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Autonomous observations of biogenic N_2 in the Eastern Tropical North Pacific using profiling floats equipped with gas tension devices

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Oxygen Deficient Zones (ODZs) of the world's oceans represent a relatively small fraction of the ocean by volume (<0.05% for suboxic and <5% for hypoxic) yet are receiving increased attention by experimentalists and modelers due to their importance in ocean nutrient cycling and predicted susceptibility to expansion and/or contraction forced by global warming. Conventional methods to study these biogeochemically important regions of the ocean have relied on well-developed but still relatively high cost and labor-intensive shipboard methods that include mass-spectrometric analysis of nitrogen-to-argon ratios (N_2/Ar) and nutrient stoichiometry (relative abundance of nitrate, nitrite, and phosphate). Experimental studies of denitrification rates and processes typically involve either *in-situ* or *in-vitro* incubations using isotopically labeled nutrients. Over the last several years we have been developing a Gas Tension Device (GTD) to study ODZ denitrification including deployment in the largest ODZ, the Eastern Tropical North Pacific (ETNP). The GTD measures total dissolved gas pressure from which dissolved N_2 concentration is calculated. Data from two cruises passing through the core of the ETNP near 17 °N in late 2020 and 2021 are presented, with additional comparisons at 12 °N for GTDs mounted on a rosette/CTD as well as modified profiling Argo-style floats. Gas tension was measured on the float with an accuracy of < 0.1% and relatively low precision (< 0.12%) when shallow ($P < 200$ dbar) and high precision (< 0.03%) when deep ($P > 300$ dbar). We discriminate biologically produced N_2 (ie., denitrification) from N_2 in excess of saturation due to physical processes (e.g., mixing) using a new tracer – 'preformed excess- N_2 '. We used inert dissolved argon (Ar) to help test the assumption that preformed excess- N_2 is indeed conservative. We used the shipboard measurements to quantify preformed excess- N_2 by cross-calibrating the gas tension method to the nutrient-deficit method. At 17 °N preformed excess- N_2 decreased from approximately 28 to 12 $\mu\text{mol/kg}$ over $\sigma_\theta = 24\text{--}27\text{ kg/m}^3$ with a resulting precision of $\pm 1\text{ }\mu\text{mol N}_2/\text{kg}$; at 12 °N values were similar except in the potential density range of $25.7 < \sigma_\theta < 26.3$ where they were lower by 1 $\mu\text{mol N}_2/\text{kg}$ due likely to being composed of different source waters. We then applied these results to gas tension and O_2 (< 3 $\mu\text{mol O}_2/\text{kg}$) profiles measured by the nearby float to obtain

the first autonomous biogenic N_2 profile in the open ocean with an RMSE of $\pm 0.78 \mu M N_2$, or $\pm 19\%$. We also assessed the potential of the method to measure denitrification rates directly from the accumulation of biogenic N_2 during the float drifts between profiling. The results suggest biogenic N_2 rates of $\pm 20 nM N_2/day$ could be detected over >16 days (positive rates would indicate denitrification processes whereas negative rates would indicate predominantly dilution by mixing). These new observations demonstrate the potential of the gas tension method to determine biogenic N_2 accurately and precisely in future studies of ODZs.

KEYWORDS

denitrification, biogenic nitrogen, oceanographic floats, gas tension device, argon

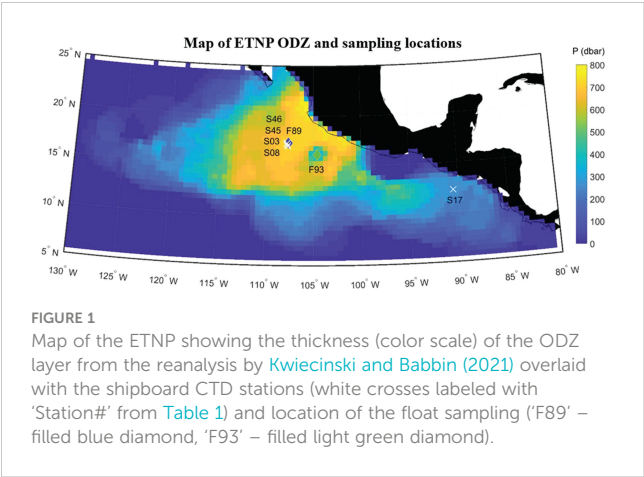
1 Introduction

ODZ's are important in ocean nutrient cycling, typically are in close proximity to large fisheries, and are susceptible to change linked to global warming (Engel et al., 2022). Occurring primarily on the eastern edge of ocean basins, ODZ's are a recognized source of the potent greenhouse gas nitrous-oxide (N_2O) to the atmosphere and an oceanic sink of nitrogen-based nutrients (nitrate and nitrite) as a consequence of microbial processes (denitrification and anammox) requiring the near absence of O_2 (Codispoti et al., 2001). Nitrogen cycling activity levels of ODZ microbial communities thus overall have a positive response to the supply of organic matter and a negative response to oxygen despite varying individual responses. The tolerance of anaerobic microbes to trace levels (nM) of dissolved oxygen (O_2) has been the focus of many prior laboratory and field studies due to its importance in modeling and prediction. The simplest model (exponential-type inhibition kinetics) assumes a threshold of O_2 at or below a few hundred nM O_2 ($\sim 0.1\%$ of background outside the ODZ) above which denitrification processes resulting in the accumulation of N_2 gas in the mesopelagic stop (Dalsgaard et al., 2014). There is considerable uncertainty associated with this threshold O_2 concentration and even the applicability of such a simplified model since microbes can adapt and respond to O_2 forcing, so a deeper knowledge of sub-threshold O_2 responses is required (Bristow et al., 2016; Garcia-Robledo et al., 2017; Zakem and Follows, 2017; Penn et al., 2019; Berg et al., 2022). Identification of ODZ waters is more commonly based on denitrification activity, rather than oxygen concentration, through removal of NO_3^- (N deficit), the presence of a secondary NO_2^- maximum (Banse et al., 2017), or the appearance of biogenic N_2 produced by denitrification. These waters have also been termed 'functionally anoxic' having O_2 concentrations of <10 's nM (Revsbech et al., 2009; Tiano et al., 2014).

From an experimental viewpoint focused on dissolved gas cycling, the study of ODZ biogeochemistry requires measuring trace level changes in O_2 and N_2 correlated in time and space with likely forcings. Prior studies of open ocean ODZs have

measured water column biogenic- N_2 and denitrification rates (ie., the production rate of biogenic- N_2) using a combination of mass-spectrometric or gas-chromatographic analysis of dissolved N_2 in discrete water samples (Chang et al., 2010; Chang et al., 2012; Fuchsman et al., 2018) or *in vitro* incubations using automated nutrient analysis for elemental stoichiometry and spiking with ^{15}N -labeled substrates to distinguish denitrification and anaerobic ammonium oxidation (anammox) processes (e.g., Babbitt et al., 2014; Dalsgaard et al., 2014). Although these types of measurements are generally highly precise and accurate, they are both costly due to the requirement for specialized collection, storage, and analysis of seawater samples at a shoreside laboratory and limited in time and space. Consequently, large research gaps have developed because of the limited types, quantity, and distribution of oceanographic data provided by existing methods. A few prior studies have investigated spatial variability in ocean denitrification rates and found evidence that eddies can be significant hotspots for denitrification (Callbeck et al., 2017; Altabet and Bourbonnais, 2019) but more statistics are needed to better understand how important eddies really are in ODZ biogeochemistry. Although short term variability in denitrification rates can be measured during a several week-long cruise, investigation of seasonal cycles in N_2 using sparse existing data sets collected using conventional approaches is far out of reach. New *in situ* methods are needed that can be deployed on moorings, floats, and gliders to allow measurement of denitrification at relevant spatial and temporal scales.

Over the last several years we have been developing new approaches for studying denitrification in the Eastern Tropical North Pacific (ETNP, see Figure 1), the largest of the three major ODZs. Our approach was motivated by the need to autonomously observe long-term changes in dissolved gas cycling within ODZs that have been observed (Horak et al., 2016; Ito et al., 2017), or are predicted (e.g., Deutsch et al., 2014), to occur due to the impacts of global warming on ODZ biogeochemistry. We use a membrane-based dissolved gas sensor known as a Gas Tension Device (GTD), which was originally developed (McNeil et al., 1995) for studying air-sea gas exchange rates and processes in oxic waters, to measure total dissolved air pressure. With the intent of providing enough



detail so other researchers can employ these methods for themselves, and likely improve upon them, we describe the dissolved gas sensors and adaptations for the sampling platforms we used, and the assumptions and data processing steps involved in calculating N_2 concentrations, biogenic- N_2 , and even denitrification rates. A limitation of our approach, based on measuring N_2 in situ, is that we cannot distinguish different processes that produce N_2 gas as other methods, such as labeled-incubations with GC-MS analysis, can and so we cannot distinguish denitrification and annamox processes (unless perhaps these processes are known to occur in different regions of the water column or under different environmental conditions).

We present and discuss vertical profiles of GTD-derived biogenic- N_2 made using data collected in the core waters of the ETNP near 17°N. The data highlights that real-time biogenic- N_2 can now be measured at sea by ships and BGC-Argo-style profiling floats and therefore this new method offers the possibility to map important oceanographic regions (e.g., ODZ boundary) or study oceanographic features of interest (e.g., eddies) within the ODZ. These data also motivate the need to better understand subtler long-term changes in ODZ biogeochemistry through a combination of combined observational and modeling campaigns.

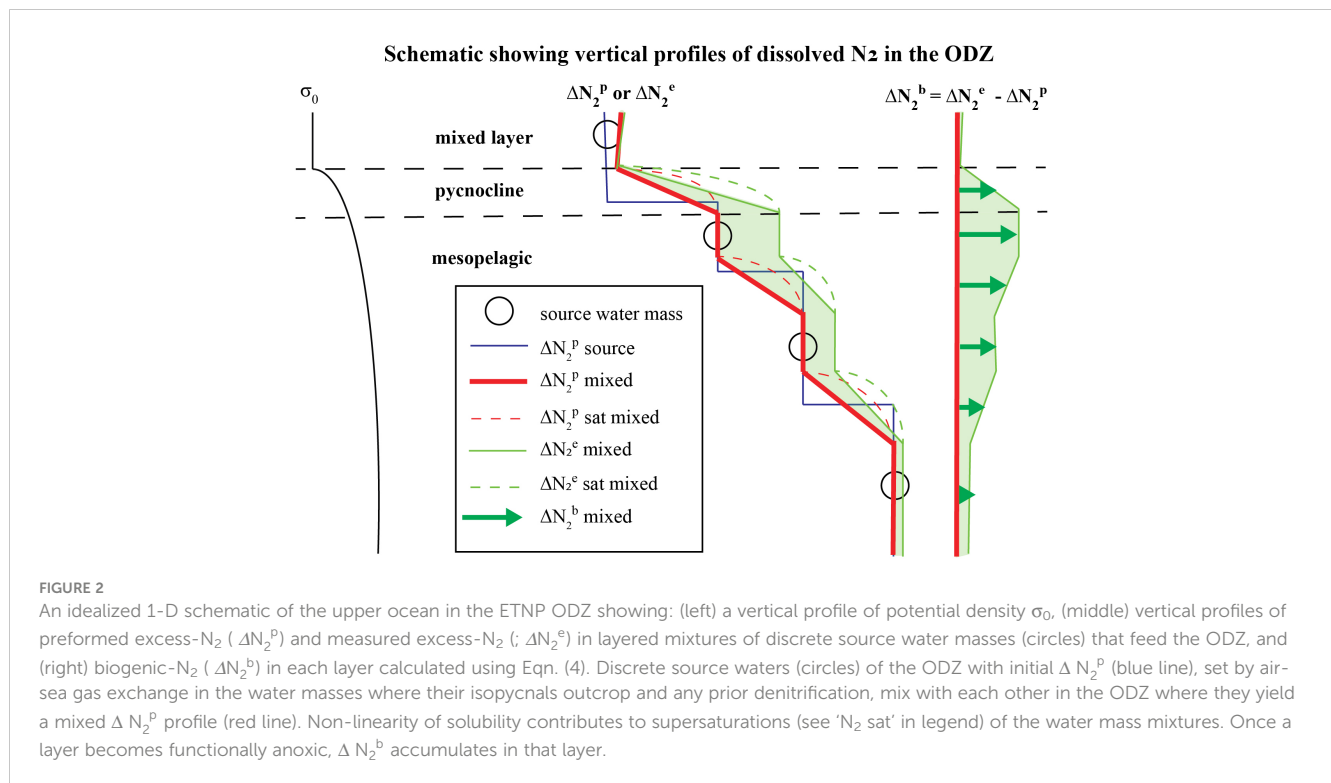
2 Materials and methods

2.1 Overview

To better study ODZ biogeochemistry our sampling objectives were: 1) to measure biogenic- N_2 profiles that result from microbial decomposition of nitrogen-based nutrients (nitrate, nitrite, ammonia) in anoxic waters at approximately the same location in the ocean (i.e., **Eulerian approach**), and 2) to measure the time rate of change of biogenic- N_2 , or the denitrification rate, within the same subsurface water mass moving with the mean flow (i.e., **Lagrangian approach**). The Eulerian approach uses a GTD, mounted either on the ship's CTD or a profiling float, held at various depths or isopycnals for sufficient time (< 1 day) to equilibrate the sensors and obtain a vertical profile. The Lagrangian approach can be used to estimate *in situ* denitrification rates by making repeat measurements on a freely drifting float at the same isopycnal to ensure sensors stay in the same water mass for sufficient time (>10 days) to detect a change in biogenic N_2 . In the ODZ core region where horizontal currents are weak, vertical gradients in biogenic- N_2 are much larger than horizontal gradients and a simple 1-D budget model for N_2 (see Figure 2) can be used to interpret both sampling approaches. A complication when applying such a simplistic N_2 budget is the need to separate biogenic and non-biogenic components (e.g., physical processes including mixing). This complication holds for any technique that budgets N_2 , regardless of the method used to measure N_2 in seawater. For example, Chang et al. (2012) encountered this same issue interpreting N_2 measurements made in the ETNP ODZ using mass-spectrometric measurements of N_2 and solved the problem by referencing their N_2 measurements made inside the ODZ region to background measurements at the same density made outside the ODZ. Their technique further benefited from using measured N_2/Ar ratios (the raw measurements of the mass spectrometer) since inert Ar provides essentially an intrinsic reference for abiotic- N_2 . We adopt and extend this same approach by considering the different source waters to the ODZ core region as importing their own preformed excess- N_2 to the ODZ and mixing vertically and horizontally along their pathways to the core of the ODZ. Rather

TABLE 1 Details on time, location, and depth range of GTD sampling by ship and float.

Cruise or Float ID [Station#, Event#]	Date [dd-mm-yyyy]	Time [GMT]	Latitude [°N]	Longitude [°W]	Depth range [dbar]
SR2011[S03, E12]	22-Dec-2020	17:37	17.289	107.775	29–859
SR2011[S08, E03]	29-Dec-2020	20:03	17.079	107.739	28–1009
SR2114[S45, E04]	14-Jan-2022	17:57	16.980	107.700	84–1010
SR2114[S46, E01]	15-Jan-2022	01:42	17.202	107.815	59–858
SR2114[S17, E05]	29-Dec-2021	01:39	12.217	90.543	83–859
F89 Drift:	01-Jan-2021	05:12	17.288	107.634	15–706 (242 ± 4)
F93	18-Apr-2022	15:06	16.119	104.613	17–1019



than having to measure dissolved gas levels outside the ODZ as Chang et al. did, we calibrate our biogenic N₂ to a concurrent but independent measurement based on nutrient stoichiometry to obtain vertical profiles of preformed excess-N₂ in the water mixtures at the ODZ core. The validity of the 1-D model is assessed using conservative mixing lines for independently measured argon.

We used the *in-situ* gas tension method as an alternative approach to measuring N₂ in seawater since it can be measured autonomously, is relatively easy, reliable, and inexpensive compared to conventional techniques. In the anoxic seawater layer found at the core of the ODZ, the gas tension method has increased accuracy due to the elimination of measurement errors associated with dissolved O₂. One prior study used this same advantage to study aerobic/anoxic N₂ cycling in the Baltic Sea (Loeffler et al., 2011). A disadvantage of the gas tension method alone is that argon levels are usually assumed, and therefore no compensatory information on abiotic processes is measured. We, however, used an advanced isotope dilution mass-spectrometric technique developed to measure absolute Ar concentrations to high accuracy. The Ar samples were drawn from the same rosette CTD used for gas tension sampling. The drawbacks of the isotope dilution technique method are high per-sample analysis cost and there are only a few laboratories that can do it, but this disadvantage is somewhat offset by the fact that argon is expected to vary predictably and smoothly (Ito et al., 2011) by conservative mixing away from boundary sources so only targeted sampling of ODZ source waters was required. We used independent shipboard measurements of biogenic N₂ from nutrient stoichiometry with concurrent gas tension derived estimates of biogenic N₂ to determine N₂ variability due solely to physical processes. We used the Ar measured by isotope dilution techniques to check our

assumption that this 'preformed' excess-N₂ followed expected conservative mixing curves. We applied the preformed excess-N₂ profiles measured in the ODZ core during the cruise to profiles of gas tension measured after the cruise by autonomous profiling floats deployed in the same vicinity to obtain the first autonomous biogenic N₂ profiles in an ODZ. Derived denitrification rates require a quantifiable increase of biogenic N₂ within the same water mass over a known period of time; this period will be longer for regions or isopycnals with lower microbial activity.

2.2 Field campaigns

Biogenic N₂ was measured in the ETNP near 17 °N (see Figure 1) during two cruises on the RV *Sally Ride* (CruiseID = SR2011 during 16-Dec-2020 to 06-Jan-2021m; CruiseID = SR2114 during 21-Dec-2021 to 21-Jan-2022). A GTD was mounted on the rosette CTD to measure vertical profiles of gas tension, along with O₂ (by SBE-43), temperature (T) and salinity (S). Water samples were taken for nutrients (nitrate, nitrite and phosphate). A Switchable Trace amount OX-ygen (STOX) sensor (Revsbech et al., 2011) was used to reference the SBE-43 O₂ data to anoxia (± 50 nM) in the ODZ core waters. Autonomous floats, also with GTD sensors, were deployed. Table 1 lists sampling details for the data presented here.

2.3 Gas tension devices

New versions of Gas Tension Devices (GTDs manufactured by Pro-Oceanus Systems, Inc., Bridgewater, NS, Canada) were used to

measure gas tension on the floats and the ship's CTD in this study. Unlike older versions of the GTD (McNeil et al., 1995; McNeil et al., 2006), the new sensor has only a weak hydrostatic response, due primarily to the use of the relatively incompressible Teflon AF 2400 gas permeable membrane introduced by Reed et al. (2018). The new GTD has the same external mechanical design as the profiling GTD developed for use on the ship's CTD (McNeil et al., 2018, see figures on pages 88 and 96) but was modified at APL/UW internally to: 1) improve the electronics, and 2) allow the sensor's pressure sensor to be easily replaced at sea even in the event of a flooded membrane. The novelty of the profiling GTD design lies in the use of a miniature pressure sensor (type Micro Electrical Mechanical System- or MEMS) with a very small (approximately 40 μ L) gas sensing volume. Since this pressure sensing volume contributes to the total sample volume behind the membrane that must exchange through the membrane, the response time of the new GTD is small: In separate laboratory tests conducted using similar high-flow pumps and plenums the characteristic (e-folding) response times

were approximately 1 minute at 20°C and 3 minutes at 5 °C. The response time of any GTD varies with seawater temperature because of the temperature dependence of gas diffusivity and solubility in the GTD's membrane. The new GTD was mounted onto a rosette 911 SBE CTD system that had a serial RS-232 ('Data Uplink') capability for real-time data and a spare SBE 5T pump running at approximately 50% of maximum capacity (ie., 0.5 A at 12 V, or 6 W); the pump flushed the GTD's membrane using a plenum attachment as described in McNeil et al. (2018). The GTD had an operating-depth rating of 1000 m.

The time required to equilibrate the new GTD to an acceptable level (e.g., <1 mbar of absolute) when profiled through strong temperature and dissolved gas gradients is not simply a function of the isothermal response time of the sensor found in the laboratory, it also depends on the prior history of exposure of the sensor during the profile and geophysical noise such as the passage of internal waves. Since GTD measurements took many hours on station, it was necessary to minimize the wait time at each sampling

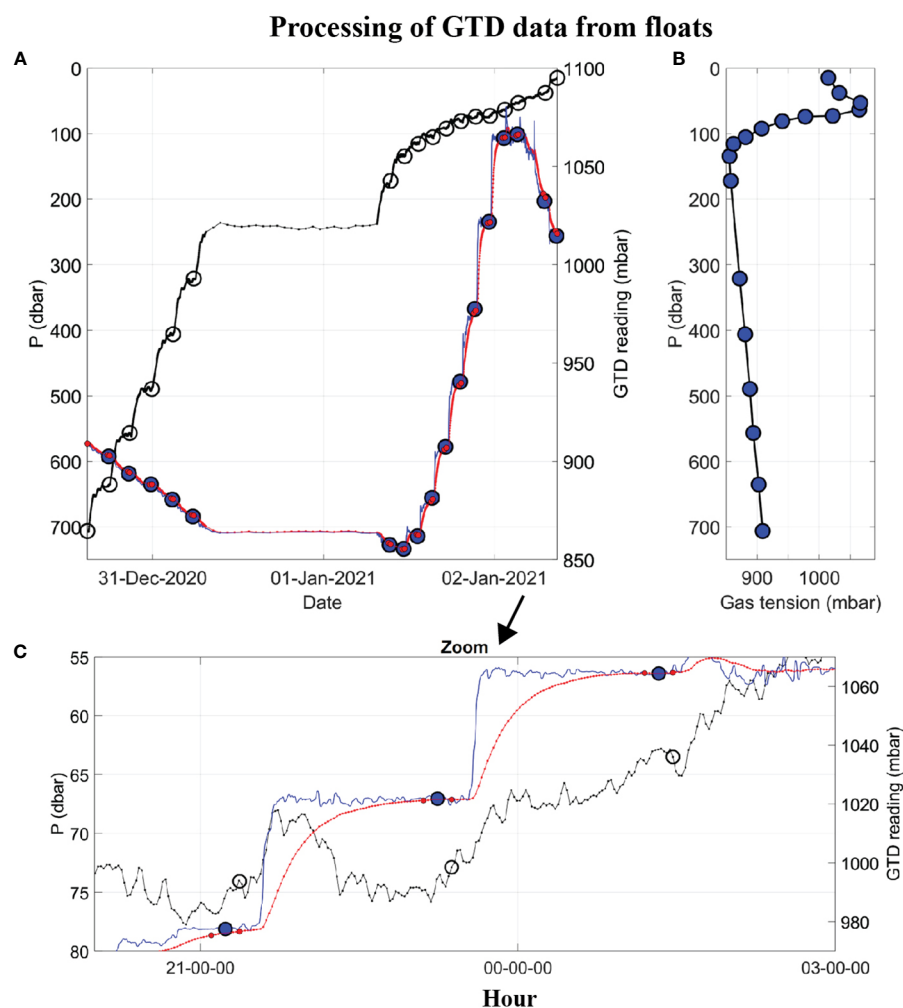


FIGURE 3

Example of raw and processed float data, from F89, showing: (A) Time series of hydrostatic pressure (left axis, black line) with time marks showing when the GTD is fully equilibrated (left axis, black circles), raw GTD reading (right axis, red line), and deconvolved GTD data (right axis, blue line), with time marks (small red dots) over which the mean equilibrated gas tension is calculated (large blue dot); (B) Vertical profile of gas tension (large blue dots); and (C) Zoom-in of the upper plot for three equilibrations, starting at 20:00 on 01-Jan-2021.

depth and provide a straightforward procedure for CTD operators to follow. At-sea trials of the new GTD on the ship's CTD were performed at the start of the cruise to determine the minimum wait-time required to equilibrate the GTD in temperature and gas before moving the CTD to the next sampling-depth. A wait-time of 10 minutes at each sampling-depth found to be sufficient except in the thermocline where the wait-time was increased to 15 minutes. Equilibrated data were chosen based on averages over the last few minutes of data prior to moving the CTD to the next depth.

2.4 Autonomous floats

As part of a larger distributed array designed to capture relevant temporal and spatial variability in denitrification rates and processes in the ETNP, a total of 10 modified Argo-style floats were deployed on the two cruises and data from two of those floats deployed in the ODZ core region are examined here. Floats (model ALTO) manufactured by Marine Robotics Systems (MRV) were modified in collaboration with MRV to carry one or more GTD sensors and two Seabird-63 optode oxygen sensors. The float's new GTD and optode membranes were flushed using a low flow rate SBE-5M pump to reduce power consumption, the trade-off being a significantly longer equilibration time of approximately 20 minutes for the GTD (i.e., approximately 20 times slower than the fastest laboratory equilibrations). One of the two optodes measured the water; the other measured in a saturated sodium sulfite solution and

acted as an anoxic reference solution subject to the same temperature and pressure as the water optode. The standard optode calibration equation is not accurate at concentrations below a few $\mu\text{mol/kg}$ as it does not account for temperature and pressure dependences of the optode response at zero oxygen. Here, oxygen was computed both from the difference between the water and reference optodes corrected for a linear pressure trend and by reformulating the calibration equation with the appropriate pressure and temperature variations. The ODZ core was clearly apparent as a minimum in oxygen concentration several hundred meters thick, with values of a few $\mu\text{mol/kg}$ or less and a linear variation of oxygen with depth of a few hundred nmol/kg . The pressure coefficient was adjusted to make oxygen constant in this region; an offset was adjusted to center this on zero. The STOX sensor on the ship's CTD, when available, was used to check the optode derived profiles. A detailed description is beyond the scope of this paper and will be reported elsewhere.

The float mission was modified from that of a standard Argo float. After an initial dive to 800–1000 m and a few hours to allow the sensors to settle, the float profiled upward in a staircase profile (Figure 3), moving upward 50–100 m, then settling at a nearly fixed depth and isopycnal for 1–2 hours to allow the GTD to equilibrate, then moving upward again. The float then drifted for up to 10 days in the OMZ just below the oxycline, about 200 m, before continuing the stair-step profile upward, surfacing, communicating, and repeating the cycle. Occasional glitches resulted in the float profiling to as deep as 1500 m with no apparent issues.

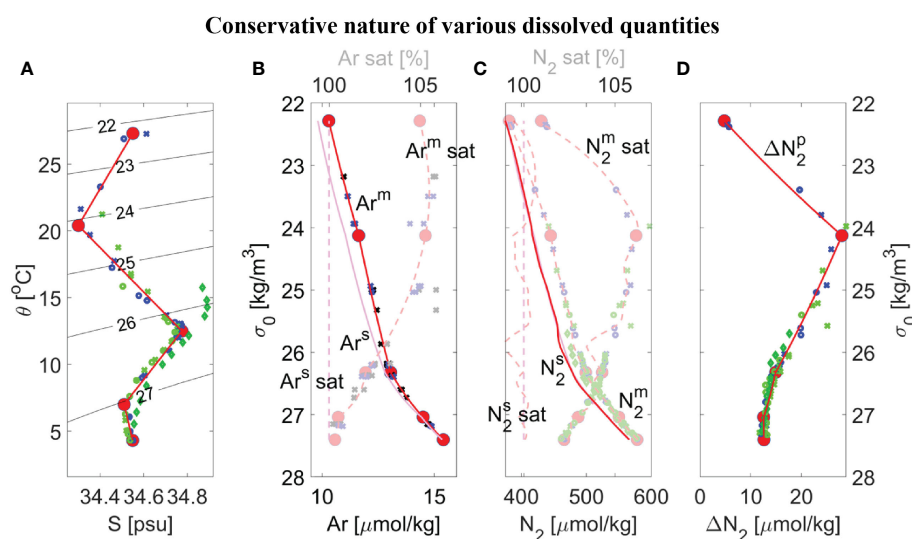


FIGURE 4

Overview of dissolved gas measurements made during duplicate casts (#1 as small circles, #2 as small crosses, and consistent with Legend shown in Figure 5B) and modeling results (red lines) for either conservative properties (opaque markers and lines) or non-conservative properties (semi-transparent markers and lines) within the ETNP ODZ core near 17°N from SR2011 (blue) or SR2114 (light green) and 12°N from SR2114 (dark green), showing: (A) Potential temperature θ versus salinity S with contours of potential density σ_0 ; (B) Dissolved argon concentration Ar^m that is either measured (blue/green symbols) or is a best-fit or mixing-model result (red solid lines), equilibrium solubility Ar^s (magenta lines), and corresponding saturation level Ar^{sat} (dashed lines) versus potential density σ_0 ; (C) Similar to panel B but GTD-measured nitrogen N_2^m and best fit (semi-transparent red line) using a smoothing spline with a residual rms of $\pm 0.15\%$ to maintain consistency with errors in solubility coefficients; (D) Preformed excess nitrogen ΔN_2^P versus σ_0 calculated using Eqn. (8). Five distinct water masses (large red dots) were chosen from panel A for the mixing model and all properties were determined by interpolating on potential density to the data. For comparison purposes, corresponding $\theta/S/\Delta \text{N}_2^P$ at 12°N from SR2114 were added to panels (A, D) (dark green diamonds) and additional argon data at $14\text{--}15.5^\circ\text{N}$ from SR2011 were added to panel (B) (black markers).

Data from two floats are reported here. The first float (F89) was deployed in the ODZ core close to the location of four CTD stations (S46, S45, S03, S08) occupied during our two cruises (see Figure 1). Only shipboard data, not all concurrent, were used to calibrate our new method for determining biogenic N_2 . After calibrations, our new method was then applied to the float data to provide the first test of the new method when applied to autonomously collected data delivered by satellite. Since the first float and CTD data were collected at the same location within the ODZ, the comparison is designed to provide information on the accumulated errors when applying our new method to float data. To demonstrate our overall goal of a long-term distributed float-based sensing observational system, we applied our new method to data from a second float (F93) chosen to be nearby and long after the shipboard calibrations work.

2.5 Dissolved N_2 from gas tension

Definitions - Dissolved N_2 partial pressure (pN_2) was calculated by the gas tension method using:

$$pN_2 = P_T - pO_2 - pH_2O - pTrace \quad \text{Eqn. (1)}$$

where P_T is the total gas tension, and pO_2 and pH_2O are the partial pressure of oxygen and water vapor, respectively. The trace gas partial pressure ($pTrace$) is mostly argon with a small contribution from pCO_2 , is relatively small, and can be accounted for in the calculations with little error (using Table 1 from Reed et al., 2018 we expect $pTrace = 9.6 - 10.9 (\pm 0.9)$ mbar for $P = 0 - 400$ dbar, which is equivalent to an uncertainty of $\pm 16\%$ in an assumed biogenic N_2 signal of 10 mbar). It should be possible to reduce uncertainties associated with pCO_2 further (we did not measure it) by estimating it from either Apparent Oxygen Utilization (AOU) or climatological empirical relationships for carbonate system parameters within the ODZ (i.e., TALK and DIC versus temperature and salinity) along with the standard seawater carbonate system calculators (e.g., CO2SYS program). We have not attempted to include any corrections here since our initial focus was understanding large scale variability in pN_2 associated with variability in actual preformed N_2 . We refer to the calculated dissolved nitrogen concentration using Henry's Law and Eqn. (1) as the '**measured**' nitrogen (N_2^m):

$$N_2^m = (P_T - pO_2 - pH_2O - pTrace) \times S_{N_2}^{(P,T,S)} \quad \text{Eqn. (2)}$$

where gas specific Henry's Law solubility coefficients (S_H) are known functions of water temperature (T), salinity (S), and hydrostatic pressure (P). Unique to the ODZ is a functionally anoxic core layer where $pO_2 = 0$. Outside of this anoxic layer small changes in O_2 are accounted for in Eqn. (2) with little error in determined N_2 by referencing the O_2 sensor data from the ship's CTD to a precise zero-point using a STOX sensor (more details below - see also Supplement in Tiano et al., 2014). We refer to nitrogen '**solubility**' (N_2^s) as the value in seawater in equilibrium with one standard atmosphere of moist tropospheric air (Hamme and Emerson, 2004). All Henry's Law solubilities are

calculated using MATLAB solubility calculators (Emerson and Hamme, 2022).

We define other nitrogen quantities as well, specifically '**excess**' nitrogen (ΔN_2^e) as:

$$\Delta N_2^e = N_2^m - N_2^s \quad \text{Eqn. (3)}$$

and '**biogenic**' nitrogen (ΔN_2^b) as:

$$\Delta N_2^b = \Delta N_2^e - \Delta N_2^p \quad \text{Eqn. (4)}$$

where the '**preformed**' excess nitrogen (ΔN_2^p) is the portion of excess N_2 associated with all processes that alter source water N_2 prior to it entering the ODZ, and include: 1) physical (abiotic) processes, such as air-sea gas exchange of the water masses during formation, and 2) denitrification in anoxic micro-environments of sinking organic matter (Bianchi et al., 2018) or continental shelf sediments (Devol, 2015). Since each source water mass to the ODZ will have its own value of ΔN_2^p with likely small interannual variability since the water formation regions are remote, we can expect mixing of the source waters within the ODZ to create stronger spatial than temporal variability in ΔN_2^p throughout the ODZ. Conservative mixing of two source waters that have different T and ΔN_2^p raises the N_2 saturation level of the mixture, a consequence of the non-linearity in the temperature dependence of Henry's Law solubility coefficients (rather than production of N_2) and must be accounted for in all mixing calculations.

Deconvolutions - To determine equilibrium values of measured quantities on the floats, which have a much slower response GTD than used on the ship CTD due to available power, we used a deconvolution model similar to Reed et al. (2018) with a suitable intrinsic (e-folding) response time scale determined by inspection of the raw data. We applied this deconvolution model to determine both raw quantities (e.g., gas tension) and derived parameters (e.g., ΔN_2^b). Initial transients can result from the fact that O_2 rapidly leaves the GTD after it was lowered into the deeper anoxic waters from the equilibrated sea surface while at the same time N_2 entered the GTD because N_2 was supersaturated in the deeper anoxic core waters due to denitrification processes making modeling of the float data difficult. To accurately deconvolve all the raw gas tension signal required use of a multiple response time deconvolution model that separately tracked equilibration of each gas. Although we developed and tested this more complex deconvolution approach, we avoided using it by simply neglecting initial transients in the derived quantities (manifest as overshoots in the raw data) and determining equilibrium values for the later phase of equilibration since the additional complexity in the end added no value to the results based on the goal of determining equilibrium end-points. However, we add this information here to document that more complex details GTD equilibration in the ODZ can be understood if, for some reason, that were an objective of the analysis. By design, the thermal equilibration time of the GTD is smaller than the gas diffusion equilibration time scales (for N_2 and O_2) so thermal transients mostly relax quickly. This much simpler approach worked well under most circumstances, except during profiling across the strongest part of the thermocline and the only solution to that problem was to hold the sensors for longer at those depths

before moving the sensor to the next sampling depth as described above.

2.6 Biogenic N₂ from nutrient deficit

We use the **nutrient deficit** (N_{def}) method (Codispoti and Richards, 1976; Codispoti et al., 2001) to calculate an independent 'nutrient' (ΔN_2^{n}) based on the stoichiometry of microbial N₂ production from nitrate. N_{def} is the difference between measured and expected dissolved inorganic nitrogen (DIN) concentrations with DIN as the sum of measured NO_3^- and NO_2^- as NH_4^+ is typically undetectable in the open ODZ:

$$N_{\text{def}} = \text{DIN}_{\text{exp}} - (\text{NO}_3^- + \text{NO}_2^-) \quad \text{Eqn. (5)}$$

$$\Delta N_2^{\text{n}} = N_{\text{def}} / \alpha \quad \text{Eqn. (6)}$$

where the divisor α accounts for conversion to di-nitrogen (N₂) gas using appropriate stoichiometry.

DIN_{exp} is commonly estimated from PO_4^{3-} given its strong linear relationship with NO_3^- outside of ODZ's. Here we use regression coefficients to estimate DIN_{exp} derived from data in the GLODAP V2 database for regions adjacent to the western Mexico EDZ:

$$\text{DIN}_{\text{exp}} = 14.58 \times (\text{PO}_4^{3-} - 0.14). \quad \text{Eqn. (7)}$$

The slope and intercept for Eqn. 7 was determined through linear regression of data from the GLODAP database (Garcia et al., 2014) for the ETNP region excluding points with evidence of denitrification with results similar to Chang et al. (2012). For these calculations we used $\alpha = 2.0$. NO_3^- , NO_2^- , and PO_4^{3-} were measured on board using a SEAL AQ 400 nutrient analyzer using standard methods. PO_4^{3-} was also estimated *via* a linear regression with AOU data also using the GLODAP database. This approach was applied to the 12°N station, and one station at 17°N, as data quality problems with the directly measured PO_4^{3-} data was evident. For those calculations, $\alpha = 1.712$ was used following the stoichiometry of Richard's equation: $(\text{CH}_2\text{O})_{106}(\text{NH}_3)_{16}\text{H}_3\text{PO}_4 + 94.4 \text{ HNO}_3 = 106 \text{ CO}_2 + 55.2 \text{ N}_2 + 177.2 \text{ H}_2\text{O} + \text{H}_3\text{PO}_4$ (Richards, 1965).

2.7 Conservative preformed excess N₂

Equating the measured biogenic-N₂ derived from the GTD-method (i.e., ΔN_2^{b} from Eq. 4) to that measured by the nutrient-method (i.e., ΔN_2^{n} from Eqn. 6) using our CTD station data - even if the CTD stations were from inside of the ODZ region as will be discussed in Section 2.9 - we obtain a measurement-based estimate of the preformed excess-N₂ from:

$$\Delta N_2^{\text{p}} = N_2^{\text{m}} - N_2^{\text{s}} - \Delta N_2^{\text{n}}. \quad \text{Eqn. (8)}$$

By interpolating (on density) the shipboard measurements to common density levels, we derived vertical profiles of ΔN_2^{p} at the ODZ core. Measurements were made at approximately the same

latitude (17°N), in duplicate casts, and repeated one year later (see Table 1) to assess interannual variability. Argon and nutrients were typically measured at the same CTD station on the prior or next cast that the GTD measurements were made due to sampling limitations but Argon data are only available here for the first cruise (samples were collected during the second cruise but analysis is not yet complete).

In the deeper stratified waters of the ODZ, which is unaffected by air-sea gas exchange processes, we expect ΔN_2^{p} to be a conservative quantity. We also expect latitudinal variability in ΔN_2^{p} since the ODZ core is fed by multiple northern and southern sources, and each source water having been formed in different regions of the ocean will likely have different ΔN_2^{p} . An objective is to reliably predict ΔN_2^{p} across the entire ODZ based on circulation and conservative mixing of ODZ source waters then use those predicted distributions along with measured profiles of gas tension, temperature, salinity and oxygen to predict, using Eqn. (4), biogenic N₂ at the floats. We assess latitudinal variability in ΔN_2^{p} comparing estimates made in the ODZ core at 17°N to measurements made at 12°N which are strongly influenced by the southern sources. Since these two stations capture a large fraction of the variability in the ETNP ODZ hydrography, we expect the comparison to provide a first look at how variable ΔN_2^{p} is within the ETNP ODZ for different source water masses. Finding no differences in the deeper water masses below the active denitrification zone would be a good internal-consistency check of our methods, noting that we cannot independently verify our method without a second independent measurement of ΔN_2^{p} .

2.8 In situ denitrification rates

Measurements of N₂ made on a float following a subsurface water mass (Lagrangian) in theory allow *in situ* production (or consumption) rates to be determined by measurement of dN_2/dt if horizontal (isopycnal) and vertical (diapycnal) mixing effects are fully compensated. However, floats do not follow water masses perfectly, especially if the water mass properties change over time due to vertically sheared mixing which will dilute over time the original water mass under study with the surrounding waters. Notwithstanding these complications, local **denitrification rates** (DNR) along the float track-line can be estimated from the production of biogenic N₂ at the float using:

$$\text{DNR} = d(\Delta N_2^{\text{b}})/dt. \quad \text{Eqn. (9)}$$

Note that for a perfectly Lagrangian float with negligible vertical fluxes of biogenic N₂ or preformed biogenic N₂, the term ΔN_2^{b} in Eqn. (9) could be replaced by ΔN_2^{m} .

2.9 Argon measurements and assumptions

Although argon is a small component of the gas tension signal and assumptions for argon have little impact on the biogenic N₂ error analysis (see Section 2.4), argon still plays a significant role in

our work as a proxy measurement of inert- N_2 which, theoretically, would be a fully conservative quantity in the ODZ mesopelagic. Since we cannot measure conservative ΔN_2^p everywhere in the ODZ, we use argon to help predict distributions of ΔN_2^p based on the stirring and mixing of source waters within the ODZ. We capitalize on the fact that argon, like potential temperature and salinity, is a third conservative tracer and therefore expect Ar and ΔN_2^p to co-vary.

Source waters for the ODZ were last at the ocean surface outside the ODZ, likely during periods of strong air-sea interaction. During water mass formation, local atmospheric pressure may be several percent different from one atmosphere and rapid cooling may have caused further disequilibrium in gas saturation levels, but these processes affect N_2 and Ar similarly (Stanley and Jenkins, 2013; Hamme et al., 2019). Dissolution of bubbles from breaking waves also creates gas supersaturation, at nearly twice the rate for nitrogen as for argon due to their differences in solubility. However, given that both preformed excess N_2 and argon are inert and conservative, we expect variability in these quantities within the ODZ to be relatively small and to vary smoothly over large spatial and temporal scales in similar ways. Precise Ar concentrations can only be measured using advanced mass-spectrometric techniques (e.g., isotope dilution, see Section 2.9) and there are only a few existing measurements in the ETNP. Prior studies did not specifically target measurements in the ODZ source waters.

One particularly relevant study of the ETNP region (Chang et al., 2012) derived a relationship for the gas concentration ratio Ar: N_2 , which is much easier to measure than argon alone, as a function of potential density and geographically separated into ‘inside’ and ‘outside’ of the ODZ. Their analysis was based on a difference approach, subtracting ‘inside’ from ‘outside’ measurements at the same isopycnal to get estimates of the ODZ impact as a function of density. Extending their approach to consider our definition of ΔN_2^p , it is unclear how Ar: N_2 ratios should vary inside the ODZ, as mass spectrometry cannot discriminate ΔN_2^p from ΔN_2^b , and would imply that ΔN_2^p can only be inferred from gas ratio measurements if those measurements are taken ‘outside’ of the ODZ. This argument is fine but breaks down if the measurements made in waters ‘outside’ of the ODZ were previously exported from the ODZ, which is somewhat unavoidable when sampling unless close attention is paid to currents and the history of the current pathway is known when sampling. Also, Chang et al.’s data were collected mostly in the northern waters, which contains information on only half of the ODZ region. Further, their smoothed density relationship derived from ‘outside’ ODZ measurements blurs important information related to mixing of source waters (i.e., ‘spice’) which is important evidence for an analysis based on conservative mixing. Combining and replotting all available argon concentration data from Chang et al. (2012) and Fuchsman et al. (2018) showed very clear linear mixing lines, similar to T/S, which convinced us that there was potential for better parameterizations of ΔN_2^p rather than density alone. Here we argue that the Chang et al.’s approach can be fundamentally improved by focusing attention on the handful of northern and southern source waters to the ODZ, rather than ‘outside’ ODZ waters, and how these source waters get mixed within

the ODZ to form ‘inside’ waters. Our approach essentially intensifies the geographical constraints of the problem to the source waters, and demands new data in the southern waters. Our cruises tracks specifically targeted sampling of northern and southern water mass sources to the ODZ and measured argon in them. Most importantly, our improved approach does not require the measurement over all density ranges of ‘outside’ ODZ waters because our conservative mixing model can recreate ‘inside’ and ‘outside’ ODZ mixtures from water mass fractional composition analysis of source waters.

Our argon concentrations were measured by isotope dilution on discrete water samples (Hamme and Severinghaus, 2007; Hamme et al., 2017) collected on the first 2020 cruise. Briefly, glass flasks with a ~180-mL volume and a double-O-ring Louwers-Hapert valve were evacuated to a high vacuum prior to shipping to the field site. The necks of the flasks were kept evacuated by a rough vacuum pump on a large manifold at sea to prevent air leaking into the flasks. Duplicate water samples were introduced from Niskin bottles into the evacuated flasks through tubing pre-flushed with CO_2 until the flasks were half-full. Back on shore, flasks were weighed, the headspace and water were equilibrated and then the water was removed. Headspace gases were flowed through ~90°C ethanol to remove water vapor, gettered to remove all except noble gases, an aliquot of ^{38}Ar added, frozen into a tube immersed in liquid helium, and brought to higher pressure with a balance gas of helium. Argon isotopes were analyzed on a dual-inlet MAT 253 Isotope Ratio Mass Spectrometer at the University of Victoria against a standard of similar composition. The standard and ^{38}Ar aliquot amount were calibrated relative to air measurements. The pooled standard deviation of the duplicates for Ar on this cruise was 0.21%.

2.10 Linear Mixing model

To aid interpretation of our station data at 17 °N in the core waters of the ODZ, a linear mixing model was developed based on the five major water types for that region. The model was initialized using shipboard observations of the three conservative tracers, namely potential temperature, salinity and argon concentration. Although argon saturation level is not conservative, its strong sensitivity to the choice of water types values helps fine tune the model to best match the observations – this sensitivity comes from the temperature and salinity dependence of the Henry’s Law solubility coefficients. Unlike argon, nitrogen gas is non-conservative in the ODZ due to N_2 production by microbial activity in functionally anoxic waters. The abiotic component of N_2 , which is equivalent to the red line in Figure 2, cannot be measured directly but can be modeled using the calibrated mixing model with conservative preformed excess N_2 from Eqn. 8 as inputs.

To determine conservative property values for each water type, a ‘broken-stick’ best fit model was initially applied to the TS data and then to argon concentration. A segmented linear regression (or ‘broken-stick’) routine called SML (Shape Language Modeling) and developed in MATLAB was used with unconstrained setting (i.e., ‘free knots’) to find the location in T-S space of the four water type values. These values were then used in the linear mixing model. The

process was repeated for argon and salinity then water type values for each conservative property were further refined by tuning the model to fit all observations, including argon saturation levels, simultaneously. The entire process proved too complicated to fully automate, so the values reported here are approximations derived ultimately by visual best fit between the model and the data.

Once the model is best fit to the observations, the model provides a convenient way to determine smoothed versions of any quantity (e.g., of ΔN_2^P or ΔN_2^b) versus potential density that satisfies the model's conceptual framework (ie., dilution of 5 water types). The results at 17 °N were then compared to results at 12 °N.

3 Results

3.1 Gas tension

An example of raw and processed time series data from an early float (Table 1, F89) mission are shown in Figure 3A. After the float sunk to its maximum depth (700 m) it slowly ascended to the sea surface, stopping at many (17) targeted depths (or isopycnals) to equilibrate the GTD. The wait time at each targeted depth was 2 hours but this was extended to 24 hours at the 'Lagrangian drift' depth (242 ± 4 m) on this dive. The total submerged time was approximately 2 days. Data were reported by satellite after surfacing. A constant e-folding response time of 20 minutes was used to deconvolve these raw GTD data (see Figures 3A, C and compare red and blue lines). This is a factor of 10–20 times slower than the fastest response of the sensor achievable using a fast flow rate pump but draws little energy from the float's batteries since it uses a small flow rate pump. The equilibrated gas tension values at

each hold-depth used a 15-minute average of the deconvolved time series just before the float is moved to the next hold-depth. The standard deviation, SD, of the data during the 15 min averaging period provides a good indicator of uncertainty. We found for $P > 300$ dbar that $SD < 0.03\%$ and for $P < 200$ dbar that $SD < 0.12\%$ (see Figure S1 in Supplementary Materials). Internal waves in the strongly stratified thermocline of the ODZ were the leading cause of geophysical noise.

Processing of the shipboard data was much simpler and didn't require deconvolving the raw data because the hold time of the GTD at any depth (10–15 minutes) was chosen to be much longer than the sensor response time (2–3 minutes, using the high flow rate pump). The noise level of the raw ship GTD data was better than ± 0.2 mbar and the end-point detection typically within ± 0.5 mbar assessed by variability from steady readings (see Figure S2 in Supplementary Materials).

To summarize, vertical profiles of gas tension (see Eqn. 1), temperature, salinity and dissolved O_2 were made at 12 °N and 17 °N by ship and floats during two cruises. Additional processing produced estimates of dissolved N_2 (see Eqn. 2), the results of which are described next.

3.2 Regional water mass analysis

Shipboard measured θ , S, Ar, and GTD-derived N_2 at the ODZ core near 17 °N are shown in Figure 4. Five regional water types are readily identified on the conservative θ -S diagram (Figure 4A, red dots) as 'break-points', defined where the slope of the plot changes rapidly. In theory, any water mass, or layer in this stratified system, that lies between two water types on the θ -S diagram can be

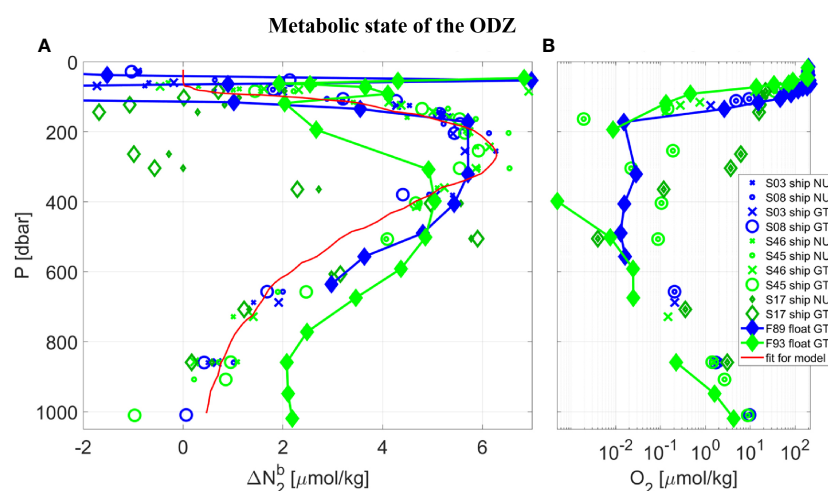


FIGURE 5

Vertical profiles taken in the core of the ETNP near 17 °N, from SR2011 (blue) and SR2114 (green), of: (A) biogenic N_2 , and (B) dissolved O_2 using a logarithmic scale. Following Table 1 and the Legend, data collected at CTD Station ('S') by ship or float ('F') used either the traditional nutrient method ('NUT') or the new gas tension method ('GTD') to measure biogenic N_2 . Note that all data labeled 'ship NUT' are used to calculate the best fit of $\Delta N_2^P(\sigma_\theta)$ so the resulting data labelled 'ship GTD' are calibrated and therefore must agree. Data labelled 'float GTD' also use $\Delta N_2^P(\sigma_\theta)$ but their gas tension and oxygen are independent. Since 'F89 float GTD' data were taken during the cruise and in the vicinity of the float the good agreement serves to validate the GTD method. The 'F93 float GTD' data were measured autonomously 2 months after the cruise and much further east (see Figure 1) with significantly different hydrography and are not expected to necessarily agree with 'F89 float GTD' data. A second order Gaussian best-fit curve (red line) to the shipboard nutrient data is used for mixing-model calculations.

described as a mixture of two adjacent water types. The linear mixing model (see Section 2.10) recreates all possible mixtures of water types that could occur between adjacent water types to produce the model result (Figure 4A, red line). The model can only produce straight lines between water types when quantities are conservative, such as θ , S, and Ar. The model easily reproduces the same 'broken stick' shape for conservative Ar (Figure 4B, solid red line) that matches the shipboard data. Derived model quantities that are non-conservative can also be calculated from the conservative model output. Gas saturation levels are one such calculation. Since Henry's Law solubility coefficients are non-linear functions of θ and S, a model Ar saturation level is calculated using the modeled θ , S, and Ar. Since Ar saturation level is non-conservative, the model produces curved lines between water types on a conservative plot and also on density (see Figure 2). The model Ar saturation level reproduces the observed Ar saturation level curved-shape (Figure 4B, semi-transparent red dashed line) and, as mentioned in Section 2.10, is sensitive to the choice of end-member values and provides an additional constraint for tuning the model to the data.

The now calibrated model is used to test the assumption that excess pre-formed N_2 derived from the GTD-measured N_2 (Figure 4C) is conservative. The model ΔN_2^P versus σ_θ (Figure 4D, red lines) shows breakpoints in density that correspond well with equivalent breakpoints on both θ -S and Ar curves which used both 2021 and 2022 station data due to the limited amount of calibrated data available. These broad similarities provide corroborating evidence that preformed excess N_2 is indeed conservative and values for each conservative property can be estimated for each water type. Based on this result, the water type values can be further refined so that preformed excess N_2 also is included in the optimization approach to determine water type properties. Our best estimate of water properties for all five water types is summarized in Table 2 and compared to previously

published values (Evans et al., 2020). As previously noted, our results do not use an automated 'broken stick' detection algorithm and hence errors in the fits are not fully quantified.

For comparison purposes, additional argon data at 14–15.5 °N from cruise SR2011 were added to Figure 4B to help expand the limited information on argon beyond the site. In addition, corresponding $\theta/S/\Delta N_2^P$ at 12 °N from SR2114 were added to Figures 4A, D to provide some additional testing of these new methods. On closer inspection, lower values of ΔN_2^P , by up to 1 $\mu\text{molN}_2/\text{kg}$, were observed at 12 °N in the potential density range of $25.7 < \sigma_\theta < 26.3$ compared to those values at 17 °N indicating water masses (or mixtures of water types) at 12 °N that were not found at 17 °N. Additional evidence is seen on the θ S plot (see Figure 4A, dark green diamonds) where these same water at 12 °N are found to have higher salinity than any waters near the same density range at 17 °N. Since it is impossible for the mixing model calibrated at 17 °N to reproduce water masses at 12 °N that have higher salinity than water types found at 17 °N, the anomaly is understandable. Application of the model calibrated to 17 °N to data at 12 °N manifests as a latitudinal dependence to ΔN_2^P . Considering the regional water masses identified in prior work (Evans et al., 2020), a likely explanation is that we are seeing purer 13CW water at 12 °N than at 17 °N. The obvious solution is to expand the analysis to other latitudes to cover all θ -S space in the ODZ.

In general, all additional data compare favorably with the main results from 17 °N to provide additional confidence that these methods are not simply unique to the core region and have broader applicability over the entire ODZ region, as we would expect given the model framework is based on conservative mixing.

3.3 Biogenic N_2 and O_2

The model's best fit of ΔN_2^P (σ_θ) at 17 °N (Figure 4D, red line) was then used to determine GTD-derived estimates of ΔN_2^b using

TABLE 2 Comparison of water types and properties of the ETNP identified by Evans et al. (2020) and water types or water mass mixtures identified in this work (bold text) at 17 °N: Upper Circumpolar Deep Water (UCDW), Antarctic Intermediate Water (AAIW), North Equatorial Pacific Intermediate Water (NEPIW), 13°C Water (13CW), Pacific Subarctic Upper Water (PSUW), Equatorial Surface Water (ESW), and Tropical Surface Water (TSW).

Row	Evans et al. (2020)			This work			
	Source Water Type or Water Mass Mixture	Theta (°C)	S (psu)	Theta (°C)	S (psu)	ΔN_2^P ($\mu\text{mol/kg}$)	Ar ($\mu\text{mol/kg}$)
1	UCDW	1.42	34.87	-	-	-	-
2	mix of UCDW + AAIW	3.725	34.785	4.31	34.55	12.60	15.40
3	AAIW	6.03	34.7	-	-	-	-
4	mix of AAIW + NEPIW	7.795	34.75	7.0	34.51	12.55	14.50
5	NEPIW	9.56	34.8	-	-	-	-
6	13CW	12.81	34.965	12.5	34.77	14.91	13.06
7	mix of NEPIW + PSUW	12.94	34.465	-	-	-	-
8	PSUW	16.32	34.13	-	-	-	-
9	ESW	22.27	34.5	20.4	34.3	28.0	11.63
10	TSW	26.95	34.06	27.3	34.55	4.78	10.30

Eqn. 4. This conveniently ensures that any water mass composed of a mixture of two adjacent water types uses the correct and smoothed mixed water mass value for ΔN_2^b . The results for biogenic N_2 alongside O_2 are shown in Figure 5. The biogenic N_2 peaks at just below the oxycline (within 100 m) at 17 °N. Note that since ΔN_2^P (σ_0) was derived by matching the ship GTD-derived biogenic N_2 estimates to the nutrient-derived estimates at 17 °N, the shipboard nutrient and GTD results must agree and the residuals simply provide information on the errors in the data and analysis procedure. A regression fit to a double Gaussian that is constrained to zero at the surface and deepest waters (see Figure 5A, red line) shows an R-squared of 0.96 and RSME of $0.54 \mu M N_2$. The first independent float-based measurements of ΔN_2^b using the same method applied to gas tension and oxygen data from Float 89 within 26 km of the shipboard station generally agrees with the shipboard calibration measurements over $130 < P < 630$ dbar or $O_2 < 3 \mu M$: Density interpolated measurements show a mean ($n=7$) difference of $-0.42 (\pm 0.78 \text{ RMSE}) \mu M N_2$, or $10 (\pm 19) \%$. This result provides a first quantitative assessment of fully autonomous measurements of biogenic N_2 using GTD-equipped floats. ODZ core water O_2 near Float 89 was uniform over $160 < P < 700$ dbar and assumed anoxic since shipboard STOX O_2 was ± 25 nmol/kg and float optode O_2 was likewise near zero and uniform (after correcting for the anoxic optode pressure and temperature coefficients).

For comparison to the results from the first float, we present additional measurements made by a second float (see Table 1, float F93) several months after the end of the second cruise and located approximately 350 km to the southeast of the core station. The second float showed lower O_2 and higher biogenic N_2 in the deeper waters (see Figure 5, float F93 results in light green filled diamonds) consistent with expected higher denitrification rates in less oxygenated waters but less biogenic N_2 at comparable O_2 in the shallower waters. This could indicate slightly different hydrography between the two locations, with the deeper waters being older at the second float, or variations in denitrification forcing, we don't know. We also have no independent comparison of Float 93 optode-derived O_2 to shipboard STOX due to the fact that the second float data were collected several months after the ship left the region, but expect the absolute value of oxygen is uncertain to several hundred nmol/kg or more. The ΔN_2^b and O_2 profiles from both floats are otherwise fairly consistent which independently supports the results from the first float.

For additional observations with which to compare and contrast with the first float, which was located near 17 °N in the ODZ core, we look to shipboard data from lower latitudes and specifically southern waters which will have different histories of biological processing than those in the core region. Indeed, the largest deviations were observed were near 12 °N where the upper oxycline was much deeper than at the core station with distinctly different water mass properties (see Figure 4A, dark green non-filled diamonds). Remarkably, the same characteristic of biogenic N_2 peaking just below the oxycline was reproduced at 12 °N even though the oxycline at 12 °N was approximately 200 m deeper than at 17 °N. Most important, the apparently low ΔN_2^b values at $P = 265$ and 305 dbar correspond with significantly high CTD measured

O_2 ($3.6 < O_2 < 6.2 \mu M/kg$) compared to other mesopelagic waters during the same cast (ie., it likely isn't a measurement error as the CTD O_2 goes well below $1 \mu M$ during the same cast) along with observed low and near zero ΔN_2^b . Since these hypoxic waters are typically considered well above the threshold for functional anoxia, we would not expect these waters to have previously hosted denitrification processes leading to N_2 production, leading us to believe these hypoxic waters are the southern source waters we had set out to find. The remaining oxygen in the waters ($4.9 \pm 1.3 \mu M/kg$) would need to be lost to respiration processes before becoming functionally anoxic and capable of hosting a robust denitrifying microbial community. This rather exciting discovery also lends support to our overall approach. We also note, however, that no truly independent measurement of biogenic N_2 has yet been made at either location and hope to make those measurements in the future.

3.4 Lagrangian drift

An example of various estimates of dissolved N_2 variability measured during a 20 hour drift period of a float at a mean potential density of 26.474 kg/m^3 are shown in Figure 6. Variability in dissolved N_2 (Figure 6, blue line) measured by the gas tension technique is very small, about $1 \mu M/kg$, or 0.18%. This is comparable to the uncertainty in Henry's Law solubility coefficients. Variability in biogenic N_2 (thick red line) is even smaller, at $<250 \text{ nmol/kg}$ or $<0.05\%$. The impact of excluding dissolved O_2 in the calculation of biogenic N_2 is nearly undetectable given that measured O_2 variability is at the limit of

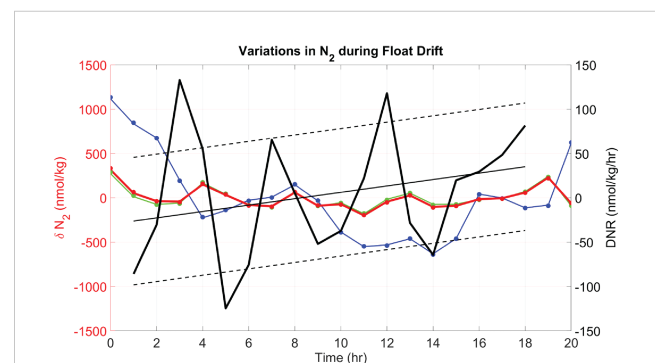


FIGURE 6

Deviations in water mass N_2 from the mean (δN_2) during the Lagrangian drift of a float on mean isopycnal of 26.474 kg/m^3 and $P = 242 \pm 4$ dbar (F89 starting 31-Dec-2020 at 09:00 GMT). Different calculations of δN_2 , using a specified ΔN_2 and O_2 , are shown: 1) ΔN_2^b with measured O_2 (red line), 2) ΔN_2^e with measured O_2 (green line), 3) N_2^m with measured O_2 (blue line), and 4) similar calculations as (1) - (3) but assuming anoxia (dots, using same color code). Note that since the plots show deviations from the mean, the mean values of δN_2 zero. The Lagrangian rate of change of ΔN_2^b is shown (left), calculated as local denitrification rate DNR (thick black line) using Eqn. 9, with corrections based on the local vertical gradient. A linear regression fit (thin black line) highlights slowly varying changes in the rates, with ± 1 std deviation of the points (black dashed lines) as an indicator of variability. Notice that DNR is negative during the first half of the time series and changes to positive values during the second half of the time series.

detection (LOD) and indistinguishable from anoxic conditions. Calculation of rates (Figure 6, black lines) allow trends in ΔN_2^b to be detected at $\pm 100 \text{ nMN}_2/\text{day}$: Positive rates would imply active denitrification after compensating for mixing effects and slippage of the float into different waters (which would explain negative rates). These rates and trends seem to be well above the LOD for ΔN_2^b since internal wave effects with periods of approximately 2 hr are well resolved, implying that longer term averaging (assuming Gaussian distribution) over 16.5 days could reduce these uncertainties to $\pm 20 \text{ nM N}_2/\text{day}$. This would be suitable for detection of active denitrification in coastal waters of the ETNP where rates of up to $20 \text{ nM N}_2/\text{day}$ have been reported (Babbin et al., 2020), but further resolution would be required to detect active denitrification in the offshore waters with reported rates of up to $4 \text{ nM N}_2/\text{day}$ (Babbin et al., 2020).

4 Discussion

The amplitude ($6 \pm 1 \mu\text{molN}_2/\text{kg}$), profile shape, and depth distribution of the biogenic N_2 with a peak at 300 m just below the oxycline agree very well with the prior detailed study of Fuchsman et al. (2018) for the ETNP ODZ core. This same study also noted significant and unexplained mismatches in excess and biogenic N_2 of $1\text{--}2 \mu\text{molN}_2/\text{kg}$ near 15°N . The authors hypothesized that water mass mixing did not following their empirical ‘background’ (or ‘outside ODZ’) values. Similarly, we found these anomalies at lower latitudes, specially 12°N , but our water mass mixing approach provides a path toward resolving them. Southern source waters to the ODZ have different preformed excess N_2 values at the same density range as waters further north around 17°N yet span the same density range. A simple density parameterization of ‘background’ excess N_2 is thus not sufficient to explain all variability across the entire ODZ. Our approach to reconstruct the equivalent abiotic excess N_2 from a mixture of various water types with different preformed excess N_2 values should ultimately provide improved biogenic N_2 estimates at all latitudes by either method. The relatively small latitudinal dependence of ΔN_2^b we found by comparing our method applied at 17°N to observations from 12°N is expected to be resolved by properly accounting of all water types in the model at 12°N . Correctly predicting ΔN_2^b over the entire ETNP ODZ based on a linear mixing model of different water types is an overarching goal.

Our primary nutrient-based estimates of biogenic N_2 were used in calibrating our novel gas tension-based estimates, and so a current weakness is the lack of a second independent estimate for comparison. Since many more hydrocasts were taken during both cruises that are not yet fully analyzed, we plan on using these shipboard data at other stations to both calibrate (i.e., quantify preformed excess N_2) and validate (i.e., quantify absolute accuracy). Accuracy and calibration drift is not expected to be limited by the raw measurement of gas tension, although dissolved O_2 could introduce comparable errors at either high O_2 levels or if the O_2 sensor has a significant zero reading offset ($>2 \mu\text{M}$). To perform the extended analysis over the entire ETNP ODZ and synthesize all our observations, we want to step back and refine a procedure that identifies water types and their values from the ‘broken stick’

analysis of multi-conservative parameter data. Work is in progress to fully automate this procedure with the expectation that the water types and property values can be determined with quantified confidence intervals, a non-trivial task given the fitting needs to be done on scattered data in multi-dimensions with many unconstrained break-points. An alternative approach would be to use a regional numerical model of the ODZ that includes dissolved Ar and a representation of abiotic N_2 , conservatively redistributes these quantities within the model domain, and is tuned to the observations inside and outside the anoxic core. This would facilitate inclusion of N_2/Ar measurements and make better connections between this method and those of Fuchsman et al. (2018) and Chang et al. (2012). The water masses created by mixing of source water types would be subject to parameterization of diapycnal and isopycnal mixing in the model.

Regarding the potential of floats to directly measure *in situ* denitrification rates (DNR), our first attempt amounts to a test of the procedure given the paucity of the data presented here and the low denitrification rates of the offshore region. It seems clear, however, that improved mapping of ΔN_2^b and understanding of mixing at the float are prerequisites to inspecting additional data we have collected for scientifically interesting results. There will likely be tradeoffs while trying to reduce uncertainty by using longer averaging periods versus the potential for the float to slip out of the water mass under study and into a different water mass of similar density during the drift. Analysis of longer duration float-drift data is clearly required, but the initial results are encouraging since they suggest that the higher denitrification rates typical of the more productive coastal waters of the ETNP (up to $20 \text{ nmolN}_2/\text{kg/day}$) should be resolvable using this method by extending the float drift period to more than 16 days. This is a great start.

Lastly, regarding our experiences using the shipboard GTD at sea, we note that each hydrocast took approximately 2.5 hrs longer due to the time required to equilibrate the GTD at the 12 sampling depths – this could be reduced to approximately 1 hr if analysis relied upon deconvolution of the data but more time would need to be spent determining the response time characteristics of the mounted GTD. The same ship’s GTD used on cruise SR2011 was used on the first half of the cruise SR2114 but succumbed to a small crustacean that got trapped in the plenum against the thin membrane and punctured it and had to be swapped out for a spare.

5 Summary

We report for the first time direct autonomous vertical profile measurements of gaseous dissolved biogenic N_2 and dissolved O_2 made on two profiling floats. The profiles were collected near 17°N in the ETNP ODZ. The data show peaks in biogenic N_2 just below the oxycline in the mesopelagic and continues to decrease until oxygen begins to increase in the deeper intermediate waters. Our measurements are the result of a new method (calculated using Eqn. 4) we have developed to autonomously determine biogenic N_2 on floats, and potentially other oceanographic platforms, based on measurement of N_2 by the gas tension method. Conservative ‘linear-mixing’ of multiple water types is used to predict abiotic dissolved N_2 in

a mixed water mass (i.e., a mixture of several source water types). We have defined, measured, and used a new conservative quantity which we call ‘preformed excess N_2 ’. We calibrated our method (see Eqn. 8) at 17 °N in the anoxic core region using shipboard measured gas tension, temperature, salinity, dissolved oxygen, and nutrients, and used argon as a constraint. The RMSE of the new method was assessed at $\pm 19\%$. The method can be improved by reducing and quantifying uncertainty in how water types and their conservative preformed N_2 are determined. Our shipboard sampling targeted and successfully found northern and southern source waters to the ETNP ODZ. The southern waters near 12 °N had not been sampled before for argon, gas tension or preformed excess N_2 and were found to contain very low O_2 ($5 \pm 1.3 \mu\text{mol/kg}$) and near zero biogenic N_2 . Our new method can also be applied to Lagrangian isopycnal data to determine *in situ* biogenic N_2 production rates. Our first attempt shows that our new method has a signal-to-noise ratio that should be useful for assessing productive nearshore waters of the ETNP ODZ but needs improvement for offshore waters with significantly lower denitrification rates.

Data availability statement

The original contributions presented in the study are included in the article/Supplementary Material. Further inquiries can be directed to the corresponding author.

Author contributions

The first three authors have worked equally to formulate the methods and apply them to the data. All authors have collaborated over the last 6–8 years to measure dissolved gases in the ETNP during three cruises. CM was responsible for gas tension derived N_2 , EAD for floats with O_2 , MA for nutrients and at sea mass-spectrometry as well as being Chief Scientist on the cruises, RH for Ar, and EG-R for STOX O_2 . Each have contributed to data planning, collection, processing, and QC of their data and sharing for this manuscript. All authors contributed to the article and approved the submitted version.

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Conflict of interest

CM is a co-founder and Vice President of the company Pro-Oceanus Systems, Inc. PSI who manufactures the GTDs used in this work – his potential conflict of interest is managed by a plan developed by the UW’s Office of Research Compliance.

The remaining authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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Supplementary material

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fmars.2023.1134851/full#supplementary-material>

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