

Characterizing the Evolution of Extreme Water Levels with Long Short-Term Memory Station-based Approximated Models and Transfer Learning Techniques

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9 Abstract

10 Extreme water levels (EWLs) resulting from tropical and extratropical cyclones pose significant
11 risks to coastal communities and their interconnected ecosystems. To date, physically-based
12 models have enabled accurate characterization of EWLs despite their inherent high computational
13 cost. However, the applicability of these models is limited to data-rich sites with diverse
14 morphologic and hydrodynamic characteristics. The dependence on high quality spatiotemporal
15 data, which is often computationally expensive, hinders the applicability of these models to regions
16 of either limited or data-scarce conditions. To address this challenge, we present a computationally
17 efficient deep learning framework, employing Long Short-Term Memory (LSTM) networks, to
18 predict the evolution of EWLs beyond site-specific training stations. The framework, named
19 LSTM-Station Approximated Models (LSTM-SAM), consists of a collection of bidirectional
20 LSTM models enhanced with a custom attention layer mechanism embedded in the model
21 architecture. Moreover, the LSTM-SAM framework incorporates a transfer learning approach that

22 is applicable to target (tide-gage) stations along the U.S. Atlantic Coast. The LSTM-SAM
23 framework demonstrates satisfactory performance with “transferable” models achieving average
24 Kling-Gupta Efficiency (KGE), Nash-Sutcliffe Efficiency (NSE), and Root-Mean Square Error
25 (RMSE) ranging from 0.78 to 0.92, 0.90 to 0.97, and 0.09 to 0.18 at the target stations, respectively.
26 Following these results, the LSTM-SAM framework can accurately predict not only EWLs but
27 also their evolution over time, i.e., onset, peak, and dissipation, which could assist in large-scale
28 operational flood forecasting, especially in regions with limited resources to set up high fidelity
29 physically-based models.

30 **Keywords:** long short-term memory networks, transfer learning, extreme water level, tropical
31 cyclones

32 1. Introduction

33 About 11% of the world's population (890 million people) currently resides in low-lying areas,
34 and according to the Intergovernmental Panel on Climate Change, this number is projected to
35 exceed 1 billion by the year 2050 (Pörtner et al. 2019; Glavovic et al. 2022). Low-lying areas are
36 particularly vulnerable to weather and climate disasters which are responsible for severe
37 socioeconomic and environmental impacts (Zscheischler et al. 2020; Rainey et al. 2021). The
38 United States, accounting for 1.6% of the current global population (129 million people) in low-
39 lying areas (Office for Coastal Management, 2024), has reported more than 377 weather and
40 climate disasters since 1980 (NOAA-NCEI 2024). In the same period, total reported losses exceed
41 \$2.67 trillion when adjusted for the 2024 Consumer Price Index (NOAA-NCEI 2024). Among
42 these disasters, six of the world's costliest hurricane events resulted in over \$50 billion in damages
43 in the United States (Douris and Kim 2021; Sanders et al. 2022). Hurricanes (or tropical cyclones)

44 are responsible for coastal flood hazards characterized by extreme water levels (EWLs) and
45 exacerbated by climate-related impacts, including regional sea level rise and anthropogenic
46 activities (Hino and Nance 2021; Khojasteh et al. 2021).

47 EWLs in estuarine and coastal systems arise from various flood drivers, including
48 precipitation, river discharge, storm surge, tides, and waves. Yet, these drivers do not necessarily
49 act in isolation but rather synergize resulting in compound flooding (Wahl et al. 2017; Muis et al.
50 2019; Parker et al. 2023). Compound flood (CF) hazards, and their associated risk to coastal
51 communities, are particularly severe when flood drivers co-occur or unfold in close succession
52 (Arns et al. 2020; Almar et al. 2021). For example, storm surge can co-occur with extreme
53 precipitation events during tropical cyclones (TCs) (Wahl et al. 2015; Bevacqua et al. 2019), high
54 tide can coincide with the peak of a storm surge (Thomas et al. 2019; Marsooli and Wang 2020),
55 peak river flow and storm surge can co-occur along estuarine systems (Moftakhari et al. 2019;
56 Muñoz et al. 2020), and waves and storm surge can interact nonlinearly (Rueda et al. 2016; Serafin
57 et al. 2017a); thereby amplifying the effects of CF events. It has been noted that changes in
58 storminess would also play a major role in future EWLs (Santiago-Collazo et al. 2019). This is
59 corroborated by the increasing frequency and intensity of TCs along with the rise of sea levels and
60 ocean temperatures over the past 35 years (Anderson et al. 2021; Ghanbari et al. 2021;
61 Bloemendaal et al. 2022).

62 TCs have been responsible for 60% of flood-induced population displacements in the United
63 States (1985 to 2021), especially in densely inhabited coastal cities along the Gulf of Mexico and
64 the Atlantic Coast (Brakenridge 2021; Tate et al. 2021; Wing et al. 2022). Recognizing the
65 heightened risks to coastal communities, it has become imperative for researchers and practitioners
66 to rely on either physically-based or data-driven modeling approaches to characterize EWLs in

67 terms of peak magnitude and timing. Physically-based models are commonly used to estimate
68 EWLs based on simplified hydrometeorological processes governed by the conservation of mass
69 and momentum equations (Santiago-Collazo et al. 2019; Bates 2023). The accuracy of these
70 models depends on the availability and quality of several spatiotemporal datasets to appropriately
71 characterize input and forcing conditions, topography and bathymetry, land surface roughness, and
72 other key morphologic characteristics (Jafarzadegan et al. 2021; Alipour et al. 2022; Bates 2022a).
73 Nevertheless, such models are often constrained by limited spatial scope and/or high
74 computational demands necessary to solve large-scale flood dynamics (Bilskie et al. 2021; Muñoz
75 et al. 2021). While physically-based models developed with a lower spatial resolution (e.g., cell-
76 grid size and digital elevation model resolution) can cover broader areas and reduce computational
77 time, they can lead to less accurate predictions due to a lack of detailed spatiotemporal information
78 around key morphological and hydrodynamic variables in narrow tidal inlets and river channels
79 (Saksena and Merwade 2015; Fraehr et al. 2022).

80 In contrast, data-driven models such as neural networks (NN) can discern intricate or hidden
81 patterns in large datasets and predict storm surges and EWLs with reduced computational demands
82 when compared to those of physically-based models (Muñoz et al. 2021). Importantly, data-driven
83 models offer rapid and efficient forecasting solutions at large scales (Lee et al. 2021; Hamitouche
84 and Molina 2022; Hamidi et al. 2023) and have the ability to generalize or identify patterns from
85 the data they are trained on. In addition, these models can be updated over time which improves
86 their predictions as more and new information becomes available. The fact that NN models are
87 inherently adept at capturing nonlinear associations in complex systems makes them reliable
88 candidates for EWL prediction (Tedesco et al. 2023). Particularly, deep learning approaches like
89 long short-term memory (LSTM) networks, a variant of the recurrent neural network (RNN), learn

90 nonlinear relationships and patterns from sequential time-series data, to enhance prediction
91 accuracy in hydrological and coastal contexts (Li et al. 2021; Zhang et al. 2022).

92 A growing body of research is demonstrating that LSTM networks can predict EWLs and aid
93 in flood susceptibility assessments, barrage integrity, riverine flood level forecast, and surge
94 prediction (Tiggeloven et al. 2021; Fang et al. 2021; Kardhana et al. 2022; Merizalde et al. 2023a;
95 Liu et al. 2023). LSTM networks are designed to recognize sequence-to-sequence patterns and
96 selectively retain information over time, which in turn enhances its predictive accuracy by utilizing
97 memorized patterns (Hewamalage et al. 2021; Lindemann et al. 2021). On the global scale, LSTM
98 networks outperformed other NN models designed for surge prediction at 92% of 1,276 tide
99 stations across regions of Europe, Africa, Australia, the Pacific, and the United States (Tiggeloven
100 et al. 2021). In addition, LSTM networks can be integrated with spatial NN algorithms to improve
101 the modeling of geographical correlations (Gavahi et al. 2021) and even designed to capture spatial
102 flood characteristics (Fang et al. 2021). Unlike physically-based models that are typically confined
103 and developed using site-specific information, LSTM networks can be trained using geographical
104 characteristics, morphological and hydrodynamic features, and forcing drivers at different scales.
105 The resulting learned patterns can then be generalized and applied to neighboring regions through
106 transfer learning (TL) techniques.

107 TL addresses the challenge of either data scarce or insufficient training data by leveraging
108 gained knowledge from data-rich training domains and applying it to other (target) domains that
109 share similar characteristics or features (Shen 2018; Tan et al. 2018; F. Zhuang et al. 2021).
110 Moreover, TL can be used to expedite decision-making processes and circumvent time constraints
111 associated with the development and training of NN models. Several studies have implemented
112 TL techniques in NN models to estimate urban flood levels (Zhao et al. 2021; Seleem et al. 2023),

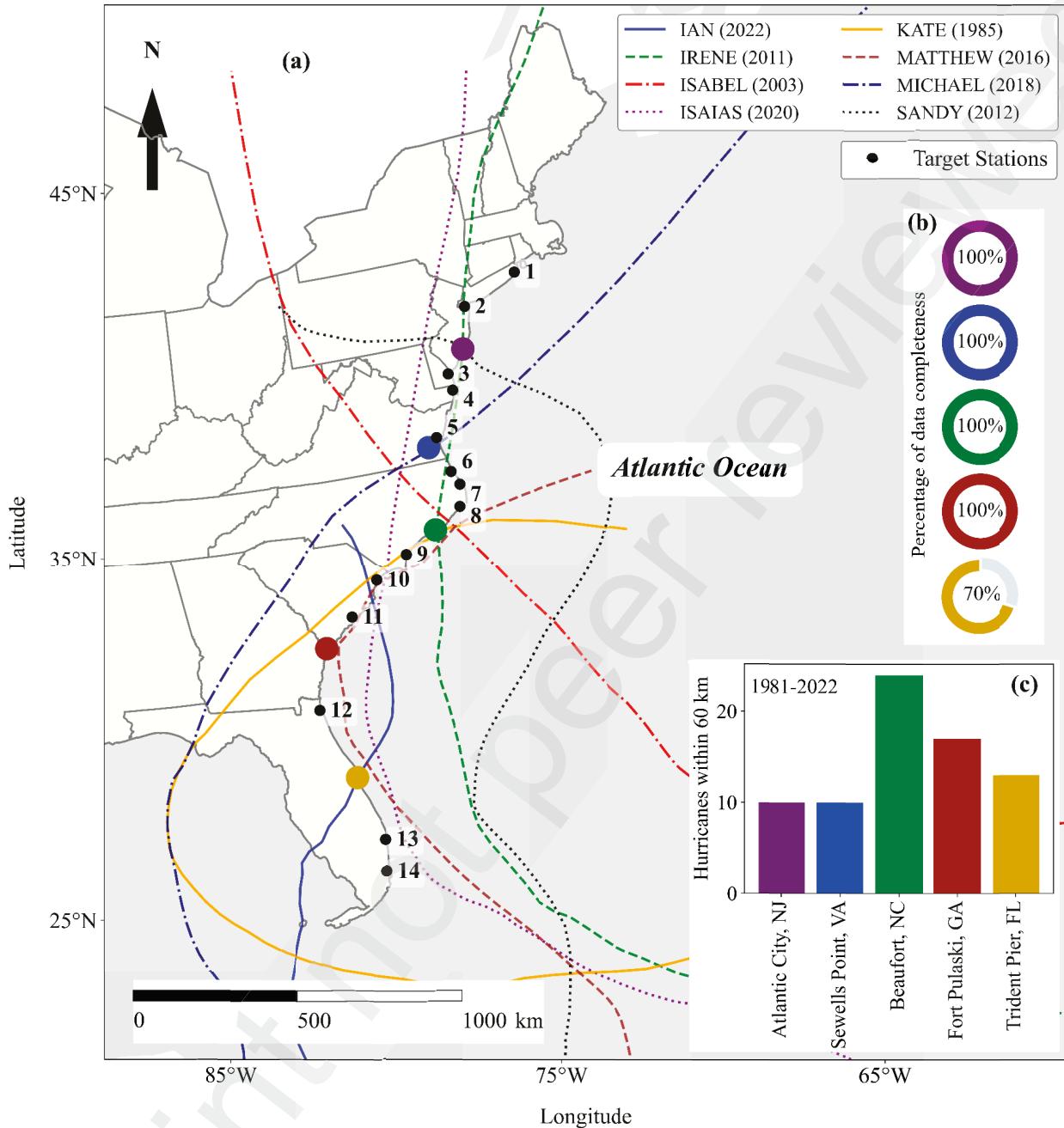
113 predict significant wave height (Obara and Nakamura 2022), conduct land cover mapping
114 (Mahdianpari et al. 2018), and compound flood hazard characterization of nearby regions to the
115 training domain (Muñoz et al. 2021). However, creating a data-driven model with effective
116 generalization capabilities beyond its training domain still remains a significant hurdle
117 (Bentivoglio et al. 2022; Bates 2022b). While maintaining consistency in location enhances the
118 accuracy and lead-time of model predictions (Altunkaynak and Kartal 2021), this limits the
119 geographical areas suitable for effectively applying TL techniques. Therefore, NN models should
120 learn patterns from nonlinear interactions among inputs features and further benefit from
121 mechanisms that ensure accurate model predictions at target domains.

122 In the present study, we introduce a comprehensive framework that (i) accurately predicts the
123 evolution of extreme water levels beyond training domains, and (ii) addresses the underlying
124 limitations attributed to transfer learning techniques. The proposed framework, named LSTM -
125 Station Approximated Models (LSTM-SAM), achieves these two objectives by gathering learned
126 patterns from neighboring tide-gage stations of the U.S. Atlantic Coast and optimizing the LSTM
127 models with an attention layer mechanism during the training phase. The remainder of the
128 manuscript continues as follows. Section 2 presents the study area, data availability, data
129 processing, and model architecture. Results of the proposed LSTM-SAM framework are shown in
130 section 3 and discussed in section 4. Lastly, section 5 presents the conclusions of this study as well
131 as future work.

132 **2. Methods**

133 **2.1 Study area**

134 The proposed LSTM-SAM framework is trained using time-series data from 5 strategically
135 selected tide-gage stations located along the U.S. Atlantic Coast. These stations are Atlantic City,
136 NJ (NOAA ID: 8534720), Sewells Point, VA (8638610), Beaufort, NC (8656483), Fort Pulaski,
137 GA (8670870) and Trident Pier, FL (8721604) (Figure 1). The training stations are selected based
138 on two criteria: (i) they have been hit by hurricane events within a radius of 60 km of the landfall
139 location, and (ii) they have over 70% consecutive water level (WL) data spanning at least 40 years.
140 The latter ensures that the training stations contain EWLs attributed to either TCs (hurricanes) or
141 extra-TCs (Nor'easter winter storms) to effectively train and validate the LSTM-SAM framework.
142 We then implement a TL approach in the framework and transfer nonlinear patterns from training
143 to target stations in order to predict the evolution of EWLs. Most of the target stations are directly
144 exposed to the Atlantic Ocean and located in-between the training stations (Figure 1). Those
145 stations include: 1) Montauk, NY (NOAA ID: 8510560), 2) Sandy Hook, NJ (8656483), 3) Lewes,
146 DE (8557380), 4) Ocean City, MD (8570283), 5) Kiptopeke, VA (8632200), 6) Duck, NC
147 (8656483), 7) Oregon Inlet Marina, NC (8652587), 8) USCG Station Hatteras, NC (8654467), 9)
148 Wrightsville Beach, NC (8658163), 10) Springmaid Pier, SC (8661070), 11) Charleston, SC
149 (8665530), 12) Mayport, FL (8720218), 13) Lake Worth Pier, FL (8722670), and 14) Virginia
150 Key, FL (8723214).



151

152 **Figure 1:** Location of training and target stations along the U.S. Atlantic Coast. (a) Selected
 153 training and target stations (numbered from 1 to 14) are shown with colored and black circles,
 154 respectively. (b) For each training station, percentage of water level data completeness obtained
 155 from the NOAA's Tides & Current portal. (c) Relevant hurricane's best tracks within a 60 km
 156 radius of the hurricane's landfall locations.

157 **2.2 Data availability**

158 We retrieve WL data from the National Oceanic and Atmospheric Administration (NOAA)'s
159 Tides & Currents portal (<https://tidesandcurrents.noaa.gov/map/index.html>) and complement these
160 with legacy data from the University of Hawaii Sea Level Center
161 (<https://uhslc.soest.hawaii.edu/data/?rq>), particularly for stations where WL records are not
162 available. Meteorological and wave data are obtained from the European Centre for Medium-
163 Range Weather Forecasts Reanalysis dataset (ERA5, version 5) produced by the Copernicus
164 Climate Change Service (<https://cds.climate.copernicus.eu/>). ERA5 dataset has a spatial resolution
165 of 31 km that allows for accurate representation of extreme climate events at large scale including
166 those driven by TCs (Bian et al. 2021). Specifically, we use hourly wind speed and direction at 10
167 m elevation, atmospheric pressure, sea level pressure, sea surface temperature, air temperature,
168 precipitation, wave direction, and wave height. In addition, we retrieve data from the U.S. Army
169 Corps of Engineers (USACE)'s Wave Information Studies (WIS) portal that provides consistent,
170 hourly, and long-term wave climatology along the U.S. coastlines
171 (<https://wisportal.erdc.dren.mil/#>). These aforementioned datasets have been successfully applied
172 to other NN models that predict hourly non-tidal residuals at tide stations on a global scale with
173 satisfactory results (Bruneau et al. 2020).

174 **2.3 Data processing**

175 The required data length to effectively train NN models depends on the response time of the
176 system under analysis. For coastal systems, previous studies recommend at least six years of
177 training data consisting of complete consecutive sequences (10 days) in order to achieve consistent
178 proficiency in NN models (Bruneau et al. 2020; Tiggeloven et al. 2021). Following this, we

179 conduct data quality control over the training stations and ensure that the time-series contain
180 complete data sequences to train the LSTM models. Then, we decompose the time-series data of
181 WL into seasonality, trend, predicted tides, and non-tidal residual (NTR) components using the
182 Seasonal-Trend decomposition using LOESS (STL) and Unified Tidal Analysis and Prediction
183 (UTide) packages in Python (Cleveland et al. 1990; Codiga 2011). The STL analysis, adept at
184 time-series analysis for its outlier resilience, flexible seasonal adjustment, and trend adaptability,
185 provides comprehensive insights into long-term and seasonal dynamics (Chen et al. 2020). UTide
186 employs a decision tree algorithm, a recognized method for automatically selecting the most
187 relevant constituents from 147 tidal constituents, and offers tide prediction correction for records
188 spanning up to one full (18.6-year) nodal cycle (Codiga 2011; Tiggeloven et al. 2021; Tedesco et
189 al. 2023). We consider a window size of 40 days and a time step of 3 days for time-series
190 decomposition in order to ensure that at least one full lunar cycle is covered (Figure S1,
191 Supplementary material), including both spring and neap tides and the independence of large storm
192 events by selecting the maximum NTR on a stepped basis (Serafin and Ruggiero 2014; Rashid et
193 al. 2024; Moftakhari et al. 2024). The time-series decomposition aids to improve deep learning by
194 distinguishing clear, recurring patterns from irregular variations; thereby refining the models'
195 ability to learn from the data and enhancing the accuracy of their predictions (Parker et al. 2023).

196 In addition to the WL components, we extract meteorological and wave data from the closest
197 grid pixel of ERA-5 dataset to tide-gage stations. For this, we calculate the minimum square
198 difference between the latitudes and longitudes of the data points and the specified location, that
199 is, within a radius of 15.5 km. Next, we use the time-series of WL components, meteorological,
200 and wave data as relevant input features to the LSTM-SAM framework in order to predict the
201 target variable (e.g., EWLs and their evolution over time). Both input and target variables are first

202 scaled using the “minmaxscaler” function from the sklearn library in Python. This technique
203 normalizes the range of multisource data and ensures that all features have an identical scale,
204 typically between 0 and 1 (de Amorim et al. 2023). Also, we create a RNN dataset function that
205 preprocesses data for LSTM by taking normalized input features (X_{norm}), corresponding target
206 values (y_{scaled}), and a specified look-back period to construct a dataset suitable for sequence
207 prediction. Moreover, we consider two look-backs of 6 and 24 hours to train the LSTM and
208 evaluate the effects of different time steps on the model’s prediction performance, i.e., the number
209 of previous time steps in hours used to predict the next time step.

210 X_{norm} and their corresponding y_{scaled} are sequentially split into a training size of 80% and
211 a testing size of 20%. Finally, we capture the evolution of EWLs in the training and testing datasets
212 by focusing our analysis on historic hurricane events and Nor'easter winter storms within a 7-day
213 window centered around the peak WL. Results of a sensitivity analysis show that longer time-
214 windows favor model’s performance metrics due to multiple non-extreme WLs being accounted
215 for, whereas shorter ones could not effectively capture the evolution of EWLs across all stations
216 as they focused more on the peak WL.

217 **2.4 Model architecture**

218 The LSTM-SAM framework consists of bidirectional LSTM (Bi-LSTM) models that are
219 garnering significant interest within the domain of WL prediction (Bai and Xu 2021; Fang et al.
220 2021; Zhang et al. 2022). Unlike traditional LSTM models that only rely on previous timesteps,
221 the advantage of Bi-LSTM models is that input sequences are processed in both forward and
222 backward directions (Equations 1 to 13 in the Supplementary material). The reader is referred to
223 the study of Ahmed et al. (2022) for a more detailed explanation of Bi-LSTM. This dual viewpoint

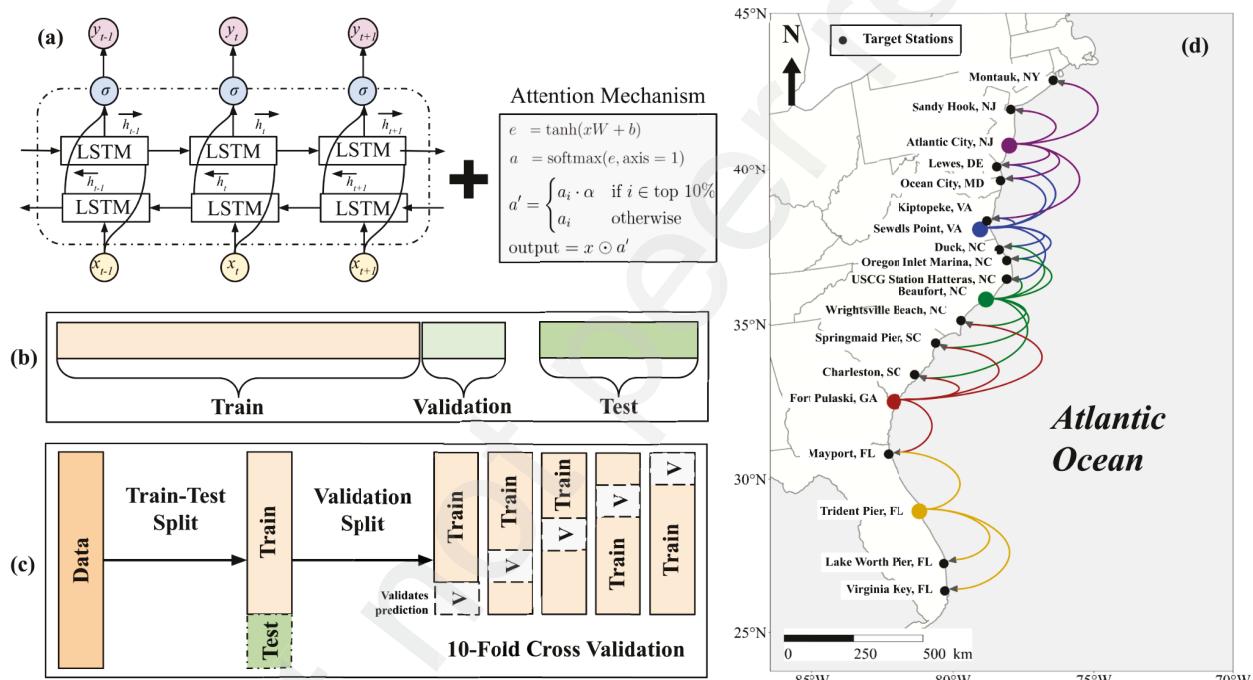
224 helps identify and learn features that may not be apparent when the sequence is analyzed in only
225 one direction.

226 **2.4.1 Bidirectional LSTM network models**

227 The model architecture consists of three LSTM units that are set to vary in intervals of 32 since
228 higher units tend to increase computational complexity (Figure 2a). We consider a “L2”
229 regularization method to prevent overfitting by penalizing large weights. Also, we include dropout
230 rates in the model architecture to prevent overfitting by randomly disabling a subset of neurons
231 during the training process, thereby allowing the LSTM network to develop a more generalized
232 understanding of the data and improve its performance on new and/or unseen data. The LSTM
233 units use the “tanh” and hard “sigmoid” recurrent activation functions. In addition, we add a
234 standard dense (fully connected) layer to output the final prediction. The loss functions consist of
235 both mean absolute error (MAE) and mean squared error (MSE) for different variants of the
236 models with the “Adam” optimizer as suggested in similar WL prediction studies (Huang et al.
237 2020). We reserve 30% of the training data for validation of the model’s learning ability during
238 the training process. An early stopping callback is also employed to monitor the validation loss,
239 stop training if no improvement is observed for five consecutive epochs, and ultimately prevent
240 overfitting and/or unnecessary computations.

241 We consider two training strategies in the LSTM algorithm: (i) train-test (TT) split, and (ii)
242 time-series cross-validation (CV) split. The first method involves a single split into training (80%)
243 and testing sets (20%) whereas the second method includes multiple training and testing sets
244 created sequentially for a more comprehensive evaluation of the model’s performance across
245 different periods (Figure 2b and 2c). Unlike traditional CV strategies, here CV fold does not shuffle
246 the data and therefore keeps the time sequence invariant (Kingphai and Moshfeghi 2022). We

247 consider 10-fold CV to check the model's performance and potentially improve the prediction
 248 accuracy. Moreover, the LSTM model is trained using the training set for that specific split for
 249 each iteration of the loop. As the loop progresses, the size of the training set increases whereas the
 250 validation set consists of data points that come after the training set in time. As a result, the training
 251 and validation process involves learning from past data and validating the model's performance on
 252 unseen future data, respectively. Once all splits are processed, the final model is trained using the
 253 entire dataset.



254
 255 **Figure 2.** Schematic of model architecture, training, validation, and transfer learning approach. (a)
 256 Bidirectional LSTM network model with an attention layer mechanism to improve pattern
 257 recognition. The model employs two data training approaches: (b) train-test split, and (c) time-
 258 series cross-validation for model development. (d) Transfer learning approach to predict extreme
 259 water level evolution at target stations (black circles) using models developed in the closest
 260 training stations (colored circles).

261 **2.4.2 Hyperparameter tuning**

262 We conduct hyperparameter tuning to identify optimal values of LSTM units, dropout rate,
263 and learning rate within specified ranges to train the models (Table 1). We set the tuner search to
264 a maximum of 300 trials, after which the best hyperparameters are used to train the models. The
265 model architecture relies on a Bayesian optimization technique for hyperparameter tuning that
266 inherently functions in a sequential manner and leverages data from previous evaluations to inform
267 subsequent runs (Wang et al. 2023). Such technique efficiently balances the exploration of new
268 areas in the hyperparameter space with an emphasis on known suitable regions. This is particularly
269 useful when each training iteration is computationally intensive since the aforementioned
270 optimization technique can identify optimal hyperparameters with less time than methods like grid
271 or random search (Marco et al. 2022). Additionally, its capacity to handle high-dimensional
272 hyperparameter spaces and integrate prior knowledge about potential hyperparameters makes it a
273 versatile choice (Bischl et al. 2023). Its proven success in real-world applications and its efficiency
274 in finding robust hyperparameters with limited evaluations position it as a top choice for many
275 practitioners (Wang et al. 2023).

276 Hyperparameters, identified through a rigorous tuning (or calibration) process on site-specific
277 training data, tend to yield models that perform optimally within particular training domains.
278 However, these models may not necessarily exhibit the same level of effectiveness across other
279 target domains, even if both domains share similar morphological and hydrodynamic
280 characteristics. We address this limitation by focusing on the range of values considered for
281 hyperparameters. First, we define the values for specific parameters such as batch sizes (32, 64,
282 128, and 256), look-back times (6 and 24), loss functions (MAE and MSE), and data training
283 strategies (TT and CV), while allowing other parameters to be determined through hyperparameter

284 tuning. We then ensure that combinations of options from the aforementioned parameters occur
285 precisely once, which in turn facilitates the creation of distinctive models with a unique set of
286 hyperparameters across the five training stations (Figure 2).

287 Hence, this tuning process produces a set of suitable Bi-LSTM models with comparable
288 performance for a given training station. We evaluate model's performance using several metrics
289 that are recommended for models predicting WL dynamics (Abbaszadeh et al. 2020; Lee et al.
290 2021; Muñoz et al. 2022a). Those include the coefficient of determination (R^2), Mean Bias Error
291 (MBE), Root Mean Square Error (RMSE), Kling-Gupta Efficiency (KGE) (Gupta et al., 2009),
292 and Nash-Sutcliffe Efficiency (NSE) (Nash and Sutcliffe, 1970). Based on the evaluation metrics,
293 there might be a set of “transferable” LSTM models among the suitable ones from the training
294 station for which inherent pattern recognition capabilities would be adequate for the target stations
295 (Section 2.4.4). However, an increase in tuning trials poses a risk of overfitting, where models
296 become excessively tailored to the training data and lose their predictive ability on new datasets.
297 Here, we specify an appropriate number of tuning trials while employing different combinations
298 of hyperparameters to generate a spectrum of suitable models for a training domain (Table 1).
299 Specifically, we develop a total of 32 models at each training.

300 **Table 1:** Range of values considered for hyperparameter tuning.

Hyperparameter (3 LSTM units)	Range of Tested Values
LSTM Units	32 - 512 (step of 32)
Dropout Rate	0.10- 0.50 (step of 0.1)
Activation	Tanh, Sigmoid
L2 Regularization	1e-6 - 1e-3 (log sampling)
Learning Rate (Adam Optimizer)	1e-4 - 1e-2 (log sampling)
Maximum number of trials	300

Batch Size (b)	32 - 256 (step of 32)
Loss Function	MAE, MSE
Look-back time (h)	6, 24
Epochs	500 (with early stopping)
Validation Split	30% of training data

301 **2.4.3 Attention layer mechanism**

302 Attention layers are used to address inherent limitations of conventional RNNs (and LSTM)
 303 such as the tendency to lose information from earlier segments of extended sequences and
 304 difficulties to train models in areas of sharp and extreme changes (Rithani et al. 2023). These layers
 305 scan through the data, identify key features, and increase their influence in the training process.
 306 Here, we incorporate an attention layer in the model architecture for better model generalization
 307 during the training process. Specifically, the attention layer computes attention scores for each
 308 time step using a weight matrix (Glorot) and bias (zeros) initialization (Equations 14 to 15 in the
 309 Appendix). In addition, we customize this layer using a factor that amplifies the top 10% of the
 310 attention scores (Equation 16). This factor allows the model to focus more on crucial parts of the
 311 sequence, which could be abrupt changes of high or low levels in the data.

312 The choice of the Glorot initializer for the weight matrix in the attention layer is appropriate
 313 due to the use of tanh and sigmoid activation functions in the LSTM units (Glorot and Bengio
 314 2010; Evangelista and Giusti 2021). The initializer keeps the scale of the gradients approximately
 315 the same in all layers of the LSTM network. Starting with zero biases ensures that all neurons in a
 316 layer initially produce outputs of roughly the same magnitude, which can be a good starting point
 317 for symmetric activation functions like tanh.

318 **2.4.4 Transfer learning**

319 We conduct TL to predict the evolution of EWLs at selected target stations by leveraging
320 “gained knowledge” from the closest training stations (Figure 2d). Such knowledge includes
321 hidden sequential patterns and nonlinear associations among input and target data features that are
322 stored as model weights (Muñoz et al. 2021; Zhao et al. 2021). Note that most of the target stations
323 are located in-between two training stations; except Montauk, NY and Sandy Hook, NJ as well as
324 Lake Worth Pier, FL and Virginia Key, FL that are close to a single training station. Here, the TL
325 approach consists in training, tuning, and validating Bi-LSTM models at each training station and
326 subsequently saving the corresponding model weights. We leverage all available Bi-LSTM models
327 (e.g., 32 models) and transfer them to the target stations in order to predict the evolution of EWLs
328 within the predefined 7-day window (Section 2.3). Among these models, we identify
329 “transferable” Bi-LSTM models based on the criteria that both KGE and NSE are above a
330 threshold value of 0.70 at the target stations. This threshold ensures that each transferable model
331 adequately accounts for the magnitude and timing of EWLs while also keeping its inherent pattern
332 recognition capabilities on new input data that has not yet been observed at the target stations (e.g.,
333 those associated with future extreme events). Lastly, we assess the performance of transferable
334 models at the target stations with identical evaluation metrics including R^2 , MBE, RMSE, KGE,
335 and NSE.

336 **3. Results**

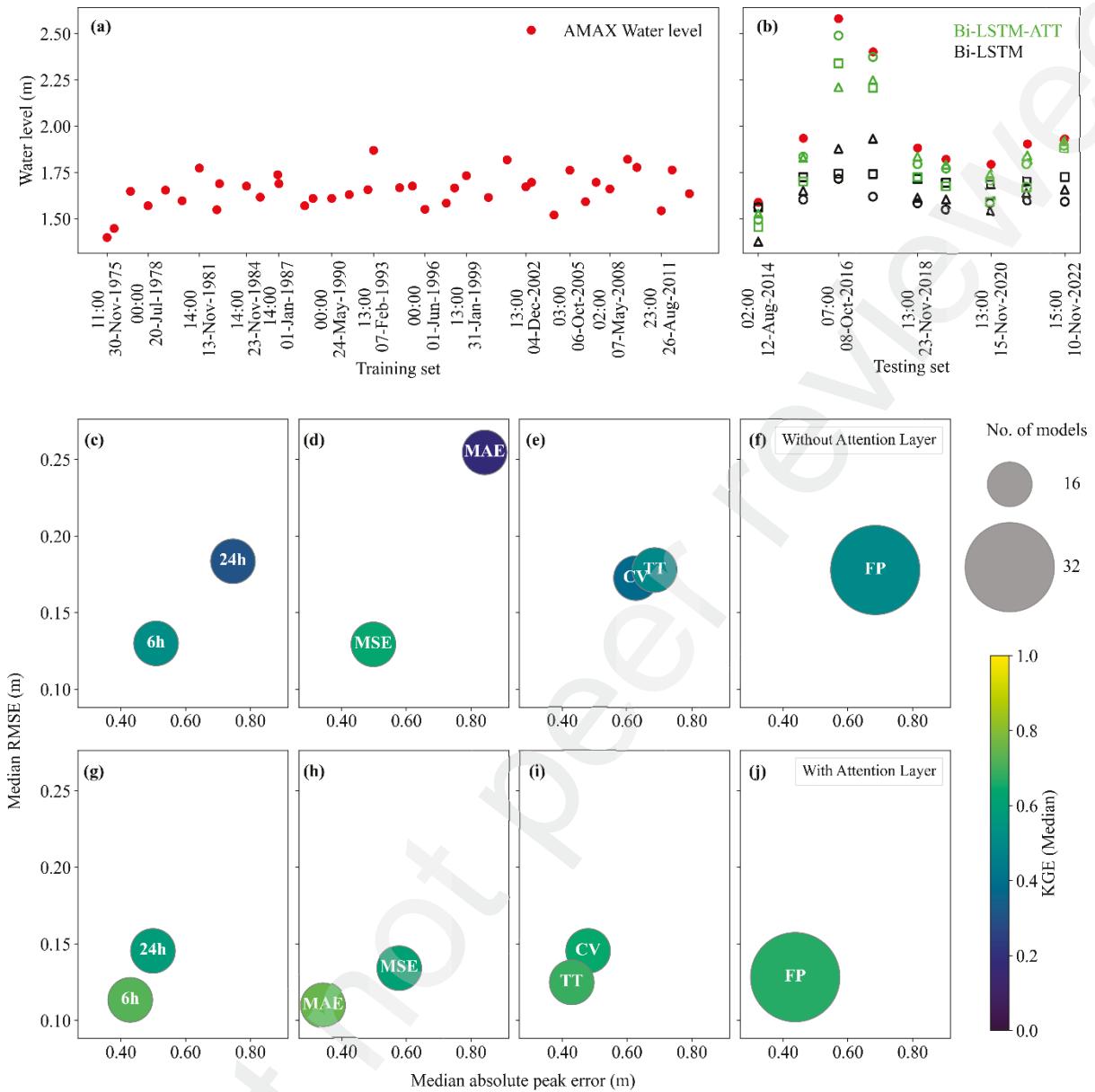
337 **3.1 Assessment of bidirectional LSTM network models**

338 We first assess the performance of Bi-LSTM models at the training stations with and without
339 the attention layer mechanism incorporated in the model architecture. For this, we consider the

340 models' ability to capture: (i) peak and timing of EWLs, and (ii) evolution of EWLs for selected
341 historic TCs (hurricanes) and extra-TCs (Nor'easter winter storms).

342 **3.1.1 Extreme water levels**

343 For convenience, we will focus on model predictions obtained from Fort Pulaski (FP), GA
344 since this training station has been reporting complete consecutive hourly data from 1975 to
345 present (Figure 3). Predictions of EWLs and associated RMSE, average peak error, and KGE
346 metrics suggest that Bi-LSTM models with the attention layer (e.g., hereinafter referred to as Bi-
347 LSTM-ATT) can capture the magnitude and timing of peak WLs with a higher accuracy than those
348 without this layer (Figure 3b). Note that this includes the two most EWL events contained in the
349 testing set (e.g., Hurricane Matthew (2016) and Hurricane Irma (2017)). More than half of Bi-
350 LSTM models achieve a median RMSE, absolute peak error, and mean bias of 0.18 m, 0.68 m,
351 and -0.11 m, respectively (Figure 3c to 3f and Table S1). Also, these models achieve very low to
352 moderate performances with a median KGE and NSE of 0.50 and 0.15, respectively (Table S1).
353 In contrast, the models' performance substantially improves after integrating the attention layer
354 mechanism in the model architecture. In that regard, half of Bi-LSTM-ATT models show a
355 reduction in the median RMSE, absolute peak error, and mean bias by 27% (0.13 m), 36% (0.44
356 m), and 55% (-0.05 m) with respect to the Bi-LSTM models only (Figure 3g to 3j and Table S2).
357 Also, these models achieve a median KGE and NSE of 0.67 and 0.56, respectively (Table S2).



358

359 **Figure 3.** Assessment of models' performance in the testing set at Fort Pulaski (FP), GA. (a, b)
 360 Annual maximum water levels in the training set (80%), and testing set (20%) in addition to
 361 predictions with/out the attention layer mechanism. The Bi-LSTM and Bi-LSTM-ATT models are
 362 categorized into (c, g) look-back times of 6 h and 24 h, (d, h) loss functions focused on Maximum
 363 Absolute Error (MAE) and Mean Square Error, (MSE) (e, i) Train-Test (TT) and Cross Validation
 364 (CV) fold strategies, and (f, j) a collective unit to compare their overall performance.

365

366 At the remaining training stations, we evaluate predictions of EWLS derived from Bi-LSTM-
367 ATT models only (Table S2). The models at Atlantic City (AC), NJ station achieve satisfactory
368 model performance with a median KGE of 0.84 despite the relatively low median NSE of 0.36.
369 The median RMSE and mean bias in this training station are 0.07 and 0 m, respectively. Training
370 station Sewells Point (SW), VA has the best performing models with a relatively high median
371 KGE and NSE of 0.92 and 0.80, respectively. These models achieve a median RMSE and mean
372 bias of 0.05 m and -0.02 m, respectively. Similarly, results at Beaufort (BF), NC station perform
373 satisfactorily with median KGE and NSE of 0.86 and 0.69, respectively. Also, this training station
374 shows a median RMSE and mean bias of 0.05 m and -0.01 m, respectively. Lastly, the models of
375 training station Trident Pier (TD), FL achieve a moderate to satisfactory performance with median
376 KGE and NSE of 0.78 and 0.51, respectively. The models show a median RMSE and mean bias
377 of -0.07 m and -0.03 m, respectively.

378 **3.1.2 Evolution and peak of extreme water levels**

379 Next, we assess the performance of Bi-LSTM-ATT models to predict the evolution of EWLS
380 including the magnitude and timing of the peaks (Figure 4). The top-two Bi-LSTM-ATT models
381 of AC predict the evolution of EWLS associated with a Nor'easter storm with relatively high
382 accuracy. These models achieve KGE of 0.89 and 0.82 and NSE of 0.96 and 0.94 (Figure 4a). In
383 contrast, the top-two models of SW capture the evolution of EWLS and almost perfectly match the
384 magnitude of the peak associated with Hurricane Joaquin (2015) (Figure 4b). These models
385 achieve KGE of 0.98 and 0.98 and NSE of 0.99 to 0.99. Similarly, the top-two models of BF
386 achieve satisfactory KGE of 0.96 and 0.95 and NSE of 0.98 and 0.98. These models capture the
387 evolution and almost perfectly match the magnitude of the peak associated with Hurricane Dorian
388 (2019) (Figure 4c). The top-two models of FP can predict the evolution of EWLS but underpredict

389 the peak associated with Hurricane Matthew (2016) (Figure 4d and 4f). These models achieve
390 KGE of 0.87 and 0.85 and NSE of 0.91 and 0.88.

391 Lastly, the top-two models of TD capture the evolution of EWLs even though the second model
392 overpredicts the magnitude of the peak associated with Hurricane Nicole (2022) (Figure 4e). These
393 models achieve KGE of 0.72 and 0.70 and NSE of 0.84 and 0.93. Regarding the peak time
394 difference (Figure 4f), all Bi-LSTM-ATT models developed for AC and BF stations show a 1-h
395 lead difference and a perfect match with respect to the observed peak, respectively. Six models of
396 SW station show a 1-h lag difference whereas two models of TP station show a 1-h lead difference
397 with respect to the observed peak. Thirteen and ten models developed for FP stations show a 1-h
398 lag difference and 1-h lead difference with respect to the observed peak, respectively. Overall, the
399 time difference between observed and predicted peak WL is ± 1 h for all trained models.



400

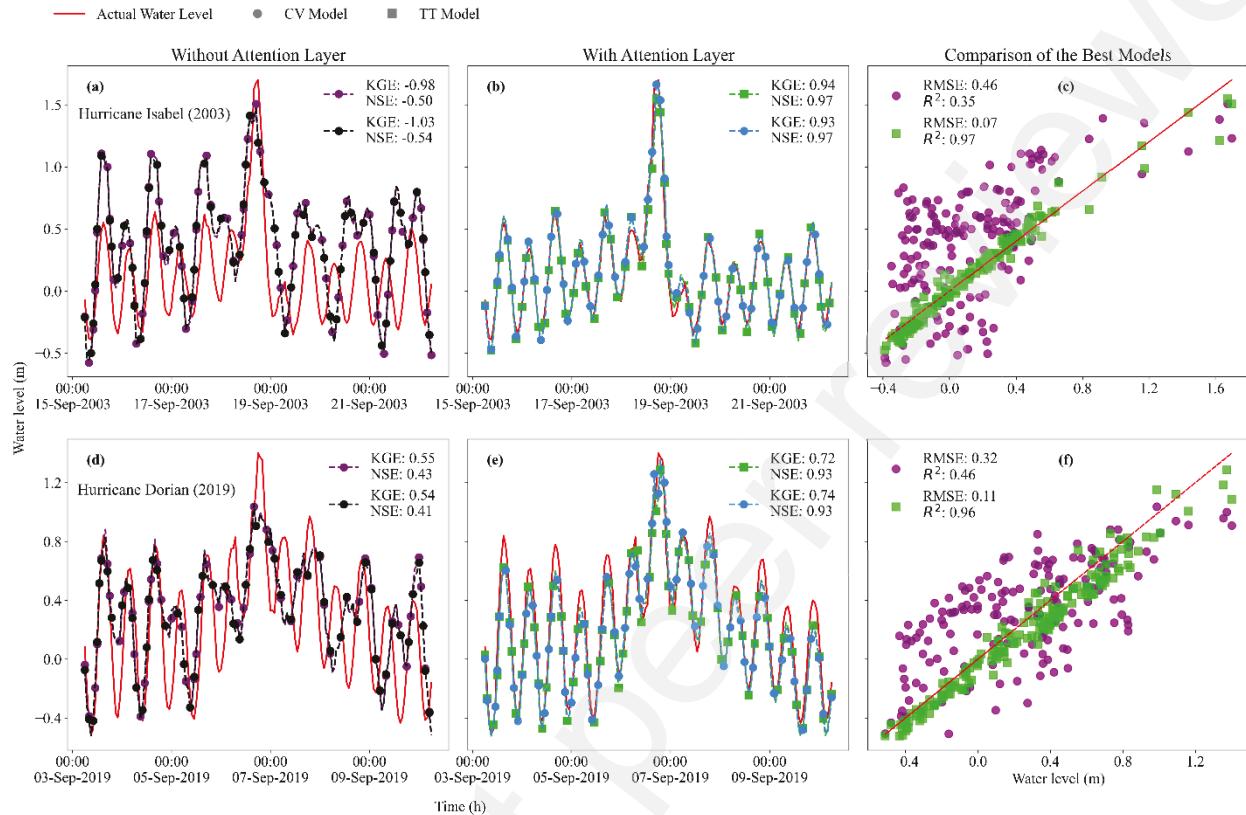
401 **Figure 4.** Assessment of models' performance to predict the evolution and peak of extreme water
 402 levels in the testing set. Predictions of the top-two Bi-LSTM-ATT models from the Train Test
 403 (TT) and/or Cross Validation (CV) split strategies at (a) Atlantic City, NJ (AC), (b) Sewells Point,
 404 VA (SW), (c) Beaufort, NC (BF), (d) Fort Pulaski, GA (FP), and (e) Trident Pier, FL (TD). (f)
 405 Observed and predicted peak time differences among the five training stations and 32 Bi-LSTM-
 406 ATT models.

407 **3.2 Assessment of transfer learning approach**

408 After evaluating the models' performance, we proceed with the assessment of transferable
409 models from the closest training stations to target stations. For example, we transfer the pretrained
410 Bi-LSTM and Bi-LSTM-ATT models and their associated model weights from Sewells Point, VA
411 to the target station at Duck, NC (Figure 2d). Since most of the extreme events at Sewells Point,
412 VA, including Hurricane Isabel (2003), Irene (2011) and Sandy (2012), are in the training set, the
413 performance of all Bi-LSTM models are satisfactory, with metrics comparable to Bi-LSTM-ATT
414 models (Table S3). Here, the goal is to predict the evolution of EWLs for relevant extreme events
415 such as Hurricane Isabel (2003) and Dorian (2019) (Figure 5). Based on the threshold value of
416 0.70 (Section 2.4.4), there are no transferable Bi-LSTM models that can predict the evolution of
417 both storm events. For Hurricane Isabel, the two best models achieve low KGE of -0.98 and -1.03
418 and NSE of -0.50 and -0.54 (Figure 5a). Although these two models show a better performance
419 for Hurricane Dorian, they still achieve moderate KGE of 0.55 and 0.54 and NSE of 0.43 and 0.41
420 (Figure 5d). In contrast, the top-two transferable Bi-LSTM-ATT models achieve high KGE of 0.94
421 and 0.93 and NSE of 0.97 and 0.97 when predicting EWLs triggered by Hurricane Isabel (Figure
422 5b). For Hurricane Dorian, the top-two transferable Bi-LSTM-ATT models achieve a relatively
423 high KGE of 0.72 and 0.74 as well as NSE of 0.93 and 0.93 (Figure 5e).

424 Furthermore, we assess the models' performance in terms of R^2 and RMSE for both extreme
425 events and compare model predictions from the best transferable Bi-LSTM and Bi-LSTM-ATT
426 models using a one-to-one plot (Figure 5c and 5f). Bi-LSTM models have poor generalization of
427 EWLs with low predictive accuracy ($R^2 < 0.50$) and high error (RMSE > 0.30). In contrast, the Bi-
428 LSTM-ATT models can predict the evolution of EWLs with a high predictive accuracy ($R^2 > 0.95$)
429 and low error within an acceptable range (RMSE < 0.15 m). In general, RMSEs below 0.20 m are

430 desirable for hurricane storm surge modeling (Muis et al. 2016). Following this analysis, we
 431 hereafter present the results derived from Bi-LSTM-ATT models only.



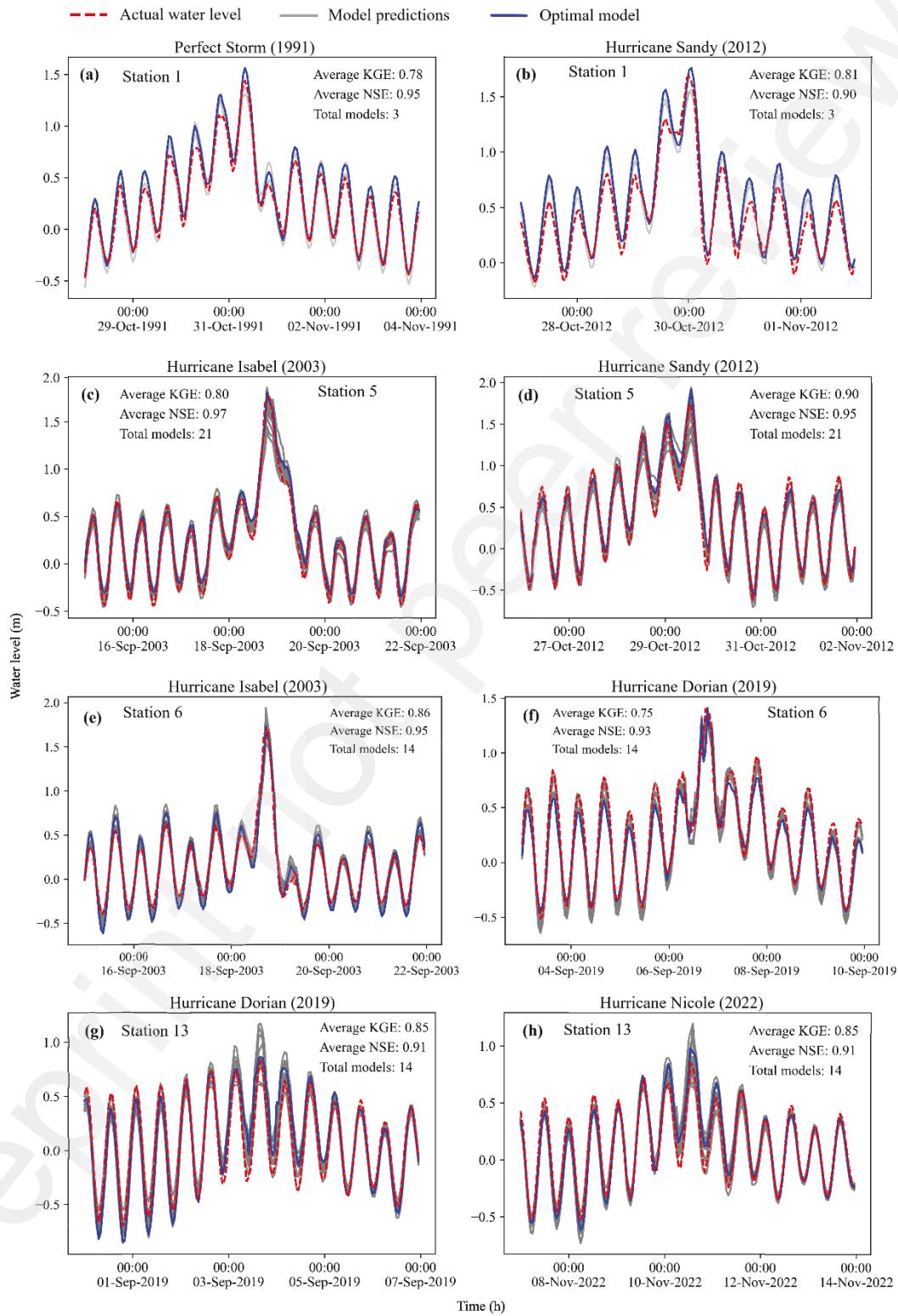
432
 433 **Figure 5.** Assessment of transfer learning approach from Sewells Point, VA (training station) to
 434 Duck, NC (target station). Prediction of extreme water level evolution using the top-two (a, d) Bi-
 435 LSTM, and (b, e) Bi-LSTM-ATT models for Hurricane Isabel (2003) and Dorian (2019). (c, f)
 436 One-to-one comparison of top-two model predictive capabilities based on the aforementioned
 437 hurricane events.

438 **3.3 LSTM-SAM framework**

439 Once the Bi-LSTM-ATT models and TL approach have been assessed at selected training and
440 target stations (Section 3.1 and 3.2), we introduce the LSTM-SAM framework to accurately
441 predict the evolution of EWLs at target stations in the U.S. Atlantic Coast (Figure 2d). As
442 mentioned before, we set a time window of 7-day centered around the peak to characterize the
443 evolution of EWLs at the target stations and leverage the 32 Bi-LSTM-ATT models developed at
444 each training station. Note that instead of considering the top-two transferable Bi-LSTM models
445 at the target stations (Figure 5), the LSTM-SAM framework identifies a set of transferable models
446 for which KGE and NSE are above 0.70 when evaluated with respect to TCs or extra-TCs events
447 (Figure 6). For practical flood prediction purposes and decision-making support, Bi-LSTM-ATT
448 models achieving the smallest peak deviation among the transferable models are considered as the
449 optimal ones at each target station (e.g., models with the closest prediction to the peak WL).

450 There are 3 transferable models from training station AC to the target station in Montauk, NY
451 that accurately predict EWL evolution of The Perfect Storm (1991) and Hurricane Sandy (2012)
452 (Figure 6a and 6b). These models achieve satisfactory performances such as an average KGE and
453 NSE above 0.75. Likewise, there are 21 transferable models from AC and SW that predict EWL
454 evolution of Hurricane Isabel (2003) and Sandy (2012) at the target station in Kiptopeke, VA
455 (Figure 6c and 6d). These models achieve an average KGE and NSE above 0.80. At the target
456 station located in Duck, NC, LSTM-SAM identifies 14 transferable models from SW and BF that
457 predict EWL evolution of Hurricane Isabel (2003) and Dorian (2019) (Figure 6e and 6f). These
458 models show satisfactory performances with average KGE and NSE above 0.75. Similarly, the
459 framework identifies 14 transferable models from TD to the target station in Lake Worth Pier, FL

460 that accurately predict EWL evolution of both Hurricane Dorian (2019) and Nicole (2022) (Figure
 461 6g and 6h). The models achieve average KGE and NSE above 0.85.

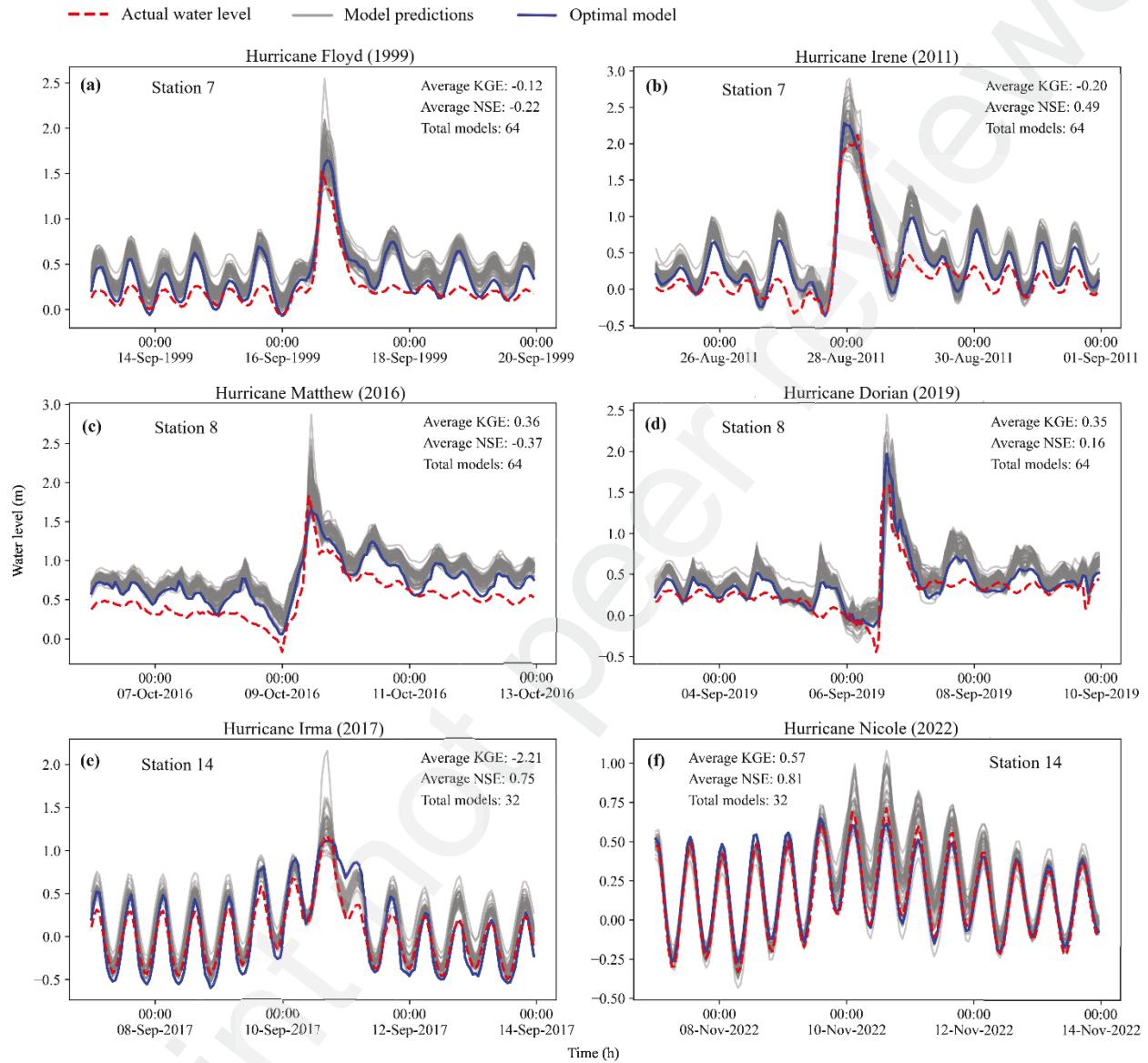


462

463 **Figure 6.** Extreme water level prediction for relevant hurricanes and Nor'easter winter storms in
464 the U.S. Atlantic Coast. Each row panel shows two extreme events at the target stations, the
465 number of transferable models, and their average performance in terms of KGE and NSE. These
466 stations are (a, b) Montauk, NY, (c, d) Kiptopeke, VA, (e, f) Duck, NC, and (g, h) Lake Worth
467 Pier, FL. The dashed red, blue, and gray lines represent observed water levels, optimal model, and
468 water level predictions from all transferable models.

469 Results of the remaining target stations show average KGE and NSE ranging from 0.70 to 0.99
470 (Figure S2, Supplementary material). It is worth noting that we leverage WL data from the training
471 station AC to predict the complete evolution of Hurricane Sandy (2012) at the target station in
472 Sandy Hook, NJ (Figure S2b). This demonstrates the transfer model capability to predict EWL
473 evolution even when tide-gauges fail or become inoperative. However, there are three target
474 stations for which no Bi-LSTM-ATT models are completely transferable given the criteria that
475 both NSE and KGE should be above 0.70 within the 7-day window (Section 2.4.4). Specifically,
476 the LSTM-SAM framework does not identify any transferable models from SW and BF stations
477 to the target station in Oregon Inlet, NC that can accurately capture the evolution of EWLS of
478 Hurricane Floyd (1999) and Irene (2011) (Figure 7a and 7b). Similarly, the framework does not
479 identify transferable models from SW and BF stations to USCG Station Hatteras, NC for Hurricane
480 Matthew (2016) and Dorian (2019) (Figure 7c and 7d). Lastly, there are no transferable models
481 from TD station to Virginia Key, Florida for Hurricane Irma (2017) and Nicole (2022) (Figure 7e
482 and 7f). However, note that some Bi-LSTM-ATT models can effectively capture the peak WL
483 within a shorter time window centered around the peak (e.g., ± 1 day); hence, the relatively low
484 KGE and NSE at the target stations are explained by an overprediction of WLs occurring before
485 and after the peak WL. Therefore, the LSTM-SAM framework considers the model with the

486 highest KGE and NSE as the optimal models for those three target stations (Figure S2,
 487 Supplementary material).



488

489 **Figure 7.** Extreme water level prediction for relevant hurricane events in the U.S. Atlantic Coast.
 490 Each panel shows two extreme events at the target stations, total number of models, and their
 491 average performance in terms of KGE and NSE. These stations are (a, b) Oregon Inlet Marina,
 492 NC, (c, d) USCG Station Hatteras, and (e, f) Virginia Key, FL. The dashed red, blue, and gray

493 lines represent observed water levels, optimal model, and water level predictions from all available
494 Bi-LSTM-ATT models developed at the corresponding training stations.

495 **4. Discussion**

496 Physically-based models can accurately predict EWLs; however, they are site-specific and not
497 transferable to other domains, even with similar characteristics, due to their need for detailed
498 topographic and bathymetric (topobathy) data (Santiago-Collazo et al. 2019; Bates 2022b). A
499 feasible alternative to overcome this limitation consists in leveraging *state-of-the-art* deep learning
500 models, such as LSTM networks, given their effectiveness for learning dynamic and/or sequential
501 data including nonlinear interactions and hidden patterns from hydrometeorological input features
502 (Tedesco et al. 2023). Conveniently, LSTM networks enable time-series prediction even in
503 absence of geographical information or catchment characteristics that may remain quasi-invariant
504 for a relatively long time (e.g., average slope, length, width, catchment size, among other input
505 features). Although model transferability is still challenging (Zhao et al. 2021; Kratzert et al.
506 2024), adequate feature engineering procedures (Merizalde et al. 2023) and improvement of LSTM
507 models' architecture (e.g., attention layer mechanisms) can increase the effectiveness of TL
508 approaches over untrained (target) sites having similar hydrodynamic and morphologic features to
509 the corresponding training areas. This in turn will help advance flood prediction efforts in large
510 scale domains with high accuracy and less computational time (Ding et al. 2020; Li et al. 2021;
511 Nearing et al. 2024). Here we present one of the first applications of TL to EWL prediction in
512 coastal domains, and to our knowledge, the first outside the hydrological context.

513 **4.1 Application of the LSTM-SAM framework**

514 The proposed LSTM-SAM framework is trained on sequential WL data from tide-gage stations
515 and follows a time-series decomposition in order to obtain hidden patterns and nonlinear
516 associations such as seasonality, trend, harmonic tides, and NTR components (Section 2.3).
517 Moreover, these input features are complemented with hydrometeorological data from reanalysis
518 datasets. As a result, the framework is capable of applying learned knowledge to selected target
519 stations and predicting the evolution of EWLs (Section 2.4.1). We ensure that Bi-LSTM networks
520 are robustly calibrated through hyperparameter tuning, which resulted in 32 distinctive models
521 across five training stations and a unique set of hyperparameters per model (Section 2.4.2). We
522 argue that this method preserves inherent pattern recognition capabilities for each model (e.g.,
523 model weights) and increases the chances for identifying effective transferable models to target
524 stations.

525 Conventional Bi-LSTM models show poor performance to predict EWLs at the training station
526 particularly when there exist more extreme events in the test set (Figure 3a and 3b). This is partly
527 due to a sequential training (80%) and test (20%) split given that the frequency and magnitude of
528 EWLs is expected to increase by the end of the century (Santiago-Collazo et al. 2019; Bloemendaal
529 et al. 2022; Boumis et al. 2023). In contrast, leveraging randomly selected subsamples in each
530 batch during the calibration process facilitates a quicker model convergence (De la Fuente et al.
531 2024) and prevents anomalies in the training data. Nevertheless, training Bi-LSTM models on
532 sequential data ensures that temporal relationships are fully considered in the learning process.
533 Since these models are in general most effective in capturing changing trends of cyclic patterns
534 (Wang et al. 2022), they do ignore some nuanced information of rare and abrupt changes during
535 the training process. Therefore, there is a higher chance of incorrect estimation of equally rare but

536 more extreme data values in testing sets which limits the model's ability to accurately predict
537 EWLs.

538 Following this, we introduce a custom attention layer in the Bi-LSTM architecture (Section
539 2.4.3), which significantly improves the models' ability to capture TCs and extra-TC events
540 (Figure 3b). Since hyperparameter tuning is predominantly inclined toward identifying the optimal
541 set of model parameters that result in the most accurate EWL prediction in the training set, there
542 is no guarantee that this optimal set leads to improved performance on unseen (test) sets (Tran et
543 al. 2020). Therefore, by amplifying the top 10% of the attention scores, the model's ability to
544 internally focus on the most relevant time steps is substantially improved and leads to more
545 accurate EWL predictions. Moreover, the attention layer mechanism allows the models to perform
546 identical operations consistently beyond the training set and therefore generalizing unseen data
547 with similar characteristics in the test set. Since Bi-LSTM-ATT outperforms conventional Bi-
548 LSTM models (Figure 3f to 3h), the proposed LSTM-SAM framework demonstrates the ability to
549 effectively predict the evolution of EWLs even when higher WLs attributed to more frequent TCs
550 and extra-TCs are expected.

551 We also observe comparable performance across all metrics for all Bi-LSTM-ATT models
552 developed on the specified options of lookbacks, batch sizes, loss functions, and data training
553 strategy (Table 2). It is worth noting that a higher batch size of 128 sample points shows the highest
554 performance of all models in terms of KGE and NSE, suggesting that they capture both peak and
555 timing of EWLs with higher accuracy. However, more than half of the models developed on a
556 smaller batch size of 32 sample points have very low NSE scores, which tend toward inaccurate
557 timing of EWLs' predictions. In addition, lower performance for models developed with 24-hour
558 lookbacks compared to 6-hour lookbacks suggests that considering more previous time steps does

559 not necessarily improve the model's predictive accuracy. Furthermore, the identical performance
560 for both data training strategies as well as the loss functions suggests that either of these options
561 are suitable to train the models.

562 Although most of the trained Bi-LSTM-ATT models can predict EWLs at the nearby target
563 station, we observe that models with KGE and NSE above 0.70 demonstrate robust generalization
564 capabilities of EWL evolution from onset, peak, to dissipation (Figure 4). NSE show a higher
565 accuracy for predictions with correct timing despite the over- or underprediction of WLs (Figure
566 7e and 7f). In addition, an extended window size renders minor discrepancies in the timing of peak
567 WLs such as 1-h lead or 1-h lag negligible in the overall prediction performance. On the other
568 hand, KGE improve when the magnitude of predictions closely aligns with actual events, even if
569 the prediction timing is inconsistent with actual WL observations (Figure 7c and 7d).

570 The low performance at target stations located at Oregon Inlet Marina, NC (NOAA ID:
571 8652587) and USCG Station Hatteras, NC (NOAA ID: 8654467) might be attributed to the
572 geographic location of the tide-gages (Figure 2d). Unlike target stations that are directly exposed
573 to the Atlantic Ocean, these stations are located behind Bodie's and Cape Hatteras' islands of the
574 Outer Banks barrier island chain. Coastal areas surrounding the target stations experience about 1
575 m of mean tidal range on the ocean side and 0.3 m behind the island (Velasquez-Montoya et al.
576 2020). In addition, these areas benefit from vast coastal wetlands and protection infrastructure such
577 as the Herbert C. Bonner bridge that alters tidal dynamics and attenuates storm surges and waves
578 (Velasquez-Montoya et al. 2021, 2022). Consequently, these site-specific conditions significantly
579 differ from those of nearby training stations including Sewells Point, VA (NOAA ID: 8638610)
580 and Beaufort, NC (NOAA ID: 8656483).

581 Nevertheless, Bi-LSTM-ATT models have the potential to capture the peak WL which is
582 crucial for supporting flood emergency response and decision-making. In fact, the optimal models
583 for those target stations correctly capture the peaks of Hurricane Floyd (1999), Irene (2011),
584 Matthew (2016), and Dorian (2019) (Figure 7a to 7d). Similarly, the evolution of EWLs at Virginia
585 Key, Florida (NOAA ID: 8723214) is overestimated (Figure 3e and 3f). Nevertheless, the models
586 can effectively capture the peaks of Hurricane Irma (2017) and Nicole (2022). The tide-gage is
587 located at the entrance of Biscayne Bay close to the Bear Cut bridge, which is most likely
588 responsible for WL being less representative of extreme events. The relatively lower water surface
589 elevation of this station during Hurricane Irma compared to Trident Pier, FL (NOAA ID: 8721604)
590 has also been noted in another study (Alarcon et al. 2022).

591 **4.2 Limitations and future work**

592 The waves and atmospheric variables obtained from ERA5 have a spatial resolution of
593 approximately 31 km. This level of resolution may not correctly account for local variability; hence
594 higher resolution data might improve the performance of LSTM-SAM framework. There are
595 instances where wave components for target stations like Lewes, DE, and Charleston, SC needed
596 to be complemented with those derived from the WIS portal data. As a result, some errors could
597 have been introduced in the input features, reducing the accuracy of EWL predictions.
598 Interestingly, these target stations have the lowest number of transferable models from the training
599 stations (Figure S2). More advanced deep learning models like Transformers have attention
600 mechanisms built into their design (Boussioux et al. 2022) and could be a worth-exploring
601 alternative to the proposed Bi-LSTM-ATT models for predicting EWL evolution in coastal areas.
602 We plan to extend the LSTM-SAM framework to inland target stations by taking into account the

603 contribution of river discharge for accurate prediction of total WLs in coastal to inland transition
604 zones (Serafin et al. 2017b; Bilskie and Hagen 2018; Muñoz et al. 2022b; Yin et al. 2022). Future
605 work should focus on predicting spatiotemporal WL variability and flood inundation extent by
606 combining Bi-LSTM-ATT and Convolutional Neural Networks (Gavahi et al. 2021).

607 **5. Conclusion**

608 In the present study, we characterize the evolution of extreme water levels (EWLs) at tide-gage
609 stations distributed along the U.S. Atlantic Coast. To achieve this, we identify 5 training stations
610 that were hit by historic hurricane events and contain complete consecutive hourly data spanning
611 at least 40 years. Then, we leverage available WL and hydrometeorological time-series data to
612 train bidirectional Long Short-Term Memory (Bi-LSTM) network models for each training station.
613 Furthermore, we incorporate an attention layer mechanism in the model architecture and a transfer
614 learning (TL) approach with the goal effectively predicting the evolution of EWLs at target (tide-
615 gage) stations. The models highlight the significance of a longer batch size in enhancing model
616 performance, while challenging the assumed benefits of longer look-back periods. The collection
617 of models with the attention layer mechanism and TL approach is referred to as the LSTM-Station
618 Approximated Models (LSTM-SAM) framework and is effectively applied to 14 target stations.

619 The LSTM-SAM framework predicts the onset, peak, and dissipation of multiple EWL events
620 emerging from tropical cyclones (hurricanes) and extratropical cyclones (Nor'easter storms) with
621 high accuracy. For this, the framework identifies "transferable" models based on KGE and NSE
622 above 0.70 in order to ensure an accurate generalization of EWLs. Under these criteria, the LSTM-
623 SAM framework demonstrates satisfactory performance with transferable models achieving
624 average KGE, NSE, and RMSE ranging from 0.78 to 0.92, 0.90 to 0.97, and 0.09 to 0.18 at the

625 target stations, respectively. Following these results, we conclude that the LSTM-SAM framework
626 can accurately predict not only EWLs but also their evolution over time, which could assist in
627 large-scale operational flood predictions like the National Water Model (NWM) or Coastal
628 Emergency Risk Assessment (CERA). Future work should focus on predicting spatiotemporal WL
629 variability and flood inundation extent by combining Bi-LSTM-ATT and Convolutional Neural
630 Networks.

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633 Climate Program (Award # 480948).

634 **References**

635 Abbaszadeh P, Gavahi K, Moradkhani H (2020) Multivariate remotely sensed and in-situ data
636 assimilation for enhancing community WRF-Hydro model forecasting. *Advances in Water
637 Resources* 145:103721. <https://doi.org/10.1016/j.advwatres.2020.103721>

638 Ahmed S, Saif AF, Hanif MI, et al (2022) Att-BiL-SL: Attention-Based Bi-LSTM and Sequential
639 LSTM for Describing Video in the Textual Formation. *Applied Sciences* 12:.
640 <https://doi.org/10.3390/app12010317>

641 Alarcon VJ, Linhoss AC, Kelble CR, et al (2022) Coastal inundation under concurrent mean and
642 extreme sea-level rise in Coral Gables, Florida, USA. *Natural Hazards* 1–30

643 Alipour A, Jafarzadegan K, Moradkhani H (2022) Global sensitivity analysis in hydrodynamic
644 modeling and flood inundation mapping. *Environmental Modelling & Software*
645 152:105398. <https://doi.org/10.1016/j.envsoft.2022.105398>

646 Almar R, Ranasinghe R, Bergsma EWJ, et al (2021) A global analysis of extreme coastal water

647 levels with implications for potential coastal overtopping. *Nat Commun* 12:3775.

648 <https://doi.org/10.1038/s41467-021-24008-9>

649 Altunkaynak A, Kartal E (2021) Transfer sea level learning in the Bosphorus Strait by wavelet

650 based machine learning methods. *Ocean Engineering* 233:109116.

651 <https://doi.org/10.1016/j.oceaneng.2021.109116>

652 Anderson DL, Ruggiero P, Mendez FJ, et al (2021) Projecting Climate Dependent Coastal Flood

653 Risk With a Hybrid Statistical Dynamical Model. *Earth's Future* 9:e2021EF002285.

654 <https://doi.org/10.1029/2021EF002285>

655 Arns A, Wahl T, Wolff C, et al (2020) Non-linear interaction modulates global extreme sea levels,

656 coastal flood exposure, and impacts. *Nature Communications* 11:1–9.

657 <https://doi.org/10.1038/s41467-020-15752-5>

658 Bai L-H, Xu H (2021) Accurate estimation of tidal level using bidirectional long short-term

659 memory recurrent neural network. *Ocean Engineering* 235:108765

660 Bates P (2023) Fundamental limits to flood inundation modelling. *Nat Water* 1:566–567.

661 <https://doi.org/10.1038/s44221-023-00106-4>

662 Bates P (2022a) Uneven burden of urban flooding. *Nature Sustainability*.

663 <https://doi.org/10.1038/s41893-022-01000-9>

664 Bates PD (2022b) Flood Inundation Prediction. *Annu Rev Fluid Mech* 54:287–315.

665 <https://doi.org/10.1146/annurev-fluid-030121-113138>

666 Bentivoglio R, Isufi E, Jonkman SN, Taormina R (2022) Deep learning methods for flood

667 mapping: a review of existing applications and future research directions. *Hydrology and*

668 *Earth System Sciences* 26:4345–4378

669 Bevacqua, Maraun D, Vousdoukas MI, et al (2019) Higher probability of compound flooding from

670 precipitation and storm surge in Europe under anthropogenic climate change. *Science*
671 *Advances* 5:eaaw5531. <https://doi.org/10.1126/sciadv.aaw5531>

672 Bian G-F, Nie G-Z, Qiu X (2021) How well is outer tropical cyclone size represented in the ERA5
673 reanalysis dataset? *Atmospheric Research* 249:105339

674 Bilskie MV, Hagen SC (2018) Defining Flood Zone Transitions in Low-Gradient Coastal Regions.
675 *Geophysical Research Letters* 45:2761–2770. <https://doi.org/10.1002/2018GL077524>

676 Bilskie MV, Zhao H, Resio D, et al (2021) Enhancing Flood Hazard Assessments in Coastal
677 Louisiana Through Coupled Hydrologic and Surge Processes. *Front Water* 3:.
678 <https://doi.org/10.3389/frwa.2021.609231>

679 Bischl B, Binder M, Lang M, et al (2023) Hyperparameter optimization: Foundations, algorithms,
680 best practices, and open challenges. *Wiley Interdisciplinary Reviews: Data Mining and*
681 *Knowledge Discovery* 13:e1484

682 Bloemendaal N, de Moel H, Martinez AB, et al (2022) A globally consistent local-scale assessment
683 of future tropical cyclone risk. *Science Advances* 8:eabm8438.
684 <https://doi.org/10.1126/sciadv.abm8438>

685 Boumis G, Moftakhari HR, Moradkhani H (2023) Coevolution of Extreme Sea Levels and Sea-
686 Level Rise Under Global Warming. *Earth's Future* 11:e2023EF003649.
687 <https://doi.org/10.1029/2023EF003649>

688 Boussioux L, Zeng C, Guénais T, Bertsimas D (2022) Hurricane Forecasting: A Novel Multimodal
689 Machine Learning Framework. *Weather and Forecasting* 37:817–831.
690 <https://doi.org/10.1175/WAF-D-21-0091.1>

691 Brakenridge GR (2021) Global active archive of large flood events, 1985-present. Available from
692 DFO Flood Observatory, University of Colorado(Accessed on October, 2021 through

693 personal communication)

694 Bruneau N, Polton J, Williams J, Holt J (2020) Estimation of global coastal sea level extremes
695 using neural networks. *Environmental Research Letters* 15:074030

696 C. Huang, J. Zhang, L. Cao, et al (2020) Robust Forecasting of River-Flow Based on
697 Convolutional Neural Network. *IEEE Transactions on Sustainable Computing* 5:594–600.
698 <https://doi.org/10.1109/TSUSC.2020.2983097>

699 Chen D, Zhang J, Jiang S (2020) Forecasting the Short-Term Metro Ridership With Seasonal and
700 Trend Decomposition Using Loess and LSTM Neural Networks. *IEEE Access* 8:91181–
701 91187. <https://doi.org/10.1109/ACCESS.2020.2995044>

702 Cleveland RB, Cleveland WS, McRae JE, Terpenning I (1990) STL: A seasonal-trend
703 decomposition. *J Off Stat* 6:3–73

704 Codiga DL (2011) Unified tidal analysis and prediction using the UTide Matlab functions

705 de Amorim LBV, Cavalcanti GDC, Cruz RMO (2023) The choice of scaling technique matters for
706 classification performance. *Applied Soft Computing* 133:109924.
707 <https://doi.org/10.1016/j.asoc.2022.109924>

708 De la Fuente LA, Ehsani MR, Gupta HV, Condon LE (2024) Toward interpretable LSTM-
709 based modeling of hydrological systems. *Hydrol Earth Syst Sci* 28:945–971.
710 <https://doi.org/10.5194/hess-28-945-2024>

711 Ding Y, Zhu Y, Feng J, et al (2020) Interpretable spatio-temporal attention LSTM model for flood
712 forecasting. *Neurocomputing* 403:348–359. <https://doi.org/10.1016/j.neucom.2020.04.110>

713 Douris J, Kim G (2021) The Atlas of Mortality and Economic Losses from Weather, Climate and
714 Water Extremes (1970-2019)

715 Evangelista LGC, Giusti R (2021) Short-term effects of weight initialization functions in Deep

716 NeuroEvolution. *Evo* 21

717 F. Zhuang, Z. Qi, K. Duan, et al (2021) A Comprehensive Survey on Transfer Learning.
718 Proceedings of the IEEE 109:43–76. <https://doi.org/10.1109/JPROC.2020.3004555>

719 Fang Z, Wang Y, Peng L, Hong H (2021) Predicting flood susceptibility using LSTM neural
720 networks. *Journal of Hydrology* 594:125734.
721 <https://doi.org/10.1016/j.jhydrol.2020.125734>

722 Fraehr N, Wang QJ, Wu W, Nathan R (2022) Upskilling Low-Fidelity Hydrodynamic Models of
723 Flood Inundation Through Spatial Analysis and Gaussian Process Learning. *Water
724 Resources Research* 58:e2022WR032248. <https://doi.org/10.1029/2022WR032248>

725 Gavahi K, Abbaszadeh P, Moradkhani H (2021) DeepYield: A combined convolutional neural
726 network with long short-term memory for crop yield forecasting. *Expert Systems with
727 Applications* 184:115511. <https://doi.org/10.1016/j.eswa.2021.115511>

728 Ghanbari M, Arabi M, Kao S-C, et al (2021) Climate Change and Changes in Compound Coastal-
729 Riverine Flooding Hazard Along the U.S. Coasts. *Earth's Future* 9:e2021EF002055.
730 <https://doi.org/10.1029/2021EF002055>

731 Glavovic B, Dawson R, Chow WT, et al (2022) Cities and settlements by the sea
732 Glorot X, Bengio Y (2010) Understanding the difficulty of training deep feedforward neural
733 networks. *JMLR Workshop and Conference Proceedings*, pp 249–256

734 Gul S, Khan GM, Yousaf S (2022) Multi-step short-term \$\$PM_{\{2.5\}}\$\$ forecasting for enactment
735 of proactive environmental regulation strategies. *Environmental Monitoring and
736 Assessment* 194:386. <https://doi.org/10.1007/s10661-022-10029-4>

737 Gupta, H.V., Kling, H., Yilmaz, K.K., Martinez, G.F., 2009. Decomposition of the mean squared
738 error and NSE performance criteria: Implications for improving hydrological modelling.

739 Journal of hydrology 377, 80–91.

740 Hamidi E, Peter BG, Muñoz DF, et al (2023) Fast Flood Extent Monitoring With SAR Change
741 Detection Using Google Earth Engine. *IEEE Transactions on Geoscience and Remote
742 Sensing* 61:1–19. <https://doi.org/10.1109/TGRS.2023.3240097>

743 Hamitouche M, Molina J-L (2022) A review of ai methods for the prediction of high-flow extremal
744 hydrology. *Water Resources Management* 36:3859–3876

745 Hewamalage H, Bergmeir C, Bandara K (2021) Recurrent neural networks for time series
746 forecasting: Current status and future directions. *International Journal of Forecasting*
747 37:388–427

748 Hino M, Nance E (2021) Five ways to ensure flood-risk research helps the most vulnerable. *Nature*
749 595:27–29

750 Jafarzadegan K, Alipour A, Gavahi K, et al (2021) Toward improved river boundary conditioning
751 for simulation of extreme floods. *Advances in Water Resources* 158:104059.
752 <https://doi.org/10.1016/j.advwatres.2021.104059>

753 Kardhana H, Valerian JR, Rohmat FI, Kusuma MS (2022) Improving Jakarta’s Katulampa Barrage
754 Extreme Water Level Prediction Using Satellite-Based Long Short-Term Memory (LSTM)
755 Neural Networks. *Water* 14:. <https://doi.org/10.3390/w14091469>

756 Khojasteh D, Glamore W, Heimhuber V, Felder S (2021) Sea level rise impacts on estuarine
757 dynamics: A review. *Science of The Total Environment* 780:146470.
758 <https://doi.org/10.1016/j.scitotenv.2021.146470>

759 Kingphai K, Moshfeghi Y (2022) On time series cross-validation for deep learning classification
760 model of mental workload levels based on EEG signals. Springer, pp 402–416

761 Kratzert F, Gauch M, Klotz D, Nearing G (2024) HESS Opinions: Never train an LSTM on a

762 single basin. *Hydrol Earth Syst Sci Discuss* 2024:1–19. <https://doi.org/10.5194/hess-2023-275>

764 Lee J-W, Irish JL, Bensi MT, Marcy DC (2021) Rapid prediction of peak storm surge from tropical
765 cyclone track time series using machine learning. *Coastal Engineering* 170:104024.
766 <https://doi.org/10.1016/j.coastaleng.2021.104024>

767 Li W, Kiaghadi A, Dawson C (2021) Exploring the best sequence LSTM modeling architecture
768 for flood prediction. *Neural Computing and Applications* 33:5571–5580.
769 <https://doi.org/10.1007/s00521-020-05334-3>

770 Lindemann B, Müller T, Vietz H, et al (2021) A survey on long short-term memory networks for
771 time series prediction. *Procedia CIRP* 99:650–655.
772 <https://doi.org/10.1016/j.procir.2021.03.088>

773 Liu Y, Yang Y, Chin RJ, et al (2023) Long Short-Term Memory (LSTM) Based Model for Flood
774 Forecasting in Xiangjiang River. *KSCE Journal of Civil Engineering* 27:5030–5040.
775 <https://doi.org/10.1007/s12205-023-2469-7>

776 Mahdianpari M, Salehi B, Rezaee M, et al (2018) Very Deep Convolutional Neural Networks for
777 Complex Land Cover Mapping Using Multispectral Remote Sensing Imagery. *Remote
778 Sensing* 10:1119. <https://doi.org/10.3390/rs10071119>

779 Marco R, Ahmad SSS, Ahmad S (2022) Bayesian hyperparameter optimization and Ensemble
780 Learning for Machine Learning Models on software effort estimation. *International Journal
781 of Advanced Computer Science and Applications* 13:

782 Marsooli R, Wang Y (2020) Quantifying Tidal Phase Effects on Coastal Flooding Induced by
783 Hurricane Sandy in Manhattan, New York Using a Micro-Scale Hydrodynamic Model.
784 *Frontiers in Built Environment* 6:149. <https://doi.org/10.3389/fbuil.2020.00149>

785 Merizalde MJ, Muñoz P, Corzo G, et al (2023) Integrating geographic data and the SCS-CN
786 method with LSTM networks for enhanced runoff forecasting in a complex mountain
787 basin. *Frontiers in Water* 5:1233899

788 Moftakhari H, Muñoz DF, Akbari Asanjan A, et al (2024) Nonlinear Interactions of Sea-Level
789 Rise and Storm Tide Alter Extreme Coastal Water Levels: How and Why? *AGU Advances*
790 5:e2023AV000996. <https://doi.org/10.1029/2023AV000996>

791 Moftakhari, Schubert JE, AghaKouchak A, et al (2019) Linking statistical and hydrodynamic
792 modeling for compound flood hazard assessment in tidal channels and estuaries. *Advances*
793 in Water Resources

794 128:28–38. <https://doi.org/10.1016/j.advwatres.2019.04.009>

795 Muis S, Lin N, Verlaan M, et al (2019) Spatiotemporal patterns of extreme sea levels along the
796 western North-Atlantic coasts. *Scientific Reports* 9:3391. <https://doi.org/10.1038/s41598-019-40157-w>

797 Muis S, Verlaan M, Winsemius HC, et al (2016) A global reanalysis of storm surges and extreme
798 sea levels. *Nature communications* 7:11969

799 Muñoz DF, Abbaszadeh P, Moftakhari H, Moradkhani H (2022a) Accounting for uncertainties in
800 compound flood hazard assessment: The value of data assimilation. *Coastal Engineering*
801 171:104057. <https://doi.org/10.1016/j.coastaleng.2021.104057>

802 Muñoz DF, Moftakhari H, Moradkhani H (2020) Compound Effects of Flood Drivers and Wetland
803 Elevation Correction on Coastal Flood Hazard Assessment. *Water Resources Research*
804 56:e2020WR027544. <https://doi.org/10.1029/2020WR027544>

805 Muñoz DF, Muñoz P, Moftakhari H, Moradkhani H (2021) From local to regional compound flood
806 mapping with deep learning and data fusion techniques. *Science of The Total Environment*
807 782:146927. <https://doi.org/10.1016/j.scitotenv.2021.146927>

808 Muñoz DF, Yin D, Bakhtyar R, et al (2022b) Inter-Model Comparison of Delft3D-FM and 2D
809 HEC-RAS for Total Water Level Prediction in Coastal to Inland Transition Zones. *JAWRA*
810 Journal of the American Water Resources Association 58:34–49.
811 <https://doi.org/10.1111/1752-1688.12952>

812 Nash, J.E., Sutcliffe, J.V., 1970. River flow forecasting through conceptual models part I—A
813 discussion of principles. *Journal of hydrology* 10, 282–290.

814 Nearing G, Cohen D, Dube V, et al (2024) Global prediction of extreme floods in ungauged
815 watersheds. *Nature* 627:559–563. <https://doi.org/10.1038/s41586-024-07145-1>

816 NOAA-NCEI (2023) National Centers for Environmental Information (NCEI) U.S. Billion-Dollar
817 Weather and Climate Disasters. <https://doi.org/10.25921/stkw-7w73>

818 Obara Y, Nakamura R (2022) Transfer learning of long short-term memory analysis in significant
819 wave height prediction off the coast of western Tohoku, Japan. *Ocean Engineering*
820 266:113048. <https://doi.org/10.1016/j.oceaneng.2022.113048>

821 Office for Coastal Management, 2024. Economics and Demographics [WWW Document]. URL
822 <https://coast.noaa.gov/states/fast-facts/economics-and-demographics.html> (accessed
823 4.1.24).

824 Parker K, Erikson L, Thomas J, et al (2023) Relative contributions of water-level components to
825 extreme water levels along the US Southeast Atlantic Coast from a regional-scale water-
826 level hindcast. *Natural Hazards*. <https://doi.org/10.1007/s11069-023-05939-6>

827 Pörtner H-O, Roberts DC, Masson-Delmotte V, et al (2019) IPCC special report on the ocean and
828 cryosphere in a changing climate. *IPCC Intergovernmental Panel on Climate Change:*
829 Geneva, Switzerland 1:1–755

830 Rainey JL, Brody SD, Galloway GE, Highfield WE (2021) Assessment of the growing threat of

831 urban flooding: A case study of a national survey. *Urban Water Journal* 18:375–381

832 Rashid MM, Moftakhar H, Moradkhani H (2024) Stochastic simulation of storm surge extremes
833 along the contiguous United States coastlines using the max-stable process.
834 *Communications Earth & Environment* 5:39. <https://doi.org/10.1038/s43247-024-01206-z>

835 Rithani M, Kumar RP, Doss S (2023) A review on big data based on deep neural network
836 approaches. *Artificial Intelligence Review* 56:14765–14801.
837 <https://doi.org/10.1007/s10462-023-10512-5>

838 Rueda A, Camus P, Tomás A, et al (2016) A multivariate extreme wave and storm surge climate
839 emulator based on weather patterns. *Ocean Modelling* 104:242–251.
840 <https://doi.org/10.1016/j.ocemod.2016.06.008>

841 Saksena S, Merwade V (2015) Incorporating the effect of DEM resolution and accuracy for
842 improved flood inundation mapping. *Journal of Hydrology* 530:180–194.
843 <https://doi.org/10.1016/j.jhydrol.2015.09.069>

844 Sanders BF, Schubert JE, Kahl DT, et al (2022) Large and inequitable flood risks in Los Angeles,
845 California. *Nature Sustainability*. <https://doi.org/10.1038/s41893-022-00977-7>

846 Santiago-Collazo FL, Bilskie MV, Hagen SC (2019) A comprehensive review of compound
847 inundation models in low-gradient coastal watersheds. *Environmental Modelling &*
848 *Software* 119:166–181. <https://doi.org/10.1016/j.envsoft.2019.06.002>

849 Seleem O, Ayzel G, Bronstert A, Heistermann M (2023) Transferability of data-driven models to
850 predict urban pluvial flood water depth in Berlin, Germany. *Natural Hazards and Earth
851 System Sciences* 23:809–822

852 Serafin KA, Ruggiero P (2014) Simulating extreme total water levels using a time-dependent,
853 extreme value approach. *Journal of Geophysical Research: Oceans* 119:6305–6329.

854 <https://doi.org/10.1002/2014JC010093>

855 Serafin KA, Ruggiero P, Stockdon HF (2017a) The relative contribution of waves, tides, and
856 nontidal residuals to extreme total water levels on U.S. West Coast sandy beaches.
857 Geophysical Research Letters 44:1839–1847. <https://doi.org/10.1002/2016GL071020>

858 Serafin KA, Ruggiero P, Stockdon HF (2017b) The relative contribution of waves, tides, and
859 nontidal residuals to extreme total water levels on U.S. West Coast sandy beaches.
860 Geophysical Research Letters 44:1839–1847. <https://doi.org/10.1002/2016GL071020>

861 Shen C (2018) A Transdisciplinary Review of Deep Learning Research and Its Relevance for
862 Water Resources Scientists. Water Resources Research 54:8558–8593.
863 <https://doi.org/10.1029/2018WR022643>

864 Shen C, Appling AP, Gentine P, et al (2023) Differentiable modelling to unify machine learning
865 and physical models for geosciences. Nature Reviews Earth & Environment 4:552–567.
866 <https://doi.org/10.1038/s43017-023-00450-9>

867 Tan C, Sun F, Kong T, et al (2018) A Survey on Deep Transfer Learning. In: Kůrková V,
868 Manolopoulos Y, Hammer B, et al. (eds) Artificial Neural Networks and Machine Learning
869 – ICANN 2018. Springer International Publishing, Cham, pp 270–279

870 Tate E, Rahman MA, Emrich CT, Sampson CC (2021) Flood exposure and social vulnerability in
871 the United States. Natural Hazards 106:435–457

872 Tedesco P, Rabault J, Sætra ML, et al (2023) Bias Correction of Operational Storm Surge Forecasts
873 Using Neural Networks. arXiv preprint arXiv:230100892

874 Thomas A, Dietrich J, Asher T, et al (2019) Influence of storm timing and forward speed on tides
875 and storm surge during Hurricane Matthew. Ocean Modelling 137:1–19.
876 <https://doi.org/10.1016/j.ocemod.2019.03.004>

877 Tiggeloven T, Couasnon A, van Straaten C, et al (2021) Exploring deep learning capabilities for
878 surge predictions in coastal areas. *Scientific reports* 11:17224

879 Velasquez-Montoya L, Overton MF, Sciaudone EJ (2020) Natural and anthropogenic-induced
880 changes in a tidal inlet: Morphological evolution of Oregon Inlet. *Geomorphology*
881 350:106871

882 Velasquez-Montoya L, Sciaudone EJ, Smyre E, Overton MF (2021) Vulnerability indicators for
883 coastal roadways based on barrier island morphology and shoreline change predictions.
884 *Natural Hazards Review* 22:04021003

885 Velasquez-Montoya L, Wargula A, Nangle J, et al (2022) Hydrodynamics of a tidal inlet under
886 gray to green coastal protection interventions. *Frontiers in Earth Science* 10:991667

887 Wahl, Jain S, Bender J, et al (2015) Increasing risk of compound flooding from storm surge and
888 rainfall for major US cities. *Nature Climate Change* 5:1093–1097.
889 <https://doi.org/10.1038/nclimate2736>

890 Wahl T, Haigh ID, Nicholls RJ, et al (2017) Understanding extreme sea levels for broad-scale
891 coastal impact and adaptation analysis. *Nature Communications* 8:16075.
892 <https://doi.org/10.1038/ncomms16075>

893 Wang X, Jin Y, Schmitt S, Olhofer M (2023) Recent advances in Bayesian optimization. *ACM*
894 *Computing Surveys* 55:1–36

895 Wing OE, Lehman W, Bates PD, et al (2022) Inequitable patterns of US flood risk in the
896 Anthropocene. *Nature Climate Change* 12:156–162

897 Yin D, Muñoz DF, Bakhtyar R, et al (2022) Extreme Water Level Simulation and Component
898 Analysis in Delaware Estuary during Hurricane Isabel. *JAWRA Journal of the American*
899 *Water Resources Association* 58:19–33. <https://doi.org/10.1111/1752-1688.12947>

900 Zhang Y, Ragettli S, Molnar P, et al (2022) Generalization of an Encoder-Decoder LSTM model
901 for flood prediction in ungauged catchments. *Journal of Hydrology* 614:128577.
902 <https://doi.org/10.1016/j.jhydrol.2022.128577>

903 Zhao G, Pang B, Xu Z, et al (2021) Improving urban flood susceptibility mapping using transfer
904 learning. *Journal of Hydrology* 602:126777. <https://doi.org/10.1016/j.jhydrol.2021.126777>

905 Zscheischler J, Martius O, Westra S, et al (2020) A typology of compound weather and climate
906 events. *Nature Reviews Earth & Environment* 1:333–347. <https://doi.org/10.1038/s43017-020-0060-z>