





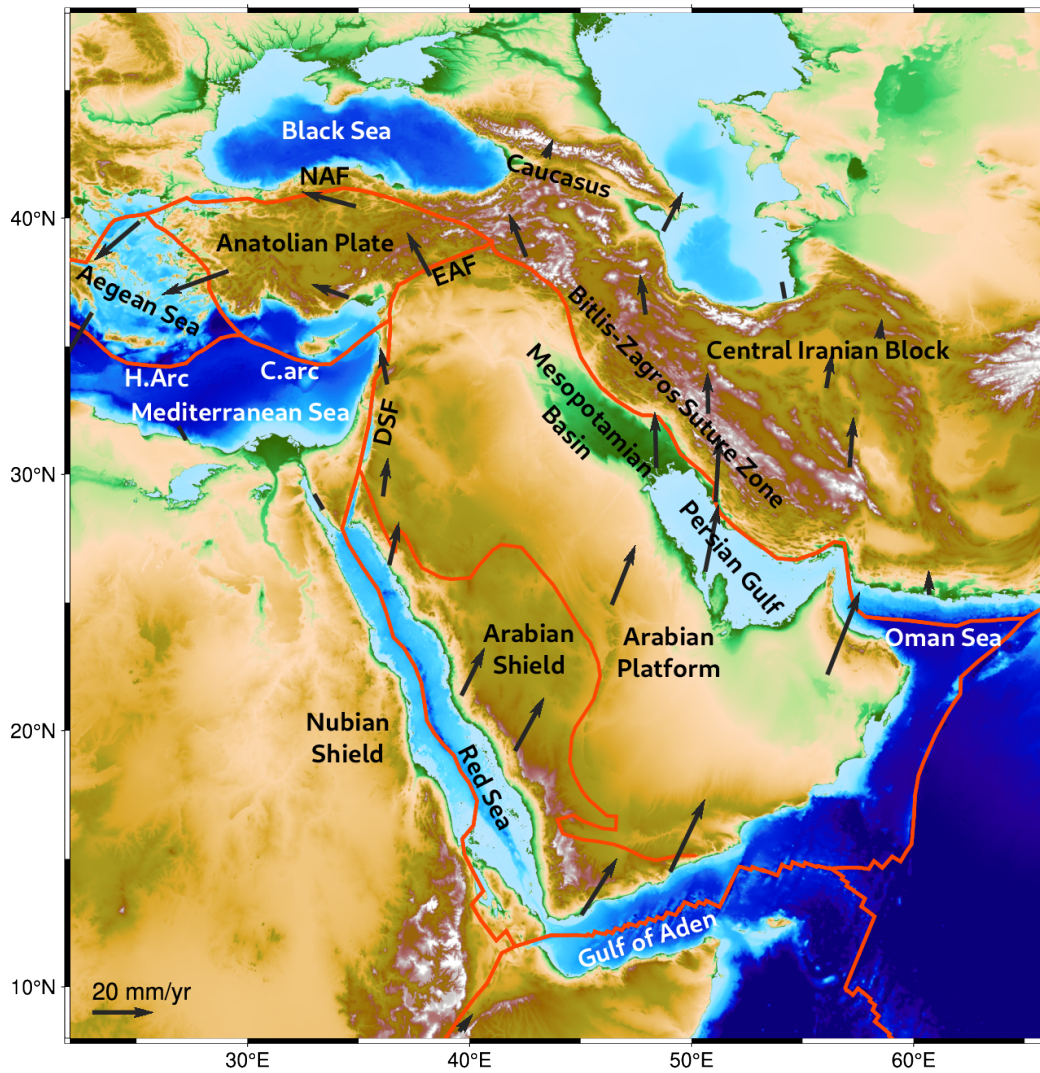


Figure 1 Topographic map of the study region with major tectonic units. NAF: North Anatolian Fault, DSF: Dead Sea Fault, H. Arc: Hellenic Arc, C. Arc: Cyprus Arc. Red lines indicate plate boundaries. A scale bar for GPS vectors is shown on the map.

studies in the region have taken a number of approaches to improve the resolution of seismic tomography. In the last decade, a number of studies have used high-resolution seismic tomography to provide higher resolution images of the crust and upper mantle. These studies have used a variety of methods, including joint inversion of P and S wave speeds, and have provided valuable insights into the tectonic structure of the region. For example, the study by [Author et al., 2010] used joint inversion of P and S wave speeds to provide a high-resolution image of the crust and upper mantle of the Middle East. This study found that the crust is thicker than previously thought, and that there is a significant amount of lateral variation in crustal thickness. Other studies have used tomography to study the structure of the upper mantle. For example, the study by [Author et al., 2012] used tomography to study the structure of the upper mantle of the Middle East. This study found that the upper mantle is hotter than previously thought, and that there is a significant amount of lateral variation in upper mantle temperature. These studies have provided valuable insights into the tectonic structure of the Middle East, and have helped to improve our understanding of the region's geology.



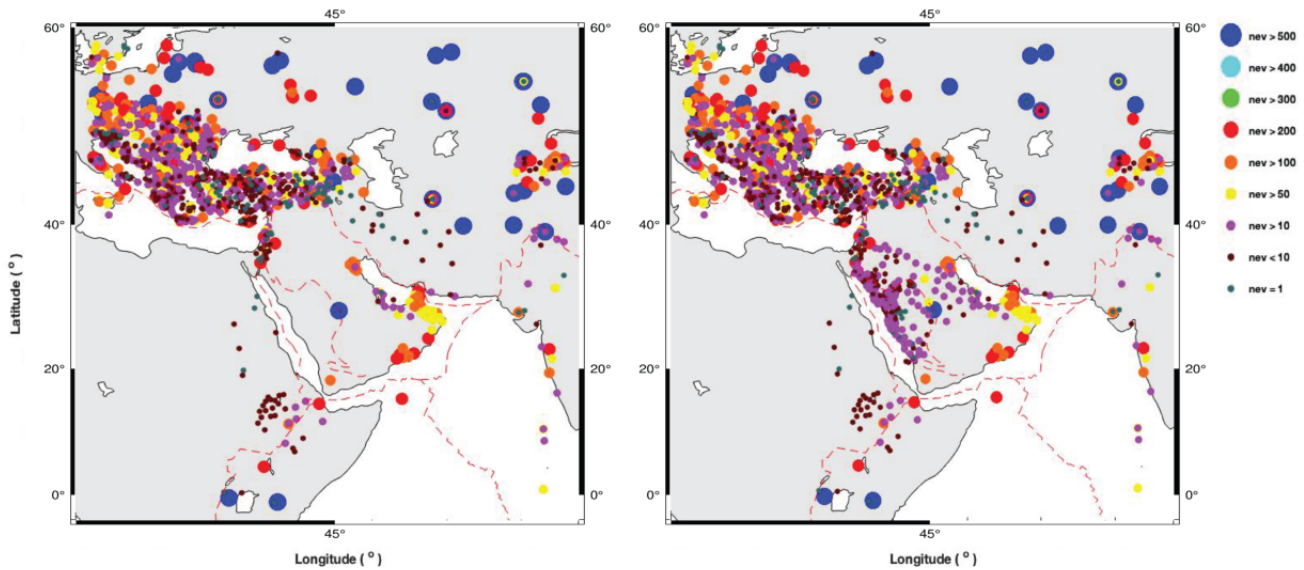


Figure 3: Distribution of P-wave arrival times from the global IASG network. The same as the figure with the addition of the onset-time readings from the SGS network on the Arabian Peninsula.

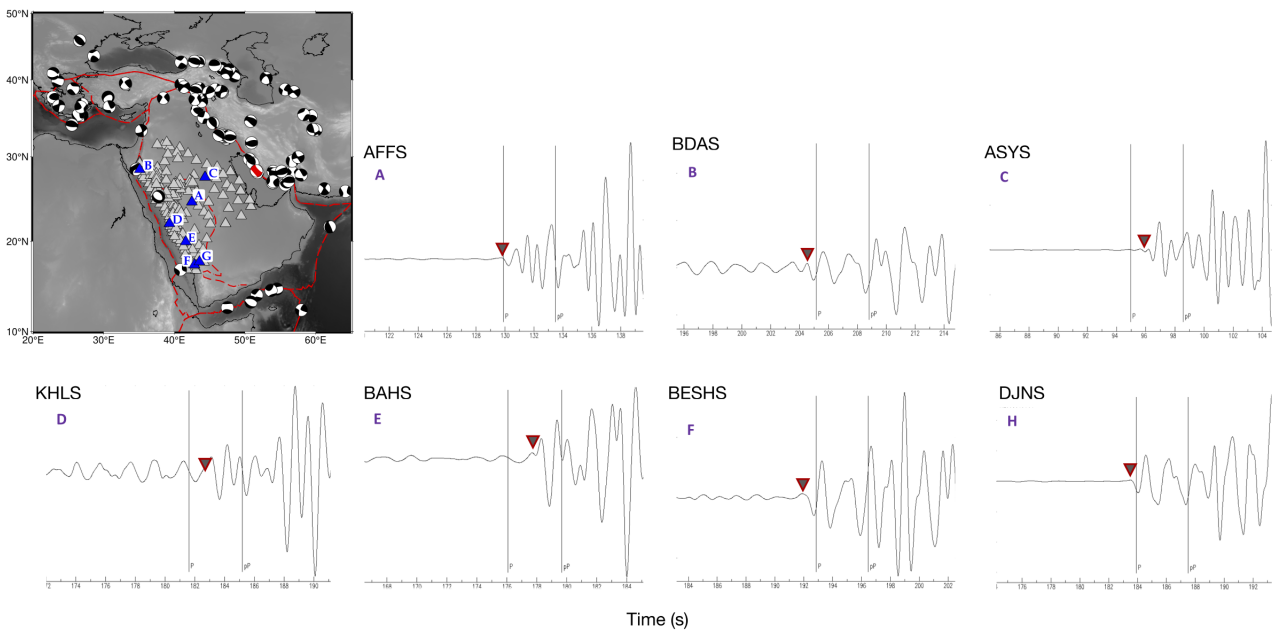


Figure 4: Vertical-component seismograms recorded at sample SGS stations for 2013. The map shows the location of the stations (A through H) and the P-wave arrival times (red beachball). The seismograms show the vertical-component seismograms recorded at the stations. The x-axis is Time (s) and the y-axis is the amplitude. The theoretical predictions of the first-arrival P-waves are shown by the blue triangles. Inverted times. Time delay input for the inversion is then defined as the

sharp onset times distinguishable from background noise. After experimenting with various other filters (i.e., combinations of causal / zero-phase, 2.3 Hz to 10 Hz), the frequency bands tested with the sharp high-pass filter (0.25 Hz to 0.5 Hz) and the onset times (0.25 Hz to 0.5 Hz). Sample vertical-component seismograms from a regional earthquake (M<sub>w</sub> = 6.4, hypocenter at 11.4 km depth) recorded by a set of seismic stations in the Arabian Peninsula with onset waveforms from regional earth

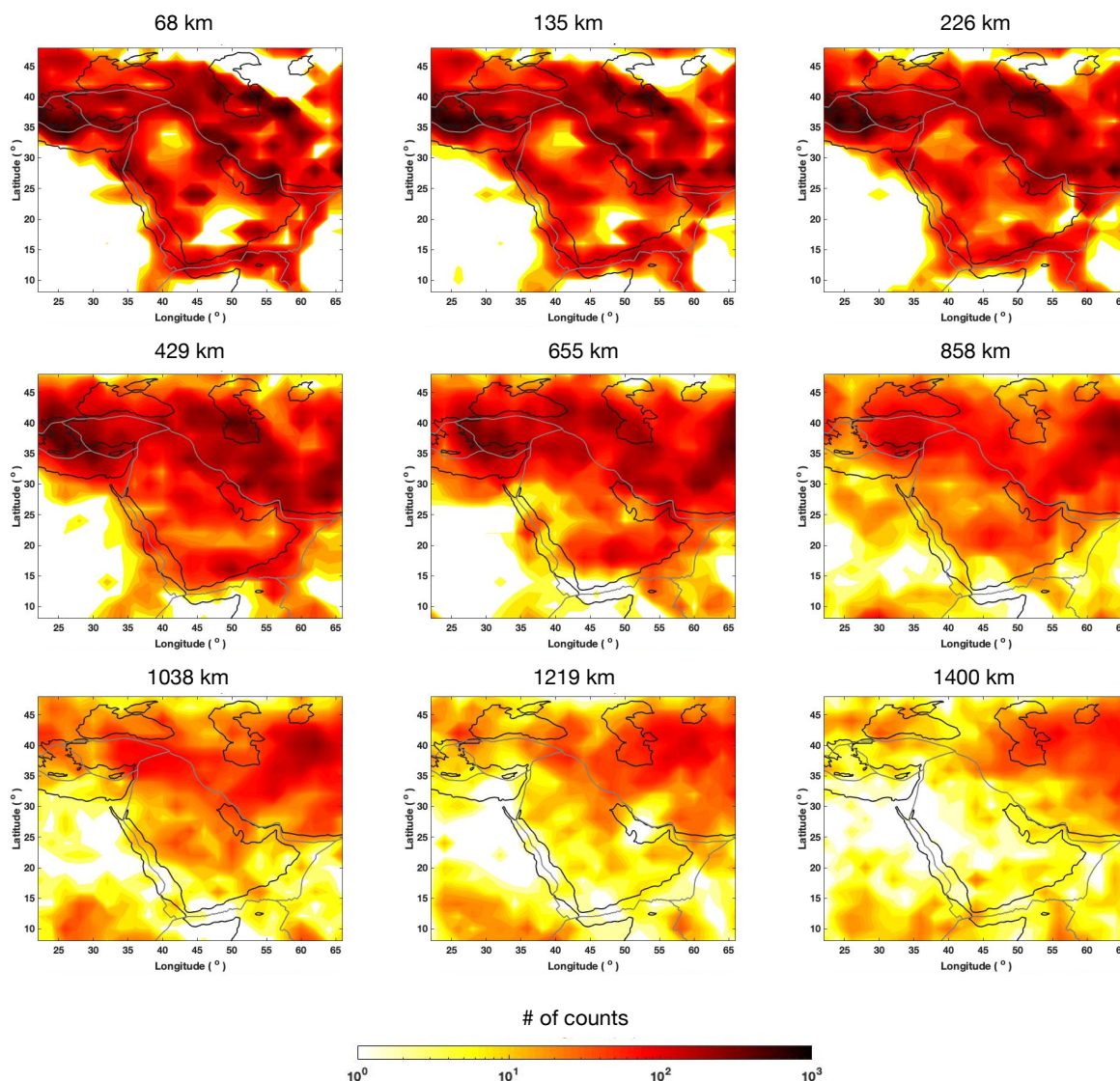


Figure 5. Density of hit counts of ray paths used in the inversion, in the Arabian Peninsula and a last section at various depths.

provement in data coverage is observed, and the new results of the first data sphere down to 400 km underneath the Himalaya belt times series is shown in Figure 10. The new results are obtained by a 1D ray tracing algorithm. The

In this study, we use the travel time inversion software package **EDAS** (Niu et al., 2008) based on ray theory to invert for the compressional wave speed perturbations with respect to the 1D reference model AK135 (Kennett, 1995). Since we use only the first-order readings, ray theory should be valid where the first-order scattering effects can be neglected. We perform a 1D linearized inversion of the travel time data to construct a P-wave model of the study area.  $V_p$  and  $V_p^0$  are the P wave speeds in the background and the perturbation models, respectively.  $\Delta T = V_p - V_p^0$ . For the

inverse problem we aim to obtain model parameters  $m_i = (\frac{\Delta V_p}{V_p^0})_i$  defined on a cubic Earth grid (Am = d, (1) et al. 2013) which has size volume pixels (voxels) on the surface  $\sim 3.6$  m. It is not to be confused with the model for the global from surface to the core-mantle transition, as proposed by Weijay, and ary (CMB). The observed P-wave delay time  $\delta t$  is a function of the partial difference between observed and expected delay time  $\delta t$  and parameters. We rival times by the 1D reference model of Aka13e5, in a region where model



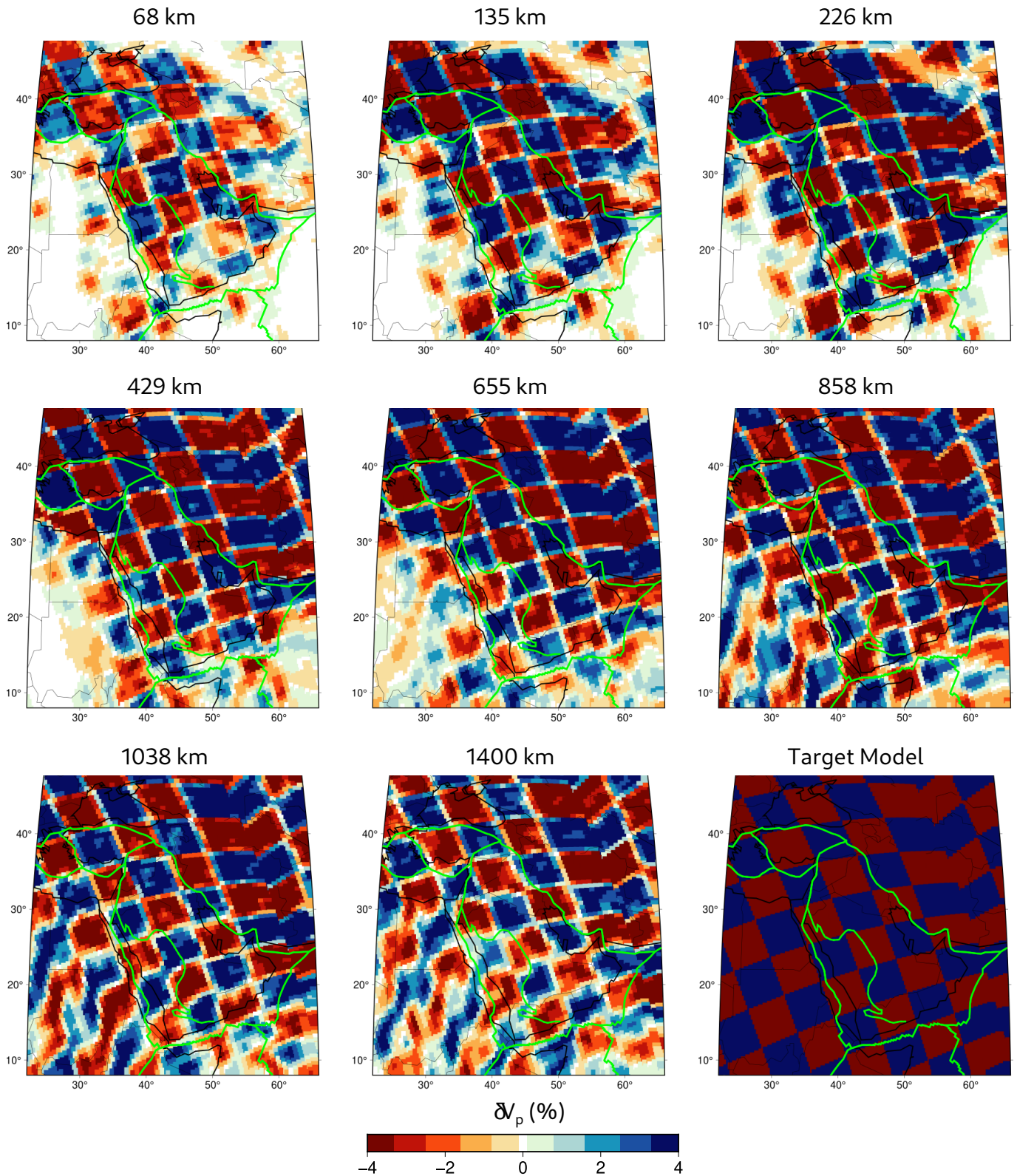
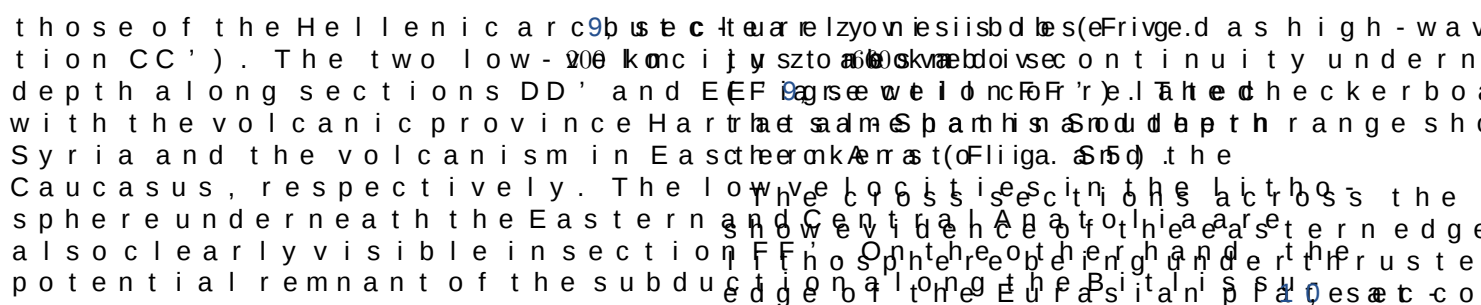


Figure 6: Horizontal cross-sections of P-wave velocity anomalies ( $\delta v_p$  (%)) at various depths. The maps show the Middle East region, with latitude ranging from 10° to 40°N and longitude from 30° to 60°E. A color scale at the bottom indicates the velocity anomaly percentage, ranging from -4% (red) to 4% (blue).

ing from Afar extends northward to the Red Sea (Ciaucias, 2001), and so on, which is well correlated with the global ISCM model. The Red Sea, in the Arabian shield and Eastern Anatolia, is a low-velocity anomaly along the Red Sea, which is shifted to the East toward Harrats around 70 km, which is observed just below the Red Sea around 200 km depth. The results are consistent with previous models (e.g., Engdahl et al., 2007). The resolution on the western part of the shallower Cyprus



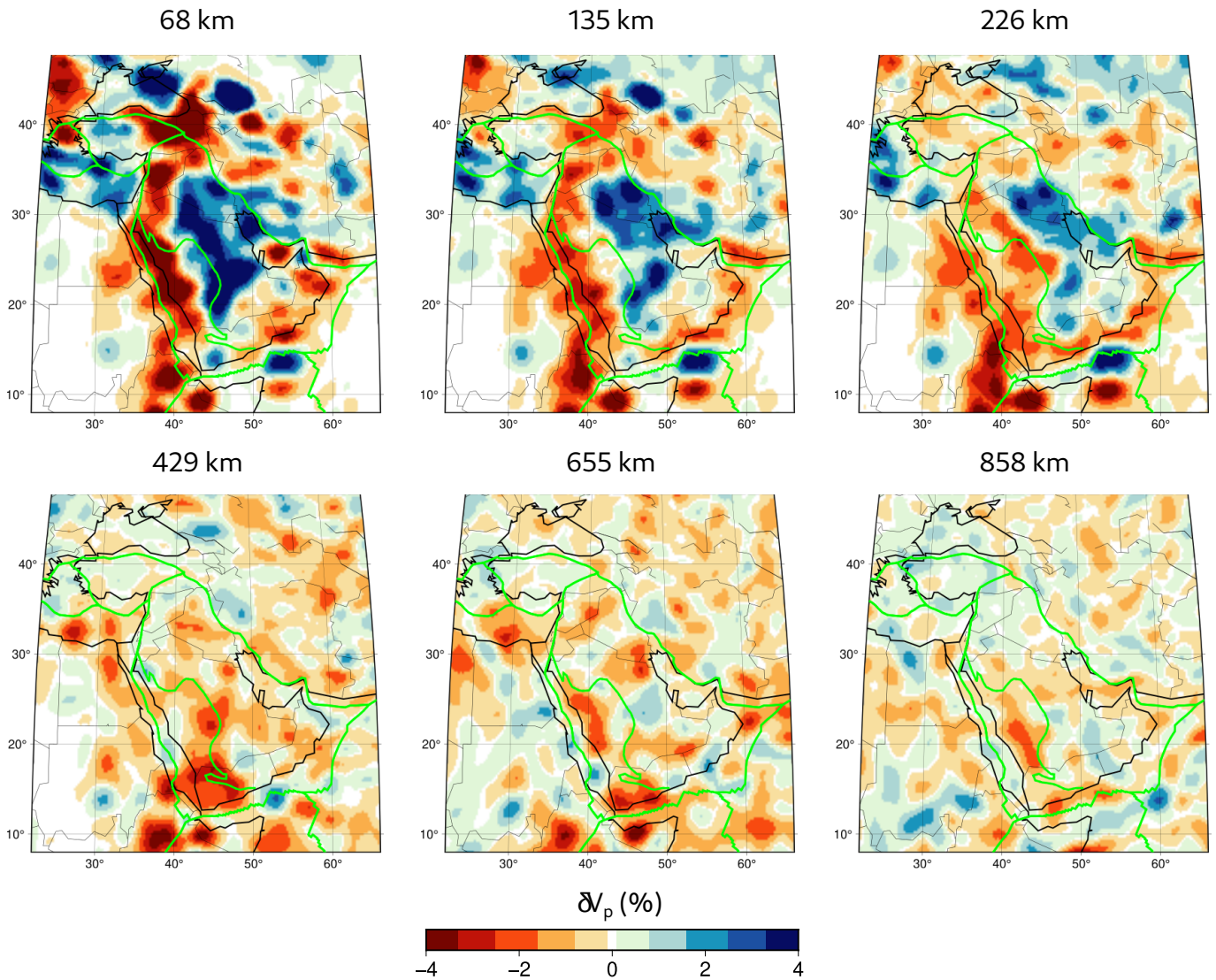


Figure 8. Horizontal cross-sections of P-wave speed perturbations at depths. ISC data combined with onset-time readings from waveform version.

tions AA' - DD'). From sections AA' - DD', the Arabian plate is seen to form the Tethys Ocean underneath Iran along the slow velocity zone in the southern part. This is visible at 300 km depth. The subduction zone in the southern part of the Arabian plate is observed to penetrate the lower mantle-related with the volcanic activity with a steeper angle as we go towards the Red Sea volcanic field. The checkerboard tests, in general, show good resolution of the velocity perturbations along the Zagros suture zone and the East African Rift, where the velocity perturbations to about 600 km depth, support the checkerboard test. However, the resolution decreases at greater depths. The section CC' - DD' (Fig. S6). The western part of the section AA' to DD' highlights the Red Sea and volcanism under the Harrats and Yemen, with low-velocity perturbations in the upper mantle extending to the lower mantle where the decreased checkerboard resolution is observed correlated to the data coverage. The cross-sections along the Red Sea and the Arabian plate, respectively, show similar low-velocity perturbations with respect to the channelled low-velocity zone in the upper mantle. The section GG' - HH' also shows similar low-velocity perturbations in the upper mantle. The section II' - JJ' shows high-velocity perturbations in the upper mantle, where the velocity regions are better resolved.





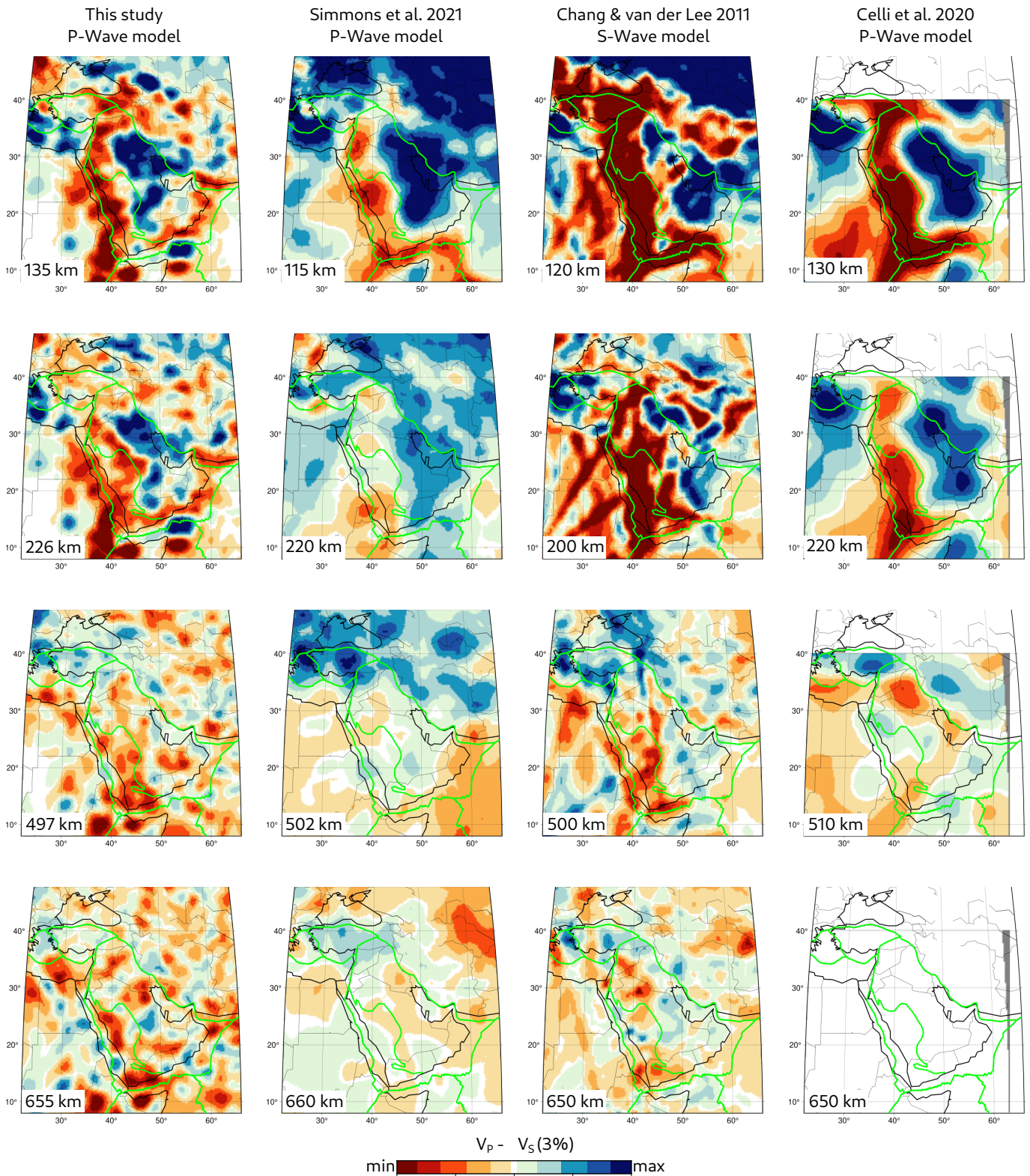


Figure 6. Comparison of the horizontal sections of P-wave and S-wave speed perturbations at various depths (135 km to 655 km) for four different studies: This study, Simmons et al. 2021, Chang & van der Lee 2011, and Celli et al. 2020. The P-wave speed perturbations are plotted with respect to the mean P-wave speed cross-sections plotted with respect to AK135.

The attenuation structure of the region based on wave speed observations and measurement showed that the highest attenuation is observed in the upper mantle underneath the Arabian Plate and the surrounding regions. The attenuation is characterized by a low  $Q$  value (around 0.2) in the platform compared to those

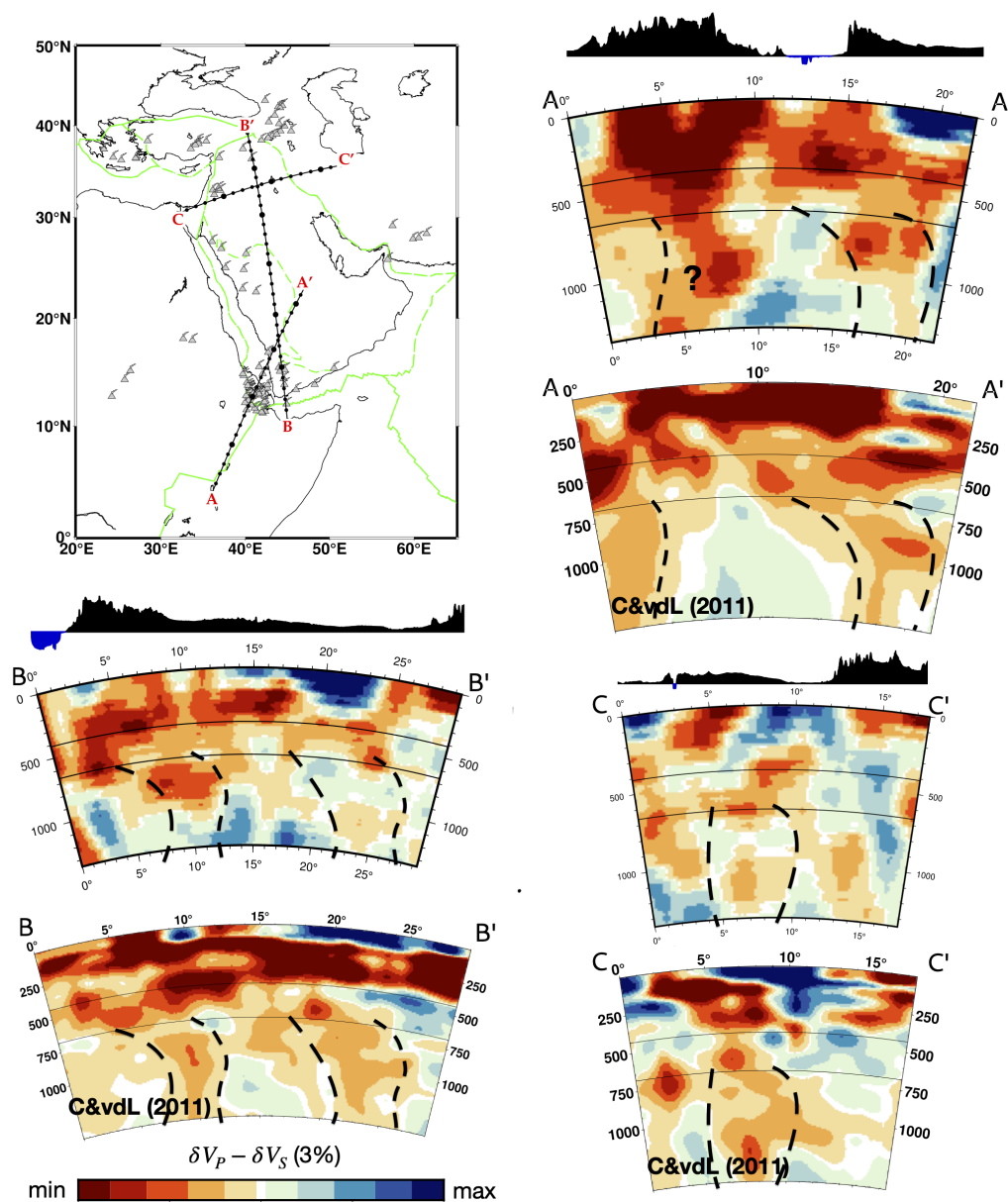


Figure 12. Comparison of the vertical sections of the P-wave arrival-time tomography with the reference model from C&vdL (2011). Perturbations are plotted with a color scale from -3% to 3%.

serve in the upper mantle might be related to the presence of a large-scale low-velocity zone. The slower P wave speeds in the upper mantle are likely related to the volcanic plume models in the Arabian plate (e.g., S&S, 2001; S&S, 2002; S&S, 2003; S&S, 2004; S&S, 2005; S&S, 2006; S&S, 2007; S&S, 2008; S&S, 2009; S&S, 2010; S&S, 2011; S&S, 2012; S&S, 2013; S&S, 2014; S&S, 2015; S&S, 2016; S&S, 2017; S&S, 2018; S&S, 2019; S&S, 2020). Our tomographic model is consistent with the theory of spreading the region. However, the cross-sections suggest the potential of a large-scale low-velocity zone in the upper mantle.

There is no consensus on the origin of the volcanic activity in the region. The Harrats, the largest continental volcanic province in Saudi Arabia extending from the South to Syria, where there are records of historical eruptions close to major cities such as Medina (e.g., Moufi and N&S, 2016). The low velocities along the Red Sea and Gulf of Aden are consistent with the continental rifting and spreading. On the other hand, there is evidence for a channel of flow

lated to a far-off the African superplume (Zhang and Zhao, 2012). In the northern part of the African superplume, the African superplume is also observed in the recent studies (Kühnert et al., 2017). It is discussed that the superplume underneath the African superplume does not support a local plume (Zhang and Zhao, 2012). The African superplume is also observed in the Arabian plate and cannot resolve the African superplume (Zhang and Zhao, 2012). Our P-wave model seems to be different from the previous models of the African superplume in the Arabian plate and Jordan, support the African superplume in the Arabian plate (Zhang and Zhao, 2012). Our P-wave model also see vertical structure of the African superplume in Fig. S17). In addition, the model is consistent with the results of Zou et al. (2017) of the channeled low-velocity zone in the Arabian plate. The potential of the Arabian plate is observed in our model (Fig. S14), which is likely another support for the African superplume model.

[illegible]



- [illegible]

[illegible]

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