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## Contents

# Expedition 361 summary<sup>1</sup>



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## Abstract

International Ocean Discovery Program Expedition 361 drilled six sites on the southeast African margin (southwest Indian Ocean) and in the Indian-Atlantic Ocean gateway, from 30 January to 31 March 2016. In total, 5175 m of core was recovered, with an average recovery of 102%, during 29.7 days of on-site operations. The sites, situated in the Mozambique Channel at locations directly influenced by discharge from the Zambezi and Limpopo River catchments, the Natal Valley, the Agulhas Plateau, and Cape Basin, were targeted to reconstruct the history of the greater Agulhas Current system over the past ~5 My. The Agulhas Current is the strongest western boundary current in the Southern Hemisphere, transporting some 70 Sv of warm, saline surface water from the tropical Indian Ocean along the East African margin to the tip of Africa. Exchanges of heat and moisture with the atmosphere influence southern African climates, including individual weather systems such as extratropical cyclone formation in the region and rainfall patterns. Recent ocean model and paleoceanographic data further point at a potential role of the Agulhas Current in controlling the strength and mode of the Atlantic Meridional Overturning Circulation (AMOC) during the Late Pleistocene. Spillage of saline Agulhas water into the South Atlantic stimulates buoyancy anomalies that may influence basin-wide AMOC, with implications for convective

activity in the North Atlantic and global climate change. The main objectives of the expedition were to establish the role of the Agulhas Current in climatic changes during the Pliocene–Pleistocene, specifically to document the dynamics of the Indian-Atlantic Ocean gateway circulation during this time, to examine the connection of the Agulhas leakage and AMOC, and to address the influence of the Agulhas Current on African terrestrial climates and coincidences with human evolution. Additionally, the expedition set out to fulfill the needs of Ancillary Project Letter number 845, consisting of high-resolution interstitial water sampling to help constrain the temperature and salinity profiles of the ocean during the Last Glacial Maximum.

The expedition made major strides toward fulfilling each of these objectives. The recovered sequences allowed generation of complete spliced stratigraphic sections that range from 0 to between ~0.13 and 7 Ma. This sediment will provide decadal- to millennial-scale climatic records that will allow answering the paleoceanographic and paleoclimatic questions set out in the drilling proposal.

## Introduction

The Agulhas Current is a fundamental component of the South Indian Ocean subtropical gyre circulation and constitutes the stron-

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<sup>2</sup> **Expedition 361 Scientists' addresses.**

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Figure F1. Global composite of land surface reflectance and sea-surface temperature (credit: Jacques Descloitres, MODIS Land Science Team, NASA) with schematic representation of the greater Agulhas Current system and its links to the Atlantic Meridional Overturning Circulation. The greater Agulhas Current system around South Africa consists of the Agulhas Current flowing south along the African shelf. Off the southern tip of Africa, the Agulhas Current abruptly retroflects and feeds back into the Indian Ocean. During this process, huge rings of water (Agulhas rings) are separated, drifting into the Atlantic and representing “salt leakage.”

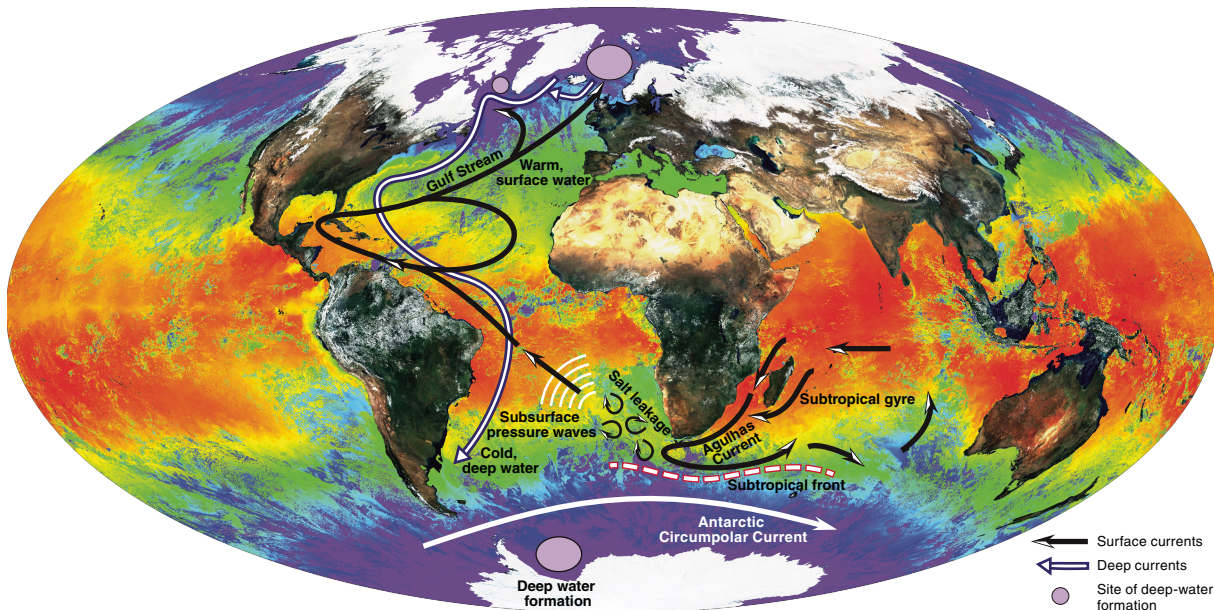
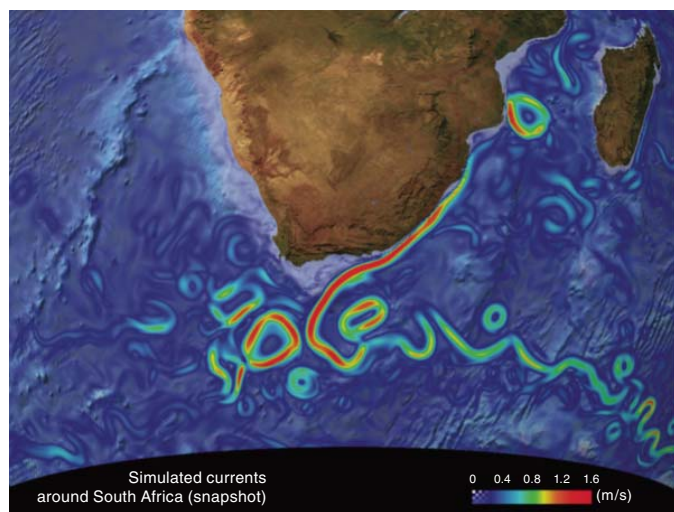


Figure F2. Surface currents of the greater Agulhas Current system represented in a high-resolution ocean model (snapshot). Upstream of the Agulhas Current, eddies around Madagascar lead to short-term offshore meanders of the main current by several hundred kilometers. Credit: Professor Dr. Arne Biastoch, GEOMAR Helmholtz Centre for Ocean Research Kiel, Germany.



gest western boundary current in the Southern Hemisphere (Lutjeharms, 2006a) transporting  $\sim 70$  Sv ( $1 \text{ Sv} = 10^6 \text{ m}^3/\text{s}$ ) of warm, saline tropical surface waters to the tip of Africa (Figures F1, F2), roughly twice that of the Gulf Stream in the North Atlantic (Cunningham et al., 2007). Like its northern counterpart, the Agulhas Current fosters exchanges of heat and moisture with the atmosphere that influence regional climates and weather systems (Reason and Mulenga, 1999). A portion of the Agulhas waters invade the South Atlantic through the Agulhas rings (Arhan et al., 2011), which form when

the current retroflects at the tip of Africa and transport between 5 and 20 Sv of warm, saline water from the Indian Ocean to the South Atlantic—the so-called interocean “warm-water route” of the global oceanic thermohaline circulation (THC). Ocean models and paleo-oceanographic data strongly suggest that the “leakage” of saline Agulhas water into the South Atlantic stimulates regional buoyancy anomalies that ultimately impact convective activity in the northern North Atlantic, plausibly acting as a contributor to the basin-wide Atlantic Meridional Overturning Circulation (AMOC) (Figure F1) (Weijer, 2001; Weijer et al., 2002; Knorr and Lohmann, 2003; van Sebille et al., 2009; Biastoch et al., 2008, 2009a). Numerical simulations specifically point to Agulhas leakage as a potential modulator of the AMOC and a rheostat for the ventilation of the deep ocean. Hence the significance of the Agulhas Current is threefold: (1) the Agulhas Current is a significant component of the global THC-carrying surface water into the Atlantic that compensates for the export of deep water from the Atlantic Basin to the world ocean, (2) the Agulhas leakage transfer of salt to the South Atlantic variably impacts the buoyancy structure there and may contribute to mode changes and variability of the AMOC and thus climate variability in the North Atlantic region and beyond, and (3) variations in Agulhas warm-water transports along the southeast African continental margin drive interannual variability with implications for southern Africa regional climate and weather systems.

Studies of the long-term evolution of ocean circulation off southern Africa, on Pliocene–Pleistocene timescales, have previously been undertaken on the Atlantic side of the continent. Deep Sea Drilling Project (DSDP) Legs 40 and 75 in 1974 and 1980 and Ocean Drilling Program (ODP) Legs 108 and 114 in 1985 and 1987 targeted the southwest African continental slope and rise north and south of Walvis Ridge to explore marine conditions during the opening of the South Atlantic, including the origin and paleo-oceanographic history of the Eastern Walvis Ridge and carbonate dissolution cycles. More recently, ODP Leg 175 in 1997 targeted the



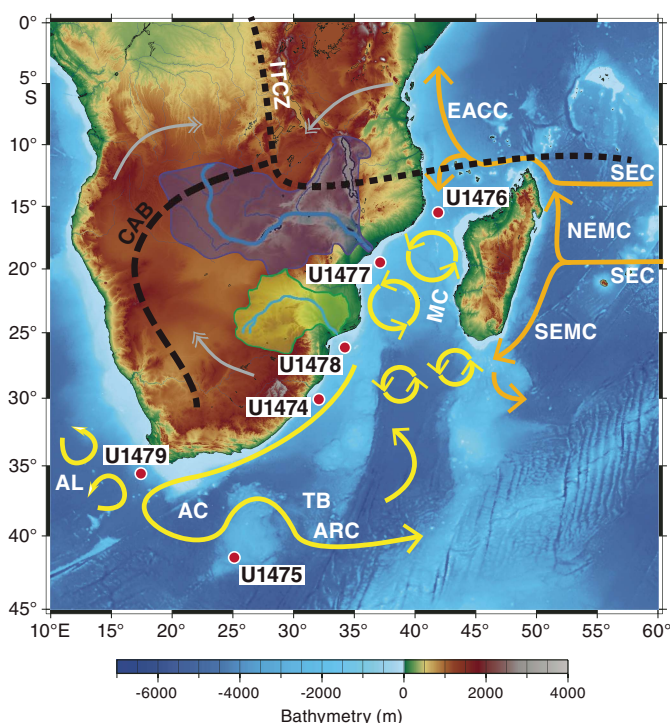
southeastern Atlantic margin between 5° and 32°S to assess the history of the Benguela Current and upwelling off Angola and Namibia. ODP Leg 208 in 2003 drilled a transect of six sites on Walvis Ridge to gain insight into paleoceanographic events such as the hyperthermal events of the Paleocene and Eocene and the early Oligocene Oi1 cold event.

By contrast, the eastern continental margin off southern Africa is poorly represented in the database of scientific drill sites. Prior to Expedition 361, only one site, DSDP Site 244, had been drilled within the direct reaches of the South Indian Ocean western boundary current. The site is positioned on Davie Ridge, at the northern entrance of the Mozambique Channel, and was occupied during Leg 25 in 1972 (Simpson, Schlich, et al., 1974) to provide data on the structure and geological history of the East African margin and western Madagascar.

Expedition 361 fills this scientific/geographic gap by drilling targets in three sectors of the greater Agulhas Current system (Figure F3):

1. The northern sector in the Mozambique Channel that provides the source waters of the Agulhas Current,
2. The central sector in the confluence region where the southwestern extension of the East Madagascar Current converges

Figure F3. Expedition 361 site locations. Also shown is a schematic representation of the main surface ocean currents (yellow and orange arrows) in the southwest Indian Ocean and atmospheric circulation over southern Africa during austral summer (December–February) with approximate position of the Intertropical Convergence Zone (ITCZ) and the Congo Air Boundary (CAB) (dashed lines; adapted from Reason et al., 2006). AC = Agulhas Current, SEC = South Equatorial Current, SEMC = South East Madagascar Current, NEMC = North East Madagascar Current, EACC = East Africa Coastal Current, MC = Madagascar Current, ARC = Agulhas Return Current, AL = Agulhas leakage, TB = Transkei Basin. Purple shaded area = Zambezi Catchment, green shaded area = Limpopo Catchments. Gray arrows = main pathways of moisture supply to the African continent from the northwest Atlantic (via Congo) and the northwest and southwest Indian Ocean.



with the Mozambique Channel throughflow to form the Agulhas Current, and

3. The southern sector in the Indian-Atlantic Ocean gateway at the tip of Africa with Agulhas retroflexion, Agulhas Return Current (ARC), and Agulhas leakage.

Analysis of the sediments from Expedition 361 will help to address the scientific objectives outlined below.

## Background

### Motivation for drilling in the greater Agulhas Current system

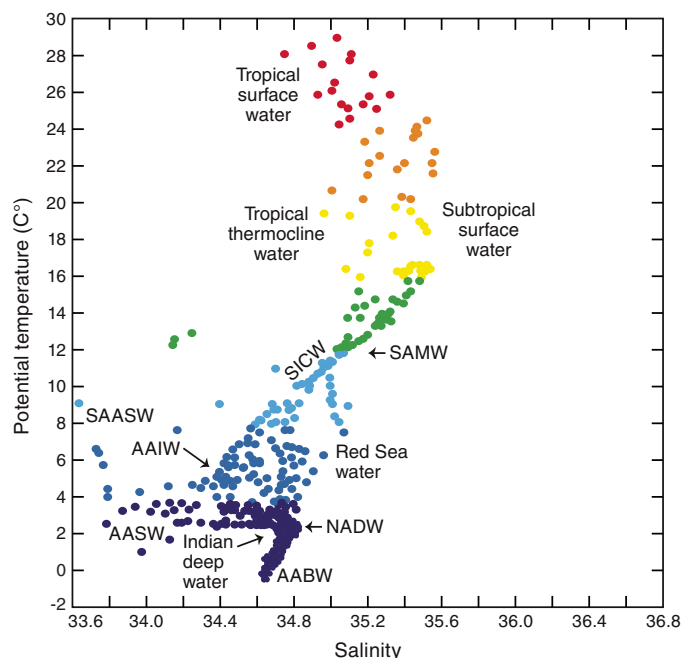
#### The Agulhas Current system—modern oceanographic circulation

The Agulhas Current constitutes the largest ocean western boundary jet in the Southern Hemisphere. The current is part of the warm-water return flow that feeds the renewal of North Atlantic Deep Water (NADW) and is a key component of the global oceanic thermohaline “conveyor belt” circulation (Gordon, 1986; Weijer et al., 1999; Lutjeharms, 2006a). The Agulhas Current carries thermocline waters from the South Indian Ocean subtropical gyre with contributions from the Red Sea and Arabian Sea and from the Indonesian Throughflow (Lutjeharms, 2006a) (Figure F4). Driven by the anticyclonic wind field over the South Indian Ocean, Agulhas Current water is advected by eddies and meanders from the equatorial Indian Ocean through the Mozambique Channel and the East Madagascar Current (EMC) (Song et al., 2004; Beal et al., 2006). Interaction between Madagascar and westward-traveling Rossby waves from the tropical Indian Ocean contribute to the formation of eddies (Schouten et al., 2002a) east of Madagascar (Biaosch and Krauss, 1999). Four to five of these eddies per year, as wide as 350 km, drift southward through the Mozambique Channel, with a mean transport of 17 Sv (Ridderinkhof et al., 2010). These eddies impose short-term circulation modulation over the entire depth and width of the channel (de Ruijter et al., 2002) and impart an interannual variability of ~9 Sv to the Agulhas Current (Biaosch et al., 2009a).

The ocean/atmosphere coupling in the tropical Indian Ocean provides important upstream forcing of the Agulhas Current (Ridderinkhof et al., 2010). The South Equatorial Current (SEC), splits into a southern and northern branch, once reaching Madagascar, forming the South East Madagascar Current (SEMC) and North East Madagascar Current (NEMC) (Figure F3). The SEMC has an estimated average transport of 20 Sv in its northern reaches (Donohue and Toole, 2003; Schott et al., 1988), increasing to 35 Sv at the southern tip of Madagascar (Nauw et al., 2008). At the latitude of Durban (30°S), the westward-flowing southern extension of the SEMC converges with the southward-flowing train of Mozambique Channel eddies to form the Agulhas Current as a narrow jet-type current that is steered by the steep South African continental margin. Fully constituted at 32°S, the Agulhas Current carries ~70 Sv of warm (24°–17°C) and saline (35.5–35.2 psu) (Gordon and Haxby, 1990; Gordon et al., 1992; Beal and Bryden, 1999; Bryden and Beal, 2001) waters of the South Indian Ocean thermocline (Bryden et al., 2005). In its deeper parts, the current also includes saline Red Sea Intermediate Water (Gordon et al., 1987) and other waters derived from far-distant sources in the Pacific Ocean (Gordon and Fine, 1996; Speich et al., 2002) (Figure F4).

The Agulhas Current’s mean trajectory is steered by the bathymetry of the narrow southeast African shelf where its steep conti-

Figure F4. Potential temperature vs. salinity for the western Indian Ocean (adapted from Gordon et al., 1987). SICW = South Indian Central Water, SAMW = Subantarctic Mode Water, NADW = North Atlantic Deep Water, AABW = Antarctic Bottom Water, AAIW = Antarctic Intermediate Water, AASW = Antarctic Surface Water, SAASW = Subantarctic Surface Water.

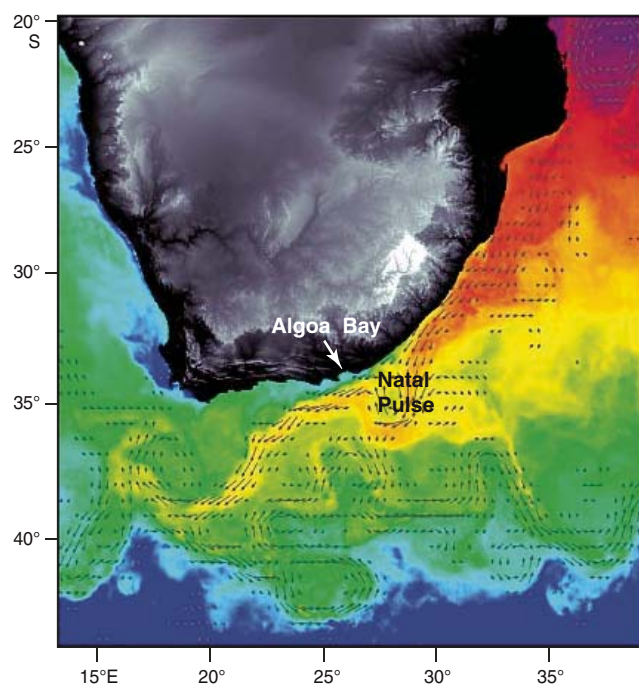


nental slope provides for maximum barotropic and baroclinic guidance (de Ruijter et al., 1999b; Lutjeharms and de Ruijter, 1996; Goschen and Schuman, 1990). Instabilities of the current are caused by the migration of eddies and meanders that interact with the bathymetry and morphology of the continental margin. An example is solitary meanders that form in the Natal Bight, the so-called Natal Pulses (Figure F5) (e.g., Lutjeharms and Roberts, 1988; Beal and Bryden, 1999; de Ruijter et al., 1999b; van Leeuwen et al., 2000; Bryden et al., 2005; Biastoch et al., 2008; Tsugawa and Hasumi, 2010; Rouault and Penven, 2011). At the tip of Africa, the current veers briefly to the west, driven by its own inertia, and then retroflects back to the east into the Indian Ocean as the ARC (Lutjeharms and Van Ballegooyen, 1988). Mozambique Channel and EMC eddies that propagate downstream with the Agulhas Current as strongly barotropic solitary meanders (Bryden et al., 2005; Biastoch et al., 2009a; Tsugawa and Hasumi, 2010) destabilize the retroflection. Four to six anticyclonic eddies (Agulhas rings) form per year with a radius of up to 400 km; they spin off the retroflection area carrying between 2 and 20 Sv of thermocline water to the South Atlantic Ocean (de Ruijter et al., 1999a; Richardson, 2007). Some researchers have suggested that Agulhas leakage is linked with interannual modes of Indian and Pacific Ocean variability, thus implying a connection with the Indian Ocean Dipole and Pacific Ocean La Niña/El Niño phases (Schouten et al., 2002b; Palastanga et al., 2006).

#### Agulhas leakage and the AMOC

The dynamics of the Agulhas leakage and its possible impact on the AMOC and global climate have been explored in a number of ocean and climate models. These studies suggest that (1) the transfer affects the buoyancy of Atlantic Ocean thermocline waters, potentially influencing NADW formation on multidecadal timescales

Figure F5. Representative SST map overlaid with geostrophic current vectors from daily Odyssea SST analysis provided by the CERSAT and the Ssalto/Duacs absolute geostrophic velocities satellite products. Flow speeds <30 cm/s were not plotted. The loop or meander in the current shows a well-developed Natal Pulse just north of Algoa Bay. Image courtesy of Marjolaine Krug (CSIR-NRE).



(Weijer et al., 2002; Knorr and Lohmann, 2003; Biastoch et al., 2009a) and (2) perturbations of planetary (Kelvin and Rossby) waves by mesoscale eddies propagate across the Atlantic Basin influence variability of the AMOC (van Sebille and van Leeuwen, 2007; Biastoch et al., 2008, 2009a).

Short-term variability that appears to be associated with Agulhas leakage is observed in the subtropical North Atlantic Ocean (de Ruijter et al., 1999a), whereas some of the interannual to decadal AMOC variability seen in the models plausibly stems from the frequency of Agulhas ring shedding that, in turn, is modulated by tropical interannual variability (de Ruijter et al., 1999a; Gordon and Haxby, 1990). The basin-scale radiation of planetary waves triggered by Agulhas leakage hence constitutes a “fast-response” mechanism that operates on interannual to decadal timescales.

The advection of saltwater from the Agulhas region to the convection centers in the North Atlantic Ocean is the likely connection between the AMOC response and Agulhas leakage. Seeding numerical Lagrangian floats into the Agulhas retroflection area in high-resolution ocean general circulation models reveals an advective transit time of 30–40 y from the Agulhas Current to the North Atlantic Ocean (van Sebille et al., 2010). These time estimates are also consistent with those obtained with lower-resolution models (Weijer, 2001). Continuing heat exchange with the overlying atmosphere augments the density anomaly associated with leaked Agulhas saltwater parcels as they drift northward (Weijer, 2001; Haarsma et al., 2011). Agulhas leakage therefore stimulates an AMOC response with a time delay dictated by the advective timescale for the excess salt to reach the North Atlantic Ocean; the new AMOC equilibrium state is achieved several hundred years later when buoyancy fluxes and mixing are in balance (Weijer, 2001; Hughes et al., 2009). Estab-



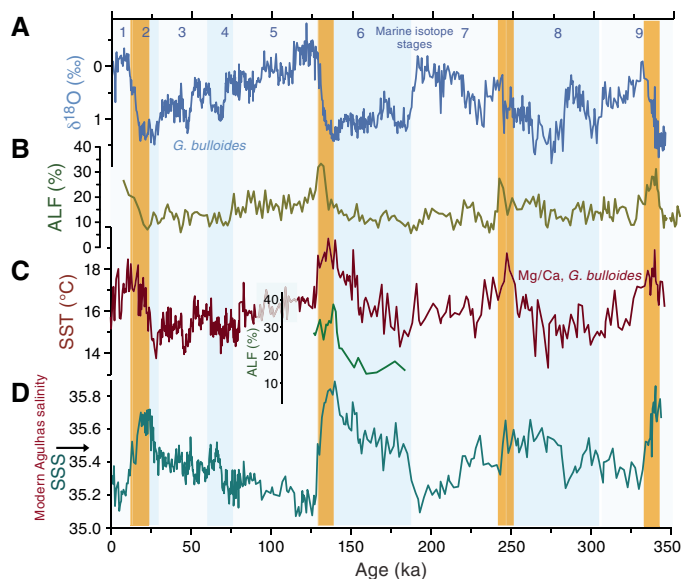
lishing the impact of wind forcing on Agulhas leakage and its subsequent effect on the AMOC is not straightforward because the strength and position of the Southern Hemisphere mid-latitude westerlies may affect the Agulhas leakage and AMOC simultaneously (Biaostoch et al., 2009b; Toggweiler and Samuels, 1995; Sijp and England, 2008), perhaps by shifting the position of the Subtropical Front (STF) south of Africa. For example, the STF is thought by some to define the width of the Agulhas leakage corridor, hence the strength of the leakage. On the other hand, recent modeling studies show a stronger dynamic relationship with the intensity of the Southern Hemisphere westerlies (Durgadoo et al., 2013). Nonetheless, the westerlies control the northward Ekman transport in the Southern Ocean, which forces deep water to ascend to the upper ocean, conceivably “pulling” the deep branch of the AMOC (Toggweiler and Samuels, 1995; Visbeck, 2007). Through their influence on deep-water upwelling and AMOC strength, the westerly winds and Agulhas buoyancy fluxes likely play important roles in the regulation of the ocean's ability to store carbon (Watson and Garabato, 2006; Backeberg et al., 2012). This connection represents an additional plausible link between Agulhas leakage and global climate variability.

### Paleoceanography of the Agulhas Current system

The Agulhas Current region is an area in which retrieval of high-quality sediment cores is difficult. The southeast African margin has a very narrow shelf and a steep continental slope, leading to frequent sediment instability and slumping. Further, the vigorous Agulhas Current and its mesoscale variability can cause sediment redistribution because meanders and eddies frequently reach the seafloor. Yet, a number of published studies and emerging work demonstrate that paleoceanographic records of excellent quality are in fact achievable for reconstructing both surface and deep-water variability in the region. A full synthesis of the paleoceanographic evidence is yet to emerge, but several aspects of the Agulhas paleo-time series reflect broad consensus. Based on published records, Agulhas leakage has been a persistent feature of the past 1.3 My, with less leakage during full glacial times than interglacial times and greatest leakage at the glacial terminations (Peeters et al., 2004; Caley et al., 2012; Petrick et al., 2015). There is also general agreement about evidence for a more northerly position of the STF during glacial times (Rau et al., 2002, 2006; Bard and Rickaby, 2009; Martínez-Méndez et al., 2010) and some suggestion that this may play a role in reducing Agulhas leakage. However, it is also possible that the positions of the westerly wind belt and STF can be decoupled (Graham and De Boer, 2013; De Boer et al., 2013; Durgadoo et al., 2013), leading to some uncertainty in the exact mechanisms of changing Agulhas leakage.

At the tip of Africa, a spliced record composed of Cores GeoB-3603-2 and MD96-2081 (termed the Cape Basin Record; Peeters et al., 2004) shows maximum abundance of the so-called “Agulhas leakage fauna” (ALF) that appears to be related to minima in precession index (maxima in northern summer insolation), leading these authors to suggest a possible teleconnection to the monsoon system. Multiple-proxy data from planktonic foraminifers from sediment Core MD96-2080 on the Agulhas Bank provide a detailed reconstruction of surface-ocean conditions in the area. Continued warming across full-glacial stages (marine isotope Stages [MIS] 2 and 6) displayed in planktonic Mg/Ca data implies increased influence of warm Agulhas water in the region and/or reduced northward advection of cold subantarctic surface water (Martínez-Méndez et al., 2010) (Figure F6). This pattern is consistent with the

Figure F6. Palaeoceanographic records from sediment Core MD96-2080 (close to Site U1479). A. Planktonic  $\delta^{18}\text{O}$  showing glacial–interglacial climates. B. Distribution of ALF from the Cape Basin records from Peeters et al. (2006). C. SST variation along Core MD96-2080 from planktonic Mg/Ca ratios. Insert displays ALF section from the same core at the Stage 6–5 transition. D. Sea-surface salinity (SSS) estimated from paleothermometry combining data from A and C. Data from Peeters et al. (2004) and Martínez-Méndez et al. (2010).



hypothesis that peak Agulhas leakage occurred during glacial terminations and plausibly aided the AMOC to shift to its full-strength interglacial mode (Knorr and Lohmann, 2003, 2007; Peeters et al., 2004). A high-resolution faunal record of the accumulation rate of tropical species *Globorotalia menardii* at ODP Site 1087 in the southern Benguela region was used to infer that Agulhas leakage strengthened during every glacial termination of the past 1.3 My (Caley et al., 2012). The interpretation of increased leakage during terminations in the Late Pleistocene is also supported by the combination of alkenone proxies ( $U_k^{37}$  and  $\delta D$ ) along with dinoflagellate cyst assemblages, sea-surface temperature (SST), and salinity estimates (Petrick et al., 2015).

SST and salinity records upstream along the southeast African margin (Core MD96-2048; 26°S) display prominent 41 ky cyclicality that was interpreted to suggest that long-term Agulhas Current variations are associated with high-latitude (as opposed to regional tropical) climate forcing (Caley et al., 2011). However, the cyclicality also suggests that some of the apparent variability in Agulhas leakage could actually be due to changes in the Agulhas Current. Thus, to circumvent the potential problem that ALF (Peeters et al., 2004) and *G. menardii* (Rau et al., 2006; Caley et al., 2012) indexes might be sensitive to the changing composition of the Agulhas source assemblage upstream, Caley et al. (2014) developed a new quantitative foraminiferal-based proxy index for Agulhas leakage, named Agulhas leakage efficiency; they suggested that Agulhas leakage efficiency represents the transfer of Agulhas water from the Indian Ocean to the Atlantic Ocean independent of the variations in the source assemblage. Using this approach, Caley et al. (2014) confirmed substantial changes in Agulhas leakage volume transports at the termination of glacial periods and suggested the possibility of a complete halt of leakage during MIS 12.

The composition of terrigenous sediment in core tops of marine sediment cores around South Africa shows significant and systematic provenance variations (Franzese et al., 2006, 2009). Provenance studies using  $^{87}\text{Sr}/^{86}\text{Sr}$ ,  $\epsilon\text{Nd}$ , and  $^{230}\text{Th}_{\text{xs}}$  flux estimates from Holocene and Last Glacial Maximum (LGM) samples from a number of cores near the tip of South Africa were used to suggest reduced advection of sediment derived from Agulhas water (and therefore possibly Agulhas leakage) during peak glacial times (Franzese et al., 2006). In a follow-up study, Franzese et al. (2009) made a more detailed geographic study of Holocene and LGM samples from a set of cores from the R/V *Charles Darwin* 154 site survey cruise as well as additional Lamont Doherty Earth Observatory repository cores from the Agulhas Plateau. The authors of these two studies infer a weaker Agulhas Current and less leakage, as well as an Agulhas trajectory very similar to modern. These results are partly at odds with models that suggest that if the wind pattern is fixed but its strength reduced, a weaker Agulhas Current results in increased leakage (De Ruijter, 1982; Dijkstra and De Ruijter, 2001; van Sebille et al., 2009). However, if the winds shift northward and weaken during glacial periods, a weaker Agulhas Current could combine with reduced leakage. The equivocal comparison between data and models calls for further work including multiproxy analyses.

A key question concerning Agulhas leakage is How important are the locations of the STF and westerly wind belts? For instance, early work from Core MD96-2080 from the western flank of the Agulhas Bank (Rau et al., 2006) showed the continuous presence of subtropical planktonic foraminifers (e.g., *G. menardii*); this observation implies persistent leakage of Indian Ocean to Atlantic Ocean surface water flow during the last 450 ky. On the other hand, transient occurrences of subpolar species (*Neogloboquadrina pachyderma* dextral) suggest intermittent incursions of cold subantarctic waters in the Agulhas corridor during glacial periods (Rau et al., 2002, 2006), perhaps the result of meridional shifts of the subtropical and subantarctic fronts. The evidence suggesting significant northward movement of the STF during some full-glacial stages (by as much as 7°; Bard and Rickaby, 2009) is consistent with this idea because Agulhas Current transports into the gateway and leakage to the South Atlantic Ocean would presumably be reduced as a consequence. However, this picture is contrasted by substantial increases of ALF (Peeters et al., 2004) in the Agulhas corridor at the termination of glacial periods, alluding to an involvement of Agulhas leakage in interhemispheric ocean and climate change, notably the resumption of NADW formation (Peeters et al., 2004). Additionally, the records of Caley et al. (2014) and Petrick et al. (2015) also demonstrate Agulhas leakage was a persistent feature of the past 500 ky, including glacial MIS 10 (Caley et al., 2014), apparently at odds with the inferred effect on Agulhas leakage caused by the significant northward migration of the STF. Again the apparent contradictions of different data sets and between data sets and theory call for further work.

Multiple-species stable isotope and alkenone data from ODP Site 1085, slightly further north on the Atlantic margin, have also been taken to suggest that Agulhas leakage variability during MIS 11 was directly related to AMOC activity and wind forcing. Site 1085 shows evidence for strengthened leakage at the end of MIS 11 that may have contributed to renewed ice-sheet build-up and the demise of warm conditions (Dickson et al., 2010).

Recent reconstructions detailing the multicentennial- to millennial-timescale hydrographic variability in the Agulhas corridor have been published for the last glacial cycle (80 ka to present; Core MD02-2594) (Dyez et al., 2014), the previous two glacial–interglacial

climate cycles of MIS 8–5 (Core MD96-2080; 265–77 ka) (Marino et al., 2013), and upstream in the Agulhas Current for the last glacial cycle (Core CD154 17-17K; Simon et al., 2013). These studies confirm a connection of leakage maxima with glacial–interglacial climate transitions and extend the connections to millennial-scale salinity anomalies that appear to systematically coincide with accelerated warming at the end of the North Atlantic cold phases (stadials). These findings lend credence to earlier hypothetical considerations (Lutjeharms, 2006a) and numerical models (Knorr and Lohmann, 2003; Weijer et al., 2002; Biastoch et al., 2008) pointing to Indian-Atlantic salt leakage as an essential modulator of abrupt climate change. However, reanalyses of a freshwater LGM hosing experiment using a fully coupled atmosphere–ocean general circulation model (Gong et al., 2013) has been recently used to suggest that the millennial-scale salinity increases in the Agulhas corridor are consistent with the presence of a quasi-interhemispheric salt seesaw. This salt seesaw is analogous to the thermal bipolar seesaw, responding to reduced cross-hemispheric heat and salt exchange during times of weakened AMOC (Simon et al., 2015a). Indeed, Simon et al. (2015a) suggest that although more upstream saline Agulhas water was propagated to the Indian-Atlantic Ocean gateway during Heinrich Stadial 1, the salt flux into the South Atlantic Ocean, through the Indian-Atlantic Ocean gateway, during that time was actually reduced. Accordingly, intervals with higher salinity in the Agulhas Current source region cannot necessarily be taken to indicate increased salt exchange through the Indian-Atlantic Ocean gateway into the South Atlantic Ocean, again calling for further study.

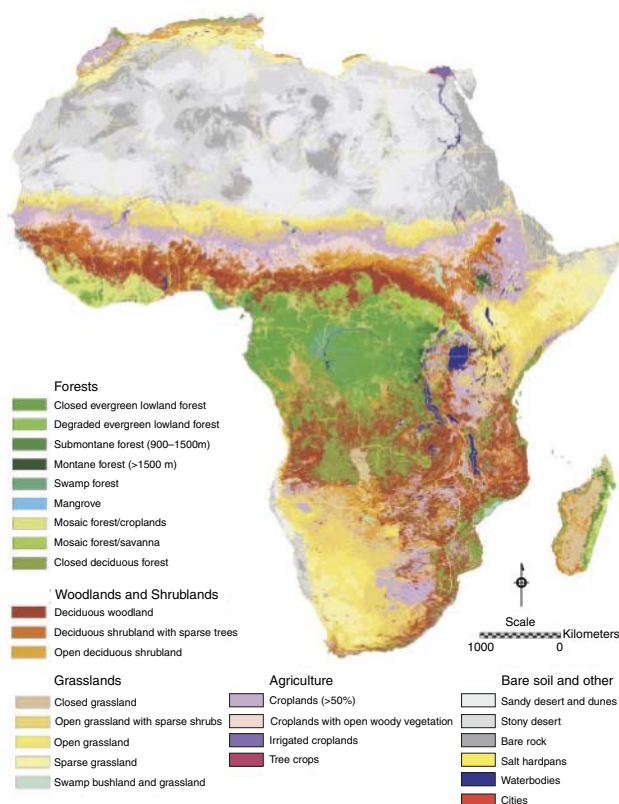
The combined analysis of these multiple proxies on the same sediments from Expedition 361 and the greater time interval available for study will help to elucidate the apparently contradictory scenarios from individual studies.

### Agulhas Current links with southern African climates and rainfall patterns

Southern Africa spans a significant range of climatic and vegetation zones (Figure F7) from the low-latitude tropical regime that is primarily under the influence of the Intertropical Convergence Zone (ITCZ) and Congo Air Boundary (CAB) to the high southern latitudes that are influenced by the northern extensions of the circum-Antarctic westerly wind belt (Tyson, 1986; Chase and Meadows, 2007) (Figure F3). The ITCZ and CAB collectively form the divide between easterly winds arriving from the Indian Ocean and westerly winds from the Atlantic Ocean (Leroux, 1983). Presently, the southernmost extent of the seasonal migrations of the ITCZ and CAB is situated over the Zambezi Catchment during austral summer (December, January, and February [DJF]). The Zambezi Catchment is the largest riverine catchment in southeastern Africa, with headwaters extending far into the continental interior. Because of the northern position of the African rainfall belt as well as the presence of mountain ranges of the East African Rift zone, annual rainfall amounts are higher across the Zambezi Catchment compared to the neighboring Limpopo Catchment to the south. Although directly connected with the ITCZ, the convergence zone over tropical southern Africa is decoupled on occasion from the divide between dry trade winds in the east and wet monsoon in the west, as occurring, for instance, over West Africa in summer. Rainfall intensity today in southeast Africa is correlated to SST variability in the southwest Indian Ocean and the Agulhas Current system (Reason and Mulenga, 1999; Walker, 1990; Jury et al., 1993; Reason, 2002; Singleton and Reason, 2007). Additionally, the interannual variability



Figure F7. Global land cover map of Africa (Mayaux et al., 2004).



ity of precipitation in the historical record of South Africa is anti-phased with that of tropical East Africa (Goddard and Graham, 1999). It is apparent from modern observations that terrestrial precipitation records have very regional patterns, and thus understanding past variability requires records from a range of archives that span this variability.

During the Late Pleistocene, shifts in African climate were paced by regional insolation changes driven by orbital precession (deMenocal et al., 1993). These orbital changes resulted in anti-phased summer insolation maxima between the Northern and Southern Hemispheres, driving an interhemispheric temperature contrast resulting in shifts of the ITCZ and leaving an imprint of the summer monsoonal rainfall (Kutzbach et al., 2008). Additionally, African climate was influenced by high-latitude forcing predominantly varying on longer obliquity and eccentricity timescales (Tuenter et al., 2003; Weber and Tuenter, 2011). Although in the low-latitude tropics African climate is dominated by the influence of the ITCZ and associated rain belts, studies from central and eastern Africa document the increasing influence of the Atlantic and Indian Oceans (Castañeda et al., 2007; Schefuß et al., 2005; Tierney et al., 2008). Whereas longer term climate evolution of the African continent is documented by a variety of terrestrial and marine records, the history of the northeastern and southeastern tropical Africa (e.g., Schneider et al., 1997; Johnson et al., 2002; Schefuß et al., 2005; Castañeda et al., 2007; Weldeab et al., 2007; Tierney et al., 2008, 2011; Ziegler et al., 2010; Otto-Bliesner et al., 2014) is better constrained than southern Africa. Southern African records are sparse, fragmented, and often suffer from a lack of robust age frameworks (e.g., Gasse et al., 2001; Partidge et al., 1997, 2004; Chase et al., 2007; Collins et al., 2014). Such information is important because variability

is likely to have significantly influenced African biomes during the Pleistocene, and it appears that this has in turn affected the distributions of species as well as human demography (Scott et al., 2008).

Lake Malawi, the southernmost of the East African Rift lakes in the southern tropics, provides the only available continuous continental climate record of the last 1.3 My in southern East Africa. Reconstruction of temperature (using the organic proxy TEX<sub>86</sub>) and aridity (using leaf wax  $\delta^{13}\text{C}$ ) has been used to suggest a variable and predominantly arid climate that has 40 ky periodicity prior to 900 ka (the mid-Pleistocene transition [MPT]), shifting to a more humid environment after the MPT (Johnson et al., 2016). After the MPT, the record has strong 100 ky periodicity, with warmer and wetter conditions in interglacial intervals and relatively cool and dry glacial intervals (Johnson et al., 2016). The inferred shift in climate conditions across the MPT is supported by additional sedimentological proxies that indicate lower lake levels prior to the MPT followed by higher amplitude lake level variability and overfilled lake conditions until ~100 ka punctuated by intervals of short-lived extreme drought (Lyons et al., 2015). A similar story for the MPT is also found in the record of African C<sub>4</sub> plant abundance from ODP Site 1077 in the tropical South Atlantic Ocean, although this record does not exhibit the shift toward overall more humid conditions for the last ~900 ky (Schefuß et al., 2003).

Millennial-scale climate variations are known from sediment cores off the Zambezi River, which drains into the Indian Ocean at ~20°S. A  $\delta\text{D}_{\text{leaf wax}}$  record in sediment Core GeoB 9307-3 off the Zambezi River documents brief and distinctly wetter intervals that correlate with the Younger Dryas and Heinrich Stadial 1 (Schefuß et al., 2011). Together with BIT index (the relative abundances of branched glycerol dialkyl glycerol tetraethers [GDGTs] and the isoprenoid GDGT, crenarchaeol [Hopmans et al., 2004; Huguet et al., 2007]) and X-ray fluorescence data, this proxy record is taken to indicate that rainfall intensities and Zambezi sediment discharge increased because of a southward shift of the ITCZ during Northern Hemisphere cold events. However, in an apparent contradiction, the  $\delta^{13}\text{C}_{\text{leaf wax}}$  record of this core shows an increased abundance of drought-tolerant C<sub>4</sub> grasses during these inferred wet intervals. This discrepancy in C<sub>3</sub>–C<sub>4</sub> plants may be attributed to shifts in source areas within the Zambezi subcatchments, as inferred from magnetic mineral inventories (Just et al., 2014), presumably due to changes in the locus of rainfall.

A more distal, deeper coring site ~200 km north of the Zambezi River mouth at ~1300 m water depth provides continuous sedimentary records spanning the last ~40 ky. A high-resolution record of neodymium isotopes ( $\epsilon\text{Nd}$ ) in the lithogenic clay fraction from one of the cores (64PE304-80) is interpreted to reflect changes in Zambezi sediment discharge relative to sediment transported from more northern sources by the strong net southward eddy-induced flow through the Mozambique Channel (van der Lubbe et al., 2016). The reconstructed intervals of enhanced Zambezi discharge are in agreement with the  $\delta\text{D}$  rainfall record obtained from a parallel core (GIK16160-30; Wang et al., 2013) after correction for the C<sub>3</sub>–C<sub>4</sub> vegetation abundances.

Collectively, the evidence is consistent with the inference that the Zambezi Catchment experienced relatively wetter conditions during MIS 2 and the late Holocene, whereby Heinrich Stadial 1 and the Younger Dryas stand out as intervals of extreme wet conditions. More arid conditions occurred during MIS 3 and between ~15 and 5 ka. The later interval is synchronous with the African Humid Period in northern Africa. Altogether, Zambezi Catchment records

document precession-paced changes in the African monsoonal rainfall belt with enhanced rainfall during periods of increased Southern Hemisphere summer insolation, superimposed by wet conditions during high-latitude Northern Hemisphere abrupt cold events.

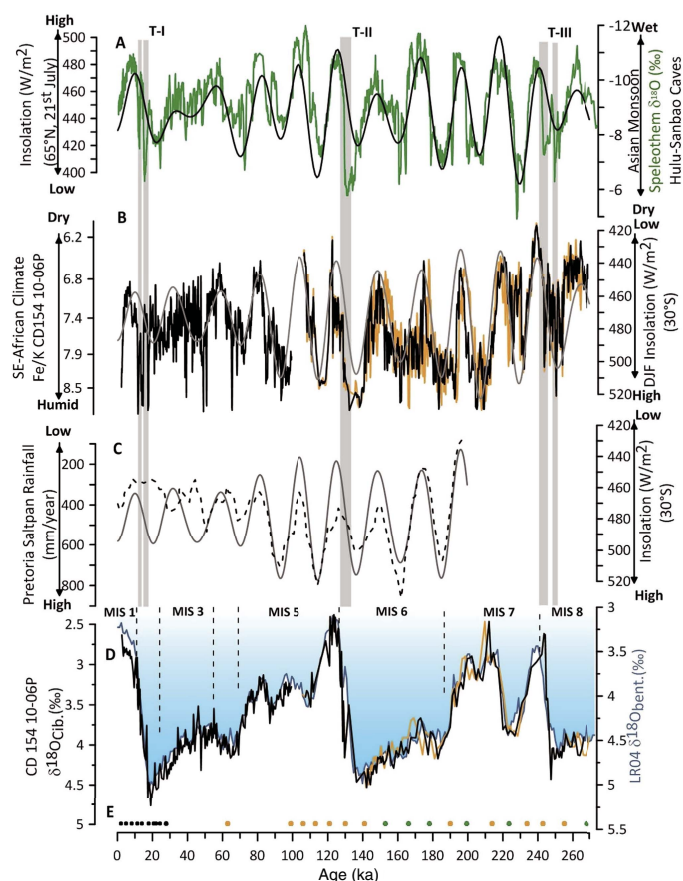
Farther south, a 342 ky pollen record from marine sediment Core MD96-2048 from 660 m water depth ~120 km south of the Limpopo River mouth shows glacial–interglacial and precession-correlated climate variations in the extent of the subtropical biome (Dupont et al., 2011). On land the sediment record from Tswaing Crater, a meteorite impact north of Pretoria, provides the longest terrestrial record of climate variability in South Africa, spanning the last 200 ky. Clastic particle size variations from Tswaing Crater have also been used to suggest that cyclic changes in precipitation were linked with the long-term effect of regional insolation changes driven by orbital precession (Partridge, et al., 1997, 2004).

Marine sediment Core CD154-10-06P located off the east coast of subtropical South Africa, KwaZulu-Natal (KZN), reveals shifts in river runoff at the southeast African margin spanning the past 270 ky that also strongly follow orbital precession pacing (Figure F8). X-ray fluorescence elemental data from Simon et al. (2015b) show humid phases that can be plausibly correlated with millennial-scale Northern Hemisphere cold events. These excursions extend those recently identified in similar data from Core CD154-17-17K, a little farther south, that were linked to periods of early modern human innovation during the middle Stone Age (Ziegler et al., 2013). Some of the millennial-scale events recorded in the eastern South Africa runoff records are coincident with periods of enhanced dune activity and mobility in the Kalahari and Namib Deserts and are thus consistent with the suggestion of shifted rainfall patterns and wind regimes (Stokes et al., 1999; Thomas et al., 2000; Stuut et al., 2002). Broadly, the conclusion drawn from the data is that past continental climate changes in southern Africa were linked with changes in ocean and atmospheric circulation that in turn varied due to precession of Earth's orbit. In some cases the changes were inferred to be associated with shifts of the ITCZ, and in other cases migration of the Antarctic Circumpolar Ocean fronts and/or the Southern Hemisphere mid-latitude westerly wind belt was likely influential.

A particularly intriguing feature of records from both stalagmites in coastal caves in South Africa (Bar-Matthews et al., 2010) and marine sediment cores at the southeast African continental margin (Scheffuß et al., 2011; Wang et al., 2013; Ziegler et al., 2013; Just et al., 2014; van der Lubbe et al., 2014, 2016; Weldeab et al., 2014; Simon et al., 2015b) is the indication of recurrent but spatially variable millennial oscillations of rainfall patterns. Such variability demands highly resolved spatial and temporal records as well as numerical simulations.

Precession-only sensitivity experiments using a high-resolution, fully coupled ocean-atmosphere model (altering the seasonal and latitudinal distribution of solar insolation) show that during precession maxima (Pmax) higher Southern Hemisphere summer (DJF) insolation causes higher temperatures and lower surface pressure over the Southern Hemisphere, especially over land (Figure F9). Specifically, lower pressure during Pmax over South Africa and higher surface pressure over the southern Indian Ocean results in stronger easterly surface winds blowing toward eastern South Africa (Figure F9). Over land, convection also increases, resulting in greater precipitation and greater net precipitation (precipitation minus evaporation) over nearly all of southern Africa during Pmax. The enhanced precipitation during Pmax and associated higher surface runoff is particularly pronounced in the KZN and Eastern Cape

Figure F8. Long-term climate variability in eastern South Africa over the past 270,000 y from Simon et al. (2015b). A.  $\delta^{18}\text{O}$  splice from Chinese speleothems (green line) (Wang et al., 2001, 2008; Cheng et al., 2009) as presented in Barker et al. (2011) showing synchronous variability of the East Asian Monsoon with Northern Hemisphere summer insolation at  $65^\circ\text{N}$ . Underlying gray bars indicate glacial–interglacial terminations (T). B. Fe/K ratios from Core CD154-10-06P (black line; five-point running mean, fine-tuned age model) indicating a more humid KwaZulu-Natal climate in accordance with varying austral summer (December–January) insolation at  $30^\circ\text{S}$ . Light orange line shows initial age model based on global benthic Stack LR04 tuning. C. Pretoria Saltpan rainfall record (Partridge et al., 1997) with December insolation at  $30^\circ\text{S}$ . D. Benthic foraminiferal (*Cibicides* spp.)  $\delta^{18}\text{O}$  record from Core CD154-10-06P (black = fine-tuned age model, light orange = initial age model) reflecting global ice volume variability and local deep-water conditions in comparison with global benthic Stack LR04 (blue) (Lisiecki and Raymo, 2005). MIS = marine isotope stage. E. Age control points for Core CD154-10-06P including radiocarbon dates (black), tuning of the foraminiferal  $\delta^{18}\text{O}$  record (light orange), and  $\delta^{18}\text{O}$  splice from Chinese speleothems on precession phase (green).



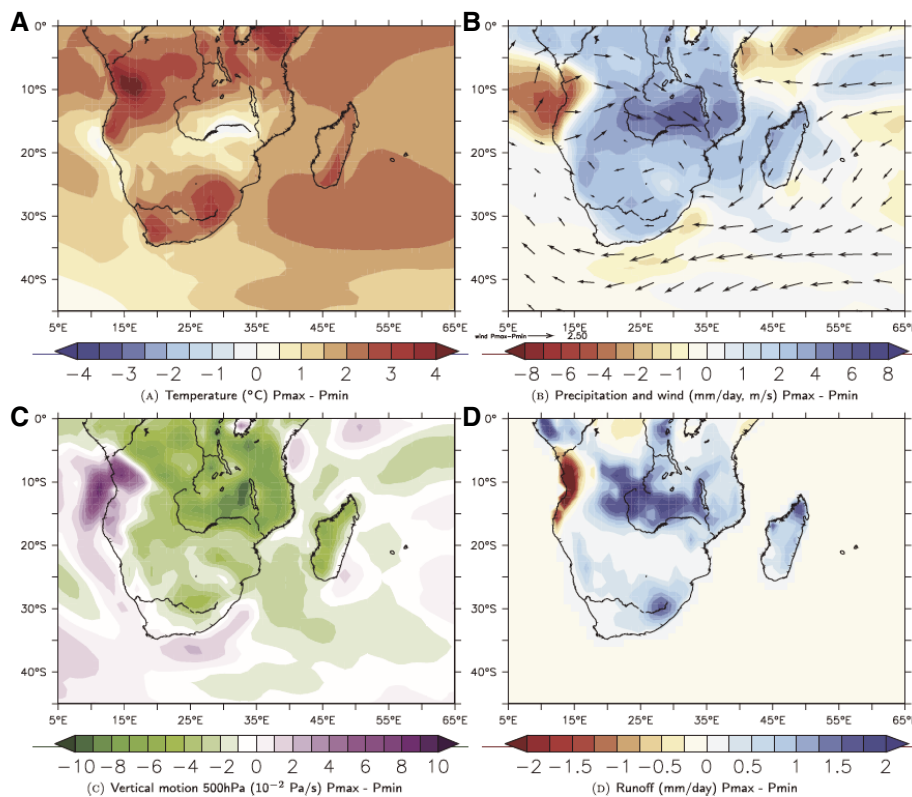
province region, compared to other regions further north or west, which may explain the considerable spatial variability observed in different core sites adjacent to land in the southwest Indian Ocean.

### Density profile of the LGM ocean

The circulation of the deep ocean at the LGM is a crucial piece of the puzzle to understand past variations in  $\text{CO}_2$  and glacial–interglacial cycles more generally. Although the distribution of sedimentary deep-ocean circulation tracers ( $\delta^{13}\text{C}$ ,  $\text{Cd}/\text{Ca}$ ,  $\epsilon\text{Nd}$ , and  $\delta^{18}\text{O}$ ) indicate that LGM ocean circulation was likely different than that of the modern ocean (e.g., Rutberg et al., 2000; Piotrowski et al., 2004, 2005; Marchal and Curry, 2008; Lund et al., 2011) the distri-



Figure F9. Results of idealized sensitivity modeling experiments showing the differences between maximum and minimum precession for December–February from experiments using high-resolution fully coupled ocean atmosphere model EC-Earth from Simon et al. (2015b). Figures were created using the model output and the program Ferret, version 6.82 for Macintosh (<http://www.ferret.noaa.gov/Ferret/>). A. Temperature. B. Precipitation. C. Vertical motion at 500 hPa (roughly 5 km). Negative values indicate upward motion; positive values indicate downward motion. D. Surface runoff over land.



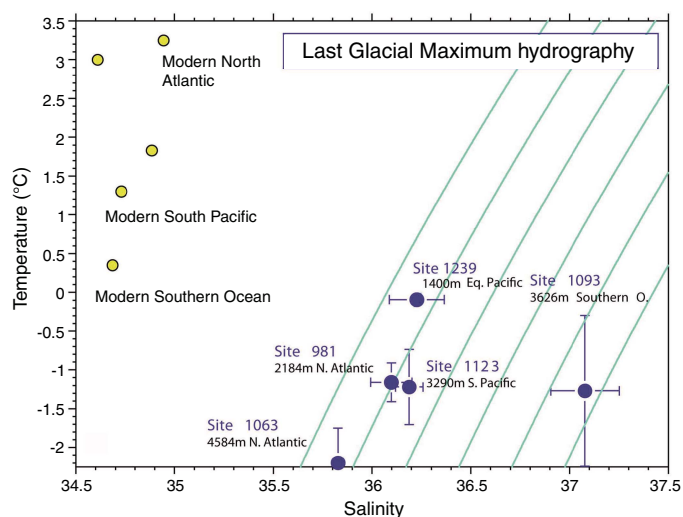
butions of these properties are inconclusive enough that they may also be consistent with modern circulation and differences in surface forcing (e.g., Gebbie, 2012; Gebbie and Huybers, 2006; LeGrand and Wunsch, 1995). Thus, to understand the role of the deep ocean in ice age climate variability, we require a more robust assessment of ice age deep-ocean variability, using a suite of tracers characterized by wide-ranging sensitivity to surface forcing.

In the modern ocean, we use the pattern of temperature and salinity to map the distribution of deep-water masses. In addition, these values inform the balance of sea-atmosphere exchange (i.e., precipitation, evaporation, heat exchange) processes that are required to create the negative buoyancy to form deep water; they are an integrated record of climate at high latitudes. It would be ideal to use the same strategy for the ice age ocean. However, there are few available means to develop a full temperature-salinity plot for the LGM. Schrag and DePaolo (1993) and Schrag et al. (1996, 2002) pioneered a strategy for quantifying the  $\delta^{18}\text{O}$  of seawater by sampling a profile of interstitial waters from deep-ocean sediment from IODP sites, making highly precise measurements of  $\delta^{18}\text{O}$  and applying diffusion models to estimate the value of LGM seawater. This strategy has been expanded to include  $[\text{Cl}^-]$ , thus allowing an estimate of salinity (Adkins et al., 2002; Adkins, 2013; Miller et al., 2015). With local  $\delta^{18}\text{O}$  data from benthic foraminifers, interstitial water information from a coring site can be turned into a single point on the LGM temperature-salinity diagram. The current version of this plot, from sampling IODP material, contains only five pairs of points (Figure F10) (Adkins et al., 2002). The relative scarcity of the observations stems from the requirements of the approach: sedi-

ment must be drilled to >150 m below seafloor to properly measure and model the diffusion profile; a well-behaved sediment column, free of complicated fluid flow and diagenesis, must be sampled; and an entire hole must be dedicated to the pursuit. Nevertheless, despite the stringent requirements this approach is currently the most direct way to estimate temperature and salinity of the ice age deep ocean.

Although sparse, the existing data imply that there were larger-than-modern density differences between deep-ocean water masses at the LGM, and most of the ocean was close to the freezing point and therefore differences in density were due primarily to salinity (see also discussion in Miller et al., 2012). Most surprisingly, the salt gradient in the modern Atlantic Ocean, consisting of fresh southern-sourced waters and salty northern-sourced waters appears to have been reversed at the LGM. This suggests the saltiest waters of the LGM ocean originated in the Southern Hemisphere, although this inference is based on a single site (IODP Site 1093). Adkins et al. (2002) concluded that Glacial Southern Source Bottom Water was the densest water mass due to its high salinity. Enhanced density differences between northern- and southern-sourced bottom waters imply greater vertical stratification. A stronger vertical stratification would help explain indications that there was a physically isolated reservoir of  $\text{CO}_2$  in the deep ocean at the LGM (Broecker and Barker, 2007). In the years since the original work, analytical measurement of  $[\text{Cl}^-]$  has been improved, several new sampling strategies have targeted increasing throughput, and the data analysis tools used to model the data have been improved. The next most important step in constraining the temperature, salinity, and density

Figure F10. Temperature and salinity plot for modern (yellow) and LGM (blue) values. Based on the type of interstitial water data, this plot shows several differences with modern values. From Adkins and Schrag (2002).



structure of the LGM deep ocean is to obtain depth transects of these properties and just generally more data from different parts of the ocean. To this end, samples taken during Expedition 361 sites ranging from 450 to 3040 m water depth will be used to help constrain the processes that could have filled the LGM deep ocean and controlled its circulation.

## Objectives

Expedition 361 Pliocene–Pleistocene objectives include the following:

- To establish the sensitivity of the Agulhas Current to changing climates of the Pliocene–Pleistocene in association with transient to long-term changes of high-latitude climates, ocean gateways, tropical heat budgets, and the monsoon system;
- To determine the dynamics of the Indian–Atlantic Ocean gateway circulation during such climate changes in association with changing wind fields and migrating ocean fronts;
- To examine the connection between Agulhas leakage and ensuing buoyancy transfer and shifts of the AMOC during major ocean and climate reorganizations during at least the last 5 My; and
- To address the impact of Agulhas variability on southern Africa terrestrial climates and notably rainfall patterns and river runoff.

Specific scientific objectives include the following:

- To determine the variability of the greater Agulhas Current system under contrasting climatic regimes: Agulhas Current variability under early Pliocene warm conditions, during the mid-Pliocene expansion of Northern Hemisphere ice sheets, and through the MPT;
- To establish the role of Agulhas leakage in transient and long-term shifts of the AMOC during these climatic transitions: proxy records of mode/strength of the Agulhas leakage during periods of major climatic reorganizations and for AMOC variations through the same intervals;
- To establish the response of the Agulhas Current system to Antarctic climate variability and bipolar linkages: evidence for vari-

ation of the Agulhas leakage in highly resolved temporal records will be compared to Antarctic and bipolar climate changes, as well as with evidence for shifts of Southern Ocean fronts bordering the Agulhas corridor at the southern tip of Africa, and position and strength of the westerlies;

- To understand the remote tropical origins of Agulhas Current variability: evidence for variation of the Agulhas leakage in highly resolved temporal records will be compared to records of Indonesian Throughflow and monsoon modulation on the Agulhas Current source region and impacts on warm and salt-water transports along the southeast African margin and Indian–Atlantic saltwater leakage;
- To better determine southern African climate sensitivity: comparison of varying Agulhas Current warm-water transports with rainfall patterns in southern Africa, in connection with the Pliocene–Pleistocene evolution of tropical climates, ITCZ migration, and tectonic history of eastern Africa; and
- To provide further verification of glacial–interglacial variations in the Southern Ocean’s role in the carbon cycle: nutrient cycling, productivity, and  $p\text{CO}_2$  and their role in regional and global biogeochemical cycles.

Additionally, Expedition 361 will complete an intensive interstitial fluids program aimed at constraining the temperature, salinity, and density structure of the LGM deep ocean, from the bottom of the ocean to the base of the main thermocline. High-resolution interstitial water samples will be utilized for  $\delta^{18}\text{O}$  and chloride measurements to characterize the salinity of the bottom water and to deconvolve the temperature and ice volume control on  $\delta^{18}\text{O}$  of benthic foraminifers.

## Principal results

### Site U1474

Site U1474 is located in the northernmost Natal Valley ( $31^{\circ}13.00'\text{S}$ ;  $31^{\circ}32.71'\text{E}$ ) ~88 nmi south-southeast of Durban, South Africa, in a water depth of 3034 m below sea level (mbsl). The Natal Valley is a sediment-filled basin between the coast of southeast Africa and Mozambique Ridge. At its northern end, the shoaling valley passes into the extensive coastal plains of southern Mozambique (Dingle et al., 1978; Goodlad, 1986), whereas toward the south the deepening Natal Valley merges with Transkei Basin below 4500 mbsl. Terrigenous sediment supply into the basin is dominated by seasonal discharge from the numerous short but fast-flowing rivers that characterize the drainage system of the Natal coast (e.g., the Tugela River), as well as contributions from the Limpopo River and possibly the Zambezi River. Surface sediment dispersal is strongly affected by the ~100 km wide Agulhas Current flowing southward close to the continental shelf edge and transporting ~70 Sv at  $32^{\circ}\text{S}$  (Beal and Bryden, 1999) with surface velocities up to 1.5 m/s. At depth, the northward-flowing Agulhas Undercurrent hugs the continental slope, presently transporting NADW. The currents related to this northward flow have led to the formation of a network of elongate, north-south–orientated sediment drifts.

According to Lutjeharms (2006a), the Agulhas Current is fully constituted at ~ $26^{\circ}\text{S}$ , around Ponta do Ouro, close to the border between Mozambique and South Africa. Here, the East Madagascar Current converges with Mozambique Channel throughflow to form the Agulhas Current. The northern section of the Agulhas Current system, overlying Site U1474, is extraordinarily stable due to the steep, nearly rectilinear continental shelf that topographically steers



the current flow southward (de Ruijter et al., 1999a; Lutjeharms, 2006a, 2007).

Several high-resolution orbital- to millennial-timescale records from shallow-penetration marine sediment cores located near Site U1474 (Ziegler et al., 2013; Simon et al., 2013, 2015a, 2015b) have provided multiproxy reconstruction of hydrographic variability within the main flow of the Agulhas Current as well as river discharge from the catchments of the Natal coast over the past 250 ky. In particular, Simon et al. (2013) demonstrated a strong correspondence between upstream Agulhas Current hydrographic variability and that of the Agulhas leakage corridor (deduced from foraminiferal proxies for temperature, salinity, and Agulhas leakage). This system can be plausibly linked to the influence of recirculation within the southwest Indian Ocean subgyre, as predicted from physical oceanography (e.g., Gordon, 1985; Stramma and Lutjeharms, 1997). Additionally, it has been shown that changes in terrestrial climate (hydroclimate) are related to regional precession-paced insolation changes and the effects associated with high-latitude abrupt climate forcing (Simon et al., 2015b). When considered in combination with archaeological records, these records appear to suggest links between climate and human settlement in South Africa during the middle Stone Age (Ziegler et al., 2013).

Our primary objective was to recover a key Pliocene–Pleistocene sedimentary succession, including the early Pliocene warm period, mid-Pliocene expansion of Northern Hemisphere ice sheets, and the MPT, formed under the influence of the upper reaches of the Agulhas Current. Site U1474 is ideally located to provide

- A reconstruction of Agulhas Current warm-water transports close to the start of the fully constituted Agulhas Current;
- An upstream record that may allow discerning connections between Agulhas leakage and its headwater variability;
- Evidence to understand the connections between southern African terrestrial climates and southeast Indian Ocean heat budgets, notably Agulhas Current warm-water transports and associated ocean-atmosphere heat and moisture transfer;
- A record of the vigor and hydrography of NADW (or its precursors) transported to the Circumpolar Deep Water (CDW) and the southwest Indian Ocean; and
- An interstitial water profile of  $\delta^{18}\text{O}$  and chloride ion concentration that can enhance the understanding of temperature, salinity, and density of the deep water bathing Site U1474 by providing an additional data point to the presently available pore water LGM temperature and salinity inventory.

Site U1474 consists of eight holes ranging in penetration depth from 3.1 to 254.1 m drilling depth below seafloor (DSF). Overall, 111 cores were recovered at this site. The advanced piston coring (APC) system penetrated 794.0 m and recovered 809.79 m of core (102% recovery). The half-length APC (HLAPC) system penetrated 102.4 m and recovered 100.96 m of core (99% recovery). A total of 910.75 m of sediment was recovered from the 896.4 m interval cored (102% recovery) at Site U1474. The total time spent at Site U1474 was 7.6 days (Table T1).

### Sedimentology

The sediment at Site U1474 consists of one lithologic unit. Unit I (0–254.07 m CSF-A) a brown (uppermost ~0.5 m CSF-A) and greenish gray (deeper than 0.5 m CSF-A) foraminifer-bearing clay with nannofossils alternating with intervals of greenish gray nannofossil-rich clay with or without foraminifers. Nannofossil-rich clay is more common downhole. Dark gray foraminifer-bearing fine-

sand layers are common (typically 1–3 per core) and range in thickness from a few centimeters to ~2 m (most are a few centimeters thick). The composition of these sandy layers is predominantly quartz with carbonate, and many have erosional bases. Sediment includes dark gray mottling that we interpret as bioturbation. The relative intensity of the bioturbation typically increases with depth throughout. Sediment at Site U1474 also includes millimeter- to centimeter-scale green layers that contain pyrite and glauconite based on X-ray diffraction (XRD) and smear slide observations. These layers represent diagenetic alterations.

Significant drilling disturbance was observed in the first section of most of the cores and in the base of partial-stroke APC cores.

### Physical properties

Despite the relatively homogeneous lithology, physical parameters vary along the entire sequence. Porosity increases with depth from 0 to 12.8 m core composite depth below seafloor (CCSF-A), reaching a maximum of 72% followed by a decrease to <50% at the bottom of the cored section. This decreasing trend toward the bottom is likely due to compaction. Variations in gamma ray attenuation (GRA) density reflect the same compaction pattern. Natural gamma radiation (NGR) and magnetic susceptibility markedly decrease in the upper 10 m CCSF-A and increase to 130 m CCSF-A. Below 130 m CCSF-A, NGR and magnetic susceptibility values decrease to the base of the section. This decreasing trend occurs in spite of compaction and thus likely reflects decreasing abundances of clay and magnetic minerals within the sediment. Superimposed on the general trends are cyclic variations in NGR of typically ~10–15 counts/s, which may be related to the variable mixtures of terrigenous and biogenic sediment components. Fine-sand layers found throughout the cored section are clearly identified by peaks in magnetic susceptibility and GRA density values. Spectral reflectance shows a marked cyclic pattern downhole and broadly corresponds with the trends in color spectrometry values.

### Micropaleontology

Analysis of calcareous nannofossils, planktonic and benthic foraminifers, and diatoms from core catcher and split-core samples from Hole U1474A indicates that Site U1474 spans the Late Pleistocene to the latest Miocene (~6.2 Ma). Calcareous microfossils show good to moderate preservation in the upper 180 m CSF-A of the site. Below this depth, preservation and fossil group abundances decrease. A complete sequence of calcareous nannofossil zones was found from the late Miocene (Biozones NN11–NN21) to Late Pleistocene. Evidence for reworking of calcareous nannofossils is found throughout the entire sequence, marked by the persistent presence of early Miocene species. The calcareous nannofossil assemblages are typical of tropical to subtropical settings, although the temperate species, *Coccolithus pelagicus*, is found throughout the cored interval. Planktonic foraminifers also record a complete sequence of tropical biozones that span Miocene Biozone M13 to Late Pleistocene Subzone Pt1b. The planktonic foraminifer fauna is dominated by *Globorotalia inflata*, which is indicative of transitional environments between subtropical and polar water masses, and in deeper parts of the sequence the assemblage contains elements of the globocollid group that are typical of subtropical convergence zones. Benthic foraminifer fauna is diverse and includes numerous suboxic components characteristic of a lower bathyal to abyssal environment. Siliceous microfossils are characteristic of a modern Indian Ocean assemblage and make up a minor component of the mudline sample. Below the mudline sample, sponge spicules are rare to trace

Table T1. Expedition 361 hole summary. [Download table in .csv format.](#)

Hole	Latitude	Longitude	Seafloor depth (mbrf)	Cores (N)	Cored (m)	Recovered (m)	Recovery (%)	Drilled (m)	Total penetration (m)	Total depth (mbrf)	Time on hole (h)	Time on site (days)
U1474A	31°12.9995'S	31°32.7080'E	3044.6	29	254.1	264.83	104	0.0	254.1	3298.7	57.25	2.4
U1474B	31°12.9882'S	31°32.7083'E	3045.2	16	147.3	142.09	96	0.0	147.3	3192.5	21.75	0.9
U1474C	31°12.9999'S	31°32.7215'E	3050.0	1	3.1	3.07	99	0.0	3.1	3053.1	1.75	0.1
U1474D	31°12.9999'S	31°32.7213'E	3045.2	14	124.5	126.04	101	0.0	124.5	3169.7	16.75	0.7
U1474E	31°12.9995'S	31°32.7208'E	3045.2	17	112.0	114.44	102	107.0	219.0	3264.2	34.00	1.4
U1474F	31°13.0105'S	31°32.7078'E	3046.8	32	236.4	240.49	102	2.0	238.4	3285.2	51.25	2.1
U1474G	31°12.9995'S	31°32.7080'E	3044.6	1	9.5	9.97	105	0.0	9.5	3054.1	1.50	0.1
U1474H	31°12.9995'S	31°32.7080'E	3044.6	1	9.5	9.82	103	0.0	9.5	3054.1	1.50	0.1
Site U1474 totals:				111	896.4	910.75	102	109.0	1005.4		185.83	7.6
U1475A	41°25.6052'S	25°15.6440'E	2681.4	1	1.5	1.48	99	0.0	1.5	2682.9	11.25	0.5
U1475B	41°25.6052'S	25°15.6441'E	2680.5	26	243.9	250.31	103	0.0	243.9	2924.4	25.50	1.1
U1475C	41°25.5941'S	25°15.6439'E	2680.4	29	272.0	280.70	103	3.0	275.0	2955.4	29.00	1.2
U1475D	41°25.6055'S	25°15.6586'E	2679.3	16	143.0	150.86	105	0.0	143.0	2822.4	19.00	0.8
U1475E	41°25.6162'S	25°15.6439'E	2682.9	29	270.5	277.20	102	6.5	277.0	2959.9	33.00	1.4
U1475F	41°25.6054'S	25°15.6299'E	2680.4	6	57.0	55.37	97	76.5	133.5	2813.9	20.00	0.8
Site U1475 totals:				107	987.9	1015.92	103	86.0	1073.9		137.75	5.7
U1476A	15°49.2535'S	41°46.1242'E	2176.8	24	224.2	232.90	104	0.0	224.2	2401.0	29.50	1.2
U1476B	15°49.2421'S	41°46.1234'E	2177.6	16	148.4	154.42	104	0.0	148.4	2326.0	14.25	0.6
U1476C	15°49.2526'S	41°46.1355'E	2176.6	1	5.7	5.72	100	0.0	5.7	2182.3	1.50	0.1
U1476D	15°49.2525'S	41°46.1358'E	2176.6	24	228.0	237.02	104	1.0	229.0	2405.6	21.50	0.9
U1476E	15°49.2635'S	41°46.1236'E	2176.3	25	234.8	243.77	104	0.0	234.8	2411.1	31.75	1.3
Site U1476 totals:				90	841.1	873.83	104	1.0	842.1		98.50	4.1
U1477A	19°21.2930'S	36°54.8962'E	440.4	27	181.2	182.53	101	0.0	181.2	621.6	20.00	0.8
U1477B	19°21.2822'S	36°54.8958'E	440.8	26	174.4	185.73	106	0.0	174.4	615.2	13.00	0.5
U1477C	19°21.2928'S	36°54.9066'E	440.4	16	118.4	121.75	103	1.0	119.4	559.8	11.50	0.5
Site U1477 totals:				69	474.0	490.01	103	1.0	475.0		44.50	1.8
U1478A	25°49.2600'S	34°46.1593'E	499.2	32	248.4	256.96	103	0.0	248.4	747.6	19.25	0.8
U1478B	25°49.2500'S	34°46.1589'E	499.2	33	246.3	246.85	100	1.0	247.3	746.5	16.25	0.7
U1478C	25°49.2604'S	34°46.1705'E	501.2	33	238.3	242.32	102	10.0	248.3	749.5	16.00	0.7
U1478D	25°49.2706'S	34°46.1592'E	502.0	21	172.7	175.97	102	43.3	216.0	718.0	14.25	0.6
Site U1478 totals:				119	905.7	922.10	102	54.3	960.0		65.75	2.7
U1479A	35°03.5290'S	17°24.0620'E	2629.9	1	1.0	1.01	101	0.0	1.0	2630.9	8.00	0.3
U1479B	35°03.5289'S	17°24.0621'E	2629.7	32	300.7	296.73	99	0.0	300.7	2930.4	31.00	1.3
U1479C	35°03.5183'S	17°24.0621'E	2629.7	31	294.5	303.69	103	1.0	295.5	2925.2	31.50	1.3
U1479D	35°03.5286'S	17°24.0754'E	2626.5	11	101.4	102.15	101	2.0	103.4	2729.9	12.50	0.5
U1479E	35°03.5402'S	17°24.0623'E	2626.5	15	142.5	148.17	104	134.0	276.5	2903.0	21.50	0.9
U1479F	35°03.5290'S	17°24.0493'E	2626.7	1	9.5	9.50	100	0.0	9.5	2636.2	2.25	0.1
U1479G	35°03.5290'S	17°24.0493'E	2623.9	1	9.3	9.29	100	0.0	9.3	2633.2	1.00	0.0
U1479H	35°03.5289'S	17°24.0489'E	2624.6	9	81.6	70.75	87	4.0	85.6	2710.2	8.75	0.4
U1479I	35°03.5286'S	17°24.0493'E	2624.8	4	34.8	21.79	63	34.7	69.5	2694.3	27.00	1.1
Site U1479 totals:				105	975.3	963.08	99	175.7	1151.0		143.50	6.0
Expedition 361 totals:				601	5080.4	5175.69	102	427.0	5507.4		675.75	28.0

until they disappear completely below Sample 361-U1474A-10H-CC (90.19 m CSF-A) and other siliceous fossils are not found.

The age model for Site U1474 is based on the biozonation of calcareous nannofossil and planktonic foraminifers. These data reveal that the recovered sedimentary sequence is free of hiatuses and that sedimentation rates are nearly constant throughout the recovered sequence (estimated to be ~3.7 cm/ky). Near the bottom there appears to be a higher accumulation rate, but more shore-based work will be required to refine the age model.

### Paleomagnetism

Paleomagnetic and rock magnetic analyses were carried out on sediment cores from Holes U1474A–U1474H. Analyses of discrete samples taken from Hole U1474A indicate that the magnetic mineralogy is dominated by magnetite. The relatively high saturation isothermal remanent magnetization (SIRM) and high and stable S-ratios further reveal that magnetic minerals are well preserved throughout the sequence. Below ~100 m CSF-A, an increase of

magnetic minerals is expressed by high SIRM. Inclination and declination records were obtained by measuring and demagnetizing the natural remanent magnetization (NRM) of archive core halves and discrete samples. The majority of the APC cores were oriented using the Icefield MI-5 tool. The inclination record is of good quality, and nine polarity zones can be identified. The boundaries of the Brunhes, Matuyama, Jaramillo, Olduvai, Gauss, Gilbert, C3n, and C3r (sub)paleomagnetic chrons constrain the chronology of the sediment and are in good agreement with the biostratigraphic data.

### Stratigraphic correlation

Continuous and distinctive variations in magnetic susceptibility were used to make hole-to-hole correlations among Holes U1474A–U1474F. Confidence in these correlations is high because they can be cross-checked with continuous downcore measurements of color reflectance parameter  $L^*$  and NGR. A complete affine table to the base of Hole U1474A was constructed to provide tabulated offsets for all cores relative to core depths (CSF-A). A con-

tinuous splice was constructed downhole to 234 m CCSF-D using cores from Holes U1474D–U1474F where possible. For some intervals, it was necessary to use sections from Hole U1474A, but sequences in which whole-round samples had been removed were avoided. All sections of disturbed sediment were also avoided for construction of the splice. The continuity of the splice was confirmed with magnetic susceptibility, GRA density, NGR, and reflectance. Core intervals not included in the splice may be correlated to the CCSF-D depth scale using within-core variations of these data.

### Geochemistry

Interstitial water chemistry shows a moderate degree of early sediment diagenesis at Site U1474. Profiles of Mn and Fe indicate suboxic conditions from the seafloor to 5 m CSF-A. Sulfate reduction occurs throughout much of the sediment column. The overall low organic carbon content, on average <0.5 wt%, does not support complete removal of sulfate by 250 m CSF-A, and methane concentrations remain at background levels. Carbonate content of the pelagic sediment is moderate, averaging 40 wt%. The sandy layers contain significantly less carbonate than the nannofossil-rich clays, with values ranging between 20 and 30 wt%. Profiles of conservative elements (K and Mg) downhole suggest either fluid flow or clay mineral alteration reactions at depths below 250 m CSF-A.

### Site U1475

Site U1475 is located on the southwestern flank of Agulhas Plateau (41°25.61'S; 25°15.64'E) ~450 nmi south of Port Elizabeth, South Africa, in a water depth of 2669 mbsl. The Agulhas Plateau, which was formed during the early stages of the opening of the South Atlantic at ~90 Ma (Parsiegla et al., 2008), is a major bathymetric high that is variably coated with sediment (Uenzelmann-Neben, 2001). The 230,000 km<sup>2</sup> plateau, which ascends to ~2500 m above the adjacent seafloor, is bounded on the north by the 4700 m deep Agulhas Passage and is flanked by Agulhas Basin in the west and Transkei Basin in the northeast. The northern part of the plateau is characterized by rugged topography, whereas the central and southern parts of the plateau exhibit relatively smooth topography (Allen and Tucholke, 1981) and have greater sediment thickness (Uenzelmann-Neben, 2001).

A strong flow of deep water passes over the Agulhas Plateau region (Macdonald, 1993), which involves the water column from the surface downward to the upper CDW. The hydrography of the upper ocean is dominated by the ARC, which comprises the component of the Agulhas Current that is not leaked to the South Atlantic Ocean but rather flows eastward from the Agulhas Retroflexion (Lutjeharms and Anson, 2001). Antarctic Intermediate Water, below the ARC, also follows the same flow path near South Africa as the Agulhas Current and shows a similar retroflexion (Lutjeharms, 1996). During its return to the southwest Indian Ocean, the ARC crosses the Agulhas Plateau, which deflects the deep-reaching current toward the Equator (Harris, 1970).

Site survey data collected during seismic cruises with the R/V *Petr Kottsov* in 1998 (Uenzelmann-Neben, 1998) and the R/V *Marion Dufresne* in 2002 show that Site U1475 is located on a sediment drift (Uenzelmann-Neben, 2001, 2002) that is deposited on the southwestern flank of Agulhas Plateau by NADW exiting the South Atlantic Ocean to the Indian Ocean. The wedge-shaped sediment drift thickens to the west, reaching a water depth of ~2510 m at its crest. The wavy character of the seafloor topography and the sub-surface seismic reflection pattern are interpreted as sediment waves that document sedimentation by active deep-water flow. Site U1475

is located at 2669 mbsl, near the drift crest where minor buried depressions create a locally expanded sequence.

Millennial-scale multiproxy records covering the past 350 ky previously collected at this location highlight the synchrony of the regional ocean climatology and Antarctic climate variability. Evidence for regional ocean paleoclimatology includes latitudinal migrations of the Southern Ocean fronts and deposition of ice-rafted debris that document expansion of sub-Antarctic conditions to the plateau. These data serve as evidence for possible meridional shifts in the subantarctic frontal zone that borders the Agulhas leakage corridor and likely influences Agulhas leakage.

Paleoceanographic studies spanning the past 350 ky combined with seismic reflection data support the suitability of Site U1475 to achieve our primary objectives, which are to

- Recover a complete Pliocene–Pleistocene sedimentary succession from a high-accumulation sediment drift located within the southern ARC sector of the Indian–Atlantic Ocean gateway;
- Assess the linking between Antarctic climate variations, circum-polar ocean front instability, and connections with Agulhas leakage into the South Atlantic Ocean;
- Assess the vigor and hydrography of NADW (or its precursors) exported to CDW and the southwest Indian Ocean at a location proximal to the entrance of NADW to the Southern Ocean and South Indian Ocean; and
- Evaluate the possibility of advective salinity feedbacks between Agulhas Leakage and AMOC variability, particularly the possible role of the Agulhas Leakage in modulating surface-to-deep ocean coupling in the North Atlantic during the transition between climatic states.

Site U1475 consists of six holes that range in depth from 1.5 to 277.0 m DSF. The APC system cored a total of 987.9 m and recovered 1015.92 m of core (103% core recovery). Six intervals advanced without coring to provide a continuous stratigraphic sequence penetrated a total of 86.0 m. The total time spent at Site U1475 was 5.7 days (Table T1).

### Sedimentology

The sediment recovered at Site U1475 consists of two lithologic units. Unit I (0–4.75 m CSF-A) is composed of pale brown, light greenish or olive-gray, and white-gray nannofossil-rich foraminifer ooze. Unit II (4.75–277.22 m CSF-A) is composed of light greenish or pale gray to white-gray nannofossil ooze. Alternations between foraminifer-bearing or foraminifer-rich nannofossil ooze and nannofossil ooze with fine sand (foraminifers, quartz, and occasionally diatoms) were observed.

Sediment in both units includes dark gray mottling that we interpret as bioturbation (ichnofossil burrows [*Planolites*]). Thin darker bands commonly surround burrows, and macroscopic pyritized burrows are also common. Dropstones displaying coarse sand to granule grain sizes were observed in three different cores, which we interpret to represent ice-rafted debris. Sediment in both units also includes green layers that predominantly consist of pyrite and glauconite representing diagenetic alteration.

### Physical properties

Despite the homogeneous lithology, physical parameters show downhole trends as well as cyclic variations. A marked change in physical sediment properties occurs at ~110 m CSF-A. Above this depth, a change from quasi-regular short and lower amplitude variations to longer, higher amplitude, and more irregular cyclic varia-



tion is observed. This change is most clearly displayed in NGR and color reflectance ( $a^*$ ) data. In addition,  $P$ -wave velocities decrease and bulk densities increase at this depth, following a compaction trend. Prominent changes in  $P$ -wave velocity at Site U1475 can be clearly linked to seismic reflectors present in the site survey profiles.

### Micropaleontology

Analysis of calcareous nannofossils, planktonic foraminifers, and diatoms from core catchers and working-half core samples from Holes U1475B, U1475C, and U1475E reveals that the ~277 m section recovered at Site U1475 spans the latest Miocene (~6.91 Ma) to recent. Calcareous nannofossils show moderate to good preservation in all sequences with the consistent presence of well-preserved Eocene to middle Miocene (reworked) species. The nannofossil chronostratigraphy at Site U1475 spans from Biozones NN21 to NN11 and CN15 to the Subzone CN9d/CN9c boundary. Calcareous nannofossil assemblages are characterized by species typical of tropical to subtropical and temperate environments. Planktonic foraminifer biota at Site U1475 are moderately to well preserved and are mostly composed of subtropical convergence species mixed with polar species. The Pleistocene foraminifer assemblage is dominated by *G. inflata*, which typically constitutes 30%–50% of each sample. Diatoms show poor to moderate preservation and have a mixed assemblage of subtropical and Southern Ocean taxa. Diatom markers from both environmental settings are used to define datums; however, the Southern Ocean assemblage abundance is greatly reduced below 101.62 m CSF-A. Overall, the diatom chronostratigraphy spans from Biozone NTD17 to NTD12.

Sedimentation rates at Site U1475, based on biostratigraphy and paleomagnetic reversals during the Pleistocene and latest Pliocene are nearly linear at 2.8 cm/ky between 0 and 3.9 Ma. A significantly higher sedimentation rate of 9.6 cm/ky occurs between 3.9 and 5.3 Ma. Finally, older than 5.3 Ma, sedimentation rates is again lower and is ~2.5 cm/ky.

### Paleomagnetism

All archive halves from Holes U1475B, U1475C, and U1475F and Cores 361-U1475E-1H through 15H were analyzed for NRM and demagnetized at 15 and 25 mT. In total, 98 discrete cube samples were taken from Holes U1475A, U1475B, and U1475D and subjected to NRM demagnetization. Cores from Site U1475 were not oriented because the relatively high latitude of the site implies that inclination data provide a reliable record of magnetic reversals. For most of the recovered interval, magnetic susceptibility measurements indicate that the sediment is diamagnetic in character and many readings are below the detection limit of the whole-round and point magnetic susceptibility instruments. Nonetheless, taken together, the inclination records of all individual holes provide a consistent and detailed magnetostratigraphy for Site U1475. The boundaries of the Brunhes, Matuyama, Jaramillo, Olduvai, Gauss, Kaena, and Mammoth paleomagnetic chrons and subchrons constrain the chronology of the sediment and are in general agreement with the biostratigraphic data.

### Stratigraphic correlation

Select sequences from Holes U1475B, U1475C, U1475E, and U1475F were spliced together to create the most complete and representative section possible using RGB blue,  $b^*$  color reflectance, and NGR data. Sedimentological logs were consulted meticulously in the process of constructing the splice to avoid inclusion of any

obviously disturbed sections or sections with completely unique features (relative to the other adjacent holes). The splice constitutes a continuous sequence with a total length of ~292 m CCSF-D. However, confidence in the splice is low surrounding occasional problematic intervals, and further work on shore will be required to improve confidence.

### Geochemistry

Interstitial water chemistry shows a moderate degree of early sediment diagenesis at Site U1475. The nitrate profile indicates suboxic conditions in the upper 0.25 m CSF-A. Sulfate concentrations decrease throughout the sediment column, and sulfate is never completely consumed. Methane concentrations remain at background levels. Carbonate is the dominant sedimentary component, ranging from 75 to 85 wt%, with terrigenous material dominating the remaining fraction. Organic carbon contents range from 0.3 to 0.6 wt%, decreasing slightly downhole. Together, carbonate, organic carbon, and elemental biogenic indicator profiles suggest that export production increased into the Late Pleistocene.

## Site U1476

Site U1476 is located on Davie Ridge at the northern entrance of the Mozambique Channel (15°49.25'S; 41°46.12'E) at a water depth of 2165 mbsl. It is 3 nmi northwest of DSDP Site 242.

The Mozambique Channel is considered to be one of the most turbulent areas in the world ocean (Ternon et al., 2014), bordered by Madagascar to the east and Mozambique to the west. At its southern end, flow through the channel feeds into the greater Agulhas Current system. The Mozambique Channel is characterized by complex and variable surface and subsurface circulation. Several water masses converge within the Mozambique Channel, and the circulation is dominated by activity related to the inflow of both Indonesian Throughflow and Tasman Strait Throughflow as well as the strength of the tropical and subtropical surface gyres in the Indian Ocean (Palastanga et al., 2006; Penven et al., 2006; Schott et al., 2009; Backeberg and Reason, 2010; Ridderinkhof et al., 2010). Satellite observations (e.g., Schouten et al., 2003; Quartly and Srokosz, 2004) and ocean models (Biaostoch and Krauss, 1999) have shown that mesoscale anticyclonic eddies generally form as the channel narrows close to Davie Ridge (at about 16°S), between the northern and central basin of the Mozambique Channel. These eddies typically maintain a high rotational velocity, often >1.5 m/s (Schouten et al., 2003; Ullgren et al., 2012) as they migrate southward through the channel at a mean rate of ~4–7 eddies per year. Accordingly, processes in the Mozambique Channel have downstream implications for the Agulhas Current including leakage into the South Atlantic Ocean (e.g., Bryden et al., 2005; Biaostoch et al., 2008, 2009b). Mozambique Channel eddies may also be linked with interannual modes of Indian and Pacific Ocean variability, implying a connection with the Indian Ocean Dipole and Pacific Ocean La Niña/El Niño phases (Schouten et al., 2002a; Palastanga et al., 2006).

Limited paleoceanographic evidence currently exists to document the longer term linkage between variability in the northern source waters of the Agulhas Current and its eventual leakage into the South Atlantic Ocean. The primary objectives at Site U1476 are to

- Recover a complete Pliocene–Pleistocene sedimentary succession including the early Pliocene warm period, the mid-Pliocene expansion of Northern Hemisphere ice sheets, and the MPT;

- Reconstruct Mozambique Current warm-water transports during periods of orbitally modulated and suborbitally accelerated climate changes; and
- Assess the influence of remote upstream forcing on southeast African warm-water transport.

Specific questions that will be addressed include Did the restriction of the Indonesian Seaway and the associated change of the Indonesian Throughflow in the Pliocene affect the Agulhas Current source region in the western tropical Indian Ocean and subsequently impact warm-water and saltwater transports along the southeast African margin? Did the Agulhas Current respond to the long-term development of tropical climates and the associated weakening of the monsoons in the course of the global cooling of the Pliocene?

Site U1476 consists of five holes that penetrated from 5.7 to 234.8 m DSF. The APC system penetrated a total of 841.1 m and recovered 873.83 m of core (104%). One interval was advanced 1.0 m without coring to adjust coring gaps to provide a continuous stratigraphic sequence. The total time spent at Site U1476 was 4.1 days (Table T1).

### Sedimentology

Two lithologic units were described at Site U1476. Unit I (0–34.57 m CSF-A) is composed of light brown to greenish gray foraminifer ooze with nannofossils alternating with foraminifer-rich nannofossil ooze. Unit II (34.57–235.41 m CSF-A) is composed of greenish gray foraminifer-rich nannofossil ooze and nannofossil ooze with foraminifers, silt, and clay.

Sediment in both units includes dark gray mottling that we interpret as bioturbation. Thin darker bands commonly surround burrows, and macroscopic pyritized burrows are common. Sediment also includes green layers that predominantly consist of pyrite and glauconite, which are diagenetic alterations. Only minor coring disturbance is observed in the cores.

### Physical properties

Despite the relatively homogeneous lithology, different physical parameters show significant trends downhole as well as cyclic variations along the sediment column. High-amplitude cyclic changes occur in the uppermost 16 m CSF-A of all holes. These cyclic changes continue downward to the bottom of the holes with lower amplitude and higher frequency. Variations in color reflectance indicate sediment compositional changes and RGB data broadly correspond with color reflectance parameter  $L^*$ . Bulk densities increase downhole due to compaction, but there are large variations in the uppermost 16 m CSF-A that are also seen in the other physical properties.

### Micropaleontology

Calcareous nannofossil and planktonic foraminifer biostratigraphy indicates that Hole U1476A spans the Late Pleistocene to late Miocene. Diatoms are barren throughout the section except for tropical diatoms present in minor amounts in the mudline sample. Calcareous microfossils show very good preservation in Pliocene-to Pleistocene-age sediment; however, the late Miocene samples are characterized by an increase of fragmented planktonic foraminifers. Assemblages in both microfossil groups include tropical to subtropical forms with the periodic appearance of species typical of temperate environments. Sedimentation rates, based on calcareous nannofossil and planktonic foraminifer datums, increase from 2.3 to 3.5 cm/ky downhole. No hiatuses were observed; however, events

that are seemingly diachronous were observed within the planktonic foraminifer biochronology. These events might suggest that calibration of microfossil datums from the tropical western Indian Ocean need refinement.

### Paleomagnetism

Paleomagnetic and rock magnetic analyses were carried out on sediment cores from Holes U1476A, U1476B, U1476D, and U1476E. All together, the rock magnetic data imply that the concentration of high-coercivity minerals (e.g., hematite and titanomagnetite) is relatively stable, whereas magnetite concentrations decrease below ~150 m CSF-A. Demagnetization of NRM carried out on discrete samples and archive halves shows that the directional record carries a strong coring overprint. The uppermost two sections of each core reveal steep downward inclinations after demagnetization at maximum levels of 25 mT, suggesting that the applied field was too weak to remove the overprint. More detailed demagnetization experiments on discrete samples show changes in inclination after demagnetization of 40 mT, suggesting that sediment has an early diagenetic overprint that might be removed by demagnetization at higher fields. Because the Icefield MI-5 orientation tool malfunctioned, the declination data could not be corrected for core orientation and can therefore not be used for assigning polarity zones. A more detailed investigation of demagnetization behavior will be needed to establish a more confident paleomagnetic stratigraphy for Site U1476.

### Stratigraphic correlation

Ideal coring conditions and good signal in the magnetic susceptibility measurements allowed for near-real time correlation at Site U1476. Coring offsets were applied in Holes U1476A and U1476D so that no gaps were present downhole to the full depth of the holes (229.0 m CSF-A). Hole U1476E provided similar offsets to Hole U1476A so that a splice could be constructed primarily between Holes U1476D and U1476E. The splice was constructed using magnetic susceptibility data and confirmed using NGR and RGB (blue) data.

### Geochemistry

Interstitial water chemistry shows mild early sediment diagenesis at Site U1476. Nitrate and sulfate, two species that disappear with progressively more intense microbial respiration, persist deeper into the sediment column and at higher concentrations than was documented at Sites U1474 or U1475. Methane concentrations remain at or near background levels. Decreasing concentrations of major elements, such as potassium and sodium, reflect uptake by clay minerals. Carbonate is variable (45–75 wt%), with terrigenous sediment making up the remainder of the sediment. The sediment is organic carbon poor, with average concentrations of 0.24 wt% total organic carbon (TOC).

## Site U1477

Site U1477 is located in the western Mozambique Channel on the upper continental slope (19°21.29'S; 36°54.90'E), ~65 nmi east of the Zambezi River delta region, at a water depth of 429.2 mbsl.

Site U1477 receives fluvial material exported from the Zambezi Catchment, which is in the southernmost reach of the African monsoonal precipitation belt and drains multiple subbasins. Collectively, these subbasins span ~1.385 million km<sup>2</sup> and the river itself extends more than 2500 km. Annual precipitation across the Zambezi Catchment varies spatially with ~150 cm/y in the northern

catchment region, ~85 cm/y in the southern region, and 60–70 cm/y in the central region (Ronco et al., 2006). Peak precipitation in the Zambezi Catchment occurs during the austral summer (DJF), when the ITCZ is located at its most southerly position.

The mean outflow of the Zambezi River is 3000 m<sup>3</sup>/s (Gammelstrød, 1992), and it transports an annual sediment load of ~51 Mm<sup>3</sup>, making it the largest single source of suspended sediment supply to the Mozambique margin (Walford et al., 2005). This fluvial sediment mixes with current-transported and biogenic material and forms a major component of the extensive current-controlled sedimentary deposits that characterize the Mozambique margin. Recent work on shorter cores demonstrates that the strata here have the potential to provide comprehensive proxy records of regional oceanographic changes as well as changes in hydrological conditions of the African precipitation belt in southeastern Africa (e.g., Schefuß et al., 2011; Schulz et al., 2011; Wang et al., 2013; Weldeab et al., 2013, 2014; Just et al., 2014; van der Lubbe et al., 2014, 2016).

Site U1477 is positioned on a contourite drift deposited under the influence of the eddy-dominated water flow through the Mozambique Channel. The drift displays packages of parallel and laterally continuous subseafloor reflectors. Site U1477 is immediately southwest of the 6.51 m long sediment core GeoB9307-3 (18°33.99'S, 37°22.89'E; 542 mbsl) (Schefuß et al., 2011). The site was selected because Core GeoB9307-3 showed elevated sedimentation rates during the deglacial period to the Holocene and thus suggested the great potential of Site U1477 for providing a centennial-scale paleoceanographic record. As noted below, the site had almost 10 times higher accumulation than Core GeoB9307-3, so the primary objectives for the site are necessarily modified from the original objectives.

The Site U1477 primary objectives are to

- Recover a complete high-resolution sedimentary succession that spans the last 2 My, including the MPT, from the Mozambique margin close to the Zambezi River delta (we only recovered <200 ky because of the extremely high accumulation rate);
- Generate a high-resolution Pleistocene profile of southern African continental climate changes on orbital and suborbital time-scales;
- Establish linking between southern African terrestrial climates and southwest Indian Ocean heat budgets, notably warm-water transport along the southeast African margin and associated ocean–atmosphere heat and moisture transfer;
- Examine the relationship between such climate variability and early human evolution (this objective could not be achieved due to the extreme accumulation rate); and
- Reconstruct upstream control on Agulhas leakage through headwater variability during periods of orbitally modulated and suborbitally accelerated climate changes.

The stratigraphic extent of Site U1477 was considerably shorter than anticipated, with a basal age of <200 ka. As a result, the paleoclimate site objectives will require adjustment to focus on high-temporal timescale variability (decadal–centennial) over the last glacial cycle. The remarkably high sedimentation rates at Site U1477 promise an unprecedented opportunity to investigate hydroclimate variability in southeastern Africa and its potential links to changes in the greater Agulhas Current system. A specific question that will be addressed is whether variable Agulhas Current warm-water transport contributes to and modulates the impacts of shifting air boundaries and rainfall patterns over southeast Africa, and

possibly offsets these patterns from those over southwest Africa and the Namib (Dupont et al., 2005, 2011; Maslin et al., 2012; Castañeda et al., 2016).

Site U1477 consists of three holes that penetrated from 119.4 to 181.2 m DSE. A total of 490.01 m of core was taken over a 474.0 m cored interval (103%). The APC system penetrated a total of 296.2 m and recovered 308.07 m of core (104%). The HLAPC recovered 174.00 m of sediment over a 169.2 m interval (103%). The extended core barrel (XCB) was deployed for one core resulting in 7.96 m of sediment (97%). One interval was advanced 1.0 m without coring to offset the otherwise aligned coring gaps for obtaining a continuous stratigraphic sequence. The total time spent at Site U1477 was 1.8 days (Table T1).

### Sedimentology

Sediment at Site U1477 consists of two lithologic units. Unit I (0–0.5 m CSF-A) is composed of very dark gray to dark greenish gray sandy clay with foraminifers and nannofossils. Unit II (0.5–254.07 m CSF-A) is composed of very dark gray to dark greenish gray silty clay with intermittent to frequent occurrences of sand layers. These sand layers contain quartz and clay, and thicknesses range from centimeter to meter scale.

Bioturbation is not apparent in either unit. Macrofossils, including bivalve and echinoderm fragments, are present throughout the cores. Many of the cores show disturbance related to gas expansion.

### Physical properties

Despite the relatively homogeneous lithology, physical properties at Site U1477 show a complex pattern of high-amplitude changes that are likely caused by climate-related variations in terrigenous sediment supply and/or current sorting effects. However, some excursions to lower values in NGR and GRA bulk density data may be due to the voids caused by gas expansion. Also, collection of reliable *P*-wave data below 16.9 m CSF-A in all of the holes was inhibited by strong signal attenuation resulting from the frequent occurrence of gas expansion voids. Porosities at Site U1477 decrease downhole, and average porosities are significantly lower when compared to the previous sites (U1474–U1476).

### Micropaleontology

Site U1477 spans the Holocene to the Late Pleistocene. Two biostratigraphic datums were observed: the abundance crossover from *Gephyrocapsa caribbeanica* to *Emiliana huxleyi* at ~90 ky and the top occurrence of *Globigerinoides ruber* (pink) at ~120 ky. Sedimentation rates, estimated from biostratigraphic datums as well as radiocarbon ages tied in from a nearby piston core (van der Lubbe et al., 2014), suggest deposition rates of ~1.5 m/ky from 0 to ~30 m CSF-A in Hole U1477A that decrease to ~83 cm/ky from ~30 m CSF-A to the base of the hole. Calcareous microfossils are generally well preserved, although abundance is typically low. Planktonic foraminifer tests compose <1% of sediment particles in most samples, and fewer than 50 nannofossils per field of view were observed throughout most of the section. The planktonic assemblage includes the regular occurrence of pteropods, benthic foraminifers, ostracods, sponge and gorgonian spicules, mollusks, echinoderm plates, and fish otoliths. Siliceous microfossils make up a minor component of the mudline and core catcher samples where marine and freshwater diatoms, phytoliths, and sponge spicules are present, providing indications of both terrestrial and marine sources. Siliceous microfossils were not found in samples below the mudline.



## Paleomagnetism

Paleomagnetic measurements at Site U1477 show a clear magnetic signal. For all holes, inclinations center around  $-50^\circ$  as expected for the site location during the Brunhes Chron. Rock-magnetic data acquired on discrete samples reveal phases of high ferrimagnetic mineral content that largely match magnetic susceptibility data collected on archive section halves and whole cores. Variations in magnetic mineral compositions reflect either a change in the composition of the terrigenous fraction or preferential mineral dissolution.

## Stratigraphic correlation

Real-time stratigraphic correlation was possible at Site U1477 using magnetic susceptibility data. Holes U1477A–U1477C were used for the splice from 0 to  $\sim 85$  m CCSF-A. Below  $\sim 85$  m CCSF-A, the splice consists primarily of sediment from Hole U1477B, with a couple of exceptions where intervals of Hole U1477A were used. A continuous stratigraphic splice was not possible to achieve at this site because of substantial gas expansion that led to core fracturing and variable core stretching between holes.

## Geochemistry

Interstitial water chemistry and headspace gas concentrations show relatively intense early sediment diagenesis at Site U1477. The sulfate–methane transition appears at  $\sim 16$  m CSF-A, below which sulfate is completely consumed and methane concentrations increase rapidly to a peak of  $\sim 75,500$  ppmv at  $\sim 35$  m CSF-A. Other redox-sensitive elements, including iron, manganese, and barium, show enrichments caused by remobilization. Terrigenous sediment is the dominant constituent of the core. Carbonate contents average 5 wt%, with a few intermittent layers rich in shell material with carbonate contents as high as  $\sim 20$  wt%. The average TOC content is 0.61 wt%.

## Site U1478

Site U1478 is located in the Delagoa Bight on the Inharrime Terrace ( $25^\circ 49.26'S$ ;  $34^\circ 46.16'E$ )  $\sim 75$  nmi east of the Limpopo River at a water depth of 488 mbsl.

The Delagoa Bight is a distinct indentation of the continental margin (Lamont et al., 2010) into which the Limpopo River, the second largest eastward-draining river in Africa, deposits sediment. The hydrography of the Delagoa Bight is influenced by the southerly flowing waters from the Mozambique Channel and the East Madagascar Current (Lutjeharms, 2006a, 2006b), the confluence of which forms the Agulhas Current. The Limpopo Catchment is a little over  $410,000$  km<sup>2</sup>, considerably smaller than that of the neighboring Zambezi River. The Limpopo River has a mean annual discharge of  $\sim 170$  m<sup>3</sup>/s and delivers annual sediment loads of 33 Mt/y (Milliman and Meade, 1983) to the Delagoa Bight.

The Site U1478 primary objectives are to

- Recover a complete high-resolution sedimentary succession that spans the last 2 My;
- Generate a high-resolution Pleistocene profile of southern African continental climate changes on orbital and suborbital time-scales;
- Establish linking between southern African terrestrial climates and southwest Indian Ocean heat budgets, notably warm-water transports along the southeast African margin and associated ocean–atmosphere heat and moisture transfer;
- Examine the relationship between such climate variability and early human evolution; and

- Reconstruct upstream control on Agulhas leakage through headwater variability during periods of orbitally modulated and suborbitally accelerated climate changes.

Specific questions that will be addressed include Did the long-term climatic developments of the Pleistocene, through their impact on atmospheric circulation, alter the rainfall patterns over southeast Africa? Did variable Agulhas Current warm-water transports contribute to and modulate the impacts of shifting air boundaries and rainfall patterns over southeast Africa and possibly offset these patterns from those over southwest Africa and the Namib (Dupont et al., 2005; Maslin et al., 2012)?

Site U1478 consists of four holes that penetrated from 216.0 to 248.4 m DSF. A total of 922.1 m of core was recovered over a 905.7 m interval (102%). The APC system penetrated a total of 698.9 m and recovered 708.66 m of core (101%). The HLAPC recovered a total of 213.4 m of sediment over a 206.8 m cored interval (103%). Five intervals were advanced 54.3 m without coring to adjust coring gaps to provide a continuous stratigraphic sequence. The total time spent at Site U1478 was 2.7 days.

## Sedimentology

Sediment at Site U1478 consists of one lithologic unit. Unit I (0–248.59 m CSF-A) is composed of dark olive-gray and dark greenish gray sand with foraminifers and nannofossils that alternate with clayey or sandy silt with foraminifers and nannofossils. The sand is predominantly composed of quartz.

Bioturbation is rare; however, the coarser grained sediment at Site U1478 may obscure sedimentary structures such as bioturbation and diagenetic alterations. Macrofossils, including bivalves, gastropods, and echinoderms, are present throughout the cores.

## Physical properties

The physical properties measured at Site U1478 show a complex pattern of amplitude changes and trends. All high-resolution records derived by physical properties core logging (magnetic susceptibility, NGR, GRA bulk density, RGB color, and spectral reflectance) show cyclic changes that are likely related to the variable lithologic alternations between silt-rich and sand-rich intervals. These cyclic changes are pervasive throughout the sequence and probably reflect complex interactions between ocean circulation, sea-level fluctuations, and fluvial discharges recorded at Site U1478. Two zones of very low magnetic susceptibility values at  $\sim 89$ – $92$  and  $\sim 164$ – $168$  m CSF-A may be related to sediment diagenesis as indicated by elevated iron and manganese concentrations in the interstitial water. Porosities decrease downhole, and average porosities are similar to those at Site U1477 but are significantly lower compared to the deeper water Sites U1474–U1476.

## Micropaleontology

Biostratigraphy of calcareous nannofossils and planktonic foraminifers indicates that Hole U1478A spans the Late Pleistocene to the middle Pliocene with a basal age of  $\sim 4$  Ma. Calcareous microfossil assemblages include warm subtropical to tropical species in combination with coastal taxa. Planktonic foraminifers and calcareous nannofossils are abundant throughout Site U1478. Foraminifers dominate sandy beds, whereas calcareous nannofossils are typically more abundant in silty and clay-rich intervals. Glassy and translucent foraminifers are found throughout most of the recovered sequence, although fragmentation is apparent in the oldest parts of the sequence. Calcareous nannofossils are mostly well preserved, but the *Discoaster* group commonly exhibits fragmentation of the

delicate arms. Diatoms are present only in surface sediment, and sponge spicules and phytoliths are scarce in the lower parts of the sequence. Low levels of reworking are evident in planktonic foraminifer records. The biochronology for both calcareous microfossil groups reveals sedimentation rates of ~9 cm/ky between the Late Pleistocene and ~2.1 Ma and lower rates of 2.9–5.3 cm/ky older than 2.1 Ma. Planktonic foraminifers and calcareous nannofossils give diverging age-depth relationships between 1.9 and 3 Ma, which is possibly linked to sedimentological processes. The two fossil groups are in agreement in the lowermost ~25 m of Site U1478 and indicate a basal age of 4–4.1 Ma.

### Paleomagnetism

Rock magnetic and paleomagnetic analyses were carried out on discrete samples from Hole U1478A and archive-half sections from Holes U1478A–U1478D. SIRM and hard isothermal remanent magnetization (HIRM) measured on discrete samples largely follow downcore patterns in magnetic susceptibility. The *S*-ratio is mostly stable; however, distinct minima at ~90, 165, and 245 m CSF-A indicate a strong depletion of low-coercivity minerals. These intervals coincide with minima in the magnetic susceptibility and SIRM records. The minimum in magnetic susceptibility at ~90 m CSF-A lies below the sulfate–methane transition, and the minimum at ~165 m CSF-A is located just below an interval of elevated dissolved Fe and Mn concentrations, suggesting that in these zones early diagenetic processes affected the magnetic mineral record.

Inclination data appear to be of good quality, and the Brunhes/Matuyama boundary and the Jaramillo Subchron were identified. According to biostratigraphy, the sedimentary sequence dates to ~4 Ma; however, the Matuyama–Gauss Chron transition was not identified in the paleomagnetic data which might relate to a loss of the paleomagnetic signal because of diagenetic processes.

### Stratigraphic correlation

Four holes were drilled at Site U1478 using the APC and HLAPC. Generally good core quality and clear magnetic susceptibility signals resulted in robust and unambiguous stratigraphic ties between holes. We have high confidence that these holes constitute a nearly complete section spanning the full length of the longest hole (248 m recovered in Hole U1478A), with the exception of one probable gap at ~195 m CSF-A. The objective of obtaining a complete section was aided by the fact that the magnetic susceptibility from whole-round physical property loggers could be used for real-time correlation in Holes U1478B–U1478D; adjustments to the drilling depth were made in these holes for the purpose of avoiding potential core gap alignment. A continuous splice was constructed for the uppermost ~200 m CCSF-D, and a floating splice of another ~50 m of section is appended below the coring gap. In general, the clarity of the stratigraphic ties and the multiple cross-checks from several different data sets gives us high confidence in the integrity of these two separate spliced sequences.

### Geochemistry

Interstitial water chemistry and headspace gas concentrations show relatively intense early sediment diagenesis at Site U1478. The sulfate–methane transition zone occurs at ~70 m CSF-A, below which sulfate is completely consumed and methane concentrations increase rapidly. Several methane peaks occur below the sulfate–methane transition, with the largest of ~65,000 ppmv at ~214 m CSF-A. Other redox-sensitive elements, including iron and manganese, also show multiple peaks at depth. These pulses of iron and

manganese and associated lows in magnetic susceptibility are interpreted as paleoredox horizons related to repeated fluctuations in sedimentation. Detrital material is the dominant sedimentary component. Carbonate contents average 28 wt%, and TOC contents average 0.5 wt%; neither show any trend with depth.

### Site U1479

Site U1479 is located on a 30 km wide morphological high, rising ~200 m above the regional seafloor, on the mid-to-lower western slope of the Agulhas Bank in Cape Basin (35°03.53'S; 17°24.06'E), ~85 nmi southwest of Cape Town, South Africa, at a water depth of 2615 mbsl.

This region has been the focus of numerous oceanographic studies and a growing number of paleoceanographic studies (e.g., Beal et al., 2011, and references therein) in order to describe the present hydrography within the so-called Indian-Atlantic Ocean gateway (or “Agulhas ring corridor”; Gordon, 2003) and also to gain a better understanding of past oceanographic variations in the area and how they may be related with global climate changes. Several studies have highlighted the Agulhas Current's sensitivity to climate changes and reorganizations of atmosphere and the ocean circulation at large.

As the Agulhas Current detaches from the Agulhas Bank at the southern tip of Africa, it retroflects, shedding rings, eddies, and filaments to the South Atlantic. This leads to leakage of between 2 and 20 Sv of Indian Ocean water into the Atlantic Ocean and affects the water column to depths of 2000 mbsl (Gordon et al., 1987; de Ruijter et al., 1999a; Richardson, 2007).

The Agulhas ring corridor is an area in which retrieval of high-quality sediment cores is difficult. The vigorous Agulhas Current and its variability can cause sediment redistribution because meanders and eddies frequently reach the seafloor. However, a number of published studies and emerging work demonstrate that paleoceanographic records of excellent quality are in fact achievable for reconstructing both surface and deep-water variability in the region (e.g., Rau et al., 2002, 2006; Peeters et al., 2004; Dickson et al., 2010; Martínez-Méndez et al., 2010; Caley et al., 2011, 2012, 2014; Marino et al., 2013; Dyez et al., 2014; Petrick et al., 2015).

The Site U1479 primary objectives are to

- Recover a complete Pliocene–Pleistocene sedimentary succession, including the early Pliocene warm period, mid-Pliocene expansion of Northern Hemisphere ice sheets, and the MPT, from a high-accumulation site located within the Agulhas ring corridor;
- Reconstruct Agulhas Current warm-water transports during orbitally modulated and suborbitally accelerated climatic changes;
- Assess the linking between Antarctic climate variations, circum-polar ocean front instability, and connections with Agulhas leakage into the South Atlantic Ocean;
- Assess the vigor and hydrography of NADW, or its precursors, exported to CDW and the southwest Indian Ocean; and
- Evaluate the possibility of advective salinity feedbacks between Agulhas leakage and AMOC variability, notably the possible role of the leakage in modulating surface-to-deep ocean coupling in the North Atlantic during the transition between climatic states.

Site U1479 consisted of nine holes, ranging in depth from 1.0 to 300.7 m DSF. Overall, 105 cores were recorded for the site. A total of 963.08 m of core over a 975.3 m interval was recovered using the APC system (99% core recovery). Seven intervals were advanced

without coring over 175.7 m. The total time spent at Site U1479 was 6.0 days.

### Sedimentology

Sediment at Site U1479 consists of one lithologic unit. Unit I (0–301.03 m CSF-A) is composed of light greenish gray, greenish gray, and light olive-gray (5Y 6/2) nannofossil ooze with or without foraminifers.

Moderate to strong bioturbation is the only secondary sedimentary structure observed and is widespread throughout the sequence. Discrete ichnofossil burrows were commonly observed, often surrounded by thin darker bands. The occurrence of intervals rich in sand-sized quartz and foraminifers further suggests the deposition of contourites and/or turbidites.

### Physical properties

All high-resolution physical property records derived by core logging show similar cyclic changes that are likely related to alternations within the nannofossil ooze between intervals of more or less terrigenous content and thus may reflect changes in ocean circulation and/or terrigenous supply. Generally, the changes in physical properties trends broadly correlate between the parameters. A slump in the lowermost part of Site U1479 is characterized by physical properties that are distinctly different from the rest of the cored section.

### Micropaleontology

Biostratigraphy of calcareous nannofossils, planktonic foraminifers, and diatoms indicates that Hole U1479B spans the Late Pleistocene to the late Miocene with a bottom age of ~7 Ma. Calcareous microfossils include primarily subtropical convergence and coastal-to slope-water taxa with minor abundances of warm subtropical and subpolar taxa. Diatoms are mostly composed of subtropical taxa with a minor presence of Southern Ocean taxa in the Pleistocene. Planktonic foraminifers and calcareous nannofossils are abundant throughout Site U1479, whereas diatoms make episodic appearances, particularly in the upper and lower parts of the record. Preservation is good to moderate for calcareous taxa and moderate to poor for diatoms. The biochronology for both calcareous microfossil groups and diatoms is consistent with an average sedimentation rate of ~4.9 cm/ky.

### Paleomagnetism

Paleomagnetic analyses were carried out on archive halves from Site U1479 and on discrete samples from Hole U1479B. The inclination and declination data from the archive halves are relatively noisy, which might be partially explained by (1) high dilution by carbonate and (2) the alteration of magnetic minerals as suggested by downhole records of various rock magnetic parameters including the *S*-ratio, HIRM, SIRM, and magnetic susceptibility. Inclinations and declinations show an interval of normal polarity in the uppermost ~55 m CSF-A (Brunhes Chron), although the lower boundary is deeper than would be anticipated from the biostratigraphy. A distinct interval of reversed polarity between ~55 and ~125 m CSF-A corresponds probably to the Matuyama Chron, which spans from ~2.6 to 3.6 Ma. Unlike the Brunhes/Matuyama Chron boundary, the inferred Matuyama/Gauss Chron boundary is in agreement with ages derived from the biostratigraphy. Additional shore-based paleomagnetic analyses on rock magnetic properties and directional data are necessary to refine the magnetostratigraphy for Site U1479.

### Stratigraphic correlation

Stratigraphic correlation of Holes U1479B–U1479I was complicated by very low magnetic susceptibility and adverse weather conditions resulting in several partial core recoveries in all the holes, as well as soupy top sections that were generally unusable for stratigraphy. Notwithstanding, the NGR and color (e.g., RGB green) records display clear variability in all undisturbed core sections, making the alignment of cores relatively straightforward once all scanning was complete. Furthermore, comparatively constant offsets between cores in Holes U1479B and U1479C led to the development of a robust and nearly complete splice from just below the mudline to the bottom of Hole U1479B at ~300 m CSF-A. The generally clear stratigraphic relationship between Holes U1479B and U1479C only needed refinement and supplementation in scattered discrete intervals.

### Geochemistry

Interstitial water chemistry and headspace gas concentrations indicate moderate early sediment diagenesis. Major element concentrations show some evidence of incorporation into clay minerals and precipitation of authigenic carbonates. There is no clear sulfate–methane transition zone in Hole U1479B—sulfate concentrations are zero at ~185 m CSF-A, but methane concentrations are never above ~50 ppmv. Other redox-sensitive elements, such as iron and barium, show more dramatic enrichments, with iron reaching ~15  $\mu$ M at ~60 m CSF-A and barium reaching ~600  $\mu$ M at ~150 m CSF-A. Calcium carbonate ranges from ~51 to 91 wt%, with an average value of 67 wt% and no distinct trend downhole. Organic carbon ranges between 0.15 and 1.93 wt% with an average value of 0.62 wt%.

## Expedition synthesis and preliminary scientific assessment

The greater Agulhas Current system is a key location to study past ocean circulation and climate, as it represents a significant component of the global thermohaline circulation, carrying surface water into the Atlantic Ocean that compensates for the export of deep water from the Atlantic Basin to the world ocean (Figure F1). The Agulhas Current marks the western boundary of the South Indian Ocean subtropical gyre where it flows southwards along the southeastern African continental margin. At the southern tip of South Africa the Agulhas Current, under its own inertia, separates from the margin and retroflects back into the Indian Ocean flowing as the ARC. During the retroflexion the Agulhas current periodically sheds large (up to 400 km) mesoscale anticyclonic Agulhas rings (Figure F2), which represent a significant pathway for warm and high-salinity water transfer from the Indian Ocean into the South Atlantic Ocean. Modulation of this so-called “Agulhas leakage” variably impacts the buoyancy structure there and also likely contributes to mode changes and variability of the AMOC and thus climate variability in the North Atlantic region and beyond. Additionally, variations in Agulhas warm-water transports along the southeast African continental margin foster exchanges of heat and moisture with the atmosphere that influence regional climates in southern Africa.

Operations during Expedition 361 drilled six sites in the Indian–Atlantic Ocean gateway in February and March 2016 (Figure F3). In total, we recovered 5175 m of core, with an average recovery of 102%, during 29.7 days of on-site operations, from a region poorly



represented previously in the database of drill sites for scientific purposes. No holes were logged. This is considerably less than the initially planned campaign targeting 5811 m of section, including downhole logging at four sites (Hall et al., 2015). The difference amounts to ~11 days of lost operational time. A medical evacuation accounted for 1.8 days and weather a further 1.5 days of lost operations. The remainder of the lost operational time was due to delays in attaining the necessary permissions to operate in Mozambique's exclusive economic zone (EEZ) waters, which necessitated a revision of the drilling plan and 2840 nmi of additional transit to Expedition 361.

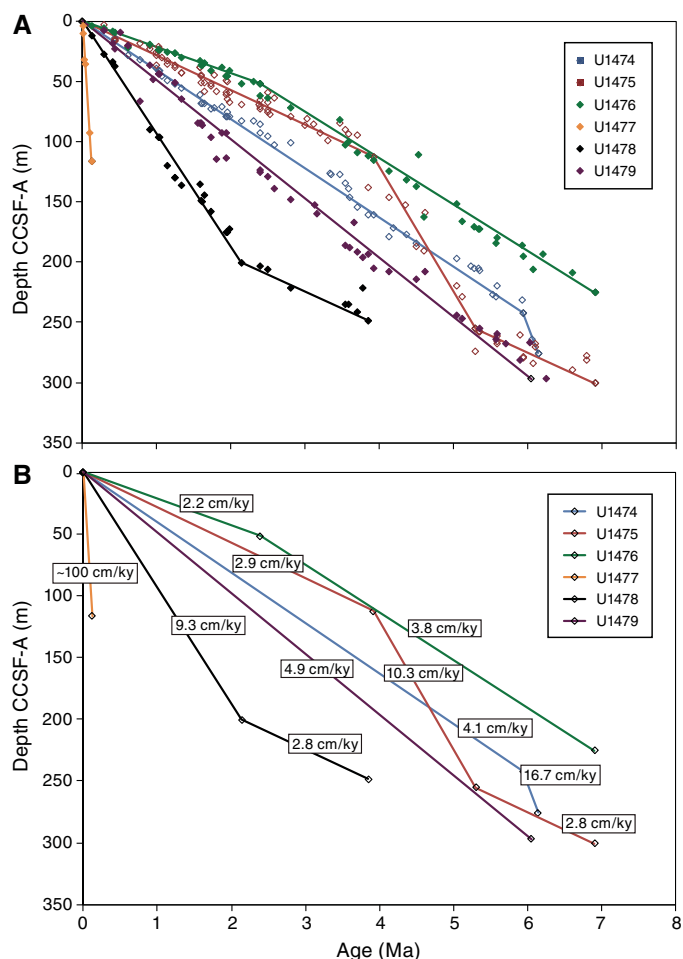
Expedition 361 South African Climates (Agulhas LGM Density Profile) had the following five fundamental objectives:

- To establish the sensitivity of the Agulhas Current to changing climates of the Pliocene–Pleistocene in association with transient to long-term changes of high-latitude climates, ocean gateways, tropical heat budgets, and the monsoon system;
- To determine the dynamics of the Indian–Atlantic Ocean gateway circulation during such climate changes in association with changing wind fields and migrating ocean fronts;
- To examine the connection between Agulhas leakage and ensuing buoyancy transfer and shifts of the AMOC during major ocean and climate reorganizations during at least the last 5 My;
- To address the impact of Agulhas variability on southern Africa terrestrial climates and, notably, rainfall patterns and river runoff; and
- To constrain the temperature, salinity, and density structure of the LGM deep ocean from the bottom of the ocean to the base of the main thermocline.

The expedition made major strides toward fulfilling each of these objectives. The recovered sequences allowed complete spliced sections to be generated that span the interval of 0 to between ~0.13 and 7 Ma (Figure F11). The sedimentary compositions (Figure F12) range from highly terrigenous near the Zambezi (Site U1477) and Limpopo (Site U1478) Rivers and in the Natal Valley (Site U1474) to very carbonate rich in the northern Mozambique Channel (Site U1476), the Agulhas Plateau (Site U1475), and Cape Basin (Site U1479). The carbonate fossil preservation is good in the Pleistocene and mostly good to moderate, tending to be a little worse near the bottom of each sequence (Figures F13, F14). A robust biostratigraphy was developed for each Site using nannofossils and foraminifers (Figures F15, F16). At Sites U1475 and U1479, diatoms are abundant enough to provide further biostratigraphic constraints, and at these sites there is potential for future paleoenvironmental work with diatoms (Figure F17). All the fossils showed a mixture of tropical and subtropical convergence and temperate or subpolar species in the southernmost sites (Figures F17, F18, F19). Changes in these assemblages through time will provide important constraints on the dynamics of the Agulhas Current system.

For Sites U1475, U1476, U1478, and U1479, sea conditions and time were sufficient that geochemical compositions were measured on a suite of solid bulk samples from the carbonate sample residue. As Ba/Al ratios may be a good X-ray fluorescence proxy for productivity, Figure F20 displays the Ba/Al ratios of bulk samples against TOC. Site U1477, with an extremely high terrigenous load, shows no correlation and has Ba/Al ratios of typical silicate detritus. Site U1478 also has mostly terrigenous Ba/Al ratios, but a few samples stand above. Site U1476 is in an area of low productivity, so the range of Ba/Al ratios is low and not correlated with TOC. However,

Figure F11. Age–depth relationships, Sites U1474–U1479. A. Time estimates based on a combination of major planktonic foraminifer, calcareous nannoplankton, diatom, and paleomagnetic datums. B. Implied sedimentation rates.

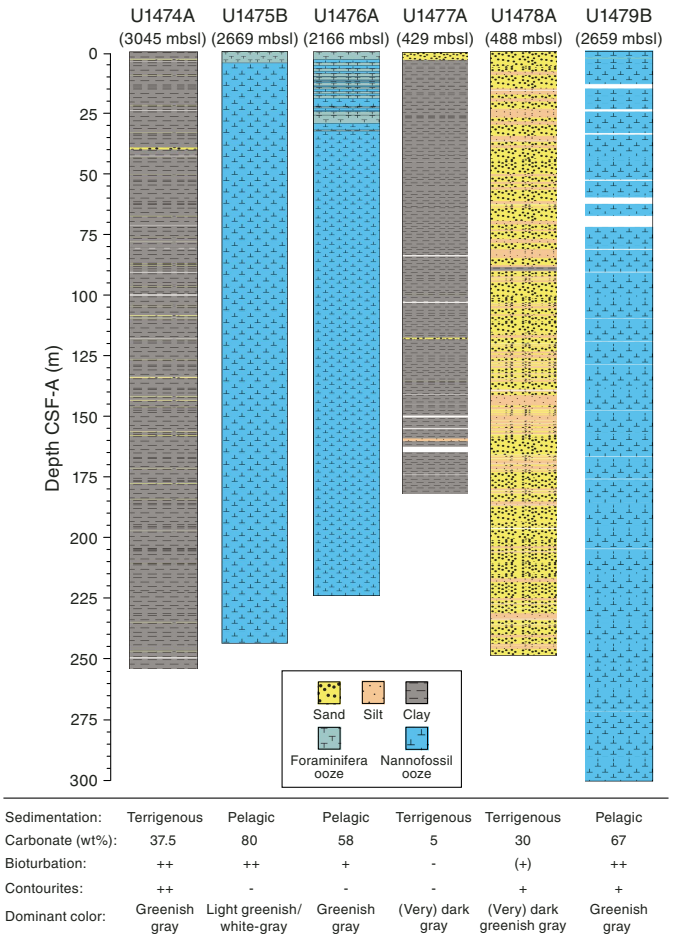


it is elevated above terrigenous, indicating at least 50% of the Ba is probably from biogenic sources. Site U1475, located on the Agulhas Plateau where we would expect variations with the position of the STF, has high Ba/Al ratios and has a generally strong positive trend with TOC with the exception of a couple of samples.

Early sediment diagenesis, driven by microbial respiration of organic matter in sediment, defines many dissolved elemental profiles. One of the characteristic profiles that indicates the extent and intensity of these reactions is dissolved sulfate (Figure F21). Where microbial respiration is intense as a result organic matter availability sulfate is rapidly consumed. Where microbial respiration is more moderate sulfate decreases more gradually and sometimes persists to the base of the measured interval. We see a range of diagenetic intensities across the Expedition 361 sites, with the most intense microbial activity and diagenesis at Site U1477 (Zambezi) and the least activity at Site U1476 (Mozambique Channel). It is likely that the observed differences between sites are driven by location and sedimentation rate rather than organic carbon content because TOC does not vary systematically between sites (Figure F20).

Operations at Site U1474 (Natal Valley; Figure F3) were scheduled to produce three APC/XCB holes to 300 meters below seafloor (mbsf), with logging in one of the deep holes and a single hole to

Figure F12. Summary of major lithologic characteristics of sediment cores, Sites U1474–U1479. Sites U1474, U1477, and U1478 are characterized by terrigenous sedimentation with carbonate content between 5 and 37.5 wt%. Sites U1475, U1476, and U1479 are characterized by pelagic sedimentation with carbonate content between 58 and 80 wt%. Sediment color ranges from white-gray (Site U1475) to greenish gray (Sites U1474, U1476, and U1479) and dark gray (Sites U1477 and U1478). Slight to moderate bioturbation occurs at Sites U1476–U1478, whereas strong bioturbation marks the sediment at Sites U1474, U1475, and U1479. Sandy layers (~0.5 cm to ~2 m thick) at Sites U1474, U1478, and U1479 are interpreted as contourite and/or turbidite deposits.



150 mbsf for interstitial water samples. This drilling objective was largely accomplished with the exception of logging and a reduction in the target section depth to 254 m CSF-A because of loss of operational time as noted above. Recovery was excellent, and triple-coring allowed compilation of a complete spliced record extending to the latest Miocene (~6.2 Ma). The site boasts an excellent paleomagnetic timescale (Figure F22) that is in good agreement with the biostratigraphy, and these together suggest uniform sedimentation rates of ~4.1 cm/ky over most of the sequence with an apparent increase near the bottom (Figure F11). This is the only site situated beneath the main flow of the fully constituted Agulhas Current and will therefore provide the opportunity for high-resolution climate reconstructions of Agulhas Current warm-water transports and upstream variability that may allow the identification of connections between Agulhas leakage and its headwater variability. The site also holds significant potential to investigate the connections between southern African terrestrial climates and southeast Indian Ocean heat budgets and notably the links to the cultural evolution of early

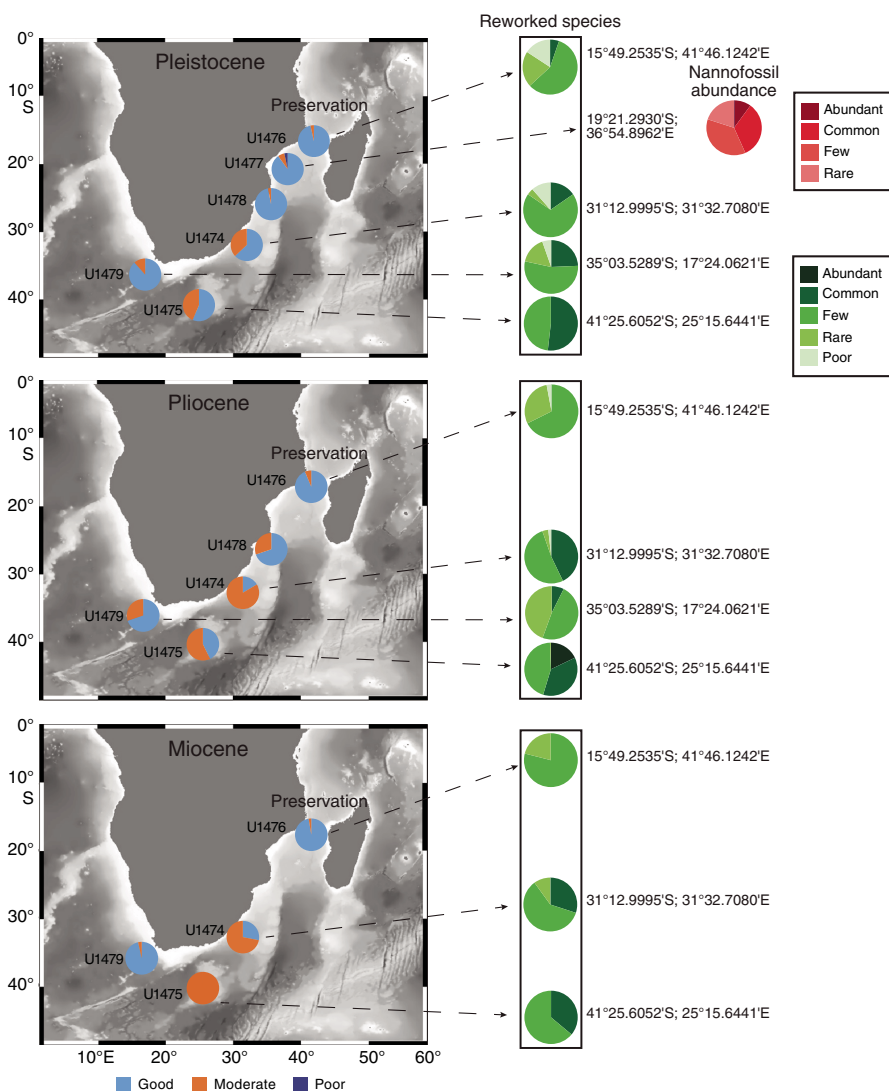
modern humans. A high-resolution program of interstitial water samples, collected from a 150 m CSF-A section, will provide an additional data point to the presently available pore water LGM temperature and salinity inventory.

Operations at Site U1475 (Agulhas Plateau; Figure F3) were scheduled for three APC/XCB holes to 300 mbsf with logging and a single hole to 150 mbsf for interstitial water samples. This drilling objective was largely accomplished with the exception of logging. Recovery was excellent, and multiple coring, including some spot coring needed because of weather-related disturbances, allowed recovery of a complete spliced record extending to the late Miocene (~7 Ma). The paleomagnetic timescale is robust to the base of the Mammoth Chron (C2An.2r at 99 m CCSF-A) (Figure F22) and is in good agreement with the biostratigraphy, but the signal was too weak to allow a paleomagnetic timescale to be extended to the bottom based on shipboard data. Biostratigraphic estimates suggest sedimentation rates of ~2.9 cm/ky in the uppermost 115 m CCSF-A, 10.3 cm/ky between 115 and 255 m CCSF-A, and 2.8 cm/ky between 255 m CCSF-A and the bottom of the section (Figure F11). This record will provide the opportunity for high-resolution climate reconstructions of the ARC and connections with the STE, productivity, and deep-water circulation. A high-resolution program of interstitial water samples, collected from a 150 m CSF-A section, will provide an additional data point to the presently available interstitial water LGM temperature and salinity inventory.

Operations at Site U1476 (Mozambique Channel; Figure F3) were scheduled for three APC/XCB holes to 300 mbsf with logging and a single hole to 150 mbsf for interstitial water samples. This drilling objective was largely accomplished with the exception of logging and a reduction in the target depth to 220 m CSF-A because of loss of operational time as noted above. Recovery was excellent, and triple coring allowed recovery of a complete spliced record extending to the late Miocene (~6.9 Ma). The site boasts excellent biostratigraphy and notably cyclic physical properties. The paleomagnetic reversal signal was complicated by apparently varying magnetic mineralogy, strong coring overprints and operational issues with the orientation tool. Shore-based work will be required to try to extract a paleomagnetic timescale at this site (Figure F22). The sedimentation rate based on biostratigraphy is ~2.2 cm/ky from the mudline to ~50 m CCSF-A and 3.8 cm/ky from ~50 m CCSF-A to the bottom (Figure F11). This site is located on Davie Ridge at the northern end of the Mozambique Channel, in a region with low productivity and a mean annual sea-air CO<sub>2</sub> flux is close to zero (Takahashi et al. 2009). The site will therefore provide the opportunity for high-resolution reconstructions of tropical faunal assemblages is promising for the identification of connections between Agulhas leakage and its headwater variability. The site also holds significant potential to investigate the connections between southern African terrestrial climates and southeast Indian Ocean heat budgets and thermocline and deep-water variability with likely links to the development of the Indonesian Throughflow as well as aridification of east Africa. Foraminifers show excellent preservation, thus also making this an ideal site for a long record of surface-ocean pH from boron isotopes. A high-resolution program of interstitial water samples collected from a 150 m CSF-A section will provide an additional data point to the presently available interstitial water LGM temperature and salinity inventory.

Operations at Site U1477 (Zambezi Catchment; Figure F3) were scheduled for three APC/XCB holes to 207 mbsf. This drilling objective was modified to end at APC refusal, which led to recovering ~166 m. Recovery was excellent, and triple coring allowed compila-

Figure F13. Summary of calcareous nannofossil preservation and reworking at Expedition 361 drill sites for the Pleistocene, Pliocene, and Miocene. Nannofossil preservation score is based on ratings of good, moderate, and poor to illustrate changes in preservation through time. Qualitative percentages of nannofossil group (red) and reworked species (green) abundances are indicated on the right with relative latitudinal and longitudinal coordinates for each site. Missing green charts from different time periods are due to the negligible number of reworked species. Site U1477 sequence covers only the Pleistocene; reworked species are sparsely recorded and it is the only site where nannofossil abundance is variable. The rest of the sites are characterized by a total abundance of calcareous nannofossils >80% of sediment particles within the three time intervals.



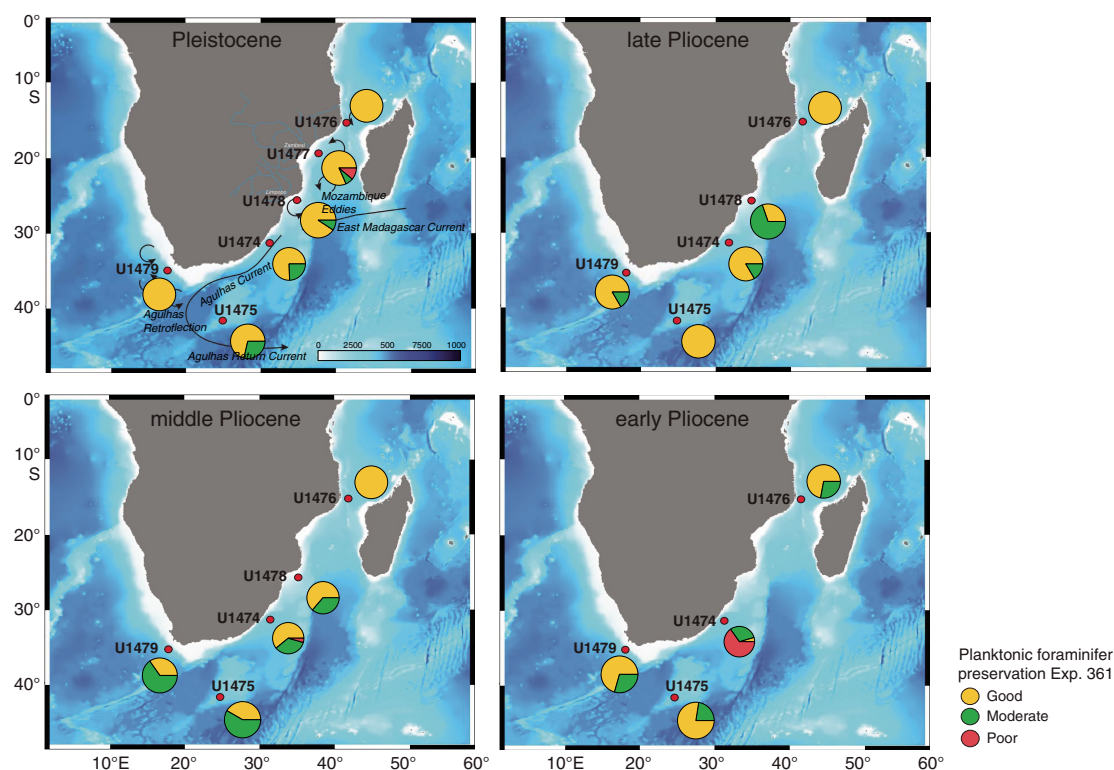
tion of a complete spliced record to ~85 m CCSF-A. The magnetic susceptibility signal is strong and could be matched to nearby  $^{14}\text{C}$ -dated cores, and two biostratigraphic markers, in agreement, pointed to sediment accumulation rates of ~1 m/ky at this site (Figure F11). This extreme accumulation rate was a complete surprise and led to only recovering ~120 ky of sediment record from the Zambezi Catchment. Therefore, the sediment sequence does not provide 2 My record as planned, but will instead provide the opportunity for extremely high resolution reconstructions of, for example, terrestrial climate and thermocline characteristics during the last glacial cycle. The site also holds significant potential to investigate the connections between southern African terrestrial climates and southeast Indian Ocean heat budgets and the links to the cultural evolution of modern humans.

Operations at Site U1478 (Limpopo Catchment; Figure F3) were scheduled to be three APC/XCB holes to 180 mbsf, with a permitted depth to 250 mbsf, and a single hole to 150 mbsf for interstitial

waters. This drilling objective was successfully accomplished, and coring was taken to the maximum allowed depth of 250 m CSF-A with APC, until refusal, followed by HLAPC coring. Recovery was excellent, and triple coring allowed compilation of a complete spliced record extending to ~4 Ma. The paleomagnetic timescale was robust for the uppermost ~100 m CSF-A, below which the signal appears to be modified by diagenesis (Figure F22). For the upper part, the magnetic reversal positions are in good agreement with the biostratigraphy, and these together suggest a sedimentation rate of ~9.3 cm/ky from the mudline to ~200 m CCSF-A and rate of 2.8 cm/ky from ~200 m CSF-A to the bottom of the site (Figure F11). This record will provide the opportunity for high-resolution climate reconstructions of faunal, biogeochemical, and terrigenous tracers that are characteristic for the upper reaches of the Agulhas Current warm-water transports that will allow connections between Agulhas leakage and its headwater variability. The site also holds significant potential to investigate the connections between southern



Figure F14. Summary of planktonic foraminifer preservation from Sites U1474–U1479 through time. Qualitative preservation ratings were tracked through four time intervals: the Pleistocene, late Pliocene, middle Pliocene, and early Pliocene. Foraminifer preservation score is based on ratings of good, moderate, and poor to illustrate changes in preservation through time.



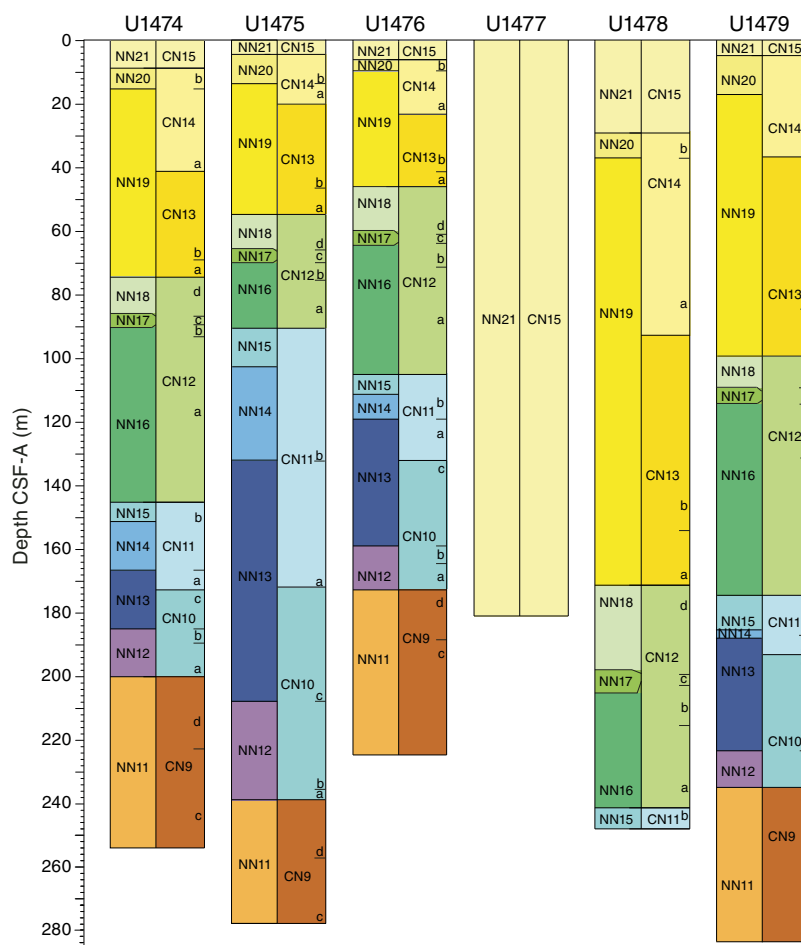
African terrestrial climates and southeast Indian Ocean heat budgets and links to the cultural evolution of early modern humans. A high-resolution program of interstitial water samples collected from a 150 m CSF-A section will provide an additional data point to the presently available interstitial water LGM temperature and salinity inventory.

Operations at Site U1479 (Cape Basin; Figure F3) were scheduled for three APC/XCB holes to 350 mbsf with logging. This drilling objective was largely accomplished with the exception of logging and a reduction in the target section depth to ~300 m CSF-A because of time and sea condition constraints and the biostratigraphy indication that we had achieved the target age. Recovery was excellent, and multiple coring (with extra cores taken because of disturbed intervals related to large ocean swells) allowed compilation of a complete spliced record extending to ~7 Ma. Although the signal is weak and therefore will need future work on shore, the paleomagnetic timescale (Figure F22) is in good agreement with the biostratigraphy, and these together suggest uniform sedimentation rates of ~4.3 cm/ky through the sequence (Figure F11). This record represents the only site situated in the immediate Agulhas leakage pathway. It will therefore provide the opportunity for high-resolution climate reconstructions of the leakage and temporal comparisons with deep-water circulation.

The following specific scientific objectives of the expedition are expected to be readably achievable with the materials recovered:

- To determine the variability of the greater Agulhas Current system under contrasting climatic regimes: early Pliocene warm conditions, during the mid-Pliocene expansion of Northern Hemisphere ice sheets, and through the MPT;
- To establish the role of Agulhas leakage in transient and long-term shifts of the AMOC during these climatic transitions: proxy records of mode/strength of Agulhas leakage during periods of major climatic reorganizations and for AMOC variations through the same intervals;
- To establish the response of the Agulhas Current system to Antarctic climate variability and bipolar linkages: evidence for variation of the Agulhas leakage in highly resolved temporal records will be compared to Antarctic and bipolar climate changes as well as with evidence for shifts of Southern Ocean fronts bordering the Agulhas corridor at the southern tip of Africa, and position and strength of the westerly wind belt;
- To understand the remote tropical origins of Agulhas Current variability: evidence for variation of Agulhas leakage in highly resolved temporal records will be compared to records of Indonesian Throughflow and monsoon modulation on the Agulhas Current source region and impacts on warm-water and saltwater transports along the southeast African margin and Indian-Atlantic saltwater leakage;
- To better determine southern African climate sensitivity: comparison of varying Agulhas Current warm-water transports with rainfall patterns in southern Africa in connection with the Pliocene–Pleistocene evolution of tropical climates, ITCZ migration, and tectonic history of eastern Africa;
- To provide further verification of glacial–interglacial variations in the Southern Ocean’s role in the carbon cycle: nutrient cycling, productivity, and  $p\text{CO}_2$  and their connections with global biogeochemical cycles.

Figure F15. Calcareous nannofossil comparative biostratigraphic chart, Sites U1474–U1479.



### Drilling operations and core handling

Drilling operations and core handling were performed, as is widely acknowledged for IODP, with total professionalism and with outstanding commitment to achieving the expedition scientific objectives. As a result, materials recovered during Expedition 361 were accordingly astonishing and promise significant future scientific research, including full-recovery APC coring to deeper than 300 m CSF-A at Site U1479; complete APC records through the late Miocene/early Pliocene to recent at Sites U1474–U1476, U1478, and U1479; and 382 interstitial water samples from Sites U1474–U1476 and U1478 to help constrain the temperature, salinity, and density structure of the LGM spanning from 488 to 3045 mbsl. At Site U1477, where degassing within the core barrel was pronounced, appropriate core handling adjustments were implemented rapidly, preserving the scientific value of the cored sections. In summary, the exceptional skills of the drillers and core technicians were the main reason why, despite setbacks, the scientific objectives of Expedition 361 could be accomplished.

### Critical aspects

#### Administrative preparation for operations within Mozambique EEZ waters

The majority of the lost operational time was due to delays in attaining the necessary permissions to operate in Mozambique EEZ waters (received some 48 days after the expedition start). These issues should be addressed and resolved prior to an expedition so as not to impact operational time and thereby jeopardize the scientific impact of the recovered material. Although such issues are relatively rare, in this case Expedition 361 had to compensate by dropping all of the planned logging operations and shortening holes (i.e., exclude XCB drilling) to ensure complete APC spliced records to target stratigraphic depths. To this end the paleoceanographic objectives and the interstitial water program, in support of the APL (Agulhas LGM Density Profile), were met. Fortunately, smooth coring operations, sequence characteristics, and a clement sea state for much of the expedition worked in our favor to produce 100% recovery. This will serve to reduce the negative impact of lost operational time.

Figure F16. Planktonic foraminifer comparative biostratigraphic chart, Sites U1474–U1479.

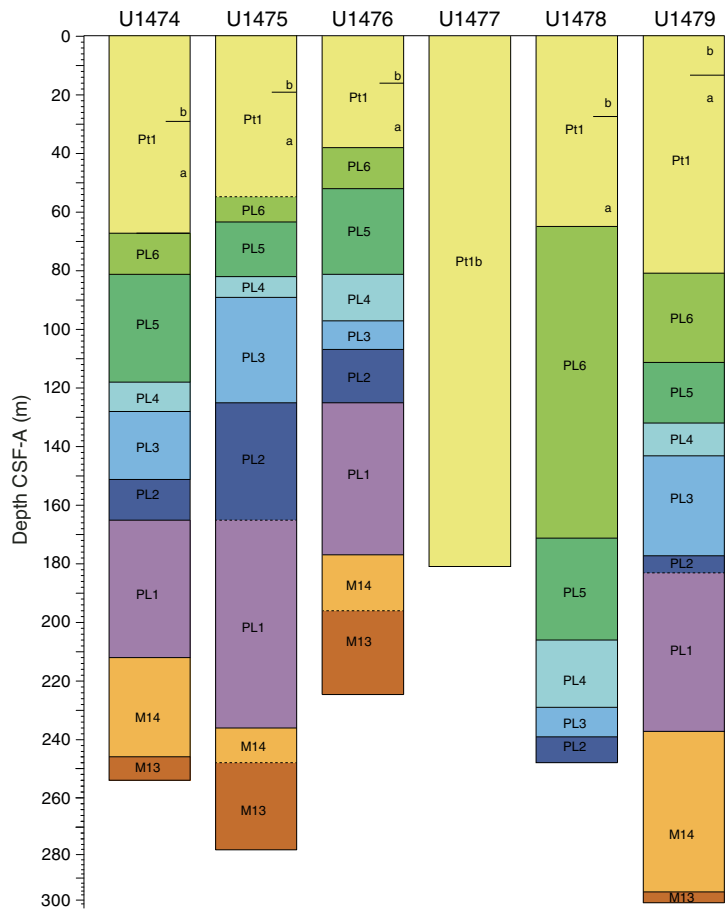


Figure F17. Summary of diatom from Sites U1474–U1479 through time. Qualitative estimated percentages were used to track changes in siliceous microfossil assemblage through time spanning the Holocene to the late Miocene.

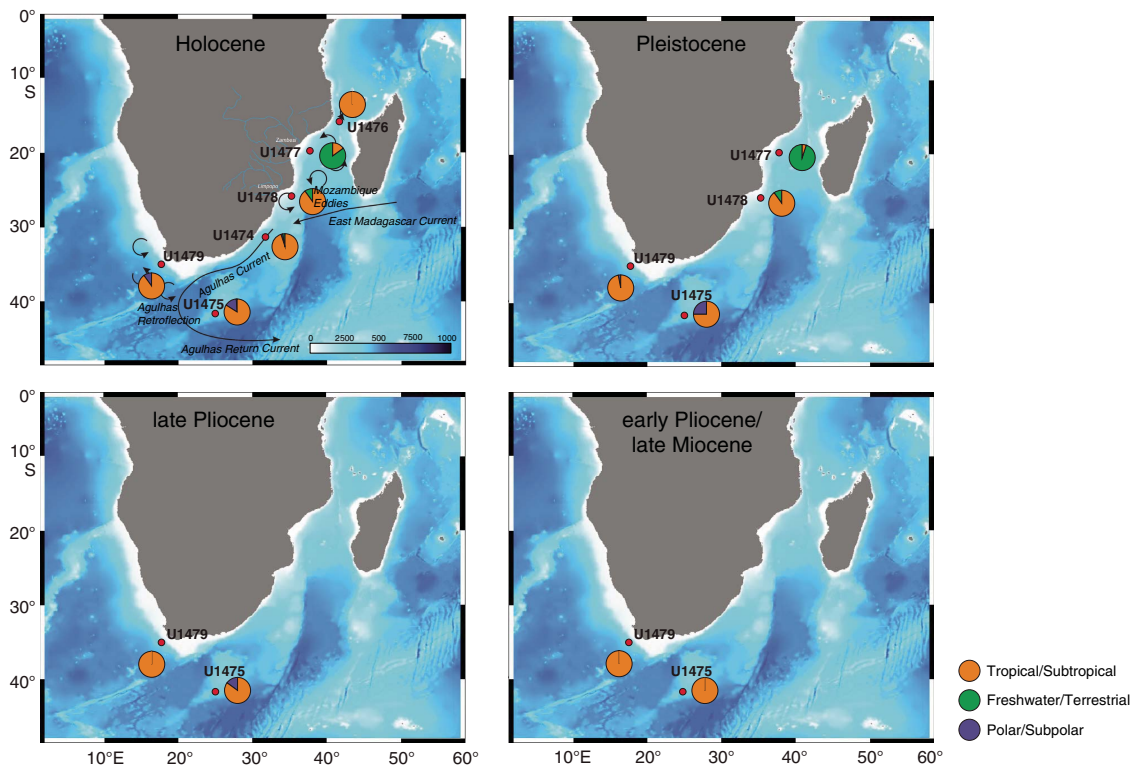




Figure F18. Summary of qualitative estimated percentages of calcareous nannofossil assemblages in sediment recovered from Sites U1474–U1479 spanning the late Miocene to Holocene (mudlines).

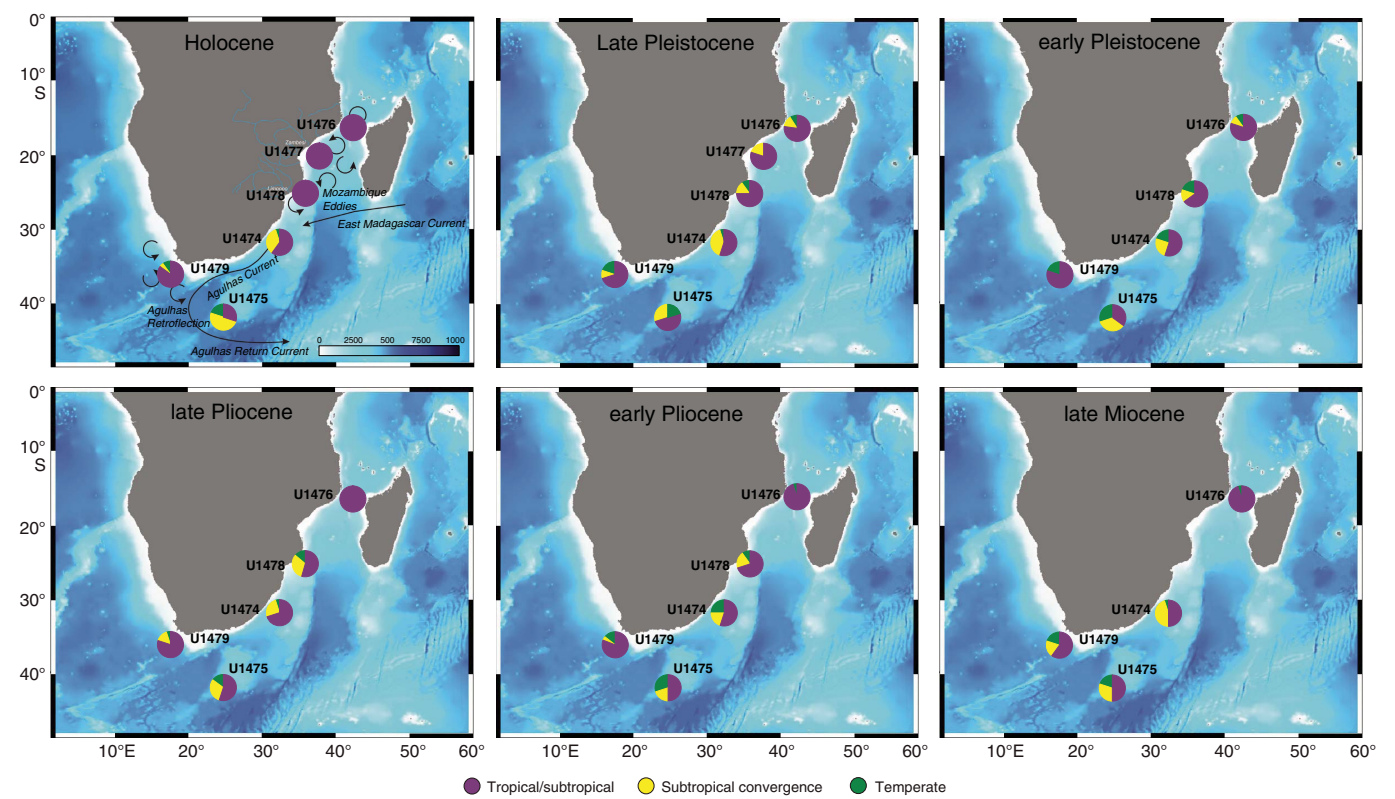


Figure F19. Summary of planktonic foraminifer abundance for Sites U1474–U1479 through time.

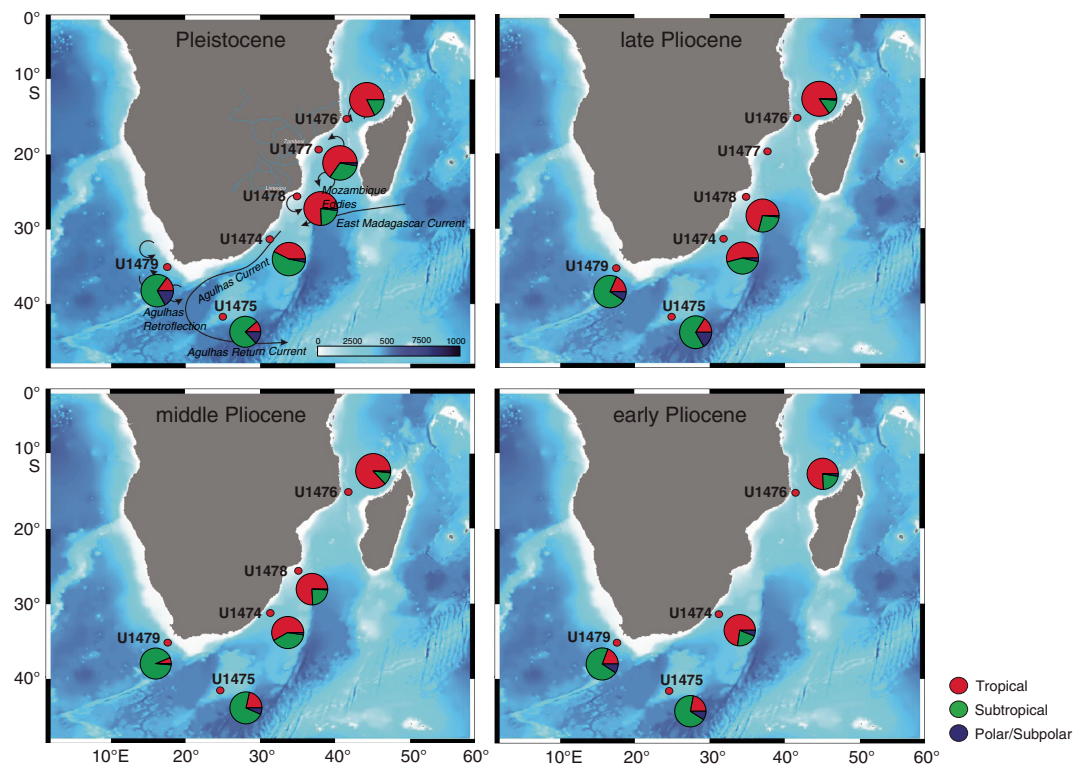


Figure F20. TOC vs. Ba/Al ratio from bulk samples, Sites U1475–U1478.

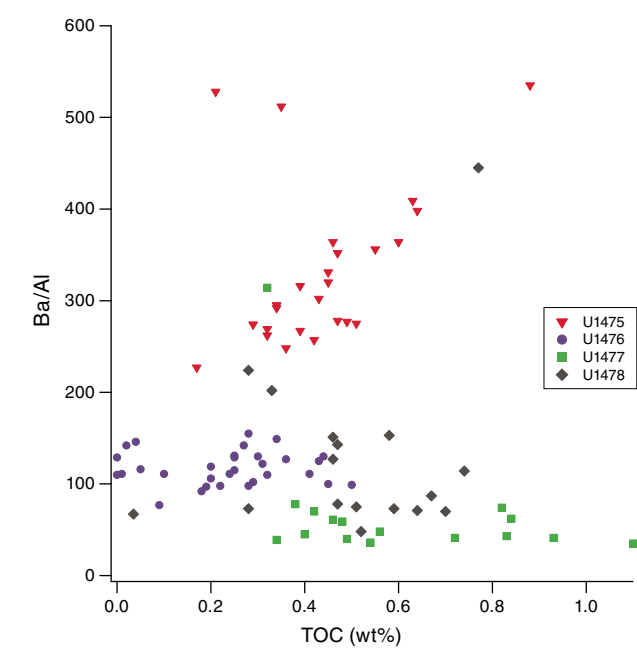


Figure F21. Dissolved sulfate, Sites U1474–U1479.

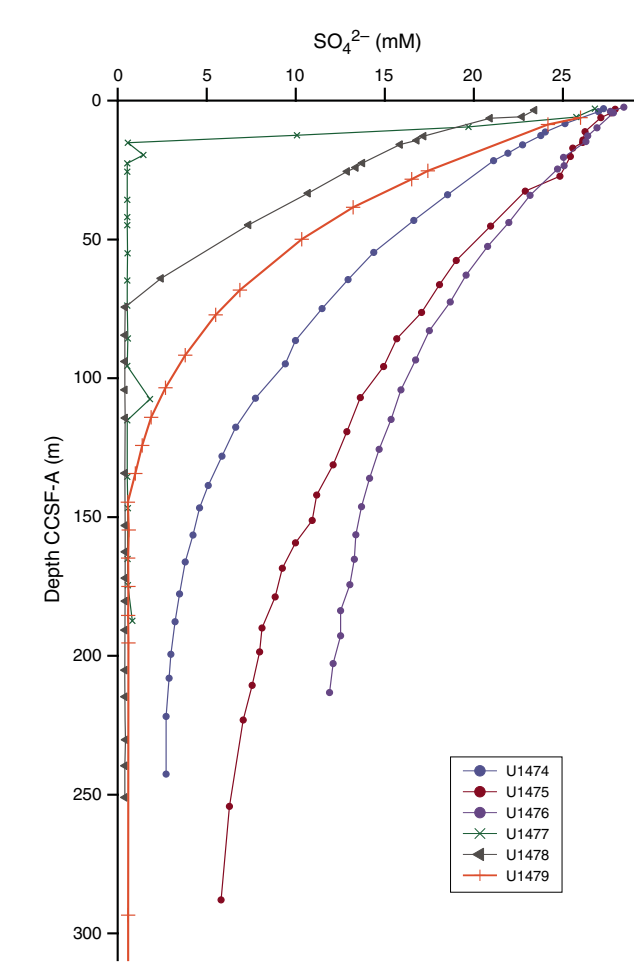
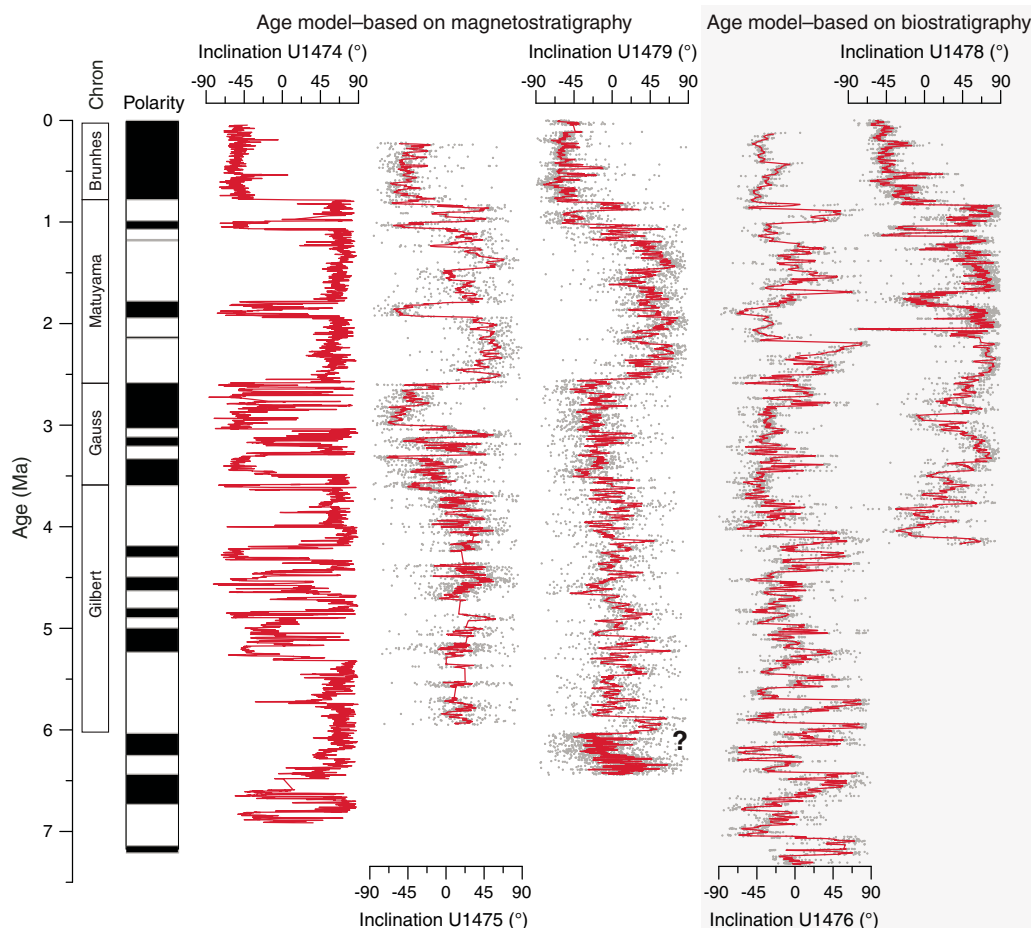


Figure F22. Summary of main shipboard paleomagnetic results (from best paleomagnetic age control [left] to uncertain age-control [right]). Accordingly, the chronologies of Sites U1474, U1475, and U1478 are based on linking the normal and reversed paleomagnetic (sub-)intervals to those of the paleomagnetic stratigraphy of Gradstein et al. (2012). The chronologies of Sites U1476 and U1478 that yield less straightforward paleomagnetic records have been established by utilizing biostratigraphic age-depth tie-points as determined onboard.



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