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3 Title: **Modeling cheatgrass distribution, abundance, and response to climate change as a**
4 **function of soil microclimate**

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18 resistance and resilience, SHAW model

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22 **Abstract**

23 Exotic annual grass invasions in water-limited systems cause degradation of native plant and
24 animal communities and increased fire risk. The life history of invasive annual grasses allows for
25 high sensitivity to interannual variability in weather. Current distribution and abundance models
26 derived from remote sensing, however, provide only a coarse understanding of how species
27 respond to weather, making it difficult to anticipate how climate change will affect vulnerability
28 to invasion. Here we derived germination covariates (rate sums) from mechanistic germination
29 and soil microclimate models to quantify favorability of soil microclimate for cheatgrass
30 (*Bromus tectorum* L.) establishment and growth across 30 years at 2662 sites across the
31 sagebrush steppe system in the western United States. Our approach, using four bioclimatic
32 covariates alone, predicted cheatgrass distribution with accuracy comparable to previous models
33 fit using many years of remotely-sensed imagery. Accuracy metrics from our out-of-sample
34 testing dataset indicate that our model predicted distribution well (72% overall accuracy) but
35 explained patterns of abundance poorly ($R^2 = 0.22$). Climatic suitability for cheatgrass presence
36 depended on both spatial (mean) and temporal (annual anomaly) variation of fall and spring rate
37 sums. Sites that on average have warm and wet fall soils and warm and wet spring soils (high
38 rate sums during these periods) were predicted to have a high abundance of cheatgrass.
39 Interannual variation of fall soil conditions had a greater impact on cheatgrass presence and
40 abundance than spring conditions. Our model predicts that climate change has already affected
41 cheatgrass distribution with suitable microclimatic conditions expanding 10-17% from 1989 to
42 2019 across all aspects at low to mid-elevation sites, while high elevation sites ($> 2100\text{m}$) remain
43 unfavorable for cheatgrass due to cold spring and fall soils.

44 **Introduction**

45 Invasive annual grasses have been linked to a worldwide decline of biodiversity and ecosystem
46 functioning in water-limited systems (D'Antonio & Vitousek 1992; Ostoja & Schupp 2009;
47 Davies 2011). Once present, these grasses accelerate fire return intervals (Fusco et al. 2019),
48 increase nitrogen storage in soils (Wolkovich et al., 2010), and compete with native vegetation
49 for soil moisture (Melgoza et al., 1990). Positive feedback loops with fire and nutrient cycles
50 have enabled invasive annual grasses to dominate substantial portions of water-limited systems
51 that now burn 2-4 times more frequently than native communities (D'Antonio & Vitousek 1992;
52 Bradley et al., 2018; Fusco et al., 2019; Pastick et al. 2021). Information about the current and
53 future distribution of these grasses is crucial to guide management decisions and wildfire
54 planning (Chambers et al., 2014).

55 Climate change will likely alter the vulnerability of different ecosystems to annual grass invasion
56 and dominance (Catford et al., 2019). In the sagebrush steppe, arid conditions are becoming
57 more prevalent (Ficklin & Novick 2017) due to increases in annual minimum temperature,
58 increased climatic water deficit, and less summer rainfall (Smith et al., 2022). These temporal
59 trends have potential to impact invasion, as colder and wetter locations that have previously been
60 classified as resistant to invasion (Chambers et al., 2014) are now becoming warmer and drier
61 (Bradford et al., 2019; Smith et al., 2022). Despite many experimental and observational studies,
62 there remains uncertainty surrounding future effects of climate change on cheatgrass (*Bromus*
63 *tectorum* L.) distribution. Large-scale studies predict range expansion of annual grasses (Pastick
64 et al., 2021; Smith et al., 2022), but results of cheatgrass-specific studies vary. While some
65 predict increases in abundance (Boyte et al., 2016), others predict no changes (Brummer et al.,
66 2016; Zimmer et al., 2021) or a dependence on future precipitation scenarios (Bradley 2009).

67 Experimental studies show both gains in fitness with warming (Compagnoni & Adler 2014,
68 Blumenthal et al. 2016), and no effect (Zelikova et al., 2013; Larson et al., 2017; 2018). The
69 signal may be unclear because of site-level factors that influence how atmospheric weather
70 impacts soil temperature and moisture. What is clear is a lack of consensus on how changing
71 climate will facilitate or inhibit cheatgrass expansion across western North America.

72 Our understanding of the climatic niche of invasive annual grass species and potential shifts in
73 distribution due to climate change could be improved by new modeling approaches that directly
74 link field observations with soil microclimate (Bradford et al., 2019; Boehm et al., 2021;
75 Hardegree et al., 2022). Current landscape-scale mapping approaches use reflectance indices
76 such as NDVI (normalized difference vegetation index) to estimate the distribution and
77 abundance of invasive annual grasses (Downs et al., 2016; Bradley et al., 2018; Pastick et al.,
78 2021). While these models are useful for mapping large areas and make it possible to study the
79 distribution or abundance of invasive annual grass species without extensive field sampling, they
80 do not directly describe species-climate relationships. Rather, one model links cheatgrass
81 abundance or distribution with reflectance, and a second analysis correlates climate variables
82 with estimates of abundance or distribution. Quantifying error propagation from the first model
83 through to the second is challenging; ignoring that error means overestimating the certainty of
84 the species-climate correlations. Previous studies linking climate to large-scale cheatgrass
85 distribution (Bradley 2009; Boyte et al., 2016) correlated climate metrics to remotely sensed
86 estimates of cheatgrass distribution that either included large uncertainty ($R^2 = 0.21$)(Peterson
87 2006) or required several years of imagery to distinguish between cheatgrass dominated sites and
88 sites dominated by other vegetation types (Bradley & Mustard 2005). An approach that directly
89 explains spatially and temporally extensive field observations with fine-scale climate indices

90 would improve inference about species-climate relationships and could also account for
91 interannual variation in abundance.

92 Persistent cheatgrass presence and abundance is largely driven by soil moisture and temperature
93 (Chambers et al., 2007; Roundy et al., 2018), but regional-scale studies are generally constrained
94 to using coarse climate data from gridded climate products. These products provide information
95 about precipitation and temperature at 0.8-4 km² spatial resolution (Daly et al., 2008; Abatzoglou
96 2013), but do not capture the shifts in soil moisture and temperature that occur with finer-scale
97 variation in topography and soil type (Hardegree et al., 2022) and which determine local
98 cheatgrass dynamics (Roundy et al., 2007, Condon et al., 2011, Bishop et al., 2019). Soil
99 moisture models use edaphic characteristics and topography, in addition to weather inputs, to
100 explicitly account for factors that modify soil microclimate (Hardegree et al., 2022) and calculate
101 soil moisture and temperature estimates at a smaller scale (10m² with current methods).

102 By combining germination models with output from a soil microclimate model, we can link
103 species-specific physiology with fine-scale information about soil temperature and moisture
104 across large spatial extents (Terry et al., 2022). Hydrothermal and thermal-germination models
105 have previously been used to predict cheatgrass response to microclimate variability at small
106 scales (Roundy et al., 2007; Rawlins et al., 2012; Hardegree et al, 2017; 2018). These models
107 predict the timing of cumulative germination response in the seedbed, but also yield rate sum
108 metrics that can be used to quantify favorability of soil microclimate for plant establishment and
109 growth (Hardegree et al., 2020). Rate sum metrics account for cumulative effects of small
110 disparate windows of germination and growth favorability that are exploited by annual plants
111 (Terry et al., 2022). Rate sum values have been shown to capture shifts in soil microclimate that
112 occur with gradients of elevation and topography, which are generally associated with ecological

113 resilience and resistance to cheatgrass invasion (Chambers et al., 2014; Roundy et al., 2018;
114 Hardegree et al., 2022).

115 The objective of this study was to combine a mechanistic understanding of cheatgrass
116 germination with soil microclimate data to predict its abundance and distribution across the
117 sagebrush steppe. Specifically, we asked, 1) can we accurately model cheatgrass distribution and
118 abundance using solely germination metrics of soil microclimate favorability? And, 2) have
119 microclimatic conditions become more favorable for cheatgrass over the last 30 years in the
120 sagebrush steppe?

121 **Methods**

122 Overview of approach

123 We used a species-specific germination model for cheatgrass to quantify favorability of soil
124 microclimate for growth as a function of soil temperature and moisture estimates in the
125 sagebrush steppe system in the western United States. We use hourly rates of germination
126 progress for cheatgrass and sum them over time to capture the favorability of soils for
127 germination and growth (Figure 1). We used monthly germination rate sums (Hardegree et al.,
128 2020) to explain field observations of cheatgrass presence (>2% cover) and abundance. We split
129 the dataset into a training and out-of-sample (OOS) testing dataset, with both spanning the
130 temporal (2002-2016) and spatial range of the dataset to test model performance. We created two
131 models, one to predict presence/absence, and one to predict abundance. After testing both models
132 on the OOS dataset, we applied the model to simulated soil conditions at all sites (training and
133 testing) from 1989 to 2019 and analyzed trends in potential distribution and abundance across all
134 sites.

135 Soil water model

136 We used the Simultaneous Heat and Water Model (SHAW) (Flerchinger et al., 2012) to generate
137 soil moisture and soil temperature estimates for each site. This model uses atmospheric, edaphic,
138 and geographic variables to model soil water and temperature as a function of soil depth (Figure
139 1). Soil texture data for each site were acquired from OpenLandMap (Hengl 2018) for the three
140 soil depths (0-10,10-30, 30-60) for each simulated field site. Other edaphic variables (field
141 capacity, bulk density, and saturated conductivity) needed for the SHAW model
142 parameterization were calculated using soil texture via Saxton equations (Saxton et al., 1986).
143 Daily temperature, precipitation, solar radiation, relative humidity, and wind variables were
144 obtained from the gridMET gridded historical climate database (4000m spatial resolution, daily
145 temporal resolution) (Abatzoglou 2013). Geographic inputs of aspect, slope and elevation were
146 derived from topographic data (10m spatial resolution) acquired from Farr et al. (2007). We did
147 not include optional vegetation inputs in the model for simplicity and for potential application to
148 post-wildfire landscapes with minimal vegetation. From this model we generated hourly
149 estimates of soil temperature (°C) and soil water potential (MPa) at 2 cm depth, a depth common
150 for seeding success in restoration settings. These estimates from the SHAW model were then
151 used as input into the wet-thermal germination models to estimate hourly and cumulative rate
152 sum values.

153 Germination model

154 Our approach utilizes germination curves that specify how hourly germination rate changes with
155 temperature when soil is wet (>-1.25 MPa). With soil moisture and temperature metrics as
156 inputs, we calculate the sum of germination rates for early spring and late fall months. We used
157 germination models with soil conditions at 2cm depth to provide hourly rate sum metrics

158 throughout the year at all sites. In this study, we used the rate sum of the 35% subpopulation of
159 seeds for our analysis (Terry et al., 2022). This metric explicitly represents the number of
160 sequential times during a given time period where conditions were sufficient for 35% of a given
161 seed population to germinate (Hardegree et al., 2020). The rate sum value for a given time period
162 is a quantitative index of seedbed favorability for germination and growth (Hardegree et al.,
163 2013; Terry et al., 2022).

164 Germination rates and rate sum calculations were based on wet-thermal germination models,
165 (Roundy et al., 2007; Rawlins et al., 2012) an approach that calculates germination rate, or the
166 relative progress toward germination during a specific time period, according to soil temperature
167 under continuously wet conditions (soil water availability < -1.25 MPa). Methods for predicting
168 germination response to temperature above threshold levels of soil water availability have been
169 previously described and validated (Roundy et al., 2007; Hardegree et al. 2018). We used data
170 from previous germination trials of cheatgrass seeds collected from 8 different field sites in the
171 sagebrush steppe (Roundy et al., 2007; Hardegree et al., 2010). We averaged germination curves
172 that explain how hourly progress toward germination varies under wet condition under different
173 temperatures. This average was done across all cheatgrass collection sites (seedlots) for the 35%
174 subpopulation to produce Equation 1. We chose the 35% subpopulation, or germination rates for
175 35% of the seeds to germinate, as this grouping captures the majority of high-quality seed with
176 the best chance for growth and fecundity (Baskin & Baskin 2014). Equation 1 calculates
177 germination rate of cheatgrass for each hour (t) as a function of soil temperature (T) when soil
178 temperature is greater than 0°C and soil water potential is greater than -1.25 MPa (Roundy et al.,
179 2007; Terry et al., 2022). Outside these conditions, germination rate was considered zero.

180 Equation 1

181 Germination Rate_t=
$$\begin{cases} 1.29 * 10^{-4} + T_t * -1.25 * 10^{-5} + T_t * 6.16 * 10^{-4}, & T > 0^{\circ}\text{C} \text{ and } MPa > -1.25 \\ 0 & \text{otherwise} \end{cases}$$
,

182

183 Cheatgrass presence and abundance data

184 We used field observations of cheatgrass presence/absence and abundance from 2662 field
185 observations collected from 2002-2016 (Appendix S1: Table S1), which was a subset of data
186 compiled by (Bradley et al., 2018). The field observations span much of the sagebrush steppe in
187 the western United States, with sites in Idaho, Utah, Nevada, California, Oregon, and
188 Washington (Figure 2). Our dataset does not include any observations in the Mojave Desert or
189 eastern portions of the sagebrush steppe in Wyoming, Montana, or Colorado. Most of the data
190 was collected using line transects, with some of the cover estimates coming from ocular
191 estimates and quadrat frames. We analyzed a subset of the data consisting of all sites with
192 measures of absolute cover (area covered by species/total geographic area) rather than relative
193 cover (% of total vegetative cover) to train and test our model to predict estimates of cheatgrass
194 cover.

195 Model Training and Testing

196 We split our data into a training set and an out-of-sample (OOS) testing set to validate the
197 accuracy of our model. We utilized 8-digit watershed units (USGS), a spatial delineation that
198 identifies watershed basins, to identify independent spatial groupings of field observations
199 (Figure 2). Altogether our data encompass 102 unique watershed units within the sagebrush
200 steppe. Field observations were randomly separated into testing (1/3 of watershed units and 29%
201 of data) and training data (2/3 of watershed units and 71% data). This resulted in a training

202 dataset composed of 884 presence points and 785 absence points and an independent testing
203 dataset composed of 533 presence points and 449 absence points. Both the training and test
204 datasets comprise field observations spanning the temporal and spatial range of our dataset
205 (Figure 2), with no repeat observations. Histograms of percent cheatgrass cover were generated
206 to ensure similar distribution of cheatgrass cover between training and testing dataset (Appendix
207 S1: Figure S1).

208 We assessed prediction accuracy based on the model's ability to predict presence (>2%
209 cheatgrass cover) and percent cover at OOS test sites in the specific year of the field observation.
210 Accuracy metrics are percentage of test sites correctly identified as present/absent and R^2 value
211 for cover predictions across OOS test sites (observed versus predicted). We chose this as our
212 comparison metric for cover predictions to allow comparison with models from previous studies,
213 which do not always report other metrics such as mean absolute error or root mean square error.

214 Cheatgrass model

215 We used a generalized additive model (GAM) from the mgcv package (Wood 2004) in R (R core
216 team 2019, version 3.6.1) to relate spatial and temporal soil favorability metrics to cheatgrass
217 presence and cheatgrass abundance. We selected a GAM modeling approach to account for non-
218 linear effects of microclimate that would require complex interaction terms in a linear model.
219 Our response variables were distribution (presence/absence) and abundance (cheatgrass cover),
220 which we analyzed in two separate models. We chose a cubic-splines smoothing approach to
221 allow knots to spread evenly throughout covariate values (Wood 2006) to ensure all
222 combinations of covariate values were considered within the smoothing terms of the model.

223 Our distribution model was a GAM binomial model, with field observations of cheatgrass cover
224 > 2% considered as species presence. Our abundance model was a GAM model with a normal
225 error distribution. We used four covariates that represent both temporal and spatial variation of
226 microclimate in the form of rate sum. Spatial covariates were mean rate sum values of spring
227 (March) and late fall (Oct-Dec) and were scaled spatially by subtracting off the mean of all sites
228 and dividing by standard deviation of all sites for each variable, such that a site with a value of 0
229 would indicate an average value relative to all the other sites. The spatial covariates describe
230 variation in climate among locations. Temporal covariates were rate sum values of spring
231 (March) and late fall (Oct-Dec) soil conditions immediately preceding the field observation.
232 These measurements were scaled temporally (across years at each site) by subtracting off the
233 site-level mean and dividing by site-level standard deviation of each variable, such that a value
234 of 0 would indicate average conditions within a given site. These temporal covariates describe
235 interannual variation in weather for each location.

236 To select the four covariates described in the previous paragraph, we first computed correlations
237 of cheatgrass cover in our training dataset with all individual monthly rate sums and groupings of
238 monthly values to seasonal sums that previous studies suggested may influence cheatgrass
239 abundance and distribution (Bradley et al., 2016; Roundy et al., 2018). After creating models
240 with the top 10 most correlative (with cheatgrass cover) rate sum metrics, we found that having
241 many non-linear parameters did not facilitate interpretation, and that by reducing the covariates
242 to the two most correlated rate sum values of late fall (Oct-Dec) and Spring (March), we could
243 retain most of the predictive power of the models while increasing interpretability. We assumed
244 that susceptibility to invasion was driven by both spatial and temporal microclimate dynamics

245 and thus included both the spatial and temporal values of these metrics to allow interannual
246 variation and average soil microclimate conditions to inform our models.

247 Models were checked for goodness-of-fit on the training data using the `gam.check` function in
248 the `mgcv` package (Wood, 2005). Specifically, we checked the basis dimensions of smoothing
249 terms to ensure that they were not so small that they force over-smoothing and checked residuals
250 for over-dispersion. While we monitored specific GAM model diagnostics, we placed more
251 emphasis on model validation with our OOS testing dataset.

252 **Results**

253 Model Performance

254 All spatial and temporal parameters within our model had significant effects ($P < 0.001$, Appendix
255 S1: Table S1 & S2) on cheatgrass distribution (presence/absence) and abundance (cheatgrass
256 cover). Accuracy metrics for the OOS testing dataset show that our mechanistic soil favorability
257 metrics covariates predicted presences with 77.8% accuracy and absences with 65.6% accuracy,
258 with an overall balanced accuracy of 71.7% (Figure 3). Soil favorability covariates predicted
259 abundance less well (R^2 0.22 testing, 0.22 training), with poor ability to distinguish high and low
260 cover sites (Figure 4).

261 Response to spatial variation in microclimate

262 Cheatgrass presence and cover responded in similar ways to spatial variation in soil microclimate
263 (Figure 5). GAM coefficient curves indicate that cheatgrass performs best in locations with
264 warmer and wetter soils in spring and fall. Cheatgrass abundance and distribution (persistence)
265 responded to average fall rate sum with a concave down shape, indicating a preference for
266 locations that on average have moderately wetter and warmer fall soils. Abundance and

267 distribution responded to spring rate sum values with a concave down but a generally positive
268 slope, indicating a preference for locations that on average have warmer and wetter spring soils
269 (Figure 5).

270 Response to temporal variation in microclimate

271 Temporal variation in fall and spring soil conditions had significant effects ($P<0.001$) on
272 interannual cheatgrass distribution and abundance (Figure 5). Wetter and warmer soils from the
273 previous fall increased likelihood of presence but benefits of warm and wet fall soils for cover
274 declined at high values, exhibiting a concave-down shape with optimal conditions occurring in
275 slightly above average years. Years with wetter and warmer spring soils were associated with
276 lower cover and probability of abundance (Figure 5).

277 Long-term trends

278 Given the reasonable predictive ability of our distribution (presence) model for our OOS testing
279 dataset, we applied our model to soil metrics at all sites ($n= 2662$) across years 1990-2019 to
280 hindcast trends in cheatgrass distribution. We found that conditions identified in our model as
281 favorable for cheatgrass presence are becoming more prevalent across our study sites (Figure 6)
282 during the period 1990-2019. Specifically, we saw evidence of climate change induced range
283 expansion, with predicted cheatgrass presence expanding 10-17% across our mid- and low-
284 elevation sites (Figure 6). Higher elevation sites remained abiotically unfavorable for cheatgrass.

285 **Discussion**

286 We developed a new modeling approach to improve understanding of an important climate-
287 species relationship and investigate how climate change has influenced the potential distribution
288 of cheatgrass over the last 30 years (1989-2019). Our results indicate that modeling the

289 distribution of an invasive annual grass species using microclimate covariates linked directly to
290 germination is a viable method to understand the relationship between climate and potential
291 distribution. The combination of site level soil-microclimate data and information about
292 germination response allowed our model to identify soil conditions in space and time that favor
293 germination and result in persistent cheatgrass presence. The relationships derived from our
294 model also allowed us to track changes in habitat suitability for cheatgrass across years.

295 Comparison to Remotely Sensed Models

296 Our model predicted the distribution of cheatgrass well, with accuracy (72%) similar to previous
297 species distribution models despite using only four microclimate-based covariates. Cheatgrass-
298 specific distribution models based on remotely-sensed covariates with similar geographic range
299 had similar performance: Downs et al., (2016) reported an overall accuracy of 71%, Bradley &
300 Mustard (2006) reported an overall accuracy of 61%, and Bradley et al., (2018) reported an
301 overall accuracy of 74%. The advantage of our approach is that it provides more direct inference
302 about the relationship between climate and cheatgrass distribution and abundance. Our model is
303 also able to capture interannual variation in climate suitability, whereas previous models that
304 correlate average climate to distribution based on several years of reflectance data (Bradley
305 2009) cannot predict variation among years.

306 Our model was unable to explain variability in cheatgrass cover as well as remotely sensed
307 models. Our model predicted percent cover with an R^2 value of 0.22, whereas other remotely
308 sensed models produced more accurate estimates of percent cheatgrass cover: Bradley et al.
309 (2018) reported R^2 value of 0.32 (3769 testing points), Peterson (2005) reported an R^2 of 0.5 (75
310 testing points), and Peterson (2006) reported an R^2 of 0.24 (806 testing points). More
311 specifically, our model failed to predict high values of cheatgrass cover, which is a similar

312 problem experienced by these other remote sensing models. The inaccuracy of our abundance
313 model probably reflects the importance of many non-climate factors, such as disturbance and
314 competition from native vegetation, in determining cheatgrass abundance.

315 Species-Climate Relationship

316 Our study indicates that areas characterized by warm and wet fall seasons (Oct-Dec) and warm
317 and wet springs (March) have the greatest abiotic potential for cheatgrass presence and
318 abundance. Fall conditions had a stronger positive effect on cheatgrass presence and abundance
319 than spring conditions (Figure 5). Previous studies have indicated hot, dry summer atmospheric
320 conditions as a key factor in cheatgrass dynamics (Bradley 2009; Bansal & Sheley 2016;
321 Brummer et al., 2016) and strong topographical effects of elevation and aspect, but our initial
322 screening of covariates indicated that fall and spring soil conditions were more informative to
323 our model than summer soil metrics. Though summer conditions are likely to affect annual
324 species indirectly by shaping competing perennial vegetation (Condon et al., 2011), summer
325 conditions should not have large effects on cheatgrass seeds that largely remain dormant until
326 germinating in the fall or early spring (Hulbert 1955; Mack & Pyke 1984). We suspect that
327 impacts of summer climate found in previous models reflect their correlation with elevation,
328 seasonal soil moisture timing, and shifts in vegetation type. Experimental findings from field
329 studies indicate that year-round warming has a positive effect on cheatgrass (Compagnoni &
330 Adler 2014; Blumenthal et al. 2016), whereas late-spring and summer warming alone had a
331 negative effect on cheatgrass cover and fecundity (Larson et al., 2017). Year-round warming
332 would increase the quantity of warm, wet conditions in the spring and fall, which is shown by
333 our model to increase suitability for cheatgrass.

334 Our results provide insight into the climate factors that generate increasing resistance to
335 cheatgrass invasion with increasing elevation. Current hypotheses link resistance to water
336 availability, soil temperature and competition (Chambers et al., 2014; 2019), but disentangling
337 the role of these abiotic factors is difficult because they are tightly correlated. Our results support
338 the conclusions of Roundy et al. (2018) that resistance to cheatgrass invasion depends on spring
339 and fall soil conditions, with colder fall and spring soils, and increasing elevation, reducing the
340 abiotic potential for cheatgrass establishment (Figure 6 & Appendix S1: Figure S2). The
341 explanatory power of our model, with explicit ties to germination, indicate further that fall and
342 spring soil conditions are important due to their influence on germination. We also anticipate,
343 that fall and spring soil conditions are linked to cheatgrass invasion, because they may allow
344 cumulative periods of growth, captured by the summing nature of our microclimate metric, that
345 can be utilized by a winter annual grass for early growth.

346 Our models generally indicate that warmer and wetter soil in the fall and spring periods, in both
347 space and time, increase probability of cheatgrass presence, and to a lesser degree, cheatgrass
348 cover (Figure 5). However, the effects of temporal anomalies in spring conditions did not follow
349 this pattern and indicated that warmer and wetter conditions decreased cheatgrass presence
350 (Figure 5). We hypothesize that this discrepancy is indicative of the complex relationships
351 between spring soil microclimate and cheatgrass dynamics. Though warm and wet spring soils
352 have been shown to be beneficial to cheatgrass, they also decrease the likelihood of fire (Pilliod
353 et al., 2017), a factor strongly linked to cheatgrass distribution (Bradley et al., 2018). Without
354 accounting for fire or spatial factors that determine the abundance of competing native flora, our
355 model predicted anomalies in spring soil microclimate to be the least informative parameter in
356 both our abundance and distribution models. We anticipate that accounting for interactive effects

357 of spatial factors of fire and native species composition with spring soil conditions would not
358 only improve model fit but would also indicate a strong positive effect of warm and wet spring
359 soils when native plant cover is low. This would support findings from Bradford & Lauenroth
360 (2006) showing that the effect of temporal conditions only becomes important in scenarios where
361 spatial factors such as total plant cover and disturbance history allow a sizeable response to
362 interannual variation in weather.

363 The contrasting accuracy of our distribution and abundance models suggests that different factors
364 control cheatgrass distribution and abundance. Our model is based on the abiotic factors that
365 directly influence germination. The success of this model in explaining cheatgrass
366 presence/absence indicates a primary role for germination and periods of wet and warm shallow
367 soil microclimate. In contrast, the low explanatory power of our model for abundance indicates
368 the importance of other biotic and abiotic factors likely unrelated to germination and shallow soil
369 microclimate. This fits well with the conclusion of (Bradford & Lauenroth 2006) that climate
370 drives susceptibility to annual grass invasion, and disturbance regime dictates severity of
371 invasion. There are many studies that indicate the positive impact of disturbance, especially fire,
372 on annual grass abundance (D'Antonio & Vitousek 1992; Condon et al., 2011; Bradley et al.,
373 2018; Gill et al., 2018; Fusco et al., 2019; Williamson et al., 2020). Failure to account for
374 disturbance history or competitive interactions may limit our model's ability to distinguish
375 between high and low cheatgrass cover, primarily because cover of competitive species and lack
376 of disturbance may limit propagule pressure and thus complicate species response of annual
377 plants species to favorable soil conditions.

378 Trends in Cheatgrass Distribution

379 Our analysis suggests that climate change has already benefitted cheatgrass and expanded its
380 potential range 10-17% across low and mid-elevation sites (Figure 6). These results are
381 consistent with trends found in remotely sensed data by Smith et al. (2022), showing an increase
382 in annual grass dominance across sites with elevation < 2100m, and Pastick et al. (2021), who
383 found similar increases in cheatgrass distribution and cover across low and mid-elevation sites.
384 Understanding new changes in distribution is critical due to potential positive feedback where
385 even small amounts of cheatgrass (<10% cover) have been linked with heightened wildfire risk
386 (Pastick et al., 2021), which can quickly lead to post-wildfire dominance of invasive annual
387 grasses and more subsequent fires (D'Antonio & Vitousek 1992; Bradley et al., 2018).

388 Benefits of our Approach

389 Our approach only considers soil moisture and temperature metrics known to directly influence
390 cheatgrass germination. This sets up a relatively simple model with few covariates to describe a
391 site's potential for cheatgrass compared to current remotely sensed models (Bradley et al., 2018;
392 Pastick et al., 2021). Using a model with fewer covariates also simplifies interpretation. Because
393 our approach explains field observations of cheatgrass distribution and abundance solely as a
394 function of microclimate covariates, we are able to understand climate constraints and
395 preferences of cheatgrass without introducing additional uncertainty that occurs when estimating
396 cheatgrass distribution based on remotely-sensed imagery. In addition, our model can explain
397 interannual variation in cheatgrass suitability at a single site, in contrast to models limited to
398 inference of mean climate conditions at each site.

399 Our results may be useful for natural resource management as it indicates not only which
400 locations that may be vulnerable to cheatgrass invasion, but also provides the tools to understand
401 which new areas may become vulnerable with current trajectories of climate that alter climatic

402 suitability for cheatgrass presence. This allows managers to mitigate activities in locations that
403 have recently become vulnerable to cheatgrass invasion but also provides time for management
404 to prepare for future invasion vulnerability.

405 Potential Limitations of our Approach

406 The soil microclimate approach we used is appropriate for systems where water availability is a
407 key limiting factor but may not provide meaningful information regarding a species' potential in
408 systems where resource availability or survival are not explicitly tied to soil moisture and soil
409 temperature metrics. Our study also focuses on a species with an annual life history, meaning
410 that annual germination and growth favorability metrics are very relevant to each year's
411 distribution and abundance. Perennial species could be less sensitive to factors regulating
412 germination and seedling performance. We suspect that distribution and abundance of perennial
413 species will have different microclimate requirements with lag-effects of favorable or
414 unfavorable conditions being important. Finally, our approach depends on lab trials to generate
415 germination curves, and thus requires more resources than remote sensing approaches to map
416 distribution, though it remains unknown if rate sum values from one species are adequate
417 estimates of soil favorability for other species. Currently, it is best suited for understanding
418 climate-species relationships or predicting susceptibility to invasion.

419 Our model indicates whether cheatgrass persistence may be a possible due to relationship with
420 climate but does not include disturbance, a major driving factor behind exotic annual grass
421 dominance (Bradley et al., 2018; Fusco et al., 2019; Pastick et al. 2021). While our approach
422 illustrates that the influence of climate alone may determine distribution and persistence of
423 cheatgrass, we encourage future effort to study how soil favorability metrics interact with
424 disturbance and competitive native plant abundance to better understand how/where climate and

425 disturbance may interact to create systems dominated by exotic annual grasses, where impacts of
426 invasion are most severe.

427

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433

434 **Author Contributions**

435 TJT and SPH conceived the study; TJT and PBA developed and ran statistical models; TJT and
436 SPH developed and ran germination models; TJT led writing with comments and multiple
437 rounds of feedback from all co-authors.

438 **Conflict of Interest Statement**

439 The authors declare no conflicts of interest.

440 **References**

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661

662 **Figure captions**

663 Figure 1. Flowchart indicating inputs used to estimate soil microclimate conditions and how they
664 are combined with germination models to produce monthly rate sum values.

665 Figure 2. Location of all field observations (right) and an example of how watershed units were
666 used to split the data into training (blue points) and testing (pink) sets. This approach allowed us
667 to test our model on a dataset that matches the spatiotemporal extent of the training dataset while
668 ensuring a degree of independence between the two datasets.

669 Figure 3. Map of model predictions for cheatgrass presence/absence using our final model at
670 both testing and training sites.

671 Figure 4. Predictions of Cheatgrass cover versus actual values on independent (OOS) test
672 dataset. Line represents 1:1 ratio.

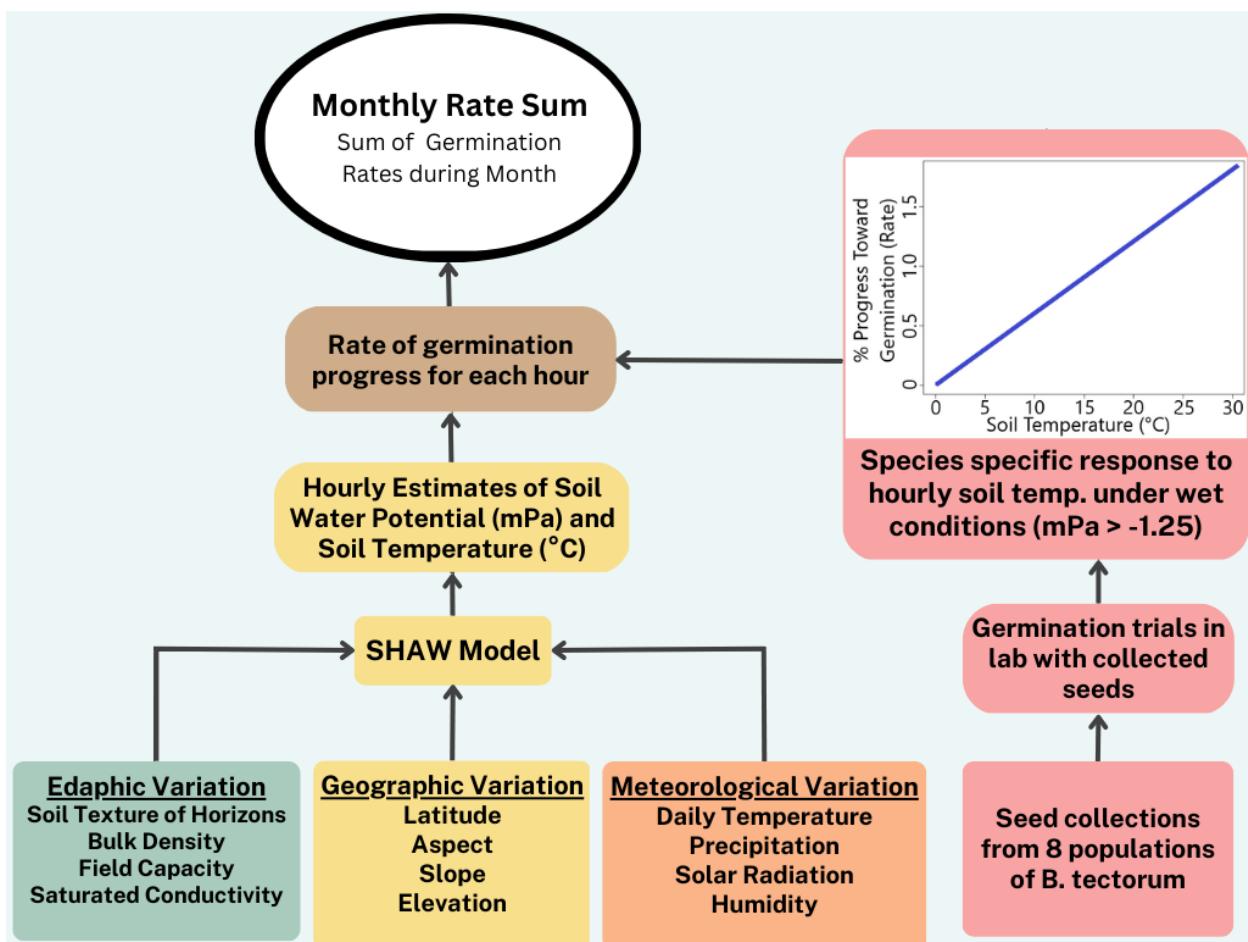
673 Figure 5. Plots showing smoothed parameter effects. Y-axis values indicate magnitude of the
674 effect, and X-axis values indicate possible parameter values within the dataset. Blue values
675 represent likelihood (log odds) of cheatgrass presence at a site. Red values indicate impacts on
676 cheatgrass cover. Top panel includes average rate sum values that were scaled spatially with
677 other sites in the dataset. Bottom panel includes recent rate sum values corresponding to the
678 annual conditions during the year of observation. Recent rate sum values were scaled within site
679 to indicate deviations from a site's respective mean.

680 Figure 6. Predicted distribution of cheatgrass across sites according to elevation and aspect. Each
681 dot represents the percent of total sites predicted to be occupied by cheatgrass in each year with

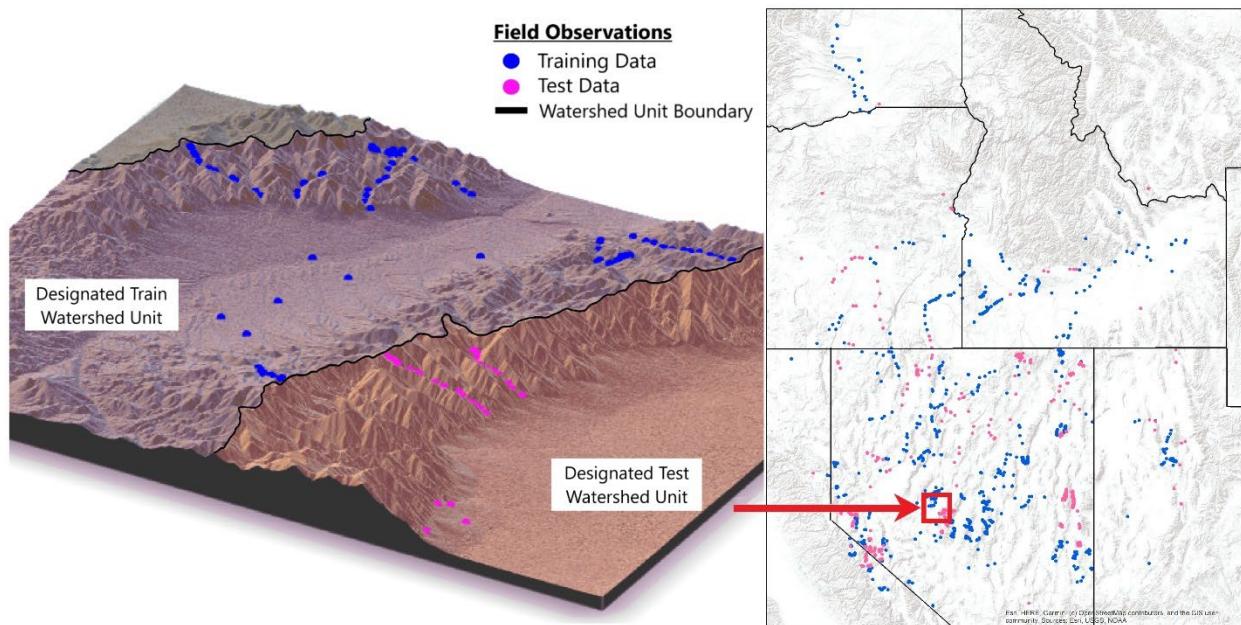
682 color indicating different aspects. The elevation categories represent the bottom, middle, and top
683 tercile of the dataset.

684

685 Figure 1



689 Figure 2

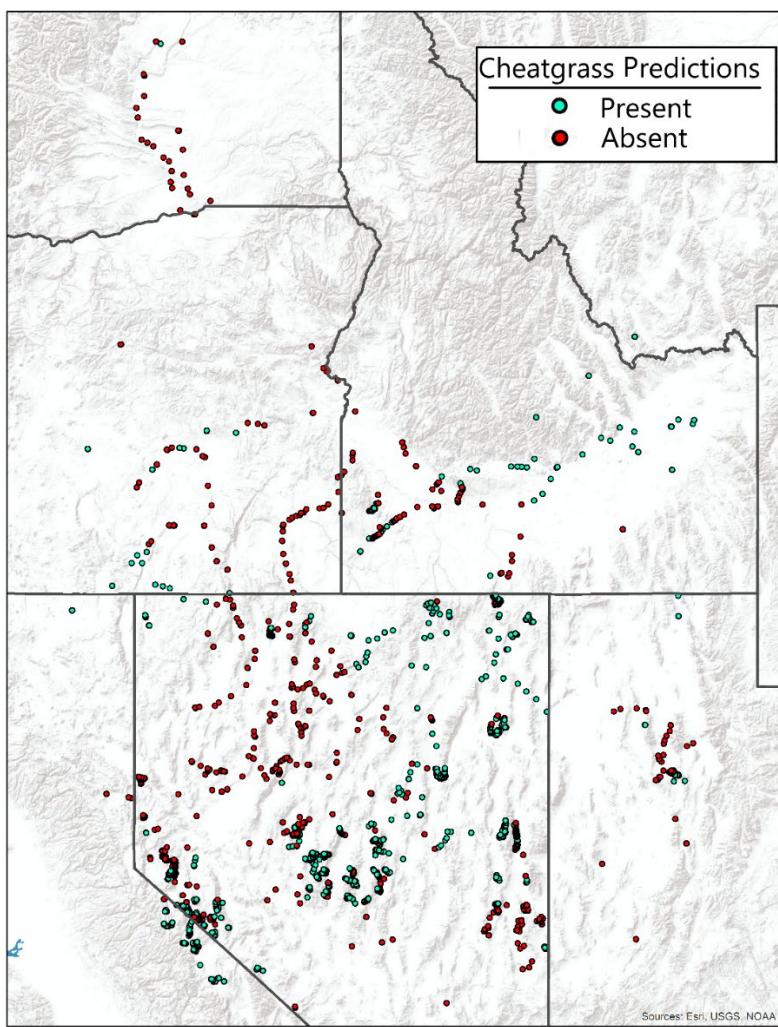


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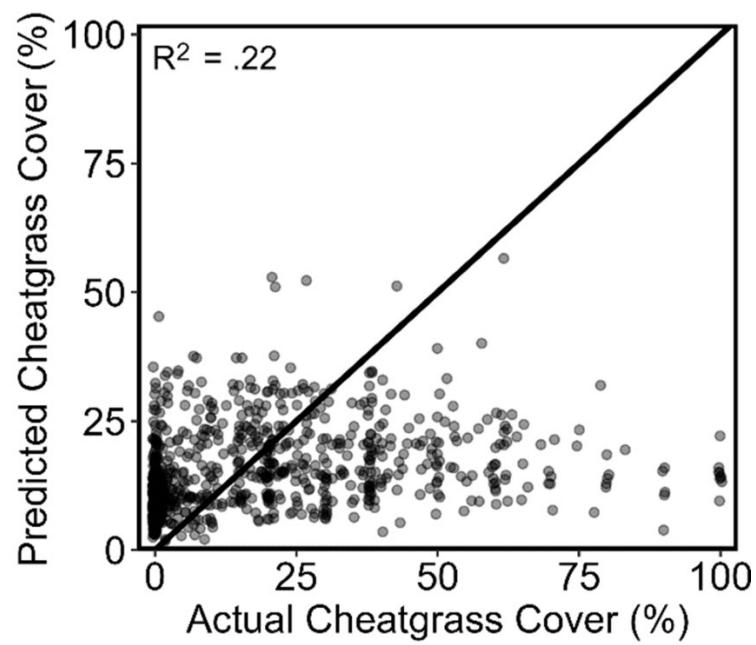
693 Figure 3



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696 Figure 4



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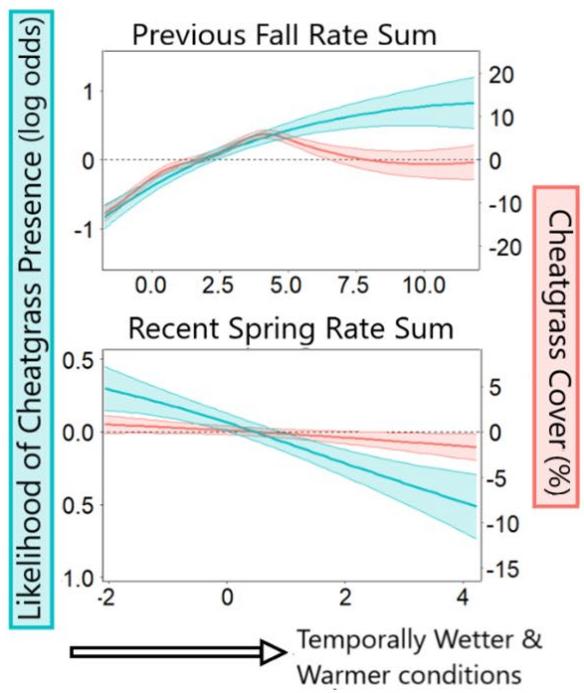
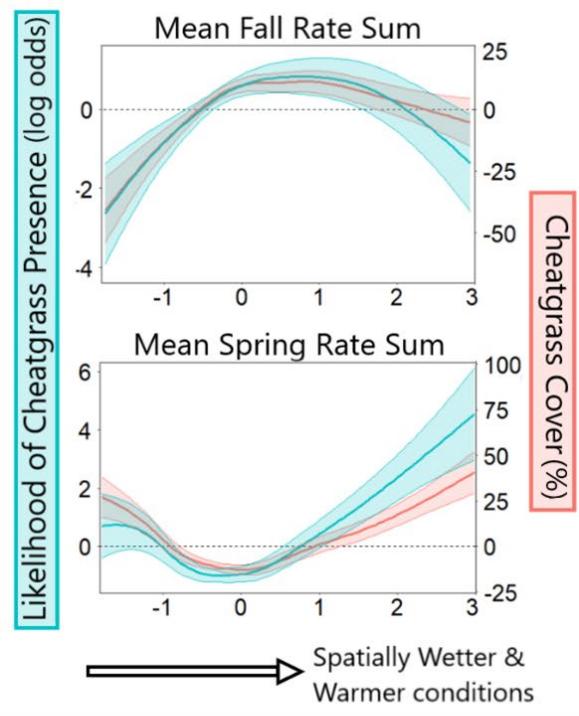
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706 Figure 5



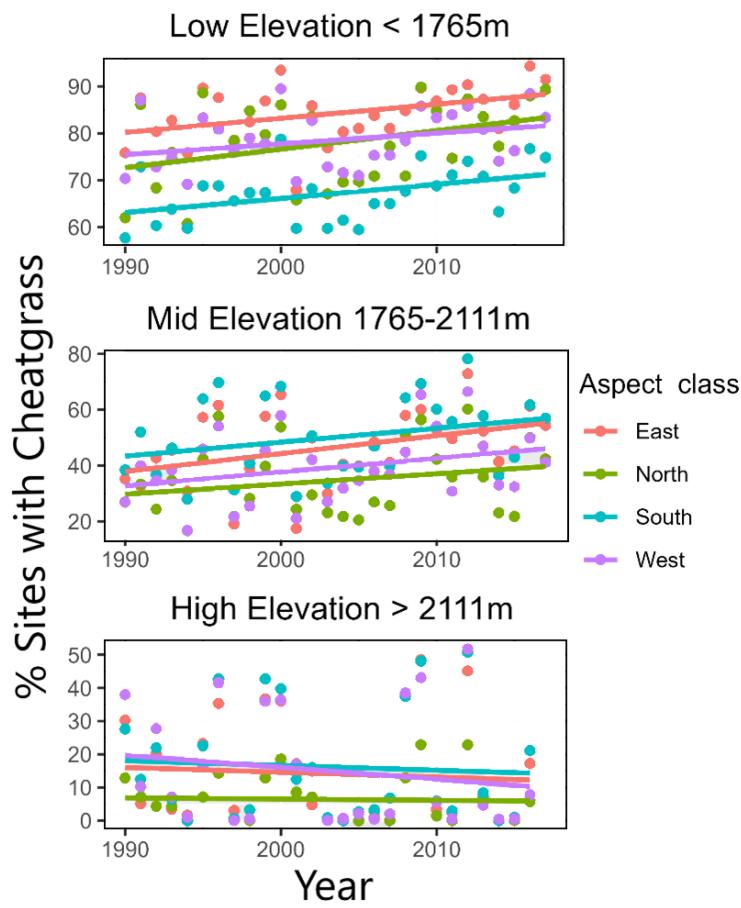
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711 Figure 6



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