

Quantifying the environmental effects on tropical cyclone intensity change using a simple dynamically based dynamical system model

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Abstract

Accurate prediction of tropical cyclone (TC) intensity is quite challenging due to multiple competing processes among the TC internal dynamics and the environment. Most previous studies have evaluated the environmental effects on TC intensity change from both internal dynamics and external influence. This study quantifies the environmental effects on TC intensity change using a simple dynamically based dynamical system (DBDS) model recently developed. In this simple model, the environmental effects are uniquely represented by a ventilation parameter B , which can be expressed as multiplicative of individual ventilation parameters of the corresponding environmental effects. Their individual ventilation parameters imply their relative importance to the bulk environmental ventilation effect and thus to the TC intensity change. Six environmental factors known to affect TC intensity change are evaluated in the DBDS model using machine learning approaches with the best-track data for TCs over the North Atlantic, central, eastern and western North Pacific and the statistical hurricane intensity prediction scheme (SHIPS) dataset during 1982–2021. Results show that the deep-layer vertical wind shear (VWS) is the dominant ventilation factor to reduce the intrinsic TC intensification rate or to drive the TC weakening, with its ventilation parameter ranging between 0.5–0.8 when environmental VWS between 200 and 850 hPa is larger than 8 m s^{-1} . Other environmental factors are generally secondary, with their respective ventilation parameters over 0.8. An interesting result is the strong dependence of the environmental effects on the stage of TC development.

41 **1. Introduction**

42 Understanding and accurately predicting tropical cyclones (TC) intensity change have long
43 been challenging to both scientific research and operational forecasting (Wang and Wu 2004;
44 Kaplan et al. 2010; Courtney et al., 2019; Hendricks et al. 2019; Tan et al. 2022). The TC
45 intensity change is controlled by complex and nonlinear thermodynamic and dynamic processes
46 interacting at and across multiple scales (Elsberry et al. 2013, Lin et al. 2021), which can be
47 classified as processes intrinsic to a TC vortex and of the TC environmental (extrinsic) effects
48 (Hendricks et.al. 2018). The effects of extrinsic and intrinsic processes on the intensity change
49 of a TC can be complementary, amplifying, inhibiting, or offsetting (Judt and Chen, 2016).
50 Previous studies have identified various environmental factors/processes that affect TC intensity
51 change, such as the large-scale vertical wind shear (VWS), mid-level dry air intrusion, mid-
52 latitude upper-level trough, the negative ocean feedback due to upwelling and vertical mixing
53 in the upper ocean induced by the TC itself, sea surface temperature (SST) gradient, and so on
54 (e.g., Gray 1968; DeMaria and Kaplan 1999; DeMaria et al. 2005; DeMaria 2009; Zeng et al.
55 2008, 2010; Tang and Emanuel 2010, 2012; Wang et al. 2015; Hendricks et.al. 2018; Fei et al.
56 2020; Li et al 2022).

57 In most previous studies, multiple linear regression has been used to identify the key
58 environmental factors by relating the selected environmental variables and the observed TC
59 intensity changes based on the TC best-track data (DeMaria et al. 2005). One of the problems
60 in those statistical studies is that the intensity changes estimated include contributions not only
61 by the environmental influences but also by the TC internal dynamics, while their respective
62 contributions are often hard to be effectively separated and quantified. This is why the
63 correlations between the environmental factors and the TC intensity changes are often small,
64 and the environmental factors can only explain a small portion of the observed TC intensity
65 changes based on the linear statistical analyses (e.g., Zeng et al. 2010; Hendricks et al. 2018).
66 Another issue is the nonlinear interactions between the internal dynamics and external
67 influences (Wang and Wu 2004; Elsberry et al. 2013), which could not be adequately considered

68 by using the linear statistical methods. One such an example is the dependence of the
69 environmental VWS effect on the stage of the TC development (e.g., Zeng et al. 2010). As a
70 result, the potential different responses of TC intensity to environmental influences at different
71 stages of TC development or lifetime could not be uniquely distinguished and evaluated based
72 on the classical statistical methods.

73 Recently, both a simple energetically based and a dynamically based dynamical system
74 models have been developed to quantify the intensification rate (IR) of a TC by Wang et al.
75 (2021a, 2021b, 2022). The energetically based dynamical system (EBDS) model was formulated
76 by viewing a TC as a Carnot heat engine, as proposed by Wang (2012, 2015) and first
77 constructed by Ozawa and Shimokawa (2015). Wang et al. (2021a) introduced an intensity-
78 dependent dynamical efficiency (E), instead of a constant percentage used by Ozawa and
79 Shimokawa (2015), to quantify the conversion of the production rate of potential energy to the
80 production rate of inner-core kinetic energy. The dynamical efficiency E depends mainly on the
81 degree of convective organization in the eyewall and the inner-core inertial stability of the TC
82 vortex as inferred from the balanced vortex dynamics (e.g., Schubert and Hack 1982). Therefore,
83 in their first version of the EBDS model, Wang et al. (2021a) parameterized E as a function of
84 the TC inner-core inertial stability. This makes the model capable of quantitatively capturing the
85 intensity-dependence of TC IR in idealized full-physics model simulations and in observations
86 (Wang et al. 2021a; Xu et al. 2016; Xu and Wang 2018).

87 The dynamically based dynamical system (DBDS) model was developed by Wang et al.
88 (2021b) based on the slab boundary-layer entropy and tangential wind budget equations and the
89 assumption of a thermodynamic quasi-equilibrium under the TC eyewall. A major advancement
90 of the DBDS model of Wang et al. (2021b) compared with the earlier time-dependent theory of
91 TC intensification developed by Emanuel (2012) is the relaxation of the moist neutral eyewall
92 ascent by introducing an *ad hoc* parameter measuring the degree of neutrality of eyewall ascent,
93 which depends on the TC relative intensity, namely, the current TC intensity normalized by its
94 maximum potential intensity (MPI, Emanuel 1986). The new model was also shown to be

95 capable of realistically capturing the intensity-dependence of TC IR in both idealized full-
96 physics model simulations and observations (Wang et al. 2021b). Interestingly, the EBDS and
97 DBDS models share the same mathematical formula for TC IR. The only difference is in that
98 the dynamical efficiency E in the EBDS model is replaced by the *ad hoc* parameter (A)
99 measuring the degree of the moist neutrality of eyewall ascent in the DBDS model. The two
100 parameters even share the same mathematical expression, as a function of the relative TC
101 intensity (Wang et al. 2021b).

102 Theoretically, without any prohibiting environmental effects, both the EBDS and DBDS
103 models give the theoretical upper bound, or potential IR (PIR), that a TC can reach under given
104 favorable oceanic and atmospheric environmental thermodynamic conditions and the current
105 TC intensity (Wang et al. 2021a, b). This was recently demonstrated by Xu and Wang (2022),
106 who showed that the EBDS model (and also the DBDS model) could skillfully reproduce the
107 observed intensity-dependence of the 99th percentile IRs of TCs in the best-track data over the
108 North Atlantic, central, eastern and western North Pacific during 1980–2020, indicating that the
109 dynamical system models developed by Wang et al. (2021a, b) can reliably estimate the PIR of
110 real TCs. More recently, the DBDS model has been extended to include the frictional dissipative
111 heating effect by Wang et al. (2022) and refined in several aspects in Wang et al. (2023). As
112 demonstrated by Wang et al. (2022), by including the frictional dissipative heating effect, the
113 skill of the dynamical system model in capturing the observed TC PIR can be further improved,
114 in particular for those extremely strong TCs in which dissipative heating can contribute
115 positively to the PIR of intense TCs and also the TC MPI (Bister and Emanuel 1998).

116 Although the EBDS or DBDS model so far developed can capture the PIR of the observed
117 TCs (Xu and Wang 2022; Wang et al. 2022) and the intensity evolution of idealized simulated
118 TCs (Wang et al. 2021a, b), it is desirable to include the environmental effects on TC intensity
119 change so that the theoretical model can be used to evaluate the effects of environmental factors
120 on the observed TC intensity change, including both intensification and weakening. This is a
121 key step toward the application of the model to TC intensity prediction. The present study

122 attempts to extend the most recent DBDS model developed in Wang et al. (2022b) by including
123 the environmental effects to allow the model to be used to estimate the effects of various
124 environmental factors on TC intensity change in observations. As mentioned in Wang et al.
125 (2021a, b), the environmental effects on TC intensity change can be included/explained by either
126 reducing the dynamical efficiency of the TC system in the EBDS model or their ventilation
127 effects to reduce the degree of the moist neutrality of eyewall ascent in the DBDS model, as also
128 briefly discussed in section 2. This allows the evaluation of the environmental effects on TC
129 intensity change, independent of the TC intensity change induced by the TC internal dynamics.

130 The main objectives of this study are to construct the DBDS model by including the
131 environmental effects and to develop a generic framework based on the Gradient Boosted
132 Decision Trees (GBDT) to quantify the relative importance of various environmental factors to
133 the observed TC intensity change based on the TC best-track data. Instead of the use of classic
134 linear statistical methods, this study develops a machine learning framework to objectively
135 quantify the relative importance of various environmental factors to the observed TC intensity
136 changes. An advantage of the framework is to allow the potential dependence of environmental
137 influences on the stage of TC development to be considered. Machine learning, artificial neural
138 network methods have been widely used to deal with systems that involve complex nonlinear
139 interactions, and have been shown to improve skills of statistical TC intensity prediction
140 schemes to some extent (e.g., Baik and Hwang 1998; Baik and Paek 2000; Lee et al. 2000;
141 DeMaria et al. 2022; Griffin et al. 2022).

142 The rest of this paper is organized as follows. The modification to the DBDS model by
143 including the environmental effects, data, and analysis methods are described in section 2. The
144 overall environmental ventilation effect and the relative importance and contributions of various
145 environmental factors to TC intensity change are analyzed and discussed in section 3. Case
146 studies for Hurricanes Katrina (2005) and Jose (2017) and Typhoon Hagibis (2019) in the study
147 period are provided in section 4 to demonstrate the validity of the results discussed in section 3.
148 The main conclusions are given in the last section.

149 **2. Model, data, and methodology**150 *a. The DBDS model including the environmental effects*

151 The DBDS model with the effect of frictional dissipative heating included recently
 152 developed by Wang et al. (2022) was extended to include the environmental ventilation effect in
 153 this study. As mentioned in section 1, Wang et al. (2022) showed that with dissipative heating
 154 included, the theoretical model can better reproduce the observed intensity-dependence of the
 155 observed PIR and also results in a high PIR for strong TCs. The simple time-dependent equation
 156 of TC intensification including the effect of dissipative heating has the following form [see Eq.
 157 (8) in Wang et al. (2022)],

$$158 \quad \frac{\partial V_m}{\partial \tau} = \frac{\alpha C_D}{h} \left\{ A V_{Empi}^2 - \left[1 - \gamma A \varepsilon \left(1 - \frac{\delta C_k}{2\gamma C_D} \right) \right] V_m^2 \right\}, \quad (1)$$

159 where $\frac{\partial V_m}{\partial \tau}$ is the rate of TC intensity change with τ being time; V_m is the near-surface maximum
 160 wind speed (referred to as the TC intensity); V_{Empi} is the MPI without the dissipative heating
 161 effect included as in Emanuel (1997); α is the reduction factor of the 10-m wind speed from
 162 the depth-averaged boundary layer wind speed; C_D and C_k are the surface drag and exchange
 163 coefficients, respectively; h is the estimated depth of the well-mixed boundary layer; $\varepsilon = \frac{T_s - T_0}{T_s}$
 164 is the thermodynamic efficiency of the Carnot heat engine (Emanuel 1986), with T_s and T_0 being
 165 the underlying SST and the outflow layer air temperature, respectively; δ is a tracking parameter
 166 to switch the possible effect of dissipative heating on surface heat flux as advocated by Edwards
 167 (2019); and γ is the percentage of the frictional dissipation converted to internal dissipative
 168 heating to warm the atmospheric surface layer; A is the an *ad hoc* parameter measuring the degree
 169 of the moist neutrality of eyewall ascent.

170 The MPI without considering the dissipative heating effect in Eq. (1) is given as

$$171 \quad V_{Empi} = \sqrt{\frac{C_k}{C_D} \varepsilon (\kappa_o^* - \kappa_a)}, \quad (2)$$

172 where κ_o^* is the saturated enthalpy of the ocean surface at a given SST (T_s), κ_a the enthalpy of
 173 the atmosphere near the surface. From Eq. (1), we can get the steady-state intensity, namely the
 174 MPI with the dissipative heating effect included, as given below

175
$$V_{mp_i} = \frac{V_{Empi}}{\sqrt{1-\gamma\varepsilon\left(1-\frac{\delta C_k}{2\gamma C_D}\right)}}. \quad (3)$$

176 The *ad hoc* parameter A in Eq. (1) without any unfavorable environmental effects was assumed
 177 as a function of the relative intensity in Wang et al. (2021b). They also mentioned that the
 178 unfavorable environmental effects can be introduced as a ventilation parameter that reduces the
 179 degree of moist neutrality of eyewall ascent. Namely, we can assume

180
$$A \cong B \left(\frac{V_m}{V_{mp_i}} \right)^n, \quad (4)$$

181 where B is a parameter ($0 < B \leq 1$) representing the ventilation effect of all unfavorable
 182 environmental factors, and $n = 3/2$ based on calibrations using results from idealized full-
 183 physics numerical simulations (Wang et al. 2021b) and observations using best-track data (Xu
 184 and Wang 2022). If there is no any unfavorable environmental effect, $B = 1.0$, indicating an
 185 intensifying TC can reach its PIR. Under more general conditions with various environmental
 186 effects, B (<1.0) can be decomposed into the following form,

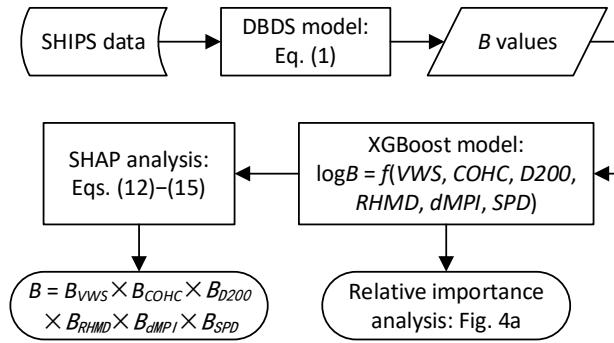
187
$$B = B_1 \times B_2 \times B_3 \cdots, \quad (5)$$

188 where B_i ($i = 1, 2, 3, \dots$) is the ventilation parameter of the i th environmental factor, such as
 189 the environmental VWS, the mid-level environmental moisture, and so on (see section 2b). The
 190 main objective of this study is to determine the environmental ventilation effects using the TC
 191 best-track data and various environmental variables from the Statistical Hurricane Intensity
 192 Prediction Scheme (SHIPS) dataset and machine learning algorithm.

193 If not otherwise stated, all parameters through Eq. (4) and constants in the DBDS model Eq.
 194 (1) are taken the same as those used in Wang et al. (2022), except for B included in A . Namely,
 195 $\delta = 1$ and $\gamma = 0.8$ were used in this study. The effect of dissipative heating on surface heat flux is
 196 included, and 80% of work done by surface friction is converted to dissipative heating (Wang et
 197 al. 2022); and values for several other parameters are $C_D = 2.4 \times 10^{-3}$, $C_k = 1.2 \times 10^{-3}$, $h =$
 198 2000 m, and $\alpha = 0.75$. These were shown to give the best fits of the results from full-physics
 199 model simulations (Wang et al. 2021a, b) and observations based on TC best-track data (Xu and
 200 Wang 2022), and all are quite reasonable under TC conditions and thus will be used in our

201 following analyses as well.

202 With the DBDS model introduced, in the following we present our approach to the
203 multiplicative decomposition of B expressed by Eq. (5) and the subsequent analyses accordingly.
204 In order to make the description easy to follow, we first present a flowchart of our approach (Fig.
205 1). Details can be found in the following subsections.



206 **Figure 1.** Workflow of the adopted approach to main objectives of this study.

207 *b. Data*

208 The data used in this study were obtained from the statistical hurricane intensity prediction
209 scheme (SHIPS) database (DeMaria and Kaplan 1999; Knaff et al, 2005), which was updated
210 on May 4, 2022. The best-track data of TCs over the North Atlantic, the central and eastern
211 North Pacific during 1982–2021 and those over the western North Pacific during 1990–2020,
212 were considered in our analysis. The SHIPS variables are from the Climate Forecast System
213 Reanalysis (CFSR) for 1982-2000 but operational Global Forecast System (GFS) analyses for
214 2001-present for the Atlantic, eastern and central Pacific, and from CFSR from 1982-2004 and
215 operational GFS for 2005-present for the western Pacific. The TC translation speed was
216 calculated from the difference between the TC location changes at 6-h intervals. To minimize
217 the influence of TC translation on its intensity, 40% of the TC translation speed was subtracted
218 from the original 6-hourly maximum sustained 10-m wind speed for all TCs, and the result was
219 used as the measure of TC intensity (V_m) as in Emanuel et al. (2004). The TC intensity changes
220 at 6-h intervals were calculated accordingly ($IR_{6h} = V_m^{t+6h} - V_m^t$). Only TCs with their V_m

222 greater than 17 m s^{-1} were included in our analysis. Only TCs south of 35°N with tropical nature
 223 and with SST greater than 25°C were considered in our analysis to avoid extratropical transition
 224 stages. All landfalling TCs were removed. As in Xu and Wang (2022), the MPIs from the SHIPS
 225 dataset were multiplied by a factor of 1.11 with the dissipative heating effect considered as in
 226 Wang et al. (2022) to consider the superintensity nature of TCs. After such modifications, all
 227 TC MPIs were greater than or equal to their corresponding intensities in the best-track data so
 228 that no negative PIR existed for any intensifying TCs in our analysis.

229 Six major environmental factors in the SHIPS dataset were selected and their effects on TC
 230 intensity changes were evaluated in this study. They are the environmental VWS defined as the
 231 magnitude of the vector wind difference between 850 and 200 hPa, the climatological ocean
 232 heat content (COHC), the upper-level divergence at 200 hPa (D200), the relative humidity (RH)
 233 between 500-700 hPa averaged between 200-800 km from the TC center, and the TC translation
 234 speed. To take into account the change in SST due to TC motion (e.g., Wood and Ritchie 2015;
 235 Fei et al. 2020), the MPI difference between t_0 and t_{0+6h} (dMPI) is considered as a proxy. Note
 236 that the effect of environmental sounding (vertical stratification of temperature and moisture)
 237 was included in the MPI calculation using the algorithm described in Bister and Emanuel (2002)
 238 and thus was not considered as an independent environmental factor herein. Table 1 lists the TC
 239 6-hourly maximum sustained 10-m wind speed and environmental variables/factors evaluated
 240 in this study.

241 **TABLE 1.** The factors analyzed in this study with their units and descriptions.

Variables	Units	Descriptions
V_m	m s^{-1}	Current TC intensity calculated by subtracting 40% of the translation speed from the best-track data
V_{Empi}	m s^{-1}	Maximum potential intensity (Emanuel 1986)
VWS	m s^{-1}	Deep-layer vertical wind shear defined as vector difference of winds averaged within 200-800 km between 850 and 200 hPa
COHC	kJ cm^{-2}	Climatological ocean heat content
D200	10^7s^{-1}	Divergence averaged within a radius of 1000 km from the TC center at 200 hPa

RHMD	%	Mean 500-700 hPa RH averaged between 200-800 km from the TC center
dMPI	m s^{-1}	MPI difference between t_0 and $t_{0+6\text{h}}$ along TC track
SPD	m s^{-1}	Translation speed of the TC system.

242 *c. Machine learning methods*

243 To quantify the environmental effects as a whole and the effects of individual environmental
 244 factors, a two-stage machine learning approach was adopted: first, eXtreme Gradient Boosting
 245 (XGBoost) (Chen and Guestrin 2016) was used to build a black-box but exact model of (log) B
 246 as a multiplication of all individual ventilation components (B_i) of the six selected environmental
 247 factors; then SHapley Additive exPlanations (SHAP) technique (Lundberg et al. 2020) was used
 248 to transform the black-box model of (log) B into an additive model, equivalent to a multiplicative
 249 model of B . The final multiplicative form of B was used to quantify the effects of all individual
 250 factors.

251 i) XGBoost

252 The XGBoost algorithm is a popular implementation of boosted regression trees (Friedman,
 253 2001). Gradient boosting optimizes a loss function by iteratively adding a set of decision trees
 254 into an ensemble. Each new tree is added sequentially such that it reduces the aggregate error
 255 from the existing ensemble of trees. At each iteration k , for the i -th sample y_i with an input
 256 feature vector \mathbf{x}_i , the estimate of y_i is updated by a decision tree $f^{(k)}(\mathbf{x}_i)$:

257
$$\hat{y}_i^{(k)} = \hat{y}_i^{(k-1)} - \alpha f^{(k)}(\mathbf{x}_i), \quad (6)$$

258 in which α denotes the learning rate, typically chosen to be less than 1, such that only a small
 259 portion of each new tree is added to the overall estimate at each iteration. To construct the
 260 decision tree $f^{(k)}$, the training data is split into left (I_L) and right (I_R) nodes based on its input
 261 features \mathbf{x} by maximizing the loss reduction, or *gain*:

262
$$G = \frac{1}{2} \left[\frac{\left(\sum_{i \in I_L} g_i \right)^2}{\sum_{i \in I_L} h_i + \lambda} + \frac{\left(\sum_{i \in I_R} g_i \right)^2}{\sum_{i \in I_R} h_i + \lambda} - \frac{\left(\sum_{i \in I} g_i \right)^2}{\sum_{i \in I} h_i + \lambda} \right] - \gamma, \quad (7)$$

263 where λ and γ are regularization parameters controlling the model complexity, $I = I_L \cup I_R$,
 264 and g_i , h_i are the gradient and hessian, respectively, with respect to $\hat{y}_i^{(k)}$ of a differentiable
 265 loss function to be minimized (e.g., the mean-squared error).

266 When constructing the decision tree $f^{(k)}$, Eq. (7) is evaluated at each node to find the best
 267 possible split gain G^* among all features in the input \mathbf{x} . Typically a split is made if the gain
 268 exceeds a certain threshold. If no split is made, the node becomes a leaf and the optimal leaf
 269 weight w_j^* can be calculated by

$$270 \quad w_j^* = -\frac{\sum_{i \in I_j} g_i}{\sum_{i \in I_j} h_i + \lambda}, \quad (8)$$

271 in which $j \in \{0, 1, 2, \dots, T\}$, with T the total number of leaves in the tree. For a particular sample
 272 y_i , the optimized $f^{(k)}(\mathbf{x}_i)$ is then simply the leaf weight w_j^* , or

$$273 \quad \hat{y}_i^{(k)} = \hat{y}_i^{(k-1)} - \alpha w_j^*. \quad (9)$$

274 A direct inference from the model fitting is the feature importance. Importance is a relative
 275 score that indicates the fractional contribution of each feature to the model performance measure,
 276 and is 100% when summed over all features. XGBoost provides a natural measure of feature
 277 importance, by first summing the gain of a feature's splits within a single tree [Eq. (7)], weighted
 278 by the number of related observations, and then averaged across all of the trees within the model.
 279 Once the model is fitted, importance is also evaluated accordingly for each feature. The XGBoost
 280 algorithm was adopted here to investigate the nonlinear relationships between the environmental
 281 ventilation parameter and environmental variables. The fitted model is much like a proxy model
 282 that encodes such relationships.

283 It is well-known that there is a bias/variance trade-off in machine learning. An overfitted
 284 model may have low bias but also poor predictive ability. For a more accurate prediction the
 285 model fitting must be controlled to allow some bias. However, it is less-known that high bias
 286 can also result in poor model interpretability (Lundberg et al. 2020). Low-bias models can better
 287 represent the true data-generating mechanism and depend more naturally on their input features,
 288 so that their interpretations of relationships in data are more stable and reliable. Since the purpose

289 of this study is to make use of the XGBoost algorithm to interpret the relationship between B
290 and various environmental factors rather than to predict B , we simply fit the model as accurate
291 as possible for the training data, without further parameter tuning as in the usual machine
292 learning practice.

293 ii) SHAP

294 SHAP is an additive feature attribution method that attributes values to each feature as the
295 change in the expected model prediction when conditioning on that feature. Its main advantages
296 are local accuracy and consistency in global model structure (Lundberg et al. 2020). Local
297 accuracy states that when approximating the original model f (e.g., a fitted XGBoost model) for
298 a specific input \mathbf{x} , the SHAP values ϕ_i for each feature i should sum up to the output $f(\mathbf{x})$:

$$f(\mathbf{x}) = \phi_0(f) + \sum_{i=1}^M \phi_i(f, \mathbf{x}), \quad (10)$$

300 i.e., the sum of feature attributions $\phi_i(f, \mathbf{x})$ matches the original model output $f(\mathbf{x})$, where
301 $\phi_0(f) = E[f(\mathbf{X})]$ is the bias term. Consistency means that if a model changes so that some
302 feature's contribution increases or stays the same regardless of the other inputs, that input's
303 attribution should not decrease.

304 These SHAP values form an additive feature attribution measure to interpret complex
305 machine learning models. SHAP values estimate contributions of each feature to each individual
306 prediction. For a given predictor and a given sample, the SHAP value is the difference in the
307 output depending on if the model is fitted with or without the predictor. For each sample, the
308 sum of all SHAP values, plus the bias term (the overall mean of predictions), equals the
309 prediction from the XGBoost model. The resulting matrix of SHAP values can be summarized
310 to understand how a predictor contributes to the predictions. The mean absolute SHAP value
311 across all samples summarizes the global feature importance, and more local model
312 interpretation is possible through exploratory data visualizations such as scatterplots of
313 individual predictors versus their corresponding SHAP values.

314 iii) Multiplicative model of B .

Combining Eqs. (1) and (3), the environmental ventilation parameter B can be calculated for all TC cases in using the TC best-track dataset, and are used as ‘observations’ to build a multiplicative model as Eq. (5). This is achieved by adopting the two-stage approach described above: we first fit a XGBoost model to $\log B$ instead of B , to capture the nonlinear relationship between $\log B$ and the selected environmental factors

$$\log B = f(VWS, COHC, D200, \dots); \quad (11)$$

321 where VWS, COHC, D200, ... indicate various environmental factors/parameters listed in
 322 Table 1. Then, by means of the SHAP values, Eq. (11) is assumed to have the following additive
 323 form:

$$\log B_i = G + \sum_{j=1}^6 S_{ij}, \quad (12)$$

325 where G is the bias term (the overall mean of $\log B$); S_{ij} ($j = 1, 2, \dots, 6$) are SHAP values
 326 corresponding to the six features in Eq. (11) for the i^{th} sample. G is an undesirable term for
 327 reaching a multiplicative model as Eq. (5). Since all parameters in Eq. (5) are between 0 and 1,
 328 the higher the value is the weaker the ventilation effect, and vice versa. As a result, all terms in
 329 Eq. (12) should be negative. By proportionally allocating G to each SHAP value according to
 330 their global feature importance $I_j = \frac{1}{N} \sum_{i=1}^N |S_{ij}|$, which is the mean absolute SHAP value across
 331 all samples, we have:

$$S'_{ij} = \frac{G \times I_j}{\sum_{j=1}^6 I_j} + S_{ij}, \quad (13)$$

333 so that Eq. (12) can be rewritten as:

$$\log B_i = \sum_{j=1}^6 S'_{ij}. \quad (14)$$

Such an allocation scheme still holds the local accuracy and global consistency properties of SHAP values. Defining $S'_{ij} = \log(B_{ij})$ where $B_{ij} \in (0,1]$, and taking the exponential function of both sides of Eq. (14), we finally obtain the sample-specific multiplicative model of the environmental ventilation parameter:

$$B_i = \prod_{j=1}^6 B_{ij}. \quad (15)$$

340 Note that SHAP values S_{ij} cannot be guaranteed to be all negative, and thus neither can

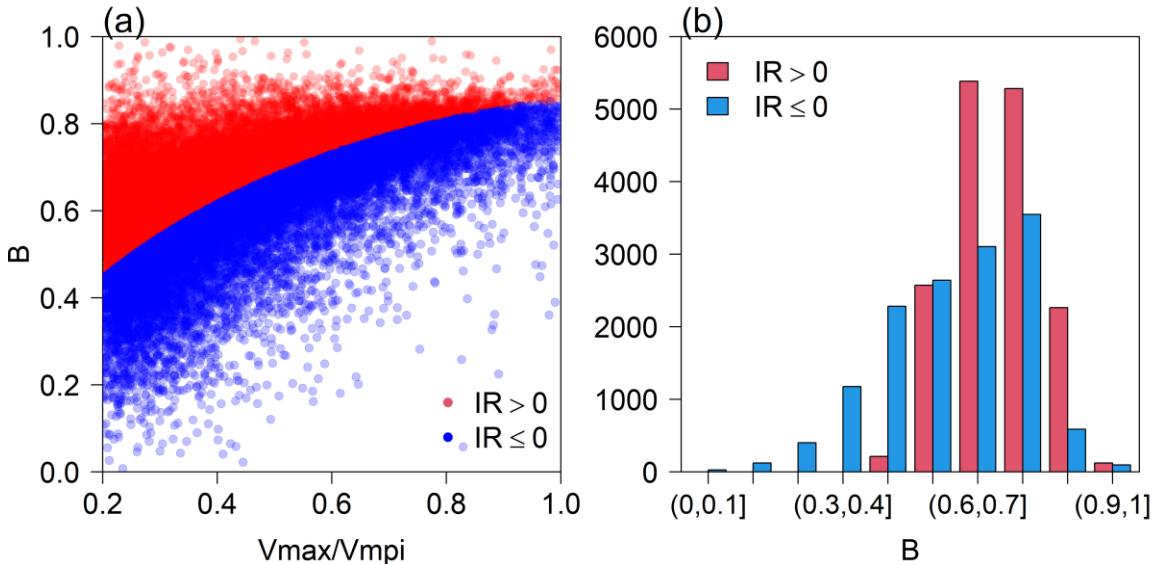
341 S'_{ij} . Consequently, for few cases with B_{ij} greater than 1 (all less than 1.2 in our analysis), we
 342 simply set such B_{ij} to be 1 in our following analysis. This does not affect the results.

343 **3. Results**

344 *a. The characteristics of the environmental ventilation B*

345 Figure 2 shows the calculated environmental ventilation parameter B in Eq. (1) against the
 346 relative intensity and compares the frequency distributions of B for intensifying and decaying
 347 TC cases using the 6-h TC best-track data. Since B is less than 1.0 for all TC cases in Fig. 2a,
 348 the environmental ventilation effect inhibits TC intensification (with $IR < PIR$) or makes TCs
 349 weaken (with $\partial V_m / \partial \tau < 0$). From Eqs. (1)–(4), we can get

$$350 \frac{\partial V_m}{\partial \tau} \geq 0, \text{ when } B \geq \frac{\left(\frac{V_m}{V_{mpi}}\right)^{\frac{1}{2}}}{1 + \gamma \epsilon \left(1 - \frac{\delta C_k}{2 \gamma C_D}\right) \left[1 - \left(\frac{V_m}{V_{mpi}}\right)^2\right]}, \quad (16)$$



351
 352 **Figure 2.** (a) Estimated B against relative intensity (V_{max}/V_{mpi}) for $IR \leq 0$ (blue) and $IR > 0$ (red) based on
 353 the TC best-track data using Eq. (1) with $\delta = 1, \gamma = 0.8, C_k = 1.2 \times 10^{-3}, C_D = 2.4 \times 10^{-3}, h = 2000$ m,
 354 and $\alpha = 0.75$, and (b) the frequency distributions of B for intensifying (red) and weakening (blue) TC cases,
 355 respectively.

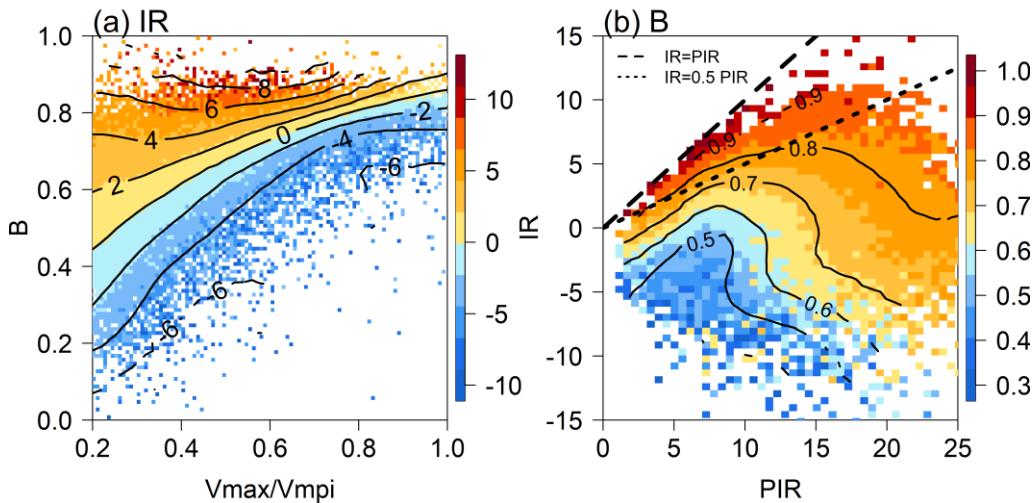
356 Eq. (16) indicates that a TC can intensify only when the ventilation parameter related to
 357 unfavorable environmental effects exceeds a critical value. Note that stronger ventilation

358 corresponds to smaller ventilation parameter B . The critical value depends on the relative
359 intensity of a TC. This can be clearly seen from Fig. 2a, in which all intensifying cases (red) are
360 located above the weakening cases (blue) for a given relative intensity. This indicates that the
361 TC weakening results primarily from strong environmental ventilation effect. From Fig. 2b, we
362 can see that overall B is greater than 0.4, with high frequency when B is between 0.6–0.8 for
363 intensifying cases, but B is evenly distributed for decaying cases between 0.3–0.8 with relatively
364 high frequency when B is between 0.5–0.7, suggesting that B can well reflect the environmental
365 ventilation effect on slowing down the TC IR or driving TC weakening. The higher B reflects
366 the more favorable environmental conditions for TC intensification. When $B = 1$ in Eq. (4),
367 Eq. (1) results in the PIR of an intensifying TC, indicating all environmental factors are favorable
368 for a TC to intensify.

369 Figure 3 further shows the TC IR (intensifying $IR > 0$ and weakening $IR < 0$) as a function
370 of B and relative intensity, and B as a function of IR and PIR, respectively. We can see from Fig.
371 3a that although IR shows a general tendency to increase with increasing B , the dependence of
372 IR on B for intensifying TC cases is much stronger than that for weakening TC cases. Particularly,
373 the rapid intensification (RI) cases with IR greater than $4 \text{ m s}^{-1}(6h)^{-1}$ for the 95th of all IR samples
374 occur with B greater than 0.7. For the weakening cases, the slow weakening cases occur with B
375 between 0.3 and 0.8, while the rapid weakening (RW) cases with IR less than $-4 \text{ m s}^{-1}(6h)^{-1}$ occur
376 with B between 0.2 and 0.7. This suggests that TCs can weaken in a large range of adverse
377 environmental conditions. Especially, as a TC approaches its MPI at its higher relative intensity,
378 the TC IR is very sensitive to the environmental effects. In those cases, even relatively weak
379 environmental effects may lead to TC weakening. However, the RW cases occur with small B ,
380 indicating that RW often results from strong adverse environmental effects, such as strong
381 environmental VWS.

382 The maximum IR occurs with B greater than 0.9 and relative intensity around 0.6. This is
383 consistent with the theoretical results in Wang et al. (2021b), which showed that the theoretical
384 maximum PIR occurs at intermediate TC intensities (roughly 60% of their MPIS). The larger

negative IR [$< -6 \text{ m s}^{-1} (6h)^{-1}$] occurs with B either being small (less than 0.4) when the relative intensity is relatively smaller than 0.5 or being between 0.6 and 0.7 when the relative intensity is relatively high around 0.8–0.9. This indicates that only strong adverse environmental effects can lead to RW of a TC in its primary intensification stage before reaching its maximum PIR, but relatively weak adverse environmental effects can lead to RW of a TC when it is close to its MPI as already mentioned above. This is consistent with the results by Fei et al. (2020), who statistically studied the RW of TCs over the western North Pacific and found that there were 86.1% of TCs undergoing their first weakening phase and about 29.4% of RW cases undergoing their first RW period within 24 hours after they reached their lifetime maximum intensity. The latter was recently studied in more detail by Zhou et al. (2022).



395 **Figure 3.** (a) Distribution of TC IR [$\text{m s}^{-1} (6h)^{-1}$, contours and shading] in B and relative intensity (V_{max}/V_{mpi})
396 space, (b) the distribution of B in IR and PIR [$\text{m s}^{-1} (6h)^{-1}$, contours and shading] space. The black dash lines
397 in (b) denote the relative IR (namely IR normalized by the theoretical PIR) of 1.0 and 0.5, respectively.
398

399 From the distribution of B in the IR and PIR space in Fig. 3b, we can see that high B , namely
400 favorable environmental conditions, is key for TCs reaching their PIRs. For example,
401 intensifying TCs with their IR reaching 50% of their PIRs or above are only observed in the
402 environment with B greater than 0.8 (short dashed line in Fig. 3b). An interesting result is the
403 quite weak dependence of RI [with IR greater than $4 \text{ m s}^{-1} (6h)^{-1}$] on B for PIR greater than 12 m
404 $\text{s}^{-1} (6h)^{-1}$. This indicates that TCs are potentially more resistant to the adverse environmental

405 influence during their intensifying stage with relatively high PIRs (often with intermediate
406 intensities as mentioned earlier, also see Wang et al. 2021a, b), but more vulnerable when their
407 PIRs are relatively low, especially, under strong adverse environmental conditions. This indicates
408 that the intrinsic vortex dynamics is key to TC intensification, while the adverse environmental
409 influence controls the weakening of TCs. Furthermore, we can see that B shows a general
410 increasing tendency with increasing IR and a decreasing tendency with the increase of IRs from
411 their corresponding PIRs. This indicates that the adverse environmental influence plays a key
412 role in limiting the TCs from reaching their theoretical PIR. This explains why very few TCs can
413 reach their theoretical PIR in observations as seen in Fig. 3b.

414 *b. XGBoost modeling and feature importance analysis of B*

415 i) Model fitting

416 The environmental ventilation effect (parameter B) discussed in section 3a results from
417 various environmental factors, such as environmental VWS, COHC, D200, RHMD, dMPI and
418 SPD as mentioned in section 2b and listed in Table 1. In this subsection, the XGBoost model
419 described in section 2c was used to quantify contributions of those individual environmental
420 factors to $\log(B)$. Each environmental factor is an input feature to the XGBoost model for all
421 TC cases. With some typical parameter settings (learning rate = 0.5 and the maximum depth of
422 a tree = 7, refer to <https://xgboost.readthedocs.io/en/latest/parameter.html> for a detailed
423 description), the root mean-square error (RMSE) of the fitted B stabilizes at 0.0023 after about
424 2000 iterations, which is 0.23% of the range of B . This result shows that the model with the
425 identified input features/factors can well reproduce B through $\log(B)$. However, the fitting error
426 does not indicate the prediction error due to the bias/variance trade-off. To examine the model's
427 prediction skill, 10-fold cross validation of the same model was further carried out. The dataset
428 was randomly divided into 10 subsamples with equal size, each of which was used as testing
429 data with all the others pooled together as training data in turn for once. For each set of testing
430 data, the mean squared error (MSE) of B predictions was calculated. In order to eliminate the

431 potential bias caused by random division, this procedure was repeated for 10 times, yielding 100
 432 MSEs. The root mean of these MSEs can be viewed as a measure of model prediction skill,
 433 which is 0.15 for the XGBoost model. Apparently, the prediction error is much greater than the
 434 fitting error.

435 Recall that, for prediction models, there is a ubiquitous trade-off between model bias and
 436 prediction variance. Linear models have the property of high bias and low variance in general
 437 (Hastie et al. 2009). We also fitted a multiple linear counterpart of the XGBoost model as a
 438 baseline for comparison:

$$439 \quad E(\log(B)) = \beta_0 + \sum_{i=1}^6 \beta_i B_i \quad (17)$$

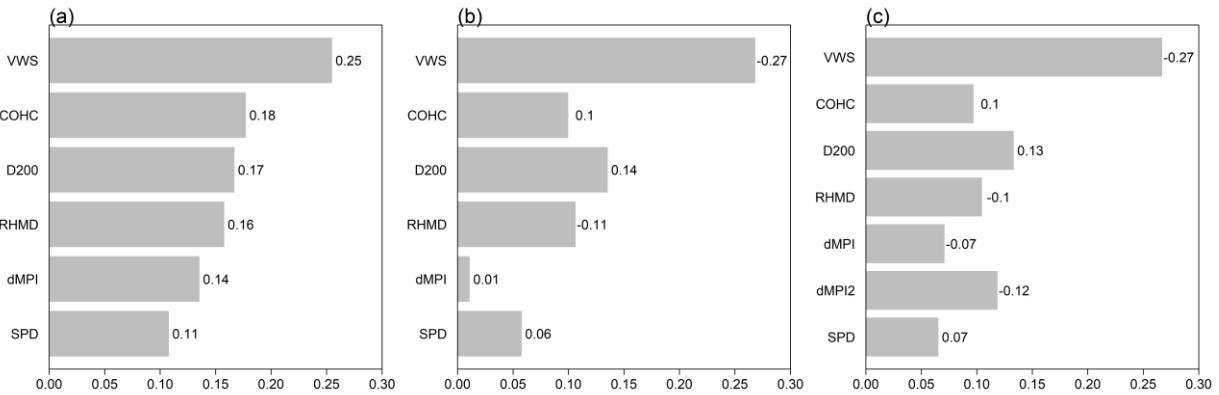
440 where β_0 is the intercept and $\beta_i, i = 1, \dots, 6$ are coefficients. Parameter estimates and their p
 441 values of significance test are shown in Table 2. It can be seen that dMPI may not be a significant
 442 linear effect at the level of 0.05 since the p value for β_5 is greater than 0.05. Using the same
 443 validation strategy, the fitting and prediction errors of B are both around 0.13. It can be seen by
 444 comparison that the fitted XGBoost model reproduces the nonlinear relationships between B and
 445 all individual environmental factors with a very small error (0.0023), much smaller than that of
 446 the counterpart linear regression model (0.13), whereas the prediction error (0.15) is a little bit
 447 larger than that of the linear model (0.13), suggesting an overfit of the XGBoost model. Since
 448 our purpose of fitting the XGBoost model is to derive a multiplicative form of B through the
 449 SHAP analysis and thus to explain which factors in B are most important, rather than to predict
 450 B for new input of environmental factors, thus an accurate or even overfitted model is acceptable.

451 **TABLE 2.** Parameter estimates and their p values for the counterpart model of Eq. (17)

Parameter	β_0 (Intercept)	β_1 (VWS)	β_2 (COHC)	β_3 (D200)	β_4 (RHMD)	β_5 (dMPI)	β_6 (SPD)
Estimate	-0.30	-8.5×10^{-4}	7.7×10^{-4}	8.5×10^{-4}	-2.4×10^{-3}	6.7×10^{-4}	3.2×10^{-3}
P values	$< 2 \times 10^{-16}$	$< 2 \times 10^{-16}$	$< 2 \times 10^{-16}$	$< 2 \times 10^{-16}$	$< 2 \times 10^{-16}$	0.059	$< 2 \times 10^{-16}$

452 ii) Feature importance analysis

453 As a direct inference of our fitted XGBoost model, the relative importance of six individual
 454 environmental factors to $\log(B)$, namely to what extent $\log(B)$ is contributed by each of the
 455 input features, are evaluated as shown in Fig. 4a. It can be seen that the environmental VWS is
 456 the most important factor and contributes 25% to $\log(B)$. Climatological ocean heat content
 457 (COHC) and the upper-level divergence (D200) contribute about 17-18% to $\log(B)$. Mid-level
 458 RH (RHMD), 6-h change in MPI along the TC track (dMPI), and translation speed (SPD)
 459 contribute, respectively, 16%, 14%, and 11% to $\log(B)$. This is broadly consistent with previous
 460 knowledge on the adverse environmental effects on TC intensity (Gray 1968; Wang and Wu 2004;
 461 Hendricks et al. 2018; Fei et al. 2020).



462
 463 **Figure 4.** (a) Relative importance of six individual environmental factors used in the XGBoost model. Factors
 464 are listed to the left (see Table 1) in descending order of their relative importance. Contributions of the
 465 individual environmental factors are given on the right of their corresponding bars. (b) Same as (a) but for the
 466 counterpart linear model of (a). Bars show absolute values of SRCs with their real values labeled to the right
 467 of their corresponding bars. (c) Same as (b) but with a quadratic term dMPI2 added to the linear model.

468 For the counterpart linear model, however, there is not straightforward to infer the relative
 469 importance. An alternative way to check the feature importance is to examine the standardized
 470 regression coefficient (SRC) (Kleijnen and Helton, 1999)

$$471 \quad SRC_i = \beta_i \frac{\sigma_i}{\sigma_Y} \quad (18)$$

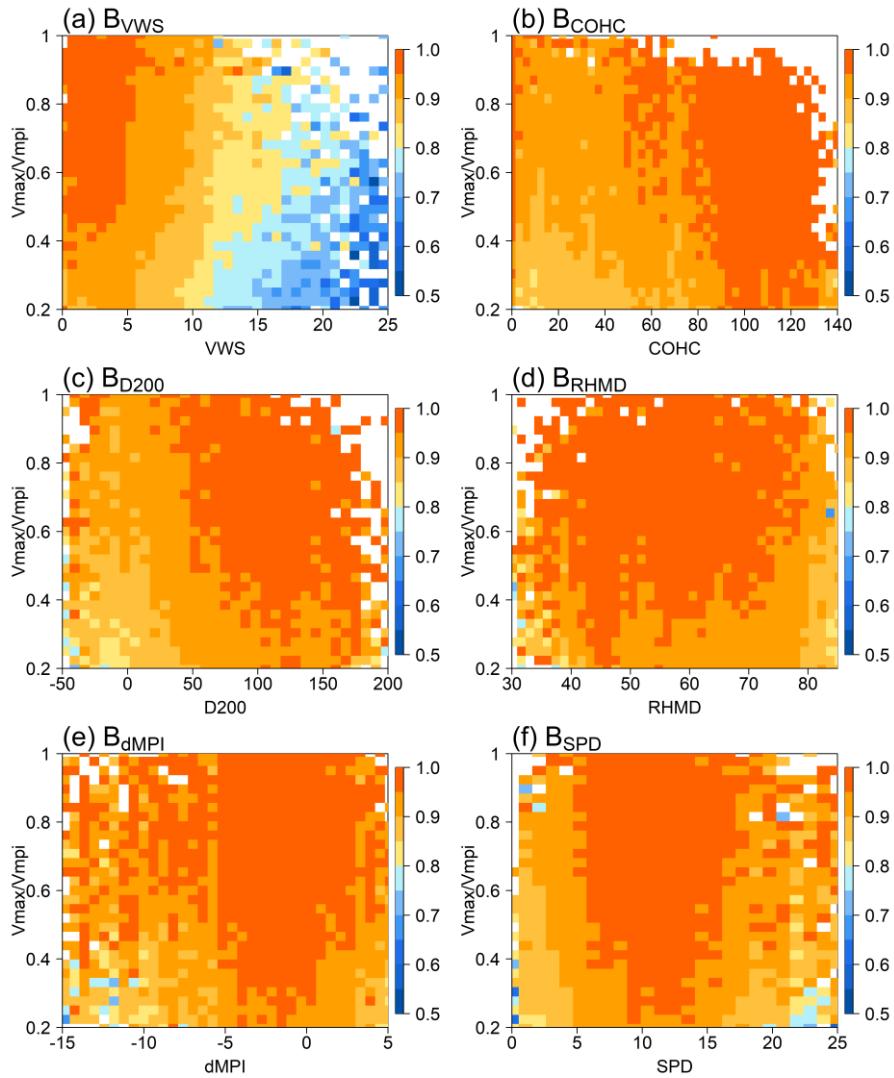
472 where σ_i and σ_Y denote the standard deviations of the model input X_i and the model output
 473 Y , respectively. This is actually a sensitivity measure representing the expected change in Y
 474 because of an increase in X_i of one of its standardized units (i.e., σ_i), with all other X variables
 475 unchanged. The absolute values of the SRCs may be compared, giving a rough indication of the

476 relative importance of the variables (but not weighted to sum to 1). Figure 4b shows the absolute
477 values of SRCs of the counterpart linear model as bars with their real values labeled. Compared
478 with Fig. 4a, VWS still has the highest relative importance; COHC, D200 and RHMD also have
479 comparable importance values to those in Fig. 4a; dMPI, however, has the least relative
480 importance. This result is consistent with the significance test of parameter estimates (Table 2),
481 by which dMPI may not be a significant linear effect. This comparison suggests that dMPI exerts
482 influence on $\log(B)$ in a nonlinear way, which has been captured by the XGBoost model. This
483 can be confirmed simply by adding a quadratic term of dMPI (denoted as dMPI2) to the right-
484 hand side of Eq. (17) and refitting the model. Significance tests of the parameters show that all
485 the seven variables, including dMPI2, have significant effects with p values less than 2×10^{-16} .
486 Relative importance of variables in the expanded model is shown in Fig. 4c. It can be seen that
487 dMPI2 gains more importance than dMPI. However, the fitting and prediction errors of B ,
488 calculated using the same method as before, are still around 0.13: reduction in each error takes
489 place only after the third decimal point digit is included, which is negligible. To sum up, the
490 multiple linear regression model can only achieve very limited improvement in the model
491 accuracy simply by adding more nonlinear terms of factors, whereas the XGBoost model can
492 reproduce the nonlinear relationship between input factors and response almost precisely,
493 without considerable loss of generalization ability. The latter merit is just what we require to
494 derive the multiplicative form of B .

495 *c. Multiplicative form of B and contributions of individual environmental factors to IR*

496 The environmental ventilation B can be expressed as the multiplication of individual
497 ventilation parameters B_i ($i = 1, 2, \dots, 6$) induced by the six environmental factors using the
498 SHAP analysis described in section 2c. Figure 5 shows the six individual environmental
499 ventilation parameters B_{VWS} , B_{COHC} , B_{D200} , B_{RHMD} , B_{dMPI} , and B_{SPD} induced by,
500 respectively, the individual environmental factors VWS, COHC, D200, RHMD, dMPI, and SPD
501 as a function of the corresponding environmental variables and relative intensity. Overall, the
502 relationship between each ventilation parameter and the corresponding variable is nonlinear and

503 depends on relative intensity of TCs.



504
505 **Figure 5.** Individual ventilation parameters (a) B_{VWS} , (b) B_{COHC} , (c) B_{D200} , (d) B_{RHMD} , (e) B_{dMPI} , and (f)
506 SPD induced by, respectively, VWS (m s⁻¹), COHC (kJ cm⁻²), D200 (10⁻⁷ s⁻¹), RHMD (%), dMPI (m s⁻¹), and
507 SPD (m s⁻¹) as a function of the corresponding environmental variables and relative intensity obtained using
508 the SHAP analysis.

509 The ventilation parameter B_{VWS} induced by the most unfavorable environmental factor
510 VWS varies between 0.5–1.0 (Fig. 5a). B_{VWS} is generally greater than 0.9 when VWS is less
511 than 7 m s⁻¹ but decreases significantly with increasing VWS afterwards. This suggests that weak
512 environmental VWS has very limited effect on TC intensity change but imposes an increasing
513 adverse effect on TC IR as VWS increases beyond 8 m s⁻¹. This agrees with previously reported
514 threshold of about 8–10 m s⁻¹ above which VWS can have a significant detrimental effect on TC

515 intensity and intensification (Zeng et al. 2010; Wang et al. 2015; Hendricks et al., 2018). B_{VWS}
516 also shows an overall slow decrease with decreasing relative intensity, implying that
517 environmental VWS is more detrimental to relatively weak TCs than to strong TCs. The
518 ventilation parameters induced by other environmental factors are generally between 0.8–1.0
519 (Figs. 5b-5f), considerably smaller than that induced by VWS, implying that they have relatively
520 weaker adverse effects on TC intensity change than VWS.

521 The ventilation parameter (B_{COHC}) induced by COHC shows a general increasing tendency
522 with increasing COHC (Fig. 5b). This is because high ocean heat content limits the upper ocean
523 cooling induced by upwelling and vertical mixing across the mixed layer base under the TC
524 (Wang and Wu 2004). Similar to COHC, the ventilation parameter B_{D200} induced by upper-
525 level divergence (D200) also varies between 0.8–1.0 (Fig. 5c). It increases with increasing upper-
526 level divergence, suggesting that upper-level divergence (convergence) is favorable (unfavorable)
527 for TC intensification. This is because the upper-level convergence or weak divergence is
528 unfavorable for eyewall ascent, and thus plays a role equivalent to the mid-level ventilation
529 induced by lateral dry-air intrusion to reduce B_{D200} . This is consistent with previous studies by
530 Kaplan et al. (2010) and Lee et al. (2015), who found that strong upper-level environmental
531 divergence is favorable for TC intensification.

532 The ventilation parameter B_{RHMD} associated with the mid-level RH between 500–700 hPa
533 is generally high (Fig. 5d), with relatively small values for both too high (greater than 75%) and
534 too small (less than 40%) RHMD. Too high RHMD implies moist mid-level environment, which
535 is favorable for active rainbands and TC size expansion, which is often unfavorable for TC
536 intensification, as demonstrated in previous modeling (e.g., Wang 2009; Hill and Lackmann
537 2009; Li et al. 2020) and theoretical (Wang et al. 2023) studies. In contrast, too low RHMD
538 makes the eyewall ascent vulnerable to any environmental perturbations by lateral dry-air
539 intrusion. Therefore, too dry mid-level environment plays a role in enhancing the environmental
540 ventilation effect (Tang and Emanuel 2010).

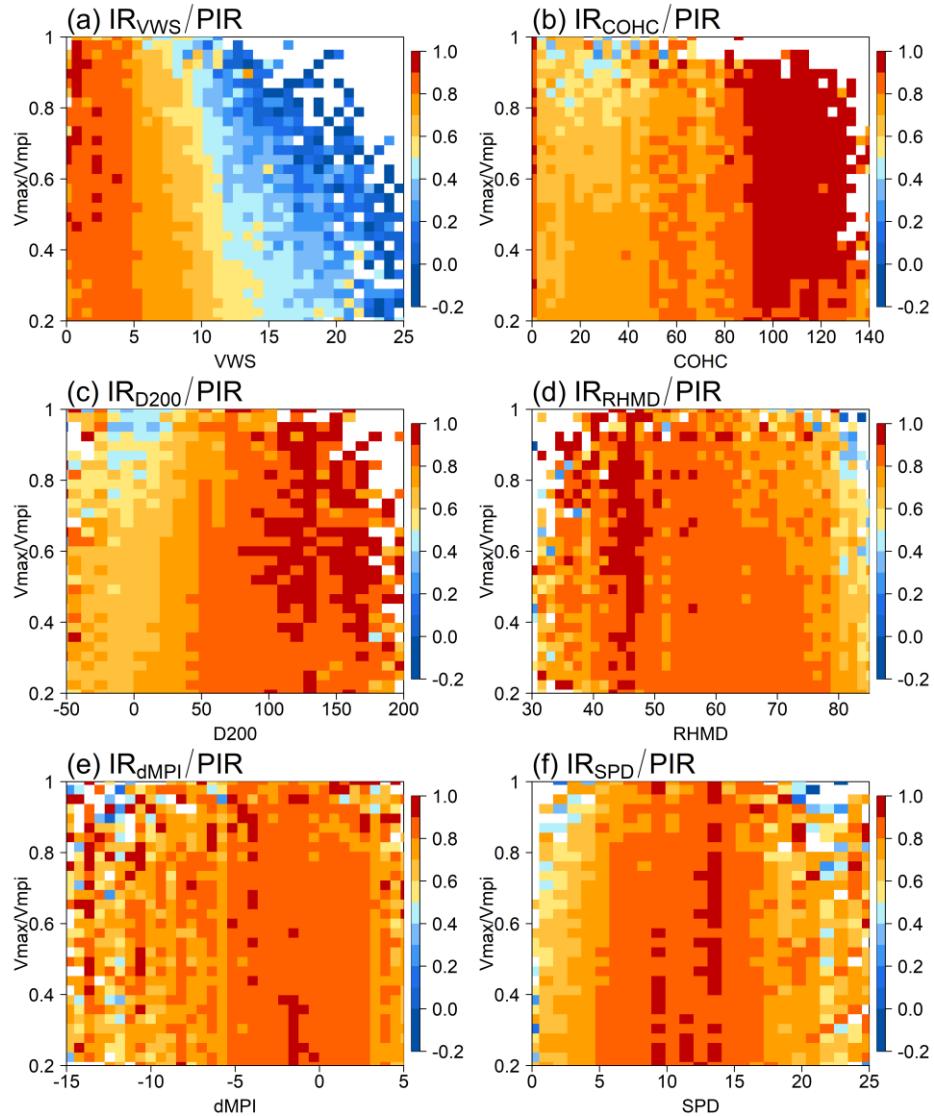
541 The factor dMPI is the change in MPI along the TC track, which is mainly determined by

542 the underlying SST gradient and the translation speed of the TC, and determines the response
543 timescale of the TC to change in the underlying SST. Positive dMPI partly reflects the potential
544 increase in eyewall convection and, consequently, the weakened ventilation (Fig. 5e). Negative
545 dMPI is equivalent to a decrease in SST, and thus increasing ventilation effect and reducing
546 B_{dMPI} . B_{dMPI} decreases with decreasing dMPI when dMPI is less than -5 m s^{-1} . As a result,
547 large negative dMPI often leads to rapid weakening of TCs, similar to the SST gradient
548 previously revealed by Wood and Ritchie and (2015) and Fei et al. (2020). However, B_{dMPI}
549 shows a decreasing trend with increasing dMPI and decreasing relative intensity for positive
550 dMPI. This may be due to the delayed response of TC intensity to the increase in SST, which is
551 more significant for weak TCs (with relative intensity less than 0.4).

552 The last factor is the TC translation speed (SPD), which has dual effects on TC intensity
553 change and thus the ventilation parameter (B_{SPD} , Fig. 5f). On one hand, too slow translation
554 (with SPD less than 6 m s^{-1}) often enlarges the negative ocean feedback due to cooling induced
555 by TC forcing. On the other hand, too fast translation (with SPD greater than 20 m s^{-1}) can induce
556 large asymmetric structure, which may lead to ventilation effect by eddy processes (Zeng et al.
557 2007, 2008). Note that fast translation has a more pronounced effect on weak TCs with relative
558 intensity less than 0.4.

559 To further quantify contributions of individual environmental factors to TC IR ($dV_m/d\tau$),
560 we calculated IR using Eq. (1) with A from Eq. (4) and B from Eq. (5). In each calculation, we
561 used the actual ventilation parameter induced by one environmental factor while keeping all
562 other ventilation factors being 1.0. For VWS as an example, the contribution by environmental
563 VWS to TC RI (IR_{VWS}), which is calculated using the actual B_{VWS} while keeping B_{COHC} , B_{D200} ,
564 B_{RHMD} , B_{dMPI} , and B_{SPD} all being 1.0, is evaluated by IR_{VWS} normalized by the PIR calculated
565 using Eq. (1) with $B = 1$ in Eq. (4). Figure 6 shows the contributions of all individual
566 environmental factors to TC IR as a function of the factor and relative intensity. The normalized
567 IR_{VWS} shows a nearly linear decrease with increasing VWS and also a decrease with increasing
568 relative intensity when VWS is larger than about $7-8 \text{ m s}^{-1}$ (Fig. 6a). This is mainly because that

569 the stronger the TCs are when approaching their MPI, there will be lower probability for them
 570 to intensify. Note that there are a few cases with negative normalized IR_{VWS} when VWS is greater
 571 than 15 m s^{-1} , consistent with the small B_{VWS} in Fig. 5a, indicating the dominant effect of VWS
 572 on TC weakening.



573
 574 **Figure 6.** Same as Fig. 5, but for the normalized TC IR induced by one of individual factors to the
 575 corresponding potential intensification rate (PIR).

576 The normalized IR_{COHC} shows a general increase with increasing COHC (Fig. 6b),
 577 indicating that high climatological ocean heat content is favorable for TC intensification. The
 578 normalized IR_{D200} (Fig. 6c) shows somewhat small values when the D200 is convergence or
 579 weak divergence, consistent with the relatively small B_{D200} value in Fig. 5c, suggesting that

580 upper-level environmental divergence reflects TC rapid intensification. The normalized IR_{RHMD} ,
581 IR_{dMPI} , and IR_{SPD} all show distributions in the parameter space similar to their corresponding
582 ventilation parameters, indicating that high middle-level RH, large negative dMPI and too slow
583 or too fast translation are all unfavorable for TC intensification. These results confirm that the
584 environmental effects on TC IR can be effectively included in our dynamical system model
585 through their corresponding ventilation parameters.

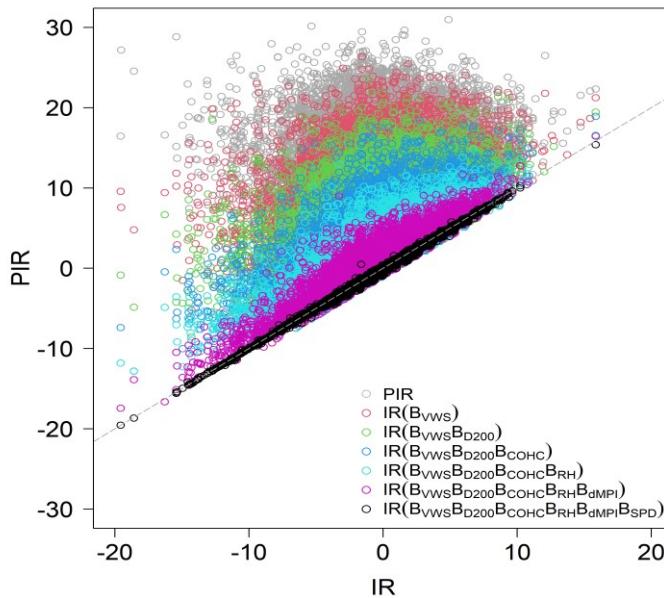
586 **4. Case studies of Hurricanes Katrina (2005) and Jose (2017) and Typhoon Hagibis (2019)**

587 In section 3, we discussed how the six environmental factors contribute to the ventilation
588 parameter B as a whole and also individually and eventually bring the theoretical PIR towards
589 the observed TC IR based on the DBDS model. This also makes it possible to objectively
590 quantify the relative contributions of various environmental factors to the observed intensity
591 change of each TC. In this section, three representative cases are used to give further insight into
592 the environmental effects on intensity change of individual TCs in terms of their lifetime
593 intensity changes including both intensification and weakening stages.

594 Before going into detailed case studies, let's first have an overview of how individual
595 environmental factors affect B and virtually bring PIR towards IR. The six environmental
596 ventilation parameters B_{VWS} , B_{COHC} , B_{D200} , B_{RHMD} , B_{dMPI} , and B_{SPD} for the whole sample
597 data can be retrieved from the database discussed in section 3. Then, we calculated a set of IRs
598 ($\partial V_m / \partial \tau$) by adding one factor each time for the six environmental effects in the above order
599 into Eq. (1) to highlight how the PIR is reduced to the actual IR ($\partial V_m / \partial \tau$) by the six individual
600 environmental factors, as shown in Fig. 7. Note that, theoretically, the final group of IRs (black)
601 should coincide with real IRs such that the dots align with the diagonal line. However, due to the
602 fitting errors from the XGBoost model propagated to the SHAP values, they scattered a bit
603 [RMSE = 0.10 m s⁻¹(6h)⁻¹]. Also note that, different orders of adding environmental ventilation
604 parameters do not make any difference in the black dots, which are only observable in Fig. 7.

605 We then retrieved the time series of the six environmental ventilation parameters B_{VWS} ,

606 B_{COHC} , B_{D200} , B_{RHMD} , B_{dMPI} , and B_{SPD} for each case individually (left column in Fig. 8), and
 607 calculated a set of IR series in the same way as done above (right column in Fig. 8). Note that
 608 the lifetime mean of individual ventilation parameters for each of the cases is given in Table 3
 609 for a quick look at the relative contributions of individual environmental factors to the observed
 610 TC intensity changes.



611
 612 **Figure 7.** Illustration of how the PIR is reduced to the actual IR by adding one of the six environmental
 613 ventilation parameters for each time. The gray dashed line is diagonal.

614 *a. Hurricane Katrina (2005)*

615 Hurricane Katrina (2005) was one of the deadliest and the costliest meteorological disasters
 616 that struck the United States on record. Katrina formed at 1800 UTC 23 August 2005 over the
 617 southeastern Bahamas. It showed few signs of weakening during its brief passage over the
 618 Florida peninsula and began to intensify shortly after moving into the Gulf of Mexico early on
 619 26 August. Two periods of RI on 26 and 28 August brought Katrina to category 5 with the
 620 maximum near-surface wind speed of 77 m s^{-1} (Knabb et. al. 2005). The environmental
 621 ventilation factors indicated a favorable environment for RI, such as weak VWS and large COHC
 622 with their lifetime mean ventilation parameters being 0.93 and 0.97, respectively (Table 3). Other
 623 environmental factors were also favorable for TC intensification, including moist RHMD,

624 positive dMPI and slower SPD than average, with their ventilation parameters being 0.95–0.96
625 (Fig. 8a). Only D200 was a little bit weaker than normal, giving rise an average ventilation
626 parameter of 0.91, which may hinder TC intensification (DeMaria and Kaplan 1999) until 0000
627 UTC 29 August. After that, the environmental VWS showed a continuous increase, leading to a
628 rapid weakening of Katrina. Figure 8b shows how the PIR was reduced to the actual IR by adding
629 one of the six environmental effects for each time, showing clearly that each *ad hoc* IR is indeed
630 an upper bound on the actual IR. The weak VWS only reduced PIR slightly, while D200 was
631 dominant in reducing PIR with the smallest B among all 6 factors (Table 3). Other factors
632 weakened PIR slightly during the intensification stage ($IR > 0$), but contributed equally during
633 the decaying stage with similar individual ventilation parameters after 1200 UTC 28 August.

634 *b. Hurricane Jose (2017)*

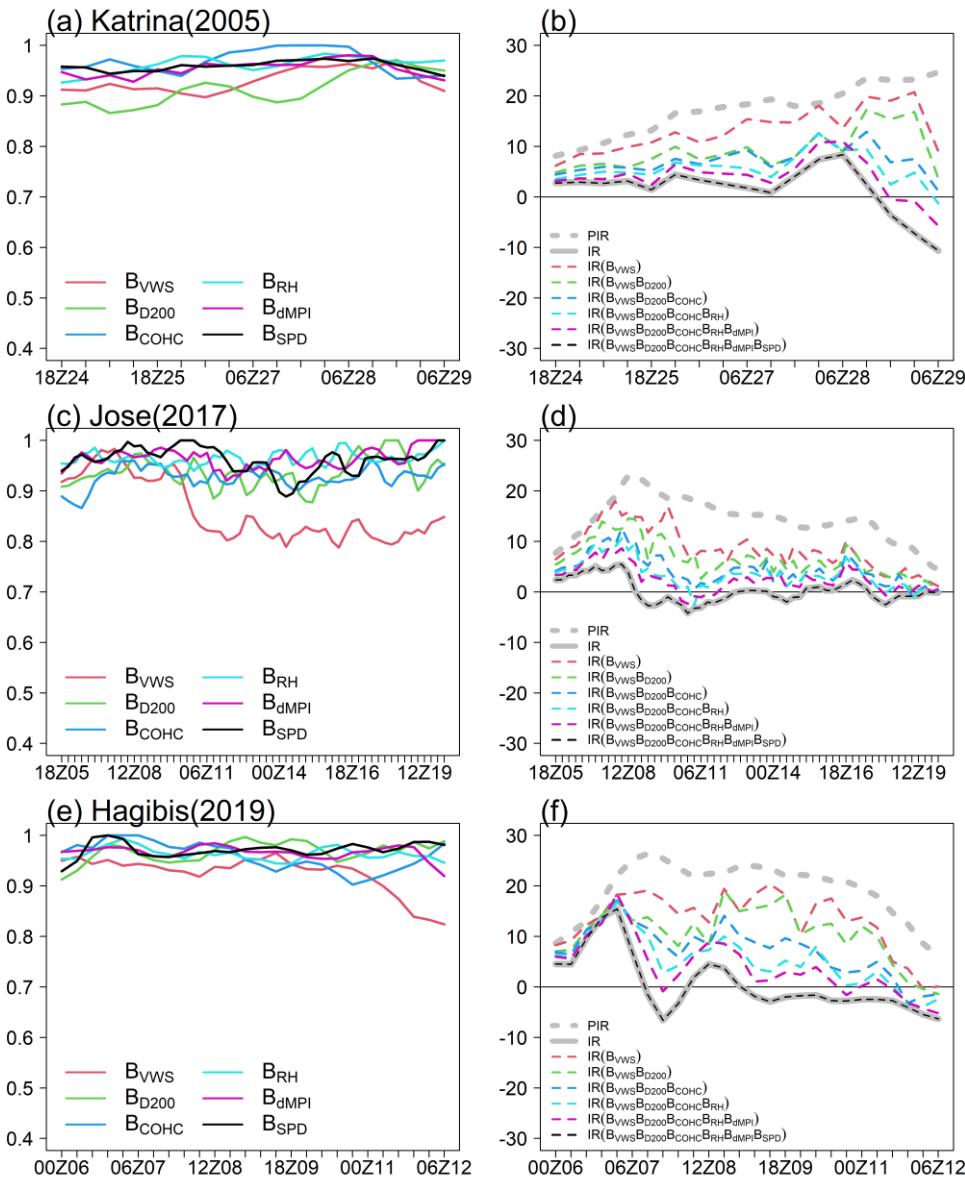
635 Hurricane Jose (2017) formed as a tropical storm by 1200 UTC 5 September west of the
636 Cabo Verde Islands, intensified to its peak intensity of 68 m s^{-1} by 1800 UTC 8, weakened and
637 then oscillated around 33 m s^{-1} for about five days, and then weakened to a tropical storm early
638 on 15 September. After re-intensifying to hurricane strength in a few days, Jose weakened to a
639 tropical storm again when it was located east of Virginia Beach and also began to take on some
640 extratropical characteristics by 1200 UTC 19 September. Jose had a long over-water lifespan of
641 a total of 14.75 days (Berg 2018). Along the long-life track of Jose, all environmental factors
642 played complicated roles in its RI, intensity fluctuation, and weakening processes. Initially, both
643 the increasing PIR and high B led to RI, making Jose attaining its lifetime maximum intensity
644 (LMI) (Figs. 8c and 8d). Jose moved northwestward after 9 through 11 September, and suffered
645 from an increasing northeasterly VWS and a partial eyewall replacement, which caused B_{vws} to
646 decrease sharply, and thus Jose weakened below hurricane intensity. For the rest of its life,
647 environmental VWS played a dominant role during its intensity fluctuation and weakening
648 processes. VWS was 3.4 m s^{-1} larger than normal average, resulting in a low B_{vws} of 0.86 (Table
649 3), which alone reduced about 42% of the PIR (Fig. 8d).

651 Super Typhoon Hagibis (2019) formed over the western North Pacific in October 2019. It
652 intensified explosively from 28 m s^{-1} at 1200 UTC 6 to 73 m s^{-1} at 1200 UTC 7 (from tropical
653 storm to Category 5), namely reached its LMI of 73 m s^{-1} . Hagibis started its RI and reached the
654 maximum IR of $15.4 \text{ m s}^{-1}(6h)^{-1}$ at 1200 UTC 7, which is very close to its PIR under a favorable
655 environment. Note that the maximum IR happened when the relative intensity (V_{max}/V_{mpi}) was
656 around 0.54, which is consistent with observation in Fig. 3a and the theoretical results in Wang
657 et al. (2021b), who showed that the theoretical maximum PIR occurs at intermediate TC
658 intensities (roughly 60% of their MPIs). After the RI, Hagibis's intensity dropped and then
659 fluctuated during 8–9 October. Actually, the environmental factors changed little during this
660 period, with individual ventilation parameters fluctuating slightly as shown in Fig. 8e. Lin et al.
661 (2020) compared the environmental conditions, such as the ocean eddy, environmental vertical
662 wind shear, and mid-level relative humidity, etc., in this period with those in the RI stage. They
663 found that some conditions, such as weak environmental VWS and warm ocean eddy were even
664 better in this period than in the RI period. As a result, they concluded that the eyewall
665 replacement cycle and the relatively large size expansion predominantly hindered Hagibis'
666 further intensification. Note that Hagibis was approaching its MPI during this period with the
667 relative intensity greater than 0.75. As we mentioned earlier, when a TC approaches its MPI,
668 there is less potential for it to intensify, and the IR becomes very sensitive to the environmental
669 effects (Fig. 3a). Hagibis terminated its strengthening at 0000 UTC 9, and turned northward and
670 moved into region with much cooler SST with relatively high VWS and low-moisture
671 environment, which led to much lower B_{vws} and B_{ohc} , as shown in Fig. 8f. During Hagibis'
672 weakening stage, the environmental factors reduced 116% of the PIR, changing from
673 intensifying to weakening. Particularly, environmental VWS alone reduced about 40% of the
674 PIR, and together with D200 and CHOC, reduced about 80% of the PIR, which dominated the
675 whole weakening period (Fig. 8f). As Hagibis moved northward toward Japan, the COHC was
676 -27.2 kJ cm^{-2} below the average (Table 3), which was also a major factor contributing to Hagibis'

677 weakening process.

678 **TABLE 3.** List of environmental factors, in standard (std) anomaly form, and individual ventilation parameters
679 B of lifetime mean of TCs Katrina, Jose and, Hagibis, respectively.

Factors	Katrina (std anomaly)/ B	Jose (std anomaly)/ B	Hagibis (std anomaly)/ B
VWS	-1.11/0.93	3.4/0.86	-0.92/0.92
COHC	37.2/0.97	-3.5/0.93	-27.2/0.96
D200	-34.4/0.91	-27.2/0.94	75.4/0.97
RHMD	3.4/0.96	-6.4/0.97	1.86/0.96
dMPI	1.81/0.95	0.59/0.97	-1.1/0.97
SPD	-1.85/0.96	-0.03/0.96	3.1/0.97



680

681 **Figure 8.** Case studies for Hurricanes Katrina (2005) (a and b) and Jose (2017) (c and d) and Typhoon Hagibis
682 (2019) (e and f), respectively. The left column shows time series of the ventilation parameter B and its
683 components due to individual environmental factors. The right column shows the PIR (grey dash, $\text{m s}^{-1} \text{ day}^{-1}$)
684 and reductions of the PIR by individual environmental ventilation parameters by B as the multiplication of
685 individual ventilation parameters B_{VWS} , $B_{VWS}B_{D200}$, ..., and $B_{VWS}B_{D200}B_{CHOC}B_{SPD}B_{dMPI}B_{RHMD}$ (colored dashes),
686 respectively, towards the observed IR (grey, solid). The effect of the dominant ventilation factor VWS is
687 highlighted by the red bold dashes for all three cases.

688 5. Conclusions and discussion

689 In two recent studies, W21a and W21b introduced a simple energetically based and a
690 dynamically based dynamical system models, or in short EBDS and DBDS models, to quantify
691 the intensification rate (IR) of a TC, respectively. Both models share the same mathematical
692 expression of TC IR as a function of the relative TC intensity and maximum potential intensity
693 (MPI). The only difference is that the dynamical efficiency (E) in the EBDS model is replaced
694 by the *ad hoc* ventilation parameter (A) measuring the degree of the moist neutrality of eyewall
695 ascent in the DBDS model. Both models have been shown to be capable of realistically capturing
696 the intensity-dependence of TC IR in both idealized full-physics model simulations and
697 observations (Wang et al. 2021b, Xu and Wang 2022). This study extends the DBDS model to
698 include the effects of various environmental factors so that the model can be used to quantify the
699 detrimental effects on IR of real TCs.

700 The environmental effect has been introduced through the environmental ventilation
701 parameter B in the DBDS model, which can be uniquely expressed as a multiplication of
702 individual ventilation parameters of various environmental factors. TC IR shows a general
703 increase with increasing B or decreasing ventilation effect. Results based on the best-track data
704 over the North Atlantic, central, eastern and western North Pacific during 1982–2021 show that
705 the dependence of TC IR on B for intensifying TC cases is much stronger than that for weakening
706 TC cases. Particularly, the rapid intensification [RI, with IR greater than $4 \text{ m s}^{-1}(6h)^{-1}$] cases
707 occur with B greater than 0.7. For the weakening cases, the slow weakening cases occur with B
708 between 0.3 and 1.0, while the rapid weakening [RW, with IR less than $-4 \text{ m s}^{-1}(6h)^{-1}$] cases
709 occur with B between 0.2 and 0.7. Especially, as a TC approaches its MPI with high relative

710 intensity, the TC IR is very sensitive to the environmental effects. In these cases, even relatively
711 weak environmental effects may lead to TC weakening. An interesting result is the quite weak
712 dependence of RI on B for PIR greater than $12 \text{ m s}^{-1} (6\text{h})^{-1}$. This indicates that TCs are potentially
713 more resistant to the adverse environmental influence during their intensifying stage with
714 relatively high PIRs.

715 Six major environmental factors in the SHIPS dataset were selected and their effects on TC
716 intensity changes were evaluated based on the TC best-track data and the SHIPS dataset during
717 1982–2021, including the environmental deep-layer VWS, the climatological ocean heat content
718 (COHC), the upper-level divergence at 200 hPa (D200), the mid-level relative humidity (RHMD)
719 between 500–700 hPa averaged between 200–800 km from the TC center, the TC translation
720 speed (SPD), and the MPI difference between t_0 and $t_{0+6\text{h}}$ (dMPI) considered as a proxy of the 6-
721 h change in SST along the TC track. The machine learning algorithm XGBoost model was
722 adopted to quantify the relative importance of the above factors, and the SHAP method was used
723 to quantify the contribution from each factor to the observed TC intensity change. Results from
724 these analyses demonstrate that VWS is the most important environmental factor, which
725 contributes 25% to $\log(B)$. COHC and D200 contribute about 17–18% to $\log(B)$. RHMD,
726 dMPI, and SPD contribute 16%, 14%, and 11%, respectively. The ventilation parameters also
727 represent their individual relative importance to the bulk environmental ventilation parameter
728 and thus their relative contributions to the observed TC intensity changes.

729 With the SHAP analysis method, the environmental ventilation parameter B can be
730 expressed as the multiplication of individual ventilation parameters of the selected
731 environmental factors. Results show that the relationship between each ventilation parameter
732 and the corresponding variable depends on the TC relative intensity. The ventilation parameter
733 B_{VWS} induced by the environmental VWS varies between 0.5–1.0. Compared with VWS, the
734 ventilation parameters induced by other environmental factors are relatively higher and vary
735 between 0.8–1.0, implying that they have relatively weaker effects on TC intensity change than
736 VWS. Consistently, the normalized IR_{VWS} decreases almost linearly with increasing VWS and

737 also with increasing relative intensity when VWS is larger than about $7\text{--}8\text{ m s}^{-1}$, largely due to
738 the little potential for strong TCs approaching their MPI. A few cases show negative normalized
739 IR_{VWS} when VWS is greater than 15 m s^{-1} , indicating the dominant effect of VWS on TC
740 weakening. The normalized IR_{COHC} shows a general increase with increasing COHC indicating
741 that high climatological ocean heat content is favorable for TC intensification. The normalized
742 IR_{D200} shows somewhat small values when the D200 is convergence or weak divergence,
743 suggesting that upper-level environmental divergence reflects TC rapid intensification. High
744 RHMD, large negative dMPI, and too slow or too fast translation are all unfavorable for TC
745 intensification.

746 Three representative cases, namely Hurricanes Katrina (2005) and Jose (2017) and
747 Supertyphoon Hagibis (2019), are chosen to give further insight into the environmental effects
748 on intensity change of individual TCs in terms of their lifetime intensity changes, including both
749 intensification and weakening stages. Results demonstrate that the individual environmental
750 ventilation parameters can well capture the detrimental effects of various environmental factors
751 on TC PIR, while the relative importance of the environmental factors varied with case and the
752 different life stages of individual TCs. In all cases, the TC weakening results primarily from
753 strong environmental ventilation effects, with strong VWS being the major detrimental
754 environmental factor.

755 We should point out that in this study it is assumed that the DBDS model can precisely give
756 the PIR that a TC can reach under all favorable environmental thermodynamic conditions. As a
757 result, the difference between the PIR and the observed intensity change is attributed to the
758 detrimental environmental effects. Since the DBDS model is highly idealized and was verified
759 based on ensemble idealized numerical simulations and best-track TC data, it could not capture
760 the short-term intensity change resulting from high-frequency convective activities. Namely, the
761 model can be used to evaluate the storm-scale intensification. In our study, therefore, we assumed
762 that the best-track data mainly reflect the storm-scale intensity change. Our results strongly
763 suggest that this assumption is acceptable. Formally the strategy we adopted here can also be

764 used to predict the TC intensity. However, for the prediction purpose, the SHAP analysis and the
765 multiplicative decomposition of B can be skipped, whereas more parameters tuning, validation
766 and testing steps should be taken for developing the XGBoost model, or any other machine
767 learning model that can model B as response to environmental factors as input features, such as
768 neural networks. In our follow-up studies, we will apply the DBDS model to estimate the PIR
769 and conduct real-time TC intensity prediction.

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774 **Data Availability Statement**

775 The SHIPS data are downloaded from [https://rammb2.cira.colostate.edu/research/tropical-
776 cyclones/ships/#DevelopmentalData](https://rammb2.cira.colostate.edu/research/tropical-cyclones/ships/#DevelopmentalData).

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