_l \left[\frac{D_m}{Re}\left(1 + \frac{1}{Pr}\right) I^2_{ml}\frac{df_l}{dt} - \frac{D_m}{Re^2 Pr} I^3_{ml} f_l\right] = -D_m \sum_n \frac{t^n}{n!} f_m. \tag{7.5}$$

Introducing the rescaled wavenumber $K = K/\sqrt{ReD_1}$, we have

$$\frac{D_m}{Re} l_{ml}^2 = K^2 \frac{D_1}{D_l}(\delta_{ml} + \epsilon_{ml}^1), \tag{7.6}$$

and

$$\frac{D_m}{Re^2} l_{ml}^3 = -K^4 \frac{D_1^2}{D_m D_l}(\delta_{ml} + \epsilon_{ml}^2) + O(Re^{-1}). \tag{7.7}$$

Finally, we note that for finite values of K (i.e., small wavelengths) where viscous effects are going to be important, the eigenfunctions are non-zero in a region whose size is $O(Re^{-1/2})$ (see Figure 3(a)). Thus, the coupling terms $\epsilon_{ml}^i = O(Re^{-1/2})$ are negligible. Therefore, the (7.5) to $O(Re^{-1/2})$ reduces to a set of decoupled equations for the coeficients f_m's, all of the form (5.5) with $D = D_m$ and $f = f_m$. Since the growth rate in the inviscid limit is proportional to the eigenvalue, we focus on the gravest mode ($m = 1$) which, by definition, has the largest eigenvalue.

Appendix C Sensitivity to grid resolution

In this Appendix we show grid-independent tests. We choose THR and DLR for illustration because they cover the maximal and minimum Reynolds numbers Re in this paper. As we can see from figures 13 and 14, as the grids are refined by a factor of 2 in both x and z directions, the growth rates are identical in the linear stage (before the growth rates drop). The only difference appears to be a slightly longer duration of the linear phase.

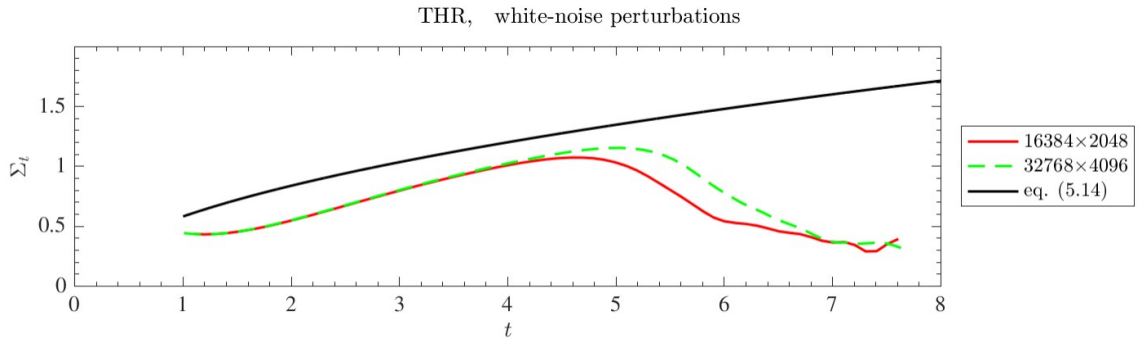

Figure 13: Growth rates as a function of time under steady radiation profile with Re = 65444. The resolution is indicated in the legend.

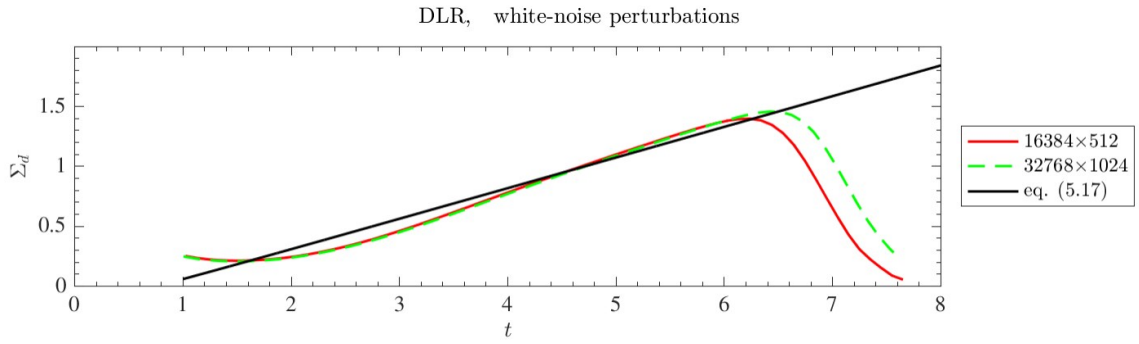

Figure 14: Growth rates under diurnal radiation profile with Re = 196. Symbols as in figure 13.

Appendix D Effects of initial perturbation magnitude

This appendix examines the effect of the magnitude of the initial perturbations. Section 6 considers random temperature perturbations normalized by T_0 uniformly distributed in the range from -10^{-2} to 10^{-2}. In this appendix, we compare different initial perturbation intensities $(-10^{-1}, 10^{-1})$, $(-10^{-2}, 10^{-2})$, and $(-10^{-3}, 10^{-3})$. We choose cases THR and DLR to examine the perturbation effect because they cover the largest and smallest Reynolds number Re considered in our numerical experiments. In both cases, for the largest initial perturbations, the linear terms are less than an order of magnitude greater than the nonlinear terms (upper panels in figures 15 and 16) and the system never experiences a linear stage. In contrast, when the initial perturbations are smaller (last two cases), the growth rate is similar and for the case with the smallest initial condition, the duration of the superexponential stage is indeed longer. This results in the stratification at the onset of the nonlinear turbulent stage being stronger.

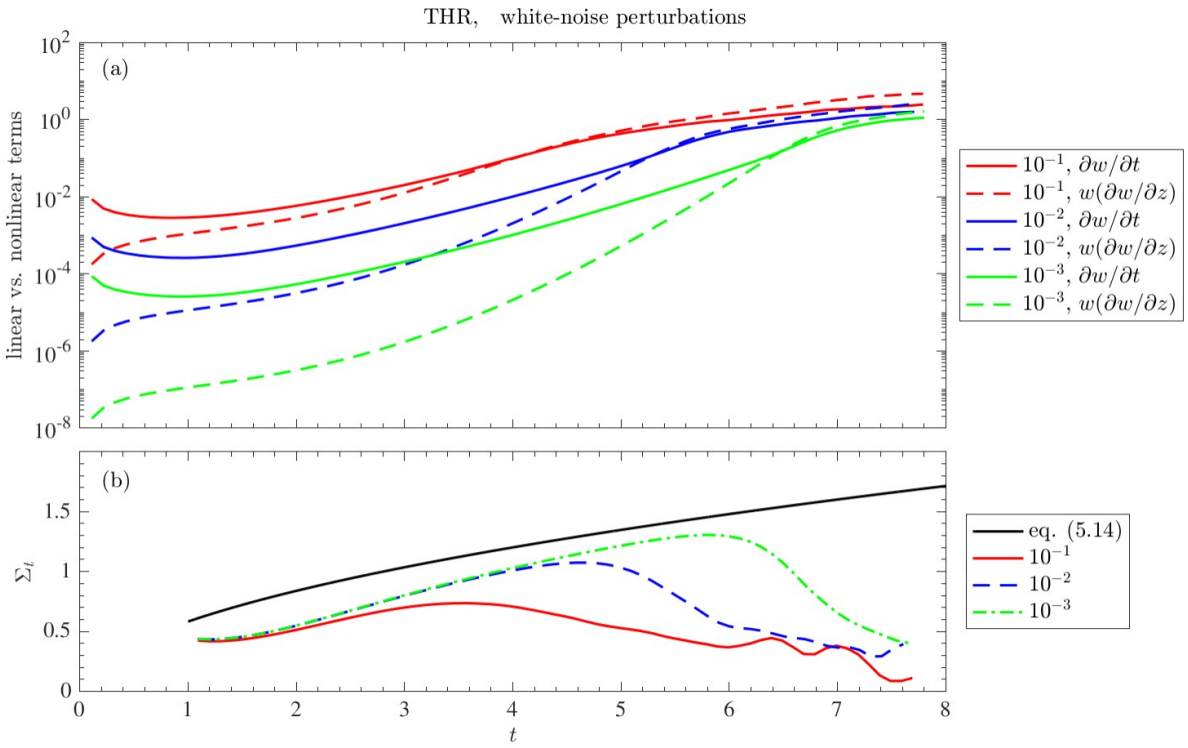

Figure 15: Evolution of RDC under time-independent radiation profile, with Re = 65444. (a) rms of linear and nonlinear terms. (b) growth rates. The numbers in the legend indicate the strength of the initial perturbations. Black curve is the theoretical value (5.14).

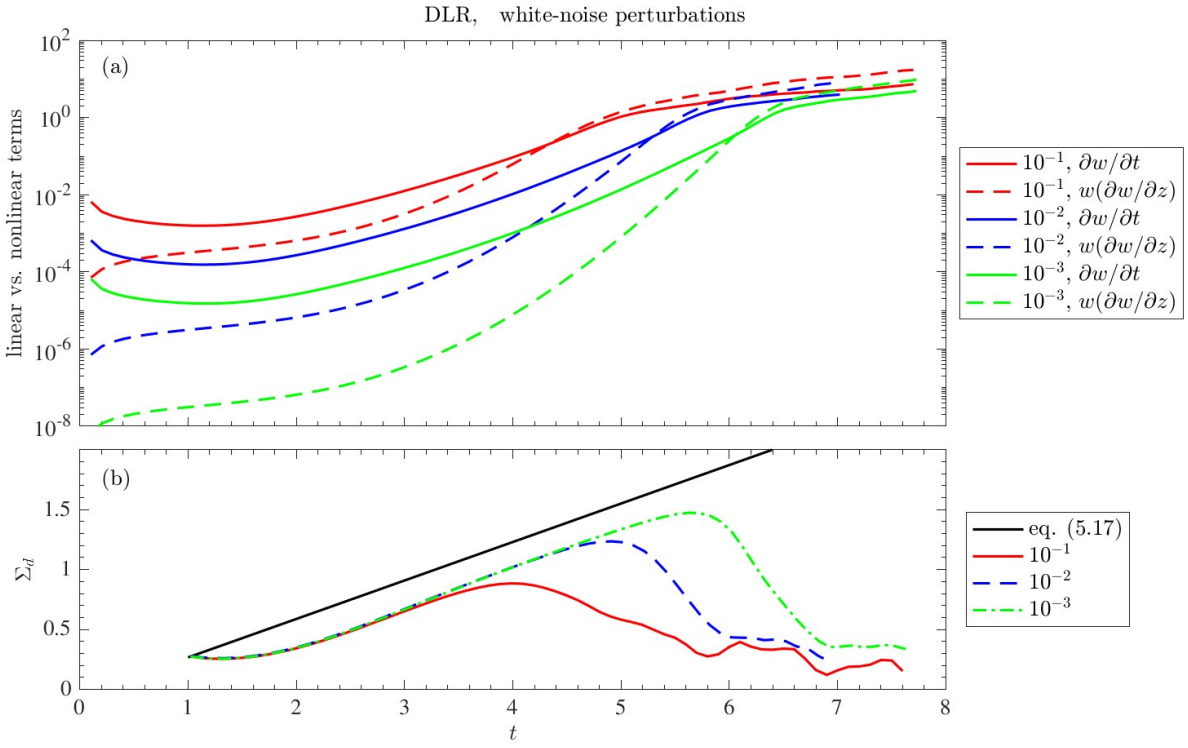

Figure 16: Same as figure 15, but under diurnal radiation profile and Re = 196.