Comparison of methods to determine extraction efficiencies of Ra isotopes and ²²⁷Ac from large volume seawater samples

- 3 Morgane Léon¹, Pieter van Beek¹, Virginie Sanial², Marc Souhaut¹, Paul Henderson³, Matthew
- 4 A. Charette³
- 5 ¹ Laboratoire d'Etudes en Géophysique et Océanographie Spatiale (LEGOS), Université de
- 6 Toulouse, CNES/CNRS/IRD/Université Toulouse III Paul Sabatier, Toulouse, France
- 7 ² Université de Toulon, Aix Marseille Univ., CNRS, IRD, MIO, Toulon, France
- 8 ³Department of Marine Chemistry and Geochemistry, Woods Hole Oceanographic Institution,
- 9 Woods Hole, MA 02543, USA

1011

1 2

Keywords: Radium, Actinium-227, Extraction efficiency, Methodology, Ocean, Tracers

Abstract

12 13 14

15

16 17

18

19

20

21

2223

24

25

26

27

28 29 30

31

32

33

34

35

36 37

38

39

Radium isotopes, other than ²²⁶Ra, and ²²⁷Ac are typically present at low activities in the open ocean. The analysis of these isotopes thus requires collection of large volumes of seawater and high sensitivity, low background instruments. To obtain the required large volumes (hundreds to thousands of liters), these radionuclides are typically preconcentrated on cartridge-style filters impregnated with MnO₂ (Mn-cartridges) deployed on *in-situ* pumps. This technique, however, requires the determination of the extraction efficiency of the Mn-cartridges for the radionuclides of interest. For Ra isotopes, we used two methods to estimate the extraction efficiency of these Mn-cartridges at two stations on the South-West Indian Ridge in the Southern Ocean (GEOTRACES GS02). Method (1) compares the ²²⁶Ra activities recovered on the Mn-cartridges versus the activities determined in Mn-fibers, through which seawater was passed at a flow rate < 1 L min⁻¹ to quantitatively sorb Ra (Mn-fiber method) while method (2) combines the ²²⁶Ra activities determined from two Mn-cartridges placed in series on in-situ pumps (A-B method). The second method is also applied to determine the ²²⁷Ac extraction efficiency. We find a relatively wide-range of Ra and ²²⁷Ac extraction efficiencies across the dataset (from 44.8 % to 99.6 % for Ra, and from 23.7 % to 77.5 % for 227 Ac). Overall, the yield of 227 Ac extraction is lower than that of Ra (mean value of 49.3 ± 19.0 % for 227 Ac, n=10; mean value of 79.2 ± 10.3 % for Ra, n= 13, using the Mn-fiber method and a mean value of 63.9 ± 12.5 %, n=11 using the A-B method). Our dataset suggests that the Ra extraction efficiencies using the A-B method are lower than those determined using the Mn-fiber method (21 % on average). However, the Ra activities (223Ra_{ex}, ²²⁴Ra_{ex} and ²²⁸Ra) determined using either one of the two methods are not statistically different, probably due to the relatively high uncertainties associated with these activities. We also show that the ²²⁷Ac extraction efficiency can be estimated from the Ra extraction efficiency allowing the use of a single Mn-cartridge. Finally, we recommend to determine the Ra and ²²⁷Ac extraction efficiencies in each individual Mn-cartridge, rather than applying a single extraction efficiency to all the Mn-cartridges, since a significant variability in the extraction efficiencies was observed between the different Mn-cartridges.

1. Introduction

40

41

42

43

44 45

46

47

48

49 50

51

52

53

54

5556

57

58 59

60

61 62

63 64

65 66

67

68 69

70

71

72 73

74

75

76

77 78

79

80

81 82

83

84

85

Radium (Ra) has four naturally occurring isotopes (²²⁴Ra, t_{1/2}: 3.66 d; ²²³Ra, t_{1/2}: 11. 4 d; ²²⁸Ra, t_{1/2}: 5.75 y; ²²⁶Ra, t_{1/2}: 1600 y) continuously produced by radioactive decay of thorium (Th) isotope parents (²²⁸Th, ²²⁷Th, ²³²Th, ²³⁰Th, respectively) in the uranium-thorium decay series. While Th isotopes are strongly reactive to particles and preferentially adsorb onto mineral surfaces in marine systems (Cochran, 1982), Ra is easily released from surfaces or particles due to the high ionic strength and is therefore mostly found in the dissolved phase (Elsinger and Moore, 1980; Ku and Lin, 1976; Li and Chan, 1979; Moore, 1987). As Th and U are mainly present in soils, sediments and rocks, Ra isotopes in marine environment find their main source from diffusion from deep-sea sediments and continental shelves (Ku and Lin, 1976; Moore, 1969). Their concentrations in the ocean are thus widely dependent on the U and Th content in these sources.

The different half-lives of the Ra isotopes allow us to study chemical and physical processes in the ocean on different temporal and spatial scales (Annett et al., 2013; Charette et al., 2007; Dulaiova et al., 2009; Sanial et al., 2014; van Beek et al., 2008). Because of its relatively long halflife, 226 Ra is the most abundant Ra isotope with a total inventory in the oceans of about 4.78 ± 0.27 ×10¹⁸ Bq (IAEA, 1988; Hanfland, 2002; Neff, 2002) and is a historical tracer of water masses used to study the global scale deep ocean circulation (Broecker et al., 1976, 1967; Charette et al., 2016, 2016; Chung, 1987; Chung and Craig, 1980; Inoue et al., 2022; Ku et al., 1980; Ku and Lin, 1976; Le Roy et al., 2018). With a shorter half-life, ²²⁸Ra is preferentially used as a tracer to study mesoscale and coastal processes at a time scale of years such as horizontal mixing between the continental shelves and the open ocean (Kaufman et al., 1973; Kipp et al., 2018a; Knauss et al., 1978; Sanial et al., 2018; Yamada and Nozaki, 1986), river inputs (Moore and Krest, 2004; Vieira et al., 2020), vertical mixing (Charette et al., 2007; van Beek et al., 2008), Submarine Groundwater Discharge (SGD; Kim et al., 2005; Li et al., 1980; Moore et al., 2008; Rodellas et al., 2017) or hydrothermal vents (Kadko, 1996; Kadko et al., 2007; Kadko and Butterfield, 1998; Kadko and Moore, 1988; Kipp et al., 2018b; Léon et al, subm). On the other hand, ²²⁷Th (direct daughter of ²²⁷Ac) and ²²⁸Th (daughter of ²²⁸Ra) are particle reactive (Cochran, 1982) and release ²²³Ra and ²²⁴Ra, respectively, in the dissolved phase by radioactive decay. With half-lives in the order of days, ²²³Ra and ²²⁴Ra are used to study coastal processes on the timescale of days or weeks, including the quantification of SGD fluxes (Bejannin et al., 2017; Garcia-Orellana et al., 2021; Tamborski et al., 2017), flushing rates in estuaries and above continental shelves (Léon et al., 2022; Moore and Krest, 2004) and horizontal or vertical mixing coefficients (Charette et al., 2007; Koch-Larrouy et al., 2015; Léon et al., subm). Aside from ²²⁶Ra, which has a long half-life relative to ocean overturning timescales, Ra isotope activities generally decrease with increasing distance from their source due to mixing with seawater and radioactive decay. As such, intermediate waters tend to have lower activities of Ra isotopes than surface or bottom waters. Because the vertical mixing rate of the ocean is much slower than the mean life of ²²³Ra, ²²⁴Ra or even ²²⁸Ra, little of the Ra isotopes can penetrate into the intermediate ocean (Moore, 1969) with the exception of regions impacted by hydrothermal vents (Kadko, 1996; Kadko et al., 2007; Kadko and Butterfield, 1998; Kadko and Moore, 1988; Kipp et al., 2018b; Léon et al, subm). Their concentrations are thus often below the detection limit in the mid water column away from the ocean boundaries (Charette et al., 2007; Sanial et al., 2015; van Beek et al., 2008). In seawater, activities of Ra isotopes usually range from ca. 0.1 to several tens of disintegrations per minute (dpm) per 100 kg of seawater.

The 227 Ac isotope ($t_{1/2}$: 21.8 years) is part of the 235 U radioactive series ($t_{1/2}$: 7.04 10⁸ years). Due to its very long residence time in the oceans (\sim 0.5 Ma; Ku et al., 1977), 235 U is uniformly distributed in the oceans (Weyer et al., 2008) leading to a constant production of 231 Pa ($t_{1/2}$: 32,760

years) by radioactive decay of ²³⁵U in the water column. ²³¹Pa is preferentially scavenged onto particles and accumulates in sediments (Anderson et al., 1983). It decays into ²²⁷Ac which is more soluble than its parents (Nozaki, 1993); it is thus partially released into the dissolved phase and then redistributed in the deepest part of the water column by mixing and transport. ²²⁷Ac is thus mainly produced in deep waters and Nozaki (1984) was the first to propose ²²⁷Ac as a deep-sea tracer. Since then, it has been used to trace deep ocean circulation on a basin scale (~100 years timescale) to quantify vertical eddy diffusivity coefficients (Geibert et al., 2002; Koch-Larrouy et al., 2015; Le Roy et al., 2023; Nozaki, 1984), to estimate upwelling rates (Geibert et al., 2002) or to trace hydrothermal system (Kipp et al., 2015; Moore et al., 2008b). Despite its recognized interest, research about ²²⁷Ac remains difficult due to the low concentrations in the ocean. Geibert et al. (2008) thus estimated the total oceanic inventory of ²²⁷Ac to be 37 mol, which is equivalent to about 8.4 kg (2.25 × 10¹⁶ Bq). Relatively few studies using ²²⁷Ac are thus reported (Dulaiova et al., 2013; Geibert et al., 2008, 2002; Geibert and Vöge, 2008; Kipp et al., 2015; Koch-Larrouy et al., 2015; Le Roy et al., 2023, 2019; Levier et al., 2021; Nozaki, 1993, 1984; Shaw and Moore, 2002).

100 101 102

103

104

105

106

107

108

109

110

111

112

113

114

115

116

117

118

119 120

121

122

123

124

125

126

127

128

129

130

131

86

87

88

89 90

91

92

93

94

95

96 97

98

99

The analysis of Ra and ²²⁷Ac isotopes in the ocean therefore requires the collection of large volumes of water (several hundred liters) followed by an extraction method. The need to preconcentrate large volumes of seawater has historically greatly limited the number of samples for Ra and Ac isotopes in the open ocean. The extraction of these radionuclides could be done either through coprecipitation of radium with barium sulfates (Gordon and Rowley, 1957), or through Fe or Mn hydroxide coprecipitation (Geibert and Vöge, 2008) or through sorption onto a media impregnated with MnO₂ (Shaw and Moore, 2002). Filtration through a media impregnated with MnO₂ is less constraining to set up and limit chemistry steps. Commonly, seawater is filtered through i) acrylic fibers impregnated with MnO₂ (Mn-fibers) at a flow rate below 1 L min⁻¹ to ensure that 100 % of Ra and Ac are sorbed onto Mn-fibers (Moore and Reid, 1973) or through ii) acrylic cartridges impregnated with MnO₂ (Mn-cartridges) mounted on In-Situ Pumps (ISP) immersed for several hours to allow the filtration of hundreds of liters of seawater (Charette and Moran, 1999; Livingston and Cochran, 1987; Mann and Casso, 1984). On the one hand, the filtration of seawater through the Mn-fibers leads to a maximal extraction efficiency of the radionuclides but requires large volumes of water to be brought on board, usually using Niskin bottles mounted on a CTD rosette or a surface pump for upper water column sampling. This technique usually prevents from building vertical profiles with a high resolution and has thus often been applied for the analysis of ²²⁶Ra (and sometimes ²²⁸Ra) that can be conducted on relatively small volumes of seawater (~10-20 L). On the other hand, the use of ISP allows the filtration of large volumes of water (> 300 L) but the flow rate is higher than the maximum value of 1 L min⁻¹ that was shown to quantitatively sorb Ra from seawater. This latter method thus requires the quantification of the Ra extraction efficiency. To do so, one possibility is to mount a Niskin bottle above the ISP to compare the ²²⁶Ra activity determined from the Niskin bottle (via Mn-fiber media) versus the activity determined from the ISP (via Mn-cartridge media) in a same cast (Charette et al., 2015). As the isotopes of a same element have identical chemical properties, the ²²⁶Ra extraction efficiency can be applied to all Ra isotopes in a given sample. Livingston and Cochran (1987) developed another method to quantify radionuclides in open ocean waters by positioning two cartridges in series in ISP. The extraction efficiency of the radionuclides such as Ra and Ac is thus achieved by comparing the activities determined on the two cartridges. The radionuclide activity in seawater is subsequently quantified by applying the efficiency to the activity observed

in the first cartridge. However, extraction efficiency can vary with the cartridge properties, the water flow rate passing through the cartridge, and also potentially with environmental conditions.

In this study, we investigated the Ra and ²²⁷Ac activities at two stations located in the Southern Ocean and visited during the GEOTRACES GS02 cruise (SWINGS cruise). At these two stations, we deployed Niskin bottles to collect ~12 L samples and ISP equipped with two Mn-cartridges placed in series to filter *in-situ* large volumes of seawater. This sampling scheme allows us to determine the yield of Ra extraction either by 1) comparing the ²²⁶Ra activities determined in Mn-fibers with the ones determined in the first Mn-cartridge and (2) comparing the ²²⁶Ra activities in the two Mn-cartridges in ISP. We therefore aimed to provide information 1) on the validity of the method based on the use of two cartridges placed in series that may sometimes be questioned due to the assumptions associated with this method and 2) on the best way to determine the Ra extraction efficiency. Subsequently, we used the second method (Mn-cartridges placed in series) to determine the extraction efficiency of ²²⁷Ac, that we compared to the Ra extraction efficiency. Such comparison exercise has rarely been done because usually one single method is used for a dedicated cruise due to large seawater volumes needed and subsequent long analytical measurements to quantify the low concentrations of Ra isotopes and ²²⁷Ac in the open ocean.

2. Material and methods 2.1.Study area

The samples were collected at two stations located above the South West Indian Ridge (SWIR) in the Southern Ocean during the GEOTRACES GS02 cruise, which took place on the R/V *Marion Dufresne* between South Africa and Heard Island (January -March 2021). Along the section, these two stations (Station 14, 1388 m, 44°51.690 S, 36°10.460 E; Station 15, 1770 m, 44°51.178 S, 36°13.841 E; Fig. 1) were suspected to be under the influence of hydrothermal vents (Léon et al., subm.; Baudet et al., subm).

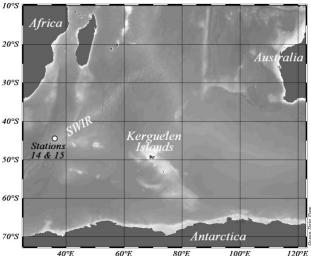


Figure 1: Map indicating the location of the stations 14 and 15 (open circle) investigated over the South West Indian Ridge in the Indian sector of the Southern Ocean.

2.2.Sampling method

First, water samples were collected using Niskin bottles mounted on a rosette and deployed at the same depths as the ISPs. These samples were designed to collect dissolved ²²⁶Ra, which displays higher activities in seawater than ²²³Ra, ²²⁴Ra and ²²⁸Ra, and can be analyzed in relatively small volumes (~12 L). Seawater from Niskin bottles were then passed by gravity through 10 g of MnO₂ impregnated acrylic fibers (Mn-fibers) placed into two small cartridges in series. The Mn-fibers were bought already impregnated to Scientific Computer Instruments and were "fluffed" to occupy the entire volume of the cartridges (Charette et al., 2012). The flow rate was set to 200 mL min⁻¹ to quantitatively adsorb Ra isotopes (Moore and Reid, 1973). Between 11.2 and 21.8 L of seawater were collected per sample (Tab. 1). The Mn-fibers were then placed into plastic bags and rinsed three times with Ra-free water, to remove salt from seawater.

Second, Mn-cartridges were impregnated in the laboratory with MnO₂ using the following protocol derived from (Henderson et al., 2013). CUNO acrylic cartridges were cut to obtain cartridges of about 77 ± 4 mm. After being dusted with compressed air, they were immersed in Ra free water for a minimum of 48 hours at room conditions (temperature of 20 ± 2°C and relative hygrometry of $60 \pm 10 \%$) in order to promote the future adsorption of manganese on the cartridges. The cartridges were then placed individually under vacuum in anti-permeation bags (SOREVAC ©: 80um thick gasproof) containing 200 mL of saturated KMnO₄ solution (0.5 M) for 24 to 48 hours at room conditions to make the solution penetrate to the core of the cartridge. The cartridges were then immersed in 3 successive tanks of Ra free water before being individually rinsed to ensure that all the excess KMnO₄ was washed out. The so-called Mn-cartridges were then stored individually in plastic bags under vacuum and in the dark until use. These Mn-cartridges were then mounted on McLane ISP to preconcentrate *in-situ* dissolved Ra isotopes and ²²⁷Ac from large volumes of seawater at different depths in the water column. A spring was placed under the Mncartridges into the ISP cartridge holder, in order to minimize seawater bypassing through the Mncartridge, especially at great depth where high pressure may reduce the size of the Mn-cartridges. Prior to passing through the Mn-cartridges, seawater was filtered through Supor (0.8 µm pore size) or QMA (Whatman© 1 µm pore size) filters mounted on the ISPs. Eight ISPs were deployed at station 14 and six ISPs were deployed at station 15. The pumping duration was set for 3 hours and thus filtering through the Mn-cartridges between 427 and 677 L of seawater (with a mean flow rate ranging from 2.3 up to 3.7 L min⁻¹; Tab. 1). The sampling resolution was increased near the seafloor because these samples were designed to trace the presence of a hydrothermal activity in the area. The yield of Ra fixation is unknown because the flow rate of seawater that passed through the Mncartridges is greater than the 1 L min⁻¹ recommended to get a yield of 100%. Except for the three shallowest pumps at station 14 (50 m, 200 m, 900 m), two Mn-cartridges were mounted in series at each depth (A Mn-cartridge followed by a B Mn-cartridge), to provide information on the yield of Ra and ²²⁷Ac fixation using the A-B method described in section 2.4 (Baskaran et al., 1993; Bourquin et al., 2008; Le Roy et al., 2019; Livingston and Cochran, 1987; Mann and Casso, 1984; van der Loeff and Moore, 1999). Each Mn-cartridges were rinsed with Ra-free water for ca. 10 minutes to remove salt from seawater. To do so, the Mn-cartridge holders were connected to two small cartridges placed in series, filled with Mn-fibers and themselves connected to the tap of the boat.

201202203

161162

163 164

165

166

167

168169

170

171

172

173

174

175

176

177

178

179

180

181

182 183

184

185

186

187

188

189

190

191

192

193 194

195

196

197

198

199

200

2.3. Analytical methods

2.3.1. Analysis of ²²³Ra, ²²⁴Ra and ²²⁷Ac using RaDeCCs

manuscript submitted to Marine Chemistry

The Mn-cartridges and Mn-fibers were analyzed using four Radium Delayed Coincidence Counter (RaDeCC, Scientific Computer Instruments, USA) systems. RaDeCCs display a very low background and are thus optimal equipment to quantify low seawater Ra and ²²⁷Ac activities. Background measurements were regularly conducted on board and later in the laboratory. Background corrections were made considering the value determined at the same period as the sample counting. The RaDeCCs were calibrated using Mn-cartridges and Mn-fibers impregnated with ²³²Th standards. Efficiencies for ²¹⁹Rn channel (used to quantify ²²³Ra and ²²⁷Ac activities) were estimated following calculations presented by Moore and Cai, (2013).

Each sample was analyzed for 6 to 24 hours with the RaDeCC system being flushed every 3 hours with air for 5 to 10 minutes before reintroducing helium. As ²²⁴Ra and ²²³Ra have short half-lives, they were measured on board within a few hours of sample collection to obtain the total ²²⁴Ra (²²⁴Ra_{tot}) and ²²³Ra (²²³Ra_{tot}) activities. A second measurement was conducted 21 days after sampling to determine the ²²⁴Ra supported by ²²⁸Th in the samples. Excess ²²⁴Ra (²²⁴Ra_{ex}) was determined by subtracting the supported activities from the ²²⁴Ra_{tot} activities. A third counting was performed approximately 90 days after sample collection to quantify the ²²³Ra supported by ²²⁷Ac. The supported activities were subtracted from the ²²³Ra_{tot} activities to determine excess ²²³Ra activities (²²³Ra_{ex}). Therefore, the activities reported for short-lived Ra isotopes are ²²³Ra_{ex} and ²²⁴Ra_{ex}. Error propagation calculations were conducted based on Garcia-Solsona et al. (2008).

To quantify ²²⁷Ac activities on A Mn-cartridges, a total of 65 measurements ranging from 6 to 25 hours, flushed every 3 hours, were performed to study the associated errors (1 triplicate, 3 quadruplicates and 10 quintuplets; Fig. 2). Thus, the external precision is associated with the standard deviation (1SD) of the replicate measurements of a given sample, while the internal precision is defined as the uncertainty calculated by error propagation according to Garcia-Solsona et al. (2008). Internal precisions are < 0.059 dpm (1 SD) and the relative standard deviation (RSD) range from 20.1 % to 44.3 % with an average of 30.9 %. As a comparison, the external precision of these replicates is < 0.032 dpm (1SD) and the RSD range from 2.1 % to 58.6 %, with a weighted average of 17.4 % (weights of 3 for triplicates, 4 for quadruplets, and 5 for quintuplets). With the exception of one sample (15_700m in Fig. 2) , all measurements have an internal precision larger than the weighted average of external precision. This suggests that the uncertainties calculated by propagation (Garcia-Solsona et al., 2008) are likely overestimated. Therefore, we prefer to report the external precision rather than the internal precision.

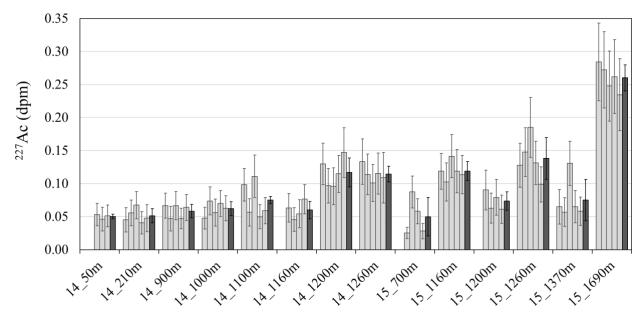


Fig. 2: Repeatability of ²²⁷Ac activities, in dpm, for the A Mn-cartridges (light grey bars). Dark grey bars show the mean values of these activities with the standard deviation of the mean (1SD). The sample IDs are listed on the x-axis as follows "Station Number depth".

2.3.2. Analysis of ²²⁶Ra and ²²⁸Ra using gamma spectrometers

Following RaDeCC counting, Mn-cartridges were ashed at 450°C for 15 to 20 minutes to reduce their volume, and let to cool down under a ventilated hood. The ashes were then placed in plastic tubes that were sealed with epoxy resin to prevent any loss of ²²²Rn. Mn-fibers were directly pressed and placed into plastic tubes. These tubes were also sealed with epoxy resin to prevent any loss of ²²²Rn. The samples were left at least 3 weeks before analysis to reach secular equilibrium between ²²⁶Ra and its daughters. The samples were analyzed using a SAGe-Well (MIRION-CANBERRA) germanium gamma spectrometer at the LAFARA laboratory in the French Pyrénées (van Beek et al., 2013) to determine the ²²⁶Ra and ²²⁸Ra activities. The detector has a volume of 450 cm³ and a well diameter of 32 mm. It is located underground below 85 m of rock to protect it from cosmic radiation, thus providing a very low background. Due to the low activity levels in the samples, both Mn-cartridges and Mn-fibers were analyzed for at least 4 days. The detector was calibrated using standards provided by IAEA (RGU1 and RGTH1). ²²⁶Ra was determined using ²¹⁴Pb peaks (294 keV and 352 keV) and ²¹⁴Bi peak (609 keV) and ²²⁸Ra using ²²⁸Ac peaks (339 keV and 969 keV). The uncertainties reported for the ²²⁶Ra activities include counting statistics and uncertainty on the detection efficiencies (1SD).

2.3.3. Analysis of ²²⁶Ra using ²²²Rn-emanation technique

The Mn-fibers (~12 L samples) were analyzed for ²²⁶Ra using the ²²²Rn emanation technique at WHOI, USA. The method involves an extraction line for ²²²Rn (daughter of ²²⁶Ra, half-life: 3.8 days) followed by a scintillation cell (Key et al., 1979b, 1979a). First, the Mn-fibers were placed in PVC cartridges (Peterson et al., 2009) and flushed with helium. The cartridges were then sealed and held for at least 5 days to allow for radioactive growth of ²²²Rn. The ²²²Rn was then flushed out and cryo-trapped in a copper tube cooled with liquid nitrogen. After 15 minutes

of ²²²Rn accumulation in the copper tube, the ²²²Rn was freed by heating the copper tube and further transferred into a Lucas cell by a helium flow. The Lucas cell was then placed into an alpha detector where it is counted from 3 to 6 hours. The counting system used for the cells is model AC/DC-DRC-MK 10-2. Uncertainties reported for ²²⁶Ra include counting statistics and uncertainty on the detection efficiencies (1SD).

2.4. Extraction efficiencies of the Mn-cartridges

2.4.1. Mn-fiber method

The first method to quantify the Ra extraction efficiency of the Mn-cartridges is to compare the 226 Ra activities determined in the Mn-fibers (~12 L from Niskin bottles) using the 222 Rn emanation technique to the 226 Ra activity determined in the A Mn-cartridges (ISP) by gamma spectrometry. The Ra extraction efficiency of Mn-cartridges E_1 (Ra), can then be determined as follows:

$$E_1(Ra) = \frac{Act_A}{Act_{Mn-fiber}} \quad (1)$$

where Acta and Act_{Mn-fiber} are the ²²⁶Ra activities (dpm 100L⁻¹) determined in the A Mn-cartridge and Mn-fiber, respectively (the volume considered here being the volume that passed onto the Mn-cartridge or the Mn-fiber, respectively). This efficiency can then be used to correct the activities determined in the Mn-cartridges for any of the other Ra isotopes that cannot be measured on small seawater volumes. The activity (dpm 100L⁻¹) of any Ra isotope in seawater (Actsw) can thus be determined using the following equation:

$$Act_{SW} = \frac{Act_A}{E_1(Ra)} \quad (2)$$

Note that in this study, the Niskin bottles and the ISPs were not deployed at the same time. We cannot exclude any temporal or spatial variability between the two samples, which may impact the estimate of the extraction efficiency. To prevent this potential bias, a Niskin bottle may be mounted directly above the ISP, so that the comparison between the two samples is more efficient (Charette et al., 2015).

2.4.2. A-B method

The second method to determine the Ra extraction efficiency (and potentially of other radionuclides such as Th and Ac) onto Mn-cartridges is based on the comparison of radionuclides activities from the two cartridges placed in series (A Mn-cartridge followed by B Mn-cartridge) mounted on the ISP (A-B method). This method was used in the past to quantify the activity of various radionuclides in the ocean (Fig. 3).

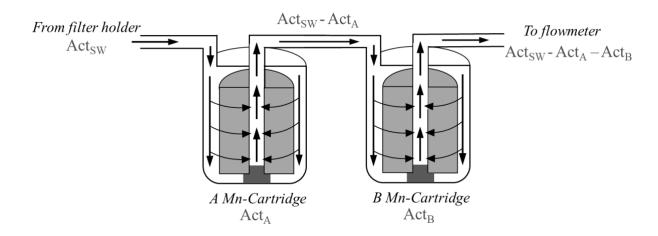


Fig. 3: Cross-section of the two Mn-cartridges placed in series on in situ pumps. The shaded zones represent the Mn-cartridges placed within the Mn-cartridges holder. Arrows indicate the seawater flow passing through the Mn-cartridges. Act_A, Act_B and Act_{SW} are the radionuclide activities (dpm 100L⁻¹) determined in the A Mn-cartridge, B Mn-cartridge and ambient seawater, respectively (the volume considered here being the volume that passed onto the Mn-cartridges.

In this case, seawater with a given activity (Actsw, dpm 100L⁻¹) first passes through the A Mn-cartridge. The activity on the A Mn-cartridge can thus be expressed as:

$$Act_A = Act_{SW} \times E_A \quad (3)$$

where Act_A is the activity (dpm 100L⁻¹) retained onto the A Mn-cartridge and E_A the extraction efficiency of the A Mn-cartridge. The seawater leaving the A Mn-cartridge – with an activity equal to Actsw minus Act_A – then passes through the B Mn-cartridge and exits the system with an activity equal to Actsw minus that absorbed onto the two Mn-cartridges consecutively (Actsw - Act_A - Act_B; Fig. 3). The activity of the B Mn-cartridge (Act_B, dpm 100L⁻¹) is thus:

$$Act_B = (Act_{SW} - Act_A) \times E_B$$
 (4)

where E_B is the extraction efficiency of the B Mn-cartridge. Assuming that the extraction efficiency for the A and B Mn-cartridges is the same ($E_A = E_B$), the extraction efficiency (E_2) for a given pair of Mn-cartridges can be determined as follows:

$$E_2 = 1 - \frac{Act_B}{Act_A} \quad (5)$$

In this study, E₂ (Ra) and E₂ (²²⁷Ac) refer to Ra and ²²⁷Ac extraction efficiencies, respectively, estimated by this A-B method. Note that in case the first Mn-cartridge becomes saturated with a radionuclide, the two cartridges may not behave similarly and the extraction efficiency may be distorted. However, the experience has shown that the radionuclide sorption capacity of Mn-cartridges is maintained after passing up to several thousand of liters of seawater through them (Baskaran et al., 1993; Charette et al., 2007; Charette and Moran, 1999). The radionuclide activity (dpm 100L⁻¹) in the water is then calculated with the following equation:

$$Act_{SW} = \frac{Act_A}{E_2} \quad (6)$$

Kemnitz (2018) has developed an alternative method which involves correcting the activity measured on the cartridges by taking into account the activity that is not retained by any of the two cartridges using the relationship:

$$Act_{SW} = \frac{Act_A + Act_B}{1 - f_{miss}} \tag{7}$$

where f_{miss} is the activity missed by both cartridges and equal to $(1 - E)^2$, where (1 - E) is the fraction missed by one Mn-cartridge. This correction method will also be used in the following as a comparison.

3. Results

3.1. Ra extraction efficiency

3.1.1. Mn-fiber method

The ²²⁶Ra activities determined in A Mn-cartridges and in Mn-fibers are shown in Fig. 4. The ²²⁶Ra activities in Mn-fibers are invariably higher than those in the Mn-cartridges because the Mn-fibers quantitatively extract Ra (Moore and Reid, 1973). The Ra extraction efficiency - E₁ (Ra) - of the Mn-cartridges can be estimated for each sample by using Equation 1 (Tab 1). The extraction efficiencies thus range from 61.8 % to 99.6 % with an average of 79.2 ± 10.3 %, where the associated error of the mean extraction efficiency is the standard deviation (1SD, n = 13; Tab. 1; Fig. 4). The uncertainty of each extraction efficiencies was estimated by propagation of the uncertainties associated with the ²²⁶Ra activities (1SD) of each Mn-cartridge and Mn-fiber (Tab.1). As a comparison, using the same method, Charette et al., (2015) estimated Ra extraction efficiencies of about 52 ± 22 % at an average flow rate of ~ 6 L min⁻¹. Using Mn-cartridges manufactured by the same protocol as in Charette et al. (2015), Sanial et al. (2018) noted that the addition of a spring in the cartridge holder resulted in a better radium extraction efficiency of $66 \pm$ 16% at a flow rate ~ 6 L min⁻¹. The use of springs has likely contributed to the improvement of the extraction efficiency. As a consequence, the standard deviation of the mean extraction efficiency has also decreased, which may point to an improved reproducibility of the process. Le Roy et al. (2019) reported an average of 60 ± 16 % (1SD, n = 15) at flow rate ranging from 3 to 6 L min⁻¹. We also note that the relatively low associated standard deviation may suggest a good reproducibility of the Mn impregnation on the cartridges. McLane ISP with slightly lower flow rates were used in this study (mean flow rate of 3.3 L min⁻¹) when compared to other studies which may have improved the extraction efficiency onto the Mn-cartridges (Charette and Moran, 1999; Henderson et al., 2013). The variability of the extraction efficiencies observed in the different studies may thus be related to differences in the average flow rate and potentially also to different Mn-cartridge properties (composition -polypropylene or acrylic, size or degree of MnO₂ impregnation). Possible causes explaining the variability of the extraction efficiencies from one Mn-cartridge to the other will be discussed below (section 4.2.).

367368

369

332

333

334

336

337

338

339

340

341

342

343344

345346

347

348349

350

351

352

353

354

355

356

357358

359

360361

362

363

364

365

366

Table 1: 226 Ra activities (in dpm $100L^{-1}$) determined in Mn-fibers and in A and B Mn-cartridges; the extraction efficiencies E_1 (Ra) (derived from Equation 1) and E_2 (Ra) (derived from Equation 5) determined

for each depth are also reported. No data were available (n.a.) at 500 m at station 15 (Mn-fibers) and at 50, 210, 900 m depth at station 14 (Mn-cartridges).

,			²²⁶ Ra		²²⁶ Ra	²²⁶ Ra	Extraction efficiency	
Station	Depth	Volume	Mn-fibers	Volume	A Mn-cartridge	B Mn-cartridge	E_{1}	E_2
	(m)	(L)	$(dpm\ 100L^{-1})$	(L)	$(dpm\ 100L^{-1})$	$(dpm\ 100L^{-1})$	(%)	(%)
14	50	11.6	11.1 ± 0.3	427	10.5 ± 0.13	n.a.	94.5 ± 2.8	n.a.
	210	12.4	13.2 ± 0.2	601	11.2 ± 0.12	n.a.	84.8 ± 1.7	n.a.
	900	12.4	15.3 ± 0.6	615	11.2 ± 0.07	n.a.	73.5 ± 3.1	n.a.
	1000	11.6	17.3 ± 0.8	528	10.7 ± 0.12	5.5 ± 0.08	61.8 ± 3.1	48.1 ± 0.9
	1100	11.5	15.5 ± 0.4	548	13.4 ± 0.14	4.7 ± 0.08	86.8 ± 2.5	65.4 ± 1.2
	1160	11.6	14.2 ± 0.7	584	11.0 ± 0.12	6.1 ± 0.08	77.4 ± 3.7	44.8 ± 0.8
	1200	11.7	16.8 ± 0.2	674	11.8 ± 0.12	4.6 ± 0.07	70.5 ± 1.0	61.3 ± 1.1
	1260	11.6	15.5 ± 0.3	646	11.2 ± 0.11	5.1 ± 0.08	72.1 ± 1.7	54.5 ± 1.0
15	700			532	10.6 ± 0.12	3.9 ± 0.07	n.a.	62.9 ± 1.3
	1160	10.9	15.8 ± 0.4	579	15.8 ± 0.15	$2.2~\pm~0.05$	99.6 ± 2.5	86.3 ± 2.2
	1200	11.3	16.6 ± 0.3	503	13.1 ± 0.14	$4.4~\pm~0.08$	79.3 ± 1.8	66.7 ± 1.3
	1260	11.1	17.5 ± 0.9	665	14.2 ± 0.14	$4.0~\pm~0.64$	81.2 ± 4.4	71.6 ± 11.3
	1370	11	16.2 ± 0.3	677	12.1 ± 0.12	4.6 ± 0.07	74.9 ± 1.7	61.7 ± 1.1
	1690	10.9	17.4 ± 0.3	630	12.8 ± 0.13	2.6 ± 0.05	73.4 ± 1.5	79.5 ± 1.6
						Mean	79.2 ± 10.3	63.9 ± 12.4

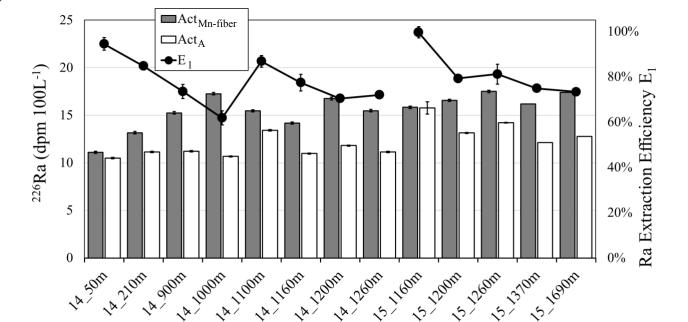


Fig. 4: Estimate of the Ra extraction efficiency $-E_1$ (Ra) - using the Mn-fiber method. The 226 Ra activities in the Mn-fibers (Act_{Mn-fiber}, in dpm $100L^{-1}$) are shown as dark grey bars and the 226 Ra activities in the A Mn-cartridges (Act_A, in dpm $100L^{-1}$) are shown as white bars. The Ra extraction efficiencies E_1 (Equation 1) determined by comparing Mn-fibers and A Mn-cartridges activities are shown as black dots. The sample IDs are listed on the x-axis as follows "Station Number depth".

3.1.2. A-B method

The ²²⁶Ra activities determined in B Mn-cartridges are always significantly lower than those determined in A Mn-cartridges, which is expected since seawater first passes through the A Mn-cartridge and then through the B Mn-cartridge (Tab. 1; Fig. 5). The E₂ (Ra) calculated using Equation 5 range from 44.8 % to 86.3 % with an average of 63.9 ± 12.4 %, where the associated error of the mean extraction efficiency is the standard deviation (1SD; n=11; Tab. 1). Uncertainties of each extraction efficiencies was estimated by propagation of the uncertainties associated with the ²²⁶Ra activities (1SD) of each Mn-cartridge (Tab.1). We note that the extraction efficiencies thus determined are slightly lower at station 14 (mean of 54.8 %) than at station 15 (mean of 71.5 %), even if average flow rates are similar at the two stations (~ 3.2 L min⁻¹). This pattern was not observed when using the Mn-fiber method (Fig.4) and is thus difficult to explain.

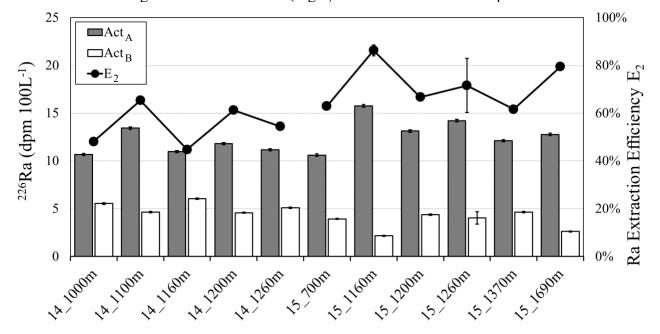


Fig. 5: Estimate of the Ra extraction efficiency - $E_2(Ra)$ - using the A-B method, The ^{226}Ra activities in the A Mn-cartridges (^{226}Ra Act_A, in dpm $100L^{-1}$) are shown as dark grey bars and the ^{226}Ra activities in the B Mn-cartridges (^{226}Ra Act_B, in dpm $100L^{-1}$) are shown as white bars. The Ra extraction efficiencies E_2 (Equation 5) determined by comparing A and B Mn-cartridges are shown as black dots. The sample IDs are listed on the x-axis as follows "Station Number depth".

3.2. ²²⁷Ac extraction efficiency

We used the A-B method to quantify the 227 Ac extraction efficiency and named it as E_2 (227 Ac). As for Ra, we compare the 227 Ac activities determined in the A Mn-cartridges with those determined in the B Mn-cartridges. Activities determined in B Mn-cartridges are always lower than those determined in A Mn-cartridges (Fig. 6; Tab.2). However, in contrast to Ra, it was not possible to compare these results with the Mn-fiber method, since 227 Ac could not be determined in our small volume samples due to its lower concentration. Fig. 6 shows E_2 (227 Ac) ranging from 23.7 % to 77.5 %, with a mean value of 49.3 ± 19 %, where the associated error of the mean extraction efficiency is the standard deviation (1SD, n = 10). The uncertainty of each extraction efficiency was estimated by propagation of the uncertainties associated with the 226 Ra activities (1SD) of each Mn-cartridge. The extraction efficiencies display a larger variability compared to the Ra extraction

efficiencies. Note that several activities reported here, especially on the B Mn-cartridges, are very low, thus resulting in relatively high associated errors. As observed with the E_2 (Ra) (Fig. 5), we note that E_2 (227 Ac) are on average lower at station 14 (34.4 %) than at station 15 (64.3 %), even if the average flow rates are the same at the two stations.

Table. 2: 227 Ac activities determined in the A and B Mn-cartridges (Act_A and Act_B, respectively, in dpm $^{100}L^{-1}$) together with the extraction efficiency - E_2 (227 Ac) - deduced from Equation 5.

			227	Ac	Extraction Efficiency
Station	Depth	Volume	A Mn-cartridge	B Mn-cartridge	$E_2(^{227}Ac)$
	(m)	(L)	$(dpm\ 100L^{-1})$	$(dpm\ 100L^{-1})$	(%)
14	1000	528	0.013 ± 0.002	0.008 ± 0.002	34.8 ± 8.3
	1100	548	0.015 ± 0.003	0.012 ± 0.002	23.7 ± 6.4
	1160	584	0.009 ± 0.002	0.006 ± 0.001	38.9 ± 11.9
	1200	674	0.018 ± 0.003	0.009 ± 0.002	50.7 ± 13.3
	1260	646	0.021 ± 0.004	0.016 ± 0.002	23.8 ± 5.5
15	1160	579	0.022 ± 0.003	0.005 ± 0.001	75.4 ± 19.2
	1200	503	0.016 ± 0.003	0.007 ± 0.002	56.4 ± 17.0
	1260	665	0.022 ± 0.004	0.009 ± 0.002	58.6 ± 14.7
	1370	677	0.013 ± 0.003	0.006 ± 0.001	53.5 ± 16.2
	1690	630	0.043 ± 0.007	0.010 ± 0.002	77.5 ± 18.1
				Mean	49.3 ± 19.0

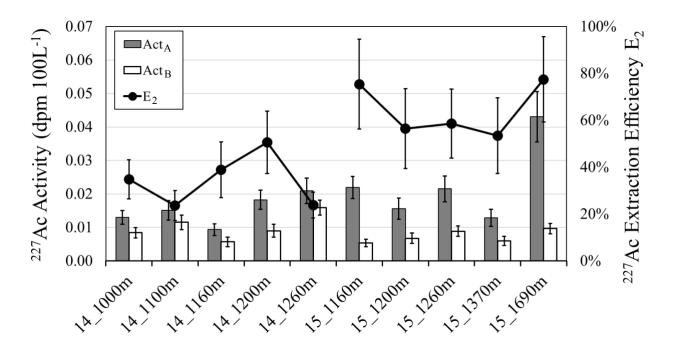


Fig. 6: Estimate of the 227 Ac extraction efficiency - E_2 (227 Ac) - using the A-B method. The 227 Ac activities in the A Mn-cartridges (Act_A, in dpm $100L^{-1}$) are shown as dark grey bars and the 227 Ac activities in the B Mn-cartridges (Act_B, in dpm $100L^{-1}$) are shown as white bars. The 227 Ac extraction efficiencies E_2 - E_2

manuscript submitted to Marine Chemistry

122 123	(²²⁷ Ac) - (Equation 5) determined by comparing A and B Mn-cartridges are shown in backs dots. The sample IDs are listed on the x-axis as follows "Station Number_depth".
124 125	Le Roy et al. (2019) also used acrylic Mn-cartridges produced with a similar protocol and reported E_2 (227Ac) that ranged from 31% to 78 %, with an average of 47 ± 12 % (1SD, n = 21).
426 427	These extraction efficiencies are consistent with the values reported in this study (Tab. 2). Kemnitz (2022) reported values of about 54 % ($n \sim 30$) using acrylic Mn-cartridges (same as in this study).
428 429	Using polypropylene Mn-cartridges, Geibert et al. (2002) and Geibert and Vöge (2008) reported extraction efficiencies of 69 ± 11 % (n = 31) and 77 ± 13 % (n = 31), respectively. Note that these
430 431 432	authors used larger Mn-cartridges. Kipp et al. (2015) proposed to estimate the ²²⁷ Ac extraction efficiency by averaging the Ra and Th extraction efficiencies at each depth. These authors thus reported efficiencies ranging from 27.3 % to 74.4 %.
132	4. Discussion
434 435	4.1. Comparison of the Ra and ²²⁷ Ac extraction efficiencies
+33 436 437	The Ra extraction efficiencies E ₁ and E ₂ , determined using the two different methods, are compared on Fig. 7 and 8.
138	
139	
140	Fig. 7 compares the Ra extraction efficiencies - E ₁ (Ra) and E ₂ (Ra) - with the ²²⁷ Ac
141 142	extraction efficiencies $E_2(^{227}Ac)$. With the exception of the sample collected at 1690 m depth at station 15, the Ra extraction efficiencies are invariably higher than the ^{227}Ac extraction efficiencies
143 144 145	determined in the same Mn-cartridges. The relationship shown in Fig. 1 in the supplementary material suggests that the E_2 (227 Ac) are on average 41 % lower than the E_1 (Ra) and 27 % lower than the E_2 (Ra).

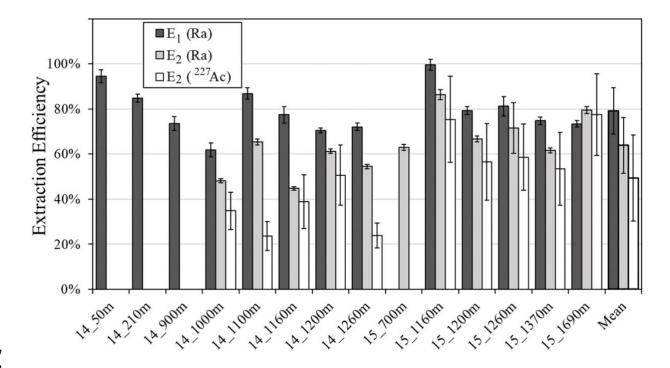
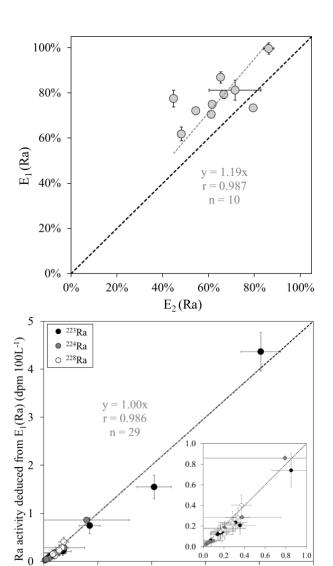


Fig. 7: Comparison of Ra and 227 Ac extraction efficiencies. E_1 (Ra) are shown in dark grey bars, E_2 (Ra) in light grey bars and E_2 (227 Ac) in white bars. The mean extraction efficiency of each method is reported in the right of the figure. The sample IDs are listed on the x-axis as follows "Station Number_depth". Only one cartridge was deployed at 50, 210, and 900 m for station 14, and at 700 m at station 15 preventing us from determining the extraction efficiency E_2 .



0.0

Ra activity deduced from E₂(Ra) (dpm 100L⁻¹)

2

0.0

3

0.2

0.4

0.6 0.8

455 456 457

458

459

460

461

0

454

Fig. 8: Left panel: Comparison of the Ra extraction efficiencies, E₂ (Ra) as a function of E₁ (Ra); right panel: Comparison of the Ra activities determined from E₁ (Ra) with the Ra activities determined from E₂ (Ra). The ²²³Ra activities are in black dots, the ²²⁴Ra activities are in grey dots and the ²²⁸Ra activities are in white dots at stations 14 and 15. The black dashed line shows the 1:1 relationship while the grey dotted line shows the linear regression forced by the origin.

5

462 463 464

465 466

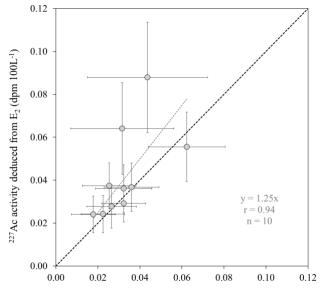
467

468

469

470

The relationship shown on figure 8 suggests that the E2 (Ra) are on average 21 % lower than the E_1 (Ra) (r = 0.987, n = 10, p-value < 0.001). Fig. 8 compares the 223 Ra, 224 Ra and 228 Ra activities determined using the two extraction efficiencies (E₁ or E₂). We observe that although E₂ (Ra) is slightly lower than E₁ (Ra), it does not have a significant impact on the final Ra activities regardless of the Ra isotope (r = 0.986; n = 29, p-value < 0.001), probably due to the relatively large uncertainties of the Ra activities determined using these methods. The correlation holds true when individual Ra isotopes are considered (r= 0.987, n= 10, p-value < 0.001 for 223 Ra; r= 0.986, n= 9. p-value < 0.001, for 224 Ra; r= 0.986, n= 10, p-value < 0.001 for 228 Ra). On figure 9, we now compare the ²²⁷Ac activities determined using either equation 6 that involves the extraction efficiencies determined from the A-B method or equation 7, following Kemnitz (2018). In the first method, the extraction efficiencies are determined in each individual Mn-cartridge (Tab.2), whereas Kemnitz (2018) use an average correction applied to all samples. Another difference is that the extraction efficiencies are applied to the sole A Mn-Cartridge to determine the 227 Ac activities for the first method, whereas, in the method of Kemnitz (2018), the 227 Ac activities are determined by summing the activities determined in the A and B Mn-cartridges. In this study, the average 227 Ac fraction missed f_{miss} is 15 ± 3 % (n=10). This value is obtained by plotting the 227 Ac activities in the B Mn-cartridges versus A Mn-cartridges (Kemnitz, 2018). Note that the average 226 Ra fraction missed f_{miss} thus calculated is 11 ± 1 % (n=11). Slightly higher activities using equation 6 are observed, but overall, we observe a relatively good agreement between the two estimates (r = 0.965, n = 10, p-value < 0.001). However, if we apply the method of Kemnitz (2018) to each individual samples (by calculating a f_{miss} factor to each sample, rather than applying an average correction to all samples) we find the same 227 Ac activities between the two methods.



 ^{227}Ac activity deduced using the f_{miss} factor (dpm $100L^{-1})$ Fig. 9: Comparison of the ^{227}Ac activities determined by correcting Act_A by E_2 (^{227}Ac) (equation 6) as a function of the ^{227}Ac activities estimated by correcting Act_A and Act_B by the missing factor f_{miss} following Kemnitz et al. (2018) (equation 7). The black dashed line shows the 1:1 relationship while the grey dotted line shows the linear regression forced by the origin.

4.2. Hypotheses that could explain the extraction efficiency variability

Overall, a relatively good agreement is observed between the two methods suggesting that the A-B method provides similar estimates of the Ra extraction efficiency (r = 0.987, n = 10, p-value <0.001). We observe, however, a high variability in the extraction efficiency estimates from one sample to the other, independently from the method used. One can invoke the following hypotheses to explain the extraction efficiency variability between the Mn-cartridges.

The Mn impregnation of the cartridges may vary from one sample to the other. We cannot exclude that CUNO acrylic cartridges may display slightly different shapes, even before being impregnated with KMnO₄. Variability in the porosity of the cartridges could lead to a different

KMnO₄ adsorption onto the cartridges, thus leading to variable Ra adsorption onto the Mn-cartridges.

The increase of pressure with increasing depth may impact the shape of the Mn-cartridge or impact its position in the cartridge holder which could in turn impact the water flow through the Mn-cartridge. No relationship, however, was observed between E₁ (Ra), E₂(Ra) or E₂(²²⁷Ac) and sample depth. As sample depth has a major influence on pressure encountered by the Mn-cartridges during the immersion of the ISP, this result thus suggests that pressure does not explain at first order the extraction efficiency variability. In this study, we placed springs in the cartridge holder below the Mn-cartridges that is expected to prevent seawater from flowing without passing through the Mn-cartridge.

Another factor that could influence the extraction efficiency is the flow rate of seawater filtered through the Mn-cartridges during *in-situ* pumping (Moore and Reid, 1973). The flow of water through the Mn-cartridges may vary from one Mn-cartridge to the other and may vary through time during filtration. Preferential pathways may also exist within the cartridge holder. Factors that can reduce the flow rate include: filter clogging at the ISP inlet, the power decrease of the ISP batteries, or pressure that could possibly reduce the size of the Mn-cartridge, making water circulation through Mn-cartridges more difficult. A decrease in the water flow would increase the extraction efficiencies.

Finally, we note that the extraction efficiencies calculated with the A-B method - E_2 (Ra) - are slightly lower than those estimated by the Mn-fiber method - E_1 (Ra). The A-B method allows us to determine extraction efficiencies based on two Mn-cartridges placed in series, assuming that they have the same extraction efficiency. This latter assumption may not hold completely true. Another explanation that could explain this almost systematic difference between both methods (E_1 (Ra) and E_2 (Ra)) is that some of the Ra present on the Mn-cartridges may partially desorb. When seawater passes through the Mn-cartridges and especially at high flow rates, some of the Ra present on the A Mn-cartridges may be removed (e.g., desorption of radium, leaching of Mn,...) and may then adsorb onto the B Mn-cartridges, potentially compensating the leaching of the Ra out of the B Mn-cartridges. This process would reduce the difference in Ra adsorbed between the two Mn-cartridges and would then reduce the extraction efficiencies determined using the A-B method (E_2).

4.3. Estimation of ²²⁷Ac extraction efficiency from E₁ (Ra)

Based on our dataset, the Ra extraction efficiencies are slightly higher than the 227 Ac extraction efficiencies (Fig. 7). Here we evaluate if the 227 Ac extraction efficiency can be determined from the Ra extraction efficiency, using E₁ (Ra) as the reference. The E₂ (227 Ac) : E₁ (Ra) ratio is on average 0.64 ± 0.23 (1SD, n=10). This suggests that the 227 Ac extraction efficiency E₃ may be estimated as follows:

$$E_3(^{227}Ac) = 0.64 \times E_1(Ra) \tag{8}$$

The ²²⁷Ac activity in seawater Act_{SW}(²²⁷Ac) may thus be determined as follows:

$$Act_{SW}(^{227}Ac) = \frac{Act_A(^{227}Ac)}{E_3} = \frac{Act_A(^{227}Ac)}{0.64 \times E_1(Ra)}$$
(9)

The uncertainties associated with E_3 (^{227}Ac) are estimated by error propagation of ^{226}Ra activities and E_1 (Ra). Fig. 10 shows the ^{227}Ac activities determined using the ^{227}Ac extraction efficiency deduced from E_3 (^{227}Ac) (equation 9) versus the ^{227}Ac activities determined from E_2 (^{227}Ac) at each depth (equation 6). We find that there is a relatively good agreement between the two activities (r=0.929, r=13, r=13

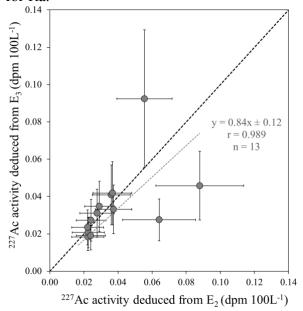


Fig. 10: Comparison of the 227 Ac activities determined from the A-B method E_2 (227 Ac) with the 227 Ac activities determined from the E_1 (Ra) extraction efficiency, corrected as described in section 4.2. The black dashed line shows the 1:1 relationship while the grey dotted line shows the linear regression forced by the origin.

5. Conclusion

Because they are present at very low concentrations in the ocean, Ra isotopes and ²²⁷Ac require the collection of large volumes of seawater, thus considerably limiting the construction of vertical profiles with a high resolution. Mn-cartridges mounted on in-situ pumps are thus commonly used to preconcentrate these isotopes from large seawater volumes (hundreds of liters). This technique, however, requires that the extraction efficiencies of the Mn-cartridges for these radionuclides are determined. Alternatively, Mn-fibers can be used to quantitatively preconcentrate Ra isotopes when seawater is passed at a flow rate < 1 L min⁻¹. This latter technique is generally used to mainly determine ²²⁶Ra activities since the determination of ²²³Ra, ²²⁴Ra, ²²⁸Ra and ²²⁷Ac

activities in open ocean waters requires larger volumes of seawater. In this work, we evaluated the extraction efficiencies of the Mn-cartridges (1) by comparing the activities of the Mn-fibers versus the Mn-cartridges (Mn-fiber method) and (2) by comparing the activities determined in two cartridges placed in series on in-situ pumps (A-B method).

We estimated Ra extraction efficiencies ranging from 61.8 to 99.6 % (mean of 79.2 ± 10.3 %) using the Mn-fiber method and ranging from 44.8 to 86.3 % (mean of 63.9 ± 12.4 %) using the A-B method. Although the Ra extraction efficiencies are slightly lower when using the A-B method, the results obtained here tend to confirm the reliability of the A-B method to estimate extraction efficiencies. We also used the A-B method to quantify the ²²⁷Ac extraction efficiency of the Mncartridges. Here, we report 227 Ac extraction efficiencies ranging from 23.7 to 77.5 (mean of 49.3 \pm 19 %) which are significantly lower than the Ra extraction efficiency (by a factor of 0.64, when compared to the mean Ra extraction efficiency). These results suggest that the ²²⁷Ac extraction efficiency may be estimated from the Ra extraction efficiency. However, more studies are required before it can be concluded that a constant factor exists between the Ra and ²²⁷Ac extraction efficiencies from Mn-cartridges mounted on ISP, since the extraction efficiencies may vary due to different factors (e.g., flow rate, size and chemical composition of the Mn-cartridges, etc...). Finally, because the extraction efficiencies of the Mn-cartridges were shown to vary from one sample to the other, we recommend that the efficiencies are quantified in each individual sample rather than using a mean efficiency that would be applied to all samples. We recognize that in this study, only a relatively small number of samples have been analyzed, and that the study of a larger number of samples would have led to statistically more reliable results. However, only few studies have been carried out to date to compare the different preconcentration methods. We recommend that more studies are conducted in the future to test the different methods for quantifying the extraction efficiencies of Mn-cartridges.

Acknowledgements

563

564

565

566

567

568569

570

571

572

573574

575

576

577578

579

580

581

582

583

584

585

586

587

588

589

590

591

592

593

594

595

596597

598

599

600

601

602

603

604

The authors thank Hélène Planquette and Catherine Jeandel, PIs of the SWINGS project and chief scientist of the SWINGS cruise. We also thank the captain and the crew of the R/V Marion Dufresne for their assistance during the SWINGS cruise. The authors also thank Emmanuel de Saint-Léger and Fabien Pérault, Marion Lagarde, Nolwenn Lemaitre, Edwin Cotard, Frederic Planchon for their help during the deployment of the *in-situ* pumps. We are grateful to Elodie Kestenare, Frederic Vivier, Gérard Eldin, Sara Sergi, Corentin Clerc and Loyd Izard for CTD data acquisition and preparation of Niskin bottles. We thank Thomas Zambardi for his help at the LAFARA underground laboratory. The SWINGS project was supported by the French Oceanographic Fleet, ANR (Agence Nationale de la Recherche; ANR) CNRS/ INSU (Centre Nationale de la Recherche Scientifique/Institut National des Sciences de l'Univers) and ISblue (ANR-17-EURE-0015). This study has been partially supported through the grant EUR TESS N°ANR-18-EURE-0018 in the framework of the Programme des Investissements d'Avenir. M.A.C. was funded by the U.S. National Science Foundation OCE-2048067.

Authors contributions

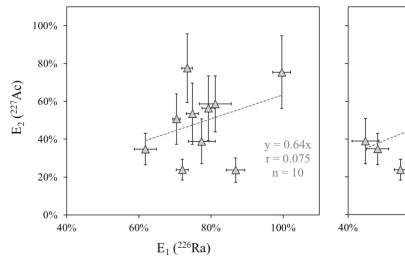
The sampling design for fieldwork was conducted by PvB and VS. PvB, MS and ML mobilized equipment and consumables for fieldwork. Samples were collected in the field by PvB, VS and ML. Sample analysis was conducted by PvB, VS, ML, MS, MAC and PH. ML, PvB, VS and MAC

analyzed and interpreted the data, ML produced the figures and wrote the paper. All authors provided comments on subsequent drafts of the paper.

607 608 609

605 606

Supplementary Material:



610 611

612

613614

Fig. 1: Comparison of the 227 Ac extraction efficiencies, E_2 (227 Ac) as a function of E_1 (Ra); on the left panel and as function of E_2 (Ra) on the right panel. The grey dotted line shows the linear regression forced by the origin.

y = 0.78x

r = 0.511

n = 10

100%

80%

 E_2 (226Ra)

60%

- Anderson, R.F., Bacon, M.P., Brewer, P.G., 1983. Removal of230Th and231Pa from the open ocean. Earth Planet. Sci. Lett. 62, 7–23. https://doi.org/10.1016/0012-821X(83)90067-5
- Annett, A.L., Henley, S.F., Van Beek, P., Souhaut, M., Ganeshram, R., Venables, H.J., Meredith, M.P., Geibert, W., 2013. Use of radium isotopes to estimate mixing rates and trace sediment inputs to surface waters in northern Marguerite Bay, Antarctic Peninsula.

 Antarct. Sci. 25, 445–456. https://doi.org/10.1017/S0954102012000892
- Baskaran, M., Murphy, D.J., Santschi, P.H., Orr, J.C., Schink, D.R., 1993. A method for rapid in situ extraction and laboratory determination of Th, Pb, and Ra isotopes from large volumes of seawater. Deep Sea Res. Part Oceanogr. Res. Pap. 40, 849–865. https://doi.org/10.1016/0967-0637(93)90075-E
- Bejannin, S., van Beek, P., Stieglitz, T., Souhaut, M., Tamborski, J., 2017. Combining airborne thermal infrared images and radium isotopes to study submarine groundwater discharge along the French Mediterranean coastline. J. Hydrol. Reg. Stud. 13, 72–90. https://doi.org/10.1016/j.ejrh.2017.08.001
- Bourquin, M., Van Beek, P., Reyss, J.-L., Souhaut, M., Charette, M.A., Jeandel, C., 2008.
 Comparison of techniques for pre-concentrating radium from seawater. Mar. Chem. 109,
 226–237. https://doi.org/10.1016/j.marchem.2008.01.002
- 632 Broecker, W.S., Goddard, J., Sarmiento, J.L., 1976. The distribution of 226Ra in the Atlantic 633 Ocean. Earth Planet. Sci. Lett. 32, 220–235. https://doi.org/10.1016/0012-634 821X(76)90063-7
- Broecker, W.S., Li, Y.H., Cromwell, J., 1967. Radium-226 and Radon-222: Concentration in Atlantic and Pacific Oceans. Science 158, 1307–1310. https://doi.org/10.1126/science.158.3806.1307
- Charette, M.A., Dulaiova, H., Gonneea, M.E., Henderson, P.B., Moore, W.S., Scholten, J.C.,
 Pham, M.K., 2012. GEOTRACES radium isotopes interlaboratory comparison
 experiment: Radium Intercomparison. Limnol. Oceanogr. Methods 10, 451–463.
 https://doi.org/10.4319/lom.2012.10.451

643

644

- Charette, M.A., Gonneea, M.E., Morris, P.J., Statham, P., Fones, G., Planquette, H., Salter, I., Garabato, A.N., 2007. Radium isotopes as tracers of iron sources fueling a Southern Ocean phytoplankton bloom. Deep Sea Res. Part II Top. Stud. Oceanogr. 54, 1989–1998. https://doi.org/10.1016/j.dsr2.2007.06.003
- Charette, M.A., Lam, P.J., Lohan, M.C., Kwon, E.Y., Hatje, V., Jeandel, C., Shiller, A.M.,
 Cutter, G.A., Thomas, A., Boyd, P.W., Homoky, W.B., Milne, A., Thomas, H.,
 Andersson, P.S., Porcelli, D., Tanaka, T., Geibert, W., Dehairs, F., Garcia-Orellana, J.,
 2016. Coastal ocean and shelf-sea biogeochemical cycling of trace elements and isotopes:
 lessons learned from GEOTRACES. Philos. Trans. R. Soc. Math. Phys. Eng. Sci. 374,
 20160076. https://doi.org/10.1098/rsta.2016.0076
- Charette, M.A., Moran, S.B., 1999. Rates of particle scavenging and particulate organic carbon export estimated using 234Th as a tracer in the subtropical and equatorial Atlantic Ocean. Deep Sea Res. Part II Top. Stud. Oceanogr. 46, 885–906. https://doi.org/10.1016/S0967-0645(99)00006-5
- Charette, M.A., Morris, P.J., Henderson, P.B., Moore, W.S., 2015. Radium isotope distributions
 during the US GEOTRACES North Atlantic cruises. Mar. Chem. 177, 184–195.
 https://doi.org/10.1016/j.marchem.2015.01.001
- 659 Chung, Y., 1987. 226Ra in the western Indian Ocean. Earth Planet. Sci. Lett. 85, 11–27. 660 https://doi.org/10.1016/0012-821X(87)90017-3

- 661 Chung, Y., Craig, H., 1980. 226Ra in the Pacific Ocean. Earth Planet. Sci. Lett. 49, 267–292. 662 https://doi.org/10.1016/0012-821X(80)90072-2
- 663 Cochran, J.K., 1982. The oceanic chemistry of the U- and Th-series nuclides.
- Dulaiova, H., Ardelan, M.V., Henderson, P.B., Charette, M.A., 2009. Shelf-derived iron inputs
 drive biological productivity in the southern Drake Passage: SHELF-DERIVED IRON IN
 THE DRAKE PASSAGE. Glob. Biogeochem. Cycles 23, n/a-n/a.
 https://doi.org/10.1029/2008GB003406
- Dulaiova, H., Sims, K.W.W., Charette, M.A., Prytulak, J., Blusztajn, J.S., 2013. A new method
 for the determination of low-level actinium-227 in geological samples. J. Radioanal. Nucl.
 Chem. 296, 279–283. https://doi.org/10.1007/s10967-012-2140-0
- Elsinger, R.J., Moore, W.S., 1980. 226Ra behavior in the Pee Dee River-Winyah Bay estuary.

 Earth Planet. Sci. Lett. 48, 239–249. https://doi.org/10.1016/0012-821X(80)90187-9
- Garcia-Orellana, J., Rodellas, V., Tamborski, J., Diego-Feliu, M., van Beek, P., Weinstein, Y.,
 Charette, M., Alorda-Kleinglass, A., Michael, H.A., Stieglitz, T., Scholten, J., 2021.
 Radium isotopes as submarine groundwater discharge (SGD) tracers: Review and
 recommendations. Earth-Sci. Rev. 220, 103681.
 https://doi.org/10.1016/j.earscirev.2021.103681
- 678 Garcia-Solsona, E., Garcia-Orellana, J., Masqué, P., Dulaiova, H., 2008. Uncertainties associated 679 with 223Ra and 224Ra measurements in water via a Delayed Coincidence Counter 680 (RaDeCC). Mar. Chem. 22.
 - Geibert, W., Charette, M., Kim, G., Moore, W.S., Street, J., Young, M., Paytan, A., 2008. The release of dissolved actinium to the ocean: A global comparison of different end-members. Mar. Chem. 109, 409–420. https://doi.org/10.1016/j.marchem.2007.07.005
 - Geibert, W., Rutgers van der Loeff, M.M., Hanfland, C., Dauelsberg, H.-J., 2002. Actinium-227 as a deep-sea tracer: sources, distribution and applications. Earth Planet. Sci. Lett. 198, 147–165. https://doi.org/10.1016/S0012-821X(02)00512-5
 - Geibert, W., Vöge, I., 2008. Progress in the determination of 227Ac in sea water. Mar. Chem. 109, 238–249. https://doi.org/10.1016/j.marchem.2007.07.012
- 689 Gordon, Louis., Rowley, Keith., 1957. Coprecipitation of Radium with Barium Sulfate. Anal. Chem. 29, 34–37. https://doi.org/10.1021/ac60121a012
- Henderson, P.B., Morris, P.J., Moore, W.S., Charette, M.A., 2013. Methodological advances for measuring low-level radium isotopes in seawater. J. Radioanal. Nucl. Chem. 296, 357–362. https://doi.org/10.1007/s10967-012-2047-9
 Inoue, M., Hanaki, S., Kameyama, H., Kumamoto, Y., Nagao, S., 2022. Unique current
 - Inoue, M., Hanaki, S., Kameyama, H., Kumamoto, Y., Nagao, S., 2022. Unique current connecting Southern and Indian Oceans identified from radium distributions. Sci. Rep. 12, 1781. https://doi.org/10.1038/s41598-022-05928-y
- Kadko, D., 1996. Radioisotopic studies of submarine hydrothermal vents. Rev. Geophys. 34, 349–366. https://doi.org/10.1029/96RG01762
- Kadko, D., Butterfield, D.A., 1998. The relationship of hydrothermal fluid composition and crustal residence time to maturity of vent fields on the Juan de Fuca Ridge. Geochim. Cosmochim. Acta 62, 1521–1533. https://doi.org/10.1016/S0016-7037(98)00088-X
- Kadko, D., Gronvold, K., Butterfield, D., 2007. Application of radium isotopes to determine crustal residence times of hydrothermal fluids from two sites on the Reykjanes Peninsula, Iceland. Geochim. Cosmochim. Acta 71, 6019–6029.
- 705 https://doi.org/10.1016/j.gca.2007.09.018

682

683

684

685

686

687

688

695

- Kadko, D., Moore, W., 1988. Radiochemical constraints on the crustal residence time of
 submarine hydrothermal fluids: Endeavour Ridge. Geochim. Cosmochim. Acta 52, 659–
 668. https://doi.org/10.1016/0016-7037(88)90328-6
- 709 Kaufman, A., Trier, R.M., Broecker, W.S., Feely, H.W., 1973. Distribution of ²²⁸ Ra in the world ocean. J. Geophys. Res. 78, 8827–8848. https://doi.org/10.1029/JC078i036p08827
- Kemnitz, N.J., 2018. Actinium-227 as a tracer for mixing in the deep Northeast Pacific (Thesis of the requirements for the degree Master of science). University of Southern California.
- Key, R.M., Brewer, R.L., Stockwell, J.H., Guinasso, N.L., Schink, D.R., 1979a. Some improved techniques for measuring radon and radium in marine sediments and in seawater. Mar. Chem. 7, 251–264. https://doi.org/10.1016/0304-4203(79)90042-2
- 716 Key, R.M., Guinasso, N.L., Schink, D.R., 1979b. Emanation of radon-222 from marine 717 sediments. Mar. Chem. 7, 221–250. https://doi.org/10.1016/0304-4203(79)90041-0
- Kim, G., Ryu, J.-W., Yang, H.-S., Yun, S.-T., 2005. Submarine groundwater discharge (SGD) into the Yellow Sea revealed by 228Ra and 226Ra isotopes: Implications for global silicate fluxes. Earth Planet. Sci. Lett. 237, 156–166. https://doi.org/10.1016/j.epsl.2005.06.011

726

727

728

729

730

- Kipp, L.E., Charette, M.A., Hammond, D.E., Moore, W.S., 2015. Hydrothermal vents: A previously unrecognized source of actinium-227 to the deep ocean. Mar. Chem. 177, 583–590. https://doi.org/10.1016/j.marchem.2015.09.002
 - Kipp, L.E., Charette, M.A., Moore, W.S., Henderson, P.B., Rigor, I.G., 2018a. Increased fluxes of shelf-derived materials to the central Arctic Ocean. Sci. Adv. 4, eaao1302. https://doi.org/10.1126/sciadv.aao1302
 - Kipp, L.E., Sanial, V., Henderson, P.B., van Beek, P., Reyss, J.-L., Hammond, D.E., Moore, W.S., Charette, M.A., 2018b. Radium isotopes as tracers of hydrothermal inputs and neutrally buoyant plume dynamics in the deep ocean. Mar. Chem. 201, 51–65. https://doi.org/10.1016/j.marchem.2017.06.011
- Knauss, K.G., Ku, T.-L., Moore, W.S., 1978. Radium and thorium isotopes in the surface waters
 of the East Pacific and coastal Southern California. Earth Planet. Sci. Lett. 39, 235–249.
 https://doi.org/10.1016/0012-821X(78)90199-1
- Koch-Larrouy, A., Atmadipoera, A., van Beek, P., Madec, G., Aucan, J., Lyard, F., Grelet, J.,
 Souhaut, M., 2015. Estimates of tidal mixing in the Indonesian archipelago from
 multidisciplinary INDOMIX in-situ data. Deep Sea Res. Part Oceanogr. Res. Pap. 106,
 136–153. https://doi.org/10.1016/j.dsr.2015.09.007
- Ku, T.L., Huh, C.A., Chen, P.S., 1980. Meridional distribution of226Ra in the eastern Pacific along GEOSECS cruise tracks. Earth Planet. Sci. Lett. 49, 293–308.
 https://doi.org/10.1016/0012-821X(80)90073-4
- Ku, T.-L., Lin, M.-C., 1976. 226Ra distribution in the Antarctic Ocean. Earth Planet. Sci. Lett.
 32, 236–248. https://doi.org/10.1016/0012-821X(76)90064-9
- Ku, T.-L., Mathieu, G.G., Knauss, K.G., 1977. Uranium in open ocean: concentration and
 isotopic composition. Deep Sea Res. 24, 1005–1017. https://doi.org/10.1016/0146-6291(77)90571-9
- Le Roy, E., Sanial, V., Charette, M.A., van Beek, P., Lacan, F., Jacquet, S.H.M., Henderson, P.B., Souhaut, M., García-Ibáñez, M.I., Jeandel, C., Pérez, F.F., Sarthou, G., 2018. The 226Ra–Ba relationship in the North Atlantic during GEOTRACES-GA01. Biogeosciences 15, 3027–3048. https://doi.org/10.5194/bg-15-3027-2018
- Le Roy, E., Sanial, V., Lacan, F., van Beek, P., Souhaut, M., Charette, M.A., Henderson, P.B., 2019. Insight into the measurement of dissolved 227Ac in seawater using radium delayed

- 753 coincidence counter. Mar. Chem. 212, 64–73. 754 https://doi.org/10.1016/j.marchem.2019.04.002
- Le Roy, E., van Beek, P., Lacan, F., Souhaut, M., Sanial, V., Charette, M.A., Henderson, P.B.,
 Deng, F., 2023. The distribution of 227Ac along the GA01 section in the North Atlantic.
 Mar. Chem. 248, 104207. https://doi.org/10.1016/j.marchem.2023.104207
- Léon, M., Beek, P., Scholten, J., Moore, W.S., Souhaut, M., De Oliveira, J., Jeandel, C., Seyler, P., Jouanno, J., 2022. Use of ²²³ Ra and ²²⁴ Ra as chronometers to estimate the residence time of Amazon waters on the Brazilian continental shelf. Limnol. Oceanogr. 67, 753–761 767. https://doi.org/10.1002/lno.12010
- Levier, M., Roy-Barman, M., Colin, C., Dapoigny, A., 2021. Determination of low level of actinium 227 in seawater and freshwater by isotope dilution and mass spectrometry. Mar. Chem. 233, 103986. https://doi.org/10.1016/j.marchem.2021.103986
- Li, Y.-H., Chan, L.-H., 1979. Desorption of Ba and 226Ra from river-borne sediments in the
 Hudson estuary. Earth Planet. Sci. Lett. 43, 343–350. https://doi.org/10.1016/0012-821X(79)90089-X
- Li, Y.-H., Feely, H.W., Toggweiler, J.R., 1980. 228Ra and 228Th concentrations in GEOSECS
 Atlantic surface waters. Deep Sea Res. Part Oceanogr. Res. Pap. 27, 545–555.
 https://doi.org/10.1016/0198-0149(80)90039-4
- Livingston, H.D., Cochran, J.K., 1987. Determination of transuranic and thorium isotopes in
 ocean water: In solution and in filterable particles. J. Radioanal. Nucl. Chem. Artic. 115,
 299–308. https://doi.org/10.1007/BF02037445
- 774 Mann, D.R., Casso, S.A., 1984. In situ chemisorption of radiocesium from seawater. Mar. Chem. 14, 307–318. https://doi.org/10.1016/0304-4203(84)90027-6
- Moore, W.S., 2008. Fifteen years experience in measuring 224Ra and 223Ra by delayed coincidence counting. Mar. Chem. 109, 188–197.
 https://doi.org/10.1016/j.marchem.2007.06.015
- 779 Moore, W.S., 1987. Radium 228 in the South Atlantic Bight. J. Geophys. Res. 92, 5177. https://doi.org/10.1029/JC092iC05p05177
- 781 Moore, W.S., 1976. Sampling 228Ra in the deep ocean. Deep Sea Res. Oceanogr. Abstr. 23, 647–651. https://doi.org/10.1016/0011-7471(76)90007-3
- 783 Moore, W.S., 1969. Oceanic concentrations of 228 Radium. Earth Planet. Sci. Lett. 6, 437–446. 784 https://doi.org/10.1016/0012-821X(69)90113-7
- Moore, W.S., Cai, P., 2013. Calibration of RaDeCC systems for 223Ra measurements. Mar. Chem. 156, 130–137. https://doi.org/10.1016/j.marchem.2013.03.002
- Moore, W.S., Krest, J., 2004. Distribution of 223Ra and 224Ra in the plumes of the Mississippi and Atchafalaya Rivers and the Gulf of Mexico. Mar. Chem. 86, 105–119. https://doi.org/10.1016/j.marchem.2003.10.001
- Moore, W.S., Reid, D.F., 1973. Extraction of radium from natural waters using manganese impregnated acrylic fibers. J. Geophys. Res. 78, 8880–8886.
 https://doi.org/10.1029/JC078i036p08880
- Moore, W.S., Sarmiento, J.L., Key, R.M., 2008a. Submarine groundwater discharge revealed by
 228Ra distribution in the upper Atlantic Ocean. Nat. Geosci. 1, 309–311.
 https://doi.org/10.1038/ngeo183
- Moore, W.S., Ussler, W., Paull, C.K., 2008b. Short-lived radium isotopes in the Hawaiian margin: Evidence for large fluid fluxes through the Puna Ridge. Mar. Chem. 109, 421–430. https://doi.org/10.1016/j.marchem.2007.09.010

- Neff, J.M., 2002. Radium Isotopes in the Ocean, in: Bioaccumulation in Marine Organisms. Elsevier, pp. 191–201. https://doi.org/10.1016/B978-008043716-3/50012-9
- Nozaki, Y., 1993. Actinium-227: A Steady State Tracer for the Deep-sea Basin-wide Circulation and Mixing Studies, in: Elsevier Oceanography Series. Elsevier, pp. 139–156. https://doi.org/10.1016/S0422-9894(08)71323-0
- 804 Nozaki, Y., 1984. Excess 227Ac in deep ocean water. Nature 310, 486–488. 805 https://doi.org/10.1038/310486a0

810

811

824

825

826

827

828

829

830

- Peterson, R.N., Burnett, W.C., Dimova, N., Santos, I.R., 2009. Comparison of measurement methods for radium-226 on manganese-fiber: Methods for ²²⁶ Ra analysis on Mn-fiber. Limnol. Oceanogr. Methods 7, 196–205. https://doi.org/10.4319/lom.2009.7.196
 - Reid, D.F., Key, R.M., Schink, D.R., 1979. Radium, thorium, and actinium extraction from seawater using an improved manganese-oxide-coated fiber. Earth Planet. Sci. Lett. 43, 223–226. https://doi.org/10.1016/0012-821X(79)90205-X
- Rodellas, V., Garcia-Orellana, J., Trezzi, G., Masqué, P., Stieglitz, T.C., Bokuniewicz, H.,
 Cochran, J.K., Berdalet, E., 2017. Using the radium quartet to quantify submarine
 groundwater discharge and porewater exchange. Geochim. Cosmochim. Acta 196, 58–73.
 https://doi.org/10.1016/j.gca.2016.09.016
- Sanial, V., Kipp, L.E., Henderson, P.B., van Beek, P., Reyss, J.-L., Hammond, D.E., Hawco,
 N.J., Saito, M.A., Resing, J.A., Sedwick, P., Moore, W.S., Charette, M.A., 2018. Radium228 as a tracer of dissolved trace element inputs from the Peruvian continental margin.
 Mar. Chem. 201, 20–34. https://doi.org/10.1016/j.marchem.2017.05.008
- Sanial, V., van Beek, P., Lansard, B., d'Ovidio, F., Kestenare, E., Souhaut, M., Zhou, M., Blain,
 S., 2014. Study of the phytoplankton plume dynamics off the Crozet Islands (Southern
 Ocean): A geochemical-physical coupled approach. J. Geophys. Res. Oceans 119, 2227–
 2237. https://doi.org/10.1002/2013JC009305
 - Sanial, V., van Beek, P., Lansard, B., Souhaut, M., Kestenare, E., d'Ovidio, F., Zhou, M., Blain, S., 2015. Use of Ra isotopes to deduce rapid transfer of sediment-derived inputs off Kerguelen. Biogeosciences 12, 1415–1430. https://doi.org/10.5194/bg-12-1415-2015
 - Shaw, T.J., Moore, W.S., 2002. Analysis of 227Ac in seawater by delayed coincidence counting. Mar. Chem. 78, 197–203. https://doi.org/10.1016/S0304-4203(02)00022-1
 - Tamborski, J.J., Cochran, J.K., Bokuniewicz, H.J., 2017. Application of 224Ra and 222Rn for evaluating seawater residence times in a tidal subterranean estuary. Mar. Chem. 189, 32–45. https://doi.org/10.1016/j.marchem.2016.12.006
- van Beek, P., Bourquin, M., Reyss, J.-L., Souhaut, M., Charette, M.A., Jeandel, C., 2008. Radium isotopes to investigate the water mass pathways on the Kerguelen Plateau (Southern Ocean). Deep Sea Res. Part II Top. Stud. Oceanogr. 55, 622–637.
 https://doi.org/10.1016/j.dsr2.2007.12.025
- van Beek, P., Souhaut, M., Lansard, B., Bourquin, M., Reyss, J.-L., von Ballmoos, P., Jean, P.,
 2013. LAFARA: a new underground laboratory in the French Pyrénées for ultra low-level
 gamma-ray spectrometry. J. Environ. Radioact. 116, 152–158.
 https://doi.org/10.1016/j.jenvrad.2012.10.002
- van der Loeff, M.M.R., Moore, W.S., 1999. Determination of natural radioactive tracers, in:
 Grasshoff, K., Kremling, K., Ehrhardt, M. (Eds.), Methods of Seawater Analysis. WileyVCH Verlag GmbH, Weinheim, Germany, pp. 365–397.
- https://doi.org/10.1002/9783527613984.ch13

manuscript submitted to Marine Chemistry

844	Vieira, L.H., Krisch, S., Hopwood, M.J., Beck, A.J., Scholten, J., Liebetrau, V., Achterberg, E.P.
845	2020. Unprecedented Fe delivery from the Congo River margin to the South Atlantic
846	Gyre. Nat. Commun. 11, 556. https://doi.org/10.1038/s41467-019-14255-2
847	Weyer, S., Anbar, A.D., Gerdes, A., Gordon, G.W., Algeo, T.J., Boyle, E.A., 2008. Natural
848	fractionation of 238U/235U. Geochim. Cosmochim. Acta 72, 345–359.
849	https://doi.org/10.1016/j.gca.2007.11.012
850	Yamada, M., Nozaki, Y., 1986. Radium isotopes in coastal and open ocean surface waters of the
851	Western North Pacific. Mar. Chem. 19, 379–389. https://doi.org/10.1016/0304-
852	4203(86)90057-5
853	