

1 Kinetics and oligomer products of the multi-phase
2 reactions of hydroxyacetone with atmospheric amines,
3 ammonium sulfate, and cloud processing

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22 interface, oligomer formation
23

24 ABSTRACT: Hydroxyacetone (HA) is an atmospheric oxidation product of isoprene and
25 other organic precursors that can form brown carbon (BrC). Measured bulk aqueous-phase
26 reaction rates of HA with ammonium sulfate, methylamine, and glycine suggest that these
27 reactions cannot compete with aqueous-phase hydroxyl radical oxidation. In cloud chamber
28 photooxidation experiments with either gaseous or particulate HA in the presence of the same
29 N-containing species, BrC formation was minor, with similar mass absorption coefficients at
30 365 nm (< 0.05 m² g⁻¹). However, rapid changes observed in aerosol volume and gas-phase
31 species concentrations suggest that the lack of BrC was not due to slow reactivity. Filter-based
32 UHPLC/(+)-ESI-HR-QTOFMS analysis revealed that the SOA became heavily oligomerized,
33 with average molecular masses of ~400 amu in all cases. Oligomers contained, on average, 3.9
34 HA, 1.5 ammonia, and 1.6 other small aldehydes, including, in descending order of abundance,
35 acetaldehyde, glycolaldehyde, glyoxal, and methylglyoxal. PTR-ToF-MS confirmed production
36 of these aldehydes. We identify C₁₇H₂₆O₅, C₁₀H₂₂O₉, C₁₅H₂₇NO₇, C₁₇H₂₃NO₅ and C₁₈H₃₂N₂O₉
37 as potential tracer ions for HA oligomers. We hypothesize that efficient oligomerization without
38 substantial BrC production is due to negligible N-heterocycle (e.g., imidazoles/pyrazines)
39 formation. While HA photooxidation is unlikely a significant atmospheric BrC source, it may
40 contribute significantly to aqueous SOA formation.

41

42 **1. Introduction**

43 Hydroxyacetone (HA) is formed by the atmospheric photooxidation of many organic precursors,
44 including isoprene.¹⁻⁵ Its atmospheric concentrations also correlate with CO due to significant
45 production from regional anthropogenic precursor emissions.¹ With a Henry's law coefficient of

46 77 mol m⁻³ Pa⁻¹),⁶ HA is the sixth most common carbonyl species found in cloudwater,⁷ where it
47 can react with aqueous phase oxidants. Due to electron density withdrawal by its α -hydroxy
48 functional group, HA's carbonyl group is more reactive than in monofunctional ketones, resulting
49 in the potential for efficient attack by nucleophiles on the carbonyl carbon. In bulk aqueous-phase
50 experiments involving HA and ammonium salts or primary amines (methylamine or glycine),
51 light-absorbing organic species known as brown carbon (BrC) are slowly formed at levels
52 comparable to aldehydes such as glycolaldehyde (GAld) and acetaldehyde (AAld),^{8,9} are rapidly
53 bleached by OH radicals,⁹ and show a time- and wavelength-dependent photobleaching →
54 browning → bleaching behavior when photolyzed over a few hours.¹⁰ However, chemical
55 reactions in aqueous aerosol particles or evaporating cloud droplets often behave very differently
56 from that observed in bulk liquid samples,^{11,12} and several aldehyde-amine reactions are known to
57 generate surface-active products.^{13,14} Several aqueous aldehyde reactions not only accelerate in
58 the presence of an air-water interface,^{12,15,16} but also exhibit enhanced photobrowning: instead of
59 the rapid photobleaching observed in bulk solution,¹⁷⁻¹⁹ photobrowning has been observed for
60 methylglyoxal- (MG)²⁰ and GAld-amine reactions in suspended aqueous aerosol.²¹ This change
61 indicates that some radical-driven BrC formation pathways occur mainly at the air-water interface.
62 HA has been shown to have a high degree of surface activity at the air-water interface,²² which
63 means that it too may exhibit distinct reactivity in multiphase (as opposed to bulk-phase)
64 experiments.

65 Due to the similarity in brown carbon formation in bulk aqueous samples by HA and aldehyde
66 species, and the great differences between bulk and aerosol reaction behavior noted above for
67 aldehydes, here we investigate pH-dependent HA reaction rates in bulk aqueous solution, and BrC
68 and SOA formation in aqueous aerosol particles during chamber cloud processing of HA. While

69 little BrC was formed in multiphase HA reactions, offline aerosol analysis revealed that the SOA
70 formed was heavily oligomerized.

71
72 **2. Materials and Methods**

73 **2.1 Chemicals and pH**. All chemicals were sourced from Sigma-Aldrich unless otherwise stated.
74 Stock solutions were made by diluting 40% w/w methylamine (MeAm) or liquid HA (95%), or
75 dissolving solid glycine (Gly) or ammonium sulfate (AS) in D₂O (99.9%-D, Cambridge Isotopes)
76 for NMR experiments or in 18 MΩ water (Millipore Milli-Q Gradient A10) for aerosol chamber
77 experiments. The pH of amine or AS samples in D₂O was adjusted downward or upward to reach
78 target levels using acetic acid-*d*6 (Cambridge Isotopes) or sodium phosphate, respectively. The pH
79 was measured before and after mixing with HA.

80 **2.2 Measuring rate constants by NMR**. D₂O solutions containing 0.5M HA and 0.5M of either
81 Gly, AS, or MeAm were pH-adjusted, mixed, placed in NMR tubes, and monitored by ¹H-NMR.
82 To quantify slow reactant losses at pH < 3.5, the NMR signals indicated in Table 1 were followed
83 for up to 62 h, with the integrated peak areas in the spectrum collected nearest *t* = 0 assumed to
84 correspond with starting concentrations of 0.5M. At higher pH levels, reactant losses could be
85 quantified with shorter monitoring periods (~90 min). Based on the behavior of other Maillard
86 reactions at ≤ 1 M concentrations of carbonyl species and amines,^{15, 23-27} reaction orders in this
87 work were assumed to be first order in HA and first order in reduced nitrogen species. Rate
88 constants are given for the ketone form of HA, rather than total HA (ketone + hydrate forms), but
89 this is a small correction, as shown below. Second-order rate constants were derived from initial
90 reaction rates using the following equation:²⁸

91 $Rate_x = k_x f_{HA} [HA]_{tot} [Am]_{tot}$ (1)

92 where $Rate_x$ represents the measured initial loss rate of reactant x in M/s, k_x is the second-order
93 rate constant calculated from the loss of x in $M^{-1} s^{-1}$; $[HA]_{tot}$ and $[Am]_{tot}$ are the total concentrations
94 in M of hydrated and unhydrated HA and protonated and unprotonated forms of the reduced
95 nitrogen species, respectively; and f_{HA} is the equilibrium fraction of HA in ketone (not hydrate)
96 form, determined to be $f_{ald} = 0.96$ using NMR signals.²⁹ (Rate constants calculated using just the
97 total concentrations of both reactant species, without f_{ald} , would thus be 4% higher than those
98 reported here.) Rate constants below $2 \times 10^{-9} M^{-1} s^{-1}$ were too slow to be measured reliably by this
99 method.

100 **Table 1:** NMR Signals Used for Quantitation of Reactant Losses

molecule	functional group	chemical shift (ppm)
hydroxyacetone	CH ₃	2.14
	CH ₂	4.40
methylamine	CH ₃	2.58
glycine	CH ₂	3.55

101
102 2.3 Chamber experiments. The CESAM chamber³⁰ is a 4.2 m³ indoor, temperature-controlled
103 chamber that was operated as a fixed volume batch reactor in this work. Before each experiment,
104 the chamber was pumped down and refilled with an 80:20 mixture of N₂ from a liquid nitrogen
105 tank and UHP oxygen (Linde, $\geq 99.999\%$ O₂, < 0.2 ppm hydrocarbons). Experimental conditions
106 are summarized in Table 2. Pressure was maintained just above ambient levels with automated
107 inward flows of the same 80:20 mixture, compensating for outward sampling flows.
108 Compensation flows were logged throughout each experiment and used for dilution corrections.
109 Seed aerosol were generated with 9 mM AS (Experiments B and E), 9 mM AS / 2 mM glycine
110 (Experiment C), and 5 mM AS / 5 mM HA (Experiment D) using a constant-output atomizer (TSI
111 3076) and were added to the dry chamber through a diffusion dryer (TSI 3062). Gas-phase HA
112 and methylamine (MeAm) were added to the chamber in Experiments A-C and D, respectively,

113 by injecting small volumes of the pure gases into the N₂ flow into the chamber. Water vapor was
 114 added to the chamber in Experiments A-D via short steam injections from a stainless-steel boiler,
 115 after five steam flushes of the boiler headspace. For cloud events, the chamber was brought to
 116 near saturation via steam injection and then pumped at ~40 L min⁻¹ for 10 min, causing cloud
 117 events of 5-10 min duration, before returning to the original chamber pressure. Each cloud event
 118 therefore diluted chamber contents by ~10%. Hydrogen peroxide (HOOH) was added in
 119 Experiments A-D by routing the O₂ inlet flow through a bubbler containing 35% HOOH.

120

121 **Table 2:** Hydroxyacetone Experiment Conditions in the CESAM Chamber

expt.	fig.	steady-state [HA] _(g) (ppb) ^a	seed aerosol type ^b	[MeAm] _(g) (ppb)	[HOOH] _{max} molec/cm ³ ^a	cloud events ^c		notes
						dark	light	
A	S1	500 ± 50	none	0	(6 ± 0.04) × 10 ¹³	0	2	No seeds
B	3	30 ± 3 ^d	AS	0	0.8 × 10 ¹³	1	2	HA(g) + AS(p) seeds
C	4	60 ± 6 ^d	AS / glycine	0	5 × 10 ¹³	1	1	HA(g) + AS/Gly(p)
D	5	40 ± 4 ^e	AS / HA	500 (2×)	3 × 10 ¹³	1	1	MeAm(g) + HA/AS(p)
E	S2	0	AS	0	0	0	0	AS dry blank

122 Notes: **a:** based on PTR-ToF-MS signals at *m/z* 75 (for HA) or *m/z* 33 (for HOOH) in dry
 123 chamber, calibrated by *in-situ* long-path FTIR absorption in range 1030-1220 cm⁻¹ (for HA) or
 124 1228-1258 cm⁻¹ (for HOOH) using reference spectra from the AERIS database.³¹ **b:** atomizer-
 125 generated and diffusion-dried before added to chamber. **c:** Runs A-D included 2-3 cloud events
 126 each, with “photolytic” indicating cloud event(s) for which solar simulator lights were turned on.
 127 **d:** Higher levels were briefly observed upon HA(g) addition under dry conditions, before
 128 humidification. **e:** HA released to the gas phase from aerosol particles. Abbreviations: fig, figure;
 129 [HOOH]_{max}, maximum hydrogen peroxide concentration; (g), gas; (p), particulate.

130

131 Gas-phase species in the chamber were quantified by dedicated CO/CO₂ (ap2e ProCeas), SO₂
 132 (Horiba APSA-370), and NO_x (Horiba APNA-370) monitors, RH and temperature sensors, in situ
 133 long-path Fourier transform infrared spectroscopy (FTIR, Bruker Tensor 37, 182.5 ± 0.5 m path
 134 length), and proton transfer reaction time-of-flight mass spectrometry (PTR-ToF-MS, KORE II,
 135 P_{reactor} = 1.38 mbar, P_{Glow Discharge} = 1.47 mbar, Temp. = 60 °C, E/N = 133 td). The optical properties

136 of chamber aerosol were characterized by cavity attenuated phase shift – single-scattering albedo
137 spectroscopy (CAPS-ssa, Aerodyne, 450 nm) and particle-into-liquid sampling (PILS) with in-line
138 waveguide UV-vis absorbance spectroscopy (Ocean Optics, 1m pathlength, 200-800 nm spectral
139 range) and total organic carbon (TOC) measurements. Aerosol size distributions were monitored
140 by scanning mobility particle sizing (SMPS, TSI model 3080 / 3772, 3 min. scan frequency). Size
141 distributions of droplets and aerosol particles larger than 0.4 μm were characterized during and
142 after cloud events with an Optical Particle Spectrometer using scattered white light (PALAS welas
143 Digital 2000, 0.4 to 15 μm range). Gas phase species signals were corrected for dilution using
144 recorded gas flows into the chamber, and SMPS measurements were corrected for both dilution
145 and size-dependent wall losses measured in an earlier control experiment.

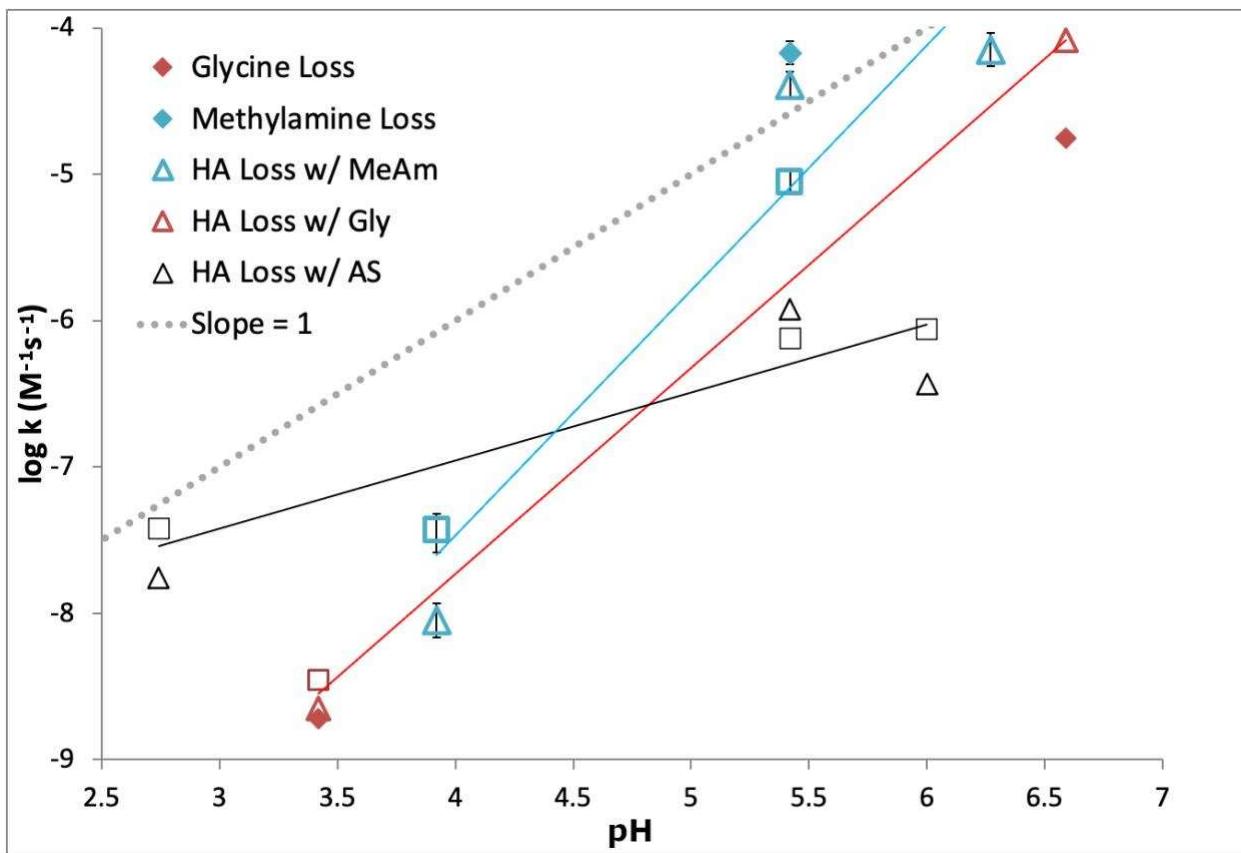
146 Time-dependent single-scattering albedo at 450 nm was calculated by dividing CAPS scattering
147 by CAPS extinction signals after correction using a daily calibration with polydisperse AS aerosol
148 diluted to achieve extinction signals between 0 and 1200 Mm^{-1} . Mass absorption coefficients
149 (MAC_{365}) for the water-soluble aerosol fraction were calculated from waveguide absorbance at
150 365 nm and TOC measurements taken in series on the outflow of the PILS sampler using the
151 equation $\text{MAC} = 2.303A / bC$, where A is the waveguide \log_{10} absorbance at a given wavelength,
152 b is optical pathlength of the capillary (1000 cm), and C is the TOC level measured simultaneously
153 in the PILS outflow ($\text{g organic carbon cm}^{-3}$). The clean, air-filled chamber was used as a daily
154 waveguide blank reference spectrum, while the TOC instrument was baselined with 18 $\text{M}\Omega$ water.

155 Offline chemical analysis of chamber aerosol was performed after the end of three experiments
156 by collecting Teflon filters (1.0- μm pore size, 47 mm diam., Tisch Sci.) overnight at 16 L min^{-1}
157 and storing them at -20°C until extraction in methanol (Optima LC/MS Grade, Fisher Sci.) by
158 sonication for 45 min. Extracts were dried under high-purity nitrogen gas, dissolved in 150 μL of

159 50:50 v/v methanol and Milli-Q water, and then analyzed by ultra-high performance liquid
160 chromatography electrospray ionization high-resolution quadrupole time-of-flight mass
161 spectrometry in positive mode (UHPLC/(+)ESI-HR-QTOFMS), as previously described in
162 detail.³² Aliquots of 5 – 10 μ L were injected onto a Waters ACQUITY UPLC HSS T3 column
163 (2.1 \times 100 mm, 1.8- μ m particle size) and eluted at 0.3 mL min⁻¹ with methanol and water solvent
164 mixtures containing 0.1% ammonium acetate (LC-MS Chromasolv-grade, Sigma-Aldrich). Data
165 were analyzed using Agilent MassHunter (Version B.06.00 Build 6.0.633.0).

166
167 **3. Results and Discussion**

168 3.1 Hydroxyacetone reaction kinetics. Reaction rate constants measured in bulk D₂O by ¹H-
169 NMR are summarized as a function of pH in Figure 1. Rate constants derived from HA loss rates
170 during reactions with AS, glycine, and MeAm are shown as open symbols and are fit with least-
171 squares lines. All three HA reactions are seen to be pH dependent over the range 2.5 to 7.
172 However, none of the reaction rate constants increase on this log(k) vs. pH plot with a slope of 1,
173 which would be expected for a reaction involving a deprotonated nitrogen-containing weak base
174 controlled only by its acid-base equilibrium. Instead, the slope was significantly less than 1 for
175 HA + AS and significantly greater than 1 for HA + MeAm or glycine reactions. In three
176 experiments where loss rates of both HA and amine were quantifiable, rate constants calculated
177 from the loss rates of the two species at a given pH were similar (i.e., the mean k_{HA} / k_{amine} ratio,
178 2.2 \pm 1.3, was not statistically distinguishable from 1). This suggests that 1:1 reactions between
179 HA and reduced nitrogen species were the norm, as expected since HA has only one carbonyl
180 moiety susceptible to nucleophilic attack by a lone pair on a nitrogen atom.



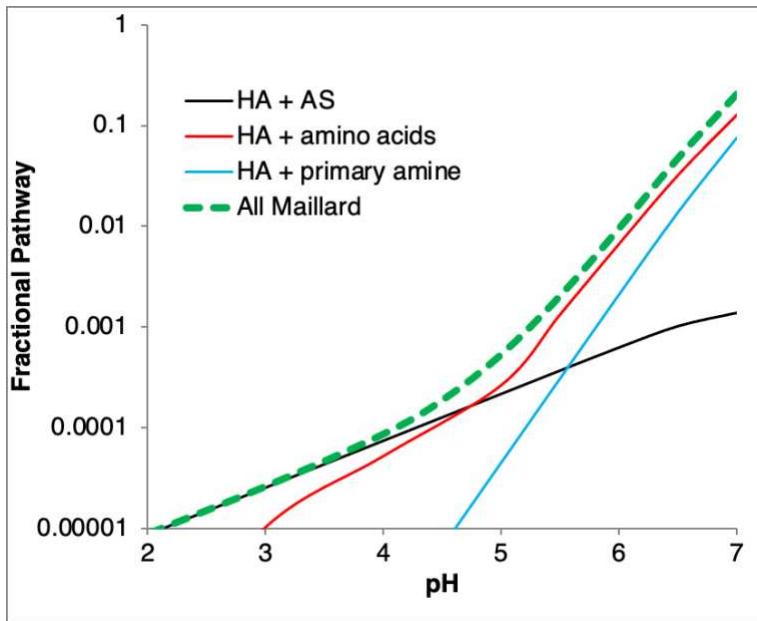
181 **Figure 1.** Measured room-temperature 2nd-order rate constants ($M^{-1} s^{-1}$) as a function of pH,
182 measured by ¹H-NMR for hydroxyacetone reactions with glycine (Gly, red), methylamine
183 (MeAm, blue), and ammonium sulfate (AS, black), initial concentrations of all reactants = 0.5 M.
184 Rate constants were calculated from initial loss rates of reactant signals, and the $\pm 1\sigma$ error bars
185 derived from uncertainties in fitted slopes shown for the methylamine reaction dataset are typical
186 for all three datasets. HA losses were followed at the CH₃ signal (2.14 ppm, open triangles) and
187 the CH₂ signal (attached to the OH group, 4.4 ppm, open squares). HA losses with each reduced
188 nitrogen species were fit with linear functions. Amine loss rate constants (filled diamonds) are
189 shown for methylamine CH₃ (2.58 ppm, blue) and glycine CH₂ (3.55 ppm, red signals) when such
190 losses were observable. AS losses could not be quantified by ¹H-NMR.

192 Measured rate constants from bulk aqueous measurements were used to estimate HA reaction
193 rates in atmospheric aqueous aerosol as a function of pH (Figure 2), assuming that all amino acids
194 react at the same rate as glycine, and all other primary amines react at the rate of methylamine. If
195 rate constants measured in bulk aqueous mixtures are applicable to atmospheric aerosol, these
196 Maillard reactions are too slow to compete with oxidation by hydroxyl radicals ($\cdot\text{OH}$) in the
197 aqueous phase at any acidic pH. The estimated summed fractional pathway for all Maillard
198 reactions involving HA was only 0.2% at pH 5.5 and rapidly decreased at lower pH. However,
199 several questions must be addressed before accepting this conclusion. First, do these reactions
200 accelerate when taking place in aerosol particles, relative to bulk liquid solutions, as has been
201 observed for analogous reactions with aldehydes?^{12, 15, 16} Second, is BrC formation significant,
202 such that this reaction, even if a minor pathway, might influence aerosol optical properties under
203 atmospheric conditions? Third, are there any multi-phase, radical-driven mechanisms by which
204 HA and amines can quickly form Maillard-like products (N-containing BrC oligomers), as
205 observed for C₂ and C₃ bifunctional aldehydes?^{20, 21} In order to address these questions we now
206 turn to chamber studies.

207

208 3.2 CESAM chamber experiments

209 As summarized in Table 2, gas-phase HA was added in Experiments A-C, while only aerosol-
210 phase HA was added in Experiment D. Experiments A and E (Figures S1 and S2) were controls,
211 performed without seed aerosol and without HA, respectively. High levels (500 ppb) of HA were
212 utilized in Experiment A in order to allow quantitation by long-path FTIR in the dry chamber. HA
213 levels were an order of magnitude lower in the remaining experiments.



214

215 **Figure 2.** Fractional pathway of HA reactivity via Maillard reactions estimated in global aerosol
 216 as a function of pH, based on extrapolation of bulk aqueous rate constants to aqueous aerosol
 217 where $[AS] = 0.081$ M (black), amino acid concentration = 0.068 M (red), primary amine
 218 concentration = 0.0034 M (blue), and total Maillard reaction fraction (green dashed line). Reaction
 219 of HA with OH radicals in the aqueous phase (with $[^{\bullet}OH] = 1 \times 10^{-13}$ M)³³ makes up almost all of
 220 the remainder of aerosol-phase reactivity.

221

222 SOA and BrC production in each experiment are summarized in Table 3. In control Experiment
 223 A where HA was added to an air-filled, seed-free chamber with freshly-cleaned walls, rapid loss
 224 of HA to the steel walls was observed, with a lifetime of 250 s (Figure S1 panel *c*). This loss rate
 225 was not increased when the chamber's solar simulator lights were turned on, as expected because
 226 HA gas does not absorb light in the actinic region. However, when the chamber was humidified
 227 to 55% RH, HA was released from the walls, and achieved a steady-state equilibrium: PTR-ToF-
 228 MS signals at *m/z* 75 rose within a few minutes to reach a constant level that was 80% higher than
 229 the initial peak when HA was first added to the dry chamber. This steady state was maintained for

230 over 30 minutes, until the equilibrium was altered by further RH increases in the chamber. This
231 behavior of uptake to dry walls and then equilibrating from the walls once a chamber is humidified
232 has been observed for some small, oxygenated molecules with carbonyl or hydroxyl groups, such
233 as the dicarbonyl compound glyoxal (GX)³⁴ or phenolic species like guaiacol.³⁵ As the RH
234 increased from 50% to ~100%, gas-phase HA concentrations decreased by a factor of 2.3 and
235 appeared to be temporarily depressed by each water vapor supersaturation period, which in the
236 absence of seed aerosol is likely due to the effect of reversible changes in the amount of wall-
237 adsorbed water on the HA gas / wall equilibrium. In the absence of seed particles, no cloud
238 droplets formed upon water vapor supersaturation in Experiment A (Figure S1 panel *b*), unlike in
239 all later experiments, which contained seed particles. At $t = 16:20$, HA(g) reached a new steady
240 state, which was maintained for an hour even after the addition of HOOH and OH radicals to the
241 gas phase. Thus, although HA and OH radicals are known to react in the gas phase with a rate
242 constant $k = (6.0 \pm 0.9) \times 10^{-12} \text{ cm}^3 \text{ mole}^{-1} \text{ s}^{-1}$,³⁶⁻³⁸ we can attribute the nearly constant HA(g)
243 concentrations to continued resupply from chamber walls of HA(g) adsorbed earlier in the
244 experiment. In later experiments, where HA(g) concentrations were an order of magnitude lower,
245 this wall source of HA is expected to be smaller. In the atmosphere, HA is produced by the
246 photooxidation of precursor gases, so a continuous daytime source is present.

247

248 **Table 3:** Summary of Experimental Results in the CESAM Chamber

expt.	fig.	Notes	SOA production ^a ($\mu\text{g}/\text{m}^3$)	$\Delta\text{MAC}_{365}^{\text{b}}$ (cm^2/gOC)	min. albedo, ^c 450 nm	ΔTOC (ppb)
A	S1	No seeds	< 1	0 \pm 200	n/a	7 \pm 10
B	3	HA(g) + AS(p) seeds	17 after dark cloud	500	0.98	60
C	4	HA(g) + AS/Gly(p)	9 after dark cloud	100	0.97	40
D	5	MeAm(g) + HA/AS(p)	19 and 12 after MeAm, 7 after dark cloud	400	0.97	70
E	S2	AS dry blank	0 \pm 3	no data	no data	2 \pm 2

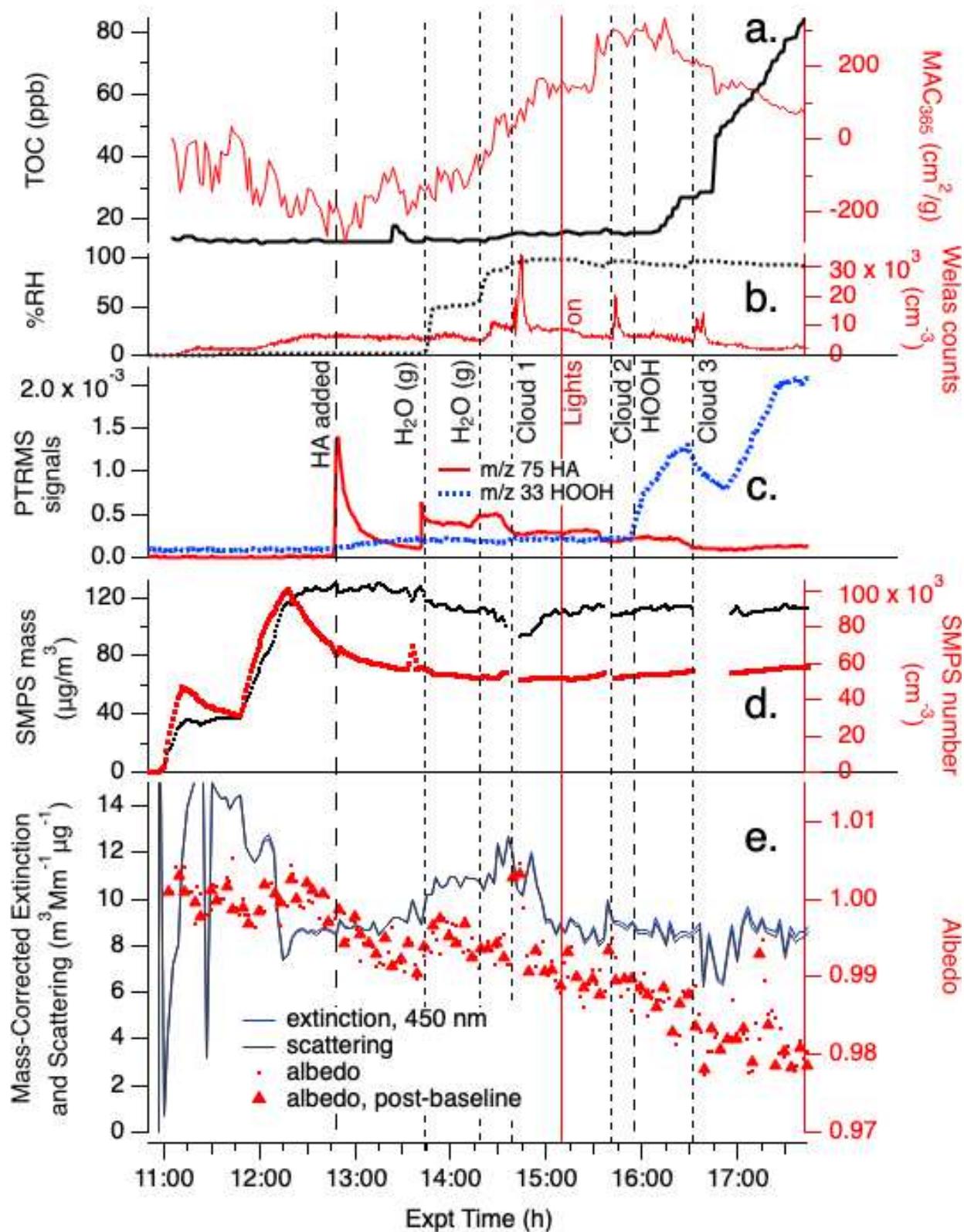
249 Notes: Uncertainties listed for Experiment A are typical for all experiments unless otherwise
 250 noted. **a:** measured by SMPS, assuming aerosol density of 1.0. All SOA production was observed
 251 only after prior equal or greater losses of aerosol mass. **b:** calculated from PILS-waveguide
 252 absorbance and in-line TOC measurements. **c:** measured by CAPS-ssa. Abbreviations: fig, figure;
 253 ΔMAC_{365} , increase in mass absorption coefficient at 365 nm; min., minimum; ΔTOC , increase in
 254 total organic carbon; (g), gas; (p), particulate.

255

256 After the addition of HOOH / OH radicals at $t = 16:30$, a few gas-phase oxidation products were
 257 observed by PTR-ToF-MS, notably $\text{C}_4\text{H}_6\text{O}_3$ at m/z 103. Some water-soluble oxidation products
 258 were taken up by chamber background aerosol particles ($\sim 100 \text{ cm}^{-3}$, mostly introduced with water
 259 vapor) and increased TOC levels by 10 ppb (Figure S1 panel *a*). SMPS measurements indicate
 260 that these particles had a dry geometric mean diameter of $\sim 30 \text{ nm}$ and totaled less than $1 \text{ }\mu\text{g}/\text{m}^3$
 261 after an hour of photooxidation from either HA or gas-phase chamber contaminant precursors
 262 (Figure S1 panel *d*). No enhancement in aerosol scattering, extinction, or absorbance was observed
 263 for these particles, as expected due to their small size and low numbers.

264 In Experiment B, HA(g) was added to a chamber containing AS seed particles. Rapid loss of
 265 HA from the gas phase was again observed by PTR-ToF-MS (Figure 3 panel *c*), with a slightly
 266 longer lifetime (400 s) than in Experiment A. If uptake to dry AS seeds was significant, we would
 267 expect instead a shorter lifetime for HA(g) loss than in Experiment A due to the additional surface
 268 sink, along with increases in TOC and SMPS mass, none of which were observed (Figure 3 panels
 269 *a* and *d*). However, after HA(g) was added to the chamber, the aerosol single-scattering albedo at

270 450 nm (Figure 3 panel *e*) began to decline and the MAC₃₆₅ value calculated from aerosol
271 absorbance rose slightly (Figure 3 panel *a*), both of which suggest that a small amount of BrC may
272 have formed in dark reactions between HA and AS at the dry aerosol surface.



2
274

275

Figure 3. Expt. B with gas-phase hydroxy-acetone, photolysis, cloud processing, and HOOH oxidation with AS seed particles in the CESAM chamber. HA addition, two water vapor additions,

276 start of chamber illumination, start of continuous HOOH addition, and 3 cloud events are labeled.
277 Panel **a**: total organic carbon readings and mass absorption coefficients at 365 nm from
278 PILS/waveguide UV-vis, color coded to axes. **b**: chamber RH and droplet spectrometer counts,
279 color coded to axes. **c**: water-corrected PTR-ToF-MS data (*m/z* 75 HA, red line; *m/z* 33 HOOH,
280 blue line). **d**: Dilution- and wall-loss-corrected SMPS aerosol total mass (assuming density = 1
281 g/cm³) and counts, color-coded to axes. **e**: CAPS-ssa data at 450 nm (mass-corrected extinction,
282 blue line; mass-corrected scattering black line; single-scattering albedo, red dots; albedo measured
283 immediately after instrument autozero to remove contributions from gas-phase absorbers, red
284 triangles.

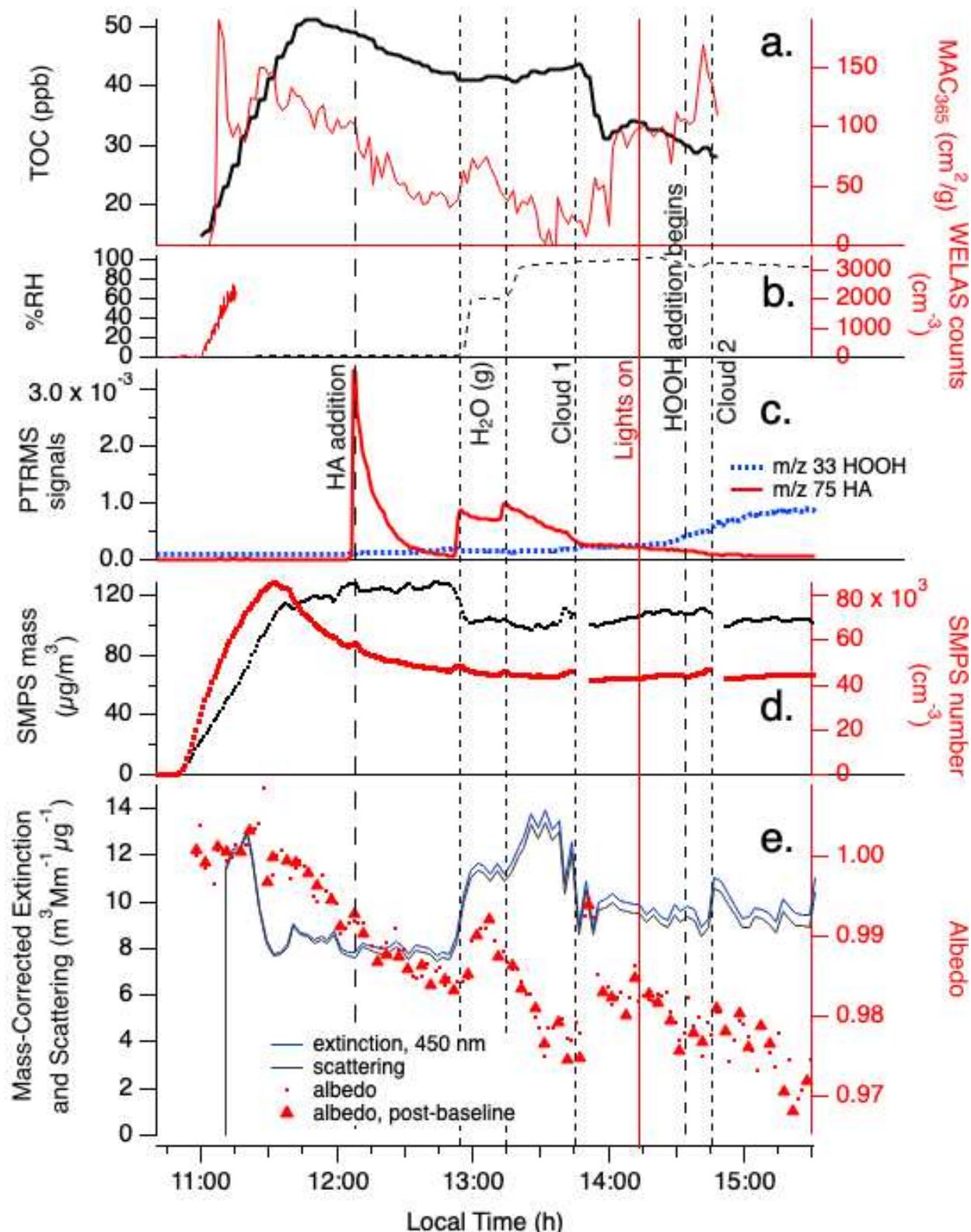
285 When the chamber was humidified to 50% RH at *t* = 13:45 in Experiment B, HA(g) re-
286 equilibrated from chamber surfaces (now potentially both walls and aerosol particles, Figure 3
287 panel *c*), but then quickly declined with a lifetime of < 1 min to concentrations that were $\sim 15\times$
288 lower than in Experiment A. At the same time, aerosol albedo and mass-normalized aerosol
289 scattering (Figure 3 panel *e*) both increased as SMPS aerodynamic dry aerosol mass dropped
290 slightly (Figure 3 panel *d*). These observations together suggest that the uptake of some HA (and
291 possibly water) restructured the AS seed aerosol particle surfaces, making the aerosol more
292 spherical and aerodynamic,³⁹ and hydrolyzing a little of the BrC formed at the aerosol surface
293 under dry conditions. Further humidification to 86% RH at *t* = 14:20 deliquesced the AS seed
294 aerosol, at which point a release of 10 ppb HA to the gas phase, an accelerated upward trend in
295 MAC₃₆₅, a slight increase in TOC, and further SMPS aerosol dry mass loss of 10.7 $\mu\text{g m}^{-3}$ were
296 seen in the data. This aerosol dry mass loss, if due entirely to HA evaporation, could account for
297 about a third of the increase in gas phase HA. Furthermore, in the absence of seed particles in
298 Experiment A, HA(g) declined sharply when the RH was increased from 50 to 85%. The HA(g)

299 increase in Experiment B therefore indicates that deliquescing AS seed aerosol particles were an
300 important source of the HA released to the gas phase at high RH, causing their SMPS-measured
301 dry mass to decline. At the same time, very small increases in MAC_{365} by $110\text{ cm}^2/\text{g}$ and TOC by
302 1.5 ppb (Figure 3 panel *a*) suggest that aqueous reactions between HA and AS were starting to
303 produce minor amounts of BrC, and therefore that not all HA had evaporated from the aqueous
304 phase.

305 After a dark cloud event in Experiment B at $t = 14:40$ (“Cloud 1”), MAC_{365} continued to increase
306 and HA(g) concentrations dropped by a factor of ~ 2 , suggesting cloud droplet and/or wall uptake
307 of HA(g) occurred without a noticeable increase in aqueous-phase BrC formation rates. Turning
308 on the chamber solar simulator lights and a photolytic cloud event (“Cloud 2”) did not change
309 these trends over the next 45 min, indicating that BrC production by the HA + AS reaction had not
310 produced any effective photosensitizing species. After the addition of HOOH(g) at $t = 15:55$ to
311 the illuminated chamber, TOC began to increase at a rate of 1 ppm / min for the next hour. At the
312 same time, MAC_{365} levels, which had been increasing for 2 h and approached $400\text{ cm}^2\text{ g}^{-1}$, began
313 a decline to $100\text{ cm}^2\text{ g}^{-1}$, following an exponential decay with a lifetime of 0.9 ± 0.2 h (Figure 3
314 panel *a*). Losses of HOOH(g) detected by PTR-ToF-MS during Cloud 3 (Figure 3 panel *c*) suggest
315 significant uptake into cloud droplets. From these observations we can infer that OH radicals,
316 generated by aqueous-phase photolysis in aerosol particles and in cloud droplets, generated new
317 water-soluble SOA material via HA photooxidation, while also oxidizing BrC that had previously
318 formed under low-oxidant conditions. We do not see evidence of sunlit BrC formation by a
319 radical-driven process, unlike analogous multiphase reactions of MG or GAld in sunlight.^{20, 21}

320 Experiment C (Figure 4) was similar to B, but seed aerosol particles were generated from a 2
321 mM glycine / 9 mM AS mixture instead of only AS. These dry aerosol particles initially contain
322 an 0.18 mole fraction of glycine, or 11.2% glycine by mass and $\sim 16\%$ glycine by volume,

323 calculated using densities of 1.16 and 1.77 g cm⁻³ for glycine and AS, respectively. As the seed
324 particles were added to the chamber, SMPS dry aerosol mass increased from 0 to 110 µg/m³
325 (assuming an aerosol density of 1 g cm⁻³, Figure 4 panel *d*), while the measured TOC increased by
326 50 ppb due to the presence of glycine (Figure 4 panel *a*). PTR-ToF-MS signals for HA(g) declined
327 rapidly after its addition to the dry chamber (Figure 4 panel *c*), with a lifetime of 430 ±30 s, the
328 same as in Experiment B. No increase in dried aerosol mass was recorded by SMPS, suggesting
329 once again that uptake of gas-phase HA in the dry chamber was mainly to the walls.



331 **Figure 4.** Expt. C with gas-phase hydroxy-acetone, photolysis, cloud processing, and HOOH
332 oxidation with AS/glycine seed particles in the CESAM chamber. HA addition, two water vapor

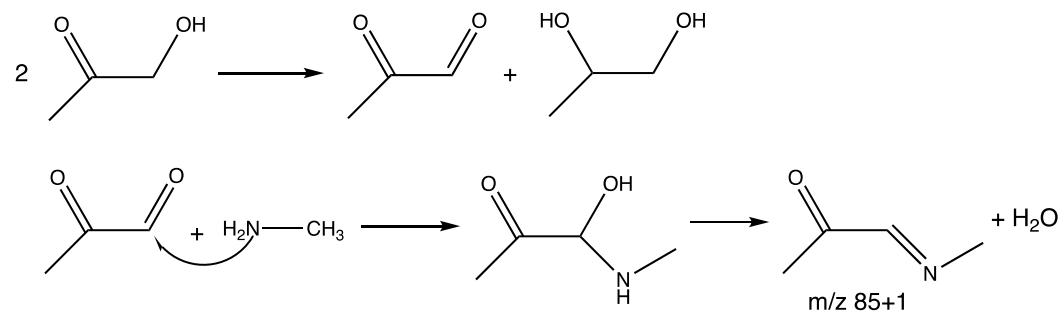
333 additions, start of chamber illumination, start of continuous HOOH addition, and 2 cloud events
334 are labeled. Panel **a**: total organic carbon readings and mass absorption coefficients at 365 nm
335 from PILS/waveguide UV-vis, color coded to axes. **b**: chamber RH and droplet spectrometer
336 counts (data ends at 11:15), color coded to axes. **c**: water-corrected PTR-ToF-MS data (*m/z* 75
337 HA, red line; *m/z* 33 HOOH frag, blue dotted line). **d**: Dilution- and wall-loss-corrected SMPS
338 total aerosol mass (assuming density = 1 g/cm³) and counts, color-coded to axes. **e**: CAPS-ssa
339 data at 450 nm (mass-corrected extinction, blue line; mass-corrected scattering black line; single-
340 scattering albedo, red dots; albedo measured immediately after instrument autozero, red triangles.

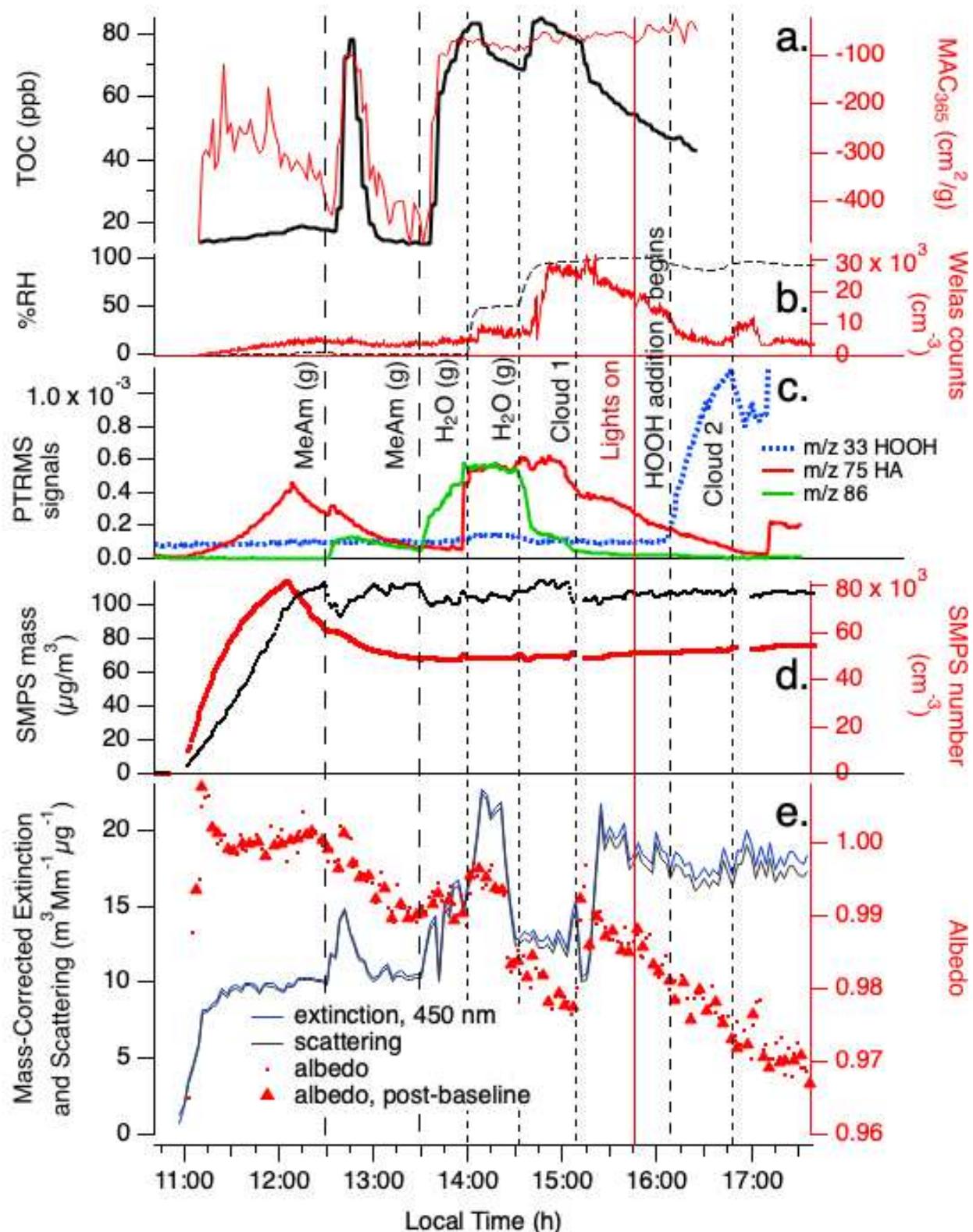
341 Upon humidification of the chamber to 58% RH in Experiment C, HA re-equilibrated from the
342 walls (Figure 4 panel *c*), as before. At the same time, SMPS mass declined by 22% within 6 min
343 (Figure 4 panel *d*), a much larger and more abrupt change compared to experiment B, while TOC
344 signals stayed constant (within 3%). After further water addition to reach 93% RH, SMPS mass
345 declined by an additional 4% in 12 min. A loss of glycine from aerosol would cause declines in
346 both SMPS mass and TOC signals, so a decline in only SMPS mass indicates a loss of mainly non-
347 carbon-containing species, such as ammonium, from the aerosol. The total observed aerosol mass
348 loss would represent a loss of more than a quarter of the AS, or essentially all of the ammonium,
349 from the seed particles. We hypothesize that an aerosol-phase reaction between HA and AS
350 generated a volatile product (rather than forming BrC), which removed a significant quantity of
351 ammonium from the aerosol phase. An SMPS-only aerosol mass loss of smaller magnitude (-
352 13%) and much slower speed took place upon humidification to 50% RH in Experiment B (without
353 glycine). Together, these comparisons suggest that the presence of aerosol-phase glycine
354 accelerated the HA + AS reaction by increasing water uptake, by deprotonating ammonium ions
355 through H⁺ exchange reactions, or both. In spite of this chemical reactivity, BrC formation

356 remained small throughout the RH increases and a dark cloud event, with an insignificant
357 ΔMAC_{365} of only $65 \text{ cm}^2 \text{ g}^{-1}$. During high-RH photooxidation in the last hour of Experiment C,
358 SMPS aerosol mass was quite stable, and unlike in Experiments A or B, TOC concentrations
359 declined after HOOH addition. This suggests that aqueous-phase OH photooxidation destroyed
360 aerosol-phase glycine faster than SOA could be formed from precursor HA. MAC_{365} levels
361 increased only marginally during photooxidation.

362 In Experiment D (Figure 5), HA was introduced into the chamber in the aerosol phase (along
363 with AS), and gas-phase methylamine was added twice. Some HA escaped to the gas phase during
364 the aerosol fill, likely as aerosol dried on the way into the chamber, resulting in initial gas-phase
365 HA concentrations that were $3 - 8 \times$ lower than in previous experiments. However, once the fill
366 ended, HA(g) PTR-ToF-MS signals declined with a lifetime of $1400 \pm 300 \text{ s}$, 4 times slower than
367 in other experiments, possibly because of continued release of HA(g) from the dry seed particles.
368 When 500 ppb methylamine gas was added to the dry chamber, gas-phase HA concentrations
369 increased slightly, MAC_{365} and TOC levels spiked significantly upward before returning to
370 baseline 30 min later, and SMPS aerosol mass dropped by $17 \mu\text{g}/\text{m}^3$ before recovering over 30
371 min. These observations suggest that MeAm was taken up by the dry aerosol even as it released
372 ammonia (and possibly small amounts of HA and/or reaction products), but that all C-containing
373 compounds produced at the aerosol surface ultimately evaporated. PTR-ToF-MS signals at m/z
374 86 ($\text{C}_4\text{H}_7\text{NO}$) increased each time MeAm(g) was added. This species could form via HA
375 disproportionation into MG and 1,2-dihydroxypropane⁴⁰, followed by imine formation, as shown
376 in Scheme 1. The fact that this species was not detected in the aerosol phase at the end of the
377 experiment (see section 3.3) suggests that its gas-particle partitioning ratio is very high, and/or that
378 it is too reactive to survive photooxidation.

379 **Scheme 1:** HA disproportionation and imine formation from methylglyoxal



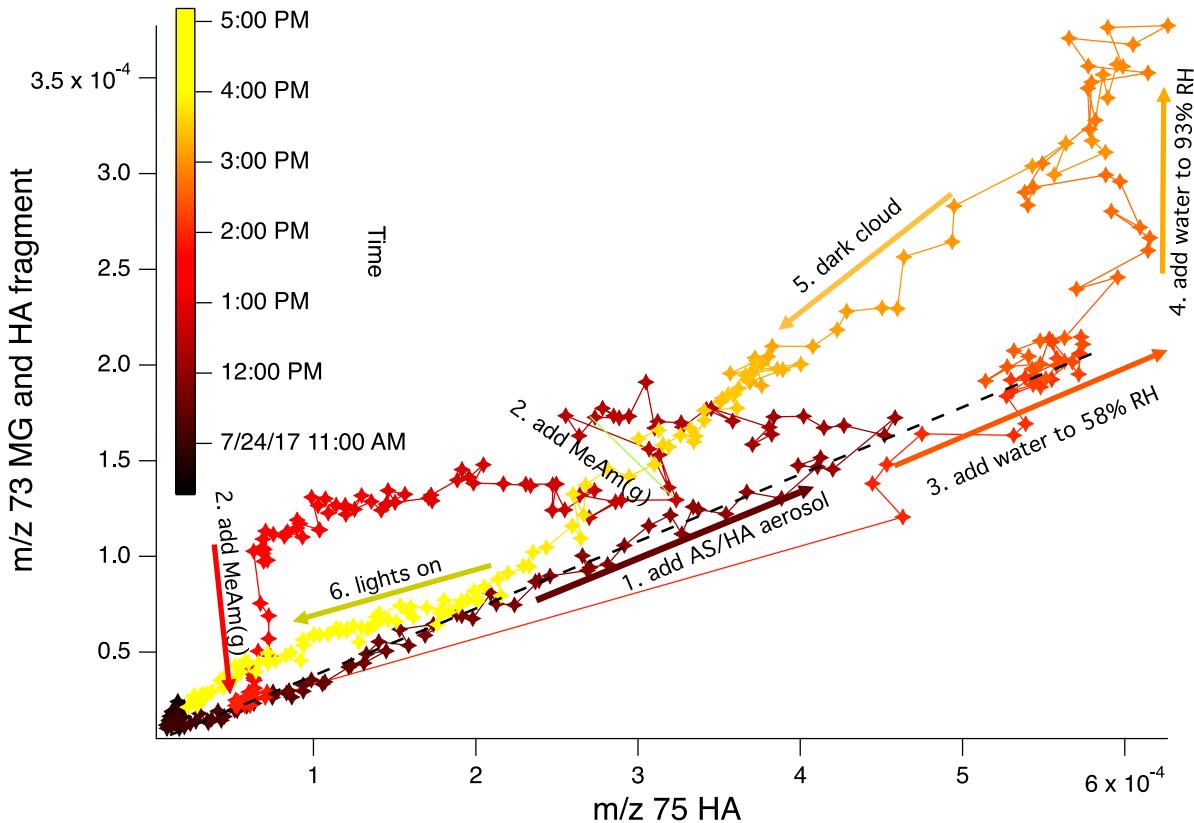


382 **Figure 5.** Expt. D with gas-phase methylamine, photolysis, cloud processing, and HOOH
 383 oxidation with AS+HA seed particles in the CESAM chamber. MeAm additions, two water vapor

384 additions, start of chamber illumination, start of continuous HOOH addition, and 2 cloud events
385 are labeled. Panel **a**: total organic carbon readings and mass absorption coefficients at 365 nm
386 from PILS/waveguide UV-vis, color coded to axes. **b**: chamber RH and droplet spectrometer
387 counts, color coded to axes. **c**: water-corrected PTR-ToF-MS data (*m/z* 75 HA, red line; *m/z* 33
388 HOOH frag, blue dotted line; *m/z* 86, green line). **d**: Dilution- and wall-loss-corrected SMPS total
389 aerosol mass (assuming density = 1 g/cm³) and counts, color-coded to axes. **e**: CAPS-ssa data at
390 450 nm (mass-corrected extinction, blue line; mass-corrected scattering black line; single-
391 scattering albedo, red dots; albedo measured immediately after instrument autozero, red triangles).

392 Evidence for HA disproportionation is seen in Figure 6, where PTR-ToF-MS signals at *m/z* 73
393 (with a contribution from both MG and a HA fragment) and *m/z* 75 (HA) are compared. The ratio
394 of *m/z* 73/75 has a lower limit of 0.36, which is interpreted as the fractional contribution of the HA
395 fragment to the *m/z* 73 signal. Ratios higher than 0.36 likely indicate a contribution from MG.
396 After the chamber is filled with AS/HA seeds (arrow 1), this ratio rises as HA signals drop and
397 *m/z* 73 signals increase slightly, reaching as high as 1.67 just before the second MeAm addition
398 (arrow 2). After each MeAm addition, the ratio returns to 0.36, suggesting that MeAm reacts
399 preferentially with MG to form products as shown in Scheme 1. Production of MG is especially
400 prominent after aerosol deliquescence (arrow 4).

401 The second addition of 500 ppb MeAm gas at *t* = 13:30 in Experiment D caused similar changes
402 as the first addition: TOC increased by 70 ppb, MAC₃₆₅ increased by 400 cm² g⁻¹, and SMPS mass
403 dropped by 14 µg/m³, but this time none of them returned to their previous values afterward,
404 indicating that at least a portion of the BrC SOA formed was stable against evaporation.
405 Production and evaporation of the imine product (PTR-ToF-MS signals at *m/z* 86) was more
406 pronounced after the second MeAm addition, with signals increasing continuously.



407

408 **Figure 6.** PTR-ToF-MS signal traces for m/z 73 vs m/z 75 recorded during Experiment D, color-
 409 coded according to time. Chamber events are noted in numbered order. For aerosol addition
 410 (arrow 1) and water addition to 58% RH (arrow 3), the data falls on the same line of an m/z 73/75
 411 ratio of 0.36 (black dashes). This signal ratio likely accounts for the fractional contribution to m/z
 412 73 of HA fragmentation. Data above this ratio likely indicates the presence of MG contributing to
 413 the m/z 73 signal.

414 Chamber humidification to 58% RH at $t = 14:00$ in Experiment D caused an increase in gas-
 415 phase PTR-ToF-MS HA signals at m/z 75 that was similar in magnitude to Experiments B and C,
 416 but the simultaneous 17% decline in TOC levels was unique to Experiment D, as expected since
 417 it was the only experiment with HA-containing seed particles. Further humidification to 95% RH
 418 allowed TOC to quickly recover to its previous level at the same time as the imine product was
 419 rapidly removed from the gas phase. SMPS aerosol mass rose slightly ($+7 \mu\text{g}/\text{m}^3$) while single-

420 scattering albedo at 450 nm declined from 0.993 to 0.977, indicating the production of BrC.
421 However, a subsequent dark cloud event at $t = 15:10$ reversed most of these changes: TOC and
422 aerosol mass declined, and albedo at 450 nm increased back to 0.99. No further SOA production
423 was observed as chamber lights were turned on and HOOH was added, even though gas-phase HA
424 was present.

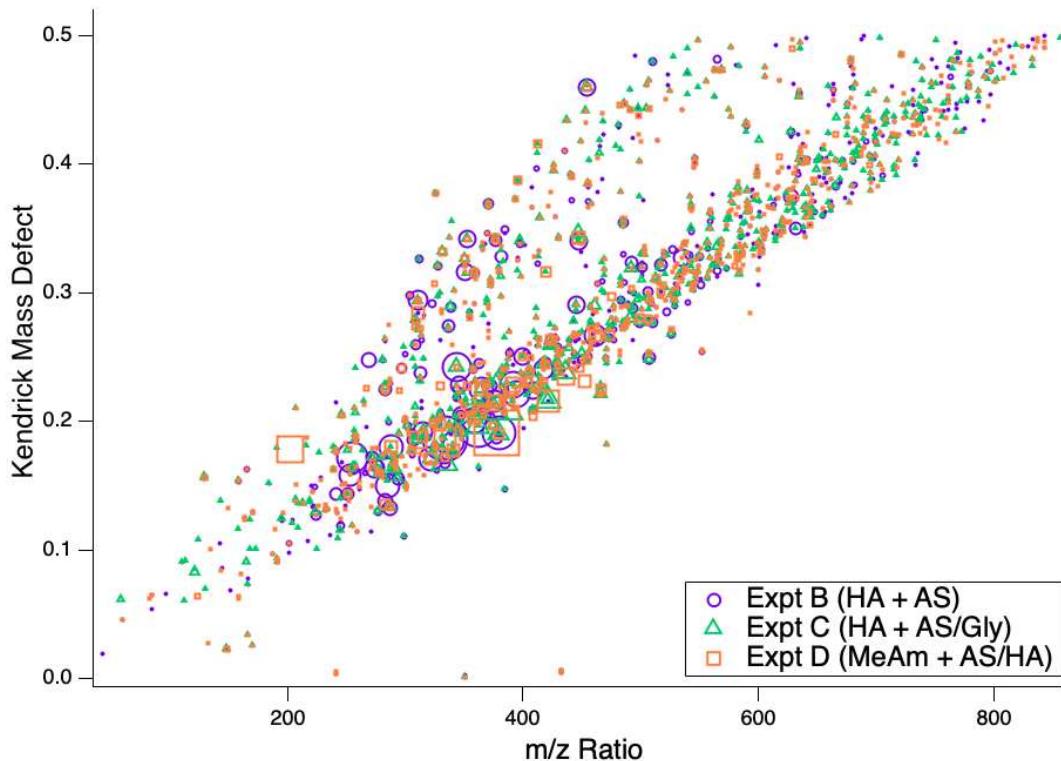
425 In summary, these chamber experiments do indicate that Maillard reactions involving HA
426 accelerate in aerosol particles and droplets, as seen by rapid changes in aerosol mass. However,
427 most of the rapid changes are mass loss, rather than SOA production, indicating that the early-
428 generation products of these reactions are prone to evaporate from aerosol particles in a multiphase
429 experimental system. Furthermore, Maillard reactions involving HA do not seem to generate
430 products capable of efficient photosensitization, and as a result net SOA and BrC generation was
431 quite minimal during the photooxidation stage of these experiments. As described in the next
432 section, however, the aerosol phase that remained after some products evaporated was itself
433 transformed by oligomerization.

434

435 3.3 Aerosol-phase reaction products. The aerosol-phase molecular composition was probed by
436 offline UHPLC-(+)ESI-HRMS analysis of filters collected from the CESAM chamber after the
437 end of Experiments B – D. Approximately 1550 species in the m/z range 40-1000 were detected
438 across the three experiments, which in each case separated into two main bands in a Kendrick mass
439 defect (KMD) plot (Figure 7): a smaller group of generally less-oxidized compounds with higher
440 KMD values and a larger group of more-oxidized compounds with lower KMD values. Peak-
441 area-weighted average molecular weights were nearly identical (407 ± 6 amu) in all experiments,
442 suggesting efficient oligomerization in the organic aerosol phase, regardless of whether HA was
443 initially added to the chamber in the gas or aerosol phase. Of these 1550 species, exact masses

444 were used to determine formulas and precursor species for a subgroup of 54 major peaks
445 representing 25% of the total peak area, focusing on masses below m/z 425 where formulas could
446 be assigned with greater confidence. SOA products (all in the elemental groups CHO, CHN, and
447 CHON) represented 39 of these major species (Table 4), with inorganic salts (11 peaks, 2% of
448 subgroup peak area), acetonitrile solvent (2 peaks, <0.2%), and two phthalate contaminants (11%)
449 responsible for the remainder. Chromatographic retention times, detected peak areas in each
450 experiment, and delta values (differences between theoretical and measured exact masses) are
451 listed for the 39 major organic products in Table S1. Kendrick mass defect plots for peaks with
452 assigned chemical formulas are shown for Experiments B-D in Figure S3.

453

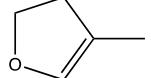
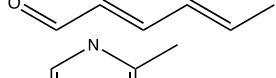
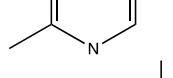
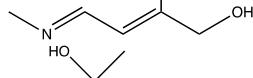
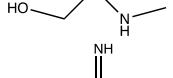
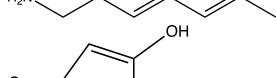
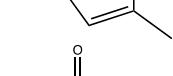
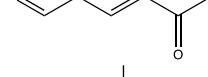
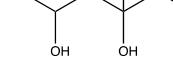
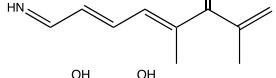
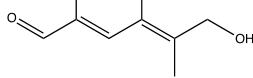
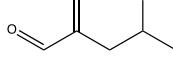
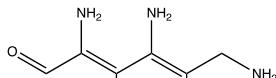
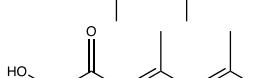
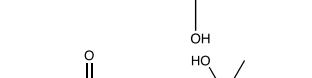


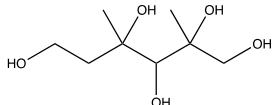
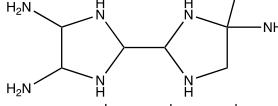
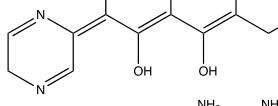
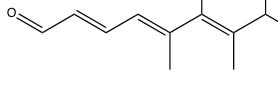
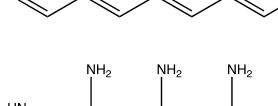
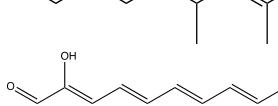
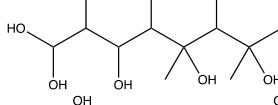
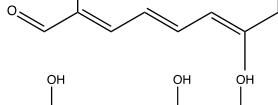
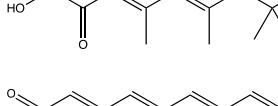
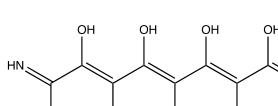
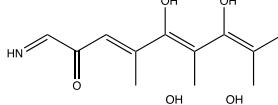
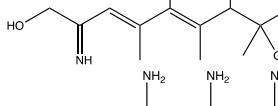
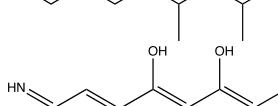
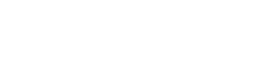
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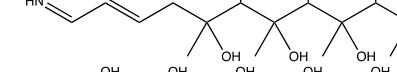
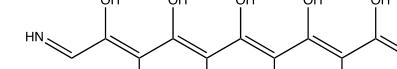
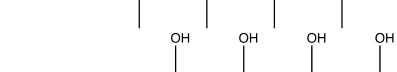
455 **Figure 7.** Kendrick mass defect plot showing differences from unit mass of each peak detected
 456 by UHPLC-(+)ESI-HRMS analysis of filter extracts of aerosol collected after Experiments B – D.
 457 Colors and symbols indicate whether ions were detected in Experiment B (purple circles),
 458 Experiment C (green triangles), and Experiment D (orange squares). Peaks with $KMD < 0$ are due
 459 to inorganic salts.

460

461 Table 4: Assigned Formulas, Precursor Species, and Proposed Structures of Detected Aerosol-
462 Phase Species in Experiments B-D

<i>m/z</i>	Formula assigned	Ion	Unsat	Precursors	Proposed Structure
85.0653	C5H8O	H+	2	2 HA - CO ₂	
97.0660	C6H8O	H+	3	3 Aald	
111.0907	C6H10N2	H+	3	2 HA + 2 NH ₃	
114.0917	C6H11NO	H+	2	HA + Aald + Glycine - CO ₂	
122.0820	C4H11NO3	H+	1	HA + Glycine - CO ₂	
125.1074	C7H12N2	H+	3	HA + 2AAld + 2 NH ₃	
133.0278	C6H6O2	Na+	4	MG + HA	
149.0233	C6H6O3	Na+	4	2MG	
159.0623	C5H12O4	Na+	0	HA + Aald	
163.1232	C10H14N2	H+	5	2HA + 2AAld + 2 NH ₃	
165.0906	C10H12O2	H+	3	2AAld + 2HA	
167.0336	C6H8O4	Na+	3	2MG	
170.1297	C8H15N3O	H+	3	GAld + 2HA + NH ₃	
187.0967	C9H14O4	H+	3	3HA	
191.0902	C8H14O5	H+	2	GAld + 2HA	

m/z	Formula assigned	Ion	Unsat	Precursors	Proposed Structure
195.1239	C8H18O5	H+	0	MG +2HA -CO ₂	
202.1784	C7H19N7	H+	2	2GX +HA +7NH ₃	
251.1436	C13H18N2O3	H+	6	2Gald +3HA +2NH ₃	
255.1720	C13H22N2O3	H+	4	2AAld +3HA +2NH ₃	
279.1601	C16H22O4	H+	6	5AAld +2HA	
283.2251	C13H26N6O	H+	4	2GAld +3HA +6NH ₃	
285.1496	C18H20O3	H+	9	GAld +8AAld	
287.1337	C10H22O9*	H+	0	2GAld +2HA	
288.1800	C14H25NO5	H+	3	GAld +2AAld +2HA +NH ₃	
294.1551	C12H20O7	NH ₄ +	3	MG +3HA	
311.1852	C17H26O5*	H+	5	HA +7AAld	
322.1712	C17H23NO5*	H+	7	GAld +5HA +NH ₃	
334.1677	C18H23NO5	H+	8	MG +5HA +NH ₃	
334.1866	C15H27NO7*	H+	3	5HA +NH ₃	
344.2425	C19H29N5O	H+	8	AAld +GAld+5HA +5NH ₃	
362.1952	C20H27NO5	H+	8	AAld +6HA +NH ₃	

m/z	Formula assigned	Ion	Unsat	Precursors	Proposed Structure
365.2250	C16H32N2O7	H ⁺	2	2AAld +4HA +2NH ₃	
378.1908	C20H27NO6	H ⁺	8	GAld +6HA +NH ₃	
378.2125	C17H28O8	NH ₄ ⁺	4	AAld +5HA	
421.2161	C18H32N2O9*	H ⁺	4	2MG +4HA +2NH ₃	

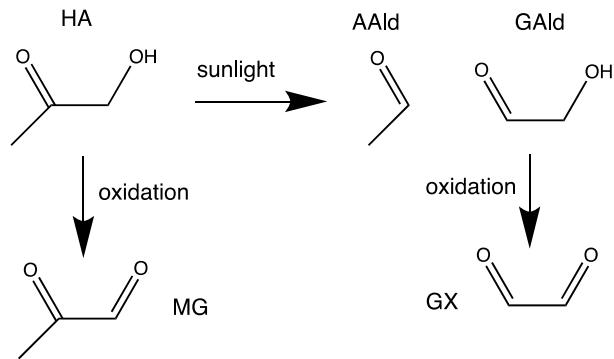
463 Notes: m/z = detected mass-to-charge ratio is positive ion ESI mode. Ion = ionization. Unsat =
 464 degrees of unsaturation. * = recommended tracer ions for HA oligomerization chemistry.

465
466 The 39 SOA molecular formulas were then used to determine which combinations of precursor
467 molecules could have formed them through oligomerization. Precursors included the molecules
468 present in a given experiment (NH₃ and HA in all experiments, plus glycine in experiment C and
469 methylamine in experiment D). However, in order to explain observed major formulas, photolysis
470 products (AAld and GAld) and oxidation products (GX and MG, produced from GAld and HA,
471 respectively) also needed to be included in the pool of possible precursors (Scheme 2). In support
472 of this scheme, using Experiment B as an example (Figures 8 and S4,), we note that PTR-ToF-MS
473 signals for the proposed photolysis products AAld and GAld increased by 41 and 19%,
474 respectively, in 20 min. after the chamber lights were turned on, while HA, GX, and MG signals
475 changed by less than 4%. This strongly supports the idea that AAld and GAld are photolytically
476 produced in this experiment. Furthermore, once HOOH was added as an OH radical source, PTR-
477 ToF-MS signal changes indicated that the gas-phase concentrations of all four of the proposed
478 aldehyde products in Scheme 2 increased by factors of approximately 2 to 4 over the next 100
479 min., while HA concentrations declined by 40%. Furthermore, PTR-ToF-MS signals decreases
480 observed for all five species during cloud 3 (Figure S4) indicate that HA and the four aldehydes

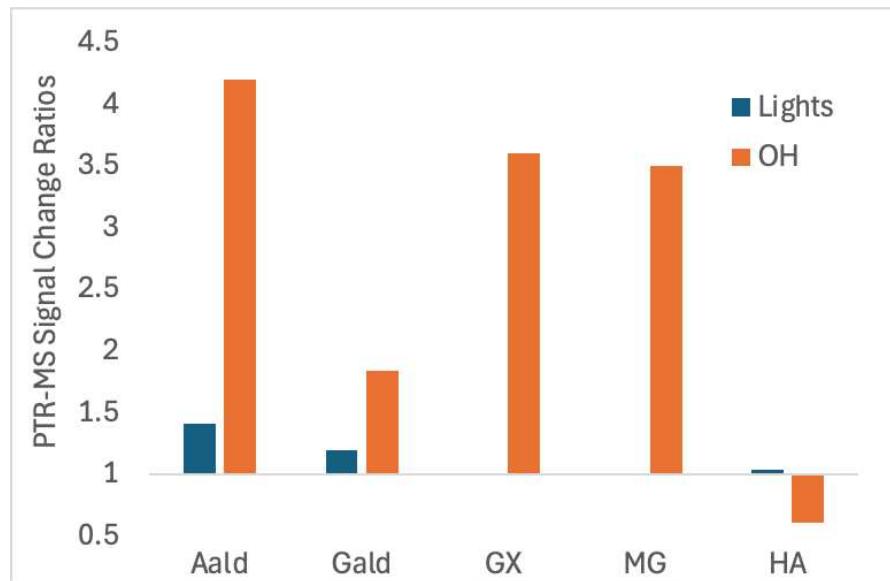
481 are all transferring to the aqueous phase (either aqueous aerosol or wall-adsorbed water).
482 Collectively, this is strong evidence that photolysis and photooxidation together are producing the
483 four aldehyde species that we identified as precursors of aerosol-phase oligomers, and that these
484 four aldehydes are available along with HA for aqueous-phase reactions.

485

486 **Scheme 2:** formation of oligomer precursors by HA photolysis and oxidation



487



488

489 **Figure 8:** Summary of PTR-ToF-MS signal change ratios in Experiment B as a result of turning
490 chamber lights on (blue) and adding HOOH as an OH radical source (orange), for acetaldehyde
491 (AAld, m/z 45), glycolaldehyde (GAlld, m/z 61), glyoxal (GX, m/z 59), methylglyoxal (MG, m/z
492 73), and hydroxyacetone (HA, m/z 75), respectively. Signal change ratios are expressed for each

493 event as (signal after) / (signal before), such that 1 = no response. GX signals remained at
494 background when the lights were turned on, and MG signals increased by less than 1% (bar not
495 visible). Changes after turning lights on and adding HOOH were measured over 20 and 100 min.,
496 respectively.

497

498 On average across all experiments (weighted by peak area), detected molecules contained 3.9
499 HA units, 1.5 NH₃ units, 0.83 AAld units, 0.45 GAld units, 0.26 MG units, and 0.10 GX units.
500 Incorporation of glycine followed by decarboxylation was needed to rationalize 2 formulas, all
501 detected only in Experiment C (which contained glycine as a reactant). GX units were needed to
502 rationalize only 1 formula, detected in Experiment D. Although a gas-phase molecule
503 incorporating methylamine was detected in Experiment D by PTR-ToF-MS at *m/z* 86, no aerosol-
504 phase molecules incorporating methylamine were found in the subgroup of 54 major peaks. Of
505 the 39 identified SOA chemical formulas, more than half could be assigned to structures generated
506 from 1-2 total units of AAld, GAld, or MG attached to 1-6 units of HA and 0-2 units of ammonia.
507 Only three formulas could be attributed to dark chemistry, such as products of HA aldol or acetal-
508 forming self-reactions or HA + NH₃ / amine Maillard reactions.

509

510 **4. Atmospheric Implications**

511 The abundance of HA oligomers incorporating other species produced by photolysis or
512 oxidation, combined with slow aqueous-phase reaction rates and a lack of “dark reaction” HA
513 oligomers and Maillard products among the identified aerosol-phase species, suggest that daytime
514 oligomerization in the HA system is primarily a radical-initiated process. A similar conclusion
515 was reached for daytime MG oligomerization.³² Direct photolysis of aqueous-phase BrC species,
516 photosensitization, or oxidation by OH radicals can produce organic radical species that initiate

517 oligomerization of closed-shell species (such as HA) that are much more numerous in the aqueous
518 phase. Due to the lack of evidence previously noted for direct photolysis of BrC or for
519 photosensitization in these chamber experiments, OH radicals are the most likely instigators of
520 oligomerization chemistry in HA-containing aqueous aerosol.

521 We observed only minor browning in our chamber experiments, with a maximum MAC_{365}
522 increase of only $500 \text{ cm}^2 \text{ g}^{-1}$, and minimum albedo at 450 nm of only 0.97, and the browning
523 observed did not correlate with photooxidation periods. From this, we conclude that there is no
524 significant pathway of radical-driven BrC production involving HA. Offline chemical analysis of
525 the organic aerosol phase, however, shows that particles were largely oligomerized by the end of
526 each experiment, and that this oligomerization appeared to be triggered by radicals produced by
527 photolysis and/or photooxidation. From this we conclude that there is indeed a significant pathway
528 of radical-driven SOA production involving HA, but that the SOA produced is very weakly light-
529 absorbing at wavelengths of 365 nm and above, unlike the brown SOA produced by GX or MG
530 undergoing multiphase photooxidation under comparable conditions (in the presence of AS and
531 amine species). We hypothesize that weak light absorption is due to the relative lack of aromatic
532 heterocycles seen in the HA oligomers identified in this work – only 1 minor peak out of 39 was
533 assigned to an aromatic heterocycle, and this was one of the three identified dark Maillard
534 products, dimethylpyrazine (m/z 111), a product also identified in dark MG + AS reactions.⁴¹

535 Aromatic heterocycles such as imidazoles^{15, 42} and pyrazines⁴¹ are formed extensively in GX,
536 MG, and GAld Maillard reaction systems, and this formation is likely key to their greater light
537 absorption. The facile formation of aromatic heterocycles appears to require two difunctional
538 aldehyde molecules to react with each other. In these experiments, the difunctional aldehydes GX,
539 MG, and GAld were all produced by HA photooxidation and all are capable of aromatic

540 heterocycle formation. However, they appear to have reacted with more abundant HA and AAld
541 molecules instead of each other.

542 Reported gas-phase HA concentrations in the atmosphere range from 0.005 to 2 ppb,⁴³⁻⁴⁵ which
543 is roughly 1 to 4 orders of magnitude lower than HA concentrations used in this work. However,
544 the fraction of HA in the aerosol phase averaged 43% in one field study, higher than any other
545 aldehyde,⁴⁴ and showed a diurnal transfer between gas and aerosol phase. Average aerosol-phase
546 HA concentrations of 97 ng m⁻³ were measured at Mt. Tai, China.⁴⁴ It is therefore clear that HA
547 is commonly present in atmospheric aerosol, but at lower concentrations than in the chamber
548 aerosol in Experiments A – D in this study. These lower atmospheric concentrations will make
549 the formation of HA oligomers and HA-based BrC less likely. However, HA in the atmosphere is
550 still likely to form AAld, GAld, GX, and MG during photooxidation, and participate in radical-
551 initiated oligomer formation with a range of other species, especially when crowded at air-water
552 interfaces with other surface-active species. The importance of the air-water interface in this work
553 is seen in the fact that at least 80% of ions detected in the photooxidized aerosol phase were not
554 detected in an earlier study of HA + AS reaction mixtures that were photolyzed in bulk aqueous
555 solution.¹⁰ Thus, while HA is unlikely to be a significant source of atmospheric BrC, it is expected
556 to contribute significantly to multiphase SOA formation in the presence of aqueous particles,
557 similar to other small multifunctional carbonyl species.

558 Certain of the HA-containing SOA products identified in this study share chemical formulas
559 with compounds detected in recent field studies, and are unique enough that they may be useful as
560 atmospheric tracers of HA oligomerization chemistry in future source apportionment studies. All
561 of the C_xH_yO_z product formulas have been identified in aerosol field studies, but most of these
562 formulas already have been linked to products of biomass burning or specific oxidation and/or
563 dimerization processes of biogenic precursor species, precluding their potential use as chemical

564 tracer ions for HA oligomerization. The exceptions are C₁₇H₂₆O₅ (*m/z* 311 with H⁺ ionization),
565 which was detected in a Zurich, Switzerland winter field study and assigned to a “non-source
566 specific” SOA principle component,⁴⁶ and C₁₀H₂₂O₉ (*m/z* 287 with H⁺), which was a “top five”
567 particle-phase HOM in the BAECC field study in Hyytiälä, Finland.⁴⁷ Few of the N-containing
568 product formulas have been reported in aerosol field studies to date, with three exceptions.
569 C₁₅H₂₇NO₇ (*m/z* 334 with H⁺) was detected by FT-ICR-MS in aerosol WSOC during summer 2010
570 at the Storm Peak Laboratory (3210 m ASL) near Steamboat Springs, CO. C₁₇H₂₃NO₅ (*m/z* 322
571 with H⁺) and C₁₈H₃₂N₂O₉ (*m/z* 421 with H⁺) were detected during the ACROSS field study of an
572 urban plume from Paris advecting over a forest.⁴⁸ These latter two compounds were assigned to
573 “terpene-isoprene HOMs” and “sunlit terpene” principal components, respectively. Since HA is
574 a common oxidation product of isoprene and other biogenic precursor species,¹ the prior
575 classifications of most of these chemical formulas as biogenic SOA in field studies are certainly
576 compatible with HA oligomerization chemistry. Thus, the five listed tracer ions (and perhaps other
577 N-containing ions from Table 4) may be useful for attributing SOA formation to HA
578 oligomerization in future field studies.

579

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591 **Data availability:** Concentration-time profiles for the large chamber experiments are freely
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593 AERIS for the benefit of ACTRIS ERIC. Expt. A: <https://doi.org/10.25326/YRJG-ES08>.
594 Expt. B: <https://doi.org/10.25326/D7Y1-S078>. Expt C: <https://doi.org/10.25326/4VYF-QG62>. Expt. D: <https://doi.org/10.25326/78ZM-8T09>. Expt. E:
595 <https://doi.org/10.25326/9C7C-X242>.
596

597
598 **Supporting Information Available**
599 Details on significant oligomer molecules detected in SOA by UHPLC/ESI-HR-QTOFMS in
600 experiments B-D, graphical summaries of Experiments A and E, individual Kendrick mass defect
601 plots for experiments B-D, and PTR-ToF-MS traces during Experiment B for HA and four
602 aldehyde species. This material is available free of charge at ACS Publications.
603

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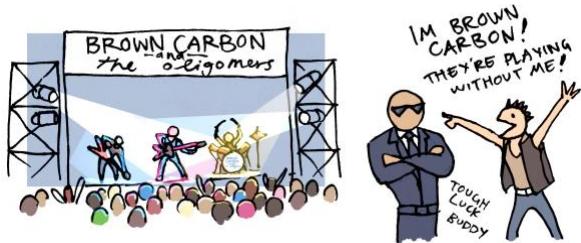
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