

# 1 Methane-Carbon Budget of a Ferruginous Meromictic Lake and Implications 2 for Marine Methane Dynamics on Early Earth

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## 12 ABSTRACT

13 The greenhouse gas methane (CH<sub>4</sub>) contributed to a warm climate that maintained liquid water  
14 and sustained Earth's habitability in the Precambrian despite the Faint Young Sun. The viability of  
15 methanogenesis (ME) in ferruginous environments, however, is debated, as iron reduction can  
16 potentially outcompete ME as a pathway of organic carbon remineralization (OCR). Here we  
17 document that ME is a dominant OCR process in Brownie Lake in Minnesota, USA, a ferruginous  
18 (iron-rich, sulfate-poor) and meromictic (stratified with permanent anoxic bottom waters) system.  
19 We report ME accounting for  $\geq 90$  and  $> 9 \pm 7\%$  of the anaerobic OCR in the water column and  
20 sediments, respectively, and an overall POC loading to CH<sub>4</sub> conversion efficiency of  $\geq 18 \pm 7\%$  in  
21 the anoxic zone of Brownie Lake. Our results, along with previous reports from ferruginous  
22 systems, suggest that even under low primary productivity in Precambrian oceans, the efficient  
23 conversion of organic carbon to CH<sub>4</sub> would have enabled a major role of marine CH<sub>4</sub>.

24 **INTRODUCTION**

25 The greenhouse gas methane (CH<sub>4</sub>), with a present atmospheric concentration of 1.8 ppmv,  
26 contributes  $\leq 25\%$  of postindustrial global warming (Etminan et al., 2016). The importance of CH<sub>4</sub>  
27 to Precambrian climate may have been considerably higher, with estimated atmospheric  
28 concentration ranging from 600-3000 ppmv in Archean and 1-100 ppmv in Proterozoic (Olson et  
29 al., 2016; Fakhraee et al., 2019; Fig. 1). Methane concentrations may have exerted multiple roles  
30 in early Earth's biogeochemical evolution, among others: in contributing to greenhouse gas  
31 warming under a faint young sun to maintain warm surface temperature, liquid water, and Earth's  
32 habitability (Haqq-Misra et al., 2008); in producing an anti-greenhouse organic haze layer (Pavlov  
33 et al., 2001); in drawing down the H<sub>2</sub>-based greenhouse warming leading to a late Archean (2.9  
34 Ga) glaciation event (Wordsworth and Pierrehumbert, 2013); in contributing to hydrogen escape  
35 to space leading to oxidation of Earth's surface environment (Catling et al., 2001); in decreasing  
36 microbial methanogenesis (ME) leading to oxygen buildup in the atmosphere (Konhauser et al.,  
37 2009); and in decreasing atmospheric CH<sub>4</sub> levels contributing to the onset of Proterozoic  
38 glaciations (Zahnle et al., 2006). All these hypotheses require an active CH<sub>4</sub> cycle, likely with  
39 biological mediation. ME is one of the oldest microbial metabolic pathways, whose origin is dated  
40 back to  $>3.5$  Ga, and is considered to have played an essential role in CH<sub>4</sub> supply to the atmosphere  
41 during Earth's early history (Kharecha et al., 2005). Ferruginous conditions were a dominant  
42 feature of Earth's early oceans (Poulton, 2021; Fig 1), and so an understanding of the role of ME  
43 under ferruginous conditions is essential to our understanding of marine carbon cycling in early  
44 Earth.

45 *Fig. 1*

46 Meromictic ferruginous lakes are considered convenient analogs to Precambrian oceans (Swanner  
47 et al., 2020). Such lakes generally have large reservoirs of CH<sub>4</sub> (1–4 mM in bottom waters; Crowe  
48 et al., 2011; Lopes et al., 2011). However, estimates of how much organic carbon (OC) is degraded  
49 by ME have only been calculated for a handful of lakes with estimated POC-to-CH<sub>4</sub>  
50 remineralization efficiency varying even within a single lake, i.e., Lake Matano, from 3 to 80%  
51 (Crowe et al., 2011; Kuntz et al., 2015). Ferruginous conditions, or rather the scarcity of sulfate,  
52 likely promote ME as the dominant pathway of organic carbon remineralization (OCR) (Friese et  
53 al., 2021). Yet, some have proposed ME plays only a minor role in ferruginous oceans (Laakso  
54 and Schrag, 2019). The efficiency of ME during OCR in ancient ferruginous oceans is thus poorly  
55 constrained. Here, we report the carbon budget for OCR and ME in Brownie Lake, a ferruginous  
56 meromictic lake with a biogeochemical analogy to Precambrian oceans (Lambrecht et al., 2018),  
57 and evaluate the implications for Precambrian CH<sub>4</sub>-carbon dynamics. Our results highlight that  
58 ME is a dominant OCR pathway in sulfate-poor ferruginous aquatic systems, suggestive of large  
59 CH<sub>4</sub> storage and fluxes from the Precambrian ferruginous oceans, thereby supporting models that  
60 invoke the importance of CH<sub>4</sub> in Earth's early climate.

## 61 **STUDY SITE**

62 Brownie Lake (44°58'04" N, 93°19'26" W) is the northernmost lake in the Minneapolis Chain of  
63 Lakes, Minnesota, USA (Fig. 2) and is characterized in detail by Lambrecht et al. (2018). It is a  
64 eutrophic lake with abundant iron in the anoxic water column and sediments. This lake has been  
65 meromictic since 1925, and its long-term water column stratification results in strong  
66 physicochemical gradients of sunlight, oxygen, and iron (Lambrecht et al., 2018). Brownie Lake  
67 currently has a maximum depth of 14 m and a surface area of 5 ha. The conductivity, dissolved  
68 O<sub>2</sub>, and temperature profiles indicate an oxic mixolimnion (0–3.5 m) and a dense anoxic

69 monimolimnion below 5 m, separated by a chemocline (4–5 m), which often coincides with the  
70 oxycline (Fig. 2). 16S rRNA sequencing revealed that Methanogens, primarily of the order  
71 *Methanobacteriales*, are abundant in the water column (Lambrecht et al., 2020). At 11-12 m water  
72 depth, Methanogen sequences accounted for ~31% of sequences, with a biogenic  $\delta^{13}\text{C}_{\text{CH}_4}$  signature  
73 (-64 ‰) and a higher  $\text{CH}_4$  concentration compared to nearest the sediment—pointing to active  
74 water column methanogenesis (Lambrecht et al., 2020). Aerobic methanotrophy is the dominant  
75  $\text{CH}_4$  oxidation mechanism (Lambrecht et al., 2020). We build on these previous results from this  
76 lake by incorporating organic carbon fluxes, sediment burial rates, and OCR rates, along with a  
77 reaction transport model, to evaluate the role of ME in OC cycling.

## 78 **METHODS**

79 Water column profiles of concentrations and stable carbon isotopes of  $\text{CH}_4$  ( $\delta^{13}\text{C}_{\text{CH}_4}$ ), dissolved  
80 and particulate organic carbon ( $\delta^{13}\text{C}_{\text{DOC}}$ ,  $\delta^{13}\text{C}_{\text{POC}}$ ), and dissolved inorganic carbon ( $\delta^{13}\text{C}_{\text{DIC}}$ ), along  
81 with concentration of major nutrients, anions, and cations in Brownie Lake were collected over  
82 several years (Swanner, 2022). This study utilizes previously reported  $\text{CH}_4$ , DIC, major nutrients,  
83 anions, and cations data (Lambrecht et al., 2018; Lambrecht et al., 2020) along with new data,  
84 including concentrations of POC, DOC, particulate organic nitrogen (PON), and dissolved organic  
85 nitrogen (DON), along with their isotopic compositions to quantify the Lake's OC budget. Primary  
86 productivity and external carbon loading were quantified using rapid light curves and the external  
87 chemical input model available for Brownie Lake (Supp. Section 2). A 1.5 m long sediment piston  
88 core was collected from the deep basin for  $^{210}\text{Pb}$  dating using the constant rate of supply model to  
89 quantify dry mass accumulation rates (Appleby and Oldfield, 1978; Supp. Section 3). To  
90 investigate and quantify the processes controlling the distribution of dissolved and particulate  
91 species as well as the turnover of C, S, Fe, and P in the water column, the data from the lake were

92 simulated with an existing biogeochemical reaction-transport model (Dale et al., 2009; Supp.  
93 Section 4). The August 2018 dataset was the most comprehensive for the above chemical species  
94 and was used for reaction transport modeling. OCR in the sediment column was constrained by  
95 mass-balancing the measured sediment OC burial and modeled OC rain rate to the lake floor.

## 96 **RESULTS AND DISCUSSION**

97 The POC loading was 107–264 mmol C m<sup>-2</sup> d<sup>-1</sup> with contribution from primary productivity (100–  
98 250 mmol C m<sup>-2</sup> d<sup>-1</sup>) and runoff (7–14 mmol C m<sup>-2</sup> d<sup>-1</sup>; Supp. Section 2). Water column profiles  
99 showed a subsurface chlorophyll maximum at 3.5 m, characteristic of ferruginous meromictic  
100 lakes, along with a positive spike in POC, PON, and DOC, and a low C:N ratio (TOC/TN  
101 mass/mass) (Fig. 3; Supp. Section 1), indicating a predominantly autochthonous labile OC flux  
102 sinking to the deeper water column. The increase in ammonium and DIC concentrations with depth  
103 in the monimolimnion (below 4–5 m) indicate OCR. Increasing  $\delta^{13}\text{C}_{\text{DIC}}$  values (–11.53‰ at 3.5 m  
104 to –2.92 ‰ at 13 m depth) and CH<sub>4</sub> concentrations (max 0.2 mM above 3.5 m to max 1.5 mM in  
105 monimolimnion) with depth indicate active ME below the chemocline. The DOC concentration  
106 profile below the chemocline did not show comparable variation with DIC and  
107 CH<sub>4</sub> concentrations, indicating that only a portion of the organic carbon is available for ME and  
108 the existence of a sizeable recalcitrant DOC pool. A depletion of  $^{13}\text{C}_{\text{DIC}}$ ,  $^{13}\text{C}_{\text{DOC}}$ , and  $^{13}\text{C}_{\text{POC}}$ , along  
109 with enrichment in  $^{13}\text{C}_{\text{CH}_4}$  at the chemocline, suggest strong aerobic CH<sub>4</sub> oxidation above the  
110 chemocline (Fig. 2; c.f. Lambrecht et al., 2020). CH<sub>4</sub> storage was estimated by integrating  
111 measured CH<sub>4</sub> concentration to water volume data per depth and lake surface area, yielding 25.85  
112 g C m<sup>-2</sup>, which is very high compared to other lakes with similar surface areas (Supp. Section 7).  
113 Our reaction-transport model-based simulation for POC remineralization in the water column  
114 returned a good fit to the measured chemical parameters (Fig. 3; Supp. Section 4). Results yielded

115 a total OCR rate of 67–224 mmol C m<sup>2</sup> d<sup>-1</sup> (62–85% of POC loading), of which 28–37 mmol C m<sup>2</sup>  
116 d<sup>-1</sup> (14–26% of POC loading) occurs in the water column (Table 1). 75±1% of water column OCR  
117 occurred anaerobically, of which ME accounted for 92–95% (19–27 mmol C m<sup>2</sup> d<sup>-1</sup>; Table 1).  
118 OCR via sulfate reduction and dissimilatory iron reduction in the water column was limited in  
119 comparison (1.6±0.1 mmol C m<sup>2</sup> d<sup>-1</sup>; 5–8% of anaerobic OCR). The excess ammonium observed  
120 below the redoxcline compared to the modeled result could be due to nitrogen fixation (Philippi et  
121 al., 2021) or dissimilatory nitrate reduction to ammonium instead of N<sub>2</sub> (Michiels et al., 2017), two  
122 processes that have been observed in ferruginous lakes. While iron reduction could be  
123 thermodynamically favorable in the Brownie Lake water column, our results suggest a minimal  
124 role of dissimilatory iron reduction in OCR. Previous studies have shown that methanogens can  
125 outcompete iron reducers during OCR under non-carbon-limited settings due to the transformation  
126 of iron oxide minerals to stable forms or due to surface passivation of reactive iron oxides by Fe(II)  
127 (Friese et al., 2021; Gadol et al., 2022). The presence of iron oxide minerals in sediments (Supp.  
128 Section 6) indicates they are escaping water column remineralization processes, thereby favoring  
129 ME as the dominant mode of OCR in monimolimnion.

130 Modeled OC rain rate at the lake floor (79–226 mmol C m<sup>2</sup> d<sup>-1</sup>; 74–86% of total OC load)  
131 combined with measured OC burial (40.38 mmol C m<sup>-2</sup> d<sup>-1</sup> for top 11 cm and 37.78 mmol C m<sup>-2</sup>  
132 d<sup>-1</sup> for top 70 cm) points to 38–186 mmol C m<sup>-2</sup> d<sup>-1</sup> OCR in the sediment column and that only 18–  
133 51% of the OC rain is being buried. This OCR in the ferruginous sediment column would occur  
134 via dissimilatory iron reduction and ME (Bray et al., 2017). The modeled CH<sub>4</sub> flux from the  
135 sediment column towards the lake floor (1–3 mmol CH<sub>4</sub> m<sup>2</sup> d<sup>-1</sup>) implies a minimum methanogenic  
136 OCR of 2–6 mmol C m<sup>2</sup> d<sup>-1</sup> in the sediment column. We emphasize that this is the minimum ME  
137 estimate in the sediment since a portion of CH<sub>4</sub> produced in the sediment column could be

138 consumed by Fe-dependent AOM (Supp. Section 6). The low C:N ratio and low  $\delta^{13}\text{C}_{\text{org}}$  in the  
139 benthic nepheloid layer and sediment column, along with highly enriched  $\delta^{13}\text{C}_{\text{DIC}}$  for the top 40  
140 cm of measured porewater (Supp. Fig. 5), support the interpretation of active ME in shallow  
141 sediments. Taken together, ME accounted for at least 13–42% of anaerobic OCR in the ferruginous  
142 water column and sediments.

143 *Fig. 2; Table 1*

144 Evaluation of the role of ME in ancient ferruginous oceans will provide critical insights into the  
145 carbon cycling dynamics during the Precambrian and Earth's early climate evolution. Comparison  
146 of Brownie Lake's  $\text{CH}_4$ -carbon budget with other published datasets (Fig 3) suggests that the  
147 reported OC loading to  $\text{CH}_4$  conversion efficiency in anaerobic OCR under ferruginous settings  
148 averages 36% (18–59%), with Brownie Lake at the lower end of this range but still significantly  
149 higher than modern (oxic) oceans with 0.1% efficiency. OC burial and ME can impact Earth's  
150 early oxygenation in different ways—the former removes a reductant from the Earth's surface. The  
151 latter injects a reductant into the atmosphere, inducing greenhouse warming and contributing to  
152 top-down oxygenation via hydrogen escape to space (after  $\text{CH}_4$  photolysis) from the atmosphere  
153 (Catling et al., 2001). A dominant role of  $\text{CH}_4$  in the Precambrian climate has been widely  
154 proposed in the past three decades of literature (Catling and Zahnle, 2020). A few recent studies  
155 have argued for a limited role of  $\text{CH}_4$  in the Precambrian climate (Laakso and Schrag, 2019), citing  
156 a case example of high OC burial in ferruginous Lake Matano (Kuntz et al., 2015). Our results  
157 from Brownie Lake rather support a lower proportion of OC being buried in sediments under  
158 ferruginous settings and a dominant role of  $\text{CH}_4$  in the Archean carbon cycle (Thompson et al.,  
159 2019).

160 In the modern oceans, an average net primary productivity (NPP) of 50 Gt C yr<sup>-1</sup> results in 2 Gt C  
161 yr<sup>-1</sup> deposited in the seafloor, leading to ~0.05 Gt CH<sub>4</sub> yr<sup>-1</sup> ME via OCR (Akam et al., 2023), with  
162 an OC loading to CH<sub>4</sub> generation efficiency of 0.1%. Estimates for a late Archean setting range  
163 from 0.1% to 14% of modern NPP (Ward et al., 2019; Farr et al., 2023). An OC to CH<sub>4</sub> conversion  
164 efficiency of 36% would yield 2 to 210 (extended range of 1–344 considering 18–59% efficiency)  
165 Tmol CH<sub>4</sub> yr<sup>-1</sup> or 0.4–56 times modern annual marine ME rates (Supp. Section 8). Interestingly,  
166 Ozaki and Reinhard (2018) modeled an increased efficiency of CH<sub>4</sub> cycling under a hybrid  
167 ecosystem composed of H<sub>2</sub> and Fe<sup>2+</sup>-based anoxygenic photoautotrophy. Under low sulfate and  
168 low oxygen surface waters, this CH<sub>4</sub> would have entered the atmosphere easily, compared to >90%  
169 CH<sub>4</sub> being oxidized in the modern ocean with high sulfate and oxygen (Habicht et al., 2002).

170 Photochemical models predict that the lifetime of CH<sub>4</sub> in a low-O<sub>2</sub> atmosphere is 5,000–10,000  
171 years, as opposed to ~12 years today (Catling et al., 2001). Hence, anoxic Archean atmosphere  
172 could hold 1000s of parts per million by volume of CH<sub>4</sub>, provided a sufficient CH<sub>4</sub> supply, and  
173 even a smaller CH<sub>4</sub> flux over time (e.g., 2 Tmol CH<sub>4</sub> yr<sup>-1</sup> in Archean over thousands of years) can  
174 increase CH<sub>4</sub>-induced warming on early Earth. The gradual oxidation of Earth's surface would  
175 have limited ME to anoxic deep water and sediment column as well as a reduced lifetime of CH<sub>4</sub>  
176 in the atmosphere, limiting their warming control (Olson et al., 2016). Our results point to efficient  
177 ME in ancient ferruginous oceans supporting the climate models, suggesting a significant climate  
178 warming by CH<sub>4</sub> (Haqq-Misra et al., 2008). In contrast, a lower role of CH<sub>4</sub>-warming was proposed  
179 recently, primarily based on high OC burial rates in Lake Matano (Laakso and Schrag, 2019). Our  
180 results hence emphasize that the ferruginous oceans were conducive to high rates of ME, and thus,  
181 the availability of OC loading and oxidants like sulfate and oxygen would have been the key  
182 controlling factor determining the marine CH<sub>4</sub> fluxes to Earth's early atmosphere. The amount of



183 CH<sub>4</sub> produced would have closely followed the trend of NPP in the early Earth until the advent of  
184 surface oxygenation, at which time the efficiency of ME with reference to NPP would have  
185 decreased gradually leading to current efficiency of 0.1%. Lastly, we highlight the need to  
186 constrain the OC budget and the relative efficiency of ME from additional ferruginous systems to  
187 improve on our understanding of the biogeochemical significance of iron-rich systems at present  
188 and in the geological past.

189 *Fig. 3*

## 190 **CONCLUSION**

191 We document that ME is a dominant OCR process in ferruginous meromictic Brownie Lake,  
192 accounting for  $\geq 90$  and  $>9 \pm 7\%$  of the anaerobic OCR in the water column and sediments,  
193 respectively, and an overall POC loading to CH<sub>4</sub> conversion efficiency of  $\geq 18 \pm 7\%$  in the anoxic  
194 zone of Brownie Lake. Our results, combined with available results from other ferruginous  
195 systems, point to a very high efficiency ( $36 \pm 21\%$ ) of POC to CH<sub>4</sub> conversion in these systems,  
196 compared to 0.1% in the modern sulfatic and oxic ocean. Hence, we conclude that even under a  
197 low primary productivity scenario, Archean oceans would have produced sufficient CH<sub>4</sub>, in  
198 agreement with climate models suggesting CH<sub>4</sub>-induced greenhouse warming in early Earth.

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312

313 FIGURE CAPTIONS

314 Figure 1: Atmospheric A) CH<sub>4</sub> concentration B) and O<sub>2</sub> concentration (Fakhraee et al., 2019). C)  
315 Spatially predominant ocean-redox conditions over Earth's history (Poulton, 2021). PAL = present  
316 atmospheric level.

317 Figure 2: A) Study site. B) Temperature, conductivity, and dissolved O<sub>2</sub> and Fe profiles of the  
318 water column. C) isotopic composition; D) measured and modeled water column concentration  
319 profiles; E) Crossplot of TOC:N and  $\delta^{13}\text{C}_{\text{org}}$  suggestive of labile carbon availability (Meyers,  
320 1994). F) Overall carbon budget schematic.

321 Table 1: Summary of modeled carbon cycling parameters in Brownie Lake.

322 Figure 3: Comparison of A) anaerobic organic carbon remineralization (OCR) rates for Archean  
323 ferruginous settings and analogs, B) efficiency of methanogenesis (ME) in OCR, and C) ME  
324 efficiency to OC loading (data from Supp. Table 7) D) simplified schematic of C-CH<sub>4</sub> cycling in  
325 ancient ferruginous oceans.

326 **[Please include this text at the end of your paper if you are including an item in the Supplemental**

327 <sup>1</sup>Supplemental Material. *[Please provide a brief description of your material.]*

328 Supplementary file contains additional information on lake setting, water chemistry measurements,  
329 Organic carbon loading calculation, sediment-coring, age-dating, mass accumulation rates, reaction-  
330 transport model description,  $\delta^{13}\text{C}$  and C:N composition for organic carbon source identification;  
331 oxidation state of iron in sediments; methane storage comparison with lakes of similar size, and overall  
332 carbon cycling schematic.